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Microwave Refractive Index Structure Function Profiles  $(C_n^2)$  Measured From a Small Aircraft

IAMES F. MORRISSEY YUTAKA IZUMI OWEN R. COTE



5 March 1987



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Microwave Refractive Index Structure Function Profiles  $(C_n^2)$  Measured From a Small Aircraft

#### 1. INTRODUCTION

In September 1985 a series of airplane flights were performed to obtain measurements of the microwave refractive index structure function  $(C_n^2)$  in and around inversion layers. This series had two goals associated with  $C_n^2$  measurement: one was to provide evidence on the validity of the measurements themselves; the other was to provide information on the variation of  $C_n^2$  in the first few km of the atmosphere. Of primary importance was the thickness and magnitude of any enhanced layers. The basic interest in the measurements of  $C_n^2$  was to understand the performance of troposcatter radio performance. The magnitude and shape of enhanced layers strongly influence the performance of these radios and can dictate the strategy of making the  $C_n^2$  measurements themselves.

The tests were carried out at Chatham, Mass. The  $C_n^2$  measurements were obtained using a microwave refractometer with the sensing cavity mounted to the wing strut of a Cessna-172 aircraft. Supporting atmospheric data were obtained from a portable radiosonde system and from synoptic radiosondes. These radiosonde flights were taken at the National Weather Service (NWS) station at

<sup>(</sup>Received for Publication 3 March 1987)

Chatham. An earlier report  $\frac{1}{2}$  compared the measurement of radio ducts by the two radiosonde systems and the airborne refractometer.

#### 2. DESCRIPTION OF MEASUREMENT

#### 2.1 Measurement Technique

A time sequence of refractive index variation measurements  $n(t_i)$  is obtained with a microwave refractometer mounted under the wing of an aircraft. This time sequence is converted to a power spectral density function  $\phi_n$  by taking the fast Fourier transform (FFT) and dividing by the frequency separation

$$\phi_n(f_i) = A(f_i)^2 / \Delta f , \qquad (1)$$

where A(f)'s are the amplitude from a FFT and  $\Delta f$  is the width between frequencies  $(f_i)$ .

Assuming the Taylor hypothesis to be valid, the time spectrum is then converted to a spatial spectrum by dividing by the average velocity of the air relative to the sensor:

$$\phi_n(\mathbf{k}_i) = \phi_n(\mathbf{f}_i)/\mathbf{V} , \qquad (2)$$

where V is the velocity of air relative to the sensor, and k is the wavenumber.

These spectra are then used to calculate  $C_n^2$  according to the -5/3 power law for three-dimensional turbulence in the inertial subrange given by Tartarski<sup>2</sup> and Wyngaard and LeMone.<sup>3</sup>

 $C_n^2 = 0.25 \phi_n^{-5/3}$  (3)

On some flights, two refractometers were used, generating two time sequences of refractive index measurements  $[n_1(t_i), n_2(t_i)]$ . The coherence spectrum and phase spectrum for these time sequences were computed using

Morrissey, J., Izumi, Y., and Coté, O. (1986) <u>Intercomparisons of Radiosondes and an Airborne Refractometer for Measuring Radio Ducts</u>, AFGL-TR-86-0143, AD A175150.

<sup>2.</sup> Tartarski, V. L. (1961) <u>Wave Propagation in a Turbulent Medium</u>, McGraw-Hill, New York.

<sup>3.</sup> Wyngaard, J. C., and LeMone, M.A. (1980) Behavior of the refractive index structure parameter in the entraining convective boundary layer, <u>J. Atmos.</u> <u>Sci.</u> 37:1573-1585.

$$Coh(f) = \frac{\{[\phi_{n_1} n_2^{(f)}]^2 + [QD_{n_1} n_2^{(f)}]^2\}^{1/2}}{\phi_{n_1}(f) \phi_{n_2}(f)}$$

and

Phase(f) = 
$$\tan^{-1} \frac{QD_{u_2}u_2}{\neq n_1 n_2}^{(f)}$$
,

where  $\phi_{n_1 n_2}$  is the cospectrum of  $n_1$  and  $n_2$ ,

 $QD_{n_1 n_2}$  is the quadrative cospectrum of  $n_1$  and  $n_2$ , and  $\phi_{n_1} \phi_{n_2}$  are the power spectra of  $n_1$  and  $n_2$ .

2.2 Flight Technique

To obtain the necessary time sequence of refractive index measurements, the sensing cavity of a microwave refractometer was attached to the wing strut of a C-172 aircraft just below the leading edge of the aircraft wing (Figure 1). The aircraft was then flown at various altitudes, maintaining altitude, air speed, and direction as constant as possible. These level data gathering runs were at ~100 knots true air speed (TAS), lasted 5-1/2 min or more, and covered in excess of 10 miles. These runs were from about 5 miles south of the Chathan Weather Service Station over Monomoy Island to about 5 miles north of the station beside Orleans (Figure 2).

The flight levels were selected for several reasons. Usually, one level was at about 3 km to characterize the area well above the inversions, if any. If there were inversions, several levels were selected to be either just above, in, or just below the inversion. On some flights, one or more of these levels were repeated. Then, when time permitted, a series of low level runs were performed at typically 15, 30, 60, 120, and 240 m altitude. The selection of the flight levels was made using information transmitted to the aircraft from the radiosonde runs. In addition, each flight began with the aircraft spiralling up from the surface to 3 km, and data from this ascent was also used in selecting levels.

(4)



Figure 1. Microwave Cavity Mounted on Strut of Cessna-172 Aircraft as Flown

#### 2.3 Measurement Sensors/System

The primary sensor is the microwave refractometer previously developed at the National Telecommunications and Information Sciences (NTIS) laboratory in Boulder, Colo.<sup>4</sup> The refractometer is a single cavity Birnbaum type. Figure 3 is a simplified block diagram of the refractometer. The cavity is continuously driven by a signal that is centered on the resonant frequency of the cavity, has a 10 kHz sweep frequency, and causes a deviation of  $\pm 40$  kHz. The center frequency is made to track the resonant frequency of the cavity by feedback from a phase sensitive detector that offsets the voltage controlled oscillator (VCXO). The output of the VCXO is averaged and scaled to produce a digital output that is displayed on the front panel. This same digital output is passed through a

Thompson, M. C., Marler, F. E., and Allen, K. C. (1985) Measurement of the microwave structure constant profile, IEEE Trans. Antennas Propag. AF-28(No. 2):278-280.







Figure 3. Block Diagram of Microwave Refractometer

digital-to-analog converter and represents one of the analog outputs. This channel has a sensitivity of 0.01 V/n units and an averaging time/update rate of 1 sec as flown. The second analog channel is obtained by taking the analog error voltage and passing it through a bandpass filter (1/100 to 10 cycles) amplifier. This channel, as flown, had a maximum sensitivity of 1.04 V/n units (1.28 V/n, refractometer No. 2) at a frequency of 2 cycles.

In addition to the refractometer, the air temperature was measured by a bead thermistor mounted just above the refractometer cavity, and pressure was measured inside the aircraft by a Rosemount Model 1201F1 pressure transducer.

A data system was designed for the  $C_n^2$  n.easurements program (Figure 4). It consists of a Euromak C computer, manufactured by Dr. Weiss GMBH, that is used to control 12 A/D (12 bit) channels. Six channels are programmed to be sampled at a rate of two times per second and six channels are programmed for 64 times per second. Only two of the fast channels were used in this series. The system is under the control of an operator using a small NEC - PC8201A computer. The operator is allowed to input titles and observations of two types; one at the start of each flight or new tape, and another in between each data run. The first is used to give the location and general observations on the test and the atmosphere. The second is used to give specific data and observations associated with the upcoming data run (i. e., air speed, altitude, heading, location, visibility, and bumpiness) as well as additions and comments on the previous run. The data were stored on a nine-track magnetic tape recorder (Kennedy 9832) during 'hese tests, but the system has since been modified to use a cassette style tape recorder (Kennedy 6455).



Figure 4. Refractometers With Associated Computers and Recorders as Flown in Tests

#### 3. DATA PROCESSING

Figure 5 symbolically shows the data processing used to obtain each estimate of  $C_n^2$ . Sixty-four seconds of the sampled high gain channel of the refractometer (4096 values) are detrended and converted to a frequency spectrum  $\psi_n(f_i)$ , which is then converted to a power density spectrum. This step required the removal of the frequency dependency,  $\frac{1}{|G_f|}$ , introduced by the pre-whitening filter in the refractometer. This frequency spectrum is converted to a wavenumber spectrum,  $\phi_n(K_i)$ , by dividing by the true air speed of the aircraft. A sample power density spectrum from the 25 September flight at 1618 EDT at Chatham is shown in Figure 6. This spectrum is then used to generate the estimate of  $C_n^2$  by using the mean value of  $C_{n_i}^2$  calculated for each wavenumber corresponding to i = 25 to i = 104. These represent wavelengths from 30 to 130 m at an air speed of 51.4 m/sec (100 knots), which is representative of the speed in these flights.



Figure 6. Sample Power Density Distribution  $\phi_n(k)$  for Flight of 25 September at 1618 EDT

During each of these  $C_n^2$  estimates, 64 sec of data, the aircraft moves about 3 km. Each level run contained five such segments. Figure 7 shows the  $C_n^2$  results from the flight on 18 September at 1525 EDT. At each level the mean, usually of five values, is shown and the standard deviation is shown as bars. The temperature and dew point profiles from the radiosonde closest in time to the aircraft flight are also shown.

On those flights where two refractometers were flown, the sampled data was processed from each refractometer separately as described above. This provides time series of dual measurements of  $C_n^2$  where the refractometers are 4.25 m apart. Figure 8 shows the  $C_n^2$  results from the 26 September flight at 1434 EDT. The results are in the order taken, with vertical lines used to separate areas of constant allitude. In addition, the two sets of sampled analog data were used to generate coherence spectra and phase spectra. This was accomplished by first calculating the cospectra and quadrature spectra in a manner analogous to that shown in Figure 5. Figure 9 shows a typical plot of these spectra from the flight of 25 September at 1618 EDT.



Figure 7.  $C_n^2$  Profile for Flight of 18 September at 1525 EDT With Temperature (T) and Dew Point Temperature (T<sub>D</sub>) From Radiosonde Flight at 1510 EDT







Figure 9. Sample of Coherence and Phase Spectra From Flight of 26 September at 1434 EDT

#### DISCUSSION

Two types of information were gained from the program; information on the measurement technique, and data on the distribution of  $C_n^2$  in the lowest couple of km of the atmosphere.

The use of two refractometers demonstrated the high degree of repeatability of the technique. This can be seen in Figure 10, which shows the digitized voltage sequences from a couple of data blocks (512 samples, 8 sec) from the flight on 25 September. This high degree of repeatability is also evident in the five figures in Appendix A that show the mean  $C_n^2$  values from two refractometers (Figures A7, A9, A11, A12, and A13). These figures show a consistent offset in the two measurements. This offset is not unexpected in that time and circumstances prevented an absolute calibration of the second refractometer. This agreement is also evident when we look at the individual  $C_n^2$  values obtained during each level portion of flight (Figure 8). Again, the offset between the refractometers is evident.

The final way that the dual refractometers were compared was by coherence and phase spectra of the two signals. Figure 9 shows the signals to be well correlated at wavelengths longer than 25 m and to have a zero degree phase shift between refractometers. At shorter wavelengths, the correlation deteriorates and the phase shift is no longer zero until, at wavelengths of less than 5 m, the picture is chaotic. This behavior is expected for two sensors separated by 4.25 m.<sup>5</sup> The signals should be well correlated with zero phase shift until the wavelength is about four times the separation, or 20 m.

The measured values of  $C_n^2$  range from a high of  $3 \times 10^{-13} \text{ m}^{-2/3}$  to a low of  $5 \times 10^{-17} \text{ m}^{-2/3}$ . Of more importance to understanding the performance of troposcatter radios is the large observed vertical gradients. One of the largest gradients can be seen at about 1600 m on the second flight of 18 September (Figure A2).  $C_n^2$  goes from  $3 \times 10^{-14}$  to  $9 \times 10^{-17} \text{ m}^{-2/3}$ , a ratio of 300 to 1, in a vertical separation of 150 m.

On several flights, we see a correspondence between the typical inversion, rising air temperature and falling dew point temperature, and a sharp decrease in  $C_n^2$  with height (Figures A2, A3, A4, A6, A10, and A14). This decrease, at times, exactly corresponds to the decrease in dew point, while, in other cases, it is displaced vertically upward a couple of hundred meters. Often there is also

Priestley, J.T., and Hill, R.J. (1985) Measuring high-frequency humidity, temperature and radio refractive index in the surface layer, <u>J. Atmos</u>. and Oceanic Technol. 2:233-251.



Figure 10. Sampled Voltage Sequences From Both Refractometers From Flight of 25 September at 1125 EDT

a sharp increase in  $C_n^2$  with increasing height just below the inversion. This results in an enhanced lupper of  $C_n^2$  right at, and sometimes slightly above, the inversion where the ratio inside the layer to outside is typically from 40 to 100 times (Figures A2, A6, A10, and A14).

On those days where there were low level runs below 200 m, we see the enhanced levels of  $C_n^2$  associated with surface friction,  $1 \times 10^{-13} \text{ m}^{-2/3}$ , which usually falls off to about  $1 \times 10^{-15} \text{ m}^{-2/3}$  by 500 m. This decrease is often not monotonic, but shows a pronounced increase with height in the first 100 m. This increase with height may be an artifact of data processing, since the altitude is less than some of the wavelengths used in calculating  $C_n^2$ .

#### 5. CONCLUSIONS

The results provide strong evidence that these measurements, obtained from a refractometer mounted on a small aircraft, can provide useful information on the small scale fluctuations of n. This can be used in the study of the refractive index structure function that influences many modern radar and propagation systems.

Several cases showed layers of very limited vertical extent (< 300 m), where  $C_n^2$  increased by 40 to 300 times. The sharp layers were usually associated with a temperature inversion and an associated sharp decrease in water vapor. The shallowness of these layers might cause them to be misrepresented in any experiment with a vertical sampling interval greater than 200 m. The shape and position of the layers could be critical to understanding the performance of various radars (airborne surveillance radars and synthetic aperture radars) and communication systems, i.e., the shifting of the layers by about 400 m could increase/decrease delay spread on a troposcatter radio link by several times.

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Appendix A

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Figure A1. Measured  $C_n^2$  Profile (18 Sept 0730) With T and T<sub>D</sub> From Radiosonde Closest in Time



Figure A2. Measured  ${C_n}^2$  Profile (18 Sept 1525) With T and T<sub>D</sub> From Radiosonde Closest in Time



Figure A3. Measured  $C_n^2$  Profile (19 Sept 0719) With T and T<sub>D</sub> From Radiosonde Closest in Time



Figure A4. Measured  $C_n^2$  Profile (20 Sept 0741) With T and TD From Radiosonde Closest in Time



Figure A5. Measured  $C_n^2$  Profile (20 Sept 1113) With T and T<sub>D</sub> From Radioconde Closest in Time



Figure A6. Measured  ${C_n}^2$  Profile (21 Sept 0747) With T and TD From Radiosonde Closest in Time



Figure A7. Measured  $C_D^2$  Profile (24 Sept 0730) With T and T<sub>D</sub> From Radiosonde Closest in Time



Figure A8. Measured  $C_n^2$  Profile (25 Sept 0653) With T and T<sub>D</sub> From Radiosonde Closest in Time







Figure A10. Measured  $C_n^2$  Profile (25 Sept 1626) With T and T<sub>D</sub> From Radiosonde Closest in Time



Figure A11. Measured  $C_n^2$  Profile (26 Sept 0707) With T and  $T_D$  From Radiosonde



Figure A12. Measured  ${C_n}^2$  Profile (26 Sept 0707) With 'T and  $T_D$  From Radiosonde Closest in Time



Figure A13. Measured  $C_n^2$  Profile (26 Sept 0707) With T and  $T_D$  From Radiosonde



Figure A14. Measured  $C_n^2$  Profile (26 Sept 1434) With T and  $T_D$  From Radiosonde Closest in Time