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NATIONAL ADVISORY COMMITTEE
FOR AERONAUTICS

TECHNICAL NOTE

No. 1200

TENTATIVE TABLES FOR THE PROPERTIES
OF THE UPPER ATMOSPHERE

By Calvin N. Warfield

for the

NACA Special Subcommittee on the Upper Atmosphere

Langley Memorial Aeronautical Laboratory
Langley Field, Va.



Washington

January, 1947

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SUMMARY

As a result of recent developments in aeronautics and ordnance, a need has arisen for tables of properties of the atmosphere at altitudes in excess of those covered by the existing standard tables (NACA Report No. 218). In order to satisfy this need, the National Advisory Committee for Aeronautics has adopted three temperature-height relationships and one composition-height relationship, and tables based upon them have been prepared for pertinent properties of the upper atmosphere (that is, from 20 to 120 kilometers in metric units, and from 65,000 to 393,700 feet in British units). In the absence of direct data, such as might be obtained by soundings with high-altitude rockets, the values adopted are based upon existing information obtained by indirect measurements of certain quantities. As a consequence, the tables are only tentative.

Two sets of tables based upon the adopted tentative standard specifications for the upper atmosphere are presented. One set of two tables is based upon the same arbitrary constant value for the acceleration of gravity as was used in the preparation of the existing standard tables for the lower levels (NACA Report No. 218). This set of tables for the upper levels of the atmosphere therefore constitutes a consistent extension of the existing standard tables. The other set of two tables takes into consideration the decrease in the acceleration of gravity with increasing altitude and therefore is more precise than the first set. Consequently, this set is presented only to satisfy the need for greater accuracy that may exist in some fields of research.

Each table is divided into separate parts for both day and night conditions at altitudes above 80 kilometers. The necessity for separate tables for day and night values is occasioned by the

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In April 1946 this Panel was superseded by the Special Subcommittee on the Upper Atmosphere which was also appointed by the NACA.

The membership of this Special Subcommittee is as follows:

- ~~Dr. Harry Hall - Navy~~
~~Dr. Joseph Kaplan, C.I.T.~~
 Dr. Harry Wexler, U. S. Weather Bureau, Chairman
 Col. D. N. Yates, Chief, Air Weather Service
 Col. Paul H. Dane, A. C., TSEAC, AAF Air Materiel Command
 Capt. H. T. Orville, USN, Office of Chief of Naval Operations,
 Navy Department
 Capt. Walter S. Diehl, USN, Bureau of Aeronautics, Navy
 Department
 ^.-Dr. Calvin N. Warfield, Langley Memorial Aeronautical Laboratory
 Dr. E. H. Krause, Naval Research Laboratory
 Dr. W. G. Brombacher, National Bureau of Standards
 Dr. L. V. Berkner, Carnegie Institution of Washington
 Dr. B. Gutenberg, California Institute of Technology
 Dr. Fred L. Whipple, Harvard Observatory, Harvard University
 Dr. O. R. Wulf, Gates and Crellin Laboratories, California
 Institute of Technology.
 Mr. Jerome Teplitz, NACA, Secretary.

This Subcommittee has considered the information available concerning temperature and composition in the upper atmosphere. On the basis of existing data obtained by balloons at altitudes up to about 32 kilometers (references 6 and 7), of indirect measurements obtained at greater heights such as those discussed in references 8 to 14, and of unpublished data resulting from similar indirect measurements, recommendations concerning temperature-height and composition-height relationships were made by the Subcommittee on June 24, 1946. The recommendations regarding temperature-height relationships cover three arbitrary sets of temperature: (1) tentative standard temperatures, (2) probable minimum temperatures, and (3) probable maximum temperatures. Also, recommendation was made that at this time no tables be prepared for altitudes in excess of 120 kilometers because of the uncertainty regarding the validity of the data in this region.

At a meeting of the executive committee of the National Advisory Committee for Aeronautics held on August 15, 1946, the previously mentioned recommendations of the Subcommittee were adopted. As a result of the adoption of the recommendations of the Subcommittee, two sets of tables for the upper atmosphere, based upon the tentative standard temperatures, have been prepared at the Langley Laboratory of the NACA.

The first set of tables provides a consistent extension of the present standard tables for the lower levels of the atmosphere

(reference 1) because the same simplifying assumption of an arbitrary constant value for the acceleration of gravity is made in both cases. Because of this consistency with the present standard atmosphere tables, and in consideration of the fact that the present standard tables (reference 1) are widely used in evaluating performance characteristics of aircraft and for design purposes, it appears that this first set of tables may also be found useful in these same fields of aeronautical engineering. In addition, in order to be consistent with present practice in the use of the terms "pressure altitude" and "density altitude" (reference 15) it appears that it may be proper to use the term "tentative pressure altitude" to designate that altitude in this first set of tables which corresponds to a specified ambient-air pressure. Likewise, the term "tentative density altitude" can consistently be used with this set of tables in connection with ambient-air densities.

The second set of tables is more precise than the first because it takes into consideration the decrease in the acceleration of gravity with increasing altitude. This set is intended primarily for use in connection with research on the properties of the upper atmosphere. Values of still greater computational precision than those listed in this second set may be obtained by means of "latitude correction factors" which have been computed and tabulated in another table.

These two sets of tables for the upper atmosphere consist of two tables each, one in the metric system of units and the other in the British system of units. The altitude range covered is from 20 kilometers and 65,000 feet, respectively, to 120 kilometers and its British equivalent of about 393,700 feet. In addition to those quantities reported in references 1 to 5, there is included the mean free path of the air molecules. This quantity has been added because of its significance at high altitudes where the molecular mean free paths may be comparable to or larger than certain dimensions of the aircraft or missiles that may be flown there.

Acknowledgement is gratefully given for the contributions made by Dr. R. G. Stone, of the AAF Weather Service, who supplied valuable data concerning maximum and minimum temperatures over the entire world to altitudes of 32 kilometers, and for the thorough technical review and excellent suggestions offered by Mr. L. P. Harrison of the U. S. Weather Bureau.

SYMBOLS

a	speed of sound
c	most probable molecular speed
\bar{c}	average molecular speed
g	acceleration of gravity
h	altitude
K	volume gradient of oxygen dissociation $\left(\frac{\Delta v}{\Delta h}\right)$
L	temperature gradient $\left(\frac{\Delta T}{\Delta h}\right)$
M	molecular weight
m	mass of a molecule
N	number of molecules per unit volume
p	pressure
R	universal gas constant
r	radius of the earth
T	absolute temperature
t	temperature
v	volume of molecular oxygen in an initial unit volume of normal air, at the same temperature and pressure
w	specific weight (gp)
γ	ratio of specific heats

λ	mean free path of molecules
μ	coefficient of viscosity
ν	kinematic viscosity (μ/ρ)
ρ	density (mass per unit volume)
σ	molecular diameter; also density ratio (ρ/ρ_0)
$\bar{\sigma}$	average molecular diameter

The following subscripts are used to refer to the indicated conditions:

O	sea level
l	lower level
a	top of region of dissociation, where oxygen is all atomic
A	base of region with constant temperature and constant composition
B	base of region with constant temperature gradient and constant composition
C	base of region with constant temperature and constant volume gradient of dissociation
D	base of region with constant temperature gradient and constant volume gradient of dissociation
g	acceleration of gravity variable
m	base of region of dissociation, where oxygen is all molecular
n	nitrogen molecules
N	non-oxygen (i. e., all constituents other than oxygen)
o	oxygen
air	mixture of molecules in atmosphere
ϕ	latitude

ADOPTED SPECIFICATIONS FOR THE UPPER ATMOSPHERE

Tentative Temperatures

Three sets of tentative temperature-height relationships have been adopted. One set gives tentative standard temperatures and the other two list values of the probable minimum and the probable maximum temperatures for the entire world. These three sets of temperatures which were originally recommended by the Subcommittee on the Upper Atmosphere are given by linear variations with altitude between the points specified in the following tabulation of temperatures.

TEMPERATURES

Altitude (km)	Probable minimum (°K) (a)	Tentative standard (°K)	Probable maximum (°K) (a)
0	225	^b 288	320
10.76923		^b 218	
11			250
17	180		
20		^b 218	
25			255
32		218	
45	200		380
50		350	
55	300		
60		350	
70			380
78		240	
80	170		300
83		240	
120	300	375	600

^aThe values of ambient air temperature listed in these two columns are not intended to represent extreme values for the entire world, and for all time, but rather values that bracket the temperatures over nearly all the earth most all the time.

^bThese values are standard, and have been used previously in references 1, 3, 4, and 5.

These temperature-altitude relationships are also shown in figure 1.

Tentative Composition

The tentative composition used in computing the tables was arrived at by taking into consideration the fact that, at altitudes below 80 kilometers in the day time and below 105 kilometers at night, the generally accepted variations in chemical composition are too small to affect appreciably the computed pressures and densities. However, it is believed that at levels above those just specified significant changes in composition result from the dissociation of oxygen molecules by solar radiation. It is furthermore known that the presence of water vapor in the atmosphere does not appreciably affect pressures and densities. As a result of such considerations, and in the interest of simplicity, the following tentative specifications for composition of the upper atmosphere were recommended by the Subcommittee and have been adopted for the purposes of computing the values in these tables:

(1) For day time, the dissociation of oxygen is such as to produce a linear volume gradient from all-molecular oxygen at 80 kilometers to all-atomic oxygen at 100 kilometers. Except for oxygen dissociation, the composition is the same as that at sea level.

(2) For night time, the dissociation of oxygen is such as to produce a linear volume gradient from all-molecular oxygen at 105 kilometers to all-atomic oxygen at 120 kilometers. Except for oxygen dissociation the composition is the same as that at sea level.

(3) At altitudes below the regions of oxygen dissociation the composition is the same as that at sea level.

(4) At altitudes above the regions in which both molecular and atomic oxygen exist, as stipulated in (1) and (2), and up to at least 120 kilometers, the composition is the same as that at sea level, except for oxygen which is in the atomic rather than in the molecular form.

The variation with altitude of the specified molecular oxygen content of the atmospheres is graphically portrayed in figure 2.

PHYSICAL RELATIONSHIPS

Basic Equations

In addition to the specifications for temperature and composition already listed, certain other assumptions are made and

serve as the basis for deriving the various equations used in computing the properties of the upper atmosphere. These additional assumptions are:

- (a) The air is dry
- (b) The air behaves as a perfect gas and hence obeys the general gas law which may be written

$$\frac{\rho}{\rho_0} = \frac{p}{p_0} \frac{T_0}{T} \frac{M}{M_0} \quad (1)$$

- (c) The air is at rest with respect to the earth and hence obeys the basic law for fluid statics

$$dp = -g_0 dh \quad (2)$$

By means of equations (1) and (2) and equations representing the adopted specifications for temperature and composition, relationships may be deduced between pressure and height. The equations representing the adopted specifications are

$$T = T_1 + L(h - h_1) \quad (3)$$

where L is the temperature gradient $\Delta T/\Delta h$, and

$$\frac{M}{M_0} = \frac{1}{1 - K(h - h_m)} \quad (4)$$

where K is the volume gradient of oxygen dissociation $\Delta v/\Delta h$. The derivation of equation (4) is given in appendix A.

In addition to the three assumptions just listed, it is necessary to make an assumption concerning the value of the acceleration of gravity. For the purpose of furnishing tables for the upper atmosphere that will be consistent with the present standard tables for the lower atmosphere (reference 1), it is necessary to make the same assumption concerning the acceleration of gravity as was used in preparing the standard tables. This assumption is

- (d) For the tables based on a constant value of g the acceleration of gravity at all altitudes is the standard sea-level value; that is,

$$g = g_0 \quad (5)$$

For those instances in which closer conformity to actual conditions is required than is inherent in these tables it is necessary to make another assumption concerning the value of the acceleration of gravity. This assumption is

- (e) For tables based on a variable value of g the acceleration of gravity varies inversely as the square of the distance from the center of the earth; that is,

$$g = g_0 \left(\frac{r}{r+h} \right)^2 \quad (6)$$

Pressure-Height Relationships

By use of the foregoing basic equations and assumptions, other equations are derived which relate pressure to altitude. Two sets of equations are used, one set based on a constant value of g as specified in assumption (d), the other set based on the variation of g that is specified in assumption (e). The deductions for the first set are indicated in appendix B and for the second set in appendix C. The equations that are based on a constant value of g are as follows:

For combination A (constant temperature and constant composition):

$$\log_e \left(\frac{p}{p_A} \right) = C_A (h - h_A) \quad (7)$$

where

$$C_A = - \frac{g_0 \rho_0}{p_0} \frac{T_0}{T} \frac{M}{M_0} \quad (8)$$

For combination B (constant temperature gradient and constant composition):

$$\log \left(\frac{p}{p_B} \right) = C_B \log \left(\frac{T}{T_B} \right) \quad (9)$$

where

$$C_B = - \frac{g_0 \rho_0 T_0}{p_0 L} \frac{M}{M_0} \quad (10)$$

For combination C (constant temperature and constant volume gradient of dissociation):

$$\log \left(\frac{p}{p_C} \right) = C_C \log \left(\frac{M}{M_C} \right) \quad (11)$$

where

$$C_C = - \frac{E_0 p_0 T_0}{p_0 K T} \quad (12)$$

For combination D (constant temperature gradient and constant volume gradient of dissociation):

$$\log \left(\frac{p}{p_D} \right) = C_D \log \left(\frac{T}{T_D} \frac{M}{M_D} \right) \quad (13)$$

where

$$C_D = \frac{-E_0 p_0 T_0 M_D}{p_0 (M_0 + M_D T_D K)} \quad (14)$$

The equations derived in appendix C, based on a variable value of g , are more complex than those listed in the foregoing and consequently they are not reproduced here.

Speed of Sound

The speed of sound at any altitude relative to that at sea level is computed by the equation

$$\frac{a}{a_0} = \left(\frac{\gamma T M_0}{\gamma_0 T_0 M} \right)^{1/2} \quad (15)$$

where the ratio of the specific heats γ , as derived in appendix A, is

$$\frac{\gamma}{\gamma_0} = 1 - \frac{128K(h - h_m)}{21M_0} \quad (16)$$

The variation with altitude of the ratio of specific heats γ for the specified atmosphere is shown in figure 3(a).

Coefficient of Viscosity

Sutherland's equation for the variation of the coefficient of viscosity with temperature is used. It is

$$\frac{\mu}{\mu_0} = \left(\frac{T}{T_0}\right)^{3/2} \left(\frac{T_0 + S}{T + S}\right) \quad (17)$$

in which, according to reference 16,

$$S = 120$$

when the T 's are in $^{\circ}\text{K}$, and

$$S = 216$$

when the T 's are in $^{\circ}\text{F}$ absolute.

A caution concerning the use of values obtained from equation (17) for the upper atmosphere is given in the section entitled "Discussion of Tables."

Molecular Mean Free Path

The ratio of the molecular mean free path at any altitude to the corresponding value at sea level is computed by

$$\frac{\lambda}{\lambda_0} = \frac{p_0 T g}{p T_0 g_0} \quad (18)$$

This equation is justified in appendix D.

BASIC CONSTANTS

In the preceding section equations are given by means of which several properties of the upper atmosphere are computed. These computations involve numerical values of the several properties at sea level. Appendix E discusses the chosen sea-level values for

each of several properties of the atmosphere and they are listed in table I in both metric and British engineering systems of units. Values are listed for each of the three specified atmospheres and in some instances the quantity is expressed in more than one unit in either the metric or British system.

The values listed in table I for the standard atmosphere at sea level are identical with those used in references 1 and 5 except in a few instances. The exceptions are noted and explained in appendix E.

DISCUSSION OF TABLES

The appropriate equation (equation (7), (9), (11) or (13) for the constant value of g , or (C3), (C5), (C10) or (C13) for the variable values of g) is used to compute the ratio of the pressure p at any height to the pressure at the base of the region to which that particular equation applies. These pressure ratios for each of the regions are then used to compute the ratio of the pressure p to the pressure p_0 at sea level. These ratios p/p_0 are given in tables II to V.

By use of the computed values of the pressure ratios p/p_0 and of the sea-level value of pressure p_0 as given in table I, the value of the pressure p is computed and then given in tables II to V. The pressures given in tables IV and V are also plotted against altitude in figure 3(b).

The remaining quantities given in tables II to V are similarly computed by means of the appropriate equation and the corresponding sea-level value given in table I. The values for these remaining quantities given in tables IV and V are also shown plotted against altitude in figures 3(c) to 3(h).

Attention is directed to the fact that all tables in this report are based on the engineering system (sometimes referred to as the gravitational system) in which the fundamental quantities are length, force, and time. The standard units for force used herein are, therefore, pounds for the British system and kilograms for the metric system.

Accuracy of Computed Tables II to V

In tables II to V all quantities except the mean free paths of the molecules are tabulated to four significant figures, and the mean free paths of the molecules are tabulated to three significant figures. All computations for table II were carried through to six significant figures and consequently the values given in this table are believed to be exact.

Most of the values for table IV were obtained from table II by use of suitable conversion factors evaluated by a graphical method described in appendix C. The errors resulting from the method, and therefore the errors in the values tabulated in table IV are believed not to exceed 0.01 of 1 percent.

A method of graphical interpolation was applied to obtain from tables II and IV the values for use at the intermediate levels tabulated in tables III and V. The accuracy of this method is such as to introduce an error of not over one-twentieth of 1 percent in the values listed in tables III and V. Consequently, whenever a discrepancy exists between the metric and British values, the metric values should govern.

Validity of Tabulated Values at the Higher Altitudes

Pressure, density, specific weight, and mean free path of molecules.- As was previously mentioned, the computations for tables II and III are based on a constant value for the acceleration of gravity g so that the values listed would be consistent with those appearing in the present standard tables for the lower levels of the atmosphere (reference 1). The errors in the computed values of pressure, density, specific weight and mean free path inherent in the assumption of a constant value for the acceleration of gravity become progressively greater with increasing altitude, being about 30 percent at 120 kilometers. However, a variation of 30 percent in pressure at 120 kilometers corresponds to a variation of less than 4 percent in altitude at this level, and at lower levels the change in altitude corresponding to the error in pressure rapidly approaches zero. It is apparent therefore that in at least some applications the values in tables II and III will be adequate and therefore useful. Furthermore, they represent an extension of the present standard tables (reference 1).

In order to satisfy the need that may exist for values that are not affected by the use of a constant value for the acceleration of gravity g , tables IV and V are presented. In these tables g is assumed to vary inversely as the square of the

distance from the center of the earth. This assumption therefore takes into consideration the variation due to gravitational attraction, but it does not allow for the effect of centrifugal force. The centrifugal force due to the rotation of the earth is known to be only a small fraction of 1 percent of the gravitational force at an altitude of 120 kilometers, and consequently this omission does not result in a significant error.

The standard value used for the acceleration of gravity at sea level (and at all altitudes for tables II and III) is 9.80665 meters per second per second. This value corresponds rather closely to the true acceleration of gravity at sea level at latitude 45° . (More specifically, it corresponds to the theoretical acceleration of gravity at sea level and at latitude $45^\circ 24'$ according to the International formula. See reference 17.) If still greater accuracy than is inherent in tables IV and V is required at latitudes far displaced from latitude 45° , an estimate of the latitude effect upon pressure and density may be obtained by use of the equation

$$\log \frac{p_\phi}{p_0} = \frac{\epsilon_{0\phi}}{\epsilon_0} \log \frac{p}{p_0} \quad (19)$$

where p_ϕ is the pressure at altitude h and at latitude ϕ , and $\epsilon_{0\phi}$ is the acceleration of gravity at sea level and at latitude ϕ . A similar equation (replacing p 's with ρ 's) applies to densities.

By means of equation (19) it can be shown that a latitude correction factor (L.C.F.) defined by

$$\text{L.C.F.} = \frac{p_\phi}{p} \quad (20)$$

can be computed by

$$\text{L.C.F.} = \left(\frac{p}{p_0} \right)^{\frac{\epsilon_{0\phi} - \epsilon_0}{\epsilon_0}} \quad (21)$$

If values of $\epsilon_{0\phi}$ from reference 17 are used, the following values for the exponent $(\epsilon_{0\phi} - \epsilon_0)/\epsilon_0$ are obtained:

Latitude (deg)	$\frac{\epsilon_{0\phi} - \epsilon_0}{\epsilon_0}$	Latitude (deg)	$\frac{\epsilon_{0\phi} - \epsilon_0}{\epsilon_0}$
0	-2.66758×10^{-3}	50	0.42175×10^{-3}
10	-2.50922	60	1.28372
20	-2.05299	70	1.98732
30	-1.35337	80	2.44701
40	-0.49405	90	2.60670

The foregoing exponents when applied to the values of pressure ratio p/p_0 tabulated in tables IV and V give the values of the latitude correction factor described by equations (20) and (21). For latitudes at increments of 10° and for altitudes at increments of 10 kilometers the latitude correction factors that are applicable to the pressures given in tables IV and V have been computed and are presented in table VI. By means of table VI it is therefore possible to obtain computed values of pressure which take into consideration the variation with latitude of the sea-level value of the acceleration of gravity ϵ_0 . This computation may be made by use of equation (20) which may be written $p_\phi = (\text{L.C.F.})p$.

Coefficient of viscosity and kinematic viscosity. - The Sutherland formula (equation (17)) is strictly applicable only to a gas of constant composition and to pressures which are not too small, and consequently the tabulated values for the coefficient of viscosity and for the kinematic viscosity are obviously not entirely reliable at the higher altitudes. However, the lack of data on the viscosity of oxygen in the atomic form does not permit at this time an estimation of the correction that is needed to allow for the specified dissociation. Furthermore, because of the fact that the effective value of the viscosity of a gas at very low pressure flowing over a body depends on the size and shape of the body, it is not practical to give a correction that will be applicable to more than one specific size and shape of a body. The values for viscosity at the higher altitudes should therefore be used with caution.

Speed of sound. - The tabulated values for the speed of sound are believed to be correct for all altitudes covered by the tables.

Caution should be exercised, however, in using the tabulated values for the upper altitudes in connection with Mach numbers because at high altitudes where the mean free paths of the air molecules are large in comparison with the dimensions of the body moving through them, the laws of fluid dynamics do not apply and the laws of particle dynamics must be used. When aerodynamic forces, for example, are computed for these conditions by use of the laws of particle dynamics the most probable speed of the air molecules is found to be the basic quantity rather than the speed of sound.

As in the case of viscosity, the altitude range in which the most probable speed of the air molecules replaces the speed of sound as the basic quantity depends upon the size of the body under consideration. It is consequently not possible to specify a single level at which the molecular speed becomes significant in aerodynamics. For this reason values for the speed of sound are listed to 120 kilometers.

In any case in which the most probable speed of the air molecules c is needed rather than the velocity of sound a it is possible to obtain the value of c from the value of a listed in the tables by use of the appropriate factor obtained from the following tabulation:

Altitude, h		Ratio of the most probable molecular speed to the speed of sound, $\frac{c}{a} = \sqrt{\frac{2}{\gamma}}$	
		Day	Night
(m)	(ft)		
80,000	262,467	1.195	1.195
85,000	278,871	1.189	1.195
90,000	295,275	1.183	1.195
95,000	311,679	1.176	1.195
100,000	328,083	1.170	1.195
105,000	344,487	1.170	1.195
110,000	360,892	1.170	1.187
115,000	377,296	1.170	1.179
120,000	393,700	1.170	1.170

CONCLUDING REMARKS

The fact should be emphasized that the values given in the tables for the upper atmosphere are only tentative and as such may become obsolete after a sufficient number of reliable direct

measurements of certain quantities have been made available. In the meantime these tentative tables should be useful not only in serving as a basis for comparing performance characteristics and estimating limiting values of performance, but also in securing the additional data needed for revising these tentative tables for the upper atmosphere.

Langley Memorial Aeronautical Laboratory
National Advisory Committee for Aeronautics
Langley Field, Va., December 6, 1946

APPENDIX A

VARIATION WITH ALTITUDE OF MOLECULAR WEIGHT

AND RATIO OF SPECIFIC HEATS

Molecular Weight in the Region of Oxygen Dissociation

Consider an initial unit volume of normal air composed only of molecular gases, consisting of oxygen and other constituents. Let all the non-oxygen constituents be diatomic of average molecular weight M_N , and let the molecular weight of oxygen in the molecular form be M_m , and in the atomic form M_a . Then

$$M_a = \frac{1}{2}M_m \quad (A1)$$

Let the initial conditions be as follows:

v_0 volume of all-molecular oxygen at height h_m

$1 - v_0$ volume of non-oxygen components at height h_m

M_0 average molecular weight of the initial air mixture at height h_m

Then

$$M_0 = v_0 M_m + (1 - v_0) M_N \quad (A2)$$

At height h , between h_m and h_a (where h_m is height at base of region in which dissociation occurs, and h_a is height at top of the region, and where all the oxygen is in the atomic form) the volume of molecular oxygen v_m per unit initial volume of normal air is

$$v_m = v_0 \left(\frac{h_a - h}{h_a - h_m} \right) \quad (A3)$$

and the volume of atomic oxygen v_a per unit initial volume of normal air is

$$v_a = 2v_0 \left(\frac{h - h_m}{h_a - h_m} \right) \quad (A4)$$

Therefore, the average molecular weight M of the atmosphere at height h can be shown to be

$$M = \frac{M_0}{1 - K(h - h_m)} \quad (A5)$$

where

$$K = - \frac{v_0}{h_a - h_m} \quad (A6)$$

the volume gradient of molecular oxygen, $\Delta v/\Delta h$.

Ratio of Specific Heats in the Region of Oxygen Dissociation

The ratio of specific heats γ for diatomic gases is taken to be $7/5$ and for monatomic gases, $5/3$. If the ratio of the specific heats γ for the atmosphere is assumed to be given by a weighted average, according to relative masses, of the values of γ for diatomic and monatomic gases, it can be shown, by using equations (A1), (A2), (A3), and (A4) that for those regions of the atmosphere in which dissociation of oxygen occurs

$$\gamma = \frac{7}{5} + \frac{4}{15} v_0 \left(\frac{M_m}{M_0} \right) \left(\frac{h - h_m}{h_a - h_m} \right) \quad (A7)$$

The standard value for γ_0 , for the atmosphere at sea level, is $7/5$, and for M_m the standard value is 32. Therefore

$$\frac{\gamma}{\gamma_0} = 1 - \frac{128K(h - h_m)}{21M_0} \quad (A8)$$

It is estimated that in the tentative standard atmosphere the variation of γ due to pressure and temperature effects is only about 0.6 of 1 percent. For this reason the effect of pressure and temperature upon γ is ignored in computing these tentative tables.

APPENDIX B

VARIATION OF PRESSURE WITH ALTITUDE (ASSUMING THE
ACCELERATION OF GRAVITY IS A CONSTANT g_0)

The equations relating atmospheric pressure to height for all altitude ranges in all three atmospheres (minimum, standard, and maximum temperatures) are only four in number. These four equations represent all possible combinations of the two types of temperature-height relationship and the two types of composition-height relationship. The deductions of the equations are based upon the familiar hydrostatic relation

$$dp = - g_0 \rho \, dh \quad (B1)$$

and upon the general gas equation

$$\frac{\rho}{\rho_0} = \frac{p}{p_0} \frac{M}{M_0} \frac{T_0}{T} \quad (B2)$$

These two equations, when combined, give

$$\frac{dp}{p} = - \frac{g_0 \rho_0 T_0 M \, dh}{p_0 T M_0} \quad (B3)$$

The differential equation (B3) is then used for deriving algebraic equations for pressure as a function of altitude, for each of the four combinations of temperature-height and composition-height relationships previously discussed. The derivations are indicated in the following paragraphs and the resulting equations are used in the preparation of tables II and III.

Combination A (constant temperature and constant composition).

The type of atmosphere in which both the temperature and composition are constant may be represented algebraically by

$$T = \text{Constant}$$

and

$$M = \text{Constant}$$

Equation (B3) when integrated between the limits of height h_A and height h then becomes,

$$\log_e \left(\frac{P}{P_A} \right) = \frac{-\xi_0 \rho_0 T_0^M}{P_0 T_0^M} (h - h_A) \quad (B4)$$

where h_A is the base of the region in which type A conditions prevail.

Combination B (constant temperature gradient and constant composition). - For the type of atmosphere having a constant temperature gradient and constant composition, let the temperature gradient be represented by

$$L = \text{Constant} = \frac{\Delta T}{\Delta h} \quad (B5)$$

and the temperature by

$$T = T_B + L(h - h_B) \quad (B6)$$

where T_B and h_B are the respective values at the base of the region to which combination B conditions prevail. Also $M = \text{Constant}$. Equation (B3) then becomes

$$\frac{dp}{P} = \left(\frac{-\xi_0 \rho_0 T_0^M}{P_0 T_0^M} \right) \frac{dh}{T_B + L(h - h_B)} \quad (B7)$$

and when integrated between the limits of h_B and h this equation becomes

$$\log \left(\frac{P}{P_B} \right) = - \frac{\xi_0 \rho_0 T_0^M}{P_0 T_0^M} \log \left(\frac{T}{T_B} \right) \quad (B8)$$

Combination C (constant temperature and constant volume gradient of dissociation). - In the type of atmosphere where both the temperature and volume gradient of dissociation are constant

$$T = \text{Constant}$$

and an expression for M as a function of h is derived in appendix A, and it is found to be

$$M = \frac{M_0}{1 - K(h - h_m)} \quad (B9)$$

where K is the volume gradient of molecular oxygen defined by

$$K = \frac{\Delta V}{\Delta h} = \text{Constant} \quad (B10)$$

Using these relationships with equation (B3) gives

$$\frac{dp}{p} = - \frac{\epsilon_0 \rho_0 T_0 dh}{p_0 T [1 - K(h - h_m)]} \quad (B11)$$

Integrating equation (B11) between the limits of h_C and L , where h_C is the height at the base of the region in which type C conditions prevail, gives

$$\log \left(\frac{p}{p_C} \right) = \frac{\epsilon_0 \rho_0 T_0}{p_0 T K} \log \left(\frac{M_C}{M} \right) \quad (B12)$$

Combination D (constant temperature gradient and constant volume gradient of dissociation). - The type of atmosphere having both the temperature gradient and the volume gradient of dissociation constant is referred to as combination D. For this combination, the expression for molecular weight given in equation (B9) and an appropriate modification of equation (B6) give, for equation (B3), the following equation:

$$\frac{dp}{p} = - \frac{\epsilon_0 \rho_0 T_0 dh}{p_0 [1 - K(h - h_m)] [T_D + L(h - h_D)]} \quad (B13)$$

Integrating the variable part of the right-hand member, between the limits of h_D and h , gives

$$\frac{1}{(1 + Kh_m)L + (T_D - Lh_D)K} \log \frac{T_D + L(h - h_D)}{1 - K(h - h_m)} \Bigg|_{h_D}^h$$

Therefore

$$\log \left(\frac{p}{p_D} \right) = \frac{-g_0 p_0 T_0 M_D}{p_0 (M_0 L + M_D K T_D)} \log \left(\frac{T_M}{T_D M_D} \right) \quad (B14)$$

APPENDIX C

VARIATION OF PRESSURE WITH ALTITUDE (ASSUMING THE ACCELERATION
OF GRAVITY VARIES INVERSELY AS THE SQUARE OF THE
DISTANCE FROM THE CENTER OF THE EARTH)

The equations relating pressure and altitude derived herein are based upon the general differential equation derived from equation (B2) of appendix B, from the hydrostatic relation

$$dp = -g\rho dh \quad (C1)$$

and from the equation representing the inverse square variation of the acceleration of gravity

$$g = g_0 \left(\frac{r}{r+h} \right)^2 \quad (C2)$$

This general differential equation is

$$\frac{dp}{p} = \frac{-g_0 \rho_0 T_0 M r^2 dh}{p_0 T M_0 (r+h)^2} \quad (C3)$$

As in appendix B four equations are deduced for use in each of the four possible combinations of specified temperature-altitude and composition-altitude relationships. The resulting algebraic equations are used in the preparation of tables IV and V. The deductions for each combination are indicated in the following paragraphs.

Combination A (constant temperature and constant composition).-

For combination A (constant temperature and constant pressure) the algebraic equation relating pressure and altitude is obtained by integrating equation (C3) between the limits of altitude h_A and h . The result is

$$\log_e \left(\frac{p}{p_A} \right)_g = \frac{-g_0 \rho_0 T_0 M}{p_0 T M_0} \frac{r^2 (h - h_A)}{(r+h)(r+h_A)} \quad (C4)$$

(Note that in this equation and succeeding equations the subscript g is used to indicate values computed with the variation in the acceleration of gravity that is specified by equation (C2).)

Combination B (constant temperature gradient and constant composition).— For combination B (constant temperature gradient and constant composition) the differential equation is obtained by substituting in equation (C3) the value for T given by

$$T = T_B + L(h - h_B) \quad (C5)$$

The differential equation is then

$$\frac{dp}{p} = \frac{-g_0 \rho_0 T_0 M r^2 dh}{p_0 M_0 [T_B + L(h - h_B)] (r + h)^2} \quad (C6)$$

The algebraic equation obtained by integrating equation (C6) between the appropriate limits is

$$\log_e \left(\frac{p}{p_B} \right)_g = C_{B_g} \left[\frac{r(h - h_B)}{(r + h)(r + h_B)} + \frac{rL}{rL + h_B L - T_B} \log_e \frac{(r + h)T_B}{(r + h_B)T} \right] \quad (C7)$$

where

$$C_{B_g} = \frac{g_0 \rho_0 T_0 M}{p_0 M_0 \left[L - \frac{1}{r}(T_B - Lh_B) \right]} \quad (C8)$$

Combination C (constant temperature and constant volume gradient of dissociation).— For combination C (constant temperature and constant volume gradient of dissociation) the differential equation is obtained by substituting in equation (C3) the value of M given by

$$M = \frac{M_0}{1 - K(h - h_m)} \quad (C9)$$

The differential equation is then

$$\frac{dp}{p} = \frac{-g_0 p_0 T_0 r^2 dh}{p_0 T [1 - K(h - h_m)] (r + h)^2} \quad (C10)$$

The algebraic equation obtained by integrating equation (C10) between appropriate limits is

$$\log_e \left(\frac{p}{p_{C_g}} \right) = C_{C_g} \left\{ \left[\frac{K}{K + \frac{1 + Kh_C}{r}} \log_e \frac{M(r + h)}{M_0(r + h_C)} \right] - \frac{r(h_C - h)}{(r + h)(r + h_C)} \right\} \quad (C11)$$

where

$$C_{C_g} = \frac{-g_0 p_0 T_0}{p_0 T \left(K + \frac{1 + Kh_C}{r} \right)} \quad (C12)$$

Combination D (constant temperature gradient and constant volume gradient of dissociation). - For combination D (constant temperature gradient and constant volume gradient of dissociation) the differential equation is obtained by substituting in equation (C3) the values of T and M given by a slightly modified form of equation (C5) and by equation (C9), respectively. The resulting differential equation is then

$$\frac{dp}{p} = \frac{-g_0 p_0 T_0 r^2 dh}{p_0 [T_D + L(h - h_D)] [1 - K(h - h_m)] (r + h)^2} \quad (C13)$$

The algebraic equation obtained by integrating equation (C13) between appropriate limits is

$$\log_e \left(\frac{p}{p_D} \right)_g = C_{Dg} \left[\frac{a(h - h_D)}{(1 + xh)(1 + xh_D)} + \frac{b}{x} \log_e \left(\frac{1 + xh}{1 + xh_D} \right) + \frac{c}{y} \log_e \left(\frac{1 + yh}{1 + yh_D} \right) + \frac{d}{z} \log_e \left(\frac{1 + zh}{1 + zh_D} \right) \right] \quad (C14)$$

where

$$C_{Dg} = \frac{-g_0 \rho_0 T_0}{\rho_0 (T_D - T_{hD}) (1 + Kh_m)} \quad (C15)$$

$$x = \frac{1}{r}$$

$$y = \frac{L}{(T_D - T_{hD})}$$

$$z = \frac{-K}{(1 + kh_m)}$$

$$a = \frac{x^2(x^2 + yz - yx - zx)}{(z - x)^2(y - x)^2}$$

$$\frac{b}{x} = \frac{x(2yz - xy - xz)}{(z - x)^2(y - x)^2}$$

$$\frac{c}{y} = \frac{-y^2}{(y - x)^2(z - y)}$$

$$\frac{d}{z} = \frac{z^2}{(z - x)^2(z - y)}$$

Equations (C4), (C7), (C11), and (C14) were used to compute the pressure ratios at the transition levels only in the tentative standard atmosphere. By dividing these pressure ratios by the pressure ratios at the same transition levels obtained by use of the equations in appendix B based on a constant value for the acceleration of gravity, a conversion factor was obtained for each of the several transition altitudes. Since it was impractical to use these complex equations for directly computing the pressure

ratios at all the levels recorded in tables IV and V, the values at these numerous intermediate levels were arrived at as follows:

(1) For each altitude a value for the conversion factor was computed by algebraic summation from the equation

$$\log_e \left(\frac{p_g}{p} \right) = \frac{\rho_0 T_0}{\rho_0 M_0} \sum_0^h (g_0 - g) \frac{M}{T} \Delta h \quad (C16)$$

where p_g is the pressure based on the variable value of g , and p is the pressure based on a constant value for the acceleration of gravity. In equation (C16) the proper value of g , T , and of M was substituted for each region of the atmosphere, according to equation (C2), (C5), and (C9), respectively.

(2) The values of p_g/p so computed were plotted against altitude to define the shape of the curve relating pressure ratios to altitude.

(3) The accurate values for the pressure ratio computed by equations (C4), (C7), (C11), and (C14) and by equations (B4), (B8), (B12), and (B14) were also plotted and another curve was drawn through these points representing the accurately computed ratios and faired according to the curve drawn through the points obtained by use of equation (C16).

(4) The curve arrived at from step (3) was then used to obtain conversion factors for each of the altitudes recorded in tables IV and V.

APPENDIX D

MOLECULAR MEAN FREE PATHS

Ratio of the Mean Free Paths of Molecules

The conventional equation for the mean free path of the molecules λ of a gas (reference 18) is

$$\lambda = \frac{1}{\pi \sqrt{2} N \sigma^2} \quad (D1)$$

Therefore the ratio of the mean free path at any altitude to the value at sea level is

$$\frac{\lambda}{\lambda_0} = \frac{N_0}{N} \left(\frac{\sigma_0}{\sigma} \right)^2 \quad (D2)$$

But

$$Nm = \rho \quad (D3)$$

and

$$g\rho = \frac{pM}{RT} \quad (D4)$$

Therefore

$$\frac{N_0}{N} = \frac{p_0}{p} \frac{T}{T_0} \frac{g}{g_0} \quad (D5)$$

and

$$\frac{\lambda}{\lambda_0} = \frac{p_0}{p} \frac{T}{T_0} \frac{g}{g_0} \left(\frac{\sigma_0}{\sigma} \right)^2 \quad (D6)$$

For all constituents of the atmosphere except oxygen in the region of dissociation,

$$\sigma = \sigma_0$$

In the absence of available data on the diameter of atoms of oxygen relative to that of molecular oxygen, and in consideration of the fact that the small difference in these two diameters of oxygen has an even smaller effect upon the average diameter of all atmospheric constituents, and for reasons of simplicity it is herein assumed for oxygen also that $\sigma = \sigma_0$. For the purpose of computing these tables therefore equation (D6) is simplified to

$$\frac{\lambda}{\lambda_0} = \frac{P_0 \cdot T}{P \cdot T_0} \frac{g}{g_0} \quad (D7)$$

Furthermore, in those computations that are based on a constant value for the acceleration of gravity

$$g = g_0$$

whence equation (D7) is further simplified to

$$\frac{\lambda}{\lambda_0} = \frac{P_0 \cdot T}{P \cdot T_0} \quad (D8)$$

Mean Free Paths of Molecules at Sea Level

The values of the mean free path of the molecules at sea level given in table I are for nitrogen and oxygen molecules in a normal atmospheric mixture of nitrogen and oxygen. These mean free paths are designated λ_n and λ_o , respectively. A weighted average of the foregoing mean free paths, based upon the relative volumes of nitrogen and oxygen in air is also included and is designated λ_{air} .

The mean free path of the nitrogen molecules in the atmosphere at sea level was computed by the following formula (p. 99 of reference 18):

$$\lambda_n = \frac{1}{\pi \sqrt{2} N_n \sigma_n^2 + \pi N_o \sigma^2 \frac{\sqrt{\bar{c}_n^2 + \bar{c}_o^2}}{\bar{c}_n}}$$

where

N_n number of nitrogen molecules per unit volume of air

N_o number of oxygen molecules per unit volume of air

σ_n diameter of nitrogen molecules

σ_o diameter of oxygen molecules

$\bar{\sigma}$ average diameter of nitrogen and oxygen molecules

\bar{c}_n average speed of nitrogen molecules

\bar{c}_o average speed of oxygen molecules

Similarly, the mean free path of the oxygen molecules at sea level was computed by

$$\lambda_o = \frac{1}{\pi\sqrt{2}N_o\sigma_o^2 + \pi N_n\bar{\sigma}^2 \frac{\sqrt{\bar{c}_n^2 + \bar{c}_o^2}}{\bar{c}_o}} \quad (D9)$$

The values for the average speeds \bar{c}_n and \bar{c}_o were obtained from

the formula $\bar{c} = \sqrt{\frac{3RT}{M}}$. The values for σ were taken from

appendix III, column 4, of reference 18. Values of N_n and N_o , the number of molecules of nitrogen and oxygen, respectively, per unit volume were calculated from the Loschmidt number and the relative volume of the nitrogen and oxygen in air at sea level.

APPENDIX E

VALUES OF CERTAIN CONSTANTS

Tentative Standard Atmosphere at Sea Level

The standard sea-level values for various properties of the atmosphere have been listed in reference 1, and sea-level values for certain other properties are listed in reference 5. Most of these previously listed values are adopted for use in computing the tables herein, but a few changes have been made. The changes are as follows:

Speed of sound.- The values for the speed of sound have been altered slightly to avoid the discrepancy which existed between the values previously listed and the values computed by the conventional equation

$$a_0 = \sqrt{\frac{\gamma_0 p_0}{\rho_0}} \quad (E1)$$

The values for a_0 listed in table I are computed according to equation (E1) by using the appropriate values for γ_0 , p_0 , and ρ_0 that are also listed in table I.

Density.- The values for density in the British engineering system has been changed from 0.002378 to 0.0023779 slugs per cubic foot to avoid discrepancies resulting when computations are based either on the standardized value for specific weight, 1.2255 kilograms per cubic meter (reference 1), or on the derived value for density.

Molecular mean free paths and molecular weight.- In addition to the various quantities previously given in references 1 and 5, the present paper lists molecular mean free paths and the average molecular weight of normal sea-level air. Molecular mean free paths for the nitrogen molecules and oxygen molecules in the normal air mixture have been computed and a weighted average for air has been taken, as described in appendix D. The average molecular weight of normal sea-level air is taken as 28.966 in accordance with reference 19.

Pressure.- The value for pressure in the British engineering system has been changed from 407.1 or 407.2 inches of water at 15° C as used in reference 5 and reference 20, respectively, to 407.15 inches of water at 15° C. This value of 407.15 is the computed value corresponding to 760 millimeters of mercury based on the auxiliary constants and conversion factors listed in the last section of this appendix E.

Table of Sea-Level Values

The values for the various properties of the atmosphere at sea level corresponding to the adopted values for probable minimum and probable maximum temperatures are computed from the values corresponding to standard sea-level temperatures. All three sets of values used in both metric and British engineering systems of units are tabulated in table I. In some instances a quantity is listed in more than one unit; in either the metric or British system.

Auxiliary Constants and Conversion Factors

In addition to the atmospheric properties at sea level given in table I certain other basic constants and conversion factors are used in computing tables II to V. They are

Auxiliary constants:

Density of mercury at 0° C, gm/cm ³	13.5951
Standard acceleration of gravity, g ₀ , cm/sec ²	980.665
Density of water at 15° C, gm/ml	0.9991286
Radius of the earth at 45° latitude and at sea level, m	6,367,623

Conversion factors:

$$1 \text{ lb} = 453.5924 \text{ gm}$$

$$1 \text{ meter} = 3.280833 \text{ ft}$$

$$^{\circ}\text{K} = ^{\circ}\text{C} + 273$$

$$^{\circ}\text{F abs} = ^{\circ}\text{F} + 459.4$$

$$1 \text{ ml} = 1.000027 \text{ cm}^3$$

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TABLE I.—PROPERTIES OF THE ATMOSPHERE AT SEA LEVEL

Quantity	Symbol	Metric engineering system			British engineering system				
		Unit	At probable minimum temperature	At standard temperature	At probable maximum temperature	Unit	At probable minimum temperature	At standard temperature	At probable maximum temperature
Temperature	t_0	$^{\circ}\text{C}$	-48.0	15.0	47.0	$^{\circ}\text{F}$	-54.5	59.0	116.6
Absolute temperature	T_0	$^{\circ}\text{K}$	225.0	288.0	320.0	$^{\circ}\text{F abs.}$	405.0	518.4	576.0
Pressure	p_0	mm Hg at 0°C	760	760	760	in. Hg at 32°F	29.9212	29.9212	29.9212
		kg/m^2	10332.3	10332.3	10332.3	in. water at 15°C	407.15	407.15	407.15
Specific weight	γ_0	dynes/cm^2	1.01325×10^6	1.01325×10^6	1.01325×10^6	lb/ft^2	2116.23	2116.23	2116.23
		kg/m^3	1.5686	1.2255	1.1030	lb/ft^3	0.097928	0.076506	0.068855
Density	$\rho_0 = \frac{\gamma_0}{g_0}$	dynes/cm^3	1.5383	1.2018	1.0816	slugs/ft^3	0.0030437	0.0023779	0.0021401
Coefficient of viscosity	μ_0	$\text{kg}\cdot\text{sec}/\text{m}^2$	0.15995	0.124966	0.11247	$\text{lb}\cdot\text{sec}/\text{ft}^2$	3.0420×10^{-7}	3.7250×10^{-7}	4.0455×10^{-7}
		poise ($\text{dynes}\cdot\text{sec}/\text{cm}^2$)	1.4852×10^{-6}	1.8187×10^{-6}	1.9751×10^{-6}				
Kinematic viscosity	$\nu_0 = \frac{\mu_0}{\rho_0}$	m^2/sec	9.2848×10^{-6}	14.553×10^{-6}	17.561×10^{-6}	ft^2/sec	0.9994×10^{-4}	1.5665×10^{-4}	1.8903×10^{-4}
Speed of sound	a_0	m/sec	300.72	340.22	358.63	ft/sec	986.61	1116.22	1176.60
		km/hr	1082.6	1224.8	1291.1	mph	672.69	761.06	802.23
						knots	584.16	660.90	696.65
Mean free path of nitrogen molecules	λ_n	m	5.76×10^{-8}	7.38×10^{-8}	8.20×10^{-8}	ft	0.1891×10^{-6}	0.2421×10^{-6}	0.2690×10^{-6}
Mean free path of oxygen molecules	λ_o	m	5.75×10^{-8}	7.36×10^{-8}	8.18×10^{-8}	ft	0.1887×10^{-6}	0.2415×10^{-6}	0.2683×10^{-6}
Mean free path of air molecules	λ_{air}	m	5.76×10^{-8}	7.37×10^{-8}	8.19×10^{-8}	ft	0.1890×10^{-6}	0.2419×10^{-6}	0.2688×10^{-6}
Average molecular weight	M_0	----	28.966	28.966	28.966	----	28.966	28.966	28.966
Ratio of specific heats	γ_0	----	1.4	1.4	1.4	----	1.4	1.4	1.4
Relative volume of oxygen	r_0	----	0.2095	0.2095	0.2095	----	0.2095	0.2095	0.2095

TABLES II AND III

PROPERTIES OF THE UPPER ATMOSPHERE
FOR TENTATIVE STANDARD TEMPERATURES
BASED ON AN ARBITRARY CONSTANT VALUE
OF GRAVITATIONAL FORCE

The following set of two tables (tables II and III) constitutes a consistent extension of the standard tables for the lower atmosphere (NACA Rep. No. 218). Consequently, altitudes in this set of tables which correspond to specified ambient-air pressures may be referred to as "tentative pressure altitudes," and those which correspond to a specified ambient-air density may be referred to as "tentative density altitudes" (NACA Rep. No. 474).

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TABLE II. PROPERTIES OF THE UPPER ATMOSPHERE FOR TENTATIVE STANDARD TEMPERATURES BASED ON AN ARBITRARY CONSTANT VALUE OF GRAVITATIONAL FORCE - METRIC ENGINEERING SYSTEM

Table with 10 columns: Altitude, Absolute temperature, Pressure, Pressure ratio, Density, Density ratio, Specific weight, Coefficient of viscosity, Kinematic viscosity, Speed of sound, Mean free path. Data is presented in rows from 20,000m to 80,000m.

TABLE III.—PROPERTIES OF THE UPPER ATMOSPHERE FOR TENTATIVE STANDARD TEMPERATURES BASED ON AN ARBITRARY CONSTANT VALUE OF GRAVITATIONAL FORCE — BRITISH ENGINEERING SYSTEM — Continued

Altitude, h (ft)	Absolute temperature, T (°F abs.)	Pressure, p (lb/ft ²)	Pressure ratio, p/P ₀	Density, ρ (slugs/ft ³)	Density ratio, σ = $\frac{\rho}{\rho_0}$	Specific weight, γ = gρ (lb/ft ³)	Coefficient of viscosity, μ (lb sec/ft ²) (1)	Kinematic viscosity, ν = $\frac{\mu}{\rho}$ (ft ² /sec) (1)	Speed of sound, a (ft/sec)	Mean free path of molecules, λ (ft)
(b) For day only										
262,467	432.0	0.06669	3.151x10 ⁻⁵	89.93x10 ⁻⁹	3.782x10 ⁻⁵	2.893x10 ⁻⁶	3.212x10 ⁻⁷	3.372	1019	6.40x10 ⁻³
264,000	432.0	0.06241	2.949	83.75	3.522	2.695	3.212	3.835	1022	6.84
266,000	432.0	0.05724	2.705	76.33	3.210	2.456	3.212	4.208	1026	7.45
268,000	432.0	0.05257	2.484	69.65	2.929	2.241	3.212	4.612	1030	8.12
270,000	432.0	0.04829	2.282	63.59	2.674	2.046	3.212	5.051	1034	8.83
272,000	432.0	0.04438	2.097	58.07	2.442	1.868	3.212	5.531	1038	9.61
274,309	432.0	0.04081	2.070	57.26	2.408	1.822	3.212	6.010	1038	9.74
276,000	435.4	0.04082	1.929	52.67	2.215	1.692	3.232	6.136	1046	10.5
278,000	439.4	0.03760	1.776	47.77	2.009	1.537	3.257	6.818	1054	11.2
279,000	443.4	0.03469	1.638	43.42	1.826	1.397	3.282	7.559	1063	12.6
280,000	447.4	0.03200	1.512	39.45	1.659	1.269	3.306	8.380	1072	13.8
282,000	451.4	0.02956	1.397	35.91	1.510	1.155	3.331	9.273	1081	15.1
284,000	455.4	0.02736	1.293	32.74	1.377	1.053	3.355	10.25	1089	16.4
286,000	459.4	0.02535	1.198	29.89	1.257	0.9616	3.379	11.30	1098	17.9
288,000	463.4	0.02351	1.111	27.32	1.149	0.8791	3.403	12.46	1107	19.5
290,000	467.4	0.02184	1.032	25.02	1.052	0.8048	3.427	13.70	1116	21.1
292,000	471.4	0.02029	0.9588	22.91	0.9635	0.7371	3.451	15.06	1124	22.9
294,000	475.4	0.01888	0.8920	21.01	0.8837	0.6761	3.475	16.54	1133	24.9
296,000	479.4	0.01759	0.8310	19.30	0.8117	0.6210	3.499	18.13	1142	26.9
298,000	483.4	0.01639	0.7746	17.74	0.7460	0.5707	3.523	19.86	1151	29.1
300,000	487.4	0.01530	0.7228	16.32	0.6865	0.5252	3.546	21.73	1160	31.5
302,000	491.4	0.01428	0.6750	15.03	0.6322	0.4837	3.569	23.75	1169	34.0
304,000	495.4	0.01336	0.6314	13.87	0.5833	0.4463	3.593	25.90	1177	36.6
306,000	499.4	0.01250	0.5908	12.80	0.5384	0.4119	3.616	28.25	1186	39.4
308,000	503.4	0.01170	0.5531	11.82	0.4972	0.3804	3.639	30.79	1195	42.5
310,000	507.4	0.01098	0.5185	10.94	0.4599	0.3519	3.662	33.47	1204	45.7
312,000	511.5	0.01030	0.4866	10.13	0.4259	0.3259	3.685	36.38	1213	49.0
314,000	515.5	0.009671	0.4570	9.386	0.3947	0.3020	3.708	39.51	1222	52.6
316,000	519.5	0.009091	0.4296	8.705	0.3661	0.2801	3.731	42.86	1231	56.4
318,000	523.5	0.008550	0.4040	8.082	0.3399	0.2600	3.754	46.45	1240	60.5
320,000	527.5	0.008050	0.3804	7.509	0.3158	0.2416	3.777	50.30	1249	64.7
322,000	531.4	0.007585	0.3584	6.984	0.2937	0.2247	3.799	54.40	1257	69.2
324,000	535.4	0.007153	0.3380	6.504	0.2735	0.2092	3.822	58.75	1266	73.9
326,000	539.4	0.006744	0.3187	6.054	0.2546	0.1948	3.844	63.50	1275	79.0
328,000	543.5	0.006368	0.3009	5.645	0.2374	0.1816	3.867	68.48	1284	84.3
328,083	543.6	0.006353	0.3002	5.628	0.2367	0.1811	3.867	68.71	1285	84.5
330,000	547.5	0.006012	0.2841	5.244	0.2224	0.1701	3.889	73.54	1289	89.9
332,000	551.5	0.005686	0.2687	4.965	0.2088	0.1597	3.911	78.77	1294	95.8
334,000	555.5	0.005377	0.2541	4.661	0.1960	0.1500	3.933	84.38	1298	102
336,000	559.5	0.005087	0.2404	4.380	0.1842	0.1409	3.955	90.30	1303	109
338,000	563.5	0.004812	0.2274	4.114	0.1730	0.1324	3.977	96.67	1308	116
340,000	567.5	0.004556	0.2153	3.866	0.1626	0.1244	3.999	103.4	1312	123
342,000	571.5	0.004315	0.2039	3.636	0.1529	0.1170	4.021	110.5	1317	131
344,000	575.5	0.004091	0.1933	3.424	0.1440	0.1102	4.043	118.1	1322	139
346,000	579.5	0.003875	0.1831	3.220	0.1354	0.1036	4.065	126.2	1326	148
348,000	583.5	0.003674	0.1736	3.032	0.1275	0.09755	4.086	134.8	1331	157
350,000	587.5	0.003485	0.1647	2.858	0.1202	0.09196	4.108	143.7	1335	167
352,000	591.5	0.003306	0.1562	2.692	0.1132	0.08660	4.129	153.4	1340	177
354,000	595.5	0.003136	0.1482	2.537	0.1067	0.08163	4.151	163.6	1344	188
356,000	599.5	0.002978	0.1407	2.392	0.1006	0.07697	4.172	174.4	1349	199
358,000	603.5	0.002829	0.1337	2.258	0.09495	0.07264	4.193	185.7	1353	211
360,000	607.5	0.002690	0.1271	2.132	0.08967	0.06860	4.214	197.7	1358	223
362,000	611.5	0.002556	0.1208	2.013	0.08466	0.06477	4.236	210.4	1362	236
364,000	615.5	0.002429	0.1148	1.901	0.07994	0.06116	4.257	223.9	1367	250
366,000	619.6	0.002311	0.1092	1.796	0.07554	0.05779	4.278	238.2	1371	265
368,000	623.6	0.002199	0.1039	1.698	0.07142	0.05463	4.299	253.2	1375	280
370,000	627.6	0.002092	0.09887	1.606	0.06753	0.05165	4.319	268.9	1380	296
372,000	631.6	0.001992	0.09411	1.519	0.06387	0.04886	4.340	285.7	1384	313
374,000	635.6	0.001897	0.08952	1.437	0.06044	0.04624	4.361	303.5	1389	331
376,000	639.6	0.001806	0.08536	1.360	0.05720	0.04376	4.382	322.2	1393	350
378,000	643.6	0.001721	0.08153	1.288	0.05416	0.04144	4.402	341.8	1398	370
380,000	647.6	0.001640	0.07751	1.220	0.05130	0.03925	4.423	362.5	1402	390
382,000	651.6	0.001564	0.07330	1.156	0.04861	0.03719	4.443	384.3	1406	411
384,000	655.6	0.001492	0.07049	1.096	0.04609	0.03526	4.464	407.3	1411	434
386,000	659.6	0.001423	0.06724	1.039	0.04369	0.03343	4.484	431.6	1415	458
388,000	663.6	0.001358	0.06417	0.9856	0.04145	0.03171	4.504	457.0	1419	483
390,000	667.6	0.001296	0.06126	0.9352	0.03933	0.03009	4.525	483.9	1423	509
392,000	671.6	0.001237	0.05847	0.8872	0.03731	0.02854	4.545	512.3	1428	536
393,700	675.0	0.001190	0.05624	0.8491	0.03571	0.02732	4.562	537.3	1431	564

¹The values for viscosity listed in these columns are not applicable at the higher altitudes where the mean free paths of the molecules are comparable to or longer than the dimensions of the body being considered. Furthermore, the values listed are based on the conventional Sutherland formula for normal air and, consequently, no allowance has been made for the effect of dissociated oxygen in the atmosphere at the higher levels.

TABLE III.—PROPERTIES OF THE UPPER ATMOSPHERE FOR TENTATIVE STANDARD TEMPERATURES BASED ON AN ARBITRARY CONSTANT VALUE OF GRAVITATIONAL FORCE— BRITISH ENGINEERING SYSTEM— Concluded

Altitude, h (ft)	Absolute temperature, T (°F abs.)	Pressure, p (lb/ft ²)	Pressure ratio, p/p ₀	Density, ρ (slugs/ft ³)	Density ratio, σ = ρ/ρ ₀	Specific Weight, w = gρ (lb/ft ³)	Coefficient of viscosity, μ (lb-sec/ft ²) (1)	Kinematic viscosity, ν = μ/ρ (ft ² /sec) (1)	Speed of sound, a (ft/sec)	Mean free path of molecules, λ (ft)
(c) For night only										
262,467	432.0	0.06669	3.151x10 ⁻⁵	89.93x10 ⁻⁹	3.782x10 ⁻⁵	2.893x10 ⁻⁶	3.212x10 ⁻⁷	3.572	1019	6.40x10 ⁻³
264,000	432.0	0.06239	2.948	84.13	3.598	2.707	3.212	3.818	1019	6.84
266,000	432.0	0.05720	2.703	77.14	3.244	2.482	3.212	4.184	1019	7.46
268,000	432.0	0.05246	2.479	70.74	2.978	2.276	3.212	4.541	1019	8.13
270,000	432.0	0.04810	2.275	64.87	2.728	2.087	3.212	4.951	1019	8.87
272,000	432.0	0.04410	2.084	59.47	2.501	1.913	3.212	5.401	1019	9.67
274,309	432.0	0.04051	1.906	54.67	2.287	1.787	3.212	5.875	1019	9.80
274,000	435.4	0.04044	1.911	54.10	2.275	1.783	3.232	5.974	1023	10.6
276,000	439.4	0.03712	1.754	49.20	2.069	1.543	3.257	6.620	1028	11.7
278,000	443.4	0.03408	1.611	44.80	1.884	1.441	3.282	7.326	1032	12.8
280,000	447.4	0.03134	1.481	40.80	1.716	1.313	3.306	8.103	1037	14.1
282,000	451.4	0.02884	1.363	37.24	1.566	1.198	3.331	8.945	1042	15.5
284,000	455.4	0.02656	1.255	33.98	1.429	1.093	3.355	9.873	1046	16.9
286,000	459.4	0.02446	1.156	31.01	1.304	0.9976	3.379	10.90	1051	18.5
288,000	463.4	0.02256	1.066	28.34	1.192	0.9120	3.403	12.01	1055	20.3
290,000	467.4	0.02081	0.9832	25.92	1.090	0.8339	3.427	13.22	1060	22.2
292,000	471.4	0.01921	0.9079	23.74	0.9984	0.7638	3.451	14.54	1064	24.2
294,000	475.4	0.01774	0.8384	21.74	0.9142	0.6994	3.475	15.96	1069	26.5
296,000	479.4	0.01640	0.7748	19.92	0.8370	0.6410	3.499	17.57	1073	28.9
298,000	483.4	0.01517	0.7168	18.28	0.7687	0.5881	3.523	19.27	1078	31.5
300,000	487.4	0.01406	0.6643	16.80	0.7065	0.5405	3.546	21.11	1082	34.2
302,000	491.4	0.01302	0.6161	15.43	0.6489	0.4964	3.569	23.13	1087	37.3
304,000	495.4	0.01206	0.5699	14.18	0.5963	0.4562	3.593	25.34	1091	40.6
306,000	499.4	0.01119	0.5256	13.05	0.5487	0.4187	3.616	27.71	1096	44.1
308,000	503.4	0.01038	0.4806	12.01	0.5052	0.3865	3.639	30.30	1100	47.9
310,000	507.4	0.009642	0.4356	11.07	0.4654	0.3561	3.662	33.08	1104	52.0
312,000	511.5	0.008958	0.4233	10.20	0.4290	0.3282	3.685	36.13	1109	56.4
314,000	515.5	0.008327	0.3935	9.409	0.3957	0.3027	3.708	39.41	1113	61.1
316,000	519.5	0.007745	0.3660	8.686	0.3653	0.2795	3.731	42.95	1117	66.2
318,000	523.5	0.007210	0.3407	8.023	0.3374	0.2581	3.754	46.79	1122	71.7
320,000	527.5	0.006711	0.3171	7.410	0.3116	0.2384	3.777	50.97	1126	77.6
322,000	531.5	0.006253	0.2955	6.853	0.2882	0.2205	3.799	55.44	1130	83.9
324,000	535.5	0.005826	0.2753	6.337	0.2665	0.2039	3.822	60.31	1134	90.8
326,000	539.5	0.005437	0.2569	5.871	0.2469	0.1889	3.844	65.47	1139	98.0
328,000	543.5	0.005073	0.2397	5.436	0.2286	0.1749	3.867	71.04	1143	106
330,000	547.5	0.004736	0.2238	5.039	0.2119	0.1621	3.889	77.18	1147	114
332,000	551.5	0.004423	0.2090	4.673	0.1965	0.1503	3.911	83.69	1151	123
334,000	555.5	0.004133	0.1953	4.335	0.1823	0.1396	3.933	90.73	1155	133
336,000	559.5	0.003864	0.1826	4.023	0.1692	0.1294	3.955	98.31	1160	143
338,000	563.5	0.003617	0.1709	3.738	0.1572	0.1203	3.977	106.4	1164	154
340,000	567.5	0.003384	0.1599	3.474	0.1461	0.1118	3.999	115.1	1168	166
342,000	571.5	0.003166	0.1496	3.227	0.1357	0.1038	4.021	124.6	1172	178
344,000	575.5	0.002967	0.1402	3.003	0.1263	0.09663	4.043	134.6	1176	192
344,487	576.5	0.002920	0.1380	2.951	0.1241	0.09494	4.048	137.2	1177	195
346,000	579.5	0.002781	0.1314	2.777	0.1168	0.08936	4.065	146.4	1185	206
348,000	583.5	0.002611	0.1234	2.568	0.1080	0.08263	4.085	159.1	1195	221
350,000	587.5	0.002453	0.1159	2.376	0.09992	0.07644	4.108	172.9	1205	237
352,000	591.5	0.002305	0.1089	2.199	0.09248	0.07075	4.129	187.8	1215	253
354,000	595.5	0.002169	0.1025	2.039	0.08575	0.06560	4.151	203.6	1226	271
356,000	599.5	0.002041	0.09646	1.891	0.07951	0.06083	4.172	220.6	1236	290
358,000	603.5	0.001924	0.09090	1.756	0.07383	0.05748	4.193	238.8	1246	310
360,000	607.5	0.001815	0.08576	1.632	0.06864	0.05251	4.214	258.2	1256	331
362,000	611.5	0.001714	0.08098	1.519	0.06388	0.04887	4.236	278.9	1267	353
364,000	615.5	0.001618	0.07648	1.414	0.05947	0.04550	4.257	301.1	1277	376
366,000	619.6	0.001521	0.07235	1.319	0.05546	0.04243	4.278	324.3	1287	400
368,000	623.6	0.001449	0.06848	1.231	0.05175	0.03959	4.299	349.2	1297	425
370,000	627.6	0.001373	0.06488	1.149	0.04834	0.03698	4.319	375.9	1308	451
372,000	631.6	0.001302	0.06151	1.075	0.04520	0.03458	4.340	403.7	1318	479
374,000	635.6	0.001235	0.05836	1.006	0.04229	0.03235	4.361	433.5	1328	508
376,000	639.6	0.001172	0.05537	0.9409	0.03957	0.03027	4.382	465.7	1339	539
378,000	643.6	0.001113	0.05260	0.8817	0.03708	0.02837	4.402	499.3	1349	571
380,000	647.6	0.001058	0.05000	0.8268	0.03477	0.02660	4.423	535.0	1359	604
382,000	651.6	0.001007	0.04757	0.7761	0.03264	0.02497	4.443	572.5	1370	639
384,000	655.6	0.0009582	0.04528	0.7288	0.03065	0.02345	4.464	612.5	1381	676
386,000	659.6	0.0009127	0.04313	0.6851	0.02881	0.02204	4.484	654.5	1391	714
388,000	663.6	0.0008702	0.04112	0.6444	0.02710	0.02073	4.504	698.9	1401	754
390,000	667.6	0.0008296	0.03920	0.6064	0.02550	0.01951	4.525	746.2	1412	795
392,000	671.6	0.0007917	0.03741	0.5712	0.02402	0.01838	4.545	795.7	1422	838
393,700	675.0	0.0007614	0.03598	0.5344	0.02285	0.01748	4.562	839.5	1431	875

The values for viscosity listed in these columns are not applicable at the higher altitudes where the mean free paths of the molecules are comparable to or longer than the dimensions of the body being considered. Furthermore, the values listed are based on the conventional Sutherland formula for normal air and, consequently, no allowance has been made for the effect of dissociated oxygen in the atmosphere at the higher levels.

TABLES IV AND V

PROPERTIES OF THE UPPER ATMOSPHERE
FOR TENTATIVE STANDARD TEMPERATURES
BASED ON AN INVERSE SQUARE VARIATION
OF GRAVITATIONAL FORCE

The following set of two tables (tables IV and V) does not constitute a consistent extension of the standard tables for the lower atmosphere (NACA Rep. No. 218) but takes into account the inverse square law of gravitational attraction and, consequently, the values in these tables are more accurate than those in tables II and III.

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TABLE IV. - PROPERTIES OF THE UPPER ATMOSPHERE FOR STANDARD TEMPERATURES BASED ON AN INVERSE SQUARE VARIATION OF GRAVITATIONAL FORCE - METRIC ENGINEERING SYSTEM

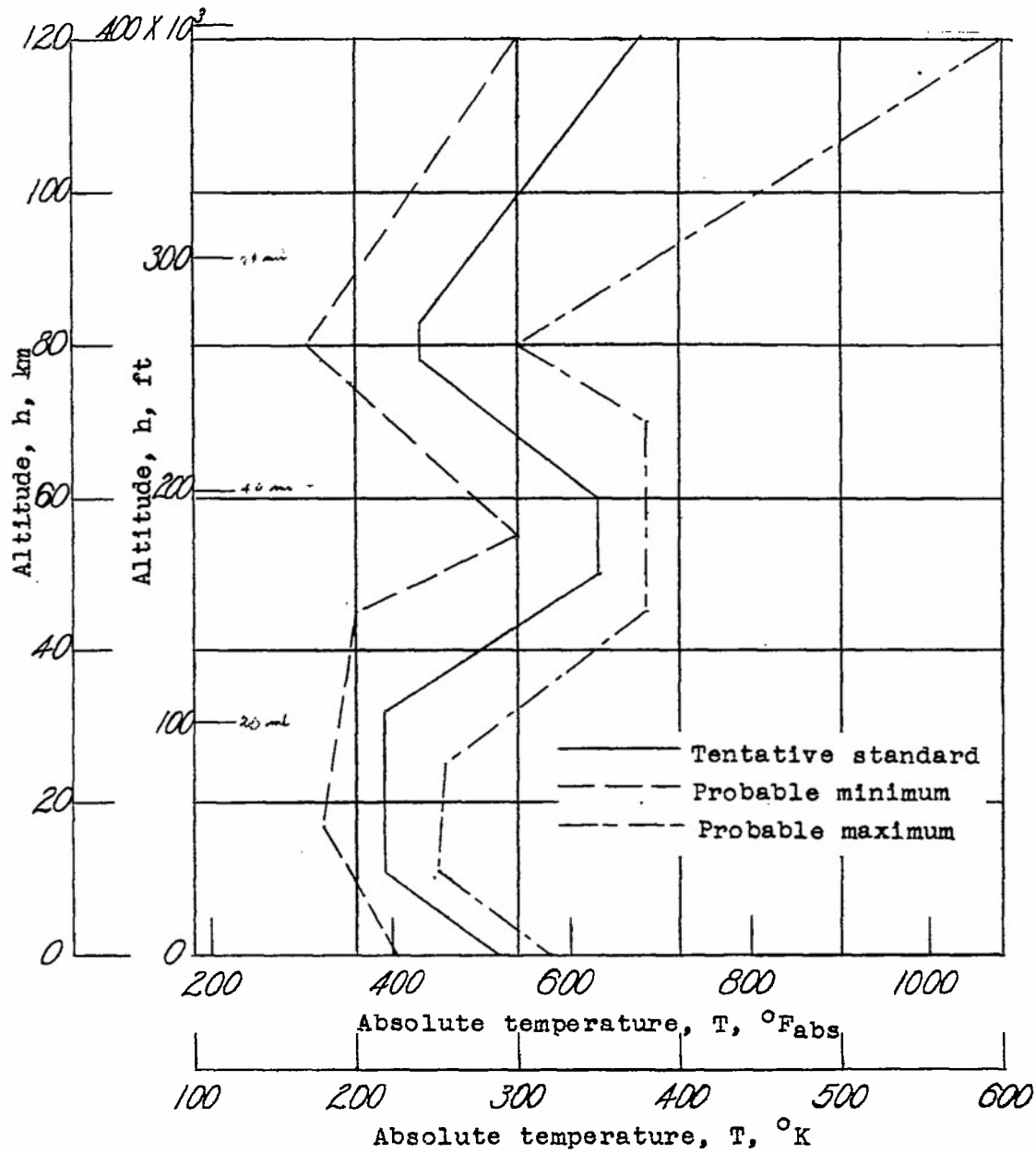
Altitude, h (m)	Relative temperature, T/T_0 (%)	Pressure, (P/P_0) (kg/m^2)	Pressure ratio, P/P_0	Density, (ρ/ρ_0) (kg/m^3)	Density ratio, ρ/ρ_0	Specific weight, (γ/γ_0) (kg/m^3)	Coefficient of viscosity, (μ/μ_0) ($\text{kg}/\text{m}\cdot\text{sec}$)	Kinematic viscosity, $U = \mu/\rho$ (m^2/sec)	Speed of sound, a (m/sec)	Mean Free Path, λ (m)
20,000	218.0	568.4	590.2	802.8	7.68	88.9	1.445	0.0159	295.0	0.0010
21,000	218.0	568.4	590.2	802.8	7.68	88.9	1.445	0.0159	295.0	0.0010
22,000	218.0	568.4	590.2	802.8	7.68	88.9	1.445	0.0159	295.0	0.0010
23,000	218.0	568.4	590.2	802.8	7.68	88.9	1.445	0.0159	295.0	0.0010
24,000	218.0	568.4	590.2	802.8	7.68	88.9	1.445	0.0159	295.0	0.0010
25,000	218.0	568.4	590.2	802.8	7.68	88.9	1.445	0.0159	295.0	0.0010
26,000	218.0	568.4	590.2	802.8	7.68	88.9	1.445	0.0159	295.0	0.0010
27,000	218.0	568.4	590.2	802.8	7.68	88.9	1.445	0.0159	295.0	0.0010
28,000	218.0	568.4	590.2	802.8	7.68	88.9	1.445	0.0159	295.0	0.0010
29,000	218.0	568.4	590.2	802.8	7.68	88.9	1.445	0.0159	295.0	0.0010
30,000	218.0	568.4	590.2	802.8	7.68	88.9	1.445	0.0159	295.0	0.0010
31,000	218.0	568.4	590.2	802.8	7.68	88.9	1.445	0.0159	295.0	0.0010
32,000	218.0	568.4	590.2	802.8	7.68	88.9	1.445	0.0159	295.0	0.0010
33,000	218.0	568.4	590.2	802.8	7.68	88.9	1.445	0.0159	295.0	0.0010
34,000	218.0	568.4	590.2	802.8	7.68	88.9	1.445	0.0159	295.0	0.0010
35,000	218.0	568.4	590.2	802.8	7.68	88.9	1.445	0.0159	295.0	0.0010
36,000	218.0	568.4	590.2	802.8	7.68	88.9	1.445	0.0159	295.0	0.0010
37,000	218.0	568.4	590.2	802.8	7.68	88.9	1.445	0.0159	295.0	0.0010
38,000	218.0	568.4	590.2	802.8	7.68	88.9	1.445	0.0159	295.0	0.0010
39,000	218.0	568.4	590.2	802.8	7.68	88.9	1.445	0.0159	295.0	0.0010
40,000	218.0	568.4	590.2	802.8	7.68	88.9	1.445	0.0159	295.0	0.0010
41,000	218.0	568.4	590.2	802.8	7.68	88.9	1.445	0.0159	295.0	0.0010
42,000	218.0	568.4	590.2	802.8	7.68	88.9	1.445	0.0159	295.0	0.0010
43,000	218.0	568.4	590.2	802.8	7.68	88.9	1.445	0.0159	295.0	0.0010
44,000	218.0	568.4	590.2	802.8	7.68	88.9	1.445	0.0159	295.0	0.0010
45,000	218.0	568.4	590.2	802.8	7.68	88.9	1.445	0.0159	295.0	0.0010
46,000	218.0	568.4	590.2	802.8	7.68	88.9	1.445	0.0159	295.0	0.0010
47,000	218.0	568.4	590.2	802.8	7.68	88.9	1.445	0.0159	295.0	0.0010
48,000	218.0	568.4	590.2	802.8	7.68	88.9	1.445	0.0159	295.0	0.0010
49,000	218.0	568.4	590.2	802.8	7.68	88.9	1.445	0.0159	295.0	0.0010
50,000	218.0	568.4	590.2	802.8	7.68	88.9	1.445	0.0159	295.0	0.0010

TABLE V... PROPERTIES OF THE UPPER ATMOSPHERE FOR TENTATIVE STANDARD TEMPERATURES BASED ON AN INVERSE SQUARE VARIATION OF GRAVITATIONAL FORCE - BRITISH ENGINEERING SYSTEM

Table with 11 columns: Altitude, Absolute temperature, Pressure, Pressure ratio, Density, Density ratio, Specific weight, Coefficient of viscosity, Kinematic viscosity, Speed of sound, Mean free path of molecules. Includes a note '(a) For both day and night'.

TABLE VI.- LATITUDE CORRECTION FACTORS FOR VALUES OF PRESSURE IN TABLES IV AND V

Latitude, deg Altitude, h		Latitude, deg									
		0	10	20	30	40	50	60	70	80	90
(km)	(ft)										
(a) For both day and night											
20	65,617	1.0078	1.0073	1.0060	1.0039	1.0014	0.9988	0.9963	0.9943	0.9929	0.9925
30	98,425	1.0120	1.0112	1.0092	1.0060	1.0022	.9981	.9943	.9912	.9892	.9885
40	131,233	1.0158	1.0148	1.0121	1.0080	1.0029	.9975	.9925	.9884	.9858	.9848
50	164,042	1.0187	1.0176	1.0144	1.0094	1.0034	.9971	.9911	.9863	.9832	.9821
60	196,850	1.0213	1.0200	1.0154	1.0108	1.0039	.9967	.9899	.9844	.9808	.9796
70	229,658	1.0242	1.0227	1.0186	1.0122	1.0044	.9962	.9886	.9824	.9783	.9769
80	262,467	1.0278	1.0260	1.0212	1.0140	1.0051	.9957	.9869	.9798	.9752	.9736
(b) For day only											
80	262,467	1.0278	1.0260	1.0212	1.0140	1.0051	0.9957	0.9869	0.9798	0.9752	0.9736
90	295,275	1.0312	1.0293	1.0239	1.0157	1.0057	.9952	.9853	.9774	.9722	.9704
100	328,083	1.0340	1.0319	1.0261	1.0171	1.0062	.9947	.9840	.9754	.9698	.9679
110	360,892	1.0364	1.0342	1.0279	1.0183	1.0066	.9944	.9830	.9738	.9678	.9657
120	393,700	1.0385	1.0361	1.0295	1.0193	1.0070	.9940	.9820	.9723	.9660	.9638
(c) For night only											
80	262,467	1.0278	1.0260	1.0212	1.0140	1.0051	0.9957	0.9869	0.9798	0.9752	0.9736
90	295,275	1.0314	1.0295	1.0241	1.0158	1.0057	.9951	.9852	.9772	.9721	.9703
100	328,083	1.0346	1.0325	1.0265	1.0174	1.0063	.9946	.9838	.9750	.9693	.9673
110	360,892	1.0374	1.0352	1.0287	1.0188	1.0068	.9942	.9825	.9730	.9669	.9647
120	393,700	1.0397	1.0373	1.0304	1.0199	1.0072	.9938	.9815	.9714	.9649	.9627



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Figure 1.- Variation of ambient temperature with altitude.

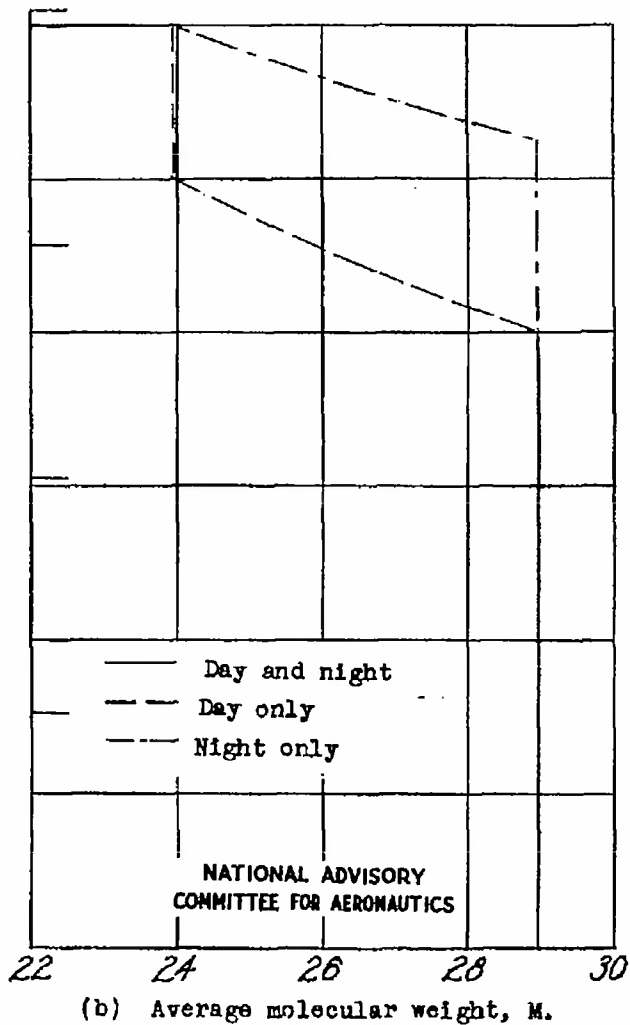
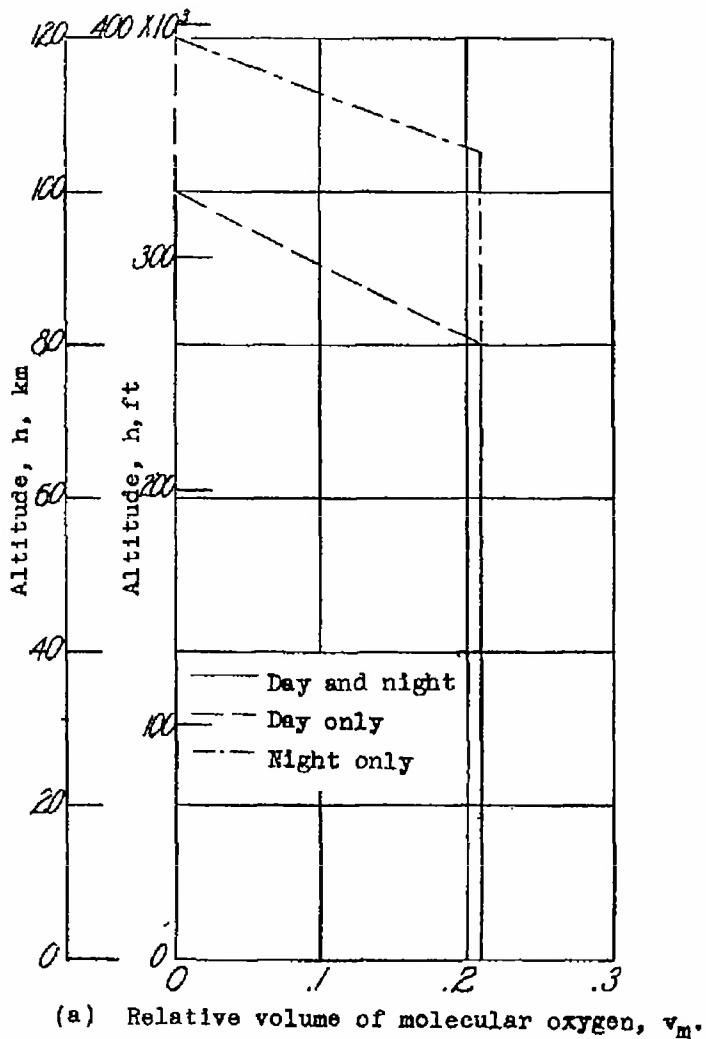
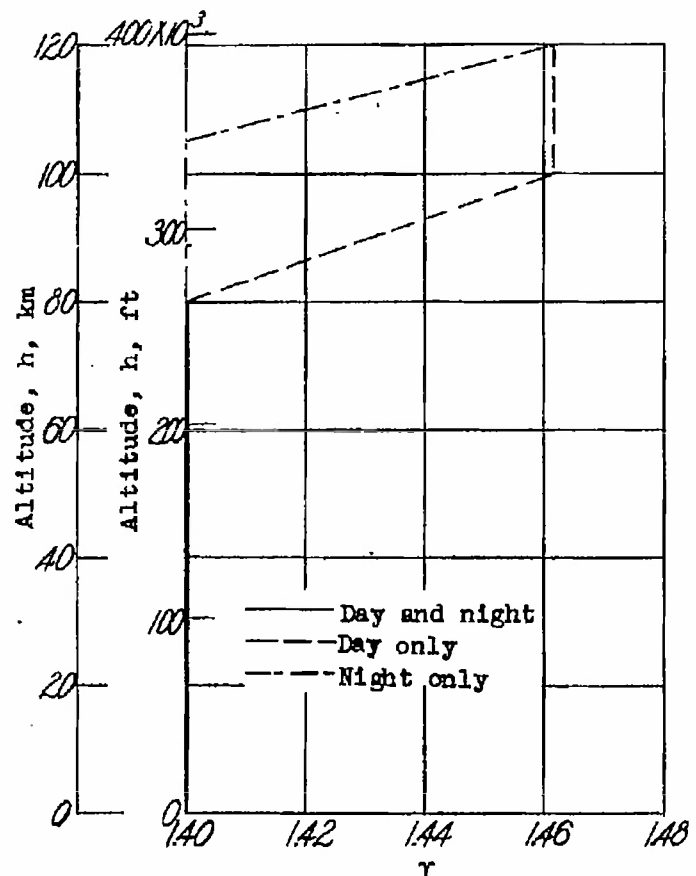
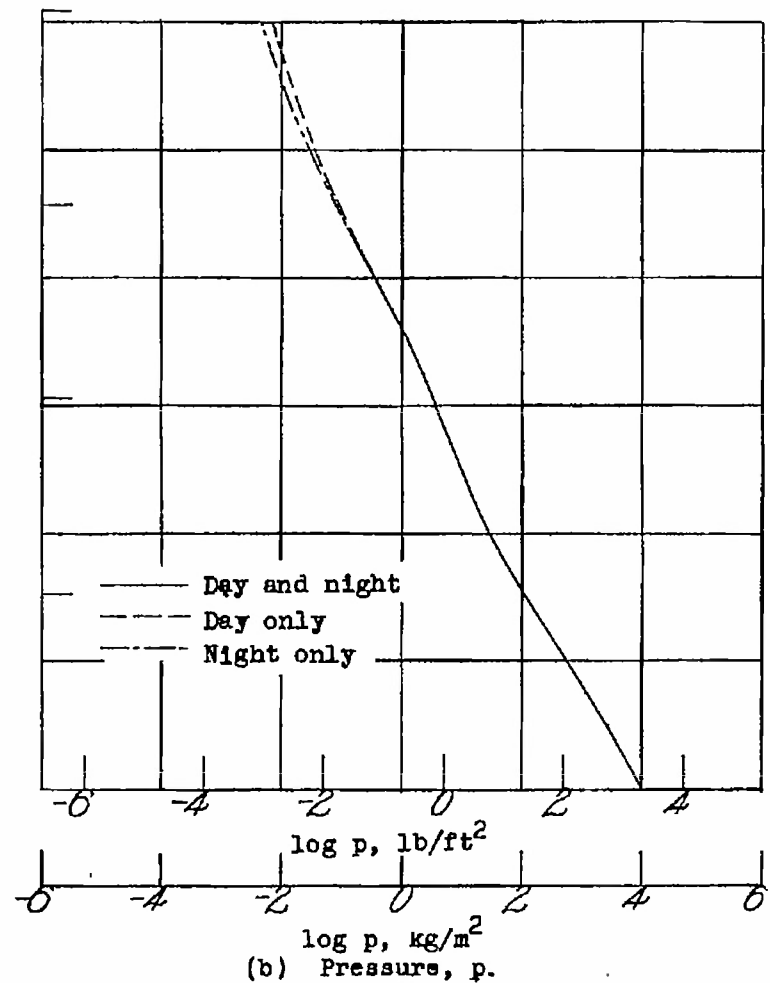


Figure 2.- Variation of composition of the tentative standard atmosphere with altitude. (The dissociation of oxygen is the only change in composition occurring in the tentative standard atmosphere.)



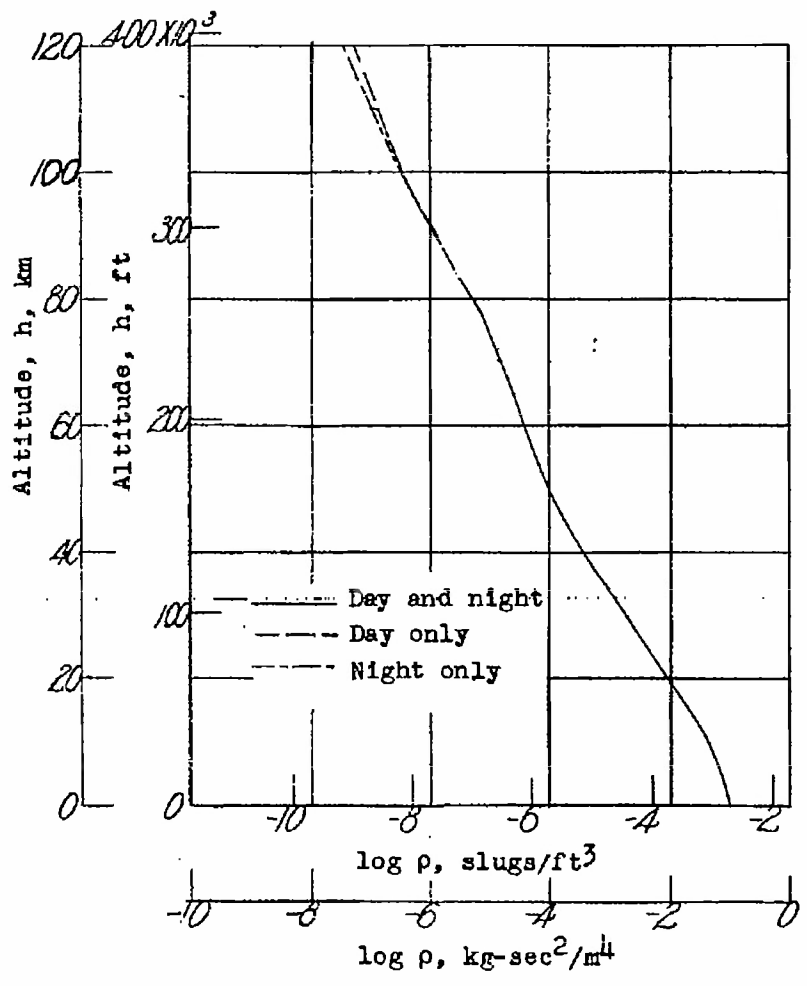
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(a) Ratio of specific heats, γ .

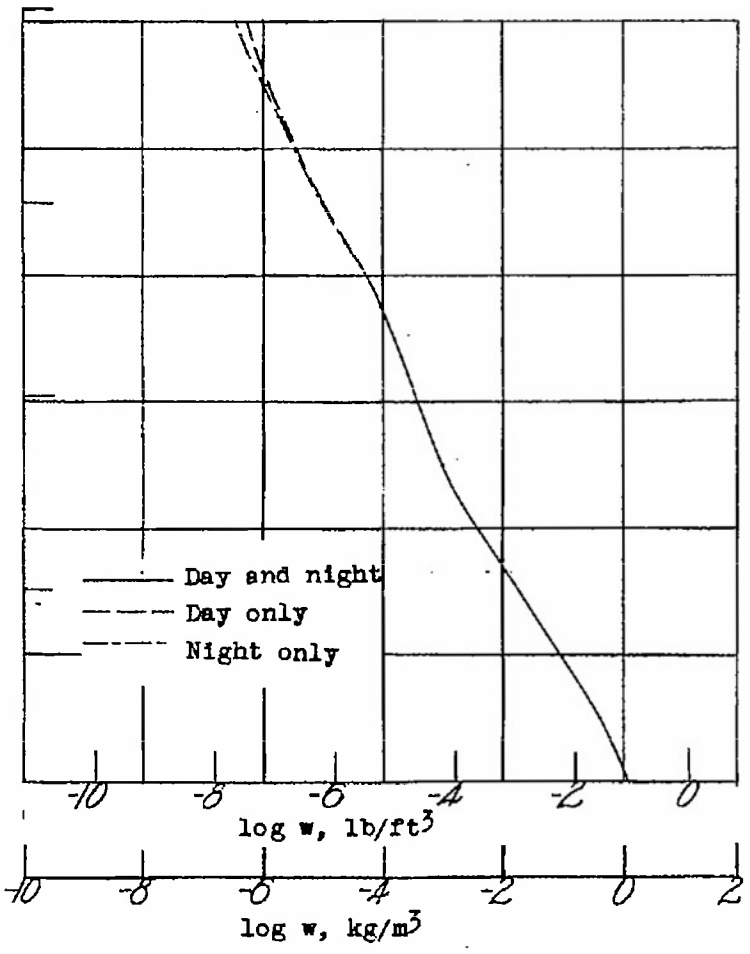


(b) Pressure, p .

Figure 3. Variation with altitude of the physical properties of the tentative standard atmosphere.



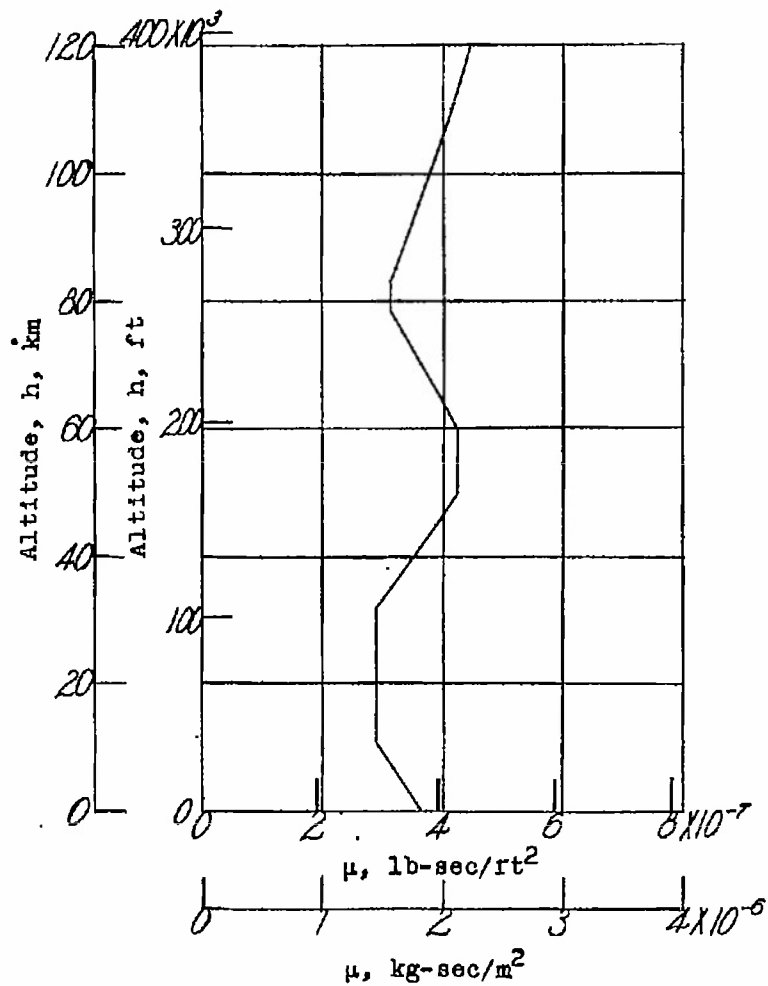
(c) Density, ρ .



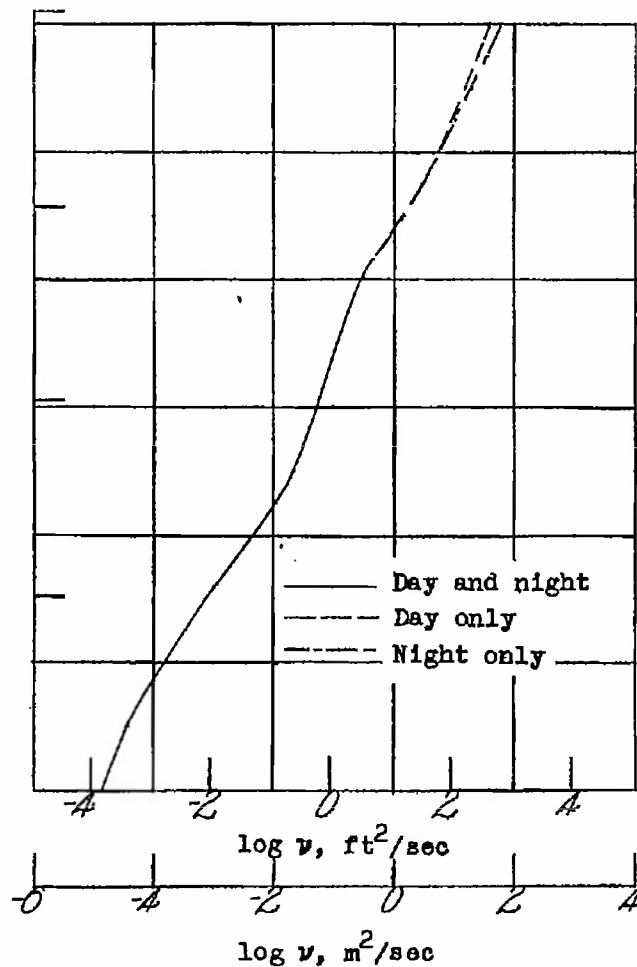
(d) Specific weight, w .

Figure 3.- Continued.

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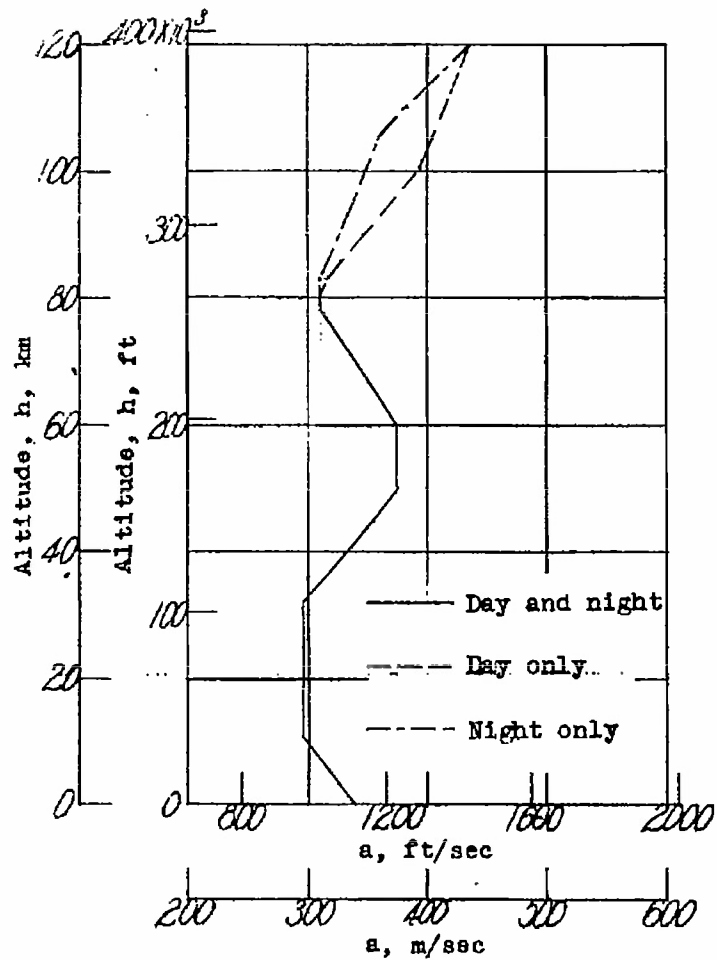
(e) Coefficient of viscosity, μ .



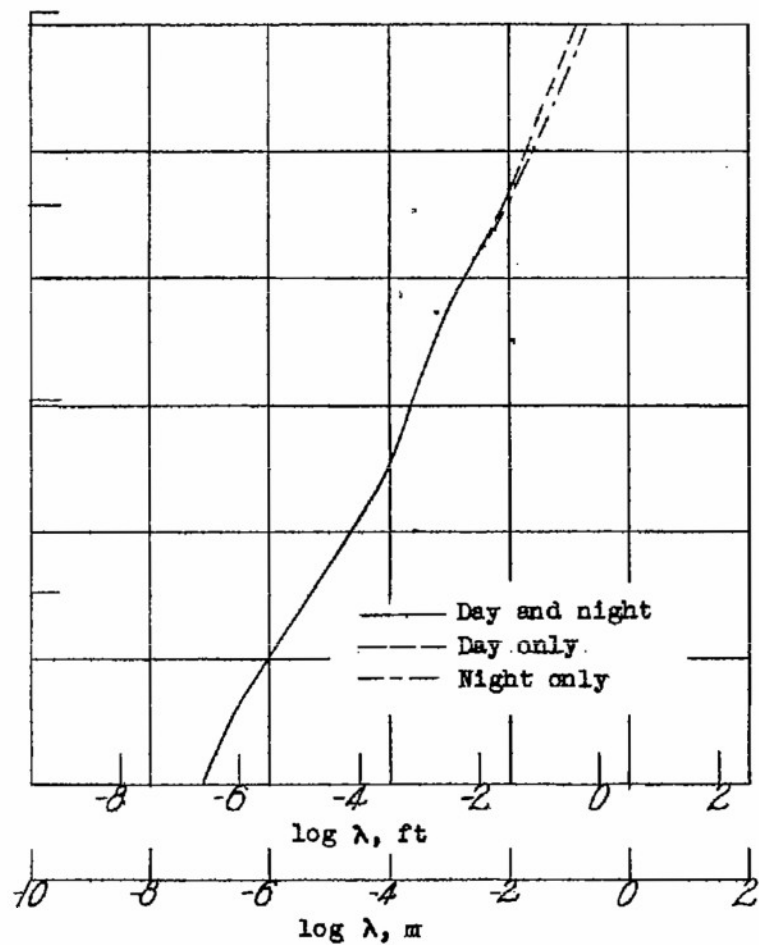
(f) Kinematic viscosity, ν .

Figure 3.- Continued.

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(g) Speed of sound, a .



(h) Mean free path of molecules, λ .

Figure 3.- Concluded.

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ABSTRACT:

Two sets of tables based upon tentative standard specifications for the upper atmosphere are presented. One set constitutes a consistent extension of the standard tables for the lower atmosphere. The other set takes into consideration the decrease in the acceleration of gravity with increasing altitude and, therefore, is more precise than the first set. All quantities listed in the tables against altitude are computed from adopted temperature-height and composition-height relationships.

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