

Modeling Beach Morphology Changes Coupled to Incident Wave Climate and Low Frequency Currents

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LONG-TERM GOAL

My longterm goal is to develop physics-based models to predict nearshore morphology changes, and to test the models against available field data.

OBJECTIVES

The objectives of the present project are to:

1. Incorporate a sediment transport and bed morphology capability in the wave-induced circulation model of Özkan-Haller and Kirby (1997, 1999)
2. Use the resulting model to study the growth to finite amplitude of bottom perturbations on initially longshore-uniform planar or barred beaches
3. Use the model to investigate the evolution of three-dimensional bed features under specific wave conditions, in comparison to ARGUS video results for the Duck FRF site.
4. Begin the development of an instantaneous sediment transport capability (on a wave by wave basis) in the Boussinesq model of Wei et al (1995).

APPROACH

Our approach is to use a robust numerical code for the modeling of 2D or quasi-3D wave-driven nearshore circulation as the basis for computing the local wave-averaged sediment transport rate and the resulting evolution of nearshore morphology. The model results will be limited in accuracy by both the accuracy of input wave information and by the accuracy of the local sediment transport model. By using data from the FRF ARGUS station and related wave information from in situ pressure gage arrays, we seek to determine whether the model will reproduce qualitative shifts in overall bed geometry which have been observed to correlate with shifts in wave conditions (Lippman and Holman, 1990).

WORK COMPLETED

We have incorporated a sediment transport formulation and bed morphology calculation into the nearshore circulation model of Özkan-Haller and Kirby (1997, 1999), and have begun to use the model to study evolution of large amplitude bed forms in river environments as well as the nearshore environments. The study of river environments has proven to be a useful intermediate step, as it has allowed us to test the robustness of the computational model before concentrating on the accuracy of the sediment

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transport formulation after waves are included. River environments provide striking examples of large amplitude bed features, including alternating bar configurations with strong shock-like structures. We have begun an evaluation of the long-term behavior of various unstable bedform patterns, whose linear stability analyses have been described in the literature (Falqués et al, 1996; Schielen et al, 1993). Work on extending the model to incorporate a Bagnold-Bailard-Bowen type model for wave-induced transport is underway.

RESULTS

Results to date have concentrated on calculations based on sediment transport modeling for slowly varying unidirectional flows, as pursued in Falqués et al. (1996) for coastal environments, or by Parker (1976), Colombini et al (1987), Schielen et al (1993) and others for river environments. In particular, we have concentrated recently on the evolution of alternate bars in river environments. The growth of these bars as a bed instability has been understood as a linear instability process since Parker (1976), and Colombini et al (1987) showed in a next-higher-order correction that the bar crests could become oblique to the channel sides. However, it has long been understood from observations that alternate bars can grow into very large amplitude structures with very abrupt, down-channel facing depth increases, taking on a shock-like appearance. Examples in Chang et al (1971) serve as a guideline for testing. An example from our own numerical work is illustrated in Figures 1 and 2. Figure 1 illustrates the evolution of a single along-channel wavelength of a nearly-fastest growing linear perturbation in a channel 200 m wide and 5 m deep. In these figures, along-channel is oriented up the page. The initially sinusoidal bars shown at $t = 0$ take on an oblique character and evolve into shock-like structures fairly rapidly, by $t = 9hr$ as illustrated on the rightmost panel. Figure 2 illustrates the continuation of this process at much later time, where the shock structure has equilibrated and is propagating along the channel. Depths in the scour holes immediately downstream of the bar crests can be twice the initial unperturbed channel depth.

Features such as these show up in a wide range of test calculations with a variety of sediment transport formulations and channel geometries, and thus are likely to self-organizing features to some degree. However, the quantitative accuracy of the morphology predictions here is almost certainly dependent on the chosen sediment transport relation as well as the limitation to depth-integrated 2-D hydrodynamics. We are attempting now to evaluate various steady-flow transport formulations by comparing numerical results to measured data. The resulting formulation will then be extended to include wave-induced moments and applied to coastal settings.

IMPACT/APPLICATION

This work is intended to provide a process-based model for predicting nearshore large scale bedform evolution over day to week time scales, as a first step to bridging the gap between short-time physical predictions and longer time (annual and longer) predictions based on parameterized physics. This gap is still immense, and will only be closed with (1) the advent of both faster computers and with (2) the ability to assess the accuracy of local sediment transport formulations applied over reasonably long time scales. We are mainly concerned with the second aspect of this problem.

TRANSITIONS

The work on morphology evolution conducted here will carry over directly into the seabed module work in the NOPP nearshore project. Tuba Özkan-Haller is working very closely with us, with model extensions and developments being shared by all investigators.

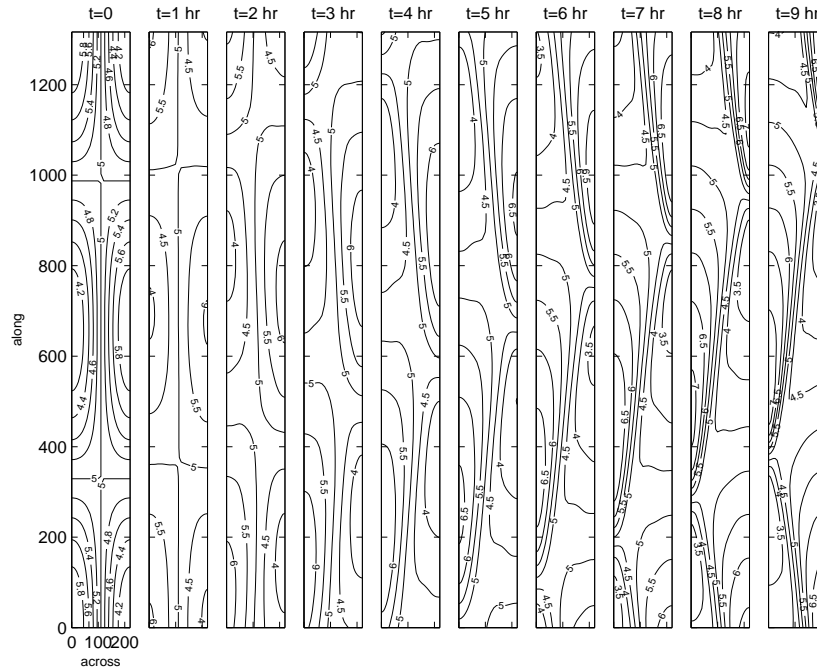


Figure 1: Initial growth of a sinusoidal alternating bar in a sloped channel. The initial bar configuration is based on linear perturbation analysis and has an along-channel length slightly longer than the fastest growing mode. Initial channel depth is 5 meters, and labelled contours are depths in meters.

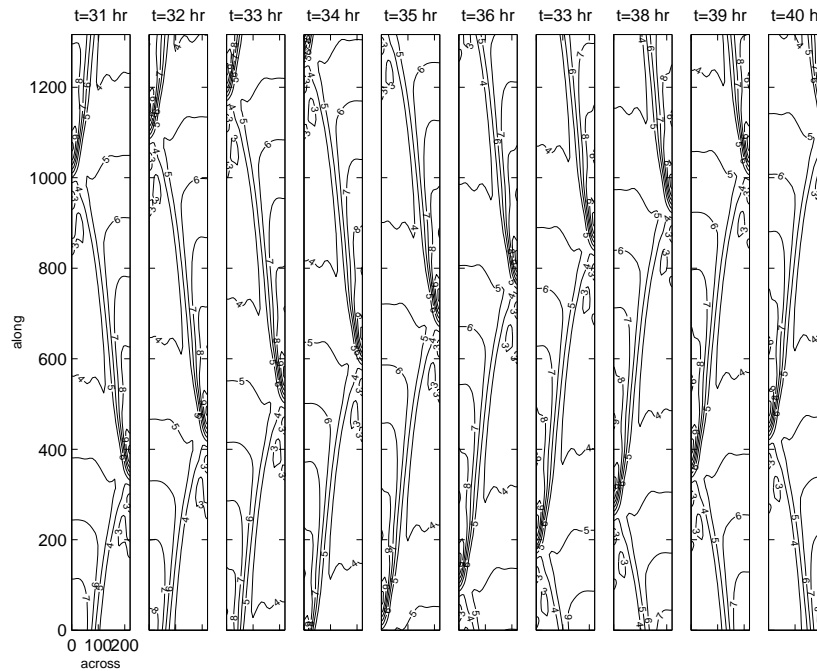


Figure 2: Later stages of alternating bar evolution, showing equilibration and along-slope migration of bars. Note deep scour holes immediately downstream (or up-page) from bar crests, as well as the steadiness of the zig-zagging shock-like structures in the bed.

RELATED PROJECTS

(1) N00014-99-1-0490 "Prediction of the Low Frequency Wave Field on Open Coastal Beaches", H. Tuba Özkan-Haller, Univ. Mich. Tuba is directly involved in this work, and will be in principal charge of transitioning results on the morphology model here to NOPP models. We are also using her work on extensions to the wave driver portion of the hydrodynamic code.

(2) N00014-99-1-1051 (NOPP) "Development and Verification of a Comprehensive Community Model for Physical Processes in the Nearshore Ocean", James T. Kirby et al, Univ. of Del. The work on morphology evolution done here will provide a foundation for work in this area in the NOPP project.

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