

# Defense Threat Reduction Agency 8725 John J. Kingman Road, Stop 6201 Fort Belvoir, VA 22060-6201



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# ECHNICAL REPORT

# Seismicity Characterization and Velocity Structure of Northeast Russia

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May 2005

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<ol> <li>ABSTRACT A seismicity catalog and associated lis</li> </ol>	t of phases for m	any events has been compile	d for northeast Ru	issia using p	ublished and unpublished data from the regional
networks operating in eastern Russia (	primarily Maga	dan, Yakutsk, and Amur), an	d international dat	a files. The	catalog contains about 39,000 events and
150,000 arrival times. The resultant ca	atalog is contami	nated by industrial explosion	s, particularly in the	ne Amur and	central Magadan districts.
The velocities obtained are generally in	was developed by	y obtaining best fit travel curv inferred tectonic regimes. T	hese travel time o	urves were t	g approximately 1300 events in eastern Russia. hen used to relocate 134 larger regional events;
26 are classified as GT10. From these	relocated events	consistent patterns of residu	als, essentially rep	resenting So	ource Specific Station Corrections (SSSCs),
show upper mantle velocities are eleva	ated under the Sil	perian platform and slower be	elow the Sea of Ol	khotsk. Seis	mic station parameters for the region were also
compiled.					
To further improve calibration capabil	ities in northeast	Russia, a small network of d	igital seismic static	ons was depl	oyed in the Magadan region, recording both
earthquakes and ground truth mine bla not sufficient to discriminate mining e	ists. Preliminary	analysis of Lg amplitude rati	os indicate that the	e ratio Lg(4- (<500km) ir	8Hz)Lg(0.75-1.5Hz) using peak amplitudes is
not sufficient to discriminate mining e.	xpiosions nome	artifquares at local and hear i	egional distances	(00011)11	HO GIOLE ICENIA
Travel Time Curves	Seismic St	ation			
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#### **EXECUTIVE SUMMARY**

#### **Overview of Study**

The primary objective of this research was to improve epicentral coordinates to better locate and identify ground truth seismic events in northeast Russia in support of Comprehensive Nuclear-Test-Ban Treaty (CTBT) monitoring. A second objective was to begin the process of explosion discrimination using digital data from stations deployed as a part of this study.

Northeastern Russia (Figure ES-1) remains one of the least researched large continental regions in the world from a seismological standpoint. This is due to geographic remoteness and lack of accessibility, as well as the absence of data available to western scientists prior to the collapse of the Soviet Union. Subsequently, access to the region's scientists and their data has allowed for the first comprehensive studies to be undertaken. Northeast Russia is a tectonically complex region formed by the amalgamation of a large number of terranes of various origins in the Mesozoic and Cenozoic (Nokleberg et al., 1994, Nokleberg et al., 2000). Thus, the crustal structure and seismic velocities are expected to vary considerably within the region.

The first step in this project was to construct computerized files of all known seismicity for the region, as well as the associated phase arrival times. Section One, The Northeastern Russia Seismicity Database, outlines the sources, data acquired, and discusses the intricacies and inconsistences involved in building this database. Assembly of this database would not have been possible without a direct knowledge of the operational procedures used in the region. The assembled seismicity data can then be used to better evaluate the region as a whole, rather than as isolated regions.

One of the problems encountered in the seismicity database was that of explosion contamination, discussed in Section Two, Explosion Contamination in the Northeast Russia Seismicity Catalog. Some regions in northeastern Russia show high levels of seismic activity, when in fact the reported events are almost entirely of anthropogenic origin. This can easily result in erroneous interpretation of seismicity distribution, seismic hazards, etc. On the other hand, known explosions are helpful for developing better travel time curves which can be used to improve epicenter locations.

Section Three, Crustal Model and Relocations of Northeast Russia Earthquakes, develops crustal velocities specifically calibrated to small regions. The calibrated travel time curves were developed in conjunction with relocating approximately 1,300 of the larger events throughout the region. Locations of events determined here are compared with the local Russian network determinations. Crustal thicknesses were determined by inverting Pg arrival times, solving for the best-fit velocities, and analyzing the Pg/Pn crossover points for 27 individual stations. The seismic velocities and crustal thickness determinations are combined into a crustal model for the Magadan region and northern Yakutia. An attempt was made to further refine the crustal model through tomography, however software problems precluded the development of a stable model.

Section Four, GT Classifications and Analysis of Teleseismic Events, carefully evaluates teleseismically recorded events to which a significant amount of previously unavailable local data were added. Results are compared to epicenters calculated by international agencies (e.g., the International Seismological Centre, ISC). Using the relocated teleseismic events, residuals essentially representing Source Specific Station Corrections (SSSCs) were mapped for arrivals at several stations.

Section Five, Digital Station Deployments, discusses operational aspects of digital stations deployed as a part of this study. Deployed stations were used to record earthquakes and obtain

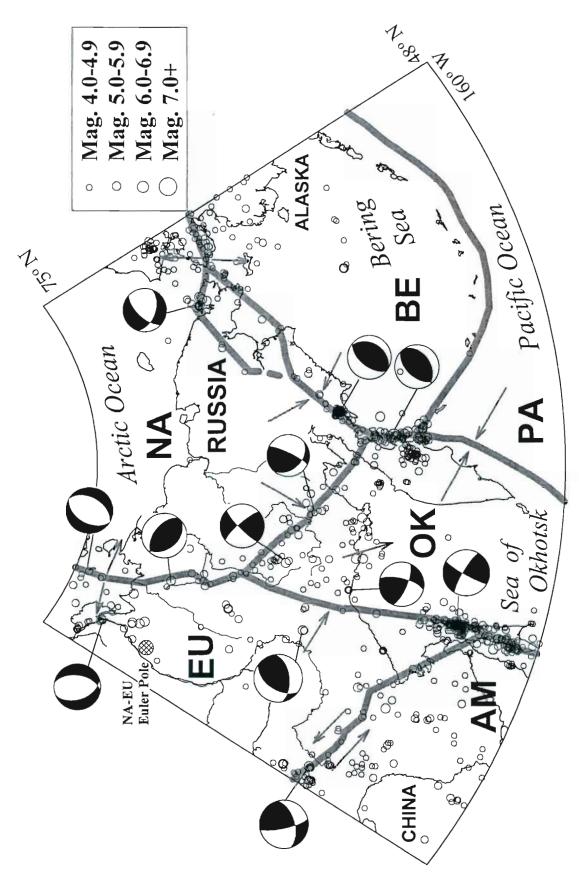


Figure ES-1. Plate tectonic map of northeast Russia with teleseismic earthquakes, representative focal mechanisms, and relative plate motions. Plates are North American (NA), Eurasian (EU), Amur (AM), Okhotsk (OK), Pacific (PA), and Bering (BE).

ground truth information on explosions in the study region. Waveforms of typical earthquakes and explosions recorded are used to investigate the possibility of discrimination between anthropogenic and tectonic sources for improved CTBT monitoring capability.

#### A Brief Overview of the Tectonic Regime and Geologic History

The tectonics of Northeast Russia result from complex plate interactions between the North American (NA), Eurasian (EU), and Pacific (PA) plates and a number of microplates between them, such as Okhotsk (OK), Amur (AM), and Bering (BE; Figure ES-1) that have been previously studied using teleseismic earthquakes (e.g., Chapman and Solomon, 1976; Koz'min, 1984; Fujita et al., 1990a,b; Imaev et al., 1990; Riegel et al., 1993; Seno et al., 1996; Fujita et al., 1997; Mackey et al., 1997; Imaev et al., 2000; Fujita et al., *in press*). The majority of the seismicity occurs along the Chersky Seismic Belt on the North America - Okhotsk boundary and the Stanovoi Seismic Zone on the Eurasia - Amur boundary.

The North American and Eurasian plates are presently converging in northeast Russia, resulting in the southward extrusion of the Okhotsk Block (Figure ES-1; Riegel et al., 1993; see also Seno et al., 1996). The northern portion of the study area includes the Laptev Sea rift system. The Laptev Sea rift system is the extension of the Arctic Mid-Ocean Ridge onto the Siberian continental shelf, expressed as a system of active grabens (Kim, 1986; Fujita et al., 1990a). To the south of the Siberian platform is the Stanovoi seismic zone, along which left-lateral motion is taking place between the Amur and Eurasian plates. Studies of the India-Eurasia collision have suggested extremely far-field extrusion effects, which results in the eastward extrusion of the Amur plate relative to the Eurasian plate (Tapponier et al., 1982; Peltzer and Tapponier, 1988). The Bering block, encompassing the Bering Sea, is driven by subduction of the Pacific Plate and extrusion of southwestern Alaska causing a clockwise rotation relative to the North American plate (Mackey et al., 1997).

The central portion of the study area is composed of a series of exotic terranes and associated island arcs which accreted in the Mesozoic forming the Kolyma-Omolon Superterrane and Kolyma Structural Loop (Figure ES-2; e.g., Parfenov, 1991; Nokleberg et al., 1994; Nokleberg et al., 2000; Layer et al., 2001). The western portion of the study area contains the Verkhoyansk range and eastern portions of the Siberian platform. The Verkhoyansk range is a fold and thrust belt of Mesozoic age associated with the terrane accretion occurring during that time period (Nokleberg et al., 2000). The Siberian platform generally consists of a flat lying Precambrian basement overlain by a few kilometers of Riphean, Cambrian and Jurassic sedimentary materials (Parfenov, 1991). Topographically the Siberian platform is relatively flat. The southern edge of the Siberian platform is separated from exotic terranes to the south by the Hauterivian to Aptian (131-110 Ma) Mongol-Okhotsk Suture (Nokleberg et al., 2000). In Chukotka, in the eastern portion of the study area, is the South Anyui Suture which probably represents the closure of the South Anyui Basin, coincident with the opening of the Canada Basin (Nokleberg et al., 2000). Beyond the South Anyui Suture lies the Chukotka Terrane, the Bering Strait and the terranes of Alaska. Superimposed over the southern edge of the Mesozoic accretionary terranes is the Okhotsk-Chukotka Volcanic Belt, dated at 67-89 Ma (Fujita et al., 1997). Farther east is the Kamchatka-Koryak Accretionary Zone, which consists of a number of Cenozoic accreted terranes (e.g., Stavksy et al., 1990) forming the Kamchatka Peninsula and the Koryak Highlands. The central part of the study area, the Chersky Range, participated in an extensional episode in the Pliocene which resulted in the formation of the Moma rift system (Grachev, 1973; Fujita et al., 1990a). Within the past 0.5 m.y., the pole of rotation for

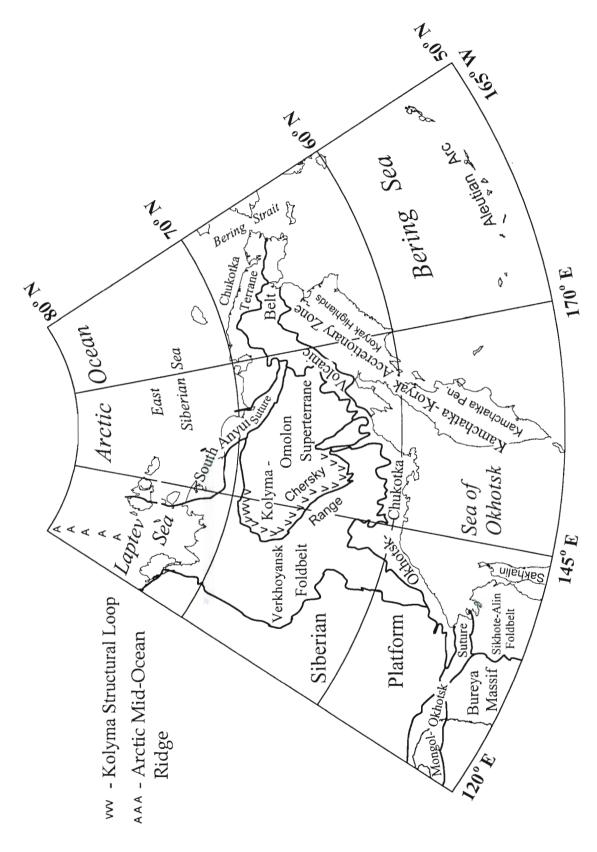


Figure ES-2. Geologic index map of northeast Russia.

the NA-EU plate (Figure ES-1) is suggested to have moved north and extensional activity along the Moma rift ceased (Cook et al., 1986). Presently, the Pacific plate is subducting along the Aleutian Arc and the eastern edge of the Kamchatka peninsula.

#### PREFACE

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#### **CONVERSION TABLE**

Conversion Factors for U.S. Customary to metric (SI) units of measurement.



10 661	BI	DIVIDE
angstrom	1.000 000 x E -10	meters (m)
atmosphere (normal)	1.013 25 x E +2	kilo pascal (kPa)
bar	1.000 000 x E +2	kilo pascal (kPa)
barn	1.000 000 x E -28	meter <sup>2</sup> (m <sup>2</sup> )
British thermal unit (thermochemical)	1.054 350 x E +3	joule (J)
calorie (thermochemical)	4.184 000	joule (J)
cal (thermochemical/cm²)	4.184 000 x E -2	mega joule.m <sup>2</sup> (MJ/m <sup>2</sup> )
curie	3.700 000 x E +1	*giga bacquerel (GBq)
degree (angle)	1.745 329 x E -2	radian (rad)
degree Fahrenheit	$t_k = (t^o f + 459.67)/1.8$	degree kelvin (K)
electron volt	1.602 19 x E -19	joule (J)
erg	1.000 000 x E -7	joule (J)
erg/second	1.000 000 x E -7	watt (W)
f∞t	3.048 000 x E -1	meter (m)
foot-pound-force	1.355 818	joule (J)
gallon (U.S. liquid)	3.785 412 x E -3	meter <sup>3</sup> (m <sup>3</sup> )
inch	2.540 000 x E -2	meter (m)
jerk	1.000 000 x E +9	joule (J)
joule.kilogram (J/kg) radiation dose		
absorbed	1.000 000	**Gray (Gy)
kilotons	4.183	terajoules
kip (1000 lbf)	4.448 222 x E +3	newton (N)
kip/inch² (ksi)	6.894 757 x E +3	kilo pascal (kPa)
ktap	1.000 000 x E +2	newton-second/m2 (N-s/m2)
micron	1.000 000 x E -6	meter (m)
mil	2.540 000 x E -5	meter (m)
mile (international)	1.609 344 x E +3	meter (m)
ounce	2.834 952 x E -2	kilogram (kg)
<pre>pound-force (lbs avoirdupois)</pre>	4.448 222	newton (N)
pound-force inch	1.129 848 x E -1	newton-meter (N-m)
pound-force/inch	1.751 268 x E +2	newton/meter (N/m)
pound-force/foot <sup>2</sup>	4.788 026 x E -2	kilo pascal (kPa)
pound-force/inch2 (psi)	6.894 757	kilo pascal (kPa)
pound-mass (lbm avoirdupois)	4.535 924 x E -1	kilogram (kg)
pound-mass-foot2 (moment if inertia)	4.214 011 x E -2	kilogram-meter² (kg-m²)
pound-mass/foot <sup>3</sup>	1.601 846 x E +1	kilogram-meter³ (kg/m³)
rad (radiation dose absorbed)	1.000 000 x E -2	**Gray (Gy)
roentgen	2.579 760 x E -4	coulomb/kilogram (C/kg)
shake	1.000 000 x E -8	second (s)
slug	1.459 390 x E +1	kilogram (kg)
torr (mm Hg, 0°C)	1.333 22 x E -1	kilo pascal (kPa)

<sup>\*</sup> The bacquerel (Bq) is the SI unit of radioactivity; 1 Bq = 1 event/sec.

<sup>\*\*</sup> The Gray (Gy) is the SI unit of absorbed radiation.

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#### SECTION 1

#### THE NORTHEASTERN RUSSIA SEISMICITY DATABASE

#### 1.1 INTRODUCTION.

Prior to the beginning of this study, there was no comprehensive database covering continental seismicity in northeastern Russia, although seismic networks have been in operation in the region for over 30 years, and significant amounts of data were collected. Areas studied here are the Magadan, Yakutsk, Amur, and Sakhalin networks, as well as minor portions of the Irkutsk and Kamchatka networks. The network boundaries are shown in Figure 1. Unfortunately, data were usually not exchanged between adjoining networks. To better understand the seismicity of northeastern Russia, it is necessary to combine as much data as possible from all sources. The combined data can be re-evaluated to improve travel time curves, hypocenter parameters, tectonic models, better understand seismicity levels, determine regions of anthropogenic sources, etc.

There are two major sections to the database developed: a complete as possible catalog of hypocenters for northeast Russia, and a database of arrival times combined from all sources and networks. In assembly of the database, several published and unpublished sources were used. A discussion of primary sources and networks follows.

This study has compiled a list of approximately 39,000 individual earthquakes throughout the region (Figure 2). Most of the events occur between 1968-2000. For about 14,000 events (Figure 3), phase data have been acquired and incorporated into computerized files; there are presently approximately 150,000 arrival times. For all networks, historic events and data recorded teleseismically are taken from standard sources such as International Seismological Summary (ISS), International Seismological Centre Bulletin (ISC), United States Geological Survey (NEIC), etc., and are not discussed with regard to specific networks.

#### 1.2 DATA SOURCES.

#### 1.2.1 Zemletryseniya V SSSR.

This publication (Zemletryseniya V SSSR:1963-1991; Earthquakes of the USSR; referred to as Zemlet, and Zemletryseniya Severnoi Evrazii: 1992-1994; Earthquakes of Northern Eurasia; also referred to as Zemlet) contains a yearly catalog listing event parameters for the larger earthquakes which occurred within each regional network. In general, only events of K-class 8.5 and larger are listed, although this has varied from year to year, and from network to network. In addition, the cutoff was often raised for large aftershock sequences. Zemlet has historically been available in the U.S.

#### 1.2.2 Materialy po Seismichnosti Sibiri.

The bi-monthly publication (Materialy po Seismichnosti Sibiri; Materials on the Seismicity of Siberia; referred to as Materialy) contains both epicenter lists and phase data for each of the seismic networks in Siberia that investigate seismicity of continental regions for 1970-192 (Irkutsk, Magadan, Yakutsk, Amur (1979-1992), and Altai). The epicenter list provided here is generally complete, although isolated events found in the Far East Bulletin and in the unpublished data seem to be missing. The largest fraction of the assembled epicenter list comes from this Publication. Materialy also contains phase data and arrival times for events equal to or larger than a K-class of

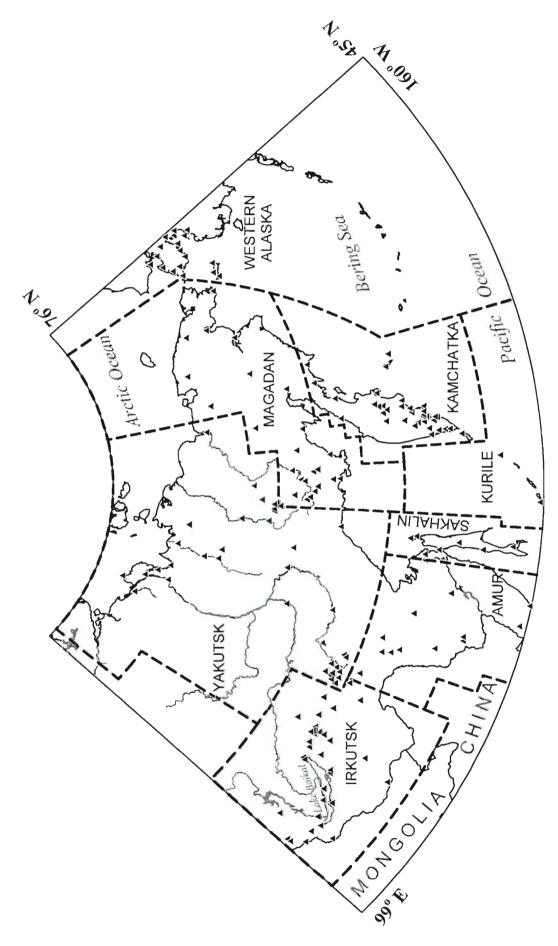


Figure 1. Seismic networks of northeast Russia. Triangles represent seismic stations that have operated at one time, most of which are now closed.

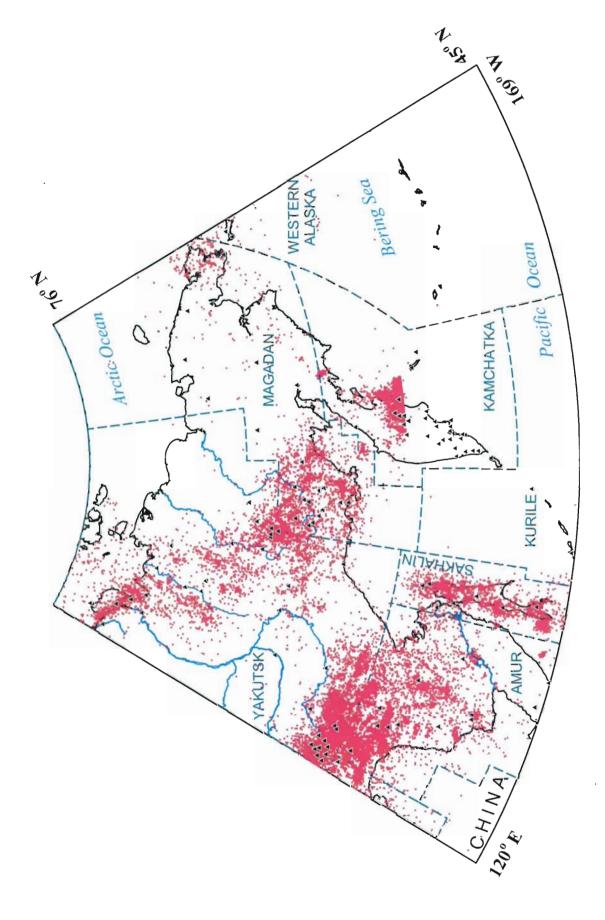


Figure 2. Seismicity of northeast Russia showing approximately 39,000 epicenters. Seismicity related to subduction of the Pacific plate under Kamchatka and the Kurile and Aleutian Islands are omitted. Seismic network boundaries and stations as in Figure 1 are shown for reference.

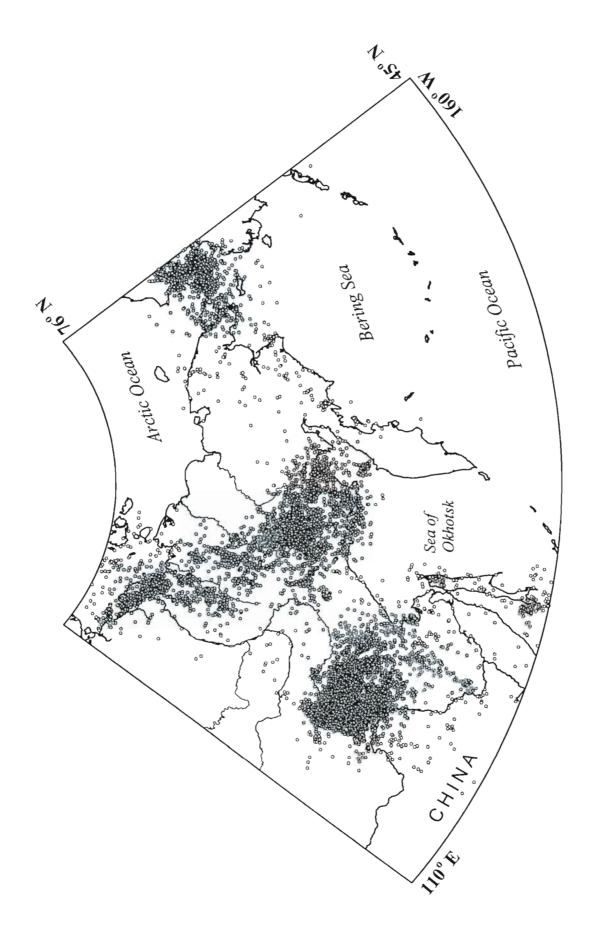


Figure 3. Events in the seismicity catalog for which phase data have been entered into the computerized database. About 14,000 events are depicted.

9.5 occurring within a network. *Materialy* was published in low print quantities with distribution generally to the seismic networks in Siberia.

#### 1.2.3 Seismologicheskii Byulleten' - Dal'nego Vostoka.

This quarterly bulletin (Seismologicheskii Byulleten' - Dal'nego Vostoka; Seismological Bulletin - Far East; referred to as Far East Bulletin) is similar in format to Materialy, and contains both epicenter lists and phase data covering the time period 1972-1990. The Far East Bulletin covers the Magadan (listed as separate networks of "Far East" and "Chukotka"), Amur, Sakhalin, Kamchatka, and Kurile networks. The epicenter list provided here is generally complete for Magadan and Amur, although isolated events found in the unpublished data seem to be missing. For the Kamchatka network, only events of K-class larger than 9.0 are listed. Coverage of phase data varies from network to network. The Far East Bulletin was also published with a low print run.

#### 1.2.4 Unpublished Magadan Network Bulletin.

The epicenter lists and phase data found in the unpublished Magadan Network Bulletin (1977-2001) for located events is identical in content and format to that found in the *Far East Bulletin*. The unpublished bulletins contain much additional material on explosions and unlocated events.

#### 1.2.5 Unpublished Yakutsk Network Draft Material.

Unpublished data from the Yakutsk network has been acquired in varying degrees of completeness (nearly complete for 1982-1997). The material includes epicenters and phase arrival times from all earthquakes within the Yakutsk network. Also included are epicenters and phase arrival times for events that occurred in adjacent networks. Many explosions with associated arrival times are given but not always located. All of the unpublished draft material is handwritten and has had to be made machine readable. For several years of the Yakutsk draft material, a supplement of unpublished Irkutsk network data is included.

#### 1.2.6 Unpublished Kamchatka Seismicity Catalog.

An epicenter catalog was obtained from Petropavlovsk listing all seismicity in the Kamchatka network from 1962 - 1998. The catalog contains approximately 60,000 located earthquakes. A limited amount of unpublished phase data was also obtained from Petropavlovsk.

#### 1.2.7 Seismograms.

Supplemental arrival times were hand-picked from seismograms in Yakutsk and Magadan, as well as develicorder film in Alaska.

#### 1.2.8 Other.

Miscellaneous data was acquired from various other publications. Earthquake data for 1920-1999 were obtained from the International Seismological Centre, the International Seismological Summary, the U.S. Geological Survey (PDE and QED listings), Alaska Earthquake Information Center (AEIC), Kondorskaya and Shebalin (1982), Andreev (1967), Avetisov (1996), Starovoit et al. (1995). Additional information on explosions was obtained from Godzikovskaya (1995).

#### 1.3 SEISMIC NETWORKS AND DATABASE ASSEMBLY.

Assembling the seismicity database of northeast Russia required a considerable effort. A more detailed analysis of the networks can be found in Mackey (1999).

In assembly of the database, it was noted that there are some problems with data supplied to global databases, such as ISC, for events occurring in the northeast Russia. Data reported to the global databases are often preliminary time picks done by station operators, and often have high residuals. This can result from poor time picks, or a secondary Pg phase being reported as a first arriving  $P_n$  phase. The published bulletins of Russian origin (*Materialy* and *Far east Bulletin*) and unpublished data obtained from several of the networks discussed below contains a more complete, and final analysis of the seismograms, done by a professional analyst. This results in discrepancies between data reported in global databases and Russian sources. In such cases, the phase and time from the Russian sources were used, as they were not derived from preliminary seismogram analysis, and were found to contain better time and phase picks.

#### 1.3.1 Northeast Russia Test Network.

In the mid 1960's a number of experimental seismic stations were established throughout northeast Russia to determine background seismicity levels to aid in developing permanent seismic networks (Mishin, 1967). The distribution of seismicity located using the test network (Figure 4) was instrumental in site selection for future seismic stations. Most of the temporary stations were deployed between 6 months and 1 year (Andreev, 1967). Arrival times and phase data for the recorded events are not available. Epicenters are listed in Andreev (1967).

#### 1.3.2 Station Iul'tin and the Magadan EMSD Chukotka Network.

Small events occurring in the Chukotka region were undetected until 1966, when seismic station Iul'tin (Figure 5) opened in Chukotka. Chukotka epicenters reported in Kondorskaya and Shebalin (1982; NCSE) from 1966 to 1974 and the annual *Zemlet* from 1966 to 1982 are single station locations from ILT. Statistical analysis of origin times of an epicenter cluster to the northwest of ILT has shown that this activity is likely due to blasting in gold mining operations around Polyarnyi being misidentified as tectonic events (See Section 2).

Events located to the east and south of Iul'tin define a rough northeast to southwest trend, which can not be associated with any known mining activity. Statistical analysis of origin times for these does not indicate any significant time biases, thus epicenters probably represent tectonic earthquakes. Analysis of some of these earthquakes with reread data from the western Alaska network indicates that the Iul'tin determined epicenters are of reasonable quality given the difficulties of single station locations.

The Chukotka regional network contained six seismic stations which were operated from 1982 to 1993 by the Magadan Experimental Methodological Seismological Division (Figure 6). Arrival times from station ILT supplemented data from the Chukotka network in analysis, and are included in data catalogs. Focal parameters from the Chukotka network were determined using the same procedures and seismic velocities as described below for the Magadan network. Most mining explosions were identified in the location process using time and location by local operators and excluded from earthquake catalogs and bulletins. Unfortunately, in 1993, all seismic stations in Chukotka were closed because of financial constraints. Bilibino (BILL) was reopened as an IRIS GSN station in 1995 and Anadyr (ANYS) was reopened in 1996 as an analog station with assistance from the Chukotka regional government.

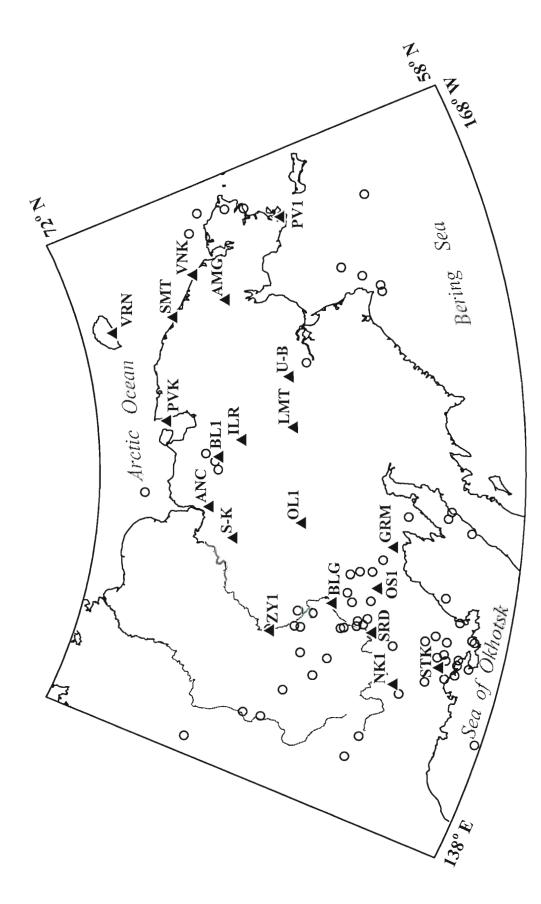


Figure 4. Seismicity and seismic stations of the northeast Russia temporary network (1962-1967).

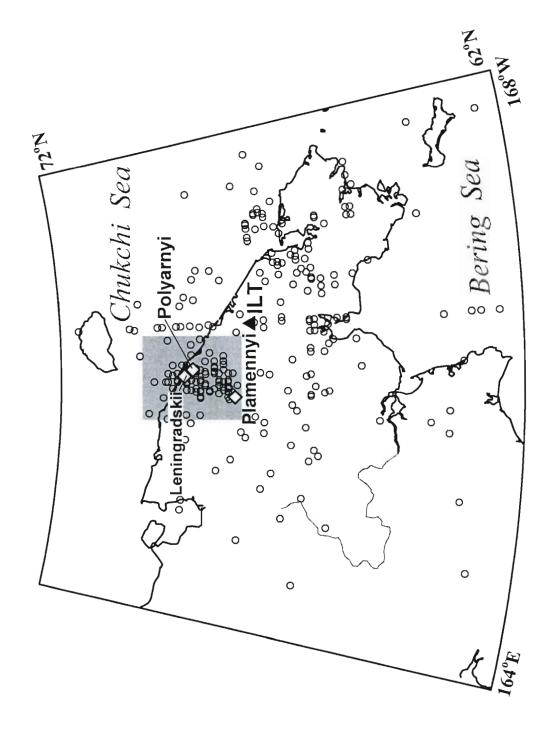


Figure 5. Seismicity located by station Iul'tin (ILT) from 1966-1982. The large cluster of seismicity to the northwest of Iul'tin (grey box) is most likely explosion contamination (see Section 2). Major mine locations are also shown as diamonds.

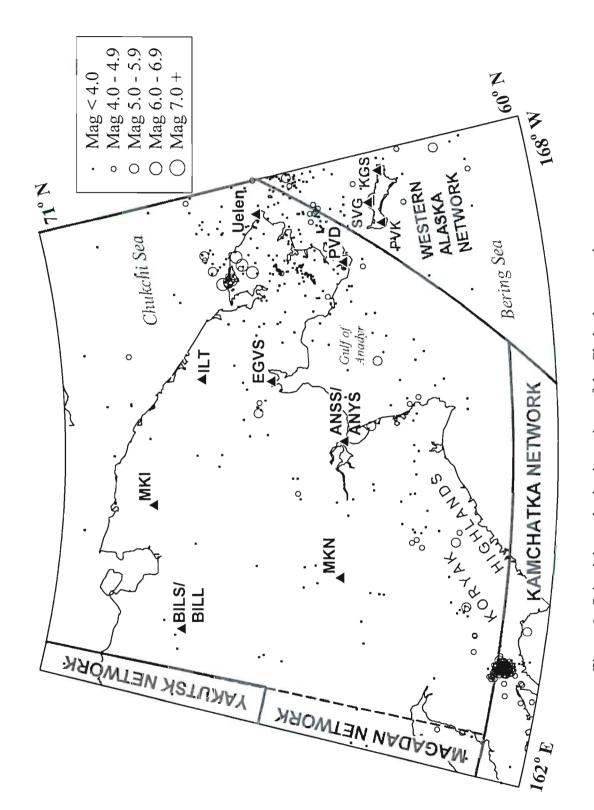


Figure 6. Seismicity and seismic stations of the Chukotka network.

Hypocenter parameters were primarily taken from *Zemlet* (1966-1982) and *Materialy* (1983-1991). Phase data and arrival times for events larger than K-class of 9.5 were taken from *Materialy* (1983-1991). Phase data and arrival time for smaller events are from the unpublished Magadan network bulletins (1983-1991). All data since 1991 is from unpublished network bulletins. For approximately 50 events, phase data and arrival times from the Chukotka network were supplemented with data reread from Alaskan stations.

#### 1.3.3 Magadan Regional Network.

The first station opened in the region was Magadan in 1952. In the late 1960's, following the deployment of the northeast Russia test network, installation of permanent seismic stations began. These stations were initially operated in conjunction with stations in the Yakutsk network. The Magadan Experimental Methodological Seismological Division (EMSD) network was formally established in December 1979, to monitor regional seismicity. The Magadan EMSD has operated a total of 15 permanent seismic stations in the region (Figure 7), excluding those in the Chukotka network, with the largest number of stations open in the late 1980's and early 1990's. Beginning in 1992, economic problems stemming from the collapse of the Soviet Union resulted in drastic cutbacks by the network and the closure of several stations. Today, eight stations are open, with Magadan (MA2) being a Global Seismograph Network (GSN; IRIS) station. Seymchan (SEY) was established as a Geoscope station in the early 1990's. Due to equipment problems, the Geoscope station in Seymchan was essentially abandoned in the mid 1990's. In the summer of 1999, four stations were converted from analog photopaper to 24 bit resolution digital recording with assistance from MSU. This also reestablished digital recording of the abandoned Geoscope instruments in Seymchan. In 2000 and 2001, additional digital stations were opened. See Section 5 below for details on digital station deployments.

There have been three methods of determining earthquake locations in the Magadan network, each using crustal Pg and Sg wave time differences. Prior to 1982, locations were computed primarily by hand and occasionally using a "Besmas 6", an electronic calculating machine. In 1982, they began using computers for locations, which were compared to arc on map determined epicenters. The computer determined epicenters were often "adjusted" to better agree with a paper location if there was a discrepancy between the two. The adjustment procedure was dropped in the early 1990's when the network switched to computer only hypocenter calculations. In the location procedure, the travel-time curves used were derived from the 1959 Magadan-Ust' Srednikan DSS (Deep Seismic Sounding; refraction) profile (Ansimov et al., 1967; Davydova et al., 1968). In the location procedure, no station corrections are used, and station elevation is not considered (Gounbina, pers. comm.).

The unpublished Yakutsk bulletin contains a significant amount of additional data for earthquakes in the Magadan Region. However, this supplemental data is not included in any of the published bulletins containing Magadan network data, as it was not exchanged with the Magadan EMSD. The Magadan network only had access to data from the Yakutsk network station in Ust'Nera.

For the database, hypocenter parameters were primarily taken from *Zemlet* and *Materialy* (1963-1991) and unpublished network bulletins (1991- June 1999). Hypocenters from July 1999 through September 2001 were calculated as a part of this study. Phase data and arrival times for events larger than K-class of 9.5 was taken from *Materialy* (1970-1991). Phase data and arrival time for smaller events are from the *Far East Bulletin* (1972 and 1974) and unpublished Magadan network bulletins (1977-1991). All data from 1992 through June 1999 is from unpublished network bulletins. Data from July 1999 through September 2001 were acquired as a part of this study (see

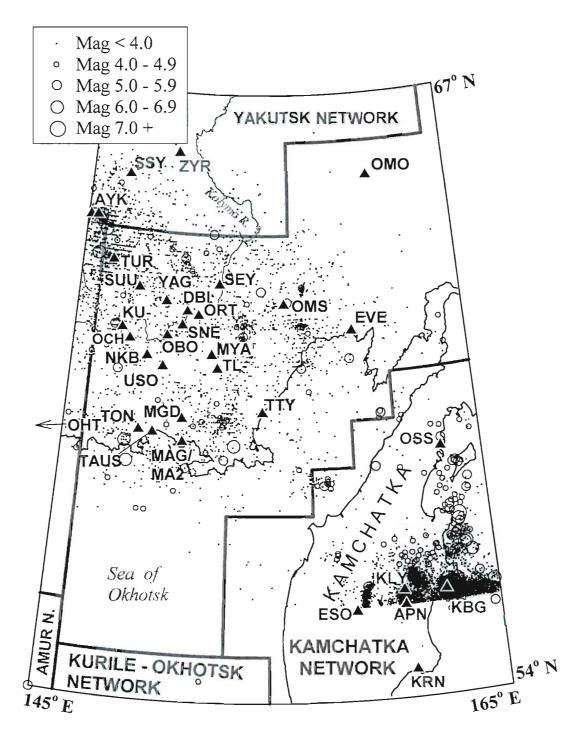


Figure 7. Seismicity and seismic stations of the Magadan network.

Section 6 of this report). All data for Magadan region events listed in the available unpublished Yakutsk bulletin were added to the data files. For several events along the boundary with the Yakutsk network, supplemental arrival times were read from Yakutsk seismograms.

#### 1.3.4 Kamchatka Peninsula Network.

Klyuchi was the first seismic station to open on the Kamchatka peninsula in 1948, followed by Petropavlovsk in 1951. Most remaining stations in the network opened in the early 1960's (Figure 8). Because of subduction of the Pacific plate under the Kamchatka peninsula, this network monitors the most seismically active region in northeast Russia. The Kamchatka network also monitors seismic activity to the north of the subduction zone into the southern portion of the Koryak Highlands, as well as activity on many of the peninsula's active volcanoes. As this study is interested primarily in non-subduction related crustal level events, only earthquakes in the region north of 56° N (the northern edge of the subduction zone) are included. Along the boundary with the Magadan network in the Koryak Highlands and in Shelikhov Bay, many events were independently located by each network, and there seems to have been essentially no exchange of phase data for anything but the largest teleseismically recorded events. Given a choice, the Magadan network hypocenter is usually preferred over that from the Kamchatka network because the Magadan network generally had better station coverage. Hypocenter locations for the Kamchatka north of 56° N were taken from the unpublished Kamchatka seismicity catalog for 1962 to 1998. A limited amount of unpublished phase data from the Kamchatka network for events bounding the Magadan network was also obtained and used in relocations.

#### 1.3.5 Yakutsk Regional Network.

The first station opened in Yakutsk in 1958, with the Yakutsk network being established in the late 1960's to monitor regional seismicity. The Yakutsk network has operated a large number of permanent seismic stations in the region, with the maximum number in the late 1980's and early 1990's. Beginning in 1993, economic problems stemming from the collapse of the Soviet Union resulted in drastic cutbacks by the network and the closure of several stations. In the early 1990's, IRIS stations were opened in Yakutsk and Tiksi. The station in Ust'Nera was moved in 1992, and converted to 24 bit digital acquisition in June 1999 with assistance from MSU.

All earthquake locations determined by the Yakutsk network are by the arcs on map method. For location purposes, the Yakutsk network is broken into northern (Figure 9), and southern regions (Figure 10). The northern region locates earthquakes on a 1:5,000,000 scale map, and includes stations north of 60° N, while the southern region includes stations south of 63° N. Stations used in determining locations are only those which are found within the bounds of the map on which the arcs are drawn. Data from the northern region were not used when locating earthquakes from the southern region, as the northern stations do not fit on the map. Data from the southern region were not used for northern region earthquakes for the same reason.

For the database, hypocenter parameters were primarily taken from *Zemlet* and *Materialy* (1963-1991) and unpublished network bulletins (1991-1997). Phase data and arrival times for events larger than K-class of 9.5 was taken from *Materialy* (1970-1991). Phase data and arrival time for smaller events are from unpublished Yakutsk network bulletins (1982-1997). All data since 1991 is from unpublished network bulletins. As in the case with the Magadan region, there is much data acquired by stations in other networks which was not exchanged with the Yakutsk network, and subsequently not included in any of the published bulletins. Historically, the Yakutsk network had access to data from the Magadan stations of Susuman, Seymchan, Kulu, and Debin (Figure 7), and occasionally others. However, overall, the acquired unpublished Magadan network

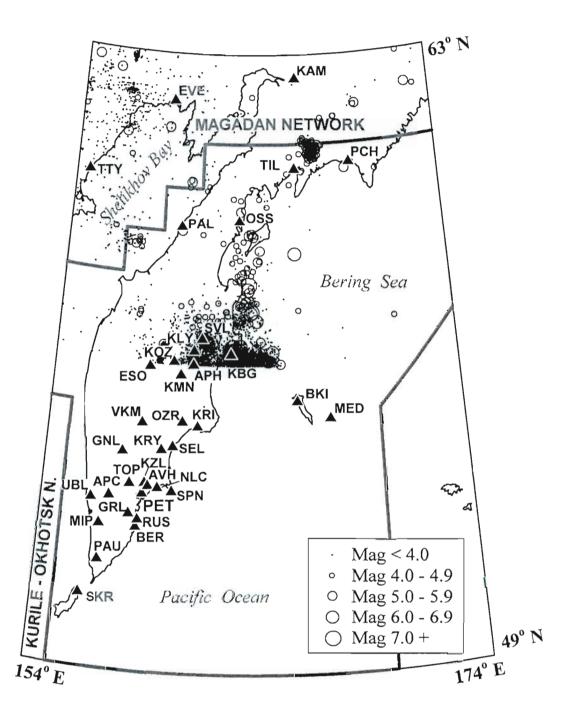


Figure 8. Seismicity and seismic stations of the Kamchatka network. Only seismicity north of 56° north are included. A few stations associated with volcano monitoring are omitted for clarity.

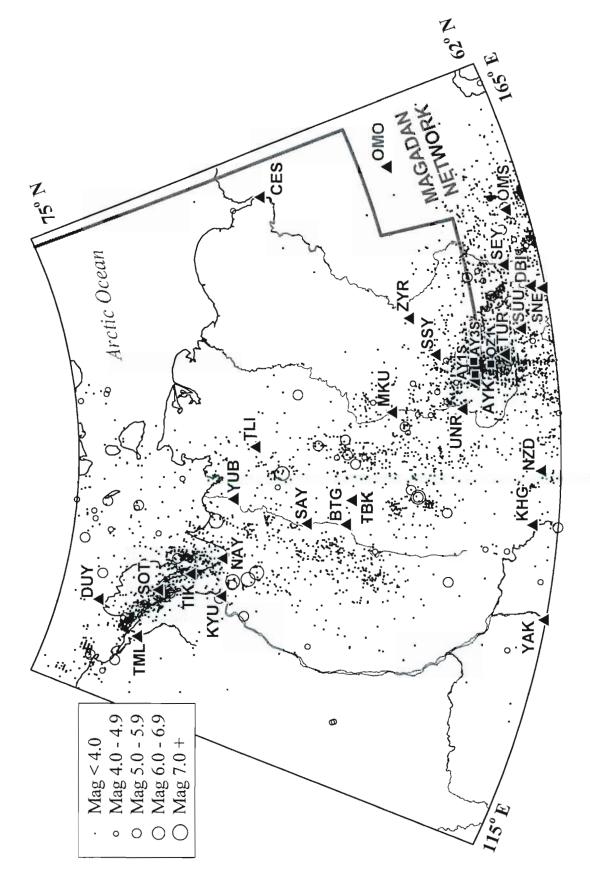


Figure 9. Seismicity and seismic stations of the northernYakutsk network. Temporary seismic stations deployed after the 1971 Aryk earthquake are shown as squares.

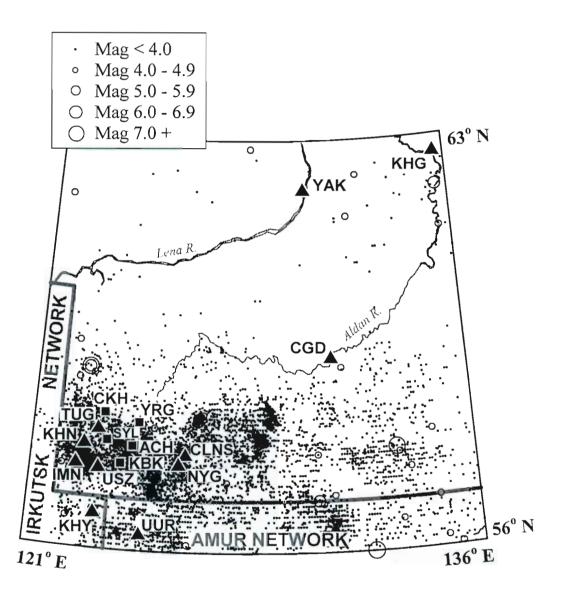


Figure 10. Seismicity and seismic stations of the southernYakutsk network. Temporary seismic stations deployed after the 1989 south Yakutia earthquake are depicted as squares.

bulletin added a significant amount of phase and arrival time data for events in the Yakutsk network. For several events along the boundary with the Magadan network, additional supplemental arrival times were read from Magadan network seismograms. Of stations in the Amur and Irkutsk regions, generally only Kirovskii (KIR), Tupik (TUP), Srednyi Kalar (SRK), and Chara (CRS) were used by the Yakutsk network.

## 1.3.6 Amur, Sakhalin, and Irkutsk Regional networks.

The Amur and Irkutsk networks monitor seismicity between the Eurasian plate and the Amur plate. The Sakhalin network monitors seismicity along the boundary between the Amur plate and Okhotsk block. Unfortunately, we have not worked directly with any of these networks, thus first hand knowledge of operational procedures and unpublished data are generally not available.

Permanent seismic stations in the Amur network region first opened in the early 1970's (Figure 11). Seismicity data from the Amur and Sakhalin networks comes primarily from the Far East Bulletin for events prior to 1979. Since 1979, event parameters are taken from Materialy, Zemlet, and the Far East Bulletin. Since 1982, many additional Amur events not found in other sources are available in the unpublished Yakutsk network bulletin.

Phase data from the Amur region is taken from the Far East Bulletin (1972-1978, 1985-1988), Materialy (1979-1990), and unpublished Yakutsk network bulletins (1982-1997). Some phase data from the Sakhalin network stations (Figure 11) are given in the published bulletins of Amur region events. Phase data from the Sakhalin network are available in the Fat East Bulletin, but were not included in the data files.

Seismic stations in the Irkutsk region were first established in the 1950's, with the exception of Irkutsk, which opened in 1901 (Figure 12). This study includes only a small portion of the Irkutsk network (south of 56° N between 120° and 122° E), thus only a relevant subset of the seismic stations are listed (those stations used to supplement Yakutsk and Amur network phase data). Epicenter data for the portion of the Irkutsk network of interest is taken from *Materialy*. Phase data for some of these events were included in the data files.

#### 1.3.7 Seismic Station Parameters.

Considerable effort was devoted to compiling as complete as possible listing of seismic stations to have operated in northeastern Russia (excluding SSD military sites). The station list covers many different seismic networks (e.g. Mackey, 1999; Mackey and Fujita, 1999), and represents what is believed to be the most reliable coordinates available (Appendix A). Several of the stations given here are listed in international station compilations but with a lower level of location precision and/or accuracy. Several stations were documented to have moved slightly one or more times, but parameters for which were never updated or new station codes assigned. The majority of the information tabulated here was obtained directly from the seismic networks, personnel who actually operated stations, station site visits, and various Russian gray literature publications. However, it is not reasonable to document all sources of information used in a report of this scope. This compilation of seismic stations is continuously updated as information becomes available.

Coordinates of most stations listed were hand checked using 1:200,000 Soviet/Russian General Staff military topographic maps. In a few cases, coordinates were adjusted if the map check showed the station to be in an unreasonable place. For example, the possible location of station Bering was found to reside in the ocean, thus the coordinates of the seismic station were adjusted to place it at the meteorological station immediately onshore; a reasonable correction given knowledge of operational procedures in northeast Russia.

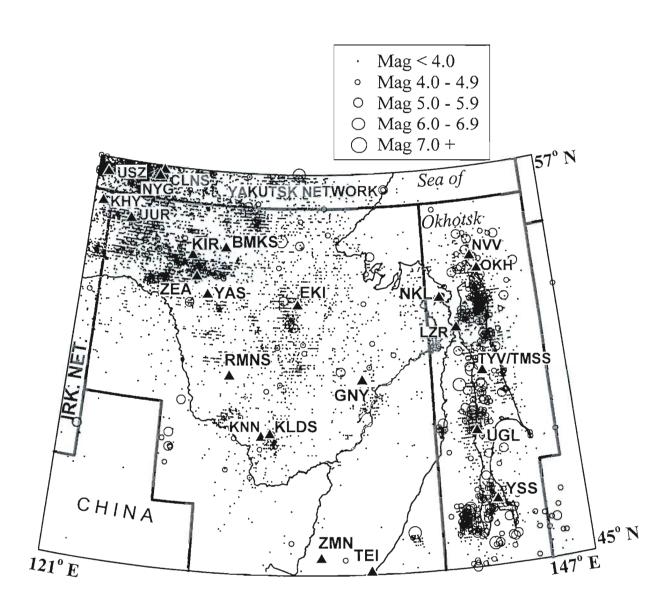


Figure 11. Seismicity and seismic stations of the Amur (left) and Sakhalin (right) networks.

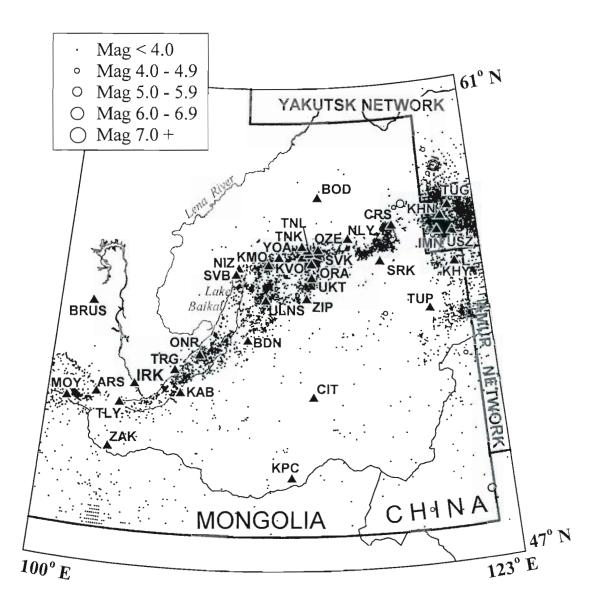


Figure 12. Seismicity and seismic stations of the Irkutsk network. Seismicity shown is primarily from 1970 through 1972.

## 1.3.8 Relationships between magnitude and K-class.

It is standard practice among Russian seismic networks to report magnitudes only for events of about  $M_b = 4.0$  and larger. However, all events, both small and large, have sizes reported as a K-class value, which is based on the logarithm (base 10) of work release in Ergs. For determining K-class, the maximum ground motion  $A_{max}$  is determined using

$$A_{\text{max}} = (A_{\text{p}} + A_{\text{s}})/T \tag{1}$$

where  $A_p$  and  $A_s$  are the respective maximum amplitudes of the P and S arrivals (in microns) and T is the period (in seconds) of the wave. If no amplitude is visible for the P phase, only the S is used. The K-class value, using  $A_{max}$  from (1) and distance, is then read off a nomogram calibrated for each particular region (Solonenko, 1974; Pustovitenko and Kul'chitskii, 1974). Unfortunately, the nomograms calibrated for different regions or networks vary significantly, resulting in inconsistent event size determinations between networks. For example an earthquake having a magnitude of 3.5 may be reported to have a K-class of 9.0 in one network, and a K-class of 10.5 from the neighboring network. To better understand the sizes of events reported in the seismicity catalog, empirical linear regressions were calculated relating K-class to magnitude for each network (Figure 13). Magnitudes used are ISC reported  $M_b$  values up to 5.5. Events larger than magnitude 5.5 use ISC or NEIC reported Ms values as body wave magnitude begins to saturate (Lay and Wallace, 1995).

#### 1.4 CONCLUSION.

The final compiled seismicity map with network boundaries and seismic stations is shown in Figure 2. In reality, the distribution of microseismicity indicates diffuse plate boundaries, probably having motion partitioned on many individual faults in a complex system, which is consistent with continental deformation (England and Jackson, 1989). Within the individual plate boundaries, there are many localized trends and clusters in the seismicity distribution which have never previously been studied in detail. However, several of these clusters and trends, particularly in the Amur region, are a result of anthropogenic sources (see Section 2). In general, the plate boundaries tend to follow the pre-existing large scale structural trends in northeast Russia, which probably represent structurally weaker, and thus easier to deform, regions. The proposed boundary between the Eurasian plate and the Okhotsk plate is somewhat problematic. There have been relatively few teleseismic events defining the location of the boundary, and there is an almost complete lack of located microseismicity in the region. The lack of located microseismicity may be entirely an artifact of the distribution of seismic networks. Note on Figure ES-1 that the region is near the intersection of the Magadan, Yakutsk, Amur, Sakhalin, and Kurile-Okhotsk networks. Furthermore, the division between the northern and southern portions of the Yakutsk networks is also in this region. There is microseismicity trending into the region from each of the Magadan, Yakutsk, and Sakhalin networks. Given the distances to individual networks stations, and the lack of data exchange between neighboring networks, small events occurring in this region are most likely simply below the detection or location threshold for any individual network. However, analysis of data from the digital station deployed in Okhotsk as a part of this study also does not show significant levels of seismicity in this region (see Section 5).

Overall, seismicity levels in northeast Russia are much higher than generally recognized in the literature. Unfortunately, economic conditions in Russia at the present time have resulted in the closure of many, if not most, of the regional seismic stations. This illustrates the importance of

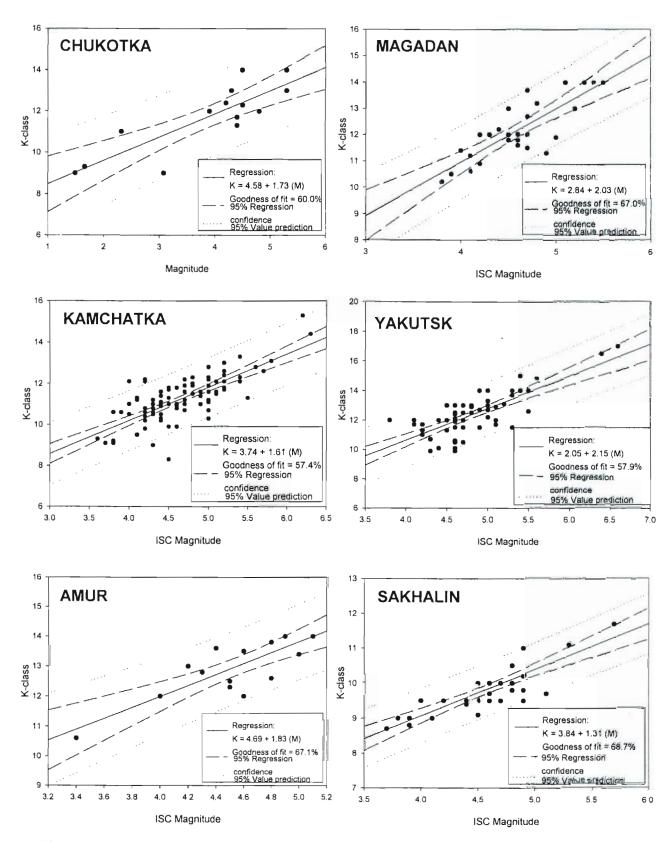


Figure 13. Linear regressions relating K-class to Magnitude for different networks in northeast Russia. Magnitudes up 5.5 are M<sub>b</sub> and greater than 5.5 are M<sub>s</sub>. Note that the derived regressions vary considerably between networks.

compiling the historic data, as further economic collapse in Russia could result in the permanent loss of the original data. It would take the installation of several tens of permanent digital stations and a minimum of two decades to duplicate the historic database using modern acquisition methods.

#### **SECTION 2**

# EXPLOSION CONTAMINATION IN THE NORTHEAST RUSSIA SEISMICITY CATALOG

#### 2.1 INTRODUCTION.

The newly compiled seismicity catalog for Eastern Russia (see Section 1) allows us to reexamine the regional seismicity of northeastern Russia to determine the level of explosion contamination and obtain a better idea of the level of natural seismicity in the region. In this part of Russia, explosions occur in tin, coal, and placer gold mines as well as in the construction of roads, railways, and dams. Many of the active mining regions are geographically associated with the seismically active regions, which can result in misidentification of mine blasts as earthquakes. Unfortunately, contamination of the seismicity catalog with explosions results in an erroneous perception of natural seismicity as well as seismic risk assessment.

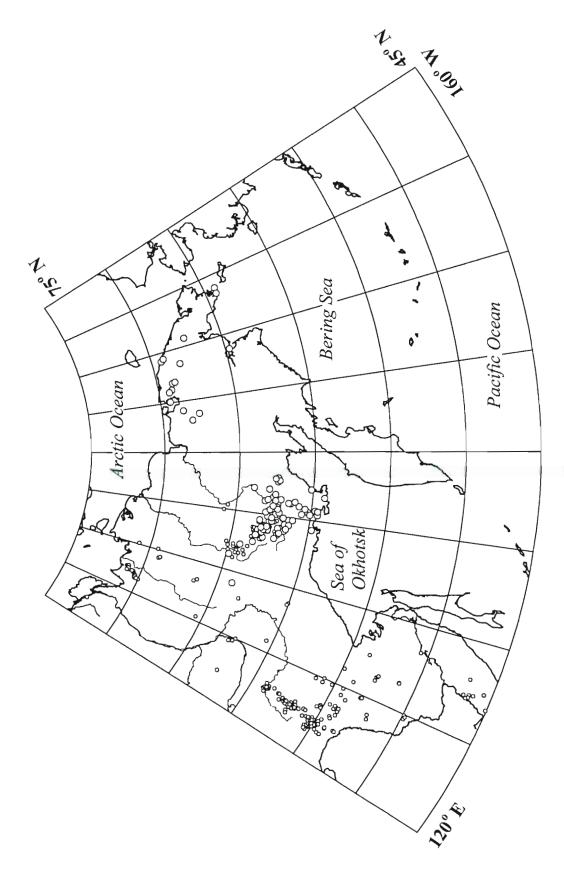
This section summarizes explosion contamination in the study area through temporal analysis of origin times. Both time of year and time of day are considered. While analysis of waveform data for all reported earthquakes would require an unrealistic re-examination of several hundred thousand analog seismograms, a qualitative estimate of the level of explosion contamination can be obtained by examining the spatial, size, and temporal characteristics of earthquakes located by the regional networks.

#### 2.2 DATA SOURCES.

The origin times of earthquakes used here come from the compiled seismicity database for northeastern Russia as described in Section 1. Information on known explosion locations comes from the unpublished Magadan and Yakutsk network bulletins. The unpublished Magadan bulletin lists explosions either by region (nearest town) or mine (Figure 14). Many mines or regions have hundreds of explosions identified. The unpublished Yakutsk bulletin lists explosion coordinates and origin times determined by locating the event with S-P time differences (Figure 14). The location procedure is the same as that used for earthquakes (Section 1). Only a small fraction of the explosions listed in the unpublished Yakutsk bulletin have been entered into the database. A small number of located explosions in the Amur region are given in Godzikovskaya (1995; Figure 14) for the time period 1990-1991. Mine locations from throughout the study region were taken from 1:200,000 scale Russian military topographic maps, as well as 1:500,000 TPC and 1:1,000,000 ONC scale U.S. air navigation charts (see NERSP-10, in prep.).

#### 2.3 PREVIOUS WORK.

Early attempts at explosion filtering from the late 1960's through mid 1970's in the Magadan region simply removed all events within a particular radius of some mining regions (q.v., Riegel, 1994; Kovalev, pers. comm.). Of course, this also removes the wanted tectonic events, and results in peculiar rings of seismicity. Beginning in the late 1970's, seismic station operators attempted to discriminate close events (up to 50 - 70 km) based on waveform characteristics and information



and large dots are towns or mines with multiple explosions. This map is not authorative for all mine explosions in the region. Explosion sources in northeast Russia listed in Russian bulletins. Small dots are individual located explosions, Figure 14.

from the mining companies. Unfortunately, not all mining companies provided information on their blasting activities for various reasons (Kovalev, pers. comm.).

Previous work on identifying explosion contamination in the seismicity catalog was undertaken by Godzikovskaya (1995). Godzikovskaya (1995) identified several regions of explosion contamination, particularly in the Zeya basin region of Amur, near the Kolyma and Ust' Srednekan dams on the Kolyma River, and the Polyarnyi mining district in Chukotka. However, many regions and trends of contamination were not identified because of an incomplete seismicity catalog. Odinyets (1996) identified the problem of explosion contamination in the Kolyma region. In this case, it was determined that a large fraction of earthquakes reported in the central Kolyma region are actually explosions. Industrial explosions locatable by the regional networks generally have magnitudes from about 1.5 up to a maximum of 3.5 (converted from Russian K-class) and occur during local day (Godzikovskaya, 1995; Odinyets, 1996). Placer deposit explosions are also concentrated during the late winter and early spring, when frozen ground is broken up for the summer processing season.

#### 2.4 DISCUSSION.

Although the regional networks operating in northeastern Russia have attempted to discriminate between industrial explosions and earthquakes, all seismicity catalogs containing events of less than magnitude 3.5 from the region remain contaminated with explosions related to the break-up of placer deposits, coal, tin, and gold mining, and construction projects.

Examination of the temporal biases in the seismicity can, in a broad sense, indicate potential regions of explosion contamination (Agnew, 1990), as blasting generally occurs during the day. A small but not statistically significant number of explosions are also known to occur during the nighttime hours in many locations throughout the study area (Godzikovskaya, 1995). In Figure 15, the study area is divided into 0.5 x 1° cells in which the percentages of daytime earthquakes are calculated. North of 64°N, cell size was increased to 0.5 x 2°. Cells containing fewer than ten events were not considered to be statistically significant, thus were not analyzed. As there are five time zones spanning the region, the 12 hour local "day" period was shifted accordingly. Light blue areas represent regions where seismicity is more or less balanced between night and day, and dark blue areas are those in which seismicity is concentrated during local night. Bias of seismicity to local night is not unexpected since most seismic stations are located in populated areas and have lower cultural noise levels during the night. This is consistent with better nighttime recording, particularly for smaller events in the outlying portions of the network. Analysis of the 1989 South Yakutia aftershock sequence confirms the better conditions of nighttime recording; of 3,492 located earthquakes, 1,815 occur during local night and 1,677 during local day.

Pink areas on Figure 15 represent regions where more than 65% of the seismicity occurs during local day. Many of the regions with predominantly daytime events are associated with discrete clusters or trends of seismicity, most of which can be associated with mining or construction related blasting. Several clusters of reported seismicity in the Amur region have more than 90% of the events occurring during local day. A few cells indicate predominantly daytime seismicity, but with no apparent relation to mining. These cells are probably a result of random statistics of small numbers as they are close to the ten event cutoff. Other areas with explosion contamination are discussed in Mackey (1999).

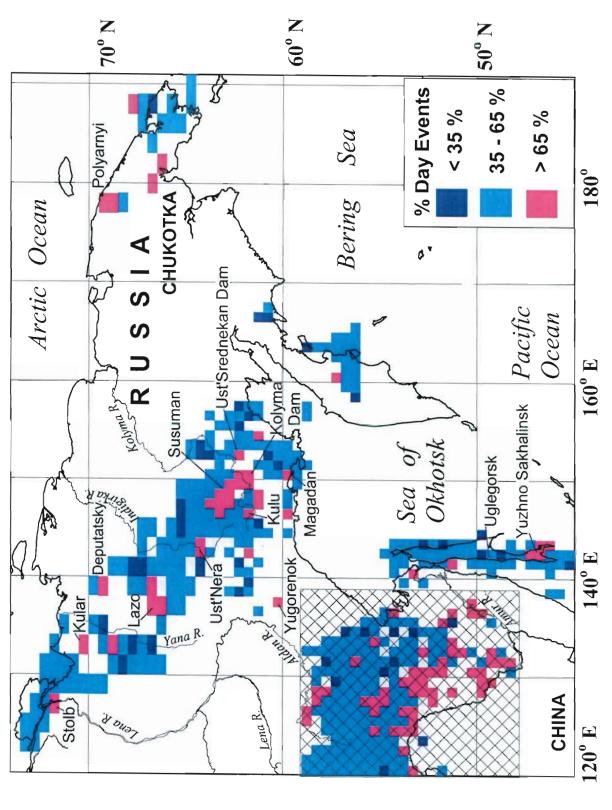


Figure 15. Percentage of seismicity occurring during local daytime. Cells containing primarily daytime events are presumed to have explosion contamination. Labeled regions are discussed in the text. Seismicity associated with subduction of the Pacific plate were not evaluated. The crosshatched area encompasses the southern Yakutia and Amur regions (Figure 16).

## 2.4.1 Amur and Southern Yakutsk District.

The clearest regions of explosion contamination are in the Amur District. When plotting daytime and nighttime epicenters separately (Figure 16), some distinct differences are apparent. There are several large clusters of daytime seismicity that correlate geographically with specific mining regions. For these regions, contamination levels also change with season. Clusters and trends composed primarily of daytime events have very few events of magnitude 4 or larger. Figures 17 through 21 show the number of events that occur during specific hours of specific months of the year. Figure 17 analyzes a reported cluster of seismicity in the northern part of Figure 16 (labeled 'TEST'). The seismicity in this region is believed to be tectonic, as the region is unpopulated and should not contain any contamination due to mining. The analysis of this region can therefore be used as a baseline to which other analysis can be compared. The area analyzed contains 399 located earthquakes, with 189, or 47.3%, having occurred during local daytime and 205 during local night. The number of summer and winter events occurring in this region are also similar.

Analysis of reported seismicity around the Raychikhinsk coal mining region shows activity only during winter months, and daylight hours (Figure 18). This is somewhat different than from the nearby Khingansk mining region where events occurring throughout the year but also only in the day (Figure 19). Other regions identified in the Amur region show similar distributions of origin times, and are associated with specific mining regions. See Mackey (1999) for additional discussion.

In the central portion of Figure 16A, there is a northwest-southeast trend of predominantly daytime seismicity extending several hundred kilometers. Temporal distribution of these events is shown in Figure 20 for the central segment. This trend correlates with the route of the Baikal-Amur mainline railroad (BAM). Explosions associated with its construction in the 1980's are listed as earthquakes in the seismicity catalogs. Note also that most events are located to the west of the track, indicating systematic errors in the location procedure.

The northern portion of Figure 16 includes the southern Yakutsk network, where two mining regions are of note. Aldan is a mining region with extensive deposits of gold and phlogopite mica (Shabad, 1969). The region is associated with a diffuse cluster of predominantly daytime reported seismicity. Although the seismicity occurs throughout the year, its concentration during daylight hours is more consistent with an anthropogenic source than tectonic. Explosions from the Aldan mining region are also located and listed in the unpublished Yakutsk bulletin.

To the southeast of Aldan is a dense cluster of seismicity near the settlement of Spokoynoi (Figure 15). Temporal analysis of the cluster shows a strong bias towards daytime and winter events, which would be consistent with placer mining (Figure 21). Soviet military 1:200,000 scale topographic maps (dated 1986) show extensive mine workings in the region, but list all the nearby settlements as uninhabited. This is inconsistent, as the events located here occur from the 1970's through the mid 1990's. The published literature lists no mention of any mining activity in this region, nor does the unpublished Yakutsk network bulletin identify any explosions in the region. The nature of activity at this location remains unclear, but may represent residual or exploratory mining. Temporal variation plots of other regions showing contamination on Figure 16 can be found in Mackey (1999).

Overall, the explosion contamination appears to be primarily confined to daylight hours, thus it is assumed that nighttime seismicity should better reflect the natural tectonic distribution of earthquakes (Figure 16b). A different, more northerly trend appears in the plotted nighttime earthquake epicenters, which probably delineates an active tectonic feature that was previously obscured with clusters and trends of explosions. When observing the locations of the larger

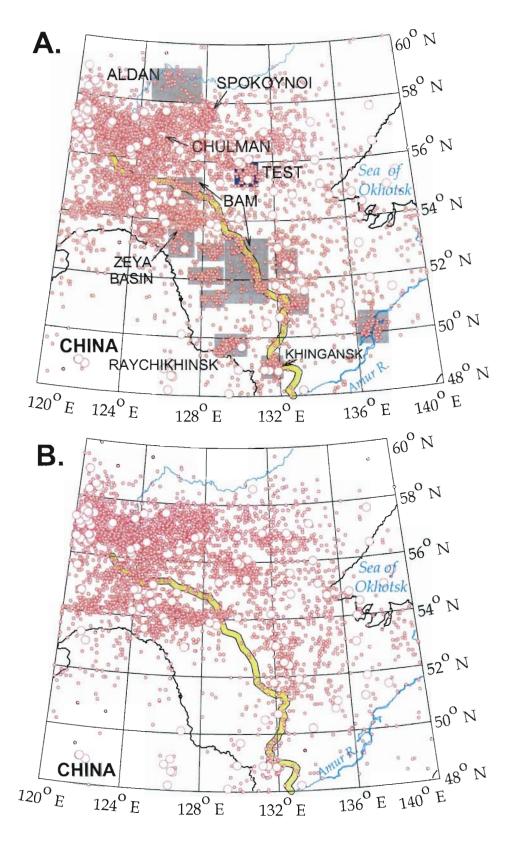


Figure 16. Day (A) and night (B) seismicity of the Amur region. Note a correlation between daytime seismicity and the Baikal-Amur (BAM) railway (yellow). Gray areas indicate clear contamination, while blue indicates the temporal variation test area (Figure 17).

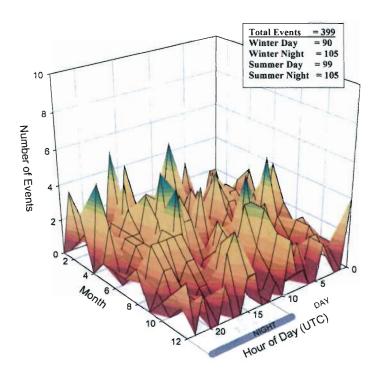


Figure 17. Temporal variation of presumed natural earthquakes from the region labeled 'TEST' on Figure 16. Note a slight bias towards nighttime events, but no seasonal bias.

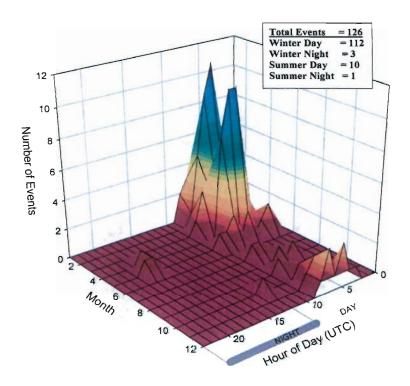


Figure 18. Temporal variation of seismicity in the region of the Raychikinsk mine. Note the strong bias towards daytime winter events.

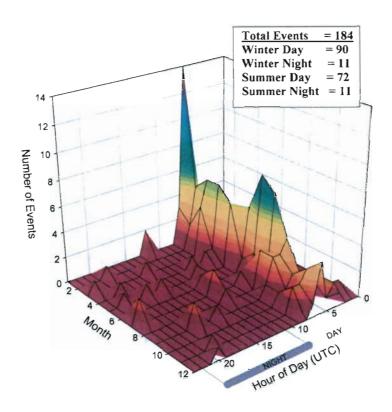


Figure 19. Temporal variation of seismicity in the region of the Khingansk mine. Note the strong bias towards daytime events.

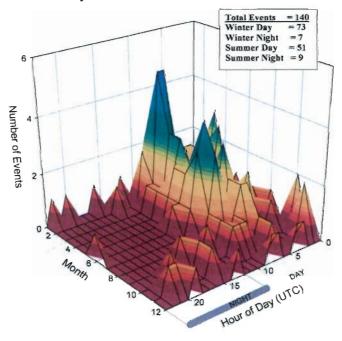


Figure 20. Temporal variation of seismicity associated with the central segment of the Baikal-Amur mainline railway. Note the bias towards daytime events. Most of these events are thought to be explosions associated with construction of the railway.

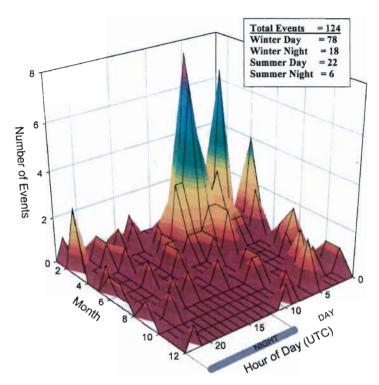


Figure 21. Temporal variation of reported seismicity in the Spokoynoi region.

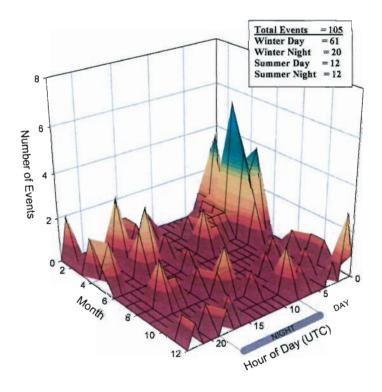


Figure 22. Temporal variation of seismicity in the Polyarnyi, Leningradskii, and Plamennyi mining region. See Figure 5.

teleseismically located events, it is found that they fall almost entirely within the regions where seismicity occurs during the night.

# 2.4.2 Polyarnyi-Leningradsky-Plamennyi.

Polyarnyi and Leningradsky are placer gold deposits located along the coast of the Chukchi Sea in Chukotka. Plamennyi was a mercury deposit located approximately 100 km to the south which operated from 1967 to 1972 (Pilyasov, 1993; see Figure 5). From 1966 to 1982, most of the events located in this area were single station locations obtained by the three-component seismic station at Iul'tin (ILT; see Figure 5). A clear bias towards winter and daytime is evident (Figure 22) for the cluster of events in the mining region. Comparison of origin times of ILT located events with the more recent known explosions from the same mining region yields a nearly identical temporal distribution.

#### 2.4.3 Kolyma Gold Belt and Northern Yakutia.

A band of predominantly daytime seismicity lies along the Kolyma gold mining belt (Figure 15). Tectonically, this region is extremely complex in that it is located just south of the Ulakhan and Darpir Fault systems (Figure 23), along which motion between the Okhotsk block and the North American plate is occurring (Riegel et al., 1993; Imaev et al., 1994). As a result of the tectonic setting, easy statistical separation of anthropogenic sources from earthquakes is more troublesome. Mining in this region is primarily placer gold, but also includes coal and other minerals. Typical large blast sizes in the placer gold deposits around Susuman may range from 50 - 75 T, while explosions in the coal mines may exceed 200 T (Leith, 1994; see Section 5). Daytime and nighttime seismicity are shown on Figure 23. Temporal analysis of the large cluster of events in the area northwest of the town of Susuman indicates there is a clear bias towards local day and winter/spring (Figure 24). This bias is consistent with the distribution of known explosions from the Magadan seismic bulletin for the Susuman region (Figure 25). Based on the day/night seismicity plots in Figures 23a,b, the eastern half of the seismicity cluster is entirely explosions, while the western half probably contains some tectonic events. Note also the reduced numbers of nighttime events within a 100 kilometer radius of Susuman, which is also probably a result of mining. About 200 kilometers southeast of Susuman is the Kolyma hydroelectric station and dam. Blasting during construction of this dam resulted in contamination of the earthquake catalog (Godzikovskaya, 1995).

The level of seismic activity in northern Yakutia is lower than the Kolyma region, with much of the seismicity in isolated clusters and regions. Several of the clusters associated with mining regions and having predominantly daytime events include Lazo, Ust' Nera, and Kular with placer gold deposits, the Deputatsky tin deposit, and placer diamond exploration near Stolb (Figure 15). The placer gold deposit at Yugorenok in east-central Yakutia is also associated with a cluster of daytime events.

Explosion contamination of the seismicity catalog has clearly affected analysis of seismic hazards in the region. Vazhenin et al. (1997), citing unpublished material by T.A. Andreev, shows an increased seismic hazard level in the region north of Susuman, around Kulu, in a trend extending south from Kulu, and near the Kolyma hydroelectric station, all of which are areas of explosion contamination.

#### 2.4.4 Sakhalin Island.

Sakhalin Island shows no clear pattern of explosion contamination except at the extreme southern tip of the island (Figure 15). The contamination here is concentrated around the city of

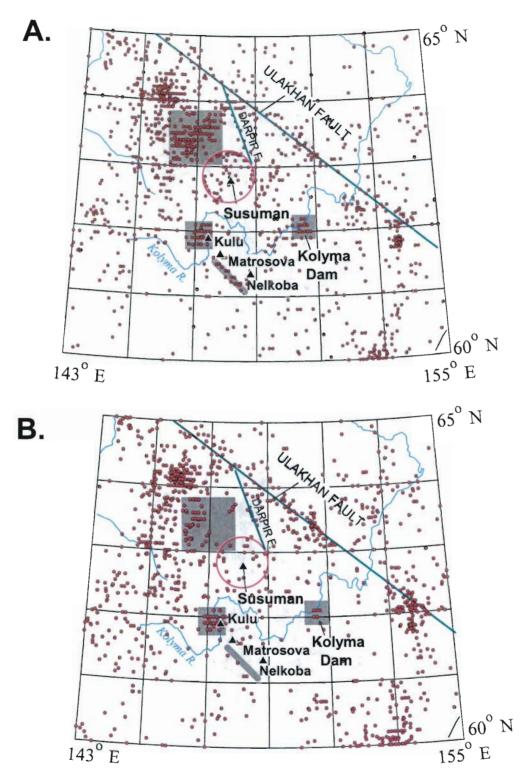


Figure 23. A. Daytime seismicity of the Kolyma gold belt. B. Nighttime seismicity of the Kolyma gold belt. Shaded areas indicate regions used in temporal analysis of event origin times and other regions of explosion contamination. Note ring of seismicity around Susuman in the daytime plot.

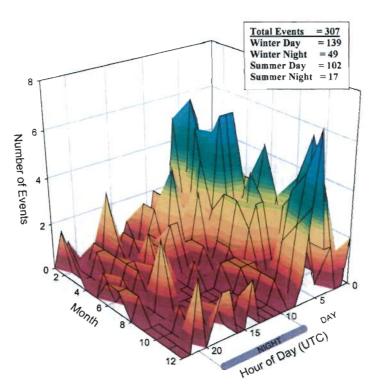


Figure 24. Temporal variation of seismicity in the region northeast of Susuman. Note the bias to winter daytime events. See also Figure 25.

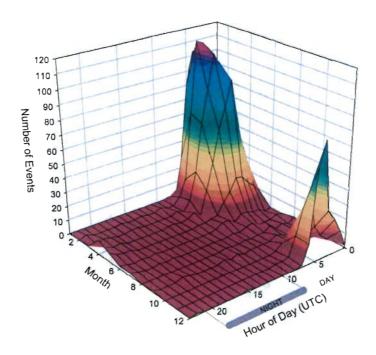


Figure 25. Temporal variation of approximately 3000 known explosions from the Susuman region. The occurrence of explosions here is consistent with the presence of explosion contamination in Figure 24.

Yuzhno Sakhalinsk where there are likely quarries for road gravel or other stone products. No bias is observed in the Uglegorsk coal mining region.

## 2.5 CONCLUSION.

Based on a very simplistic temporal analysis, it is evident that the seismicity catalog of northeast Siberia is heavily contaminated with industrial explosions. Identification of these explosions and their removal from the seismicity catalog is essential in studying and understanding the tectonics and associated natural seismicity of the region and the associated seismic risk. Phase data for known ground truth explosion sources can be used to develop better calibrated regional velocities and improve event location procedures for the area.

#### **SECTION 3**

# CRUSTAL MODEL AND RELOCATIONS OF NORTHEAST RUSSIA EARTHQUAKES

#### 3.1 INTRODUCTION.

The ability to accurately locate seismic events is of utmost importance for monitoring of the CTBT. Of course, high quality locations depend on using the correct velocities in the location procedure. The existing database contains several thousand events which have been located by several seismic networks, each employing different location methodology, travel time curves, and phase data (see Section 1). Hypocenter determinations throughout the study area can be improved by using a single methodology, calibrated travel time curves, and combined phase data from adjoining networks. This section summarizes the results of relocating the larger events in the study area in conjunction with developing best-fit crustal travel time curves. This section also discusses the failed attempt to develop upper mantle P<sub>n</sub> tomography models for northeastern Russia. Of particular interest for this study was the Laptev Sea and northern Yakutia region, where the velocity model developed indicates sharp velocity contrasts.

## 3.2 DISCUSSION.

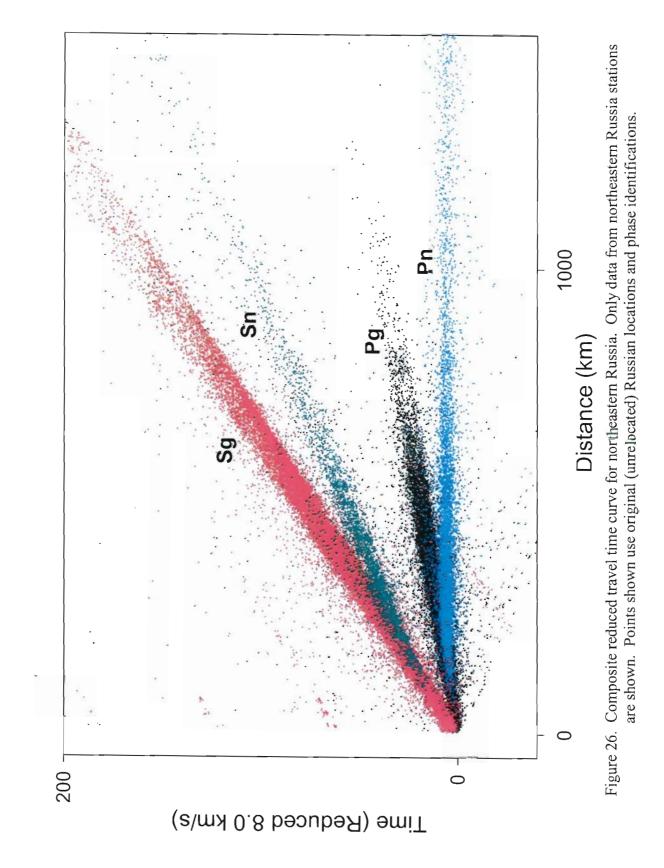
## 3.2.1 Earthquake Relocations and Crustal Velocities.

To gain a better understanding of the Russian computed locations, a composite reduced travel time curve based on the original epicenter and origin times was produced (Figure 26). Relocations computed here build on relocation work presented in Mackey (1996), and Mackey et al. (1998). The basic location routine is a standard least squares best fit routine. Originally, this program used only P phase arrivals, but was modified for this study to accept multiple phases, specifically Pg, P<sub>n</sub>, and Sg. In keeping consistent with notation used in data sources used in this study, Lg onset is referred to here as Sg. S<sub>n</sub> arrivals are not used as the data are very noisy.

Throughout the study area, many geologic and tectonic environments exist, which combined with the physical vastness of the region, makes it likely that no single velocity or travel time curve will reflect actual seismic velocities for any particular phase. In order to overcome this problem, the study area was broken into cells, and the best fitting velocities were determined by minimizing the sum of event RMS residuals through trial of multiple travel-time curves. Block sizes are generally 3° north-south by 5° east-west, although this was adjusted in areas with sparse activity in order to increase the number of useable events. For any given block, only events containing P<sub>n</sub> phase arrivals (generally 2 or more) were used. This selects only the larger events, which contain more arrivals and have better azimuthal coverage of receiving stations.

Travel-time curves for Pg and Sg phases were calculated assuming a straight-ray approach. For hypocenters at depth, the travel path is assumed to be the hypotenuse of the triangle made by the depth and the surface distance. Event depths are restricted to a maximum of 33 km, as all events are assumed crustal in nature. Events for which depth tends above the surface are restricted to 0 km. This is reasonable, as the database is known to be contaminated with explosions (see Section 2).

In order to determine the best crustal velocities for locating events within a given cell, the selected events were first located with a first approximation of crustal velocities for the region. The velocities used in the first approximation were generally the best fit velocities from an adjacent cell.



The resulting output was analyzed for high residual arrivals, generally those greater than 3.5 seconds. High residuals are generally a result of either typographical errors, misassociated phases, or bad time picks. Typographical errors and misassociated phases were corrected, if possible, while bad time picks were flagged to be omitted from use in further locations. Events with many high residual arrivals, such that the solution is unstable, and events with four or fewer recording stations, were omitted from further analysis. Overall, less than 5% of the originally selected events were omitted due to stability problems.

In order to determine the better fitting travel time curve to use for  $P_n$  arrivals in the location procedure, the Jeffreys-Bullen (JB; Jeffreys and Bullen, 1970) P wave travel time curve was compared to the IASPEI 91 (I-91) curve (Kennett, 1991). Results for test regions in Amur, northern Yakutia, Magadan, and Chukotka indicated that the JB table did a better job of fitting the  $P_n$  arrivals in the region. Overall, when comparing RMS residuals for events located using the two travel-time curves, approximately 80% of the events have a lower RMS residual when the JB table is used as opposed to the I-91 table; the Pg and Sg velocities used in each comparison were the same. Therefore, the JB table was used in all relocations computed in this study.

Following removal of large residual time picks and correction of misassociated phases, the remaining selected events for a given block were located multiple times using different crustal velocities. Crustal velocities tested for each block range from 5.875 km/sec to 6.350 km/sec in 0.025 km/sec increments for Pg, and 3.470 km/sec to 3.650 km/sec in 0.020 km/sec increments for Sg. In this manner, the crustal velocities which best fit the events in the cell were found. The newly found best fitting velocity for each cell is then used to relocate the events a second time. After locating with the new velocity, any additional high residual arrivals (generally those over 3.0 seconds) were omitted or corrected. If any arrivals were removed or had phase associations changed, the events in the cell were again subjected to a search for the best fit crustal velocities. Three-dimensional plots of a cell's cumulative residuals for varying Pg and Sg velocities are useful to illustrate how the residuals change with differing velocities. Figures 27 and 28 illustrates how the residuals change for two cells, one in the southern Yakutsk region, the other in Magadan. Final crustal Pg velocities for each cell are shown on Figure 29 (Table 1). Sg velocities show a similar pattern (Figure 30; Table 1). In conjunction with developing the crustal velocities, a total of 1311 earthquakes were relocated in 44 geographic cells (Appendix B).

Crustal velocities in northern Yakutia are somewhat problematic, with adjacent cells alternating between high and low velocities (Figures 29 and 30). Therefore, northern Yakutia was re-evaluated using a 5° x 3° moving window, shifting the window in 1° increments. For each cell, the best fitting crustal velocities were determined by trial of multiple Pg and Sg velocities as discussed above. This resulted in a similar velocity structure as determined earlier, for both Pg and Sg arrivals, but with greatly smoothed velocity shifts and better resolution (Figures 31 and 32).

#### 3.2.2 Results of Relocations.

Crustal velocities determined in the location process correlate well with the regional tectonic provinces. Generally the highest velocities (Pg velocities ranging from 6.225 to 6.300 km/s and Sg from 3.61 to 3.65 km/s) occur in the western portion of the study region, which is associated with the Siberian Platform. South of the Siberian platform, velocities decrease across the Mongol - Okhotsk suture (Figure ES-2). Velocities also drop sharply in the Verkhoyansk foldbelt, along the eastern edge of the Siberian platform. From the Verkhoyansk foldbelt and east through the Mesozoic terrane assembledges (Kolyma - Omolon superterrane) to the Bering Strait, crustal velocities are consistently in the 6.00 - 6.05 km/s range, and Sg velocities in the 3.51 - 3.55 km/s range, with only a few cells deviating slightly. The final analysis for velocities in northern Yakutia

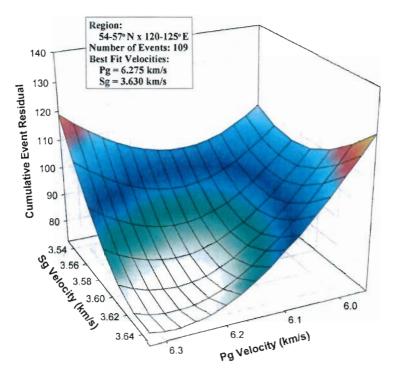


Figure 27. Pg-Sg velocity residual graph showing fast velocities for one cell in southern Yakutia.

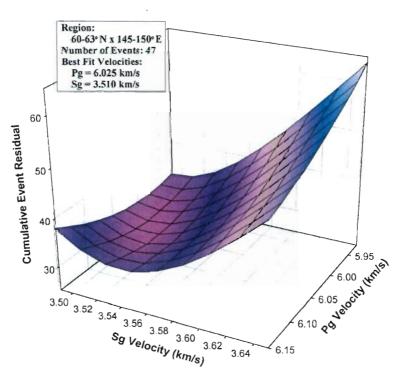


Figure 28. Pg-Sg velocity residual graph showing slow velocities for one cell in the Magadan region.

Table 1. Best fit velocities and number of events per geographic region.

Region	Best Velocities	km / sec 1	Number of events
48-51N 120-125E	Pg=6.040	Sg=3.530	2
48-51N 125-130E	Pg=6.025	Sg=3.530	12
48-51N 130-135E	Pg=6.100	Sg=3.550	28
48-51N 135-140E	Pg=6.100	Sg=3.550	4
51-54N 120-125E	Pg=6.100	Sg=3.570	14
51-54N 125-130E	Pg=6.125	Sg=3.570	32
51-54N 130-135E	Pg=6.125	Sg=3.570	47
51-54N 135-140E	Pg=6.075	Sg=3.550	20
54-57N 120-125E	Pg=6.275	Sg=3.630	109
54-57N 125-130E	Pg=6.175	Sg=3.590	69
54-57N 130-135E	Pg=6.150	Sg=3.590	40
54-57N 135-140E	Pg=6.100	Sg=3.570	19
57-60N 120-125E 57-60N 125-130E	Pg=6.275	Sg=3.630	149
57-60N 125-130E 57-60N 130-140E	Pg=6.300	Sg=3.650	44
57-60N 130-140E	Pg=6.275	Sg=3.630	9
57-60N 140-145E	Pg=6.075	Sg=3.570	7
57-60N 145-150E	Pg=5.950	Sg=3.510	42
57-60N 155-160E	Pg=6.040	Sg=3.510	33
60-63N 135-140E	Pg=6.025 Pg=6.050	Sg=3.510	35
60-63N 140-145E	Pg=6.050	Sg=3.570 Sg=3.530	4
60-63N 145-150E	Pg=6.025	Sq=3.530	35 47
60-63N 150-155E	Pg=6.025	Sg=3.510	47
60-63N 155-160E	Pg=6.06	Sg=3.510	63
60-63N 160-165E	Pg=6.05	Sq=3.51	11
60-65N 165-170E	Pg=5.975	Sq=3.470	19
60-65N 170-180E	Pg=6.000	Sg=3.470	19
60-66N 120-135E	Pg=6.225	Sq=3.610	7
63-66N 135-140E	Pg=6.050	Sg=3.550	11
63-66N 140-145E	Pg=6.050	Sg=3.530	31
63-66N 145-150E	Pg=6.025	Sg=3.510	103
63-66N 150-155E	Pg=6.025	Sq=3.530	29
63-66N 155-160E	Pg=6.025	Sg=3.510	10
63-69N 180-160W	Pg=6.050	Sg=3.530	48
65-69N 165-180E	Pg=6.060	Sg=3.530	11
66-69N 125-130E	Pg=6.275	Sg=3.610	6
66-69N 130-135E	Pg=6.075	Sg=3.570	14
66-69N 135-140E	Pg=6.200	Sg=3.550	8
66-69N 140-145E	Pg=6.175	Sg=3.570	17
66-69N 145-150E	Pg=6.040	Sg=3.53	2
69-72N 125-135E	Pg=6.050	Sg=3.550	18
69-72N 135-145E	Pg=6.200	Sg=3.590	14
72-75N 120-130E	Pg=6.225	Sg=3.570	15
72-75N 130-135E	Pg=5.900	Sg=3.510	5

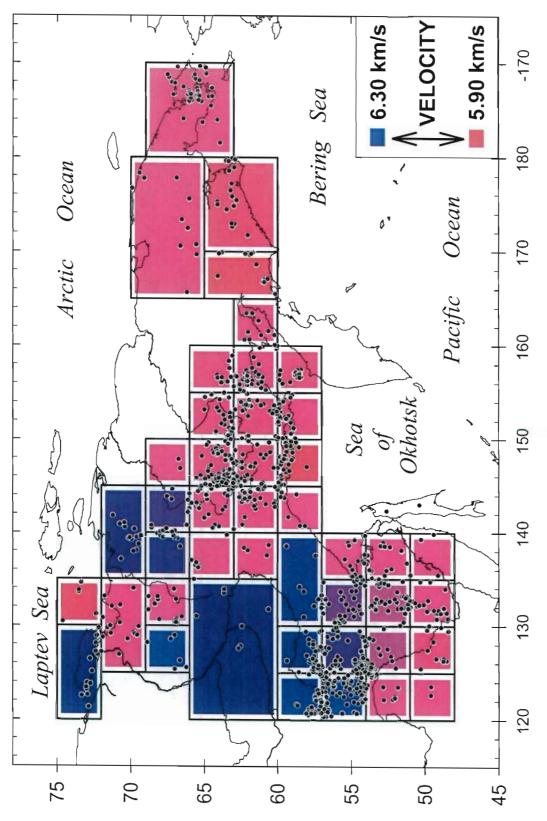


Figure 29. Grid of calibrated Pg velocities. The velocities depicted were calculated in conjunction with relocating the epicenters shown. Velocities found generally correlate with the geology, with the Precambrian Siberian Craton in the west having the highest velocities.

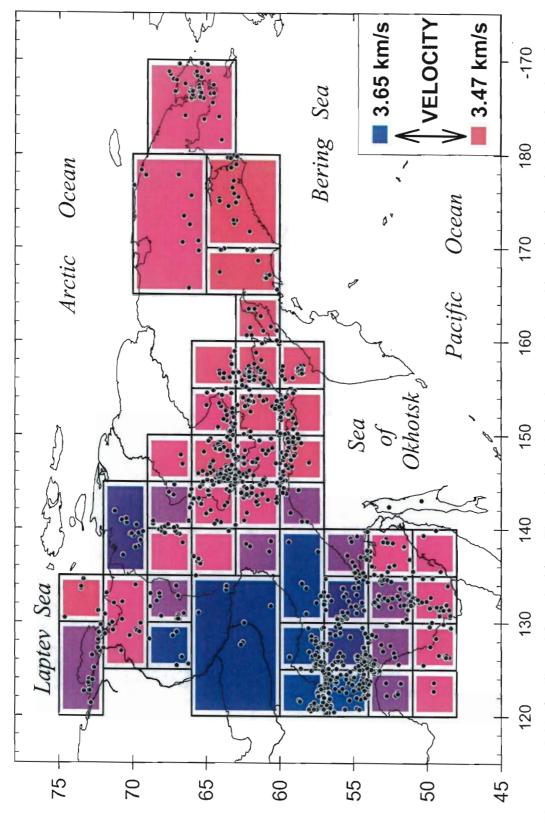


Figure 30. Grid of calibrated Sg velocities. The velocities depicted were calculated in conjunction with relocating the epicenters shown. Velocities shown generally correlate with the geology, with the Precambrian Siberian Craton in the west having the highest velocities.

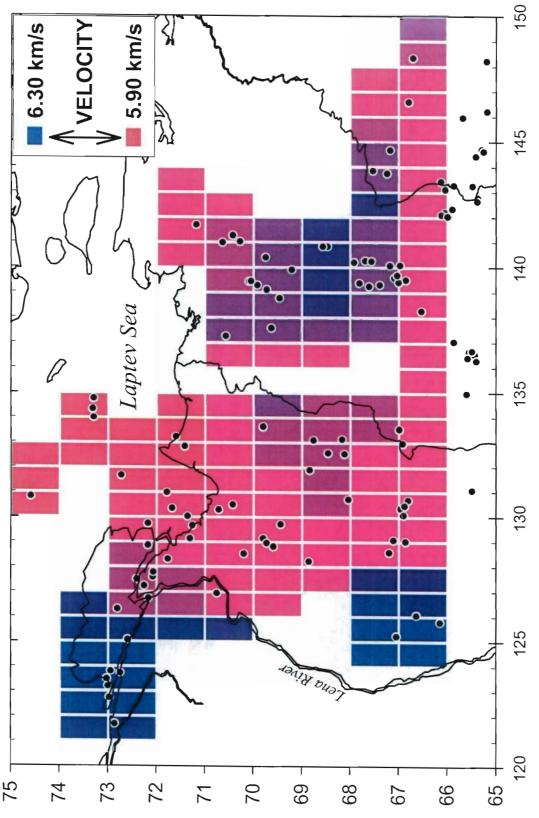
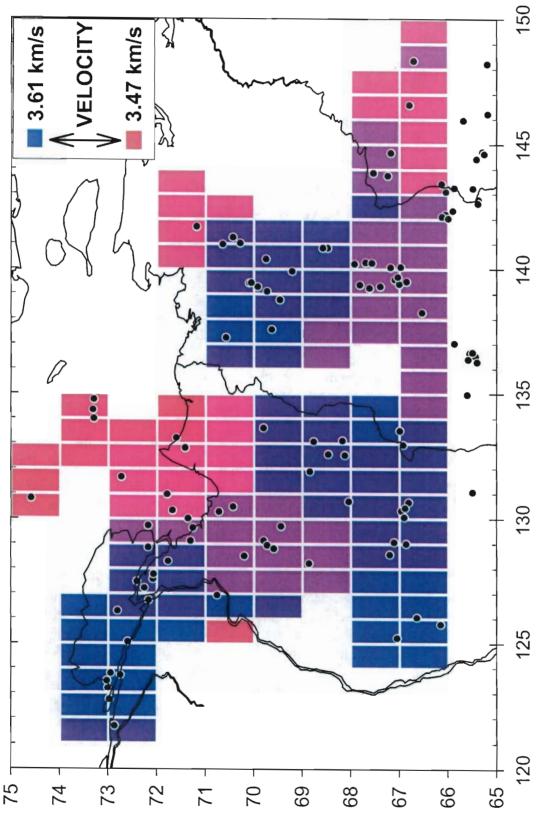


Figure 31. Pg velocities for northern Yakutia determined by using a moving window. Epicenters shown are those used in conjunction with determining velocities. Velocities found generally correlate with the geology, with the Precambrian Siberian Craton in the west having the highest velocities and the Laptev Sea Rift in the north-central portion of the figurehaving the lowest velocities.



Sg velocities for northern Yakutia determined by using a moving window. Epicenters shown are those used in conjunction with determining velocities. Velocities found generally correlate with the geology, with the Precambrian Siberian Craton in the west having the highest velocities and the Laptev Sea Rift in the north-central portion of the figure having the lowest velocities. Figure 32.

indicate that the highest velocities are associated with the Siberian platform along the western edge of the region. The lowest velocities occur in the Laptev Sea and correspond to the active grabens of the continuation of the Arctic Mid-Ocean Ridge into northeast Asia (Figures 31 and 32). The velocity shifts in northern Yakutia are probably a result of rapidly changing velocity gradients associated with presently active rifting adjacent to the Siberian platform and other older tectonic structures. The low velocity region in the Laptev Sea extends into the continent, where it generally follows the strike of the grabens (Fujita et al., 1990a; Drachev, 1998; Sekretov, 1998).

Plots comparing original and relocated epicenters show clear improvement of relative locations. In the Amur region, there is improvement on a seismicity trend extending through the Zeya basin and tightening of seismicity clusters throughout the region (Figure 33). Relocated epicenters in the Magadan region show a tightening of several clusters of seismicity and a slight improvement of events along the trace of the Ulakhan fault (Figure 34). The cluster near 62° N x 157° E is due to an aftershock sequence following an event on February 11, 1987. Clusters of seismicity in the eastern portion of Chukotka are reduced to much smaller lineations (Figure 35). Unfortunately, the seismicity trend in the Koryak Highlands shows no apparent improvement, although this may be due to poor statistics resulting from a small number of events. Original and relocated epicenters for northern Yakutia are shown in Figure 36. Although epicenters clearly moved in the relocation process, the lack of clear clusters and trends makes it difficult to evaluate any improvement in relative locations. Lack of apparent improvement in northern Yakutia may be a result of apparently high lateral velocity gradients.

Travel time curves resulting from the relocated events show a distinct improvement. A composite travel time curve derived from all relocated events (Figure 37) shows a decrease in level of scatter when compared to those derived from the original Russian locations (Figure 26). Because each cell used distinct calibrated crustal velocities, the composite travel time curve (Figure 37) is somewhat smeared, as it reflects all velocities used in the relocation process. Note that arrivals identified as  $S_n$  are not plotted on Figure 37 as they were not used in the relocation process. Travel time curves from individual cells better show the improvement obtained. Figures 38 and 39 compare respective travel time curves for individual cells in the southern Yakutsk the Magadan regions (Figures 27 and 28). The level of scatter is reduced for all phases using the relocations, consistent with improved hypocenter parameters. These resulting curves for individual cells are also distinguishable from each other when compared (Figure 40).

#### 3.2.3 Crustal Thickness.

Initial work to determine crustal structure in the region, undertaken independently of the crustal velocities and relocations addressed above, used relocations of 75 larger events that occurred throughout the Magadan and northern Yakutsk networks. Full details of this study are found in Mackey (1996) and Mackey et al. (1998). Relocations computed here were calculated using only Pg arrivals, assuming a crustal velocity of 6.00 km/s. High residual arrivals were omitted from the location routine and several misassociated phases were corrected. In order to determine the first-order crustal structure, the method of Ruff et al. (1994), is used. This method assumes that regional travel time curves for multiple earthquakes can be approximated with a regional average crustal model. Given a set of earthquakes, the Pg travel time data are simultaneously inverted, solving for the best fit velocity for the set of events, and new origin times for individual events, assuming fixed hypocenters. For a given station, any systematic variation from the regional average should be a result of local variations in crustal structure. The process of relocating the events and inverting the arrival data result in a significant reduction of scatter in the travel time curves for both Pg and P<sub>n</sub>

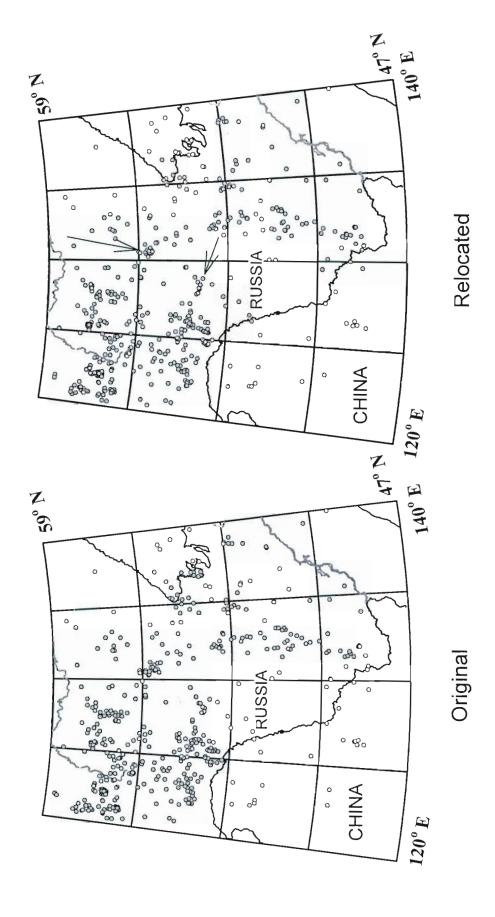


Figure 33. Original vs. relocated epicenters for the Amur region. Arrows indicate locations of improved definition of some seismicity clusters and trends.

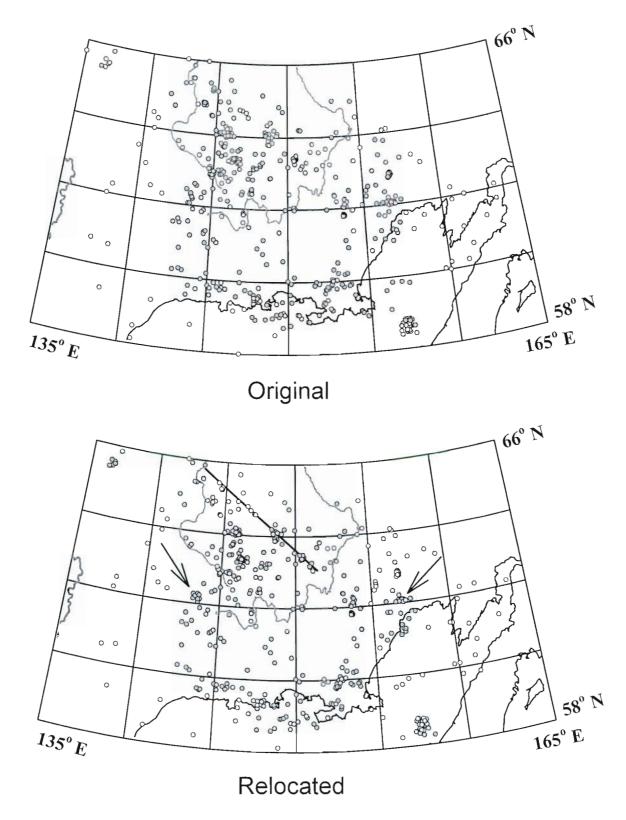


Figure 34. Original vs. relocated epicenters for the Magadan region. Ulakhan fault shown by heavy line. Note improvement in relative locations of clusters indicated with arrows, as well as many other smaller clusters.

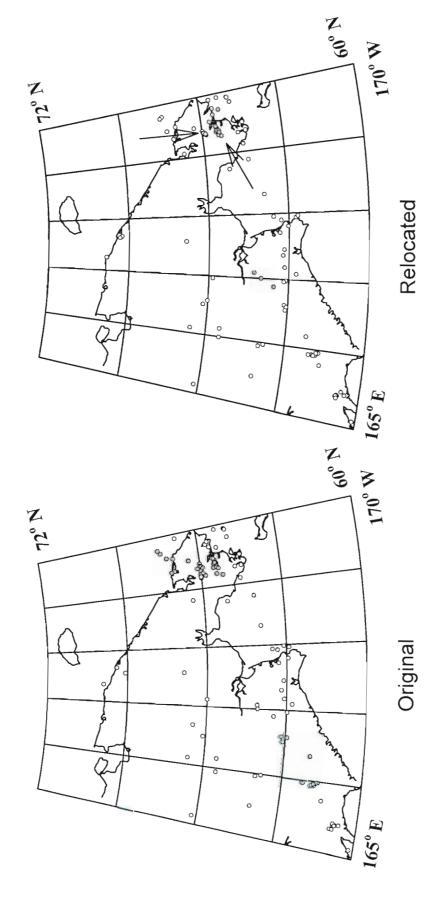


Figure 35. Original vs. relocated epicenters for Chukotka. Arrows indicate improved lineations.

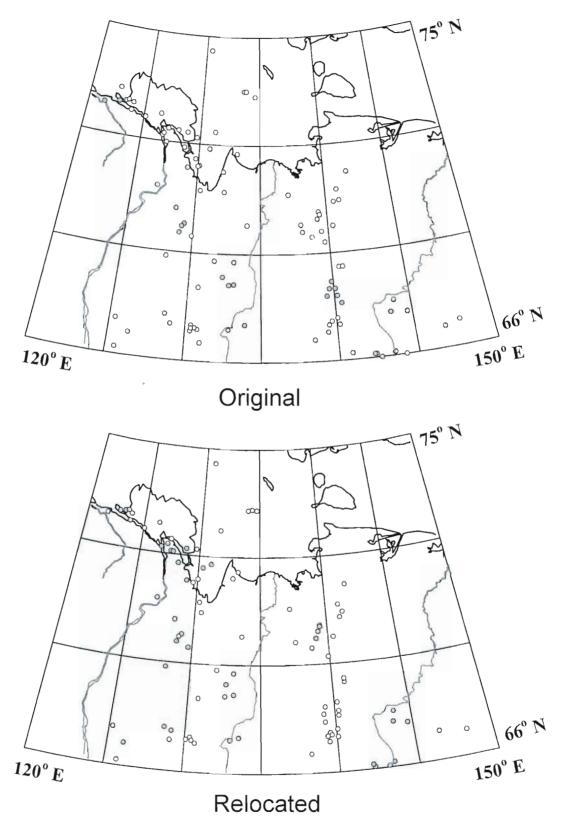


Figure 36. Original vs. relocated epicenters for northern Yakutia.

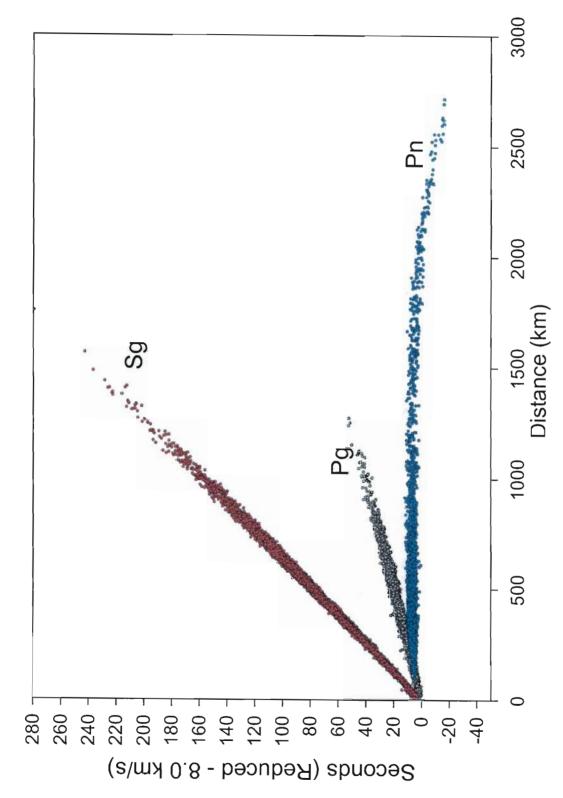


Figure 37. Composite regional reduced travel time curve for northeast Russia using hypocenter parameters from relocations. Sn data are omitted.

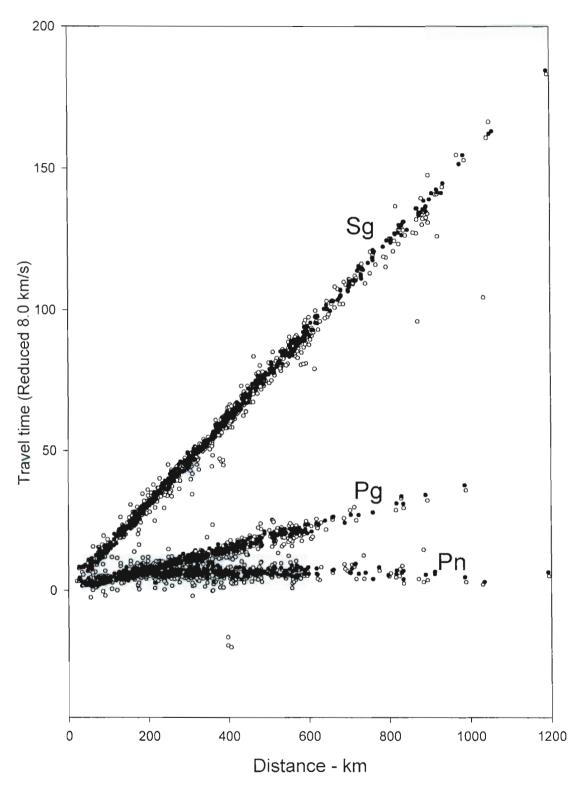


Figure 38. Travel time curve for a cell in the southern Yakutsk region (54-57°N x 120-125°E) comparing original (open circles) with relocated event parameters (closed circles). Note the significant reduction in scatter of data points. Sn data are omitted.

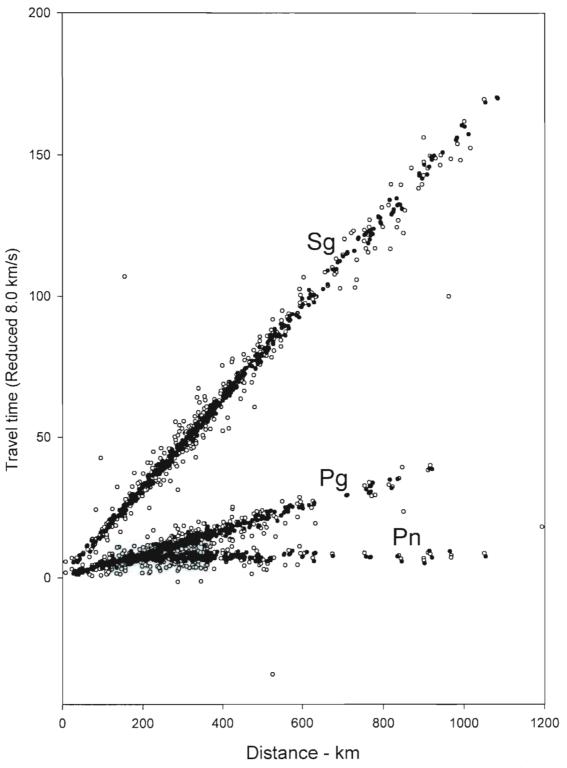


Figure 39. Travel time curve for a Magadan region cell (60-63°N x 145-150°E) comparing original (open circles) with relocated event parameters (closed circles). Note the significant reduction in scatter of data points. S<sub>n</sub> data are omited.

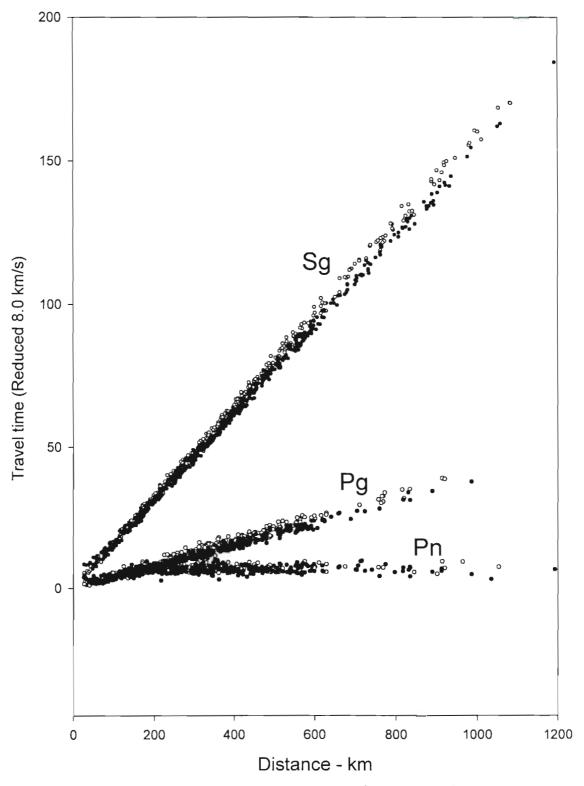


Figure 40. Travel time curve comparing the regions 60-63°N x 145-150°E (open; Magadan) and 54-57° N x 120-125° E (solid; south Yakutia). Note increased velocities for relocated events in south Yakutia. Sn data are omitted.

arrivals (Figures 41 and 42 respectively). Figure 43 depicts the Pg-P<sub>n</sub> composite travel time curve derived from the final Pg inversion new origin times.

As a result of this inversion, the apparent velocities, crossover distance and approximate crustal thickness are obtained for the study area as a whole (Figure 43). Results of P wave velocities for the Magadan and northern Yakutsk regions indicate a generally simple structure with a 5.992 km/sec crustal layer overlying a 7.961 km/sec upper mantle. Average crustal thickness is about 37 km. Examination of individual stations extracted from the composite travel time curve allow determination of crustal structure at individual stations. Sufficient data were available for 27 stations, and present data suggest that structural variations between stations in the study area are resolvable with this method. Based on final results, errors in determining crustal thickness appear to be  $\pm$  4 km, Pg velocities  $\pm$  0.03 km/s and P<sub>n</sub> velocities  $\pm$  0.1 km/s. Although formal errors are smaller, those cited here allow for unmodelable effects and biases such as hidden low- or highvelocity layers. Figure 44 depicts travel time curves and crustal models for 5 stations, representing the full range of thick (Khandyga) and thin (Yubileniya) crust, as well as both poorly resolved (Sasyr) and well resolved models (Omsukchan). Figure 45 contours the crustal thicknesses determined in this study. Table 2 compares crustal velocities determined here to results of several other studies. In most cases, Pg velocities determined here (Table 3) are slower than those determined independently above, although the general pattern of high and low velocities is consistent. Based on inverted data, Sg velocities appear to be  $3.5 \pm 0.04$  km/s. For most of the area analyzed using the inversion, Sg velocities determined above by the best fitting velocities method earlier in this section are 3.51 or 3.53 km/s, consistently faster than those determined here by the inversion method, but within this error estimate.

#### 3.2.4 Crustal Model.

The Pg and Sg velocities determined above in conjunction with relocation of earthquakes, and the  $P_n$  velocities and crustal thickness determinations from the above inversion study can be combined to establish a more complete crustal model for northeast Russia. Table 4 combines the cells in which Pg and Sg velocities are determined (Table 1) with crustal thickness and upper mantle  $(P_n)$  values interpolated from the inversion study (Figure 45).

## 3.3 TOMOGRAPHY

#### 3.3.1 Previous work.

The shallow velocity structure of northeastern Russia is essentially unstudied by tomographic methods. The only previous attempt of  $P_n$  wave tomography was done by Wallace and Tinker (1998). The study investigated the portion of Siberia between 70° E and 180° E using events recorded at 11 digital broadband stations installed in the 1990's (primarily IRIS stations). The study obtained 237 arrivals from 43 earthquakes in and around Siberia. Given the geographic size of the area and the small number of raypaths, conclusive results were not obtained.

## 3.3.2 Tomography Code.

The tomography method used an expanded form of the Time Term Method developed by Hearn (1984) and Hearn (1991). The tomography code used was written by D. McNamara for investigating  $P_n$  tomography in the Tibetan Plateau (McNamara, 1995).

In this tomography code, assumptions are made regarding average crustal thickness, average crustal velocity and average Moho velocity. From these assumptions, the static corrections (mean  $P_n$  residual) associated with each event are determined, and new residuals are calculated. The event

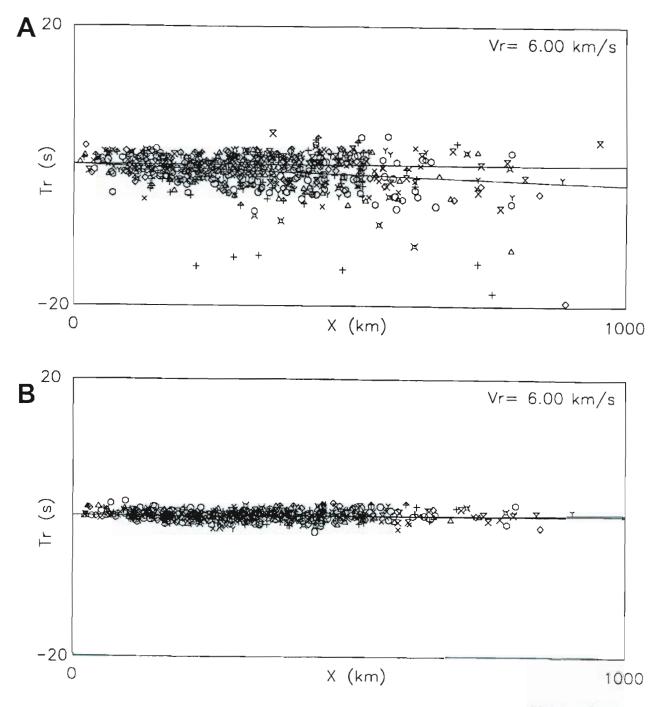


Figure 41. Reduced Pg travel time curves for determining crustal thickness. A. Using epicenters and origin times reported in *Materialy*. The best fit velocity is 6.10 km/s. B. Using relocated epicenters and new origin times from the inversion. This travel time curve indicates a Pg velocity of 5.996 km/s. Reduction velocities (*Vr*) are noted on figures. Individual events are represented by different symbols.

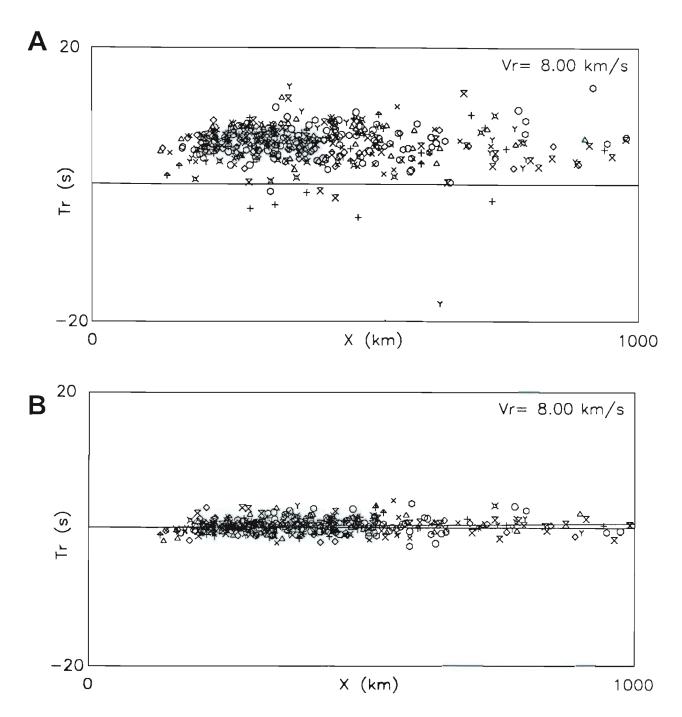


Figure 42. Reduced P<sub>n</sub> travel time curves for determining crustal thickness. A. Using epicenters and origin times reported in *Materialy*. B. Using relocated epicenters and new origin times from the inversion. This travel time curve indicates a P<sub>n</sub> velocity of 7.961 km/s. Reduction velocities (*Vr*) are noted on figures. Individual events are represented by different symbols.

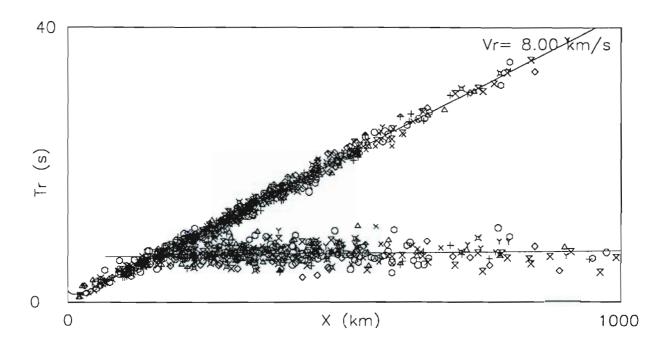


Figure 43. Reduced travel time curves for Pg and  $P_n$  data. All data plotted use the relocated epicenter and origin times from the inversion program after high residual arrivals were removed. The Pg- $P_n$  crossover point yields a regional crustal thickness of 37 km. Pg velocity plotted is 5.996 km/s, and Pn velocity is 7.96 km/s. Reduction velocity (Vr) is noted on the figure. Individual events are represented by different symbols.

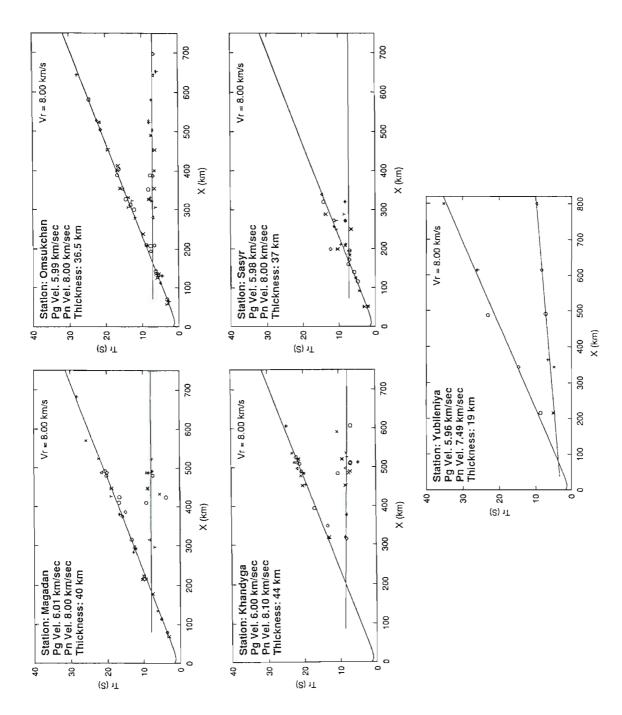


Figure 44. Travel time curves and crustal models for individual stations. For each station, Pg, Pn and reduction (Vr) velocities as well as crustal thickness are noted on figures. Individual events are represented by different symbols.

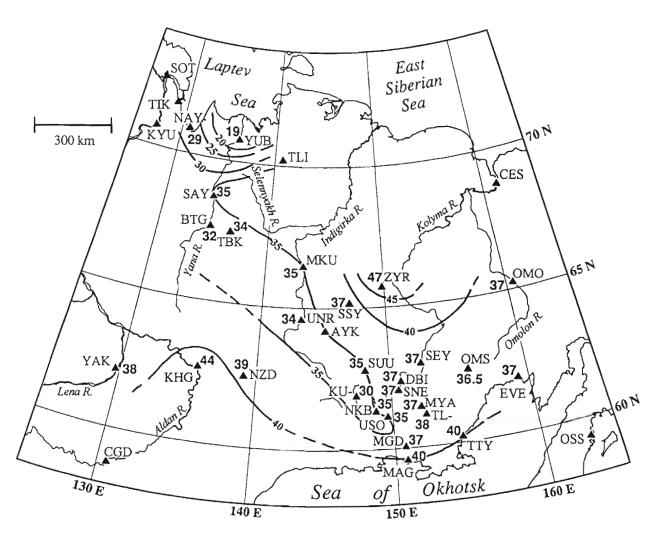


Figure 45. Summary of crustal thickness determinations for the study area contoured at 5 km intervals. Triangles denote seismic stations.

Table 2. Comparison of crustal thickness determinations for northeastern Russia

72	Station	Belyaevsky (1974)¹	Belyaevsky and Borisov (1974)	Bobrobnikov and Izmailov (1989)²	Suvorov and Komilova (1986)	Vasneniiov (1979) <sup>s</sup>	Misnin and Dareshkina (1966)	mma (1989)	and Parfenov (1985)	
ш	3atagai			6	26±3					
- (21)	svensk	343	403			33.8 (36.0)				
لا يعر	Khandyga				44±8				42	
~ ~	Magadan		38		30±2					
~	Moma				29±7					
4 4	Myakit Jaiba									
-, -	valba Vel'koba	39	40	38.9		38.9 (41.2) 38.9				
. ∠	Vezhdaninskoe		:		40±4			37	40	
J	Omolon	38	38			38.9 (40.1) 38.9				
J	Omsukchan	30	35	30.2	38±3	30.2 (31.8) 30.2				
V)	Saidy						40			
<i>S</i> 2	Sasyr									
V)	Seymchan			34.0	35±13					
<i>V</i> 1	Sinegore									
V2	Stekoľnyi	31	37	31.0		31.0 (32.9) 31.0				
<i>V</i> 2	Susuman		45 (?)	50.0	33±3					
,	Fabalakh							40		
_	<b>Fakhtoyamsk</b>									
_	Falaya									
٦	Jst Nera				24±3			40		
ب	Jst' Omchug				29±3					
~	Yakutsk				37±6				42	
_	Yubileniya									
Z	Zvrvanka							41		

<sup>1</sup>Atrobuted to work by Nikolaevsky.

<sup>2</sup>Values reported by Bobrobnikov and Izmailov (1989) attributed to Mishin and Dareshkina (1966). Bobrobnikov and Izmailov (1989) consider the value for Debin to be an error.

<sup>3</sup>Data from Belyaevsky (1974) and Belyaevsky and Borisov (1974) for Garmanda, about 20 km north of Evensk.

<sup>4</sup>Suvorov and Kornilova (1986) used Magadan as a calibration point based on the 1959 deep seismic sounding line.

<sup>5</sup>First column is attributed to Mishin and Dareshkina (1966), while values in brackets are from P-S conversions determined with a different formula.

Table 3. Pg and  $P_n$  velocities determined by inversion for specific seismic stations.  $P_n$  velocities determined by Suvorov and Kornilova (1986) are also included.

Station	Pg This Study	P <sub>n</sub> This Study	P <sub>n</sub> Suvorov and Kornilova (1986)
Batagai	5.97	7.94	8.1
Debin	5.99	7.97	
Evensk	6.01	8.07	
Khandyga	6.00	8.10	8.1-8.2
Kulu	5.98	7.73	
Magadan	6.01	8.00	
Moma	6.03	7.98	8.1
Myakit	5.99	7.90	
Naiba	5.96	7.70	
Nel'koba	5.97	7.87	
Nezhdaninskoe	5.98	7.98	
Omolon	5.98	7.98	
Omsukchan <sup>1</sup>	5.99	8.00	
Saidy	6.01	8.04	
Sasyr	5.98	8.00	
Seymchan	5.98	8.00	7.9-8.1
Sinegor'e	6.00	8.00	
Stekol'nyi	5.98	8.04	
Susuman	6.00	7.96	7.9-8.0
Tabalakh	6.01	7.94	
Takhtoyamsk	6.02	8.10	
Talaya	5.96	8.10	
Tenkeli	5.98		
Ust' Nera	5.99	7.95	8.0-8.1
Ust' Omchug	6.05	7.90	7.9-8.1
Yakutsk	6.03	8.00	
Yubileniya	5.96	7.49	
Zyryanka	5.98	8.22	
REGIONAL	5.99	8.00	

<sup>&</sup>lt;sup>1</sup> Mishin and Dareshkina (1966) calculated a Pg velocity for Omsukchan of 6.02 km/sec.

Table 4. Crustal model for northeast Russia.

Lat.	Long.		Best	Fit Velocities	(km/s)	Crustal	Thickness
48-51N	120-125E	Pg=6	.040	Sg=3.530			
48-51N	125-130E	Pg=6	.025	Sg=3.530			
48-51N	130-135E	Pg=6	.100	Sg=3.550			
48-51N	135-140E	Pg=6	.100	Sg=3.550			
51-54N	120-125É	Pg=6	.100	Sg=3.570			
51-54N	125-130E	Pg=6	.125	Sg=3.570			
51-54N	130-135E	Pg=6	.125	Sg=3.570			
51-54N	135-140E	Pg=6	.075	Sg=3.550			
54-57N	120-125E	Pg=6	.275	Sg=3.630			
54-57N	125-130E	Pg=6	.175	Sg=3.590			
54-57N	130-135E	Pg=6	.150	Sg=3.590			
54-57N	135-140E	Pg=6	.100	Sg=3.570			
57-60N	120-125E	Pg=6	.275	Sg=3.630			
57-60N	125-130E	Pg=6	.300	Sg=3.650			
57-60N	130-140E	Pg=6	.275	Sg=3.630			
57-60N	140-145E	Pg=6	.075	Sg=3.570			
57-60N	145-150E	Pg=5	.950	Sg=3.510			
57-60N	150-155E	Pg=6	.040	Sg=3.510	$P_{a} = 8.05$	40	km
57-60N	155-160E	Pg=6	.025	Sg=3.510			
60-63N	135-140E	Pg=6	.050	Sg=3.570	$P_n = 8.04$	42	km
60-63N	140-145E	Pg=6	.050	Sg=3.530	$P_n = 7.97$	38	km
60-63N	145-150E	Pg=6		Sg=3.530	$P_{n} = 7.92$	35	km
60-63N	150-155E	Pg=6	.025	Sg=3.510	$P_n = 7.99$		km
60-63N	155-160E	Pg=6	.06	Sg=3.51	$P_{n} = 8.06$	38	km
60-63N	160-165E	Pg=6	.05	Sg=3.51			
60-65N	165-170E	Pg=5	.975	Sg=3.470			
60-65N	170-180E	Pg=6	.000	Sg=3.470			
60-66N	120-135E	Pg=6	.225	Sg=3.610	$P_n = 8.05$	38	km
63-66N	135-140E	Pg=6	.050	Sg=3.550	$P_{n} = 7.99$	36	km
63-66N	140-145E	Pg=6	.050	Sg=3.530	$P_{n} = 7.96$		km
63-66N	145-150E	Pg=6	.025	Sg=3.510	$P_{\rm n} = 8.06$	39	km
63-66N	150-155E	Pg=6	.025	Sg=3.530	$P_{n} = 8.06$	39	km
63-66N	155-160E	Pg=6	.025	Sg=3.510	$P_{n} = 7.99$	37	km
63-69N	180-160W	Pg=6	.050	Sg=3.530			
65-69N	165-180E	Pg=6	.060	Sg=3.530			
66-69N	125-130E	Pg=6	.275	Sg=3.610			
66-69N	130-135E	Pg=6		Sg=3.570	$P_n = 7.99$	34	km
66-69N	135-140E	Pg=6		Sg=3.550	$P_{n} = 7.9$	7 34	km
66-69N	140-145E	Pg=6		Sg=3.570	$P_n = 7.98$	3 35	km
66-69N	145-150E	Pg=6	.040	Sg=3.53	-		
69-72N	125-135E	Pg=6	.050	Sg=3.550	$P_n = 7.74$	1 24	km
69-72N	135-145E	Pg=6	.200	Sg=3.590	$P_{n} = 7.74$	1 24	km
72-75N	120-130E	Pg=6		Sg=3.570			
72-75N	130-135E	Pg=5	.900	Sg=3.510			

delays calculated are dependent on crustal thickness and velocity variations, as well as errors in hypocenter depth and origin time. The individual effects of these parameters will trade off with one another, thus the actual event statics cannot be interpreted in a meaningful way (Hearn et al., 1991).

Static corrections for receiving stations are next determined using the mean residuals for each station and new residuals are again calculated. Static corrections for the stations are primarily dependent on crustal velocity and crustal thickness. Crustal thickness can affect station statics the most (about 1 second per 10 km change in crustal thickness; Hearn et al., 1991), thus they can be interpreted as such.

Finally, cell slownesses are estimated from the weighted mean of the apparent slowness of all rays passing through each cell. The weighting for slowness in a given cell is the product of the rays traversing the cell, adjusted for the fraction of each rays total length that is included within the cell. In each iteration, the full residual remaining after the static corrections is applied to the Moho leg of the travel path. To remove any cells with extreme slowness values due to poorly sampled cells, the model is smoothed between each iteration by averaging each cell with its eight neighbors. This is, in effect, the only dampening that occurs in the process. Once the new model is determined, an updated average Moho velocity for the entire study region is calculated. New residuals are again calculated based on the updated average Moho velocity and the process beginning with event static corrections is repeated until the data converge. Iterations are usually stopped when the change in RMS residuals from one iteration to the next becomes less than 1 % (McNamara et al., 1997).

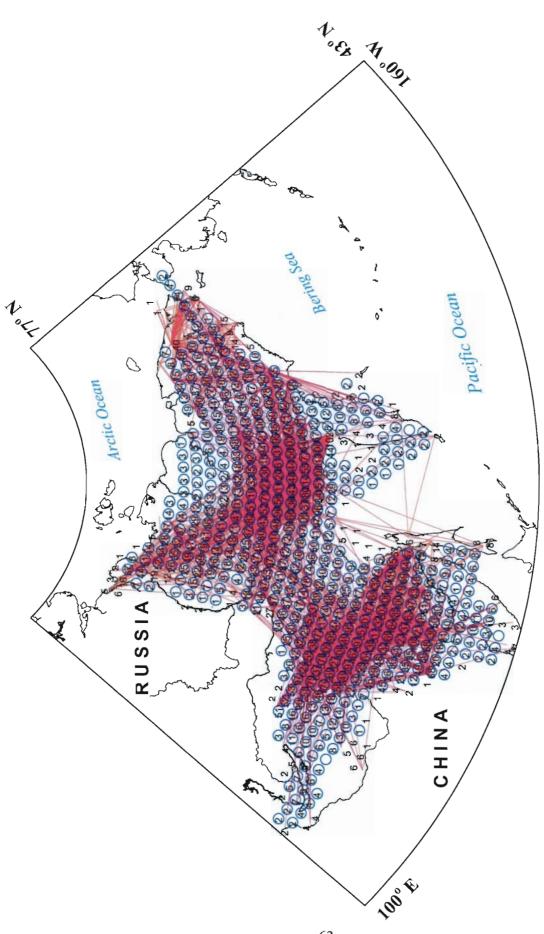
#### 3.3.3 Results.

Considerable effort went into developing reasonable tomographic models for Northeast Russia using the code of McNamara (1995). Attempts were made to calculate models using the original Russian determined epicenters for the entire study area as well as relocated epicenters from Section 3 above for both the entire study area and a much smaller Magadan-only region, where raypath coverage was much more uniform.

For all model attempts, consistent results with converging residuals and reasonable static corrections were not achieved. Throughout the effort to develop the models, several input/output errors of the tomography code were discovered and workarounds were successfully developed. This still did not result in consistent models. Specific models developed in this process are discussed at length in Mackey (1999), but not discussed individually here due to the final determination of a negative result. The following illustrate the software variables used and changed in the attempts to develop reasonable models:

- Varying the range of acceptable velocities for data selection
- Changing the distance range for acceptable data
- Changing the number of smoothing passes between each iteration
- Changing the minimum number of P<sub>n</sub> phases per event or station for usable data
- Varying the cell size used in the model
- Removal of some receiving stations
- Varying the assumed average crustal thickness
- Varying the assumed average Moho velocity

It was ultimately determined after looking at secondary outputs of the tomography code that there are severe indexing problems with the handling of data in the code used (McNamara, 1995). Figure 46 illustrates the apparent indexing problems with the code. The figure illustrates a mismatch between the location of raypath traces, the cells where velocity perturbations were



hits and the number of hits. Blue circles indicate cells where velocity perturbations were calculated. Note that none of Figure 46. Raypath coverage (red lines) for the tomography study using relocated hypocenters. Black numbers indicate cells with the three show exactly the same coverage, particularly around the edges of the raypath coverage.

calculated, and the number of raypaths recorded as hitting each cell. For example, some cells clearly have raypaths penetrating them, but record no hits, or miscount them, while other cells have no raypaths crossing them, but record hits and calculate perturbations.

It is remotely possible that some of the indexing problems may be a result of whether the code maps raypaths assuming a flat earth or as great circles. At high latitudes it is necessary to use great circle paths, but the author of the code was not certain if this was being done. Attempts to isolate the problem in the source code were not successful.

## 3.4 CONCLUSION.

The earthquake relocation process used here has improved the quality of hypocenter locations over the original Russian determinations, as well as developed regionally calibrated, best-fit crustal velocity models. Regional crustal velocities determined are consistent with the known geologic and tectonic setting. Improvements in epicenter locations have clarified several seismicity clusters and fault lineations in northeast Russia. Teleseismically recorded events will be further evaluated and assigned Ground Truth (GT) classifications (Section 4). Development of tomographic models for northeastern Russia was attempted, but not successful. It is believed that computer code problems contributed to this result.

#### **SECTION 4**

## GT CLASSIFICATIONS AND ANALYSIS OF TELESEISMIC EVENTS

#### 4.1 INTRODUCTION.

In effort to obtain ground truth (GT) classifications in support of CTBT monitoring for continental regions of northeastern Russia, relocations were computed 134 seismic events reported in the International Seismological Centre (ISC) or Earthquake Data Report (EDR) catalogs which occurred from 1970 through 2000 (Figure 47). ISC solutions for this region utilize data from very few local or regional stations. The patterns of arrival time residuals for several stations were plotted and found to be generally consistent with the regional geologic setting. In essence, the travel time residuals found here are mapping the upper mantle velocity structure, which compliments the crustal model as outlined in Section 3.

## 4.2 METHODOLOGY.

Local and regional arrivals from Russian regional seismic networks in Kamchatka, Magadan, Yakutsk, Amur, Sakhalin, and Irkutsk were used to supplement ISC and EDR data. For many of the events, selected original seismograms were inspected to obtain arrival times not otherwise recorded. Often this was for stations located in one network for an event which occurred within the area covered by another network, since data were not normally exchanged between networks.

In the process of combining local network data with ISC arrivals, it was found that for stations up to 10° distance, phases reported in ISC are sometimes incomplete or misidentified. The most common misassociation is the secondary Pg or Sg phase identified as a P<sub>n</sub>/P or S<sub>n</sub>/S arrival. This misidentification of phases can result in poor hypocenter determinations by ISC, particularly for events with few receiving stations or poor azimuth distribution to stations. These errors arise due to the fact that arrival times and phases reported to the ISC for many of the Russian stations are the result of preliminary data analysis. The local data bulletins contain the final analysis of arrival times and phase associations. Also, in many cases, ISC reported times are rounded to the nearest second, while local bulletins report arrivals to a tenth of second. In both cases, local bulletin data take preference over that from ISC for relocations computed in this study. The inconsistencies in local data reported by ISC are significant in that they illustrate the problem of relying on international bulletins to develop local or regional velocity models for improved monitoring of the CTBT.

Throughout the study area, many geologic and tectonic environments exist. Combined with the physical vastness of the region, this makes it likely that no single velocity or travel-time curve will reflect actual seismic velocities for any particular phase. In order to overcome this problem, the study area was broken into cells, and the best fitting velocities were determined. The best fitting velocities were then used in the relocation of ISC reported events. Full details of the methodology and procedures used in the relocation of earthquakes are discussed extensively in Mackey (1999), and Section 3 above.

Each event was relocated three times using combined local and teleseismic arrivals, only arrivals within 20° and only arrivals within 10°. To best evaluate which set of data produce the best locations, particular attention was paid to aftershock sequences, and how closely events cluster using

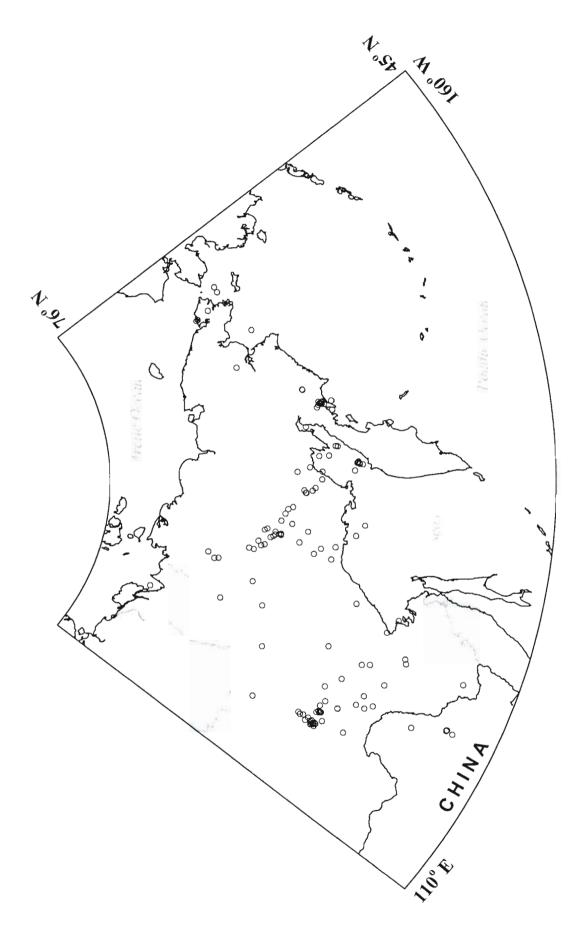


Figure 47. ISC determined for locations for teleseismic earthquakes analyzed in this study.

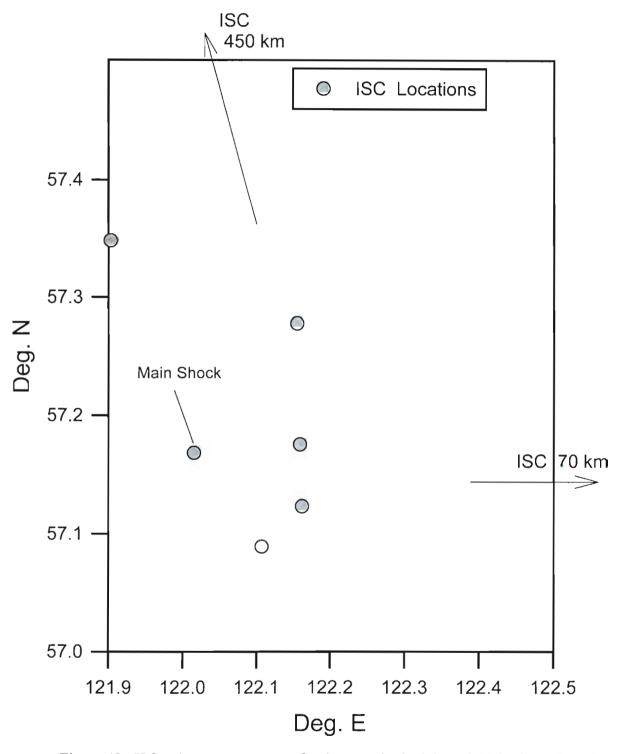


Figure 48. ISC epicenter parameters for the magnitude 6.6 south Yakutia earthquake and aftershocks. Note that 2 events are off the map.

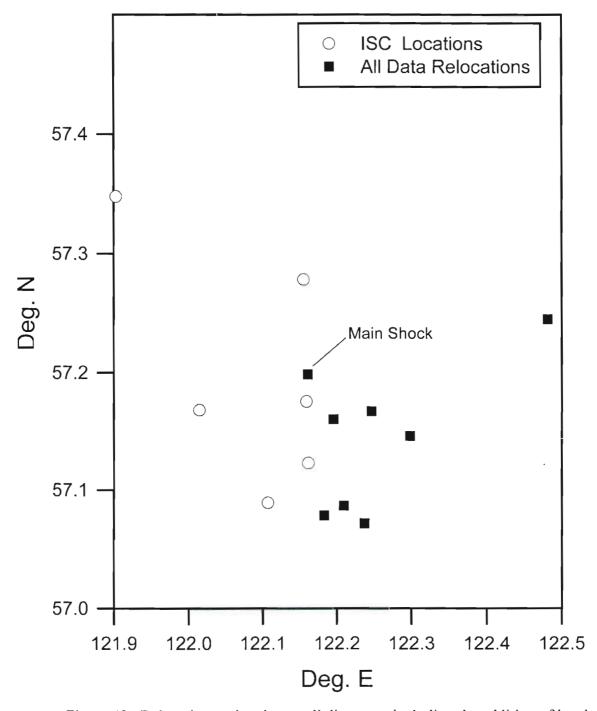


Figure 49. Relocations using data at all distances, including the addition of local and regional stations not reported in ISC. ISC locations shown for reference.

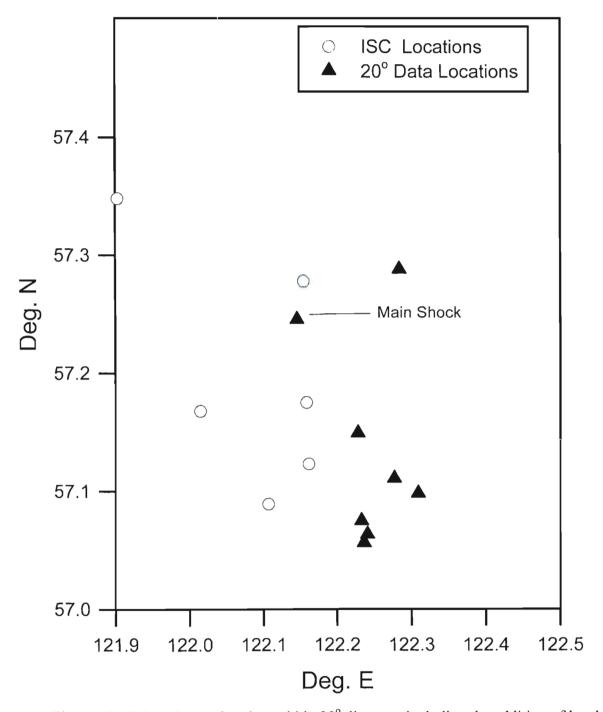


Figure 50. Relocations using data within 20° distance, including the addition of local and regional stations not reported in ISC. ISC locations shown for reference.

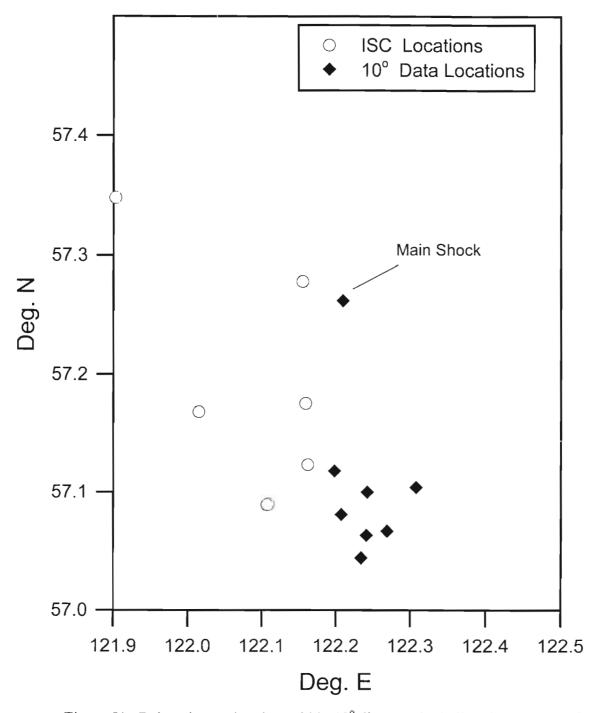


Figure 51. Relocations using data within 10° distance, including the addition of local stations not reported in ISC. ISC locations shown for reference.

the different data sets. Figures 48-51 show the differences in locations determined for an aftershock sequence from south Yakutia in 1989 (see South Yakutia Sequence section below). In this sequence, locations computed using only data either within 10° or 20° cluster best, although locations including local and teleseismic data are clearly improved over those from ISC alone. In general, for events with good azimuthal coverage of stations within 5°, the best locations are those computed with only the local (<10°) and regional (<20°) data, while events having poor azimuthal coverage of close-in stations are better located by including teleseismic data. Based on the general distribution and residuals of stations, a preferred location was selected for each event analyzed (Figures 52 and 53; Appendix C).

All relocated events are assigned GT levels, with 26 events meeting or exceeding the GT10 criteria (Figure 54). Criteria used for acceptance of GT10 events state that the event must be recorded by a minimum of five stations within 3° distance, and that for stations within 5°, the largest azimuthal gap must be 180° or less (Yang and Romney, 1999). An example of residuals and station distribution for a GT10 event is presented in Figure 55 and Table 5. A few events which formally met these criteria were excluded from GT10 classification due to high station residuals for some close in stations, or other odd station distribution problems which can result in poor epicenter control. It should also be noted that many relocated events miss the formal GT10 criteria, but it is suspected that the location determined here is within 10 km of the true epicenter based on proximity to fault traces, etc. (see discussion below). Most of the remaining relocated events meet GT25 criteria. Although no events formally qualify as GT5, many events are likely within that accuracy.

Overall, compared with ISC locations, events occurring within aftershock sequences cluster tighter, and event alignment with known active faults improves.

#### 4.3 RESULTS.

# 4.3.1 Comparison with ISC Epicenters.

134 events listed in the ISC bulletin were relocated incorporating regional data. Most (90%) of the relocations differ by less than 45 km relative to the ISC parameters, with a few being significantly different (one event with poor station coverage in the ISC moved 500 km). As noted, the new locations are likely to be an improvement due to the large number of local and regional phases utilized. However, there is no large-scale spatial pattern to the azimuth or magnitude of change in epicenters calculated by ISC relative to these relocations (Figure 53).

Although there are some areas where there appears to be some consistent directional mislocation bias (e.g., the larger events in Chukotka, the Laptev Sea, and also perhaps in the north-central Verkhoyansk Range), the most consistent bias in the epicenter shift is "radial," with aftershocks and other events tending to move into tighter clusters or into lineaments (e.g., central Chersky Range). No specific areas stand out as having large differences between ISC and solutions determined here. Examination of events whose locations have the greatest differences, by more than 45 km, suggests that the relocations are much better than the ISC solutions, aside from a few exceptions at the edges of the study area.

ISC solutions are generally worse where they have no close stations or the stations fall into a narrow azimuthal range. Similarly, a few of relocations at the edges of the study area are worse because of a narrow azimuthal range without using teleseismic data; this may also have affected the travel time curve calibrations for these areas (Section 3).

An area where the relocations show clear improvement is Chukotka, where the use of both Russian and American data provides better azimuthal coverage. Throughout the remainder of the

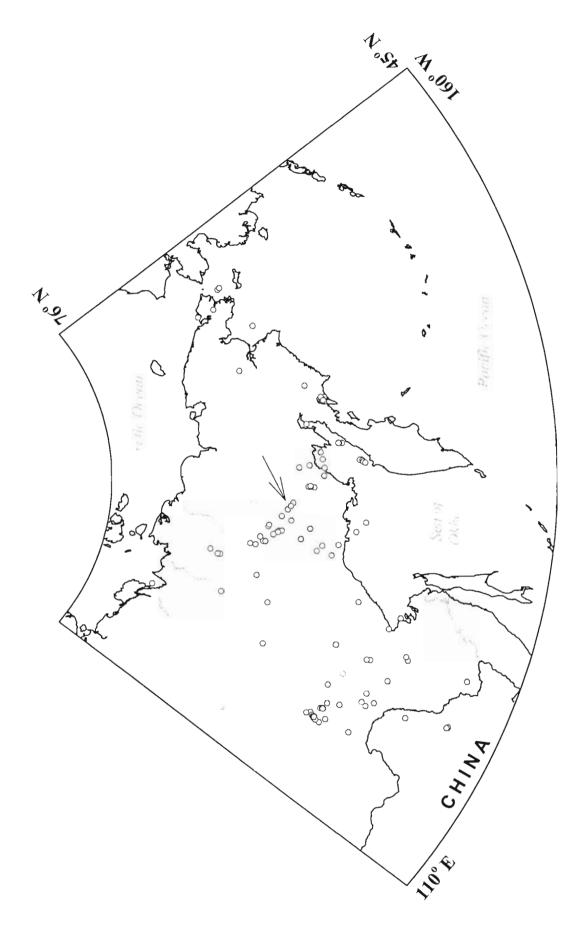


Figure 52. Relocations determined in this study using ISC data supplemented with local and regional time picks. Arrow indicates linear trend associated with the Ulakhan fault,

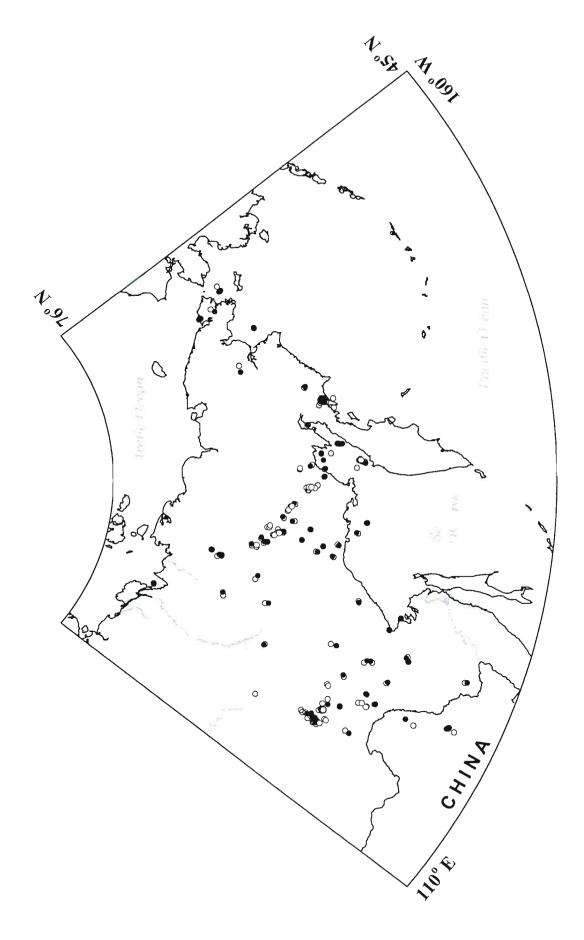


Figure 53. Comparison of ISC epicenters (open circles) and relocations (filled circles). Gray circles represent relocations that meet or exceed GT10 criteria. Note the relocations cluster tighter and form better lineaments near the center of the

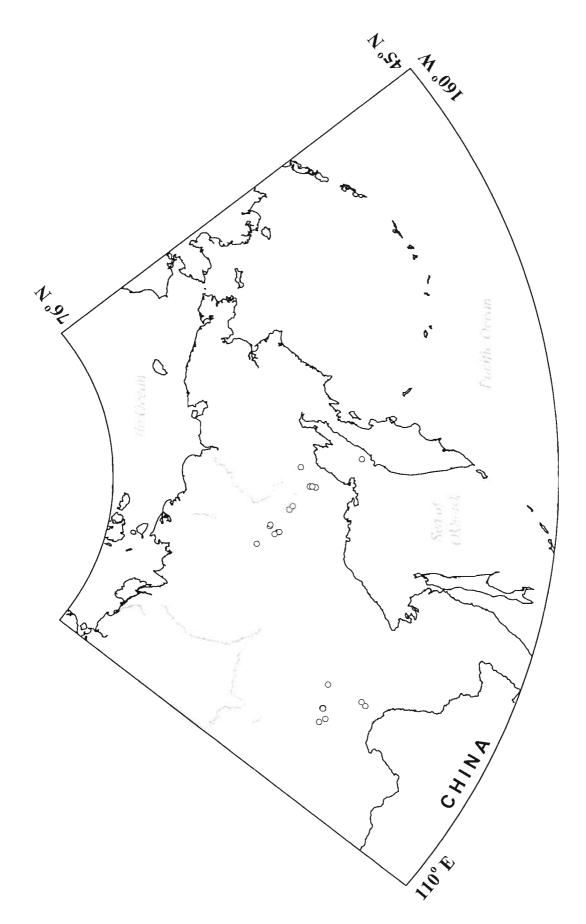
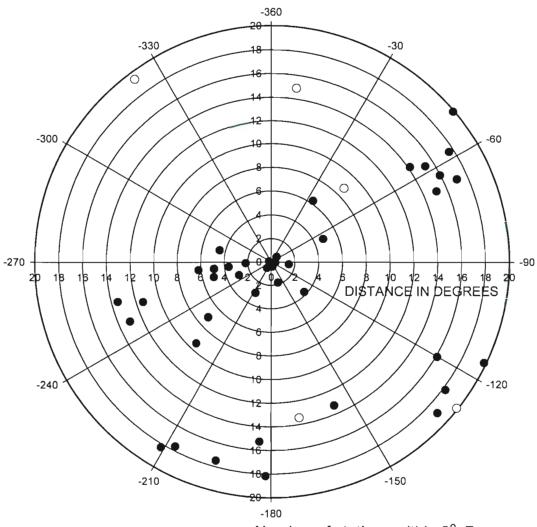


Figure 54. GT10 events.



Origin time: 05 04 35.84 utc

Latitude: 57.057° N Longitude: 122.237° E

Depth: 30.7 km

Magnitude: 5.6

Number of stations within 2°: 7 Number of stations within 3°: 10

Largest Azimuthal gap within 5°: 117.5°

Classification: GT-10

Figure 55. Distribution of stations used in the 20° distance relocation of the May 17, 1989 event. Stations shown in gray were omitted from the location computation due to high residuals (see Table 5).

Table 5. Relocation output of the May 17, 1989 (05:04 UTC) earthquake using only stations within 20 degrees distance. Arrivals in bold are those added in this study, supplementing ISC listings. See also Figure 55.

NUMBER	OF ARR	IVAL	S =	= 56				
	RIGIN T				DEG±KM)	LONG. (DEG±KM)		RMS RESIDUAL
	04 35.8				500	122.2400		
1 05	04 35.8	4±.1	7	57.0	560±1.62	122.2367±1.24 122.2365±1.23	30.69 ±1.96	
2 05	04 35.8	$4\pm.1$	7	57.0	562 <u>+</u> 1.62	122.2365±1.23		
						122.2365±1.23		
							30.69 ±1.96	
5 05	04 35.8	4±.1	7	57.0	562 <u>±</u> 1.60	122.2365±1.23	30.69 ±1.96	. 84
STATIO	N PHASE	Т	IMI	Ξ	DISTA	NCE AZIMUTH	RESIDUAL	WEIGHT
SYL	PG	5	4	42.2	. 2		.24	1.0
ACH	PG	5	4	44.4	.3	5 97.35	.57	1.0
KBK	PG	5	4	44.5	.4	165.34	.02	1.0
YRG	PG	5	4	49.0	.63	3 45.80	.94	1.0
USZ	PG	5	4	47.6	. 6	1 215.82	22	1.0
CLNS	PG	5	5	3.2	1.4	97.33	.66	1.0
CLNS	SG	5		22.0	1.4	97.33	02	.5
UUR	P	5		6.0	1.8	162.30	.13	1.0
CRS	P	5		12.1	2.1		1.45	1.0
SRK	P	5		22.0	2.9		.32	1.0
TUP	IP	5		21.5	2.9		04	1.0
TUP	PG	5		27.5	2.9		71	1.0
TUP	SG	5		5.5	2.9		98	.5
NLY	ΙP	5		32.0	3.64		.54	1.0
KIRS	₽	5	5	32.8	3.7		33	1.0
BOD	P	5	5	44.0	4.5		.41	1.0
BOD	PG	5		54.7	4.5		-1.06	1.0
BOD	SG	5		50.0	4.5		-3.98*	. 0
CGD	IP	5		48.5	4.7		.91	1.0
SVK	IP	5		50.0	4.8		.83	1.0
SVK	PG	5		6.0	4.8		3.35*	. 0
UKT	IP	5		51.7	5.04		.26	1.0
UKT	PG	5		5.5	5.04		.04	1.0
YAK	P	5		6.7	6.25		-1.45	1.0
YAK	PG	5		26.4	6.2		23	1.0
YAK	SG			46.0	6.2		-1.40	. 5
KMO	P	5		8.1	6.2		.19	1.0
CIT	P	5		18.3	7.1		-2.35	1.0
CIT	PG			46.0			3.69*	.0
CIT	SG	5		14.0	7.13		51	.5
KHG	P	5		20.5	8.73		-22.28*	.0
KPC	P	5		51.0	9.43		-1.40	1.0
IRK	P	5		20.0	11.42		.27	1.0
ZAK	P	5		42.5	13.03		1.06	1.0
MDJ	P	5		44.0	13.28		62	1.0
MOY	P	5		47.3	13.47		. 07	1.0
CN2	P	5		43.0	13.43		-3.49*	. 0
SUUS	P	5	7	54.9	14.14	55.37	85	1.0

Table 5 (Cont.)

TIK	P	5	8	1.0	14.90	8.19	-4.56*	.0
DBI	P	5	8	10.3	15.25	57.89	.10	1.0
SNY	IP	5	8	10.0	15.27	176.19	53	1.0
MGD	P	5	8	8.1	15.11	66.67	34	1.0
TL-S	EP	5	8	18.3	15.94	62.62	79	1.0
YSS	Þ	5	8	22.0	16.09	120.12	.91	1.0
TTY	EP	5	8	33.8	17.08	65.82	.40	1.0
BJI	EP	5	8	38.0	17.48	195.66	58	1.0
OMS	EP	5	8	39.4	17.59	58.01	46	1.0
HHC	P	5	8	40.5	17.63	207.64	13	1.0
DL2	IP	5	8	48.0	18.17	181.52	.90	1.0
BTO	P	5	8	50.0	18.28	210.92	1.34	1.0
ASAJ	EP	5	8	50.0	18.21	126.66	2.45	1.0
MRRJ	ΕP	5	8	56.4	18.94	132.58	21	1.0
NRI	P	5	8	55.3	19.38	323.22	-5.89*	. 0
KUR	EP	5	9	5.0	19.83	115.67	-1.40	1.0
HOOJ	EP	5	9	4.6	19.90	128.56	-2.61*	.0
OMO	EP	5	9	7.3	19.91	50.06	.11	1.0

<sup>\*</sup> Arrivals excluded from the location solution due to high residuals.

study area the aftershock relocations are better, with solutions falling in closer proximity to the mainshock, a result of a better data set using larger numbers of close-in stations. Because the magnitude of the largest event does not exceed 7.0, aftershock areas are not expected to be large.

The area where relocation solutions may have some problem is in northern Yakutia and the Laptev Sea. Here, if only local and regional arrivals are used, there are no stations to the north, resulting in poor azimuthal. Rapid variations in crustal seismic velocities in the Laptev Sea region are suspected of further degrading the quality of the relocations (see Section 3). Because of poor azimuth distributions and poorly resolved crustal velocities, teleseismic earthquakes occurring in the Laptev Sea are excluded from this study even though additional local and regional data are available.

# 4.3.2 South Yakutian Sequence.

Many of the best located events are from the aftershocks of the South Yakutian earthquake of 1989 (M<sub>w</sub> 6.3-6.4) where five temporary stations were deployed by the Yakut regional network in the aftershock zone (Koz'min et al., 1993). Figure 56 shows relocations of the main shock and seven aftershocks based on local, regional, and all (including teleseismic) data. For these events, the regional (<20°) relocations are suspected to be best, with the exception of one aftershock, where the local (<10°) location is better. The relocated aftershocks cluster in a zone with dimensions of about 7 x 10 km, with the mainshock located approximately 12 km northwest of the aftershocks (Figure 56). Since all seven aftershocks meet GT10 criteria, based on the tight clustering of the aftershocks, it is likely that the accuracy of these solutions may be better. The data for one of the aftershocks is presented in Table 5 and the variation in solutions for that event is shown in Figure 57 (see Figure 55 also). In contrast, three of the original ISC epicenters are located 30, 90, and 500 km from the cluster (Figures 48 - 51) and the remaining four aftershock solutions from ISC are offset about 10-15 km to the northwest of the relocated cluster. In general, the ISC solutions that differ greatly from the relocations are probably the result of poor time picks or little data, while the four events showing the 10-15 km offset probably better represent a true location bias in the ISC solutions. The aftershock sequence from this event continued for many months and over three-thousand events were recorded.

The relocation of the mainshock was somewhat problematic. There was considerable variability between the solutions and all solutions tended to be 15 to 20 km to the north of the aftershock cluster and away from inferred faults in the area; this was most pronounced for the regional solution (Table 6). However, this may be due to the higher seismic velocities along ray-paths to stations on the northern edge of the Siberian platform (see discussion on travel-time variations, below). Exclusion of these stations yielded more reasonable solutions and the mainshock and three of the aftershocks align along a possible northwest-southeast striking fault (Figure 56). The remaining four aftershocks cluster about 5 km to the south.

#### 4.3.3 Ulakhan Fault.

All of the relocated events in the Chersky Range relocate into essentially straight lines which parallel the Ulakhan and Chai-Yureya faults. These faults are among those thought to represent major breaks along which the strain between the Okhotsk and North American plates are partitioned (Imaev et al., 1990, 1994). The Ulakhan fault (Figure 58) is well defined in the topography between the Indigirka and Kolyma Rivers. It shows clear river offsets (Figure 59) at several levels, but with a probable maximum of 40 km of left-lateral movement since the establishment of the river network (inferred as mid-Pliocene based on Grossheim and Khain, 1967).

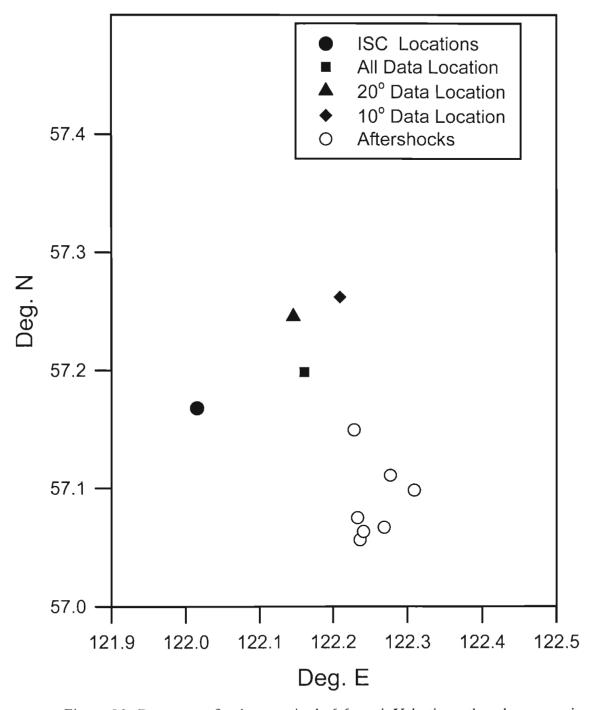


Figure 56. Parameters for the magnitude 6.6 south Yakutia earthquake comparing ISC parameters with relocation solutions calculated using different data sets. Preferred locations of aftershocks are shown in grey.

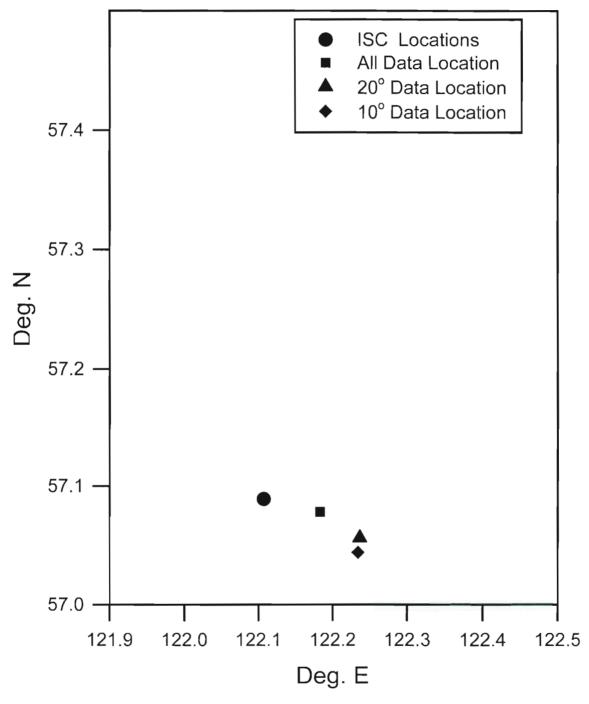


Figure 57. Epicenter calculations for the May 17 aftershock of the south Yakutia event comparing ISC parameters with relocation parameters using different data sets. The 20° solution is preferred. See also Table 5 and Figure 55.

Table 6. Comparison of 1989 South Yakutia Earthquake mainshock solutions. Bold entries in the table are relocations determined in this study.

Solution	Date	Time	Lat.	Long.	Depth
All Data	1989 04 20	22 59 53.39	57.198	122.161	$22.5 \pm 2.5$
Regional	1989 04 20	22 59 53.63	57.245	122.146	$29.9 \pm 4.1$
Local	1989 04 20	22 59 53.52	57.262	122.209	$32.1 \pm 6.0$
ISC	1989 04 20	22 59 54.24	57.168	122.015	$26.0, 32 \pm 2.1 \text{ pP}$
Yakut Net	1989 04 20	22 59 54.8	57.17	122.31	27
NEIC	1989 04 20	22 59 54.0	57.17	121.98	26
Moscow	1989 04 20	22 59 54.8	57.19	122.08	32 sP, 35 pP
Doser (1991)	1989 04 20		57.08	122.12	$24 \pm 4$ synth
Harvard CMT	1989 04 20	23 00 01.4	57.03	121.23	$48 \pm 1.8$
Beijing	1989 04 20	22 59 53.6	57.15	122.44	35

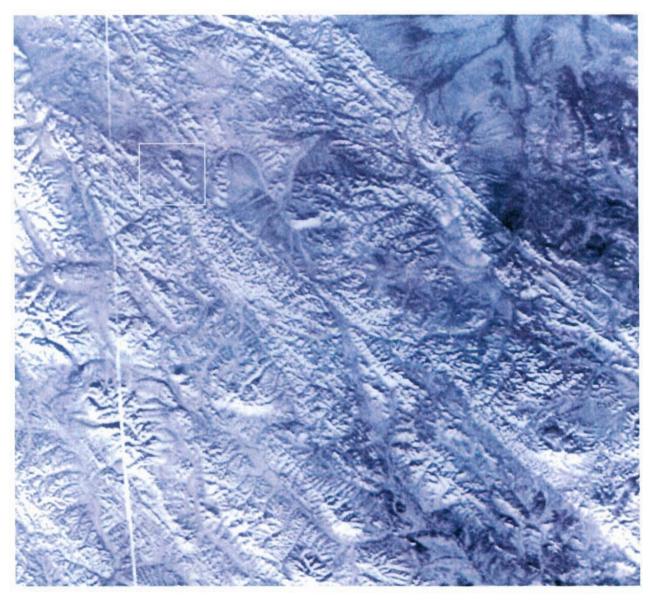


Figure 58. Meteor satellite image of the central portion of the Ulakhan fault, extending from the lower right corner to the upper left corner of the image. This image depicts a section of the fault several hundred kilometers in length. Box indicates location of map in Figure 59.

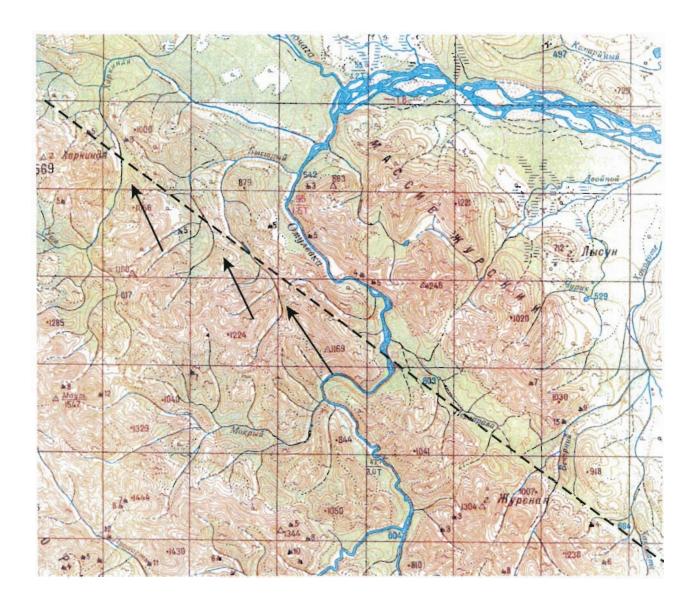


Figure 59. Topographic map of a portion of the Ulakhan fault, depicted by the dashed line. Note large river offsets showing a left lateral sense of motion along the fault (arrows). One square on this map represents approximately 4 km. The base map used here is a 1:200,000 Russian military topographic map.

This age and offset is consistent with current estimates of motion of about 0.8-1.0 cm/yr (Imaev et al., 1990; Seno et al., 1996).

The seismicity along the Ulakhan can be broken into two linear segments to the east and west of the Darpir-Yuryakh River, with a very slight change in strike. While these segments closely follow the surface trace of the Ulakhan fault, they remove the arcuate curve to the north that is observed in the surface topography. The seismicity is furthest offset from the surface expression of the Ulakhan to the east of the Darpir-Yuryakh River. This offset region is associated with the presence of Early Paleozoic rocks at the surface. To the west, Parfenov et al. (1989) have interpreted the Paleozoic rocks as being rootless and representing a large klippe. If this is also true in the Omulevka Mountains, it is possible that the main trace of the Ulakhan fault is in fact straight at depth below the Paleozoic and offset from the surface expression of the fault, which is located at the edge of the Paleozoic. This conjecture is supported by the fact that the microseismicity becomes more diffuse through the region of the Omulevka Mountains (McLean et al., 2000). In any case, except at the point of maximum deviation of the seismicity trend from the surface expression of the fault, all events are within 10 km of the mapped fault, and in several cases within 5 km. A second satellite image, slightly to the northwest of Figure 60 shows another section of the Ulakhan fault and the proximity of two relocations to the fault. A third event is further north on a parallel fault (Figure 60).

The Chai-Yureya fault is thought to represent a splay off the Ulakhan and strikes in a more southeasterly direction just east of the town of Artyk. Relocations of seven events that fall along the Chai-Yureya; a straight line can be fit with an error of less than two km through six of the events (the seventh is the most poorly located event of the set). Relocated epicenters also fall almost directly on the surface expression of the Chai-Yureya fault, with a maximum deviation among the GT10 and GT25 events of 6 km from the trace, suggesting that epicentral locations are accurate to 10 km or better (Figure 52). One of the events is a magnitude 7 earthquake, the largest in the Chersky Range, which occurred along this fault in 1971. This event, while nominally GT25, falls almost directly on the surface expression of the fault.

## 4.3.4 Travel-time variations in northeast Russia.

To investigate further improvement in hypocentral locations by determining Source Specific Station Corrections (SSSC's), travel-time residuals relative to JB for nine regional stations from events in northeast Russia were investigated. The stations have epicentral distances from the events of about 1 - 20°.

The results for Tiksi (TIK, TIXI; Figure 61), based on all the events studied, are characteristic. High negative residuals, indicating fast velocities, are observed from events in south Yakutia. These rays pass through sub-Siberian shield mantle, which is expected to have higher velocities. Slightly negative residuals are obtained from the Chersky Range, indicating some higher velocity upper mantle exists in the region of the Kolyma-Omolon superterrane. Events from the eastern Sea of Okhotsk and Koryakia have slightly positive residuals, and suggest there may be slower mantle due either to lithospheric rifting (Penzhina and Parapol basins; Worrall et al., 1996) or to effects of the supra-subduction mantle wedge. Similarly, a few events with slightly positive residuals in northeast Amur District may be related with extension along the western Sea of Okhotsk (Shantar-Liziansky basin) also proposed by Worrall et al. (1996). One event in the Yana River valley gives a high positive residual, which may be related to the present or recent extension of the Arctic Mid-Ocean ridge into the continent. The restricted number and distribution of GT10 events provides only information from southern Yakutia and the Chersky Range, but leave no doubt as to the faster velocities to Tiksi from those two regions. Travel paths under the Siberian platform to

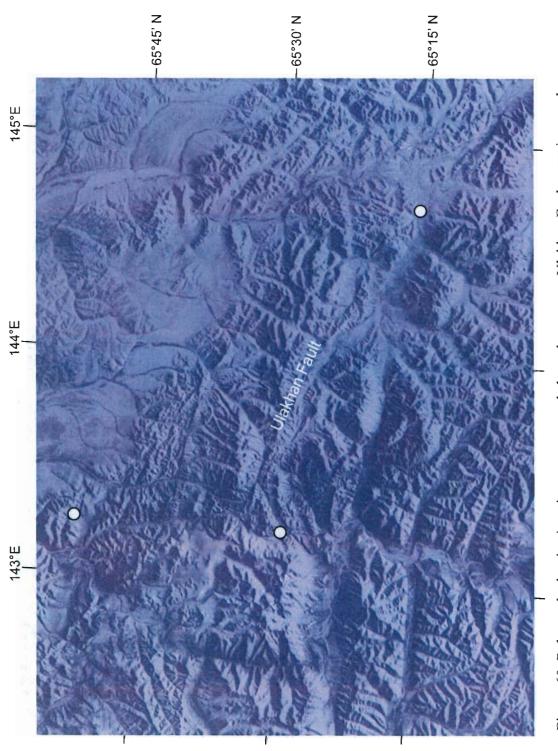


Figure 60. Relocated teleseismic epicenters near and along the western Ulakhan Fault superimposed on a Landsat image. Epicenters relocate to within five to ten km of the fault. The third epicenter epicenter locates along a parallel fault 30 km north of the Ulakhan.

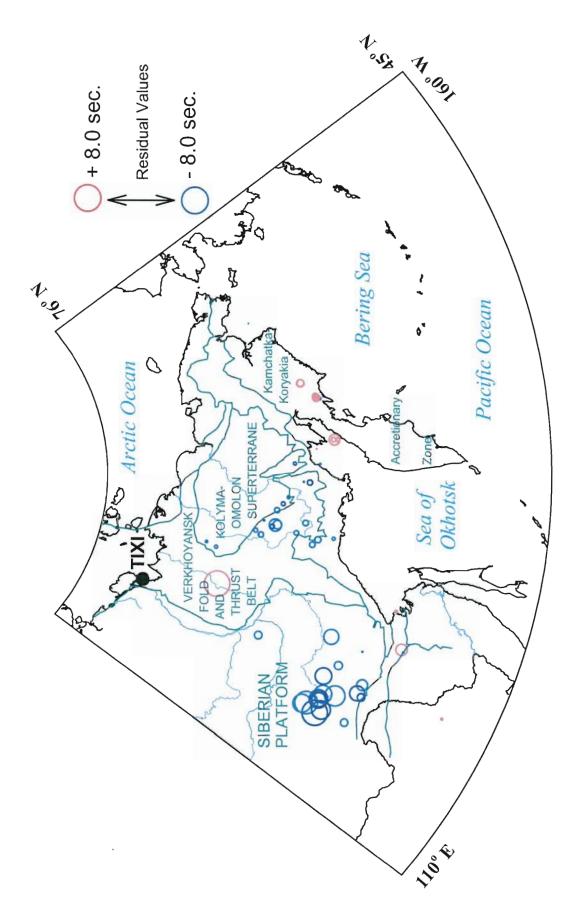


Figure 61. Tiksi (TIK, TIXI) residuals for all GT analyzed events. Solid colors are from GT10 events while the lighter shade represents GT25 or poorer classification. Note the consistently large negative residuals for paths associated with the Siberian platform, and lower residuals elsewhere. Residuals are especially consistent among the GT10 events and essentially represent Source Specific Station Corrections (SSSCs)

Norilsk also yield consistently negative residuals of a magnitude similar to those at Tiksi (Figure 62).

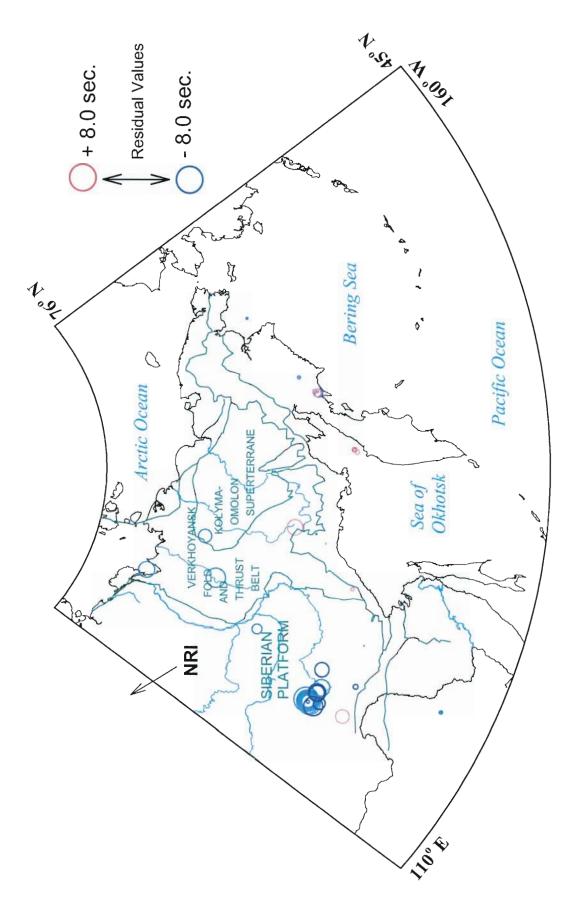
Arrivals at Yuzhno Sakhalinsk confirm that ray paths through the sub-northern Sea of Okhotsk mantle yield high positive residuals (Figure 63). Events from the Amur region with origins in the Siberian platform west of the proposed rift again yield slightly positive residuals, suggesting that their respective higher and lower velocities tend to cancel each other out. All arrivals at Petropavlovsk are late, indicative of passage through the supra-subduction zone upper mantle (Figure 64). Similar to Yuzhno Sakhalinsk, as events occur farther into the higher velocity Siberian platform relative to Petropavlovsk residuals become less negative.

Bilibino also yields slightly positive residuals from the northern Sea of Okhotsk and negative residuals for the Kolyma-Omolon superterrane and Koryakia, suggesting that upper mantle velocities in far northeast Russia are slightly fast (Figure 65). Peleduy, however, yields late arrivals from the Kolyma-Omolon superterrane. Arrivals from the northern Sea of Okhotsk are again late, and early arrivals are observed on Siberian platform crossing paths from the Verkhoyansk Range (Figure 66).

Residual patterns for Seymchan and Magadan are similar (Figures 67 and 68). The arrivals from Shelikhov Bay (northeast Sea of Okhotsk) and the Amur region are again all late, while a mix of residuals is observed from the Chersky seismic belt. From these stations, Chukotka paths are slightly slow. Finally, Yakutsk shows a wide mix of arrivals with generally early arrivals from the Aldan shield and the Verkhoyansk Range and very late arrivals from Shelikhov Bay and Chukotka (Figure 69). Surprisingly the arrivals from the south Yakutia events of 1989 show a mix of residuals for both GT25 and GT10 events.

## 4.4 CONCLUSIONS.

By determining locally calibrated travel time curves and relocating regional earthquakes recorded both teleseismically and by regional networks operating in northeast Russia, epicenters have been greatly improved relative to the ISC Bulletin. GT classifications have been assigned to 134 events, with 26 events meeting or exceeding GT10 criteria. Using the relocated events, travel time residuals for several internationally reported seismic stations located in northeast Russia were plotted. The resulting residual distributions are consistent with high velocities under the Siberian Platform, and low velocities under the Sea of Okhotsk. Taken as a whole, it is clear that ray paths traversing the upper mantle under the Siberia platform are fast, while those originating from, or traversing, the supra-subduction zone or rift regions of the Sea of Okhotsk are very slow. There is an indication of low velocities under the northern Amur district, although this is less distinct. Velocities under Chukotka are less certain. The situation in the Chersky Range is also less clear, although ray paths traveling greater distances may be marginally faster than those to nearby stations. These observations will help in constraining upper mantle velocities and developing better Source Specific Station Corrections (SSSCs) for the region.



consistent among the GT10 events and essentially represent Source Specific Station Corrections (SSSCs). Figure 62. Norilsk (NRI) residuals for all GT analyzed events. Solid colors are from GT10 events while the lighter shade represents GT25 or poorer classification. Note the consistently large negative residuals for paths associatedentirely with the Siberian platform, and lower residuals elsewhere. Residuals are especially

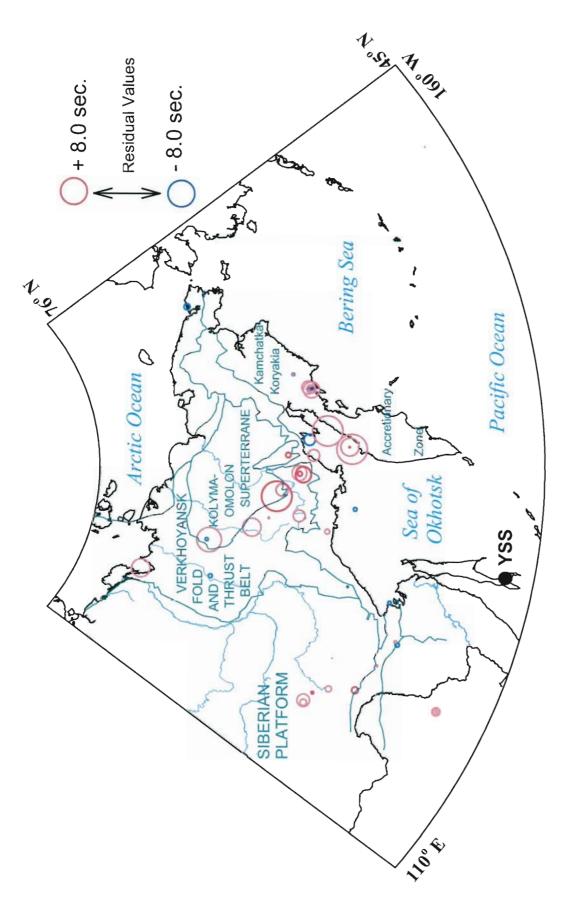


Figure 63. Yuzhno Sakhalinsk (YSS) residuals for all GT analyzed events. Solid colors are from GT10 events while paths crossing the Sea of Okhotsk, while paths originating in the Siberian platform are smaller. Residuals the lighter shade represents GT25 or poorer classification. Note the generally large positive residuals for are essentially representing Source Specific Station Corrections (SSSCs)

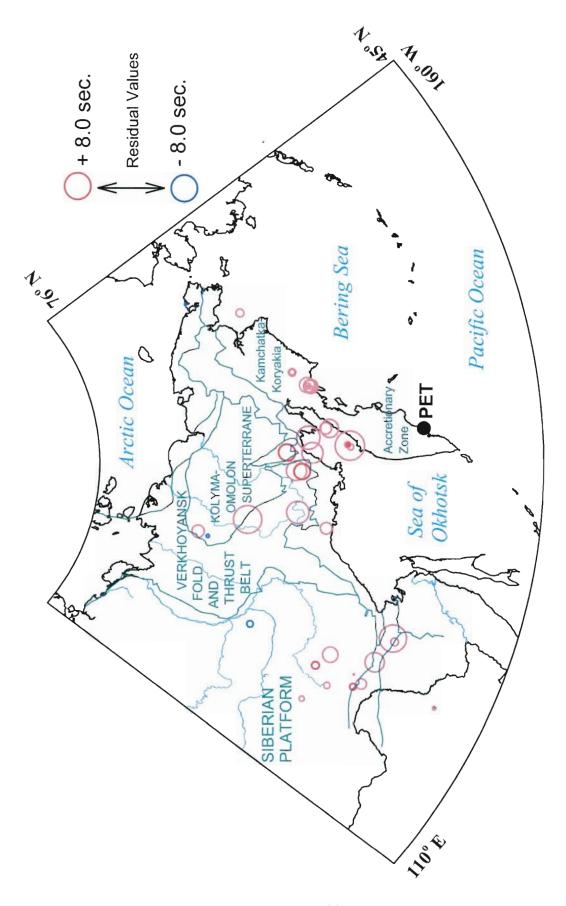


Figure 64. Petropavlovsk (PET) residuals for all GT analyzed events. Solid colors are from GT10 events while the lighter progressively further into the Siberian platform, residuals decrease, possible a result of high platform velocities shade represents GT25 or poorer classifications. Note that most residuals are large and positive. For events countering low velocities under the Sea of Okhotsk. Residuals essentially represent Source Specific Station Corrections (SSSCs).

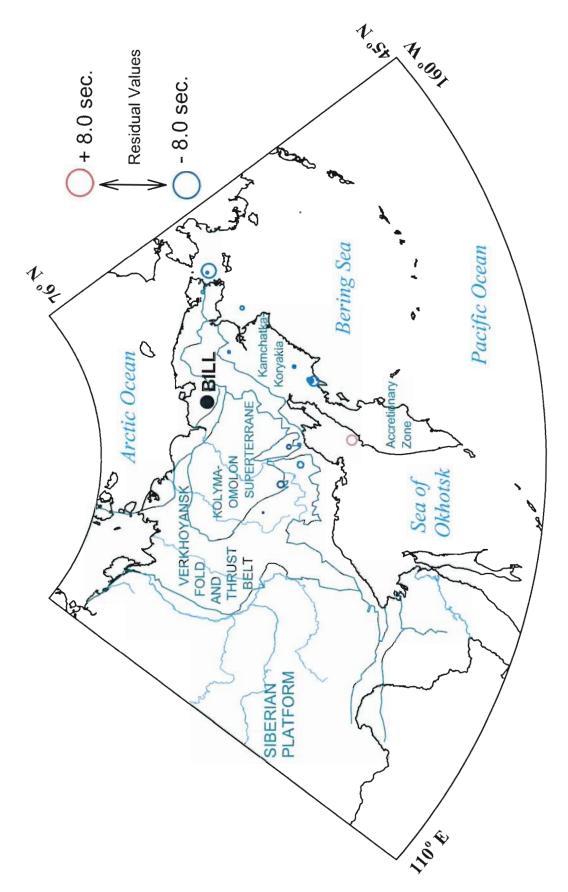


Figure 65. Bilibino (BILL) residuals for all GT analyzed events. Solid colors are from GT10 events while the lighter shade represents GT25 or poorer classifications. Residuals essentially represent Source Specific Station Corrections (SSSCs).

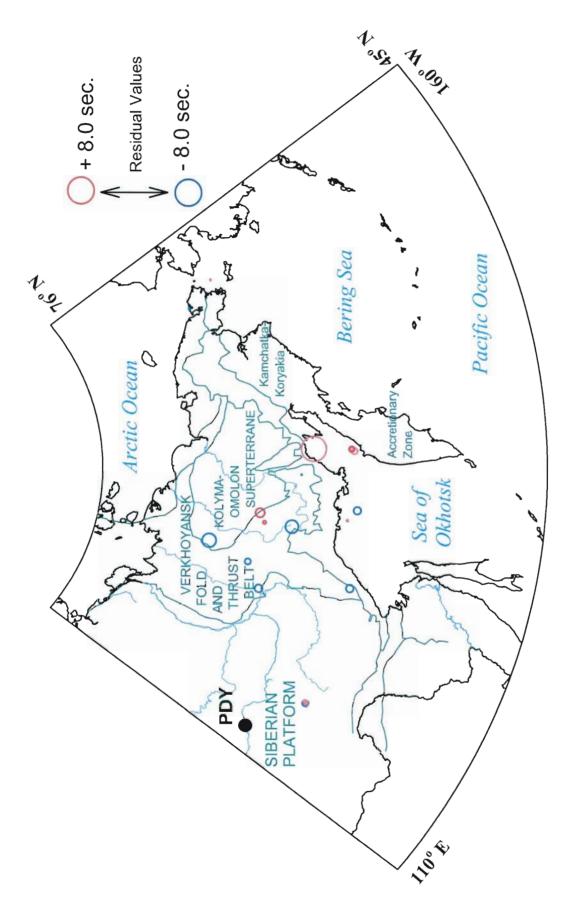
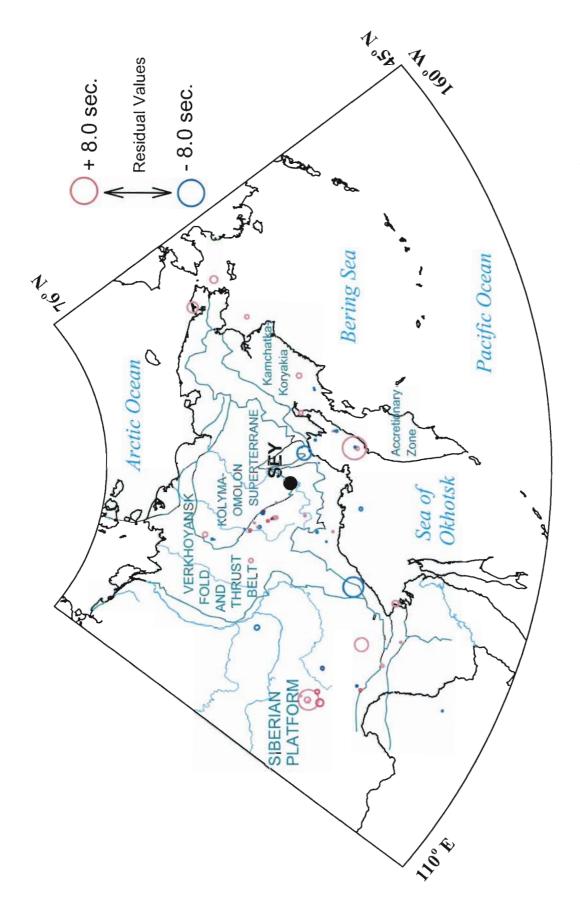


Figure 66. Peleduy (PDY) residuals for all GT analyzed events. Solid colors are from GT10 events while the lighter shade represents GT25 or poorer classifications. Residuals essentially represent Source Specific Station Corrections (SSSCs).



lighter shade represents GT25 or poorer classifications. Residuals essentially represent Source Specific Station Corrections (SSSCs). Figure 67. Seymchan (SEY) residuals for all GT analyzed events. Solid colors are from GT10 events while the

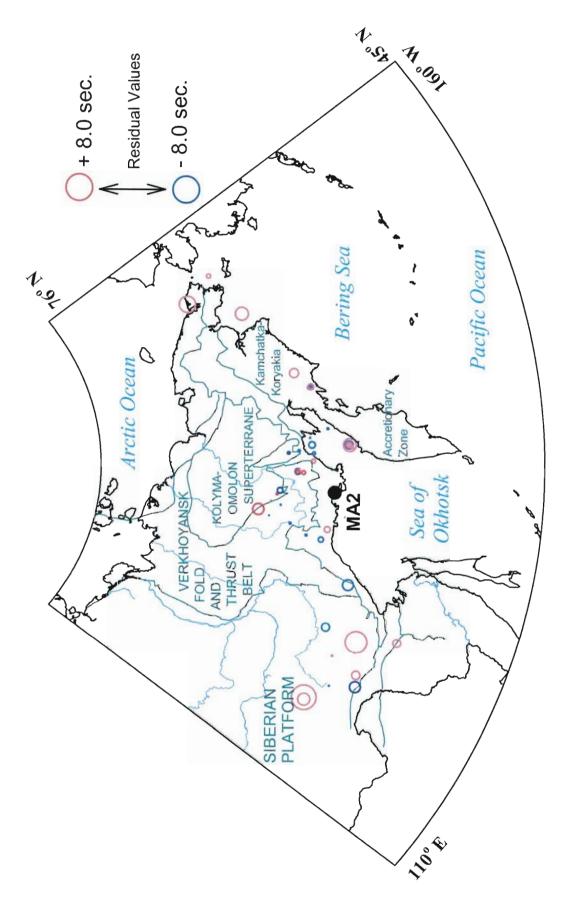
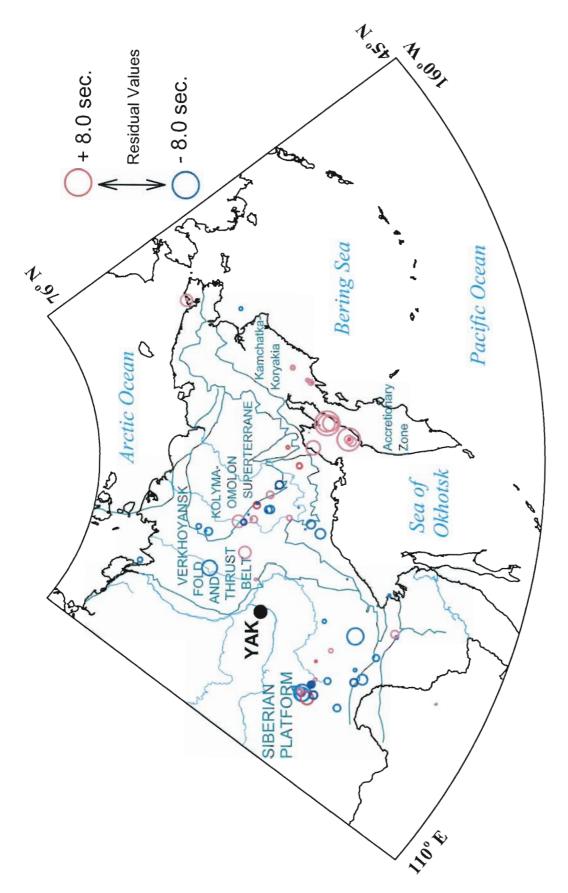


Figure 68. Magadan (MAG, MA2) residuals for all GT analyzed events. Solid colors are from GT10 events while the lighter shade represents GT25 or poorer classifications. Residuals essentially represent Source Specific Station Corrections (SSSCs).



Siberian platform have negative residuals while those from the Kamchatka-Koryakia accretionary zone lighter shade represents GT25 or poorer classifications. Note that most residuals traveling through the show positive residuals. Residuals essentially represent Source Specific Station Corrections (SSSCs). Figure 69. Yakutsk (YAK) residuals for all GT analyzed events. Solid colors are from GT10 events while the

## **SECTION 5**

# DIGITAL STATION DEPLOYMENTS

# 5.1 INTRODUCTION.

In the summer of 1999, a program began to develop seismic discriminants for CTBT monitoring by digitally recording local explosions and earthquakes. To obtain better records for study, existing photopaper analog stations in the Magadan and Yakutsk networks were converted to digital data acquisition. At present, seven digital stations are deployed in northeast Russia, most in the Magadan region (Figure 70; Table 7). Since the initial deployment, a couple of stations have had to be moved due to logistical reasons. A few stations were also deployed in temporary locations for the purpose of recording known explosions or earthquake aftershocks. A description of each station location is given in Appendix D. All stations are 3 component, with three stations recording broadband instruments (Seymchan - STS-1, Okhotsk - Guralp CMG-40T, and Omsukchan - CMG-40T). The remaining stations use short-period Russian seismometers (SKM-3, SM3 or SM3-KV). All stations recording SM3 or SM3-KV seismometers are equipped with anti-aliasing filters to cut frequencies above 10 hz. All seismometer outputs are digitized at 30 samples per second with a 24bit A/D converter and logged on a PC. Digital data acquired as a part of this study is supplemented with records from the IRIS stations at Magadan (MA2) and Bilibino (BILL), and other regional analog photopaper stations. Using waveforms of both tectonic events and mine blasts, explosion discrimination using amplitude ratios of various Lg frequencies has been addressed. Preliminary results indicate that the ratio Lg(4-8Hz)/Lg(0.75-1.5Hz) using peak amplitudes may not be sufficient to discriminate mining explosions from earthquakes in northeast Russia.

## 5.2 RESEARCH PERFORMED.

# 5.2.1 Earthquakes.

Approximately 250 local and near regional earthquakes have been located by the network from August, 1999 through August, 2001 (Figure 70). Arrival times for all local and near regional earthquakes are routinely picked from the digital waveform data and a complete catalog is maintained. Catalog development includes calculating hypocenters for all locatable earthquakes, as well as identification of signals from anthropogenic sources. The stations are located close to areas of both tectonic seismicity and active mining, thus record signals from both. Figure 71 depicts digital records from one local tectonic event. Many smaller events are also recorded, but are not locatable. The magnitude cutoff for locatable events for the central portion of the study is about  $M_L = 2.5$ , and about  $M_L = 3.0$  outside the aperture of the deployed stations. Some variation exists depending on specific location relative to operating stations. The largest event from eastern Russia recorded by the network was the  $M_w$  6.8 Uglegorsk earthquake from the southern part of Sakhalin Island, which occurred at 21 hours on August 4, 2000 (Figure 72). This event occurred approximately 1000 km south of the network, at the edge of the scope of this project. The stations have also recorded several hundred other regional and teleseismic earthquakes.

In a qualitative sense, both Pg and Lg phases appear to propagate well throughout the region enclosed within the network. Clean arrivals from both phases are often well recorded beyond 700 km, the aperture of the deployed digital network. With a few exceptions,  $P_n$  arrivals are generally much lower in amplitude, or not visible at these near regional distances. Almost no  $S_n$  arrivals are

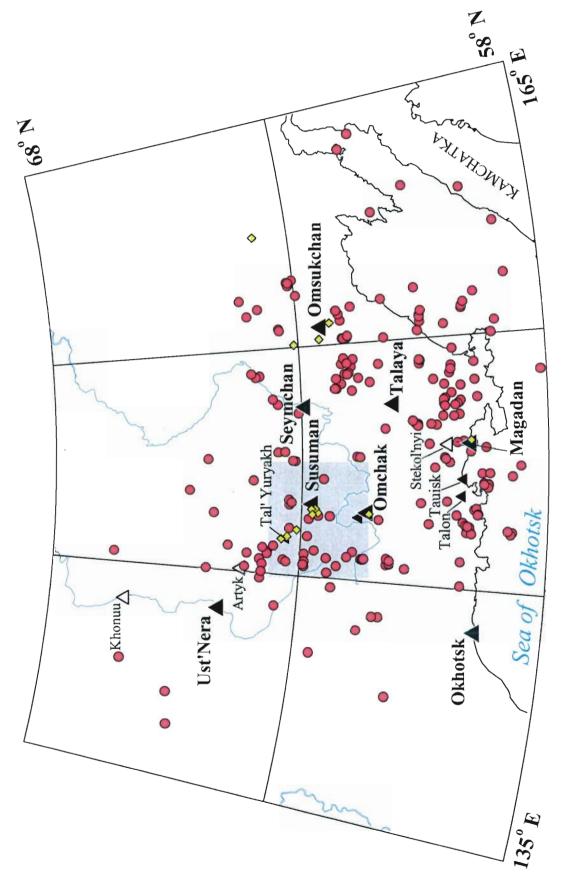


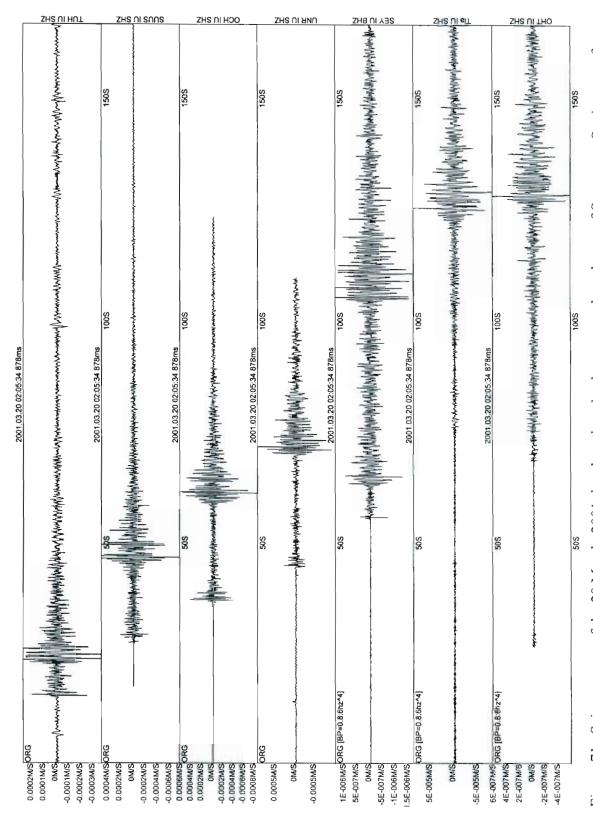
Figure 70. Digital seismic network deployed by MSU in eastern Russia (permanent stations - large triangles, temporary stations small triangles, and supporting analog stations - open triangles). Also shown are earthquakes located from August 1999 through August 2001 (red circles) and explosion sources (diamonds). Shaded area shown in Figure 75.

Table 7. Digital seismic stations deployed or used in this study.

Station Name	e Code	Lat.	Long	Instrument	Notes
Anadyr	ANYS	64.734°N	177.496°E	CMG-40T	3/00 - 4/01*
Bilibino	BILL	68.039°N	166.271°E	STS-1	IRIS
Magadan	MA2	59.5756°N	150.768°E	STS-1	IRIS
Matrosova	MATR	61.643°N	147.820°E	SM3-kv	7/99 - 7/99
Nel'koba	NKB	61.338°N	148.813°E	SM3-kv	7/99 - 9/99
Okhotsk	OHT	59.359°N	143.331°E	CMG-40T	7/00 - OPEN
Omchak	OCH	61.67°N	147.87°E	SM3-kv	9/99 - OPEN
Omsukchan	OMS	62.52°N	155.77°E	CMG-40T	7/01 - OPEN
Seymchan	SEY	62.933°N	152.368°E	STS-1**	9/99 - OPEN
Susuman	SUUS	62.780°N	148.163°E	SM3-kv	7/99 - OPEN
Stokolviya	STV	61.8475°N	147.6598°E	SM3-kv	7/99 - 7/99
Talaya	TL-S	61.129°N	152.392°E	SM3-kv	7/99 - OPEN
Talon	TON	59.757°N	148.661°E	SM3-kv	1/01 - 2/01
Tal-Yuryakh	TUR	63.307°N	146.634°E	SM3-kv	3/01 - 6/01
Tauisk	TAUS	59.729°N	149.335°E	SM3-kv	1/01 - 4/01
Ust'Nera	UNR	64.565°N	143.242°E	SKM	7/99 - OPEN

<sup>\*</sup>Open and close dates of the digital station. Station did not acquire data for most of this time.

\*\*Seismometers and support equipment left from non-operational Seymchan GEOSCOPE station



bottom, Tal-Yuryakh, Susuman, Omchak, Ust'Nera, Seymchan, Talaya, and Okhotsk. All stations depict the vertical Seismograms of the 20 March, 2001, local earthquake that occurred northwest of Susuman. Stations are, from top to component. Figure 71.

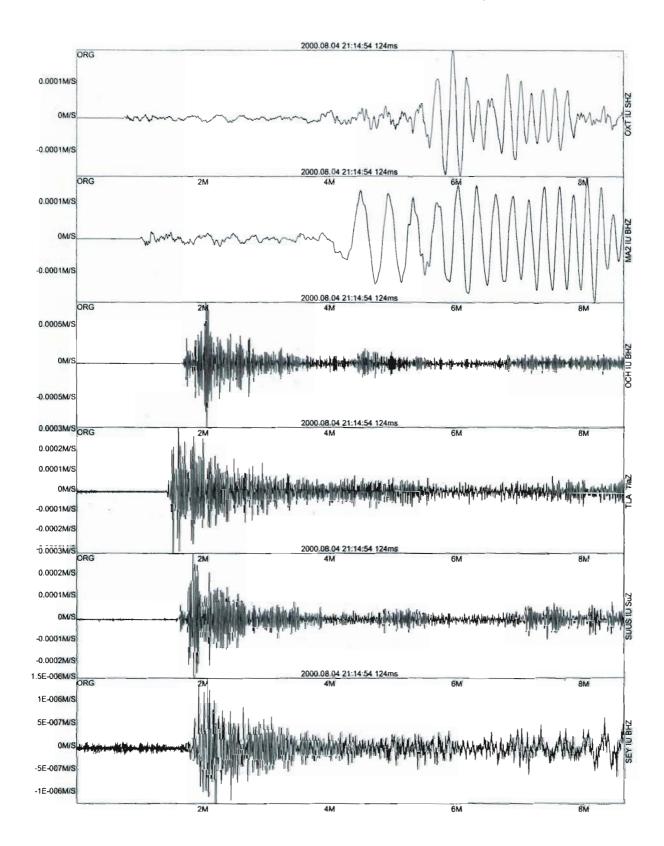


Figure 72. Seismograms of the August 4, 2000 Mw 6.8 Uglegorsk earthquake. Stations are, from top to bottom, Okhotsk, Magadan, Omchak (time out), Talaya, Susuman, and Seymchan. Instrument types are given in Table 7

visible in the digital records.

Initial analysis of waveforms indicates that the Lg arrival sometimes appears to have two distinct arrivals, Lg<sub>1</sub> and Lg<sub>2</sub>. Lg<sub>1</sub> being a slightly faster phase penetrating deeper and having less high frequency attenuation, and Lg<sub>2</sub> being a slower, shallower phase with attenuation of high frequencies (Hartse, pers. comm.). This characteristic is most prominent at Susuman (Figure 73).

For some events at local distances, the mantle reflection phases PmP and SmP appear to be prominent, occasionally greatly exceeding the amplitudes of Pg and Sg (Figure 74). It should be noted that the moveout of these suspected mantle reflection phase arrival times has not been investigated to confirm positive identification.

An effort was undertaken to acquire improved ground truth data for earthquakes. In January, 2001, two moderate earthquakes occurred approximately 100 km west of Magadan. Two temporary stations (Talon and Tauisk) were deployed to investigate the aftershock sequences of both events. About 15 events were locatable, although several more were recorded.

# 5.2.2 Explosion ground truth and discrimination.

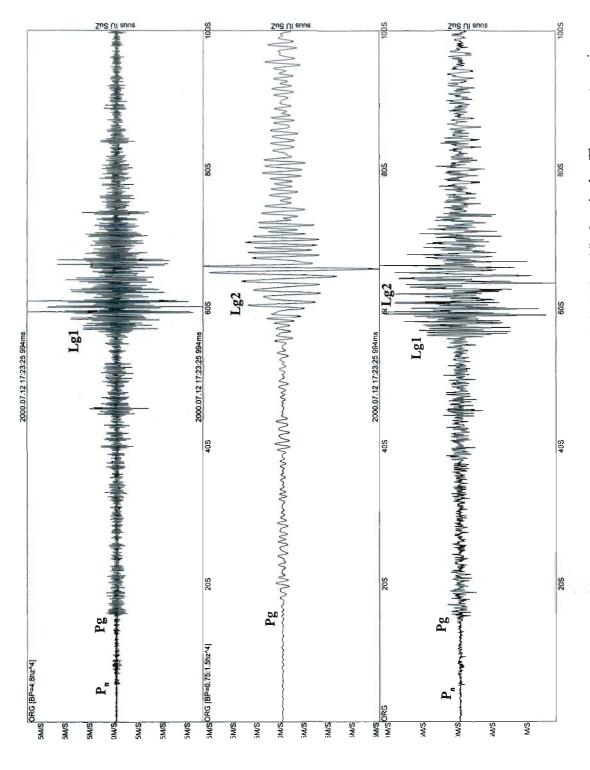
Identification of local and regional mine blasts has been an integral part of the data analysis of the digital network data. Three stations are located in active mining regions, which lends itself to acquisition of ground truth data directly useful for calibration of several IMS Alternate Stations in the region. To obtain ground truth information on explosions, mine locations were either visited or temporary stations deployed to obtain information on blast size, time and location. Blasting schedules for some other mines were also obtained, although mine sites have not yet been field checked for accurate locations. Specific mine locations and mine types in the Susuman region are plotted on Figure 75. Full data on explosions recorded between late 1999 and late 2001 can be found in the annual data compilation reports issued separately as a part of this study (Mackey and Fujita, 2001a,b,c).

Many mine blasts were recorded from many locations throughout the study region. The most active region is around Susuman where blasting originates in both gold and coal mines. Explosions in the placer gold deposits, used to break up frozen ground, have explosive yields ranging up to or exceeding 40 T. Figure 76 shows one 40 T mine blast which occurred at the Neksikan placer gold deposit, about 25 km from Susuman in April 2001. This seismogram also shows a clear acoustic arrival, which is visible on the Susuman seismograms for about 25 percent of large blasts from Neksikan. There were approximately 15 similar blasts detonated in April and May, 2001, to form the currently active placer mining pit in Neksikan (Figure 77).

Preliminary analysis of mine blasts show distinctive Rg arrivals at short distances of less than about 50 km (Figure 76). Lg arrivals from several of the 40 T Neksikan blasts were recorded at Magadan (MA2) and the farthest station, Okhotsk, over 400 km distant (Figure 78). Lg arrivals recorded at Magadan and Okhotsk are characterized by low frequency arrivals (0.7-1.5 Hz) with higher frequencies (4.0-8.0 Hz) being absent or completely obscured by background noise.

The largest explosion in the field area during the period of study is believed to have occurred 10 June, 2000, at one of two pits at the Kadakchan coal mine northwest of Susuman. This explosion was identified by location, time of day, similarity of waveforms to known blasts from Kadakchan, as well as a large acoustic arrival recorded at Omchak. The actual tonnage of the event is unknown, although it had a K-class of 9.5, which corresponds to a magnitude of about 3.3. This explosion was well recorded at the farthest stations in the network, including Magadan (MA2; Figure 79). Appendix E expands the discussion on this event with additional seismograms and maps.

A few temporary stations were deployed at or within a few kilometers of mine sites to obtain ground truth locations and origin times of explosions. The longest deployment was at the Tal-



bandpass filtered 4-8 Hz, and shows an early high frequency maximum amplitude. The middle trace is bandpass filtered 0.75-1.5 Hz, and shows that the low frequency maximum amplitude is later. The lower trace shows the full waveform Figure 73. Susuman seismograms of a tectonic earthquake showing characteristic Lg1 and Lg2 arrivals. The upper trace is where the low frequency arrival is visible about seven seconds after the initial onset of Lg.

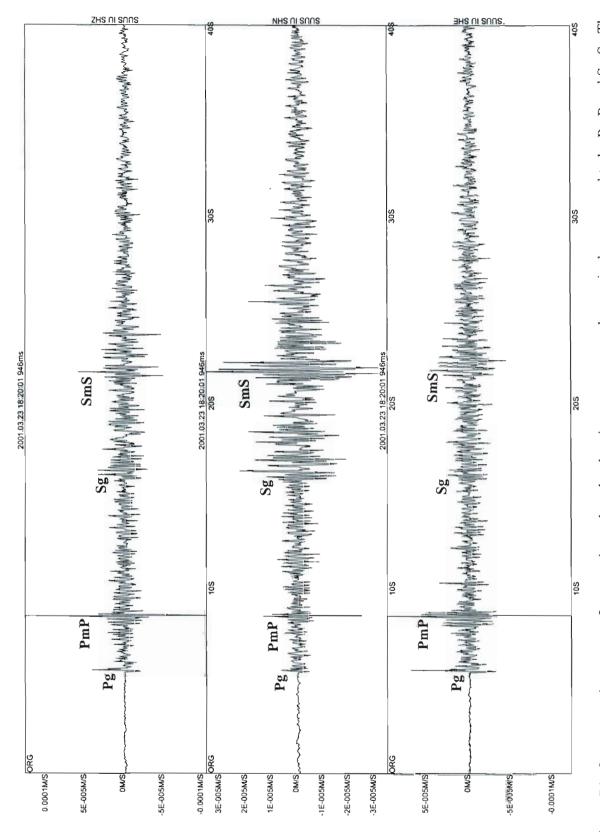


Figure 74. Susuman seismogram of a tectonic earthquake showing strong secondary arrivals, presumed to be PmP and SmS. The amplitudes of the secondary arrivals exceeds the first arriving Pg and Sg phases.

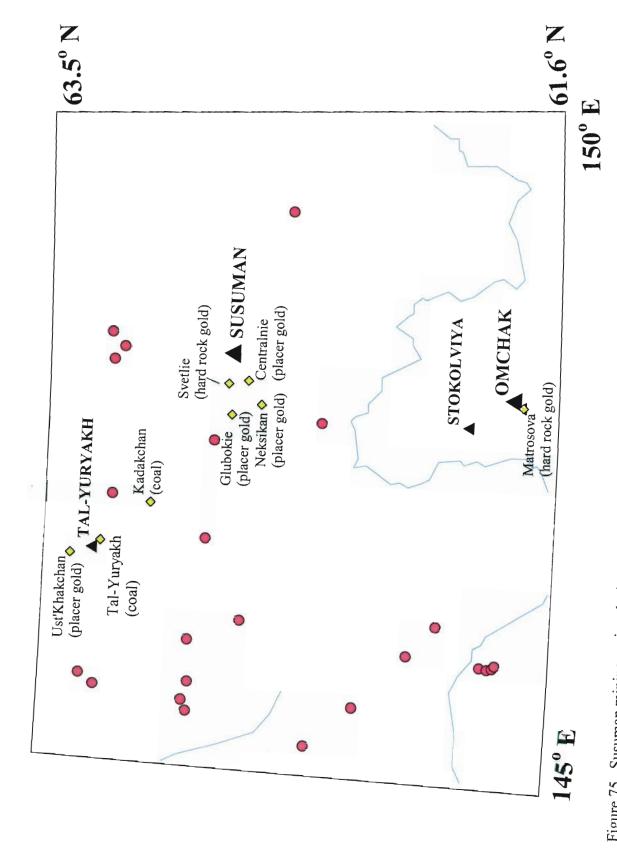


Figure 75. Susuman mining region depicting located tectonic events and specific known active mines. Symbols as in Figure 70.

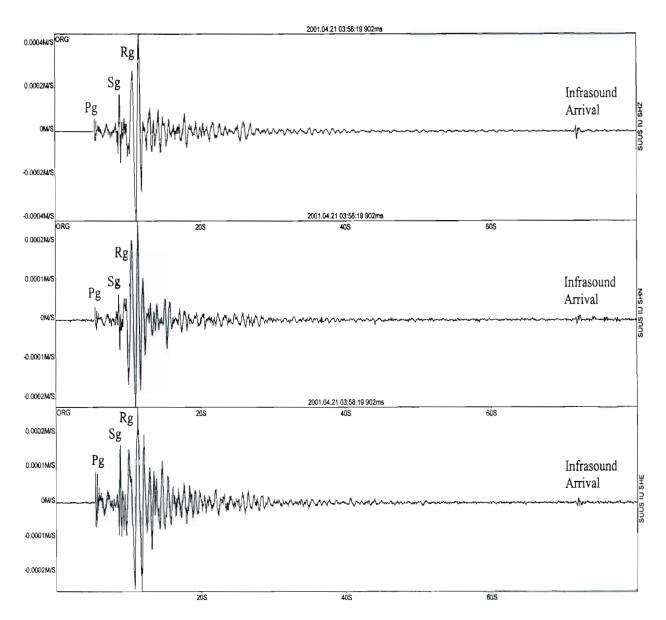


Figure 76. 3-component seismogram (top - Z, middle - N-S, bottom - E-W) of a 40,000 kg explosion recorded at Susuman. This explosion occurred April 21, 2001, in a placer gold mine near the town of Neksikan, about 21 km southwest of Susuman. Note the clearly developed Pg, Sg, and Rg arrivals, as well as a clear acoustic arrival recorded on all seismometer components about 70 seconds after Pg.

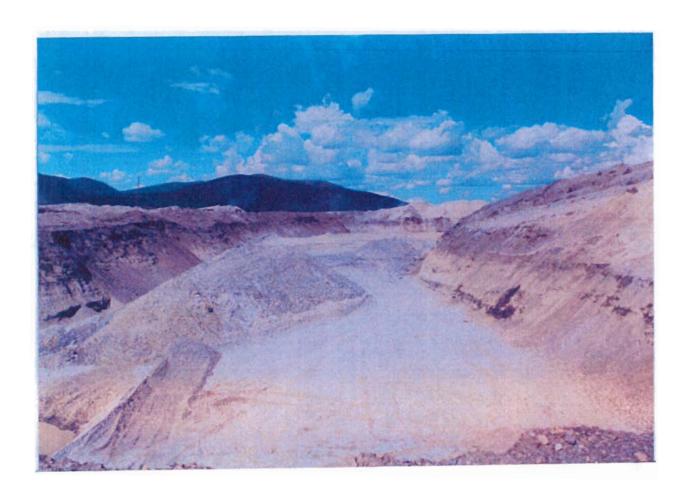


Figure 77. Placer gold mine at Neksikan. Approximately fifteen blasts in the 30,000 to 40,000 kg range were detonated in April, 2001, to prepare frozen sediments for the excavation of the pit seen here. Some of these blasts were recorded throughout the deployed digital network (see Figures 76 and 78).

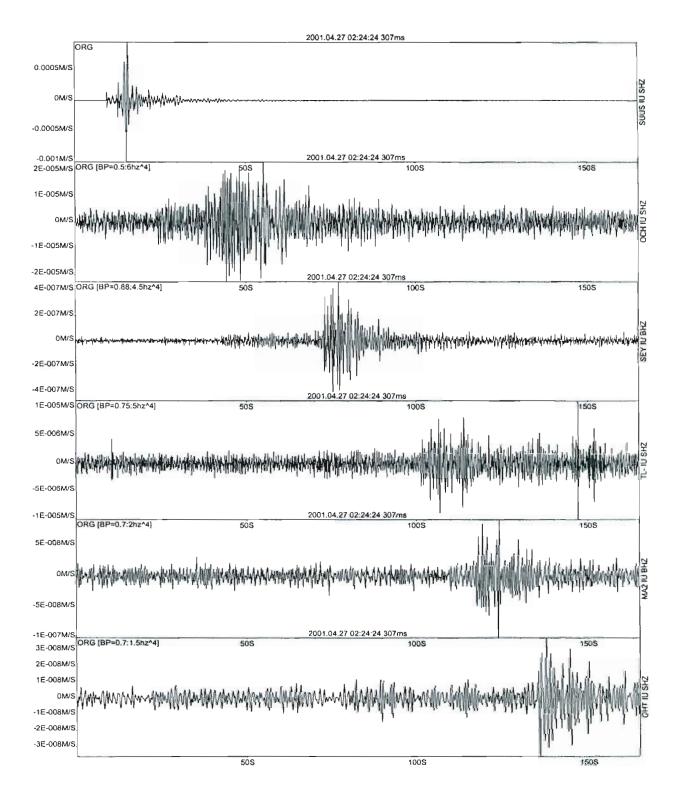
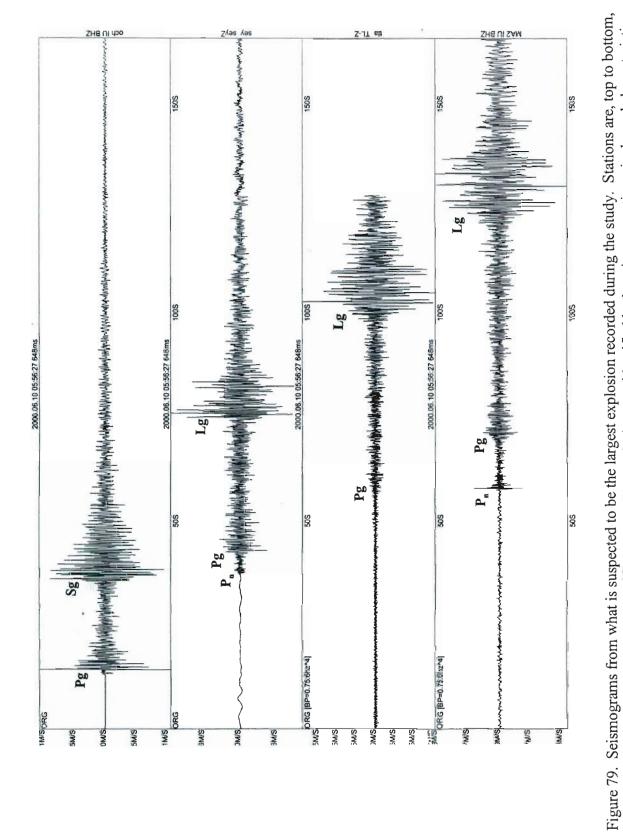


Figure 78. 40,000 kg blast of April 27, 2001 as recorded by several of our stations (top to bottom: Susuman (22 km), Omchak (111km), Seymchan (236 km), Talaya (297 km), Magadan (MA2; 381 km), and Okhotsk (441 km). Note the well developed low frequency Lg arrival at Okhotsk. All records except Susuman are bandpass filtered.



waveform to have occurred at the Kadakchan coal mine (see Figure 70). This explosion is discussed in detail in Appendix Omchak, Seymchan, Talaya, and Magadan. The explosion was identified by location, acoustic arrivals, and characteristic

Yuryakh coal mine, northwest of Susuman (Figure 75). This temporary station operated from March through May, 2001, and recorded several explosions from coal mines at both Tal-Yuryakh and Kadakchan. The largest explosion recorded from this deployment was about 24 T. A sample seismogram from Tal-Yuryakh is shown in Figure 80. Temporary stations were also deployed at the Matrosova mine and near the Magadan granite quarry.

Careful examination of explosions originating at several mines and recorded at several stations indicates that the correlation of waveforms is very high when viewing the waveforms from the same explosion source/station pair. In many cases, waveforms can be matched swing for swing, including relative amplitudes of individual peaks. Work is currently underway to construct a catalog of characteristic waveforms for all explosion source/station pairs.

As a large number of earthquakes and explosions from throughout the study region have been recorded, it is possible to attempt to discriminate between explosions and earthquakes using various Lg amplitude ratios. Preliminary results indicate that the ratio Lg(4-8Hz)/Lg(0.75-1.5Hz) using peak amplitudes is not sufficient to discriminate mining explosions from earthquakes at local and near regional distances (<500km) in northeast Russia (Figure 81). In general, the overall relationship (slope and intercept) between earthquake Lg ratios vs. distance is nearly identical to that of HIA where the ratio has been used to effectively discriminate between nuclear explosions and tectonic earthquakes (Hartse, 2000). It should be noted that this represents only the most simple approach to an Lg discriminant for northeast Russia. Additional more comprehensive research is underway to look at the RMS amplitudes of the full Lg wave packet (Hartse, 2000) as opposed to peak amplitudes. Additional phase amplitude ratios (Pg/Lg, P<sub>n</sub>/Lg, etc.) are also planned to be investigated.

## 5.3 CONCLUSION.

In order to obtain ground truth seismic events, a small network of digital seismic stations was deployed in the Magadan district, including at or near active mines. The network has recorded and located a large number of both earthquakes and explosions. Several temporary stations were also deployed to record mine blasts of aftershocks from larger earthquakes. Preliminary analysis of Lg amplitude ratios indicate that the ratio Lg(4-8Hz)/Lg(0.75-1.5Hz) using peak amplitudes is not sufficient to discriminate mining explosions from earthquakes at local and near regional distances (<500km) in northeast Russia.

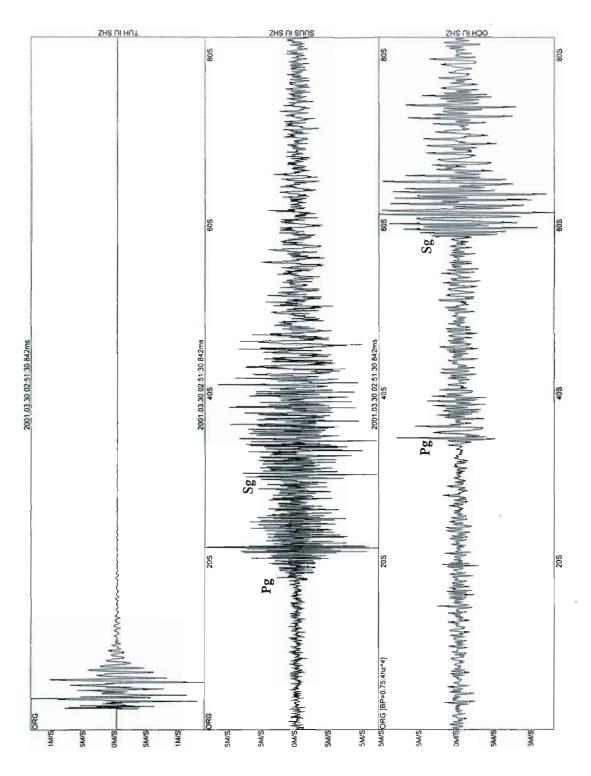


Figure 80. Explosion from the Tal-Yuryakh coal mine recorded by the temporary seismic station deployed at the mine, as well as Susuman (center) and Omchak (bottom). The yield of this blast was 4590 kg. The complex waveform recorded at Susuman is characteristic of Tal-Yuryakh explosions and can be used as a discriminant for this mine. The complexity of the waveform is much reduced in the filtered seismogram from Omchak.

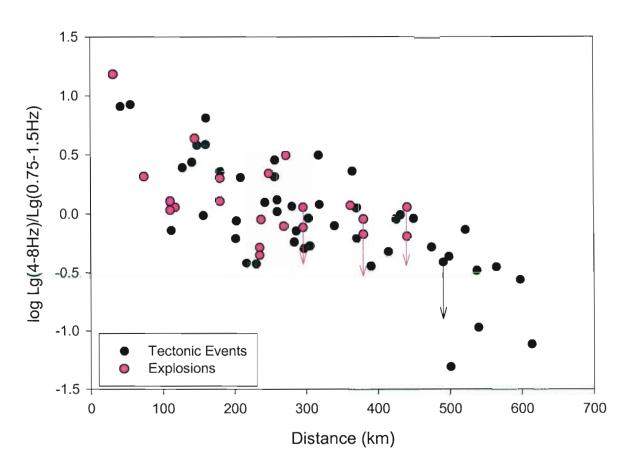


Figure 81. Composite plot of Lg spectral ratio vs. distance for tectonic events and explosions using all stations except Omsukchan. Lg amplitudes used were maximum peak to peak. There is no apparent difference in the plotting of earthquakes and explosion, thus discrimination is difficult here using this method at these distances. Points with arrows represent maximum possible ratio for arrivals with a clear low frequency Lg arrival (see Okhotsk, Figure 78) but no apparent high frequency arrival. In these cases, the high frequency Lg is substituted with the high frequency background.

#### SECTION 6

## **CONCLUSIONS**

A catalog covering historic seismicity and associated phases and arrival times compiled for northeast Russia has proven useful for seismic characterization of the region. The resultant catalog was found to be contaminated by industrial explosions, particularly in the Amur and central Magadan districts where dramatic differences between daytime and nighttime seismicity were found. In the process of developing the seismicity catalog, operational procedures, network histories and seismic station parameters were researched.

Phase arrivals were used to develop a crustal velocity model by obtaining best fit travel-time curves over 3 x 5 degree regions. The velocities obtained are generally in agreement with inferred tectonic regimes. These travel-time curves were then used to relocate larger regional events, improving relative locations as aftershock clusters are tighter and earthquakes better align along faults..

Using the developed crustal velocities, 134 seismic events reported in the International Seismological Centre (ISC) were relocated. All events were assigned GT levels, with 26 events classified as GT10. From relocated events, consistent patterns of residuals, essentially representing path corrections, show upper mantle velocities are elevated under the Siberian platform and slower below the Sea of Okhotsk...

To further improve calibration capabilities in northeast Russia, a small network of digital seismic stations was deployed in the in the Magadan region, recording both earthquakes and ground truth mine blasts. Preliminary analysis of Lg amplitude ratios indicate that the ratio Lg(4-8Hz)/Lg(0.75-1.5Hz) using peak amplitudes is not sufficient to discriminate mining explosions from earthquakes at local and near regional distances (<500km) in northeast Russia.

## **SECTION 7**

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APPENDIX A

Alphabetized list of northeastern Russia seismic stations.

Station Name	_	Russian Code	Lat. (°N)	Long. (°E)	Elev. (m)	Date Open	Date Close	Qual.
Aku			56.46	120.91	700	/68	/68	1
Aldan	ALDR		58.67	125.40		<b>-/99</b>	OPEN	1
Alla		AJI		110.82	550	10/63	5/70	1
Alygdzher	ALY			98.218	920	1/66	1/67	1
Amedichi	ACHS	АМП	57.03	122.85	930	05/89	08/89	1
Amguema	AMG	' '	67.05	178.88W	150	11/65	4/66	2
Ammonl'naya	MMS		64.55	143.18	540	/61	/62	1
Anadyr	ANYS	$AH\Pi$		177.496	55	4/89	7/93	1
1 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2						9/96	OPEN	
Anadyr-1	ANSS	АНП	64.77	177.57	40	11/80	1/89	3
Angarakan	AGK			113.67	1430	11/76	8/81	1
Anyuisk		AHC	68.34	161.56	10	6/64	3/65	1
Apacha	APC		52.925	157.131		2/90	OPEN	1
Apakhonchich		АПХ	56.00	160.84	700	/64	-/80	1
Arshan		APIII	51.908	102.433	840	<b>-/58</b>		1
Artyk	AYKS	AP	64.18	145.13	700	06/71	/71	1
,			64.183	1145.1347	700	/88	OPEN	G
Avacha - old	AVH	АВЧ	53.07	158.5		<b>-</b> /63	10/76	-
Avacha - new	AVH	АВЧ	53.265	158.738	900	7.76	OPEN	1
Babushkin	BAU		51.717	105.867	470	6/66	9/66	1
Baikal'sk	BKK		51.522	104.133	460	1/64	2/66	1
Balygychan		БЛГ	63.91	154.09	139	7/63	6/64	1
Barluk	<b>BRUS</b>		54.533	101.717	525	11.60		1
Batagai	<b>BTGS</b>	БТГ	67.653	134.630	127	/75	<b>OPEN</b>	1
Bazovskii			56.53	123.42	1080	/70	/70	1
Berezovaya	BER	БРЗ	52.27	158.433		<b>-</b> /81	<b>-</b> /94	1
Bering (Nikol'skoe)	BKI	БРН БРГ	55.195	165.99 400	<b>-</b> /62	OPEN	-	
Bilibino GSN	BILL	BILL	68 065	166.452	299	8/95	OPEN	1
Bilibino GSN		БЛБ		166.449	283	8/81	4/92	1
Pillolilo	BILS	БЛН	08.039	100.449	203	0/01		1
Bilibino-1	BL1	БЛБ	68.04	166.44	260	8/64	1/65	1
Bodaibo	BOD	БДБ	57.807	114.03	245	11/60		1
Bodon	BDN	БОД БДН	53.713	110.1	540	11/69	/83	1
Bogachevka	BGC	r 1	54.850	160.900		<b>-</b> /64	<b>-</b> /65	1
Bomnak	BMKS	БМН		128.847	325	3/74		1
Bykov				142.567	40	6/68	10/68	-

Chagda	CGD	ЧГД	58.75	130.60	185	/68		1
Chara	CRS	ЧР	56.9	118.267	710	11/61		1
Cherskii	CES	ЧРС	68.75	161.33	10	/79	/88	1
Chil'chi	020	11 0	56.06	122.33	500	/70	/70	1
Chita	CIT	ЧТ		113.55	790	6/70	770	1
	CH	71					0/0//75	
Chochurdakh	~~~~~	*****	72.83	116.25			9/26/75	
Chokchoi	CKHS		57.65	121.72	240	/89	/89	1
Chul'man-1		ЧЛМ	56.85	124.90	650	/62	/86	1
Chul'man-2	CLNS	ЧЛМ	56.84	124.90	760	/86		1
Davsha	DAS		54.538	109.503	460	1/64	6/65	1
Debin	DBI	ДБН	62.339	150.751	332	/74	6/86	G
		, ,	62.343	3150.7594	332	6/86	/92	G
Dimnoe				142.400		3/74	4/74	_
Dovochan	DOV			117.533	1094	7/62	9/63	1
Dunai	DUYS	пц	73.92	124.49	5	11/89		1
	D013	7411	56.60	121.13	460	/67	/67	
Dyrynmakit	T-CX 10	OTD						1
Egvekinot	EGVS			179.127W	18	<b>-/90</b>	/94	1
Ekimchan	EKI	EKM		132.945	485	11/79		1
Emegachi	<b>EMG</b>			118.158	960	9/62	4/63	1
Esso	ESO	ЭСС	55.925	158.700	490	-/65	OPEN	1
Esutoru (Uglegorsk)	ESU		49.083	142.033	100	12/39	/45	-
Evensk	<b>EVES</b>	ЭВН	61.92	159.23	22	/80	7/93	1
Firsovo			47.65	142.567	20	8/79	11/79	_
Ganali	GNL	ГНЛ	53.942	157.620	1200	1/88	<b>OPEN</b>	1
Garmanda	GRM	ГРМ	62.18	159.08	140	12/66	5/67	1
Gorely	GRL	ГРЛ		158.080	1250	7/80	OPEN	1
Gornotaezhnoe	GRZ	ГРТ	43.70	132.15	220	7/90		3
Gomotaczinioc	GRT		73.70	152.15	220	1170		5
Compayadaaa		ГРВ	43.70	134.733	270	7/88		1
Gornovodnoe	GRD	IFD	43.70	134.733	270	//00		1
0 11	GRV		16.565	1.41.05	50	10/71	(170	
Gornozavodsk	~~~	TTTT		141.85	50	12/71	6/72	-
Gornyi	GNY	ГРН		136.455	500	12/78		2
Gusinoozersk	GOO	ГСН		106,517	600	11/71	2/72	1
Ilimei	ILR	ИЛР	67.26	167.96	350	10/64	10/65	1
Imangra-1			56.75	121.24	395	07/67	08/67	1
Imangra-2	<b>IMNS</b>	ИМГ	56.62	120.71	540	/75	/79	1
Institut	INS		53.066	158.605	175	11/81	<b>OPEN</b>	1
Vulkanologii								
Irkana	IKN		55 867	111.253	480	1/64	3/66	1
Irkutsk	IRK	ИРК		104.31	467	12/01		1
Iul'tin	ILT	ИЛТ	67.87	178.74W	235	3/66	7/93	1
Kabaktan		КБК	56.68	122.42	1010	05/89		1
Kabaktan	KDKS		30.08	122.42	1010	03/69	00/09	1
Walaansi.	TZAD	KBT	52.05	106 659	165	1/51		1
Kabansk	KAB	КБ	52.05	106.658	465	1/51	00/10/2	1
Kalgannakh	****	T73 /TY	71.83	114.33			08/10/7	
Kamenistyi	UL2S	KMH	65.41	144.83	670	/88	/88	1

Kamenistaya	KMN		55.76	160.240	1100	10/90	OPEN	1
Kamenskoe	KAM		62,456	166.210		<b>-/94</b>	OPEN	1
Karam	KRMS	1		107.583	600	1/66	3/66	1
Karymski - old	KII	KPM?		159.480	790	7/74	<del>-</del> /86	1
Karymski - olu	KII		54.050	133.400	790	///4	-/80	1
Y7 1.1	T/DX/	KAP?	54006	150 440	000	0.100	0222	
Karymski - new	KRY			159.449	900	9/89	OPEN	1
Khaim	KAIS	XM		108.085	480	10/69	5/70	1
Khandyga	KHG	ХНД	62.65	135.56	125	/69	/94	1
Khani	KHNS	XH	57.04	121.01	390	/67	/67	1
						/75?	/76?	
Khapcheranga	KPC	ХПЧ	49.707	112.392	950	12/68		1
Khatystyr	KHY	XTC	55.71	121.57	475	/68	/68	1
Linutyouji	1111	1110	33.71	121.57	175	/75	/82?	1
Khingansk	KNN	ХНГ	40 122	131.192	520	7/80	4/84	1
Killigalisk		AIII	49.122	131.192	320	//80	4/04	I
77' '1 11	KNG		70.067	120.067		c / = 0	0.450	
Kigilyakh		****		139.867		6/73	9/73	-
Kirovskii	KIRS	KPC		126.983	440	4/74		1
Klyuchi	KLY	КЛЧ	56.313	160.852	80	<b>-/48</b>		1
						2/89	OPEN	
Kobdi	AY1S		64.20	145.51	800	06/71	/71	1
Kolokol'chik						-/67	<b>-</b> /67	
Korito	KRT		55.966	160.222	1000	10/97	OPEN	1
Korsakov						/51	1/52	_
Koryak	KRK	KPK	53 292	158.636	1050	7.75	OPEN	1
Kotelnyi	KIKIK	KLIK		137.60	1030	8/72	9/72	1
•					1.0			-
Kotikovo	*****	TENTE		144.25	10	7/69	9/70	-
Kovokta	KVO	KBK		113.05	1180	/81		1
Kozelskaya	KZL	КЗЛ		158.894	950	/76	-/84	1
Kozyrevsk	KOZ	КЗР	56.05	159.87	40	<b>-</b> /62	<del>/89</del>	1
Kozyr	KZY		56.070	159.900	450	11/89	OPEN	1
Krestovaya	KTB		52.665	106.395	560	7/71	9/71	1
Krestovskii	KRS		56.214	160.558	1200	7/87	<b>OPEN</b>	1
Kronoki	KRI	KPH	54.596	161.134	5	8/66	<b>OPEN</b>	1
Krutoberegovo	KBG	КБГ		162.705	10	-/68	OPEN	
	2200	КБ				,		
		КРБ						
Kul'dur	KLDS		40.205	131.642	425			1
		, ,				1 /00	10/02	
Kulu	KU-S	КЛ	01.889	147.431	655	1/80	10/92	G
		КЛУ						
Kumora	KMO	KMP		111.208	475	9/66		1
Kurbulik	KBK		53.708	109.038	460	1/64	9/65	1
Kurul'ta			56.90	121.11	495	/68	/68	1
Kyakhta	KYA	KXT	50.35	106.45	760	3.52	5/70	1
Kyubyme			63.38	140.95	950	/74	/74	2
Kyusyur	KYUS	KCP	70.68	127.37	20	/85	08/89	1
Lamutskoe	LMT	ЛМТ	65.54	168.85	178	4/65	10/65	1
Lapri		VAA-14	55.69	124.91	640	/72	/73	1
Lupii			55.07	147.71	0-10	112	113	1

Lazarev	LZR	ЛЗР	52.2	141.493	120	12/80		2
Lesogorsk			49.442		40	7/69	10/69	-
Lopatino			46.6	141.825	40	4/69	10/69	-
Lurbun	LRB		56.63	117.883	780	8/62	9/63	1
						5/67	9.68	1
Magadan	MAG	МГД	59.560	150.805	78	1/52	1/92	1
Magadan-GSN	MA2	MA2	59.575	150.768	339	9/93	<b>OPEN</b>	1
Maiskii	MKI	МАЙ	68.97	173.71	261	8/82	6/94	1
Mal. Ipelka	MIP			156.758	370	8/97	OPEN	1
Maritui	MRU			104.217	520	-/08	<del>-/18</del>	1
Markovo	MKV	MPK	64.68	170.41	25	10/86	4/92	1
1714111010	MKN	1,111	0 1.00	170.41	23	10/00	7/ / 2	
Matrosova	MATE	)	61 643	147.82		7/99 -	7/00	G
Mednyi	MED	МДН		167.556		-/73	-/75	1
Milkovo		МЛК			155			
MINKOVO	MLK	MIJIK	54.70	158.63	155	-/62	-/63	1
M IZI	3.4777.10		CC 17	1.42.22	100	10/89	93	1
Moma-Khonuu	MKUS	SMOMA	66.47	143.22	192	/83	 5 (50	1
Moneron				141.25	40	9/71	5/72	-
Mondy	MOY	мНД		100.993	1300	/58		1
Murino	MUO			104.408	470	8/66	9/66	1
Myakit	MYA	MKT		152.093	670	/83	/88	G
Mys Kozlova	MKZ			161.730				1
Mys Khvoinova			74.267	140.883		4/76	6/76	-
Mys Nerpichii			75.833	143.333		4/76	6/76	-
Mys Diring-Ayan			75.950	139.917		5/76	6/76	-
Mys Shmidta	SMT	ШМТ	68.88	179.38W	5	4/65	1/66	1
Nagornyi Sta		$H\Gamma P$	55.92	124.97	920	/77	/77	1
Nagomyi			55.95	124.92	840	/69	/69	1
Naiba	NAYS	НБ	70.85	130.73	5	/85		1
Nalychevo	NLC	НЛЧ	53.171	159.345	5	<del></del> /67	<b>-/67</b>	-
						3/69	12/69	-
						3/84		1
Naminga	NMG		56.60	118.517	1380	5/67	4/69	1
Naminga-1	NMG1	Į	56.70	118.583	1160	/63	/63	1
Nel'koba		НЛБ	61.34	148.81	531	9/83	6/97	1
				148.813	531	6/97	9/99	G
Nel'koba	NK1	НЛБ	61.34	148.81	531	6/63	1/64	1
Nelyaty	NLY	НЛТ	56.492		470	1/61		1
Nesterikha	NSR	HCT		109.708	480	7/70	9/70	1
Neryungri	NYGS		56.68	124.66	760	/77	/78	1
rveryungn	14100	1111	50.00	124.00	700	/80	/82	1
			56 657	124.723	840	11.01	OPEN	
Nezhdaninsk	MADO	шул						1
Nikola	NKO	НЖД		139.06 104.827	603 460	/80	9.71	1
Nikolaevsk-Amur		НКЛ				7.71		1
	NKL NIZ			140.783	25	9/70		l 1
Nizhnii Angarsk	NIZ	H-A		109.55	487	10/61	1/60	1
Nizhnii Armudan			50.817	142.533	150	9/66	4/69	-

Nogliki			51.817	143.15	25	10/64	12/64	-
•						/88		-
Novaya Sibir			75.050	147.000		4/75	6/75	_
Nyvrovo	NVV	HBP	54.317	142.617	5	11/81		-
O. Vrangelya	VRN	BPH	70.94	179.62W	10	2/66	4/66	3
Obo	OBO		61.80	149.77	440	/77	<del></del> /77	1
Ogon'ki				142.383	70	6/68	9/68	_
Oimur	OIM			106.833	460	10/59	-/60	1
Okha (New)	Onvi		00.000	1001020		/65		_
Okha	OKH	OXA	53.55	142.933	24	12/58	/65	_
Ol'khon	OLK	OAA	53.20	107.342	490	7/69	9.69	1
		OXT		7143.3308	430		OPEN	
Okhotsk	OHT				920			
Omchak	OCH	ОМЧ	61.67	147.87	820	9/99	OPEN 7/02	
Omolon	OMO	ОМЛ	65.23	160.54	260	6/82	7/93	1
Omolon-1	OL1	ОМЛ	65.25	160.52	260	12/63	1/65	1
Omsukchan	OS1	OMC	62.52	155.77	527	1/63	1/64	1
Omsukchan	OMS	OMC	62.52	155.77	527	12/67	OPEN	1
Onguren	ONR	ОНГ	53.233		500	<b>-</b> /88		2
Ootomari (Korsakov)	TOO		46.65	142.767	36	/09	/45	-
Oran	ORA	OPH	55.933	113.667	705	9/79		1
Orlik	ORA	ОРЛ	56.295	113.983	620	9/78		1
Orotukan	ORT		62.26	151.34	470	/77	/77	1
Ossora	OSS	OCC	59.25	163.065	10	<b>-/73</b>	<b>OPEN</b>	1
Ostrov MednyiMED		МДН	54.786	167.566		<b>-/73</b>	/75	_
Otiai (Bykov?) OTI		/		142.783	20	2/34	/45	_
Ozernaya	OZE	ОЗН		113.983	620	9/78		1
Ozernaya	AY2S	O3P	63.75	146.11	875	06/71	/71	2
Ozero	OZR	O3P		160.392		10/66	_/77	1
Ozhidaevo	OZIC	001		142.392	220	/76	/77	_
Pakhach	PCH			169.125	220	-/92	_/94	1
Palana	PAL			159.963		<del>-/94</del>	<del>-/</del> 96	1
Pauzhetka	PAU	ПЖТ		156.810	110	11/61	OPEN	1
	PET	ПТР		158.650	150	<del>-/51</del>	OPEN	
Petropavlovsk				170.27	20	5/65		1
Pevek	PVK	ПВК						
Podkova	PDK	ПДК		160.780	800	<del>-/83?</del>	OPEN	_
Polovinka	PLK			104.35	470	8/66	9/66	1
Pravda		ADD.		142.008	40	9/71	11/71	-
Provideniya	PVD	ПРВ		173.226W	25	9/80	12/93	1
Provideniya-1	PV1	ПРВ		173.18W	20	1/65	6/66	2
Romny	RMNS	SPMH		129.4	210	10/78	<b>-/87</b>	2
Russkaya	RUS	PYC	52.432	158.507	75	12/87	OPEN	1
Saidy	SAYS	СД	68.70	134.45	88	/80		1
Sasyr	SSYS	CCP	65.16	147.08	580	/86		1
Savino	SOV		52.543	102.15	720	9/68	8/69	1
Sedlovina	SDL		53.278	158.884	1235	9/91		1
Semlyachik	SEL	СМЛ	54.12	159.98		11/61	7/74	1
•								

C D-11-1-1-	CVDC	CT	55 64 100 25	505	/70	/90	1
Severo Baikalsk	SVBS		55.64 109.35	505	<b>-/78</b>	/89	1
Severo Muisk	SVK	C-M	56.183 113.533	850	11/76		2
Seymchan	SEY	СМЧ	62.9339152.3844	211	4/69	OPEN	
Shamanka	SHMS		53.125 105.6	700	8/59	2/63	1
Shara-Tagot	SRTS		53.005 106.717	500	10/59	-/60	1
Shebunino			46.433 141.858	40	9/71	11/71	_
Shikka	SKK		49.233 143.117	2	/28	/45	-
(Shikuka; Poronai	sk)						
Shimki	ŚMKS		51.675 102.012	765	11/66	11/67	1
Shipunski	SPN	ШПН	53.107 160.011	170	<b>-</b> /62	OPEN	1
Sinegor'e	SNES	СНГ	62.09 150.52	400	/76	/88	1
Slyuda	DIALO	CIII	56.33 124.12	1080	/73	/73	2
Solontsovaya	SOL	СЛЦ	54.17 108.35	458	-/79	-/87	1
•	SOL	СЛЦ	48.967 140.283	50	6/69	11/70	1
Sovetskaya Gavan	0.17	CV					1
Srednekolymsk	S-K	C-K	67.46 158.71	30	4/64	12/64	1
Srednii Kalar	SRK	КЛР	55.86 117.38	716	-/61		1
Srednii Sakutan	SDK		56.898 118.095	750	2/63	10/63	1
Stekol'nyi	STK	CTK	60.046 150.730	221	7/64	5/66	1
Stekol'nyi	MGD	MA1	60.046 150.730	221	3/71	<del>-</del> /94	1
		МГД-1	(60.046)(150.730)		/94	OPEN	1
		CTK					
Stokolviya	STV		61.8475147.6598		7/99 -	7/99	G
Stolb	SOTS	СТБ	72.40 126.82	10	/85		1
Susuman	SUUS	CMH	62.7786148.1503	640	8/69	/95	G
	2002	CCM	62.7786148.1481	640	/95	/98	G
		0.01/1	62.779 148.163	640	/98	OPEN	
Sutam			55.96 127.59	700	/69	/69	1
Suvo	SUVS	CVR	63.655 110.008	1000	-/84		ı 1
Syllakh	SYLS	СЛХ	57.12 121.86	600	/89	/89	1
•		CJIA	56.065 117.222	1000	6/62	9/63	1
Syul'ban Tabalalah	SYB TBKS	трп	67.54 136.52	200	/80		1
Tabalakh	IBK2		07.34 130.32	200	/80		1
m ' 1	m (r c	TB	70 (1 101 00	60	10.6		
Taimylyr	TMLS	ТМЛ	72.61 121.92	60	/86		1
		TMP					
Takhtoyamsk	TTY	TXT	60.20 154.68	11	9/87		1
Talaya 1and 2	TL-	ТЛА	61.1325152.3952	730	1/89	OPEN	G
	TLAR	ΤЛ					
Talaya	TAL	ТАЛ	51.681 103.644	579	11/82	OPEN	1
Tal-Yuryakh	TUR	TUR	63.307 146.634	900	3/01	6/01	G
Talon	TON	TON	59.757 148.661	20	1/01	2/01	G
TasYuryakh-1			56.64 121.33	395	02/67	03/67	1
TasYuryakh-2			56.62 121.41	415	07/67	08/67	1
Tauisk	TALIS	TAUS	59.729 149.335	5	1/01	4/01	Ğ
Tenkeli	TLIS	THK	70.18 140.78	110	/84	/93	1
Ternei	TEI	TPH	45.067 136.6	30	7/82		2
		1111	60.433 166.075	50	-/94	-/96	1
Tilichiki	TIL			150			1
Tikhmenevo			49.2 142.9	150	6/69	10/69	-

Tiksi	TIK	TKC	71.632	128.863	38	3/56	/93	1
TIXI	Tiksi-(	GSN	71.64	128.87	30	/95	OPEN	1
Tokarikan			56.10	126.42	800	/72	/73	3
Tokhoi	TKH	TX	51.361	106.608	640	11/71	4/73	1
Tonnel'nyi	TNL	ТНЛ		113.35	820	11/76		î
Topolovo	TOP	ТОП		158.041	155	11/61	<b>-</b> /93	1
Topolovo	101	ТПЛ	33.230	150.041	133	11/01	-193	1
T1 (V1)		11111				/42	/45	
Toyohara (Yuzhno)	G.D.	TYOTY	<del>-</del>			/43	/45	-
Tsipikan	ZIP	ЦРК		113.35	1110	/75	2/86	1
Tsiveluch	SVL	ШВЛ	56.583	161.225	900	10/80	OPEN	1
Tungurcha-1	TUG1		57.33	121.48	440	/70	/70	1
Tungurcha-2	TUG	THT	57.27	121.48	315	/78		1
Tungusskii	AY3S		64.20	146.38	1080	/71	/71	1
Tupik	TUP	ТПК	54.425	119.933	630	11/61		1
Turan	TRNS			101.666	870	12/66	11/68	1
Turikan	TRKS	TPK		113.108	695	8/81		1
Tymovskoe	TYV	TMC	50.85	142.65	100	4/69		_
Tyttlovskoc	TMSS	11410	30.03	142.03	100	7/07		
Timdo		ТНД	55 122	122 717	610	7/70	1/70	2
Tynda	TYD	, ,		123.717	610		1/72	2
Tyrgan	TRG	ТРГ		106.342	600	1/60	/0.0	1
Tyubelyakh	UL1S	ТЕБ	65.37	143.15	380	/88	/88	3
Uakit	UKT	УКТ		113.62	1140	12/62	/75	2
Udokan	UDK		56.75	118.305	810	4/67	4/69	1
Udzha			71.25	117.17		/27/75 C	9/19/75	-
Uelen			66.16	169.84W	5	/81	<del>-</del> /82	1
Uglegorsk	UGL	УГЛ	49.083	142.083	20	-/51		-
Ulyukchikan	UCK	УJIК	53.87	109.598	490	7/70	9/70	1
Ulyunkhan	ULNS	УЛХ	54.867	111.07	560	7/89	<b>OPEN</b>	2
Uoyan	YOA	УН	56.13	111.77	520	<b>-/79</b>		1
Ust' Nera-1	UNRI	У-НР	64.566	143.230	485	/62	/92	1
Ust' Nera-2	UNR	У-Н		143.242	485	/92	OPEN	G
Ust' Nyukzha	USZ	У-Н	56.56	121.59	415	/64		1
Ust' Urkima	UURS		56.31	123.16	540	8/81	3/99	1
Ust' Bolsheretsk	UBL	JIK		156.308	20	11/61	<del>-</del> /64	1
		У-Б	65.51	173.28	20		5/67	
Ust' Belaya	U-B					11/66		1
Ust' Srednikan SRD	1100	СРД	62.44	152.32	580	12/62	11/63	l
Ust' Omchug	USO	У-ОМ	61.13	149.63	580	/68	/83	l
Utesnoe		~~~~	46.6	143.075	20	7/73	9/79	-
Vankarem	VNK	BHK	67.84	175.85W	10	3/66	6/66	1
Verkhene Kamchatsk	VKM		54.627	158.473	170	10/66	<del></del> /75?	1
Vladivostok	VLA	ВЛД	43.12	131.893	75	<b>-/29</b>	-/31	-
Vodopadnii	VDP	ВДР	55.770	160.220	1060	<b>-/77</b>	<b>-/91</b>	1
Vzmor'e			48.85	142.517	20	7/82	12/82	-
Yablochnyi			47.167	142.067	20	6/68	9/68	-
Yagodnoe	YAG		62.53	149.62	480	/77	/77	1
Yakutsk-1	YAK1			129.722	90	10/57	/62	1
Yakutsk-2	YAK	як		129.677	91	/62	OPEN	
I WILLIAM L	T : T T T	7 A A A	02.050	L 4 7 . U / /	/ L	/ 02		

Yaruga	YRGS	ЯРГ	57.49	123.07	780	/89	/89	2
Yasnyi	YASS	ЯСН	53.29	127.983	310	1/75		1
Yubileniya	YUBS	ЮБЛ	70.74	136.10	10	/86	/93	1
·		ЮБТ						
Yuzhno Sakhalinsk	YSS	ЮСХ	46.958	142.762	100	/57	OPEN	1
Yuzhno Sakhalinsk	YSS1		47.02	142.717	40	10/47	/57	-
(Novo Aleks-andro	vsk)							
Zakamensk	ZÁK	ЗКМ	50.383	103.292	1125	12/60		1
Zapadnyi	ZAP		56.613	118.433	1600	4/67	8/67	1
Zarech'e	ZARS		52.550	107.15	460	7/59	/60	1
						7/69	9/69	
Zarya	ZRY		57.24	118.917	655	10/59	9/68	2
Zemlya Bunge			74.833	142.583		4/75	6/75	-
Zeya	ZEA	ЗЕЯ	53.755	127.293	270	6.76		1
Zhigalovo	ZGL		54.808	105.15	625	12/67	2/67	1
Zhuravlikha	ZRV	ЖРВ	53.517	109.375	475	7/70	9/70	1
Zimniki	ZMN	ЗМН	45.475	134.258	150	7/88		2
Zyryanka	<b>ZYRS</b>	ЗРН	65.72	149.82	120	/82	/90	1
Zyryanka-1	ZY1	ЗРН	65.74	150.89	37	1/64	10/64	1

# APPENDIX B

# Relocations of events used to determine crustal velocities

DATE	h	TIN m	ME s	LAT.	LONG	DEP.	MAG.	DATE	h	TIN m	ME s	LAT.	LONG	DEP.	MAG.
700321	11	21	10.84	56.92N	133.18E	6.8		730130	22	53	22.38	56.69N	128.83E	0.	
					127.50E			730222					123.91E		Ċ
					124.78E			730429					124.72E	9.5	
					126.75E	9.1							152.87E	0.	
710202	12	11	40.74	63.49N	146.12E	8.4		730527	05	49	09.67	62.23N	161.72E	2.5	
710220	14	05	13.95	56.41N	127.41E	Ο.		730608	11	43	47.07	63.82N	146.17E	26.1	
710318	20	08	41.2	55.77N	133.48E	3.4		730616	02	12	18.31	59.99N	152.63E	11.1	
710414	11	18	03.2	56.54N	121.07E	7.1		730618	13	59	47.7	56.84N	121.05E	9.2	
710518	22	44	41.90	64.05N	145.76E	33.1	7.1	730727	15	09	19.99	60.19N	153.55E	0.	
710518	23	09	08.04	64.14N	146.05E	1.2							145.62E	4.6	
					145.79E	4.2							148.05E		
					145.84E								139.59E	0.8	
					145.83E	0.							154.41E		
					145.68E	1.2							142.99E	0.	
					145.61E	4.7	•	731022					121.05E		
710614					123.70E			731026					123.55E		
710614					123.67E 146.13E								145.76E 125.71E	7.2	•
					127.63E	1.4	•						125.71E		
710801					127.63E								143.25E	0.	•
710812					123.35E	3.0	•						153.11E	0.	
					142.20E								153.23E	5.8	
					142.34E								156.78E		
					150.37E	5.9							127.79E	8.7	
711004	12	59	08.33	66.12N	142.11E	0.0		740516	09	05	19.6	56.16N	123.74E	7.2	
711008	03	27	30.6	56.26N	123.60E	16.5		740518	06	45	06.25	59.26N	148.66E	30.0	
720113	17	24	17.07	61.63N	147.11E	0.9		740619	03	09	36.78	63.26N	151.10E	30.0	
720113	22	20	12.45	61.86N	147.11E	9.0		740629	14	33	23.9	56.55N	121.05E	7.5	
720115	18	80	00.1	57.69N	121.10E	4.5		740702	09	19	19.70	66.85N	139.51E	20.0	
720117	09	22	49.64	61.91N	147.12E	7.9		740702	17	04	03.15	66.53N	138.27E	12.2	
					147.17E								127.23E		
					127.24E								152.49E		
					145.01E								127.24E	0.	-
720214					122.14E								139.41E		
					145.97E	6.1							140.21E		
720401					121.08E		•						125.79E	0.	
					134.83E	1.8		741217					153.15E 145.89E		
720421					151.47E 121.18E		•						150.41E	5.4	•
					147.11E	4.7	•						161.23E		•
					139.31E								129.02E		
					124.13E								122.96E		
					128.64E								146.09E		
					146.10E								132.13E		
					126.46E								126.95E		
					127.66E								122.26E		
					145.79E								122.33E		
					127.80E								143.43E		
720811	12	45	21.1	53.95N	128.19E	7.1		751103	12	06	58.01	57.70N	126.24E	0.	
720831	13	43	17.2	56.07N	124.56E	7.2		751104	12	41	06.83	59.81N	160.17E	10.3	
720908	19	48	53.4	56.00N	130.60E	0.0							144.28E		
721221	02	12	57.58	62.82N	140.45E	4.0		751227	03	54	20.52	62.06N	154.92E	10.9	

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760121 06 01 48.46 67.69N 140.27E 22.7 5.3
                                             790104 21 50 16.7 54.18N 122.96E 29.2
760216 22 06 23.82 58.17N 131.02E 4.4 4.6 790221 15 03 27.35 63.37N 146.28E 6.6
760226 17 23 03.3 56.53N 121.10E 12.8 .
                                             790427 19 38 13.8 55.76N 130.47E 12.7
760408 04 20 58.28 67.03N 139.69E 10.0 .
                                             790428 09 34 06.23 57.45N 126.92E 10.7
760426 20 12 21.41 61.77N 155.97E 13.5
                                             790502 22 10 28.88 64.19N 149.21E 2.5
760508 11 39 25.77 60.09N 152.72E 2.3 .
                                             790517 06 39 09.4 53.70N 126.09E 8.3
760520 16 53 29.2 56.36N 132.48E 0.0 .
                                             790523 17 37 33.9 53.61N 126.09E 21.2
760614 11 58 01.9 55.57N 130.72E 5.9 .
                                             790531 10 49 26.2 55.68N 130.26E 14.5
760624 17 58 01.07 59.94N 157.70E 33.2 .
                                             790531 11 52 35.3 56.51N 124.88E 13.8
760627 16 03 23.8 55.39N 131.22E 1.8 .
                                             790616 20 41 48.41 64.89N 143.67E 26.7
760704 23 23 31.80 65.60N 134.98E 25.7 .
                                            790618 22 09 57.41 57.50N 126.06E 5.3
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940128 01 05 12.10 55.06N 135.29E 6.6 5.0 960212 18 56 59.9 $7.53N 120.77E 26.2
940130 08 18 38.9 57.53N 122.37E 33.0 . 960220 23 32 54.09 61.23N 158.42E 25.4
940204 17 10 10.46 59.66N 143.93E 10.1 .
                                          960303 10 53 57.77 58.26N 158.17E 4.4
940215 17 40 57.09 57.86N 128.57E 8.2 .
                                          960303 11 54 41.62 58.26N 158.04E 33.0
940216 12 43 17.5 56.96N 123.51E 23.8 .
                                          960318 20 15 14.89 58.35N 143.97E 30.6
940223 17 56 12.2 57.30N 123.25E 17.4 .
                                         960403 23 39 29.97 62.85N 147.32E 6.1
940301 17 48 03.00 55.13N 138.47E 0. .
                                         960406 00 20 39.37 62.83N 147.63E 11.6
940317 13 55 31.6 57.06N 121.57E 26.5 . 960606 21 53 42.36 63.02N 144.73E 6.0
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980326 02 24 00.38 62.08N 156.93E 3.2
960612 17 08 01.2 56.84N 121.24E 24.9 .
960707 10 50 00.61 58.58N 157.54E 12.2 5.8
                                          980326 06 41 33.48 62.09N 156.73E 8.7
960707 14 19 52.14 58.63N 157.34E 6.3 .
                                           980327 09 29 59.56 62.07N 156.87E 5.6
960707 14 29 03.55 58.56N 157.50E 12.0 .
                                           980327 09 42 36.02 62.09N 156.95E 7.2
960707 15 39 03.11 58.28N 157.42E 31.6 .
                                           980404 14 02 34.13 62.05N 156.76E 7.6
960707 17 58 34.14 58.53N 157.49E 3.9
                                           980505 21 31 57.75 59.05N 147.14E 5.1
960707 23 23 49.02 58.50N 157.32E 16.8
                                           980508 22 32 49.88 62.02N 157.03E 4.9
960708 01 06 25.24 58.47N 157.28E 0. .
                                           980531 08 02 51.19 62.04N 156.54E 11.7
960708 13 29 51.71 58.59N 157.38E 15.1
                                           980604 12 03 23.75 62.06N 157.18E 2.3
                                      .
960709 04 07 41.83 58.43N 157.17E 8.4 .
                                            980901 10 26 39.52 63.06N 147.28E
960709 10 07 05.5 56.61N 120.95E 7.4 .
                                            980917 22 50 11.78 64.95N 149.24E 0.
960709 11 52 51.98 58.33N 157.19E 0. .
                                           981106 03 48 20.34 63.40N 158.75E 12.6
960710 05 20 11.74 58.45N 157.70E 0.8
                                           981218 14 03 48.02 63.23N 159.45E 4.2
960710 06 22 05.38 58.52N 157.34E 18.2
960710 07 15 54.93 58.67N 157.45E 15.2
960710 14 09 33.80 58.57N 157.64E 3.9
960710 18 36 37.17 58.46N 157.30E 28.1
960727 07 28 34.95 60.49N 148.66E 7.4 3.6
960803 12 32 49.64 58.65N 157.21E 29.4
960803 13 09 13.64 58.68N 157.08E 0.
960807 18 51 14.48 58.58N 157.21E 7.6
960808 17 09 40.38 58.75N 157.20E 22.2
960824 07 29 41.80 60.05N 153.05E 17.7
960902 23 37 54.6 56.60N 123.88E 19.4
960904 12 56 31.38 57.55N 128.03E 15.1
960913 15 45 07.54 58.70N 157.64E 6.5
960914 05 29 59.84 58.65N 157.41E 9.8
960914 05 41 45.02 58.73N 157.48E 8.3
960914 09 57 57.82 58.74N 157.14E 0.
960916 03 05 07.56 58.56N 157.35E 14.6
960925 01 48 26.12 63.38N 150.43E 9.9
961024 19 31 50.91 67.04N 173.08W 0. 6.0
961024 21 57 36.71 67.08N 173.20W 14.1
961126 20 18 18.31 58.66N 157.54E 5.8
961207 10 38 20.98 62.12N 153.75E 10.6
970129 15 51 38.11 58.81N 149.65E 5.4
970304 21 49 26.30 62.07N 155.92E 10.0
970607 11 59 41.19 64.19N 148.32E 0.
970614 13 31 08.39 63.87N 148.57E 2.9
970616 09 54 07.72 64.07N 148.43E 9.0
970629 21 37 46.82 59.78N 152.40E 4.6
970721 12 37 49.52 60.08N 144.68E 10.2
970825 06 00 43.24 63.54N 144.91E 8.2
970910 18 42 42.47 61.77N 156.15E 1.5
970914 14 07 54.73 61.10N 145.23E 0.
970915 05 48 00.77 59.95N 151.96E 0.
971122 11 41 15.58 61.14N 155.44E 7.7
971206 16 06 07.04 59.40N 147.91E 10.9
980103 03 15 51.98 59.81N 152.50E 0.
980130 23 11 39.97 63.48N 150.29E 20.1
980201 02 33 13.50 63.24N 150.00E 2.4
980215 04 35 42.38 61.49N 147.43E 0.
980221 22 22 47.82 60.33N 152.86E 5.4
980304 04 57 18.17 62.08N 156.99E 6.3
980309 10 38 11.91 63.87N 156.95E 0.
980311 23 46 07.01 58.46N 157.22E 0.
980312 03 15 54.71 62.10N 157.03E 1.5
980314 17 03 29.87 58.40N 157.41E 3.5
980316 05 33 08.13 58.42N 157.30E 13.7
980318 21 30 34.19 58.39N 156.98E 8.3
980320 16 50 52.81 62.06N 157.06E 7.9
980325 23 52 55.66 62.11N 156.94E 5.0
980326 02 17 29.32 62.09N 157.05E 3.4
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8.4

# APPENDIX C

Relocations and GT classifications of Teleseismically recorded earthquakes from eastern Russia. Each location was relocated using the crustal models calculated in Section 3. Each event listed had a significant amount of local/regional data added, supplementing that reported in the ISC. ISC calculated parameters are given for reference. Location types indicate which data set was considered best for calculating relocations: T - teleseismic, R - regional, L - local. See text in Section 4 for additional discussion on selection of the best data set.

DATE	MSU RELOCATIONS	ISC SOLUTION	MAG LOC GT
	ORIGIN TIME LAT. LONG DEPTH	ORIGIN TIME LAT. LONG. DEPTH	
	h. m. s.	h. m. s.	
710518	22 44 42.23 63.960 146.064 20.85	22 44 39.25 63.921 146.101 0.0	7.1 T 25
710518	23 09 09.20 64.029 146.100 2.62	23 09 09.46 64.181 146.450 0.0	. T
710614	13 48 52.96 56.200 123.636 11.91	13 48 53.66 56.183 123.594 15.0	5.4 T 25
710614	14 25 54.16 56.187 123.654 12.56	14 25 56.90 56.171 123.562 33.0	4.4 T
720113	17 24 19.62 61.902 147.034 9.74	17 24 23.17 61.940 147.038 33.0	5.3 T 25
720115	18 08 00.31 57.718 121.062 8.83	18 08 00.79 58.073 120.748 0.0	4.6 R
720330	20 20 52.45 64.004 145.969 4.98	20 20 50.17 63.769 145.994 0.0	. T 25
720613	10 45 03.27 54.849 126.435 18.75	10 45 03.19 54.911 126.460 0.0	4.9 T 25
720809	20 51 47.59 56.945 127.625 19.38	20 51 51.76 56.838 127.410 36.0	4.7 R 25
730608	11 43 47.07 63.822 146.173 26.07	11 43 48.25 63.776 146.200 33.0	. T
731102	07 31 28.80 54.094 125.697 1.63	07 31 32.86 54.042 125.746 0.0	4.9 T
740619	03 09 37.00 63.236 151.037 30.64	03 09 36.12 63.140 150.921 14.7	4.9 R 25
750629	12 24 39.55 53.110 132.134 4.74	12 24 43.28 53.074 132.111 33.0	4.8 T 25
751104	12 41 05.39 59.827 160.326 5.08	12 41 12.39 60.020 160.317 52.3	4.7 T
760121	06 01 48.56 67.731 140.194 18.98	06 01 48.51 67.734 140.034 0.0	5.0 T 25
760216	22 06 24.32 58.170 131.044 0.00	22 06 32.07 58.573 131.012 33.0	4.5 T
770816	13 56 54.67 53.878 128.875 0.00	13 56 59.80 53.926 128.715 33.0	4.2 T
771101	03 54 22.45 55.556 130.641 3.84	03 54 26.14 55.411 130.522 31.7	4.5 T
771118	21 55 36.34 60.142 143.423 9.54	21 55 39.45 60.052 143.321 33.0	4.5 T 25
780605	07 05 52.16 60.042 160.361 3.78	07 05 59.37 60.131 160.406 55.8	5.1 R 25
780605	21 01 34.09 60.089 160.385 8.06	21 01 37.40 60.163 160.390 28.2	4.5 R 25
790427	19 38 13.16 55.797 130.528 15.46	19 38 18.41 55.935 130.166 32.8	4.6 R 25
790819	07 10 06.14 61.326 159.131 29.50	07 10 06.67 61.329 159.041 30.7	5.1 T 25
791007	01 29 26.09 65.003 144.016 12.29	01 29 26.26 65.019 143.774 6.1	4.8 R 25
791026	00 26 26.18 62.238 153.783 19.01	00 26 28.60 62.258 153.319 27.9	4.5 T 10
791118	08 59 15.88 62.213 153.801 2.02	08 59 19.08 62.302 153.318 20.6	. T
800225	23 50 02.71 54.892 125.191 30.98	23 50 00.58 55.143 125.047 3.0	4.0 R 10
810522	04 59 21.03 61.098 156.699 9.22	04 59 25.89 61.148 156.624 50.6	5.1 R
811108	21 56 09.29 61.821 153.675 20.81	21 56 11.50 61.834 153.686 37.4	5.6 R 10
811110	10 07 52.32 63.954 148.697 24.38	10 07 52.73 63.862 148.484 33.0	4.7 R 25
811205	01 49 08.92 56.996 123.143 18.0	01 49 10.84 57.356 122.930 36.9	. R 25
820903	07 29 26.40 66.933 133.319 29.76	07 29 22.17 66.823 132.776 0.0	4.5 T 25
830325	10 36 55.41 63.580 149.908 5.40	10 36 56.85 63.632 149.740 15.7	4.7 R 10
830514	08 09 13.46 65.852 174.822 0.00	08 09 22.26 65.811 175.975 33.0	4.5 T 25
830515	06 00 36.88 65.861-173.216 2.58	06 00 43.30 66.031-172.540 33.0	4.2 T 25
830730	15 42 09.76 53.303 132.479 22.24	15 42 12.20 53.242 132.642 26.0	4.4 R 25
840324	01 03 19.24 54.517 136.970 23.26	01 03 17.81 54.512 137.036 3.0	5.0 R 25
840802	21 25 39.14 60.897 144.640 8.02	21 25 37.78 60.887 144.683 0.0	4.9 R 25
841029	14 06 14.27 62.113 163.756 26.73	14 06 14.59 62.049 163.736 33.0	4.6 T
841122	13 52 57.11 68.473 140.821 18.71	13 52 57.17 68.521 140.815 19.1	5.3 R 25
841202	08 35 45.04 63.383 150.502 22.06	08 35 47.94 63.461 150.525 43.9	5.2 R 10
850129	00 36 05.81 64.159 145.778 10.17	00 36 09.38 64.286 145.672 33.0	4.5 L 10
850201	08 31 45.86 62.892 127.241 12.63	08 31 45.85 62.843 127.381 3.0	4.6 T
850301	09 19 18.42 57.556 125.461 10.60	09 19 21.89 57.667 125.573 33.0	4.6 R 10

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850602
        04 08 09.37 64.846 144.052 8.26
                                          04 08 12.81 64.788 144.045 33.0
                                                                                R 25
850624
        03 54 34.78 65.257 144.698 17.58
                                          03 54 32.38 65.252 144.489 0.0
                                                                           4.6 R 25
860209
        12 42 53.73 48.770 126.467 13.77
                                          12 42 51.24 48.685 126.625 0.0
                                                                           4.8 T 25
        17 07 24.63 48.699 126.543 0.00
                                          17 07 24.42 48.645 126.656 0.0
                                                                           4.8 T 25
860228
                                                                           4.3 T 25
860815
        17 53 08.20 48.691 126.621 19.22
                                          17 53 14.51 48.181 126.434 33.0
                                          20 20 37.80 48.676 126.559 33.0
                                                                           5.1 T 25
        20 20 33.58 48.755 126.492 3.70
860815
                                                                           5.3 T 25
                                          18 30 56.86 63.897-178.693 6.4
861019
        18 30 58.82 63.894-178.546 19.79
                                          18 04 10.22 61.312 143.729 0.0
                                                                           4.6 R 25
        18 04 12.17 61.216 143.719 25.85
861218
        00 58 20.89 62.849 156.863 5.68
                                          00 58 23.59 62.877 156.699 22.8
                                                                           4.4 R 10
870211
                                          12 39 56.90 64.829-169.398 33.0
                                                                                T 25
870421
        12 39 48.43 64.922-170.428 12.58
                                                                               T 25
881013
        00 32 10.79 61.769 169.699 21.07
                                          00 32 12.87 61.853 169.651 33.2
                                                                           5.4
        04 16 22.76 59.828 145.125 14.64
                                          04 16 27.35 59.916 145.230 0.0
                                                                           4.8 R 25
890409
        22 59 53.63 57.245 122.146 29.30
                                          22 59 54.24 57.168 122.015 26.0
                                                                           5.9 R 25
890420
890421
        19 08 37.29 57.067 122.269 33.00
                                          19 08 41.00 57.168 123.768 33.0
                                                                           4.8
                                                                                L 10
                                          01 34 01.57 57.278 122.155 39.6
890424
        01 33 59.74 57.111 122.277 24.88
                                                                           4.8 R 10
                                          06 25 38.78 57.175 122.159 26.1
890429
        06 25 39.17 57.149 122.229 33.00
                                                                           5.3
                                                                                R 10
890507
        16 28 05.70 57.075 122.233 33.00
                                          16 28 05.22 57.123 122.162 27.0
                                                                           5.0
                                                                                R 10
890517
        05 04 35.84 57.056 122.237 30.69
                                          05 00 35.56 57.089 122.107 26.9
                                                                           5.6
                                                                                R 10
890517
        15 55 22.85 57.063 122.241 32.05
                                          15 55 56.12 61.804 119.599 33.0
                                                                           4.5 R 10
        20 07 46.33 57.098 122.309 33.00
                                          20 07 49.17 57.348 121.903 52.1
                                                                           4.8 R 10
890709
890723
        12 01 31.20 54.532 124.934 25.26
                                         12 01 29.97 54.498 125.054 10.0
                                                                           4.6 R 10
        03 37 25.23 48.707 132.023 18.95
                                         03 37 26.30 48.857 131.928 15.0
                                                                           4.5 R 25
900717
        21 54 02.63 64.785 146.640 6.97
                                          21 54 03.19 64.931 146.604 10.0
                                                                           4.6 R 10
901102
        08 15 37.55 56.844 120.386 21.56
                                         08 15 35.50 56.998 120.405 5.2
                                                                           4.6 R 10
910109
        20 47 08.66 60.644 166.940 0.0
                                          20 47 13.91 60.970 166.830 33.0
                                                                           4.1 T --
910217
        11 13 48.52 60.861 167.026 12.39
                                         11 13 43.75 60.089 165.784 33.0
                                                                           4.8 T --
910218
                                         04 46 37.17 60.711 167.125 33.0
        04 46 32.49 60.841 167.086 0.0
                                                                           4.5 T 25
910220
        18 12 28.02 54.705 121.041 33.00
                                         18 12 29.75 55.015 121.116 33.0
                                                                                T 25
910304
                                                                           5.2 T 25
       09 02 19.61 60.934 167.009 15.26
                                         09 02 20.45 60.815 167.087 23.1
910308
       11 36 30.89 60.928 166.988 33.00
                                         11 36 28.39 60.856 167.016 13.0
                                                                           6.7 T 25
910308
                                         11 54 57.86 60.810 167.058 10.0
                                                                           6.1 T 25
910308
       11 54 58.49 60.769 167.063 16.68
                                         12 26 46.07 60.914 167.261 14.2
                                                                           5.1 T 25
910310
       12 26 44.68 60.948 167.289 6.10
910311
       15 33 46.34 60.762 166.857 33.00
                                         15 33 46.27 60.763 166.897 33.0
                                                                           4.7 T 25
                                         19 46 46.2 60.557 166.818 5.4
                                                                           4.8 T 25
910311 19 46 50.60 60.663 166.770 33.00
                                         19 43 25.86 60.825 167.116 29.5
                                                                           5.2 T 25
910312
       19 43 25.03 60.889 167.110 23.93
       06 58 43.34 60.781 166.863 33.00 06 58 43.2 60.651 167.241 33.0
                                                                           4.2 T --
910313
        06 26 52.06 60.809 167.007 33.00 06 26 52.16 60.814 167.003 33.0
                                                                           5.2 T 25
910317
        14 28 33.45 60.619 167.438 28.14 14 28 35.27 60.890 167.022 33.0
                                                                           4.5 T --
910321
       07 55 16.90 60.728 167.166 7.31 07 55 20.83 61.078 166.481 33.0
                                                                           4.7 T --
910322
910322
        17 18 22.79 60.836 167.543 5.84 17 18 26.50 60.027 167.136 33.0
                                                                           4.4 T --
                                         13 27 52.05 60.709 166.921 32.5
                                                                           5.0 T 25
        13 27 50.44 60.723 166.993 21.40
910329
                                                                           4.2 R 10
910527
        08 55 35.05 57.088 122.209 31.16
                                         08 55 33.55 57.060 122.129 33.0
       19 08 06.24 65.477 143.229 21.87
                                         19 08 07.24 65.467 142.905 33.0
                                                                           4.3 R 10
910804
920121
        18 07 35.63 61.764 169.741 11.19
                                         18 07 39.36 61.832 169.574 33.0
                                                                           4.6 T 25
                                         06 29 17.10 65.806 142.990 33.0
                                                                           4.6 R 25
920122
       06 29 16.67 65.855 143.255 20.06
                                         22 47 22.76 56.693 121.157 33.0
920413
       22 47 21.21 56.569 121.112 14.35
                                                                           4.1 R 10
920913
        21 42 56.24 62.065 153.795 10.20
                                         21 42 57.64 61.634 154.105 21.3
                                                                           4.7 R 10
                                         09 37 54.70 58.289 120.613 33.0
930221
        09 37 53.24 57.954 120.714 24.42
                                                                           4.2 R 25
930324
        22 43 29.89 71.659 130.225 19.09
                                          22 43 29.78 71.686 130.405 20.3
                                                                           4.8 T 25
                                         07 56 34.88 64.162 145.941 10.0
                                                                           4.2 R 10
930830
        07 56 35.56 64.068 145.825 6.37
                                         01 05 13.78 55.054 135.175 15.3
940128
        01 05 13.17 55.077 135.226 12.46
                                                                           4.9 T 25
                                         16 53 17.56 63.882 133.722 10.0
                                                                           4.5 T 25
950209
        16 53 19.00 63.653 133.925 23.53
                                         17 29 38.30 50.960 125.290 6.0
                                                                           3.5 T --
        17 29 26.41 51.673 125.602 27.51
950301
950815
        01 53 23.06 58.122 148.817 0.00
                                         01 53 25.85 58.069 148.844 16.2
                                                                           4.6 T 25
                                          18 57 01.18 57.656 120.677 33.0
                                                                           4.3 T 25
960212
        18 56 59.87 57.527 120.771 26.12
                                          23 33 01.86 60.650 159.011 33.0
                                                                           3.6 T --
960220
        23 32 57.92 61.215 158.014 15.99
                                          10 49 59.59 58.585 157.697 10.0
                                                                           5.6 T 10
960707
        10 49 59.79 58.611 157.794 13.00
960707
        23 23 47.20 58.591 157.812 8.20
                                          23 23 53.59 58.689 157.838 59.0
                                                                           4.1 T 25
960708
        13 29 49.65 58.641 157.798 7.29
                                          13 29 59.58 58.730 157.738 87.7
                                                                           4.0 T 25
        05 20 14.54 58.614 157.660 8.03
                                          05 20 15.66 58.649 157.942 11.0
                                                                           3.6 T --
960710
960710
        07 15 52.00 58.637 157.780 5.86
                                          07 15 56.25 58.685 157.718 35.3
                                                                           4.5 T 25
960918
        09 51 41.77 67.906 140.171 0.0
                                          09 51 46.82 68.013 139.821 33.0
                                                                           4.6 T 25
        16 09 18.79 62.092 153.901 6.29 16 09 20.4 62.410 153.709 10.0
                                                                           4.5 R 25
961003
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961024	19	31	50.64	67.026	-173.119	0.00	19	31	54.1	66.96	-173.24	20.0	6.1	Τ	25
961024	21	57	37.27	67.068-	173.257	12.77	21	57	37.5	67.08 -	-173.28	18.0	4.9	Т	25
961103	23	24	33.77	64.754	-170.281	33.00	23	24	30.7	64.84 -	-170.41	10.0	4.9	T	25
961126	20	18	17.03	58.681	157.725	11.10	20	18	17.6	58.743	157.773	15.4	4.0	$\mathbf{T}$	
961129	04	38	50.56	64.376	145.366	10.43	04	38	53.82	64.464	145.389	33.0	3.4	Т	10
970618	06	13	43.11	64.934	137.934	11.53	06	13	46.85	65.033	137.258	38.6	4.0	T	25
970620	04	57	02.43	64.759	146.963	18.24	04	57	04.32	64.759	146.887	33.0	3.6	Т	10
970902	22	30	10.65	57.629	137.669	29.28	22	30	07.71	57.658	137.823	10.0	3.9	Т	
971019	16	58	18.17	57.321	120.600	9.79	16	58	26.43	57.759	120.216	51.1	3.8	Т	
971021	00	18	08.71	57.303	120.706	22.09	00	18	12.22	57.026	120.055	46.1	3.8	R	
971024	11	50	13.43	57.341	120.743	24.51	11	50	17.52	57.232	120.174	47.4	4.3	R	25
971024	11	58	11.55	57.316	120.680	33.00	11	58	12.78	57.117	120.226	22.3	3.6	L	
971024	12	05	46.25	57.337	120.749	12.53	12	05	46.32	57.131	120.291	10.0	4.0	R	
971024	12	19	03.53	57.356	120.560	14.07	12	19	04.73	57.338	120.590	25.0	4.1	Т	
971024	12	52	50.76	57.267	120.492	18.60	12	52	53.84	57.290	120.035	42.2	4.4	R	25
971024	13	14	04.54	57.157	120.441	33.00	13	14	03.77	57.096	120.483	31.9	3.4	Т	
971106	19	50	29.56	57.407	120.662	0.99	19	50	35.72	57.421	120.436	54.4	4.1	R	25
971106	20	24	39.83	57.342	120.656	12.61	20	24	45.03	57.348	120.229	33.0	3.5	R	
971108	15	07	08.12	57.336	120.667	9.93	15	07	12.86	58.461	120.766	0.0	3.6	Т	
971122	11	41			155.498	8.68	11	41			155.442		3.7		25
980314	17	03	29.69	58.491	157.636	9.45	17	03			157.581	28.4	4.2	$\mathbf{T}$	25
980316	05	33	07.58	58.442	157.461	10.04	05	33	11.69			44.	4.2	_	25
980318	21	30			157.307	2.67	21	30	20.2		156.670	0.0	4.1	_	
980326	06	41			157.150	0.0	06	41	31.76	62.001	157.352	10.0	4.3	_	
990827	20				147.292	3.6	20		44.32		147.21	33.0	4.1	-	
991002	07	01			148.095	5.1	07	01	27.73		148.113	43.8	4.0	_	
000927	13	26	59.59	62.443	145.227	7.6	13	26	59.6	62.433	145.193	10.0		T	

# APPENDIX D

Digital station deployment sites and observations

# **OVERVIEW**

In the summers of 1999 and 2000, as well as winter and summer 2001, the senior author traveled to Magadan to upgrade existing photo paper seismic stations to digital acquisition. This work was performed in conjunction with the Magadan Experimental Methodological Seismological Division (MEMSD), from Magadan, Russia.

A total of nine digital acquisition systems have been purchased and imported into Russia for deployment. The digital acquisition systems were manufactured by PC System Design, Palo Alto, California, and use 8 channel, 24 bit A/D cards with GPS timing. Data are recorded on PC computers, which were purchased in Russia. Except as noted, the seismometers recorded are three-component Russian (Soviet) SM3-KV short period instruments, with the free period set at 1.5 seconds. Seismometer output is amplified 1000 times, and a 10 Hz cutoff low pass filter is used. The amplifier/filter was designed and manufactured by MEMSD. Digitization of time and all seismometer components is 30 s.p.s. A description of each station follows.

# PERMANENT STATIONS

ANADYR. Anadyr is located several thousand kilometers east and a bit north of the other deployed stations. The station recorded a Guralp CMG 40T broadband seismometer, upgrading an existing analog photopaper station. The station vault consists of a concrete pad set into permafrost underneath the station building. The recording computer failed shortly after installation in March, 2000, and was logistically problematic to repair. The station was removed in May, 2001.

MAGADAN. Although not installed as a part of this study, data from data from the IRIS GSN station Magadan (MA2), has actively been used. There are two station sites in Magadan. The original Magadan station (MAG), and headquarters of the Magadan network, is located in a residential part of the city about 1 km south and a little west from the center of town. The station consists of a single story, moderately sized, wooden, office building. The seismic vault is located under the station in a cellar about 5 m below the ground surface. The vault contains a large concrete pad measuring about 2 m per side mounted in a rocky soil. This old station site in Magadan had considerable noise due to vault conditions and cultural noise. The vault is no longer used except for instrument testing purposes.

The new GSN station Magadan (MA2) is recorded at the old Magadan station site. The vault of the new Magadan GSN station is located on top of a mountain about 2.5 km to the northwest of the old station. The vault consists of a bunker set into bedrock of granitic composition. Seismometers used at MA2 are Streckeisen STS-1.

**NEL'KOBA**. Nel'koba is a small town whose primary function was a regional supply and repair center for support of the gold mining industry in the region. The seismic vault in Nelkoba consists of a concrete pad measuring approximately 1 m x 1 m set into permafrost. The concrete pad is

housed in a small wooden shed on the grounds of the Nelkoba kindergarten, on the west side of town. The current station site was constructed in the summer of 1997 and is approximately 200 m north of a previous site. At some time in the past, SKM instruments were operated at the old station site. The town of Nelkoba was permanently abandoned in late September, 1999. Our equipment was moved to establish the station in Omchak, near the Matrosova temporary station deployment site (see below).

**OKHOTSK**. Okhotsk is located on the north coast of the Sea of Okhotsk. The seismic station is located in the town's telecommunication center and operated under contract from the Magadan EMSD. The Guralp CMG 40T broadband sensor is deployed in a small shelter and set on a 1 meter square concrete pad set 1.5 meters into rocky ground on a terrace above the coast of the Sea of Okhotsk. The station at Okhotsk is a new site, opened in early July, 2000.

OMCHAK. Omchak is a mining town with a large refinery operation that supports the Matrosova gold mine a few kilometers away. The station at Omchak is a new station site using equipment moved from the station in Nel'koba after the town was abandoned. Seismometers at Omchak sit on a concrete pad directly on bedrock in an abandoned mine adit. Background noise levels at Omchak are the lowest of all stations deployed in this study. The Omchak station frequently records mine blasts from the Matrosova gold mine a few kilometers distant.

OMSUKCHAN. Omsukchan is a local administrative center in the eastern part of the Magadan region. The digital station was deployed in mid July 2001, using the Guralp CMG-40T removed from Anadyr. The station is situated on a 1.5 meter square concrete pad at the bottom of a blind stairwell adjacent to a large apartment building near the center of town. The seismometer is approximately 2 meters below ground level. Omsukchan frequently records blasts from nearby coal, gold and silver mines

SEYMCHAN. Seymchan is one of the longest running seismic station sites in the Magadan district, opening in 1969. A GEOSCOPE station was installed here in the early 1990's, which broke after about 4 years of operation due to a failed tape drive to record data. The station was not repaired, but all equipment, including the STS-1 seismometers, remains in place. The station deployed here as a part of this study digitizes an analog output from the GEOSCOPE station and logs the data on a computer. At Seymchan, the seismometers are located in an underground permafrost vault, set on a concrete pier that extends several meters into permafrost. Temperatures remain constant in the vault at around minus 1° C. The station at Seymchan is situated within the town not far from the power plant. Considerable noise is recorded by the Seymchan station when the power plant grinds coal.

SUSUMAN. Susuman is a town of approximately 5,000 people and is the center of placer gold production in the Magadan region. The seismic station is operated by MEMSD and is located at the meteorological station just west of the main town. The station is sufficiently far from the town that cultural noise is minimal. The seismic vault in Susuman consists of a large concrete block set approximately 1.5 m into a permafrost foundation. The vault is inside a small building adjacent to the meteorological station building.

**TALAYA** Talaya is a resort town of approximately 500 people; the main attraction is the natural hot springs. Two digital seismic stations have operated here, both by MEMSD. The seismic vault

at the old Talaya site consisted of a 1 m square concrete pad set into volcanic bedrock on the side of a hill. The seismic station was located on the east side of town next to the now abandoned cinema theater. This Talaya site was abandoned in September, 2000 as electricity and heat were cut off from the abandoned portion of the town where the station was located.

The new Talaya site was established approximately 300m ESE from the old Talaya site, also in September, 2000. The site is located in the town's new administrative/hospital building that is housed in the old school. The seismic vault consists of a 2 m square concrete pad set into the ground below the building, which consists of sand and rocky material. Distance to bedrock is uncertain. Background noise levels are similar to the old Talaya site.

UST' NERA. The seismic station in Ust' Nera is operated by the Yakutsk seismic network. The station in Ust' Nera has occupied its present site since 1992. The seismic vault consists of a large concrete block set into a permafrost foundation. The vault is located between two large apartment buildings, thus cultural noise can be high at times due to passing cars and children playing on top of the vault. Seismometers recorded here are Russian (Soviet) SKM short-period instruments. Seismometer output is amplified 60 db and a 30hz cutoff low pass filter is applied. The amplifier used in Ust' Nera is a USGS Prototype Series Seismic Amplifier. The amplifier was originally part of an IASPEI system installed in Ust' Nera in 1997.

# TEMPORARY STATIONS

MATROSOVA (61.6432° N, 147.8205° E). Matrosova is a small town with an operating gold mine. Seismometers were located on the concrete foundation of a building used for ore transfer. The foundation of the building was set into bedrock consisting of black shale. In Matrosova, digitization of time and all seismometer components was 120 s.p.s. The station was operated for only a few hours, to record blasting from the mine. Two large blasts of 1,300 and 1,500 kg of Ammonite were recorded, among several much smaller explosions. Several abandoned mine adits are also in and near Matrosova, which may be good sites for future permanent stations.

STOKOLVIYA (61.8475° N, 147.6598° E). Stokolviya is a hydrological research station in a remote, unpopulated region. The seismic vault consisted of a one meter pit dug into the side of a hill. Ground material consisted primarily of angular cobble sized rocks of volcanic origin. The material was consistent with the bedrock surface being close. When installing the station, it was intended that Stokolviya would be a permanent site. All permissions were obtained, and the hydrologic research station agreed to operate it prior to installation. However, the individual workers at the station were unfamiliar with computers and refused to consider operating the station. This was unfortunate, as the site is quiet. Also at Stokolviya are several boreholes, some of which exceed 200 m depth. These boreholes are not is use, thus may be ideal for future borehole instrument installation. The temporary station deployed at Stokolviya recorded two explosions from the Matrosova gold mine.

**TAUISK** (59.729°N, 149.335°E). Tauisk is a coastal village west of Magadan whose primary function is a small army frontier border patrol post. The station at Tauisk was deployed from January 18 - April 18, 2001 to record aftershocks of the January 07, 2001 Mb 5.3 earthquake that occurred about 80 km west of Magadan. The station was situated on the concrete slab in the

basement rifle firing range of the school, which is located in the center of the village. Many earthquake aftershocks were recorded in Tauisk.

TALON (59.757°N, 148.661°E). The seismic station deployed at Talon was installed in the dirt basement of a house in the south edge of the village to record aftershocks of the January 07, 2001 Mb earthquake that occurred about 80 km west of Magadan. The station operated from January 26 - February 16, 2001. The station was quite noisy and only a few events were recorded.

TAL'YURYAKH (63.307°N,146.634°E). Tal'Yuryakh was deployed a few kilometers northwest of the active pit at the Tal'Yuryakh coal mine to obtain blasting information. The station operated from March 16 - June 16, 2001. The station did not operate properly after mid May 2001 due to a power failure in the amplifier. Seismometers deployed were three component SM3 instruments located in a small pit adjacent to a mine outbuilding. Several blasts and tectonic events were recorded.

# APPENDIX E

The 10 June, 2000, Kadakchan explosion.

The largest explosion to have occurred during the period of the study is believed to have occurred at the Kadakchan coal mine northwest of Susuman on 10 June, 2000. This appendix outlines the evidence used to determine its nature as an anthropogenic event.

This event was well recorded by four of the digital seismic stations deployed as a part of this study, as well as two additional operating analog photopaper stations in the Magadan region (Table E-1). The K-class of the event was 9.5, which corresponds to a local magnitude of 3.3. Given the size of the event, it was routinely located and cataloged as a tectonic earthquake. The event was located using the calibrated crustal velocity model discussed in Section 3 above. Results of the location, including arrival residuals are shown in Table E-2.

The closest operating recording station to the event was Omchak (Susuman was down). Unfortunately, timing at Omchak was not working. In the process of locating the event, data from Omchak was used with a manual time correction applied. The time correction was determined by minimizing the sum of residuals for all Omchak arrivals, which is essentially using Sg-Pg-Pn time differences to fit the data. It should be noted that as the event was originally considered tectonic, there was no attempt to 'force' the event to any particular location, mine or otherwise.

The tectonic nature of the event was not questioned until September, 2001, when it was found that the calculated epicenter of the event falls very close to the Kadakhan mine (Figure E-1). The proximity to the Kadakchan mine, and distance to other tectonic events, flagged this event for a closer look.

Upon closer inspection of the seismogram from Omchak, a large acoustic arrival was noted about seven minutes after the P arrival (Figures E-2 and E-3). Using the calculated origin time of the event, the velocity of the acoustic arrival over the 163 km between the epicenter and the station in Omchak was 328 m/s. The existence of the acoustic arrival is most consistent with an explosion and not a tectonic earthquake.

During the 2001 field season, work was undertaken to acquire better information and waveforms from mines in northeastern Russia. Three confirmed blasts from Kadakchan were recorded at a nearby temporary station, as well as the station in Omchak. Yields of the blasts were provided by the mining company. Figure E-4 compares the Omchak waveforms of the 10 June, 2000, event with a 10,140 kg Kadakchan explosion from 23 March, 2001. The character of the waveform and relative amplitudes correlate quite well for the Pg arrival, especially accounting for the noise on the 2001 trace and the width of one sample. The Pn arrivals are also not inconsistent with each other. For both records, the Sg-Pg time differences are also the same. The correlated waveforms between the two records is consistent with the 10 June, 2000, event being a large mine blast.

Considering the  $M_L = 3.3$  size of the 10 June, 2000, event, any explosion source would have to be quite large. Leith (1994) tabulated large explosions in the Magadan region covering the time period 1989 - 1992. The largest explosion to occur took place at the Kadakchan coal mine, and had a yield of 256 tons. Given the history of large blasts at Kadakchan, an explosion source for the 10 June, 2000, event can not be ruled out. The origin time of the 10 June, 2000 event is 05 hours, 54 minutes UTC, which is the middle of local day and consistent with historic blasting in the region. Therefore, origin time can not rule out a large mine blast source.

All available lines of evidence either directly support a Kadakchan mine explosion origin

Table E-1. Original data from the 10 June, 2000, event.

STA- TION	PHASE TYPE		ARRIVAL TIME M. S.	IOD	AMPLIT IN MICRO N-S E-W	ONS		NOTES
осн	P Pg Sg	05	56 49.9 56 50.9 57 11.7					INFRASONIC ARRIVAL AT: 06 04 40.5 All OCH TIMES REFLECT MANUAL CORRECTION OF + 9.3s
SEY	eP ePg Pgm eS eSg Sgm	05	57 33.5 57 42.3		.07		K= 9.4	
TL-	ePg eSg	05	57 15.0 57 23.8 58 05.9 58 10.0	0.7 0	.37 0.30		K=10.3	
MGD	eP ePg iSg Sgm	05	57 18.5 57 28.5 58 17.0 58 27.0		.21 0.24		K=10.1	
MA2	eP ePg Pgm eSg Sgm		57 25.2 57 38.1 57 38.7 58 30.2 58 38.5	0.3	.04 0.07		K= 9.7	
OMS	ePg eSg Sgm	05	57 39.4 58 32.0 58 42.6		.13 0.11		K= 9.9	

Table E-2. Location parameters from the 10 June, 2000, event. Low residuals indicate a stable location.

Origin Time: 05 56 23.8 UTC

Latitude: 63.082°N Longitude: 146.976°E

Depth: 14 km

Infrasonic velocity to Omchak: 328 m/s

STATION	PHAS.	ARR.	TIME	DIST. (DEG.)	AZIMUTH	RESIDUAL	WEIGHT
OCH	P	5 56	49.9	1.48	163.23	.13	1.0
	PG	5 56	50.9	1.48	163.23	39	1.0
	SG	5 57	11.7	1.48	163.23	.69	.5
SEY	P	5 57	4.7	2.47	91.06	.54	1.0
	PG	5 57	9.8	2.47	91.06	.34	1.0
	SG	5 57	42.0	2.47	91.06	28	.5
TL-S	P	5 57	15.0	3.21	125.17	.25	1.0
	PG	5 57	23.8	3.21	125.17	.57	1.0
	SG	5 58	5.9	3.21	125.17	.12	.5
MGD	P	5 57	18.5	3.54	147.81	78	1.0
	PG	5 57	28.5	3.54	147.81	65	1.0
	SG	5 58	17.0	3.54	147.81	1.04	.5
OMS	PG	5 57	39.4	4.08	94.09	.22	1.0
	SG	5 58	32.0	4.08	94.09	-1.29	.5
MA2	P	5 57	25.2	3.97	150.87	18	1.0
	PG	5 57	37.1	3.97	150.87	.05	1.0
	SG	5 58	29.1	3.97	150.87	52	.5

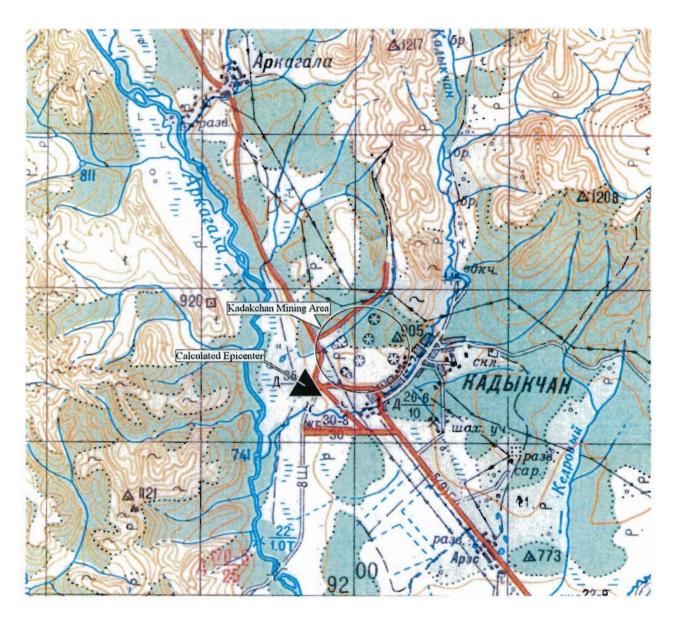
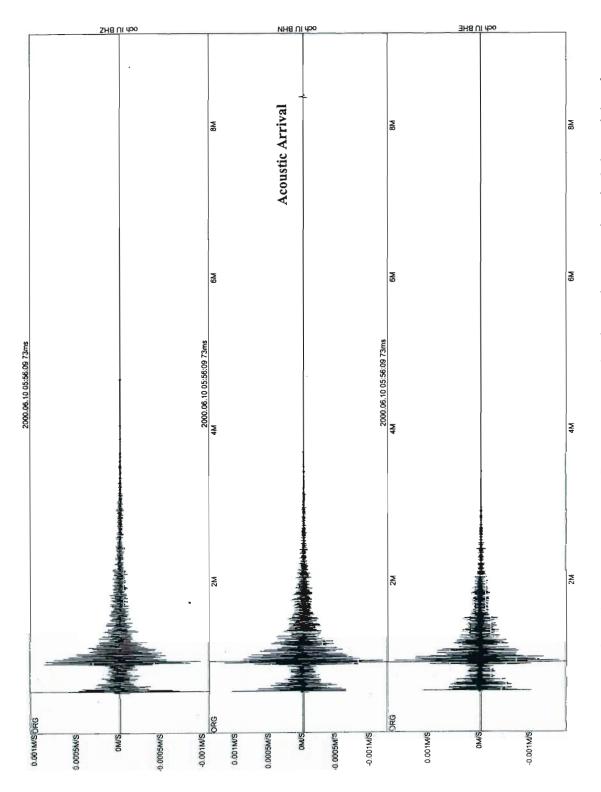
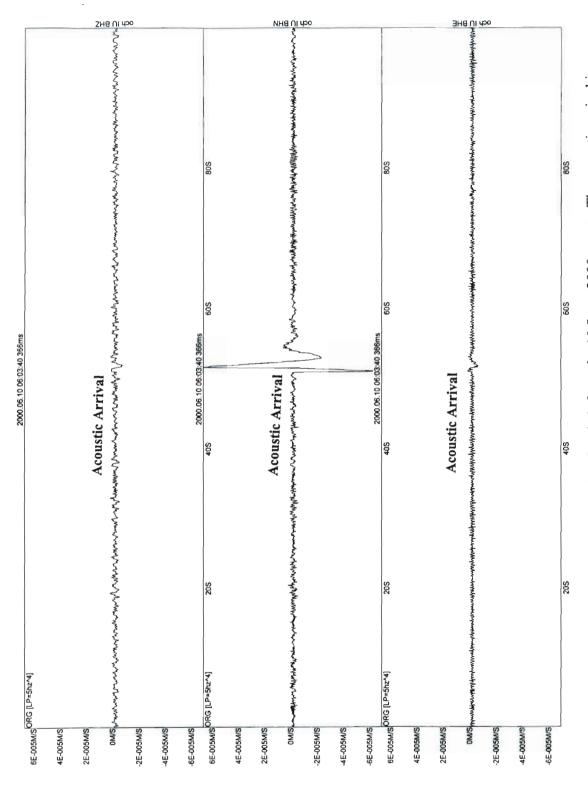


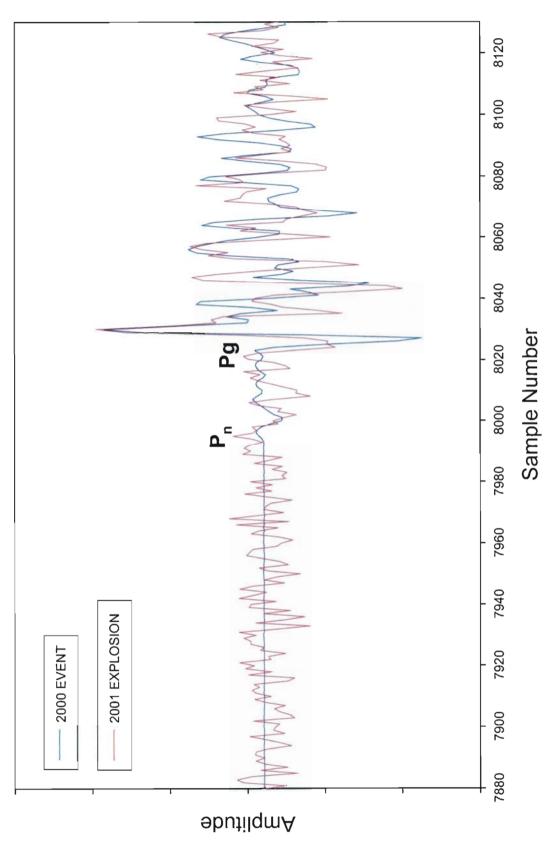
Figure E-1. Kadakchan mining region. The calculated epicenter is shown as the triangle. Note the close proximity of the calculated epicenter to the mapped mine pits enclosed by the circle. It is unknown exactly where the currently active pit is in Kadakchan, as the mine has not been field checked. One square on this map represents approximately 4 km. The base map used here is a 1:200,000 Russian military topographic map.



after the P arrival. The acoustic arrival is only visible on the N-S component at this scale. Figure E-3 enlarges the Figure E-2. Omchak short period seismograms of the 10 June, 2000 event showing a clear acoustic arrival about eight minutes acoustic arrival. Seismogram components are, top to bottom, Z, N-S, and E-W.



strongest on the N-S component, consistent with Omchak being due south of the explosion source at Kadakchan. Figure E-3. Enlargement of the acoustic arrival recorded at Omchak from the 10 June, 2000 event. The acoustic arrival is Seismogram components are, top to bottom, Z, N-S, and E-W.



correlation here suggests that the 2000 event is also an explosion from the Kadakchan coal mine. The yield of the 2001 Figure E-4. Comparison between waveforms of the 10 June, 2000, event and a known Kadakchan explosion from 24 March, 2001. Note that the waveforms correlate well, especially accounting for the noise superimposed on the 2001 (red) trace. The explosion was 10,140 kg of explosives. Amplitude of the 2001 explosion is increased for clarity of correlation.

for the 10 June, 2000, event, or can not rule it out. The size of this explosion and quality of recording at Magadan (MA2) make this event a good candidate for a ground truth event worthy of additional research. Given the limited knowledge of the exact layout of the Kadakchan mine, this event should currently be classified as GT-5. However, a future field check of the Kadakchan mine should improve this to a GT-1. In addition to MA2, this explosion may have been recorded at other globally reported regional stations, such as Yakutsk (YAK).

Figures E-5 through E-8 provide additional digital seismograms from the Magadan region stations Omchak, Seymchan, Talaya, and Magadan.

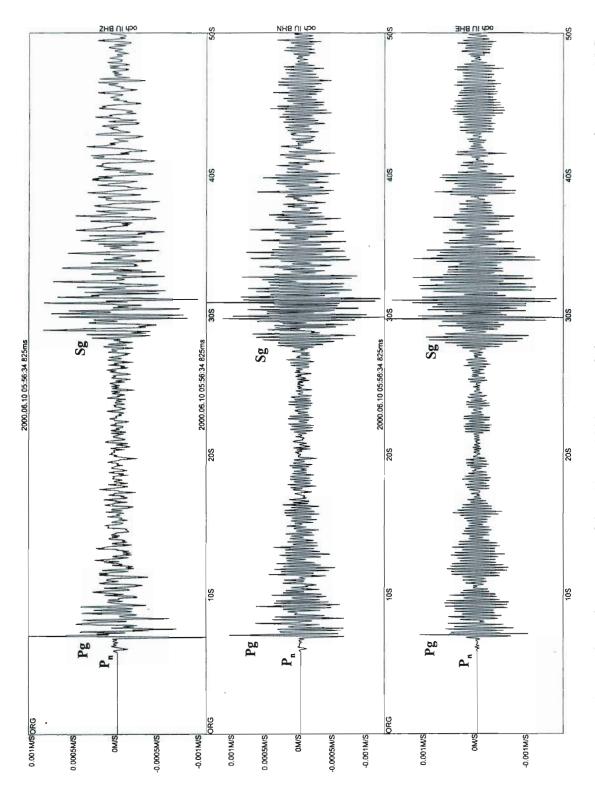


Figure E-5. Omchak short period seismograms of the 10 June, 2000 event. Seismogram components are, top to bottom, Z, N-S, and E-W.

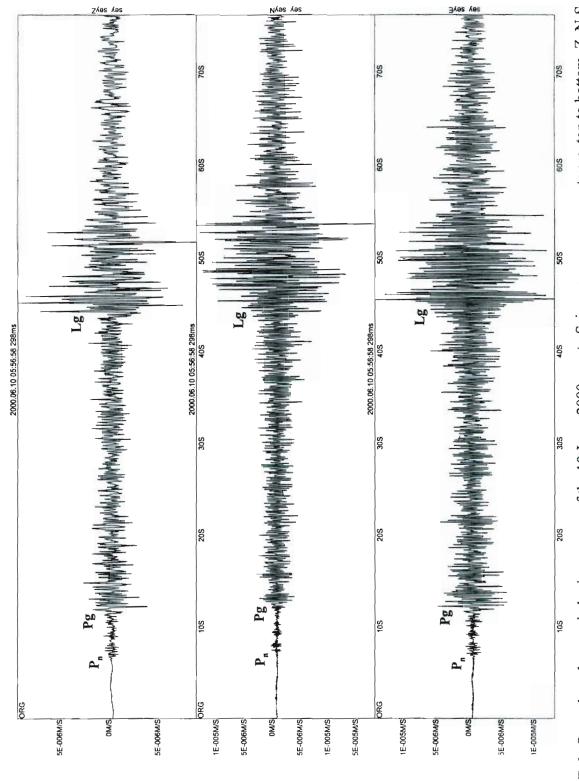


Figure E-6. Seymchan short period seismograms of the 10 June, 2000 event. Seismogram components are, top to bottom, Z, N-S, and E-W.

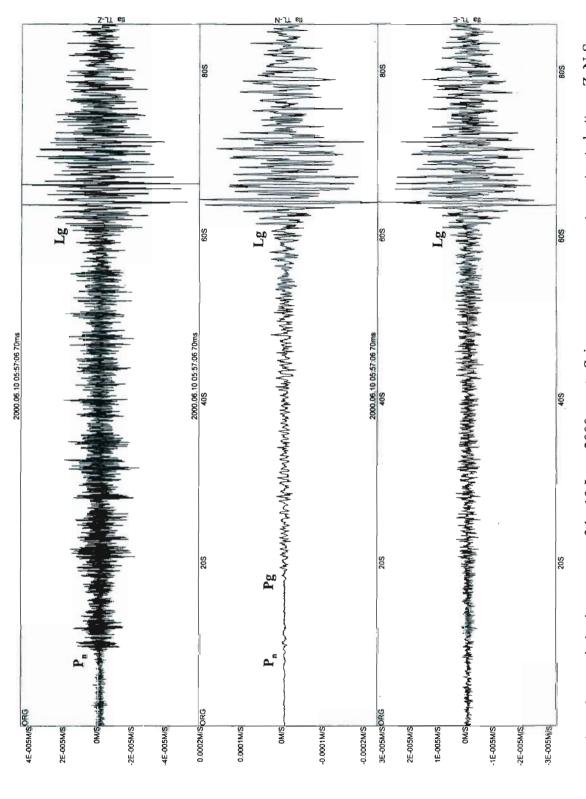


Figure E-7. Talaya short period seismograms of the 10 June, 2000 event. Seismogram components are, top to bottom, Z, N-S, and E-W.

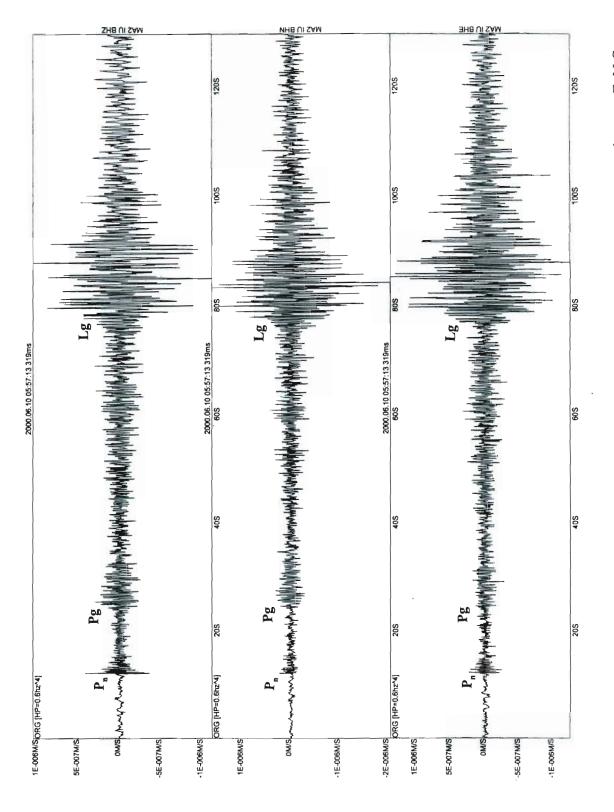


Figure E-8. Magadan broadband seismograms of the 10 June, 2000 event. Seismogram components are, top to bottom, Z, N-S, and E-W. A 0.6 Hz high pass filter was applied to these records.

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