Naval Research Laboratory

Washington, DC 20375-5320



NRL/MR/6110--04-8766

Natural CO₂ Flow from the Loihi Vent: Impact on Microbial Production and Fate of the CO₂

RICHARD B. COFFIN THOMAS J. BOYD

Chemical Dynamics and Diagnostics Branch Chemistry Division

David L. Knies Kenneth S. Grabowski

Surface Modification Branch Materials Science and Technology Division

JOHN W. POHLMAN CLARK S. MITCHELL

Geo-Centers, Inc. Lanham, MD

March 31, 2004

20040428 036

Approved for public release; distribution is unlimited.

----**.**... -

Γ

Form Approved

Т

		CUMENTAT			OMB No. 0704-0188	
maintaining the data need suggestions for reducing t Suite 1204, Arlington, VA information if it does not d	ed, and completing and revie his burden to Department of 22202-4302. Respondents sl isplay a currently valid OMB	ewing this collection of inform Defense, Washington Heado nould be aware that notwiths	ation. Send comments regar- juarters Services. Directorate	ding this burden estimate on for Information Operations I law, no person shall be su M TO THE ABOVE ADDRE		
1. REPORT DATE	(DD-MM-YYYY)	2. REPORT TYPE			ATES COVERED (From - To)	
March 31, 2004 4. TITLE AND SUB	TITIE	Memorandum	Report		FY-03 Status Report 10/01/03-12/31/03 CONTRACT NUMBER	
4. III LE ARD 30B				joa.	CONTRACT NUMBER	
Natural CO ₂ Flow from the Loihi Vent: Impact on Microbia Fate of the CO ₂			al Production and	5b.	GRANT NUMBER	
				5c.	PROGRAM ELEMENT NUMBER	
6. AUTHOR(S)				5d.	PROJECT NUMBER	
Richard B. Coffin, Thomas J. Boyd, David L. Knies,* Kenn John W. Pohlman,† and Clark S. Mitchell†			neth S. Grabowski,*	5e.	TASK NUMBER	
					WORK UNIT NUMBER	
7. PERFORMING C	RGANIZATION NAM	E(S) AND ADDRESS	i(ES)		ERFORMING ORGANIZATION REPORT	
Naval Research I 4555 Overlook A Washington, DC		0			NRL/MR/611004-8766	
9. SPONSORING /	MONITORING AGEN	CY NAME(S) AND A	DDRESS(ES)	10.5	SPONSOR / MONITOR'S ACRONYM(S)	
	Technology Laborato 10 Collins Ferry Roa				SPONSOR / MONITOR'S REPORT NUMBER(S) Project No. DE-A126-99FT40623	
12. DISTRIBUTION / AVAILABILITY STATEMENT Approved for public release; distribution is unlimited.						
 13. SUPPLEMENTARY NOTES *NRL, Surface Modification Branch, Materials Science and Technology Division †Geo-Centers, Inc., Lanham, MD 						
a suite of biogeoc 2000, to obtain a sediment chemist completed, local of Norwegian Coast Control Authority flow off the south of the Internationa Work focused on the cruise on the <i>R/V</i>	hemical parameters o thorough data set ir ries. These data were environment regulatio was conducted durin , for policy reasons r ern coast of the Big Is al Carbon Dioxide Se tracing the venting ga Ka'imikai-O-Kanaloo is final report is from	ff the coast of Kona o acluding measurement collected in order to on forced moving the g 2002. However, No egarding the CO_2 injo- sland. From December questration Experimen- ses, the impacts of the	n the Big Island of Ha tts of pH, current pro interpret a planned C project to the coast n prwegian government ection experiment. As r 3rd-13th 2002, scier ent examined the hydr e vent fluids on marin ives by the <i>PISCES V</i>	waii. The preliminar files, CO_2 concentra CO_2 injection experim ortheast of Bergen, revoked a permit, a a result, the research tists from four coun othermal venting at e organisms, and CC	997. Preliminary steps involved surveying y survey was conducted twice, in 1999 and ations, microbial activities, and water and nent. After these preliminary surveys were Norway. The preliminary survey along the pproved by the Norwegian State Pollution h team decided to monitor the natural CO_2 tries representing the Technical Committee Loihi Seamount (Hawaiian Islands, USA). D_2 influence on biogeochemical cycles. The i, and 2 at a nearby site in the lee of the Big	
Carbon Sequestra	tion; Ocean floor CO	2 plumes; Bacterial ir	npact; Rađio carbon t	racing		
16. SECURITY CLASSIFICATION OF:			17. LIMITATION OF ABSTRACT	18. NUMBER OF PAGES	19a. NAME OF RESPONSIBLE PERSON Richard B. Coffin	
a. REPORT	b. ABSTRACT	c. THIS PAGE	UL	10	19b. TELEPHONE NUMBER (include area	
Unclassified	Unclassified	Unclassified			^{code)} 202-767-0065	

Standard Form 298 (Rev. 8-98) Prescribed by ANSI Std. Z39.18

CONTENTS

.

.

•

-

-

OVERVIEW	1
INTRODUCTION	1
METHODS	2
RESULTS	3
CONCLUSIONS	5
LITERATURE CITED	7

DISCLAIMER

This report was prepared as an account of work sponsored by an agency of the United States Government. Neither the United States Government nor any agency thereof, nor any of their employees, makes any warranty, express or implied, or assumes any legal liability of responsibility for the accuracy, completeness, or usefulness of any information, apparatus product, or process disclosed, or represents that its use would not infringe privately owned rights. Reference herein to any specific commercial product, process, or service by trade name, trademark, manufacturer, or otherwise does not necessarily constitute or imply its endorsement, recommendation, or favoring by the United States Government or any agency thereof. The views and opinions of authors expressed herein do not necessarily state of reflect those of the United States Government or any agency thereof.

Natural CO₂ Flow from the Loihi Vent: Impact on Microbial Production and Fate of the CO₂

OVERVIEW

The program for International Collaboration on CO₂ Ocean Sequestration was initiated December 1997. Preliminary steps involved surveying a suite of biogeochemical parameters off the coast of Kona on the Big Island of Hawaii. The preliminary survey was conducted twice, in 1999 and 2000, to obtain a thorough data set including measurements of pH, current profiles, CO₂ concentrations, microbial activities, and water and sediment chemistries. These data were collected in order to interpret a planned CO_2 injection experiment. After these preliminary surveys were completed, local environment regulation forced moving the project to the coast north east of Bergen, Norway. The preliminary survey along the Norwegian Coast was conducted during 2002. However, Norwegian government revoked a permit, approved by the Norwegian State Pollution Control Authority, for policy reasons regarding the CO₂ injection experiment. As a result the research team decided to monitor the natural CO₂ flow off the southern coast of the Big Island. From December 3rd-13th 2002 scientists from four countries representing the Technical Committee of the International Carbon Dioxide Sequestration Experiment examined the hydrothermal venting at Loihi Seamount (Hawaiian Islands, USA). Work focused on tracing the venting gases, the impacts of the vent fluids on marine organisms, and CO₂ influence on biogeochemical cycles. The cruise on the R/VKa'imikai-O-Kanaloa (KOK) included 8 dives by the PISCES V submarine, 6 at Loihi and 2 at a nearby site in the lee of the Big Island. Data for this final report is from the last 2 dives on Loihi.

INTRODUCTION

World concern about the increase in atmospheric CO_2 has initiated investigation of a broad range of technologies for anthropogenic CO₂ sequestration. Currently, potential locations for CO_2 sequestration are in mines, forests and the ocean. The world's oceans are the largest potential sinks for deposition of anthropogenic CO₂. Work over the last few years has focused on laboratory and modeling efforts for CO₂ sequestration (Handa and Ohsumi 1995, Herzog, 1996; Omerod, 1997). Results from these efforts warrant field studies to validate and ground truth predictive data. The Project Agreement for International Collaboration on CO₂ Ocean Sequestration was initiated December 1997 to foster increased understanding of the ocean as a CO₂ sink. The initiation of the program was accomplished through a union of research efforts of laboratories from three countries; 1) Research Institute of Innovative Technology for the Earth (Japan), 2) the Norwegian Institute for Water Research (Norway); and 3) the Massachusetts Institute of Technology (USA). Recently, Canadian Institute of Ocean Science, the US Naval Research Laboratory and ABB Switzerland have joined this program. Permission to conduct this research off the coast of the Big Island and the western coast of Norway was requested denied. As a result of this regulation work on carbon sequestration was conducted at approximately 1300 m depth on Loihi, a new forming island off the coast of the Big Island. In the center of this region are active hot flows of dissolved CO₂ commonly referred to as dissolved inorganic carbon (DIC). The NRL Code 6114 contributions to this international project were; 1) measuring the impact

Manuscript approved March 12, 2004.

of the DIC on the microbial production, and 2) coupling radiocarbon isotope measurements of the DIC with pH measurements to determine the region(s) most influenced by the venting activity.

METHODS

A. PISCES V and R/V KOK

Ship board activity was on the RVKOK. This vessel was the support ship for the manned research submarine *PISCES V*. Eight dives were conducted during the research cruise. Loihi was visited 6 times and 2 control sites were sampled. The *PISCES V* was equipped with two mechanical arms for sampling. These were used to obtain sediment cores, manipulate Niskin bottles to collect water samples, and set and retrieve experimental cages. The submarine was equipped to navigate and measure salinity and temperature. The dive depths for this project on Loihi were characteristically 1360 m.

B. Radiocarbon Isotope Analysis

Radiocarbon isotope analysis (Δ^{14} C) is calculated as:

$$d^{n}C = \left[\frac{R_{s}}{R_{std}} - 1\right] \times 1000 \quad (\%)$$

where dⁿ is the radio carbon isotope ratio, R is the ¹⁴C/¹²C for radiocarbon, Rs is the sample ratio for the sample and Rstd is the ratio for the standard, which is oxalic acid. Radiocabon isotope analysis is used to determine a carbon source age and the relative input of thermogenic and new carbon sources to a variety of pools. Carbon derived from contemporary processes such as photosynthesis will have an abundant ¹⁴C signature that is termed modern. At the sediment-water interface, mineralization of detrital organic material will introduce contemporary CO₂ into the water column. The Δ^{14} C signature of this CO₂ will be vastly different from thermogenic CO₂ which has no ¹⁴C. This difference allows one to calculate the influence of vent gas to the total DIC sampled at the source and through the plume.

Radiocarbon is formed in the atmosphere when ¹⁴N is altered by cosmic rays. Carbon that is isolated from contemporary biogeochemical cycling (i.e. in long-term sediment burial), will eventually become devoid of ¹⁴C (which has a half life of 5370 years). As a result, Δ^{14} C of new organic matter recently produced in the ocean is greatly elevated relative to thermogenic carbon sources. In the ocean, Δ^{14} C ranges between approximately 100‰ and an undetectable value, near -1000‰. The wide range of dissolved organic carbon Δ^{14} C in the ocean is related not only to ocean-atmosphere coupling, but also to ocean circulation. Open ocean surface water Δ^{14} C range between – 150‰ to -258‰. Deeper ocean water, below 1000 m, characteristically have Δ^{14} C that ranges between -393‰ to -525‰. A somewhat different range occurs for dissolved CO₂, as reported in results of the World Ocean Circulation Experiment (WOCE) for the Pacific Ocean. In that work, the surface ocean waters have a Δ^{14} C range from about -50 to +150‰, while deeper waters below 1000 m have a range from about -160 to -220‰. These ranges still allow considerable resolution between the -1000‰ end member (geothermal CO₂) and typical deep ocean waters (-220‰ -525‰).

C. Bacterial Production

Bacterial productivity in seawater was determined by the leucine incorporation method (Kirchman et al., 1985) as adapted by Smith and Azam (1992). To convert the amount of ³H-Leu incorporated into protein to bacterial organic carbon demand of the assemblage (g C mL⁻¹ h⁻¹), the formula of Simon and Azam (1989) was be used:

$$g C mL^{-1} h^{-1} = \frac{\left(\frac{DPM_{sample}}{hr mL}\right)\left(\frac{mol \ leucine}{Y \ DPM}\right)\left(\frac{2 \ X \ 10^{17} \ Cells \ produced}{mol \ leucine}\right);}{where \ Y = \left(specific \ activity\right)\left(\frac{2.22 \ X \ 10^6 \ DPM}{Ci}\right)}$$

We assumed a 20% metabolic efficiency, which is similar to that measured for low molecular weight organics in other aquatic systems (Bjornsen, 1986).

RESULTS

A. Vent Plume Analysis

During the Loihi dives 7 and 8 water samples were taken with Niskin bottles for measurement of the pH and Δ^{14} C of DIC (Figure 1). Results through this sampling show a strong correlation between the pH and age of the carbon in the DIC pool. The oldest Δ^{14} C values were measured in the samples with lowered pH, taken at the mouth of the vent. The combination of data from the pH and Δ^{14} C in the DIC pool provides a good method to trace the region that is influenced by the thermogenic sources. Future work on this data set will provide a 3-D assessment to determine the volume of water that is impacted by this point source.

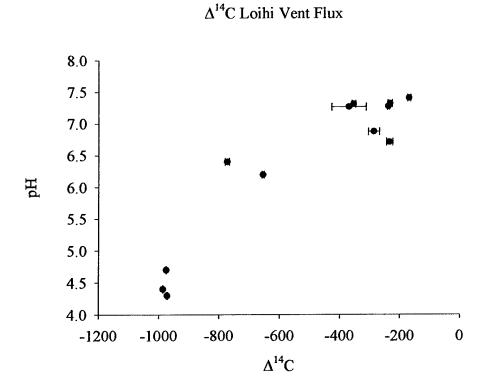


Figure 1: Samples for these data were taken from the mouth of a heat vent on Loihi. This figure compares Δ^{14} C relative to the Ph as the *PICSES V* maneuvered vertically from the vent.

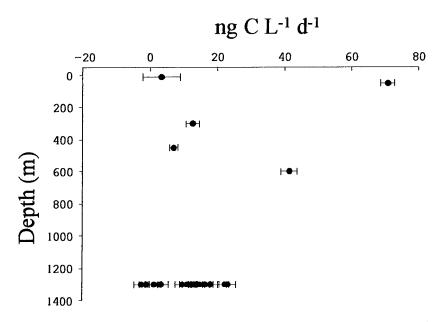


Figure 2: Two vertical profiles of bacterial production and analysis of deep ocean production measured with ³H-leucine assimilation.

B. pH Influence on Microbial Production

During the *PICSES V* dives water column samples were taken vertically to determine the natural range in bacterial production (Figure 2). In the surface to 400m waters there was a large range in bacterial production from 0 to 70 ng C $l^{-1} d^{-1}$. In the 1300 m water samples the bacterial production ranged from 0 to 22 ng C $l^{-1} d^{-1}$.

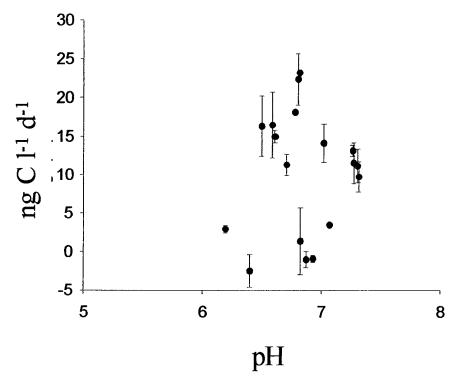


Figure 3: Bacterial production relative to pH for on samples taken as the PICSES V moved vertically from the mouth of the plume.

In the region of the CO_2 vents water samples were taken from the mouth of the plume and through 0.5-1.0 m vertical intervals. These samples were analyzed for pH and the bacterial production (Figure 3). Results indicate a moderate affect of the lower pH on bacterial production. However, similar production rates are observed at the higher pH values of 7.0 to 7.2.

CONCLUSIONS

Application of the Δ^{14} C analysis of the DIC in seawater provides valuable information on two research lines. First, the combination of Δ^{14} C and pH can provide the technology to map CO₂ release in experimental applications and definitively couple the change in pH to a particular source of CO₂. With accurate mapping of region influenced by the CO₂ release, there can be a thorough evaluation for the impact to chemical cycles and biological activity. This analysis also shows the need for re-evaluating the mean ocean carbon age. Typically it is thought that the ocean carbon pool turnover is ~6000 years. If input of dead carbon through hydrothermal venting is significant ocean-wide, then the 6000 year turnover time is overestimated. Over the past couple of decades there has been a vast amount of ocean floor exploration. Many active hydrothermal sites have been discovered inputting thermogenic CO_2 . These sources, i.e. hot vents, cold seeps, mud volcanoes, etc., need to be integrated with the ocean carbon model relative to phytoplankton production and lateral terrestrial transport. Results of such an effort will assist in a more thorough analysis of the ocean carbon cycle and the prediction/modeling of global warming.

In the deep ocean bacterial activity is an important factor in the health of the ecosystem. A part of this community is responsible for remineralization of key elements (such as nutrients) needed by other components of the food web. Also in the deep sea environment chemo-autotrophic bacteria can be a key part of the food chain. In many of these systems there are large amounts of reduced compounds (methane, ammonium, and sulfide) that are oxidized by the chemo-autotrophs and result in carbon dioxide fixation. This provides the first step in the food chain for energy. As a result of the importance of the microbial community on ecosystem health we chose to monitor the impact of pH during the dives on Loihi. Results from the bacterial production do not show a strong relationship with the variation in pH (Figure 3). While there was some indication of inhibited production at the lowest pH levels, similar results were observed at a range in pH of 7.0 - 7.2. This variation may be a factor of the growth phase of the sampled bacteria. Studies have shown that a natural bacterial population transitions from a rapid growth until the maximum carrying capacity (K_m) is met. When the growth rate slows at the K_m, bacteriovore grazing reduces the population and the bacterial growth rate accelerates again. Another consideration in this analysis is that with long term low pH the bacterial community at Loihi has adapted and is not inhibited. This work is a valuable study for understanding the impact of carbon sequestration. Future work needs to combine field and laboratory experiments.

LITERATURE CITED

Bjornsen, P. K. 1986. Bacterioplankton growth yield in continuous seawater cultures. Mar. Ecol. Prog. Ser. 30:191-196.

Kirchman, D. L., K'Nees, E., Hodson, R. 1985. Leucine incorporation and its potential as a measure of protein synthesis by bacteria in natural aquatic systems. Appl. Environ. Microbiol. 49:599-607.

Handa, N. and T. Ohsumi. 1995. Direct ocean disposal of carbon dioxide. (eds.) Terrapub, Tokyo.

Herzog, H. 1996. Proceedings of the 3rd International Conference on Carbon Dioxide Removal, Cambridge, MA, (ed.) September 1996.

Ormerod, B. 1997. "Ocean storage of carbon dioxide" proceedings of four workshops sponsored in 1996 and 1997 by IEA Greenhouse Gas R&D Programme, Cheltenham, UK.

Simon, M., and F. Azam. 1989. Protein content and protein synthesis rates of planktonic marine bacteria. *Mar. Ecol. Prog. Ser.*, 51: 201-213.

Smith, D. C., and F. Azam. 1992. A simple, economical method for measuring bacterial protein synthesis rates in seawater using ³H-leucine. Mar. Microb. Food Webs 6(2):107-114.