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LONG TERM MODELLING OF PERMAFROST DYNAMICS

Final Technical Report

by

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ABSTRACT

The PERIMOVE model is capable of simulating moisture and heat transfers within soil profiles. In addition, the simulation of these processes formed the basis for the prediction of the volume changes that moisture redistribution and freezing create within soil profiles and the resulting movement of material and loss of soil strength during freezing and thawing cycles. In order to aid the development and validation of the model a suite of data sources have been identified and assembled. These have enabled model testing at each stage of development. Areas of further work are suggested.

1 ORIGINAL OBJECTIVES

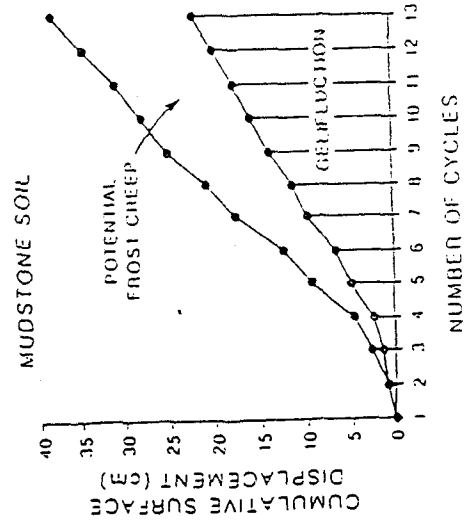
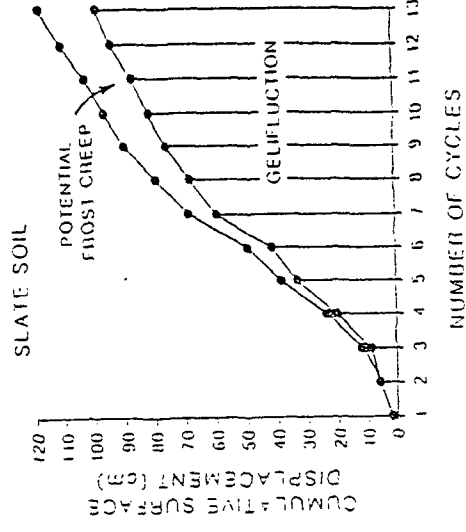
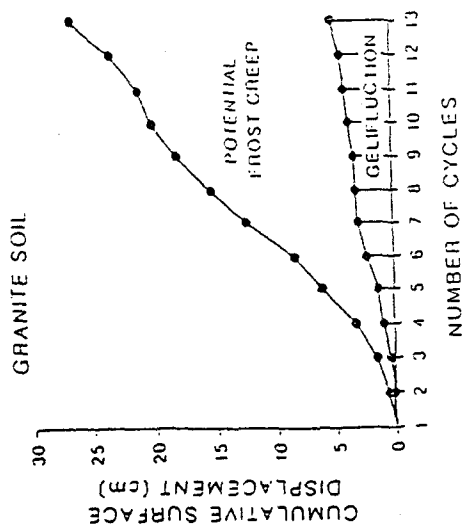
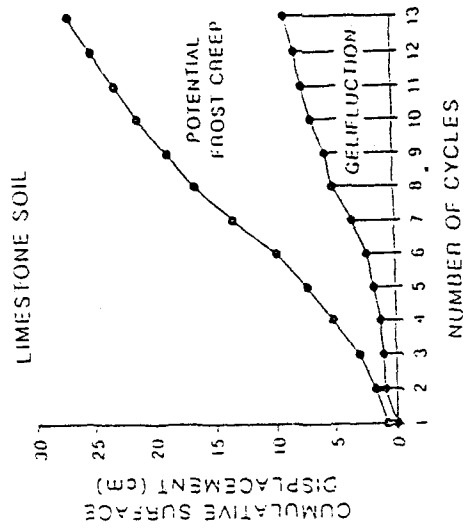
ORIGINAL OBJECTIVE:

To develop a 2-dimensional, physically based model for mass and heat transfer, driving a creep process model for cold regions slopes.

Harris *et al* (1993) describe some results of a NERC funded project involving a large scale laboratory experiment into frost creep and gelifluction processes. These processes can be identified as distinct where creep describes the downslope movement of material due to gravitational resettlement of soil during thawing as a result of heave occurring during a freeze cycle and gelifluction describes the downslope deformation of a saturated mass of soil overlying a frozen soil layer during the thaw cycle. Figure 1 shows their results for four different soil types and the relative importance of creep and gelifluction processes in controlling downslope movement for the different soil types is apparent. Because of the significance of gelifluction in producing downslope movement of material it was decided that the original objective should be altered to include gelifluction processes in addition to creep processes.

The term **solifluction** has replaced the term creep in the original objective as the former term embraces both frost creep and gelifluction movements as defined in Harris (1987). Collaboration of this project with that of Harris has taken place and is described in more detail in later sections.

An additional change in the original objective is that the model should strictly be termed a pseudo 2-dimensional model. Although the basic heat and moisture transfer model could operate in 2-dimensions the increase in complexity resulting from the incorporation of volume change make a fully 2-dimensional scheme beyond the means of this project. Ideally the model would operate within a continuously deforming 2-dimensional mesh structure. Possibilities for the extension to this structure are discussed in section 6. PERIMOVE in its current form however



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Figure 1: Relative importance of frost creep and gelifluaction processes for selected soil types (note variation in vertical scales), Harris et al (1993).

is capable of simulating the 2-dimensional transfers of material which occur over freeze/thaw cycles.

2 DEVELOPMENT OF METHODS AND TECHNIQUES

Three areas of development are outlined below. The first section outlines the structure and equation base of the PERIMOVE model. The second section outlines the data preprocessing scheme and the third section describes the data sources that have been collated to supply data which have contributed to the development and the validation of the model. Figure 2 summarises the development and validation route for the model and also suggests to further areas of development for the model.

i) Description of processes, equations and spatial discretisation approach selected for the PERIMOVE model

Spatial Discretisation

The slope profile is represented as a series of columns of equal width and breadth. Each column is further subdivided to create cells which are then treated as homogeneous units within the model. The computation point is at the centre of each cell and all thermal and hydraulic properties are assigned to that point. These block-centred computation points represent the mathematical nodes for the finite difference mass-energy calculations performed by the model. Mass and energy transfers take place on the basis of the thermal and hydrological gradients which are calculated between cells and phase change processes are governed by the within cell conditions.

A maximum of two hydrologically distinct layers may be identified (as defined by suction-moisture curves and characteristic freezing curves). The boundary between them is determined by specification of the lowest cell in the upper layer so that the rest of the column comprises the second layer.

Each slope section is modelled in isolation therefore three-dimensional attributes such as aspect, convergent and divergent slope forms are not included in the present model.

Process and Equation Base

Three major sets of equations can be identified which describe the three major groups of processes which need to be included in a modelling scheme. These are:

- processes of heat flow in soils
- processes of phase change in freezing and thawing soils

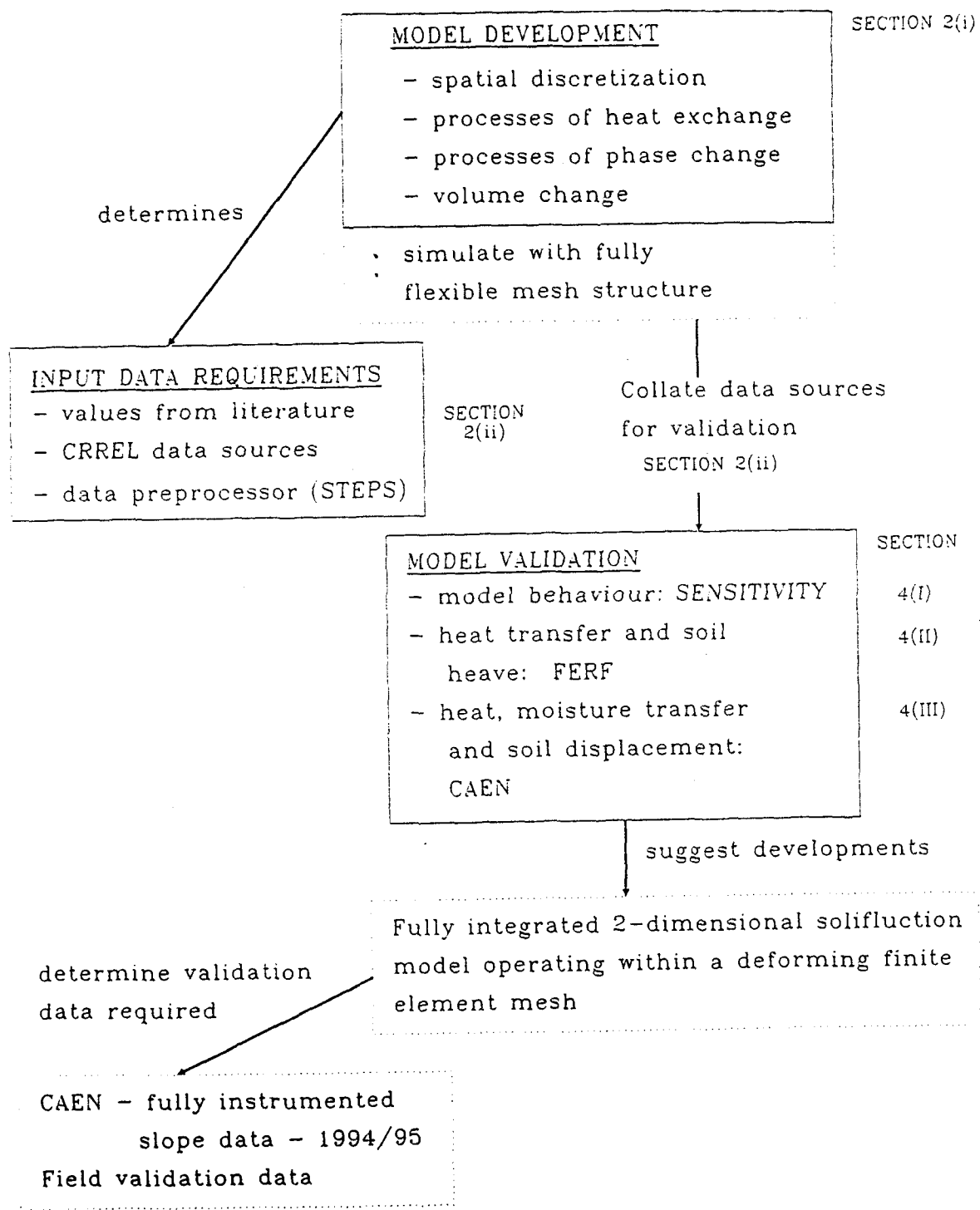


Figure 2: Summary diagram of stages in model development and application. Solid line boxes indicate the scope of this project and dotted lines indicate areas for further research.

processes of water transfer through the profile.

Although these processes all interact it is possible to consider them independently in order to outline the major equations describing each of the processes.

Processes of heat flow within soil

Heat flow within soils takes place by a variety of processes. In periglacial situations the two major processes responsible for heat transfer are conduction through the soil/water/ice medium and transfer of latent heat by liquid flow. Additional processes such as heat flow by radiation, heat generated by mechanical work of the particles and sensible heat convection are generally considered to be of minor importance and are currently excluded from the PERIMOVE model.

Thermal Conductivity is a measure of the quantity of heat that will flow through a unit area of a substance of unit thickness in unit time under a temperature gradient. It is usually measured in $W m^{-1} K^{-1}$.

In situations where ground temperatures fall below $0^{\circ}C$ the thermal conductivity of soils cannot be assumed constant as freezing soils have a temperature dependent ice content and the thermal conductivity of ice is four times that of water. In soils undergoing freeze/thaw the ice/water ratios change rapidly. In periglacial situations where soil temperatures are frequently in the range 0 to $-3^{\circ}C$, failure to account for the temperature dependence of thermal conductivity may lead to *unacceptable errors in the modelling stage* so temperature dependant relationships are incorporated in the PERIMOVE model.

DeVries (1952) developed the equation for calculation of thermal conductivity in a multi-component system:

$$K_t = \frac{\sum_{i=0}^n x_i k_{ti} F_i}{\sum_{i=0}^n x_i F_i}$$

(eq. 1)

K_t = thermal conductivity of the system ($W m^{-1} K^{-1}$)

n = number of different kinds of particles in the continuous media

x_i = volume fraction of i th kind of particle (%)

k_{ti} = thermal conductivity of i th kind of particles ($W m^{-1} K^{-1}$)

F_i = ratio of average temperature in the i th kind of particles to average temperature in the continuous media

this means that for a water, ice and soil system the equation is (by Penner, 1970):

$$Kt = \frac{x_w k_t F_w + x_i k_i F_i + x_s k_s F_s}{x_w F_w + x_i F_i + x_s F_s} \quad (\text{eq. 2})$$

subscripts: w = water, ($F_w = 1$)
 i = ice
 s = soil

PERIMOVE uses the above formula in which both ice and water thermal conductivities are considered temperature dependent below 0°C . The relationships of thermal conductivity with temperature are input to the model as a series of data pairs of conductivity and temperature and interpolation between data points is made to 'look-up' conductivity values for the current temperature of the particular computation cell.

The above equation does not include latent heat transfer effects and as stated above significant errors can occur in predictions if these are omitted. Latent heat is released during freezing and absorbed during thawing and means that a partially frozen soil behaves as if it has a greatly increased heat capacity. Latent heat is considered in the following section.

Once a thermal conductivity has been established it can be used to describe transfer of heat by conduction by application of Fourier's Law:

$$h_x = -Kt_x \frac{\delta T A}{\delta x}$$

where:

h_x = rate of energy flow in x direction (W)
 Kt = thermal conductivity in specified direction ($\text{W m}^{-1} \text{K}^{-1}$)
 T = temperature (K)
 A = area (m^2)

Fourier's law states that the rate of energy flow in the direction of temperature drop is proportional to the area through which the energy flows and the temperature gradient.

Laplace's equation combines the heat conduction theory with the continuity equation to generate a second order partial differential equation:

$$\frac{\delta}{\delta x} \left(\frac{-Kt \delta h}{\delta x} \right) + \frac{\delta}{\delta y} \left(\frac{-Kt \delta h}{\delta y} \right) + \frac{\delta}{\delta z} \left(\frac{-Kt \delta h}{\delta z} \right) = 0 \quad (\text{eq. 3})$$

where: $Kt = Kt(x,y,z)$

If the region considered is homogeneous and isotropic then Kt is independent of x, y and z , so that the above equation simplifies to:

$$\frac{\delta^2 h}{\delta x^2} + \frac{\delta^2 h}{\delta y^2} + \frac{\delta^2 h}{\delta z^2} = 0$$

(eq. 4)

Laplace's equation therefore describes heat conduction under steady state conditions. Fourier's law in combination with the continuity equation is solved in one dimension within PERIMOVE so that temperature change through time is described by:

$$C_v \frac{\delta T}{\delta t} = -\delta \left(-K_t \frac{\delta T}{\delta x} \right)$$

(eq.5)

where: C_v = volumetric heat capacity of the soil medium

and: **volumetric Heat Capacity** is the amount of heat required to change the temperature of a unit volume of a substance by one degree. The heat capacity of a mixture is equal to the sum of the heat capacities of its components.

This heat exchange with the profile is achieved by specification of Neuman boundary conditions in which the transfers of energy across the surfaces bounding the problem are known.

The thermal response of a soil requires the calculation of volumetric heat capacity. In unfrozen soils the volumetric heat capacity is dependent on dry bulk density of soil, moisture content and specific heat capacity of the constituent mineral matter, (Jumikis, 1966). The relationship can be expressed as:

$$C_v = \gamma_d (C_s + C_w W/100)$$

where:

C_v = volumetric heat capacity ($J m^{-3} K^{-1}$)

γ_d = dry bulk density of soil ($g m^{-3}$)

C_s = specific heat capacity of dry mineral matter ($J g^{-1} K^{-1}$)

C_w = specific heat capacity of water ($4.18 J g^{-1} K^{-1}$)

W = water content (%)

As soil freezes the volumetric heat capacity changes as relative volumes of ice and water change. The equation for partially frozen soil is expressed by:

$$C_{vf} = \gamma_{df} (C_s + (C_w W/100) + (C_i I/100))$$

where:

γ_{df} = frozen bulk density

C_i = specific heat capacity of ice ($J g^{-1} K^{-1}$) pure ice = $2.1 J g^{-1} K^{-1}$

I = ice content (%)

This assumes that the air phase can be ignored and the specific heats are assumed to be independent of temperature. Bi-modal values of soil bulk density where one value is used for frozen conditions and another for unfrozen conditions are employed in PERIMOVE. This represents a simplification of reality as soil structure changes during repeated freezing and thawing events producing corresponding changes in bulk density values. It may well be necessary to incorporate this continual change in bulk density at a later stage especially if the model is particularly sensitive to this.

Processes of Phase Change in Freezing and Thawing Soils

Phase change of water to ice and ice to water takes place over a range of temperatures in soils. As an example, fine-grained silt/clay soils can still contain a liquid moisture content at temperatures as low as -20°C . Therefore it is not sufficient to consider the freezing process occurring at a single interface (i.e. the 0°C isotherm). Phase changes in cooling and warming soils are determined by heat flux and water content of the cell.

Phase change in freezing soils

The heat budget for a uniform volume of soil can be described by:

$$\Delta Q = \frac{L_a \delta \theta}{\delta T} \Delta T - C_v \delta \Delta T \quad (\text{eq. 6})$$

where:

$L_a \frac{\delta \theta}{\delta T} \Delta T$ is the energy released due to phase change of moisture to ice, termed the **hydrological component**.

$C_v \delta \Delta T$ is the energy release due to thermal capacity of soil medium, termed the **thermal component**.

ΔQ = change in energy of the system

L_a = latent heat of fusion (J g^{-1})

θ = volumetric soil moisture content (%)

C_v = volumetric heat capacity ($\text{J m}^{-3} \text{K}^{-1}$)

V = volume (of cell) (m^{-3})

T = temperature (K)

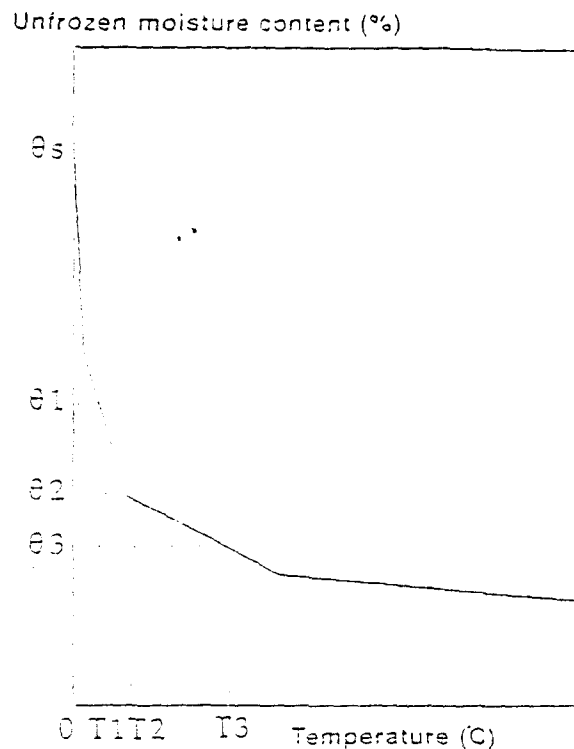
Use is made of the characteristic freezing curve for a soil which is an experimentally determined curve which describes the percentage of liquid water which remains unfrozen at a given temperature below 0°C

The characteristic freezing curve for the specified soil, the soil liquid water content and the soil temperature are used to determine the quantity of ice formed for a given drop in temperature.

The operation of the curve in the PERIMOVE model is described in figure 3.

Characteristic Freezing Curve

The characteristic freezing curve for a soil defines the relationship between liquid water content and temperatures below 273.15K (0°C). The curve can be determined experimentally and an example of a characteristic freezing curve is given below.



If the soil is initially saturated (θ_s) then as the soil temperature begins to drop below 0°C the liquid water content will fall such that at temperature T_1 the liquid water content will be θ_1 and at T_3 the corresponding liquid water content will be θ_3 . The fall in liquid water content with temperature is equivalent to the water equivalent volume of ice formed.

If, however, the soil is unsaturated at the start of the freezing process, for example θ_2 then ice formation will not be initiated until the soil temperature falls below T_2 (the temperature corresponding to a liquid water content of θ_2). Until this point is reached the liquid water content will be super-cooled. Once temperature T_2 is reached the falling liquid water content will follow the pattern of the characteristic curve.

This curve is used in subroutine FREEZE to determine a liquid water content given the soil temperature (as calculated in HEATRANS).

Figure 3: Operation of the characteristic freezing curve within the PERIMOVE model.

On freezing three possible thermo-hydrodynamic conditions are considered.

$$1) \text{ La } \frac{\delta\theta}{\delta T} \Delta T > C_v \Delta T \quad \therefore \Delta Q \text{ is positive and isothermal phase change occurs}$$

The energy generated from the phase change of water into ice is greater than the temperature decline controlled by ambient temperature change, heat conduction through the soil and heat capacity of the soil (thermal component). The soil temperature remains unchanged as the temperature decline due to soil heat capacity and temperature gradient is compensated for by energy release from phase change. Excess ΔQ is assumed to be lost from the system.

$$2) \text{ La } \frac{\delta\theta}{\delta T} \Delta T = C_v \Delta T \quad \therefore \Delta Q \text{ is zero}$$

The energy losses due to the thermal component are equivalent to the energy gain due to water phase change. The soil temperature remains constant as the gains/losses are compensatory.

$$3) \text{ La } \frac{\delta\theta}{\delta T} \Delta T < C_v \Delta T \quad \therefore \Delta Q \text{ is negative, non-isothermal phase change occurs}$$

The energy loss from the thermal component is greater than can be compensated for by release of energy which occurs as the soil cools below freezing. There is a drop in soil temperature as phase change cannot release sufficient energy. ΔQ is negative and no energy is lost from the soil system as all hydrological latent heat is used in compensating for thermal losses.

Thawing Soils:

The logic of phase change for thawing soils is based on the assumption that energy input to a cell is partitioned between:

- (1) temperature change - the **thermal component**
- (2) phase change of ice to water - the **hydrological component**

A thaw temperature is specified (normally 273.15K) and all energy input into a cell is used to raise the cell temperature until the thaw temperature is reached. Once the thaw temperature is reached the energy transferred into a cell is used to satisfy the energy requirements of the phase change from ice to liquid water. Once the cell is ice free the incoming energy is again used to raise the cell temperature. These can be expressed as three conditions.

- 1) if the cell temperature is less than the set thaw temperature then all the energy entering the cell is used to raise the cell temperature.

$$\Delta T = \Delta Q / C_v \delta \Delta T$$

2) if cell temperature is at thaw temperature and ice is present then the energy entering the cell is used to satisfy the latent heat requirement of melting ice. The incoming energy determines the volume of ice which can be melted. The temperature remains at the thaw temperature until all the cell ice is melted.

$$\Delta\theta = \frac{\Delta Q / L_a \delta\theta}{\delta T} \cdot \Delta T$$

3) if the cell temperature is at or greater than thaw temperature and there is no ice present all the incoming energy is used to raise the cell temperature.

$$\Delta T = \Delta Q / C_v \delta \Delta T$$

Processes of Water Transfer Through the Profile

In the current version of PERIMOVE only vertical movement of water is incorporated.

Darcy's law provides the basic equation for water flow through soils and for vertical movement the equation can be written as:

$$q_y = -K \frac{\delta H}{\delta y} = -K \frac{\delta (H_p - y)}{\delta y} \quad (\text{eq. 7})$$

where:

- q_y = water flux
- H = total hydraulic head
- H_p = pressure head (m)
- y = vertical distance component (in PERIMOVE scheme)
- K = hydraulic conductivity (m s^{-1})

In unsaturated soils H_p is negative and can be expressed as a suction head:

$$q_y = K \frac{\delta \psi}{\delta y} + K$$

Combining the above equation with the continuity equation yields:

$$\frac{\delta h}{\delta t} = \frac{\delta}{\delta y} \left(K \frac{\delta H}{\delta y} \right) = -\frac{\delta}{\delta y} \left(K \frac{\delta \psi}{\delta y} \right) - \frac{\delta K}{\delta y}$$

(eq. 8)

As soil moisture content and suction head are uniquely related the left hand side of the equation can be written as:

$$\frac{\delta\theta}{\delta t} = \frac{\delta\theta}{\delta\psi} \cdot \frac{\delta\psi}{\delta t}$$

Substituting this relationship into the above equation produces:

$$C \frac{\delta\psi}{\delta t} = \frac{\delta}{\delta y} \left(K \frac{\delta\psi}{\delta y} \right) + \frac{\delta K}{\delta y} \quad (\text{eq. 9})$$

where: C is the specific water capacity describing the change in water content of a unit volume of soil per unit change of matric suction.

Equations 8 and 9 are forms of the Richards equation. Conventionally these are developed for flow in isothermal saturated and unsaturated flow within ice free soil water systems.

Extension of this equation to situations where ice formation occurs within the soil is possible because of the similarity between air/water systems and ice/water systems. As temperatures fall to the point of ice nucleation a single crystal forms in the centre of a water filled pore. If this crystal is regarded as analogous to an air bubble then the flow solution can be assumed similar to that for unsaturated flow. The growth of an ice crystal can be compared to the growth of an air bubble in unsaturated soils and relationships developed for drying soils can be used to describe the hydraulic properties of freezing soils (Burt and Williams, 1976).

As ice begins to form in the upper soil layers the liquid moisture content of these layers falls. This results in significant soil water suctions developing in these cells and allows for water to be drawn up into the freezing zone and fuel the development of ice quantities which can cause profile expansion. Soil conductivity, moisture availability and rate of freezing control the extent to which water will be drawn up the profile.

The hydraulic conductivity curve with moisture content is derived using the Millington-Quirk procedure within the PERIMOVE model.

Volume change and mass movement in the PERIMOVE model

The incorporation of volume expansion and the movement of water through the profile is essential if volume expansion during freezing and subsequent supersaturation of the soil profile and loss of soil strength following thawing are to be modelled. It is known that volume

expansion of soil undergoing freezing is not due solely to the volume change that occurs when liquid water becomes ice within the soil pore system. Continued migration of water towards the freezing front as liquid water is converted to ice results in the formation of segregation ice lenses and additional expansion. On thawing this excess water is often prevented from draining away due to the frozen state of lower profile layers. The supersaturation of the soil material that results creates conditions which can lead to solifluction because of the reduction in soil strength.

For inclusion in the PERIMOVE model it is assumed that the change from liquid water to ice is accompanied by a 9% increase in volume. Within each cell the ice formation proceeds until the sum of the liquid water and ice contents exceeds the porosity of the soil. When this situation is reached the new volume of the cell and the change in cell depth (all expansion is assumed to take place in the vertical dimension) is calculated. Following this the new volumetric percentages of soil particles, liquid water and ice are calculated.

The new values are then transferred to the mass transfer calculation section of the program and transfers are made based on the hydraulic gradients between cells.

Previously the movement of water between cells was limited by the cell porosity as water could not move into a cell which had been 'filled'. This check has been removed to allow continued upward migration of water to fuel freezing and volume expansion.

The volume expansion of a cell also effects the operation of the suction-moisture curve and the characteristic freezing curves.

Characteristic freezing curve

In normal operation the characteristic freezing curve is used to determine the liquid water content of a cell at a given temperature and determine if water is available in the cell for the freezing process to proceed. With volume expansion the assumption is made that the liquid water at a given temperature is an absolute volume calculated on the basis of the original cell volume. When volume change occurs the characteristic freezing curve is used to look up the water content for the given temperature and a value for volumetric water content value is returned. In the program the value is then transformed to a value in m^3 for the freezing calculations. If the volume change is not accounted for in this calculation then the volume in m^3 will be overestimated. To avoid this the value of liquid water obtained from the characteristic freezing curve is divided by the total volume change that has occurred over the simulation.

Suction moisture curve

The suction moisture curve is used to 'look-up' values of m suction for a given soil moisture content. The curve is derived for liquid moisture contents so as ice forms in a cell liquid moisture contents fall and suctions increase.

As volume expansion occurs the same m^3 volume of water will be equivalent to a reduced volumetric water content and return a higher suction value from the suction-moisture curve. However, this interpretation may be satisfactory as an equivalent absolute volume in a larger cell volume will lead to a reduced connectivity in pore space water and thus a higher suction. This is the approach that is being taken in the current version of the PERIMOVE model.

Conductivity

The Millington-Quirk relationship is used to derive a conductivity-moisture curve from the suction-moisture curve. The conductivity-moisture curve is then used to 'look-up' the value of conductivity for a given volumetric moisture content in the mass transfer section of the program. If volume expansion has occurred the volumetric liquid water content will be lower for a given absolute volume of water and a lower conductivity will be returned than for the same liquid volume pre-expansion. This reflects the decrease in connectivity that will result when a cell expands so the values extracted from the curve are used directly in the model.

Summary

The above sections have described the model structure and equations. Figure 4 illustrates the sequence in which the processes described above are modelled within each time step of the PERIMOVE model.

The other factors to be considered in this section are the data sources that have been assembled to help construct and validate the model.

ii) Soil type driven data preparation scheme

An important part of the model development plan was that the model would successfully reproduce the differences in moisture redistribution and freezing of different soil types. To facilitate the derivation of data files for this the Soil Type Essential Property Selector (STEPS) program was written. This is based on the model written by Cochrane and Anderson (1988) and generates soil property files when given a soil class from the standard USDA soil triangle. The original model and database have been expanded to include selection of soil thermal properties and characteristic freezing curves. This provides a quick and user friendly way of selecting soil properties for the PERIMOVE model and introduces the option of wide applicability without elaborate testing of soils.

iii) Validation data sources

Two major data sources have been used in the model development and validation. They are derived from two areas of collaboration that have been established during the project.

A) Collaboration with the Cold Regions Research and Engineering Laboratory in Hanover, New Hampshire. A range of experiments conducted there have been useful in development and testing for the PERIMOVE model.

CRREL has conducted many laboratory tests to derive characteristic freezing curves. In a visit made there the various sample results were collated and soil type information matched with the relevant curves to augment the data preprocessing program STEPS described above.

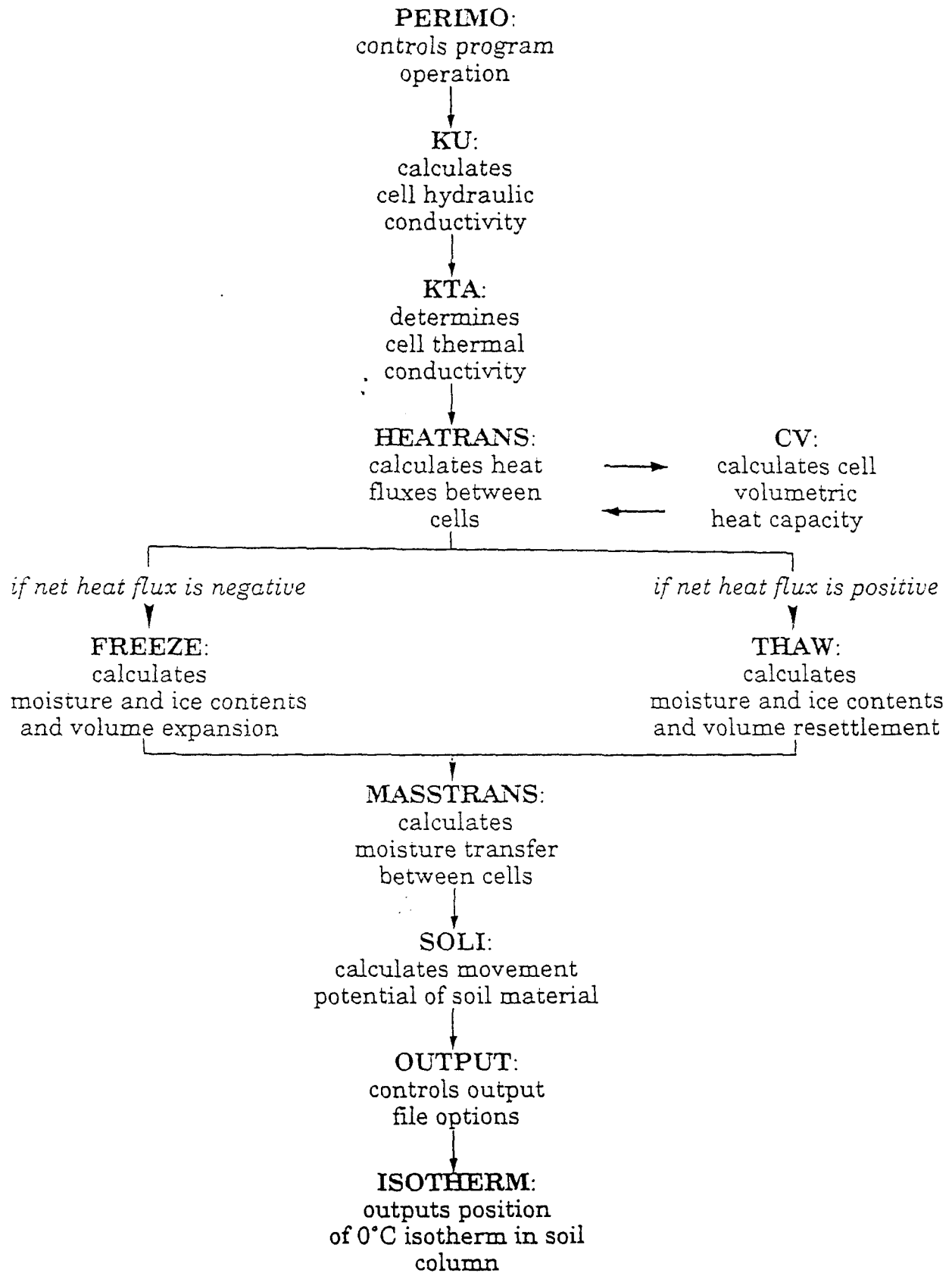


Figure 4: Sequence of operation of the subroutines within the PERIMOVE model.

Shoop (1989) describes the results of an experiment performed within the Frost Effects Research Facility (FERF) at CRREL. The FERF building is an environmentally controlled building and contains a series of test cells measuring around 20ft square and 12 ft deep. Soil freezing is instigated by placing freezing panels on the soil surface and thawing is produced by either heating the building or opening the building to the ambient air temperature. The data which was made available was collected during a freeze/thaw cycle for a silty sand. Profile temperatures were monitored and measurement of total soil heave made at the end of the freeze cycle. Moisture content was measured using tensiometers (although these cease working once freezing commences) and gravimetric moisture determinations were made on samples of soil profile removed at the end of the freeze cycle. As such the data provides a good source for validating temperature and moisture movement within the profile.

An additional data source that was used more in the early development of the model was data collected from some plot experiments conducted at CRREL. Two outdoor plots - one containing a sandy soil type and the other a loam soil type were monitored over two winter seasons. The data consists predominantly of temperature data and this was used in the early development of the model before other data sources were available. The second season has some moisture data and may be used at a later date.

B) The second source of data has come from collaboration with Harris *et al* (1993) who conducted a large scale experiment using a constructed slope within a refrigerated container (some results of which are described in section 1) using different soil types at the C.N.R.S. Centre de Géomorphologie in Caen, France. The results from this study include temperature profiles and heave measurements for a suite of freeze/thaw cycles. The heave measurements can be used to indicate the moisture redistribution occurring during freezing within different soil types and the temperature profiles to compare with model simulated temperatures. In addition measurements were made of the relative amounts of downslope movement of material that could be attributed to creep and gelifluction processes for each of the soils types. This gives an extremely good data set against which soil strength, soil supersaturation on thawing and soil movement can be calibrated within the model.

3 PROGRESS IN RELATION TO SCHEDULE OF WORK

At the current stage of development the model can be used to investigate the nature and extent of 2-dimensional movement of material downslope although as stated in section (1) the current model is pseudo 2-dimensional in structure because of the intense computational difficulties in defining a 2-dimensional deforming mesh structure. The continued development of finite element techniques will mean that it should be possible in the future to have a fully integrated 2-dimensional simulation capacity.

The collation of the range of data sources for model validation and testing has been an important

part of the project and a larger part than originally anticipated however, the collection and use of these data sources has meant the model development and validation has been more rigorous than originally planned.

4 SUMMARY OF RESULTS AND DATA TO DATE

The results summary is presented in three sections and each section highlights a particular area of results from the model. The results follow a progression through the validation and application development of the model. The first section looks at the general performance and behaviour of the model by summarising the results of the sensitivity analysis of model parameter values. The second and third sections concentrate on the application of the model to the two test data sources looking first at the application to the FERF building study which concentrates mostly on heat movement within the profile and total column heave. The extension of the application of the model to consider heave and potential for downslope movement is looked at with the application of the model to the CAEN data set.

A variety of means of handling the data output from the model have been investigated and the results presented here are created using cricket graph and UNIRAS.

1) Sensitivity Analysis

A sensitivity analysis provides an opportunity to discover the most significant parameters in determining model results and also to investigate if the model is responding to changes in parameter values in the manner that one would expect.

To perform the sensitivity analysis a test profile was created which was 50cm deep and input files were used for a silty loam soil type. A short freeze/thaw cycle was used as input to allow for a short simulation duration and permit many runs to be made. With the sensitivity analysis made here only one parameter was altered at any one time. This is very useful in giving a general idea of how the model is operating under different circumstances but means that interrelationships between parameter values which inevitably exist are not considered. An analysis of the combinations of timestep and cell size were made. As cell size is increased timestep can be increased. The time-step and cell size limits are also influenced by the hydraulic conductivities set as these have a strong influence in controlling the mass transfers which will cause instability if they are too great within any time-step length.

The functioning of the hydrological component of the model is of great importance as this is highly significant in determining water migration and therefore the major influence on the occurrence and extent of creep and gelifluction processes. If volume change in a profile occurred only because of the in situ change of water to ice then only a small quantity of heave would result. In many cases the freezing of water in the upper soil layers reduces the liquid water

content of the soil and creates a hydraulic gradient sufficient to pull water up through the profile. This process continues until the pore connectivity reduces so that the transfer rates are so slow that movement is effectively halted. In certain silts ice contents can lead to cell volumes doubling.

As expected the value of the hydraulic conductivity was important. The value for saturated conductivity is supplied to the model and this is then used to create the conductivity with liquid moisture content curve within the PERIMOVE program. An increase in hydraulic conductivity reduces the depth to which the freezing front penetrates and increases the total heave and the excess ice contents in the upper soil layers. For the test profile the conductivity range of $1 \times 10^{-6} \text{ ms}^{-1}$ to $1 \times 10^{-8} \text{ ms}^{-1}$ produced the corresponding range in heave values of 33.8mm to 1.2mm. The slowing of frost penetration with increasing conductivity occurs as a result of the increased volumes of water in the upper cells requiring a greater time period to freeze. Figure 5 illustrates the sensitivity of output to variations in hydraulic conductivity.

The value of saturated moisture content was varied. The major effect of this was in the soil heave values. As the saturated moisture content is used to calculate soil porosity the effect of increasing the saturated moisture content whilst holding the initial profile moisture contents fixed is to reduce heave as greater volumetric change in cell moisture content is required before heave can commence. As the derivation of the conductivity-moisture curve utilizes the value of saturated moisture content the conductivities for corresponding moisture contents are reduced and rate of water movement is slowed.

Bulk densities (both frozen and dry) were altered although the impact of altering these did not result in a straightforward relationship. The most significant impact was to increase the depth of frost penetration with increasing bulk densities. Therefore the impact was mostly in the thermal energy transfer component.

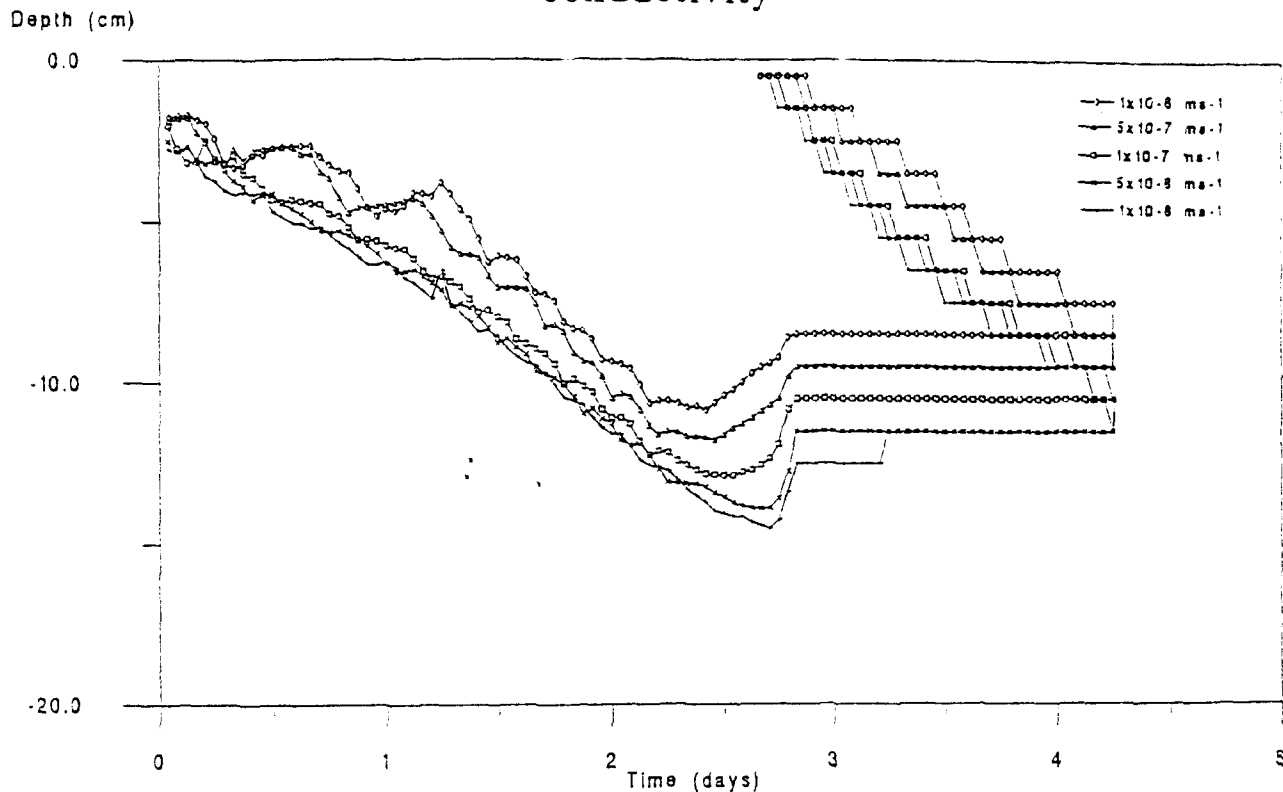
The soil particle conductivity is very significant impact on the thermal component of the model although not greatly influencing the hydrological response. Increasing the thermal conductivity increases the rate and thereby the depth of thaw penetration. Soil heave is reduced a little which probably reflects the slight reduction in water movement caused by the profile freezing earlier and therefore water migration routes being severed sooner.

The sensitivity analysis has identified the implications of parameter changes to model operation and has also providing a guide to the temporal and spatial time-step limitations of the model. This knowledge can be used to aid the application of the model to laboratory data and test the extend the model validation to practical examples.

II) FERF Building Simulations

The FERF building simulations concentrated on the simulation of profile temperatures and total

Sensitivity of 0°C isotherm to hydraulic conductivity



Sensitivity of soil heave to hydraulic conductivity

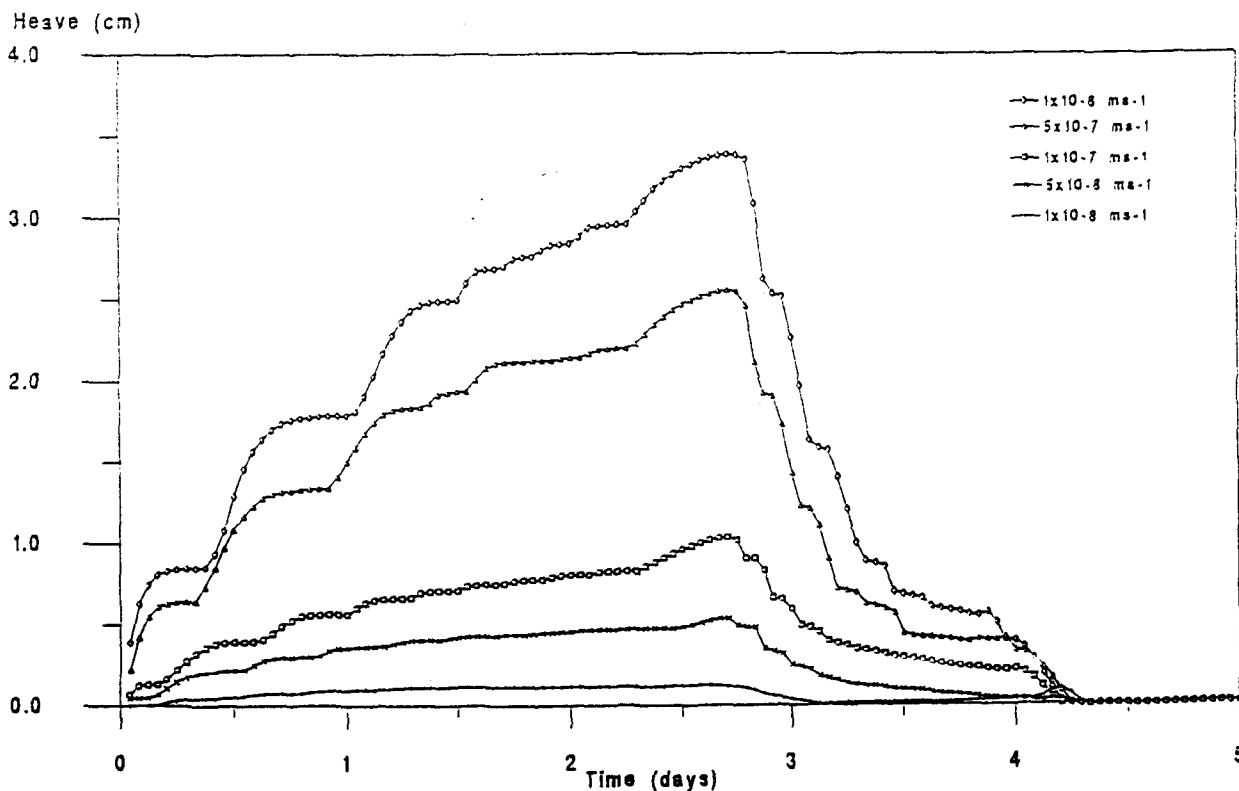


Figure 5: Sensitivity of 0°C isotherm and soil heave to changes in hydraulic conductivity parameter values in the PERIMOVE model.

soil heave values. For the simulation a test profile was created with 1cm deep cells in the upper 10cm of the profile and 2cm deep cells in the remaining 80cm to create a total depth of 90cm. An example simulation is shown in figure 6. The parameter used to indicate the depth of the frost/thaw front is the 0°C isotherm. With the model this is the temperature at which ice can begin to form but, as is also true for the observed situation, does not always indicate that ice forms a significant volumetric portion of the cell.

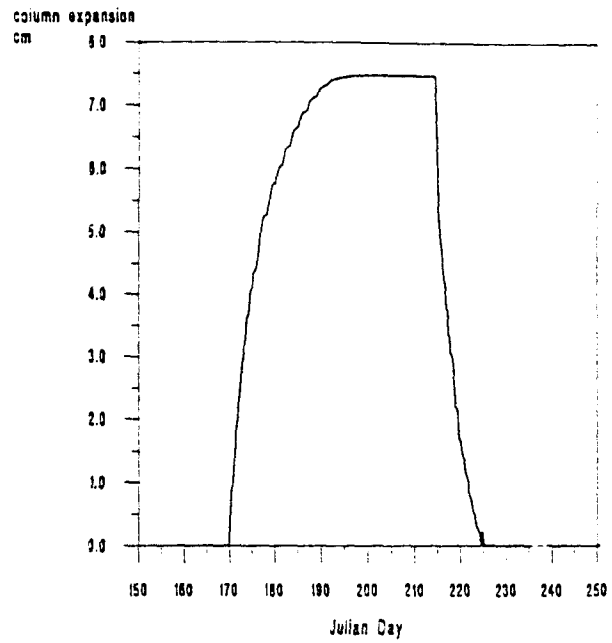
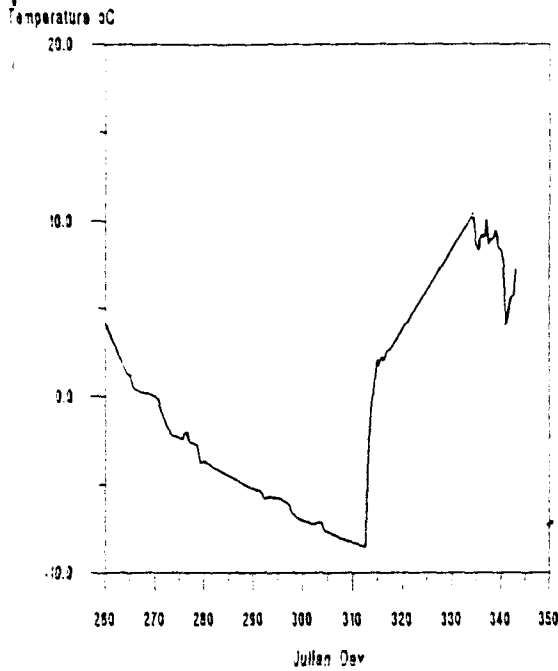
The soil heave predicted by the model is within the range observed for the study which was from 2.5cm to 11cm and the heave extent can be used as an indicator of the migration of water up through the profile to create the observed ice contents and the excess total water contents. The simulation shown uses the data files established for a silty loam soil type. A certain amount of model calibration was made by adjusting input parameter values (mostly thermal conductivities) to achieve the simulation shown and the basal temperature boundary condition was also used as an additional factor in the calibration. The depth of the profile used in the FERF building study was 1m with an underlying layer of gravelly sand of a further 2m and so the lower zone of the profile would have acted as a heat store to counter frost penetration.

III) CAEN Laboratory Simulations

The CAEN study provides additional sources of data for model development and validation. The laboratory investigation looked at the response to freezing and thawing of four soil types. For this investigation two of those soil types were selected to see if the model could successfully reproduce the differences between the soil types. The physical properties of the two soil types selected are summarised in table 1. The data was collected over a suite of freezing and thawing cycles and one of the cycles was selected for the modelling program. The ambient temperature conditions and the observed position of the 0°C isotherm are shown on figure 7. It appears that the sandy soil freezes slower than the silty soil which is not what one would expect however this may have been due to the production of large quantities of needle ice in the upper few centimetres of the profile which greatly slowed the initial frost penetration. The thaw penetrates the sandy soil faster which is as one would expect. Plots of the observed and simulated depths of the 0°C isotherm, soil heave and estimates of downslope movement of material are shown in figure 8 for the silt soil types. Figure 9 shows an alternative method of displaying the temperature profiles over time using a UNIRAS program. Water, ice and other model output can also be displayed in this way.

Because the CAEN experiment used a series of freeze/thaw cycles it will be possible to model alternative cycles with the parameter sets derived for the cycle modelled (cycle 7) to give an indication of how much calibration of input data is required and if data sets that reproduce the observed patterns over a range of cycles can be determined. In addition further work will look at the differences between soil type responses.

Heave values are calculated assuming soil expansion due to freezing occurs orthogonal to the slope and resettlement values by assuming gravitational resettlement of the heaved material.



FERF - simulated and observed frost/thaw depths

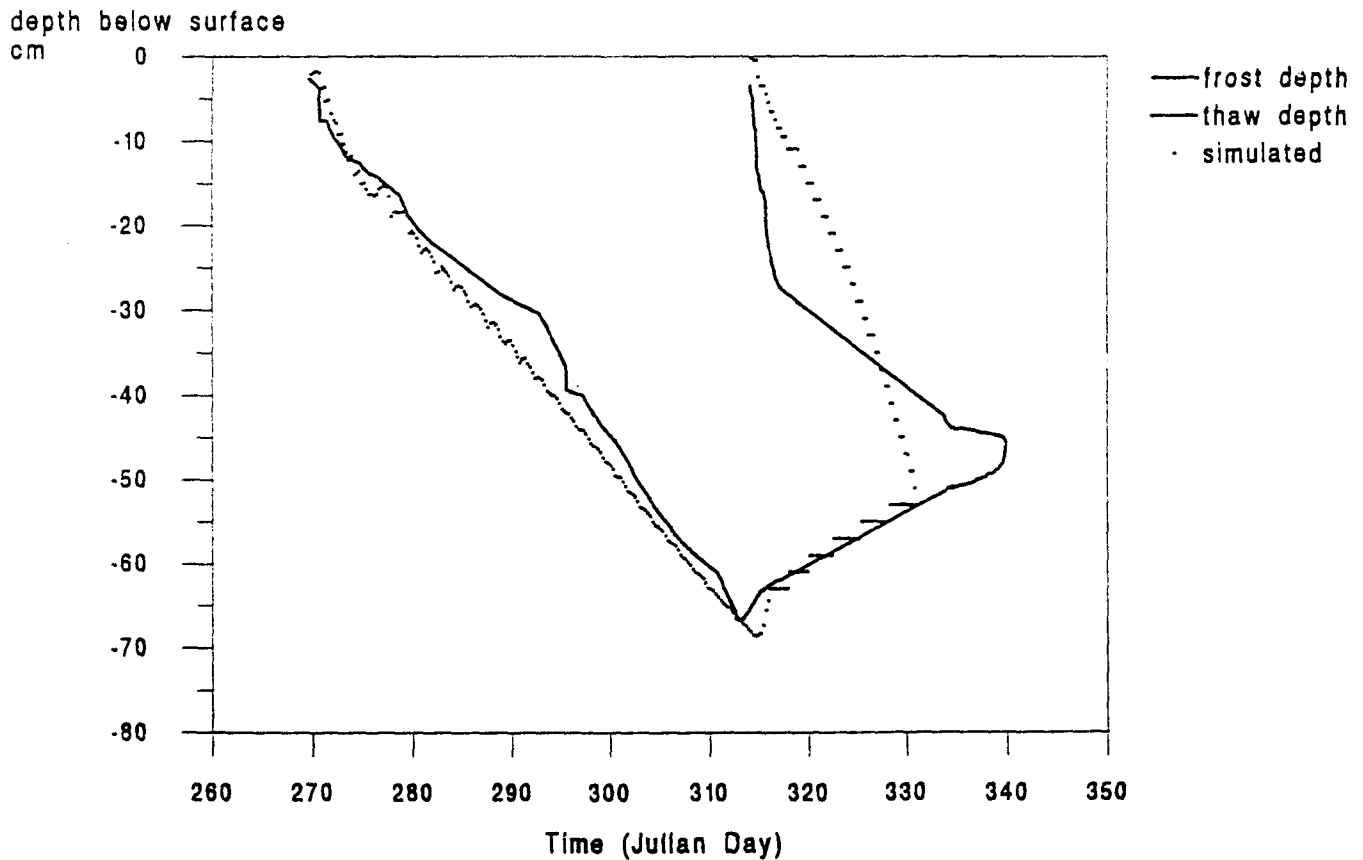


Figure 6: Example simulation of FERG building experiment using the PERIMOVE model.

	%Clay	%Silt	%Sand	%Gravel	Plastic limit	Liquid limit	c' kPa	φ' degrees
Granite	6	8	75	11	-	34	0	38
Slate	7	50	38	5	31	34	0	34
mean per cycle rates of downslope soil transport								
Granite	5.8 cm ³ per cm width of slope							
Slate	31.2 cm ³ per cm width of slope							

Table 1: Summary of physical properties of soil used in CAEN experiment

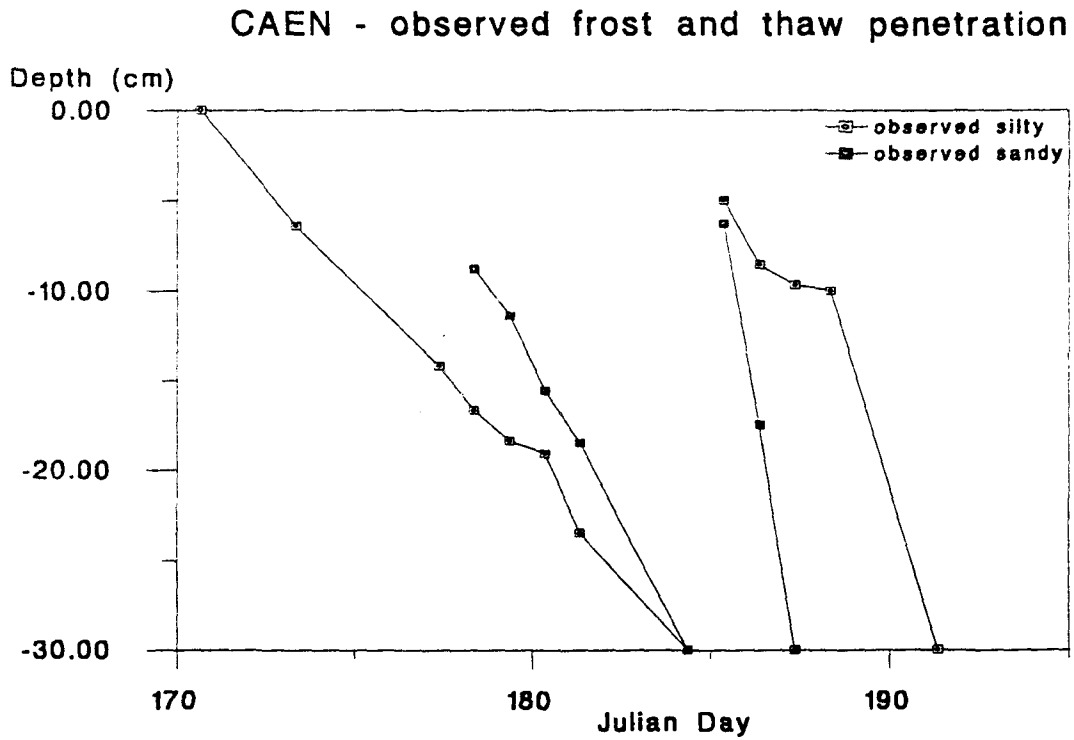
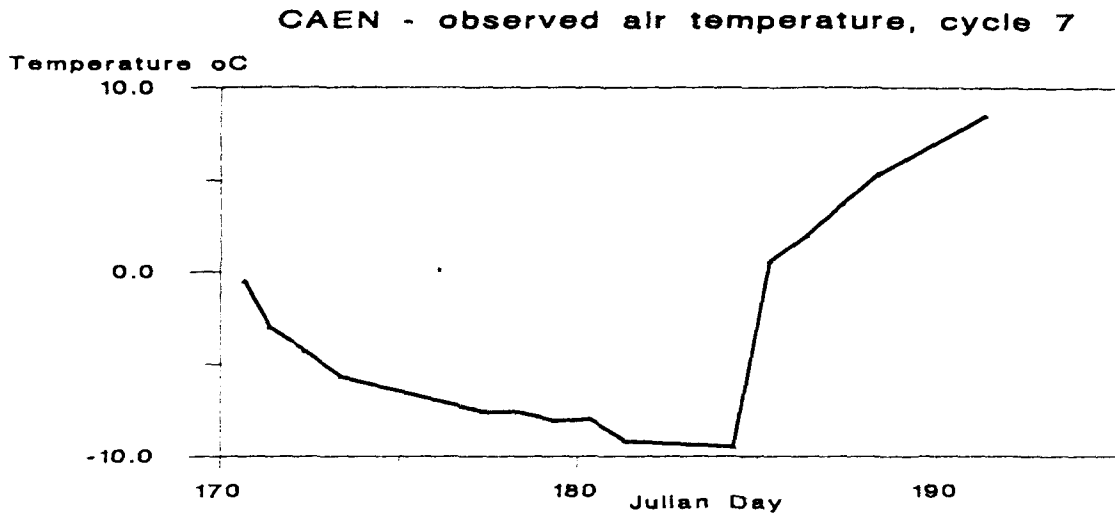
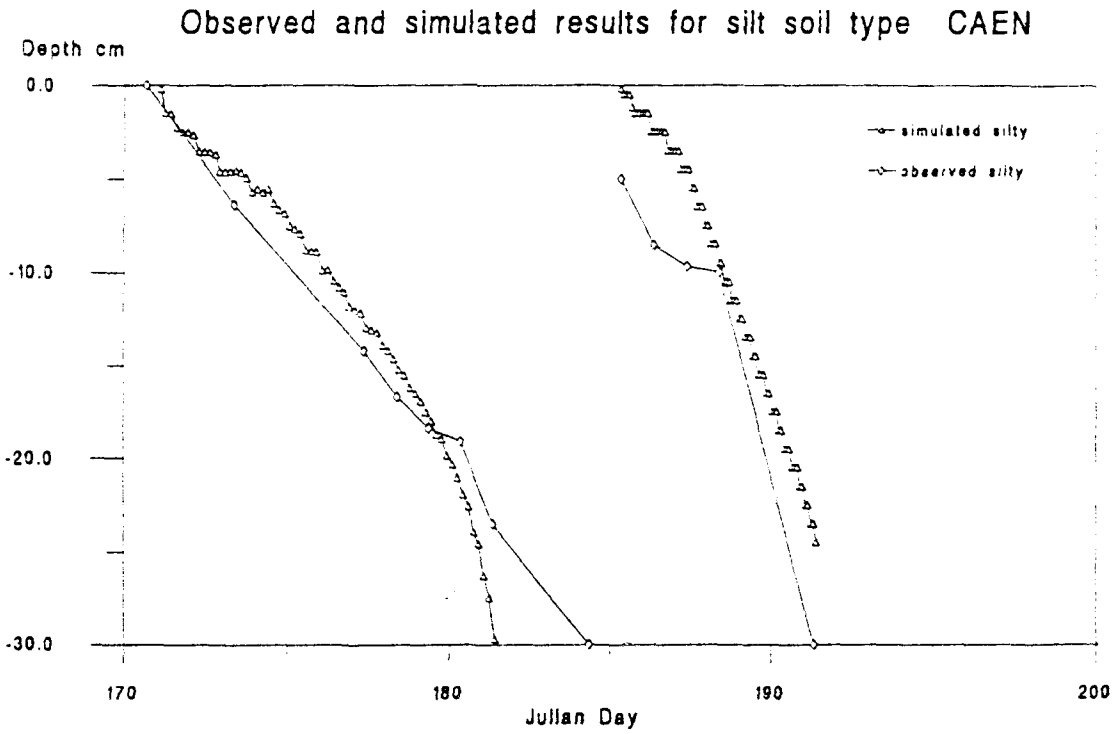
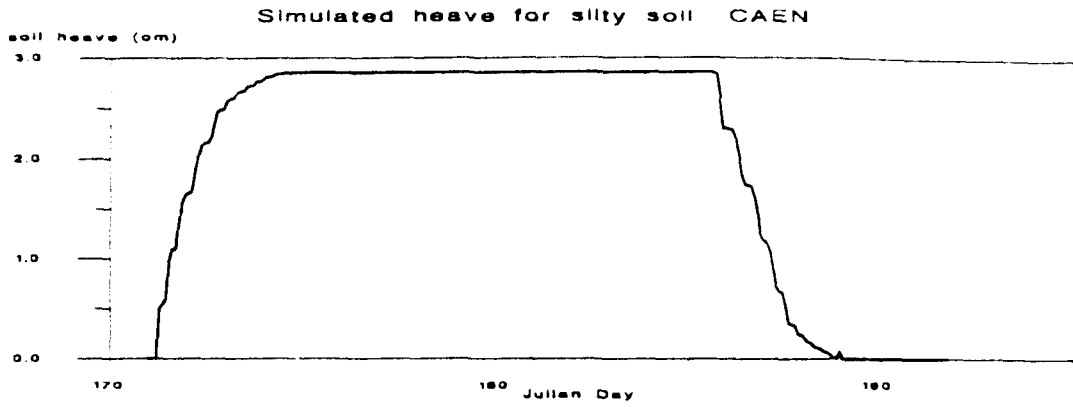


Figure 7: Observed ambient temperature and frost thaw depths for the CAEN laboratory experiment

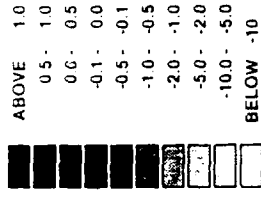


Downslope movement at the surface due to creep:	0.6 cm
Downslope movement at the surface due to gelifluction:	3.6 cm
Total downslope movement at surface for freeze/thaw cycle:	4.2 cm

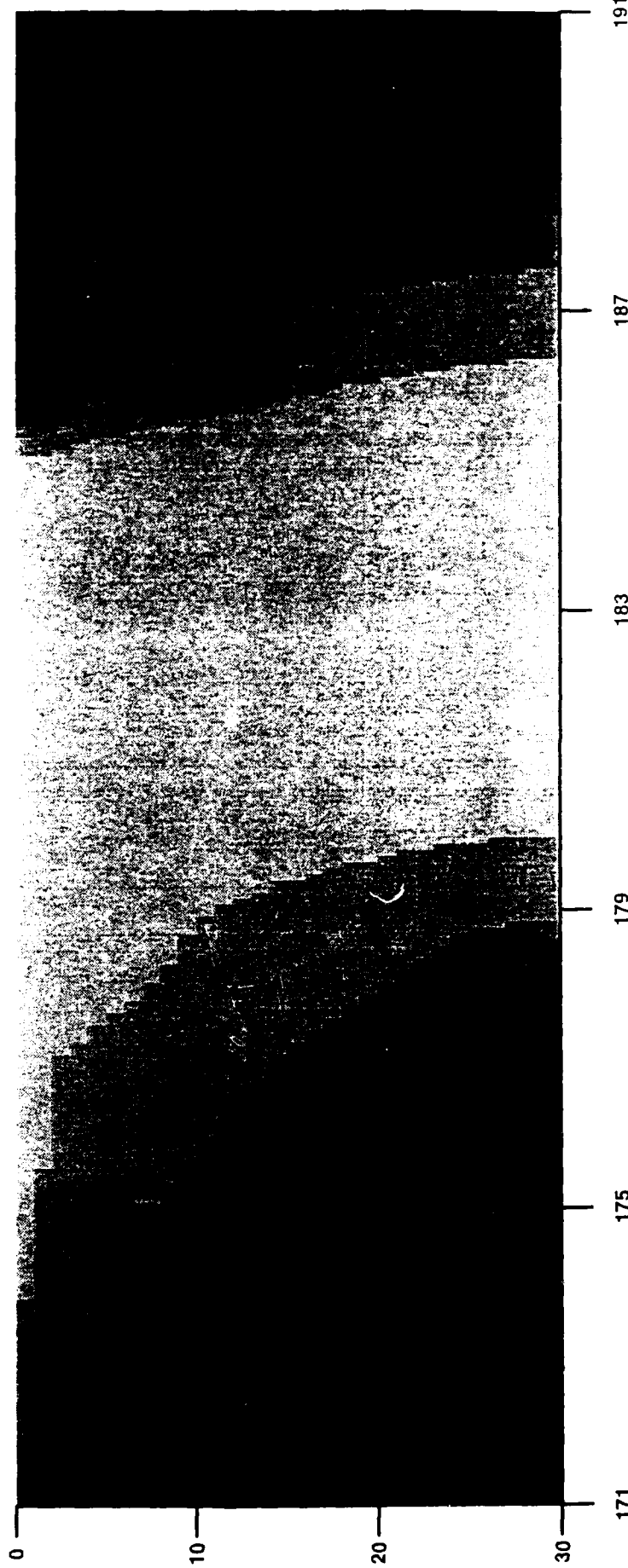
Figure 8: Results from simulation of silt soil type for the CAEN experiment

CAEN simulated profile temperatures

Soil Temperatures Centigrade



depth down profile cm



Julian Day

Figure 9. UNIRAS output showing profile temperatures over simulation period for CAEN silt soil.

For the CAEN study the slope angle was 12° . It is not really possible to compare observed and simulated values since the results from the CAEN study were used to determine relationships between excess water contents, soil heave and movement due to gelifuction. Further comparison with other cycles and additional data sources will help to ensure that relationships developed are general in nature.

5 CONCLUSIONS REACHED

- The PERIMOVE model is a complex model which is capable of simulating the important processes governing moisture and heat transfer through soils, soil heave and solifluction. The model operation and responses to changes in parameters and input data are as expected.
- very good simulations of observed data have been achieved with the model for a range of experimental conditions.
- Data collection in soil freezing and thawing environments is difficult but new techniques are being developed and many of these concentrate on methods for measuring liquid moisture contents within the soil environment. The collection of further 'internal' data will place further constraints on model validation and ensure that the model is behaving in the desired manner. Routes for further interaction with experimental projects have been identified.
- The PERIMOVE model can also be viewed as a tool to investigate the influence of changing the controlling factors governing freezing, thawing and mass movement to asses the influence of these.

6 FUTURE LINES OF RESEARCH ARISING FROM THE PROJECT

Two main lines of further research can be identified which would enhance the capability and applicability of the model. They divide into further improvement to the structure of the model and extensions to the testing and application of the model in association with experimental results.

Firstly to be considered is the extension to a fully 2-dimensional scheme incorporating a constantly deforming grid element structure. These are highly complex to program as model equations must then incorporate a term which includes the velocity due to the movement of the grid as well as the velocity due to the heat or water exchanges. The structure of such schemes has been investigated and a connection made with Mary Albert at CRREL (see for example Albert, 1984 and Albert and O'Neill 1986). Figure 10 shows diagrams of a deforming mesh from Albert (1984) where the movement of the phase boundary was modelled for a drum filled with

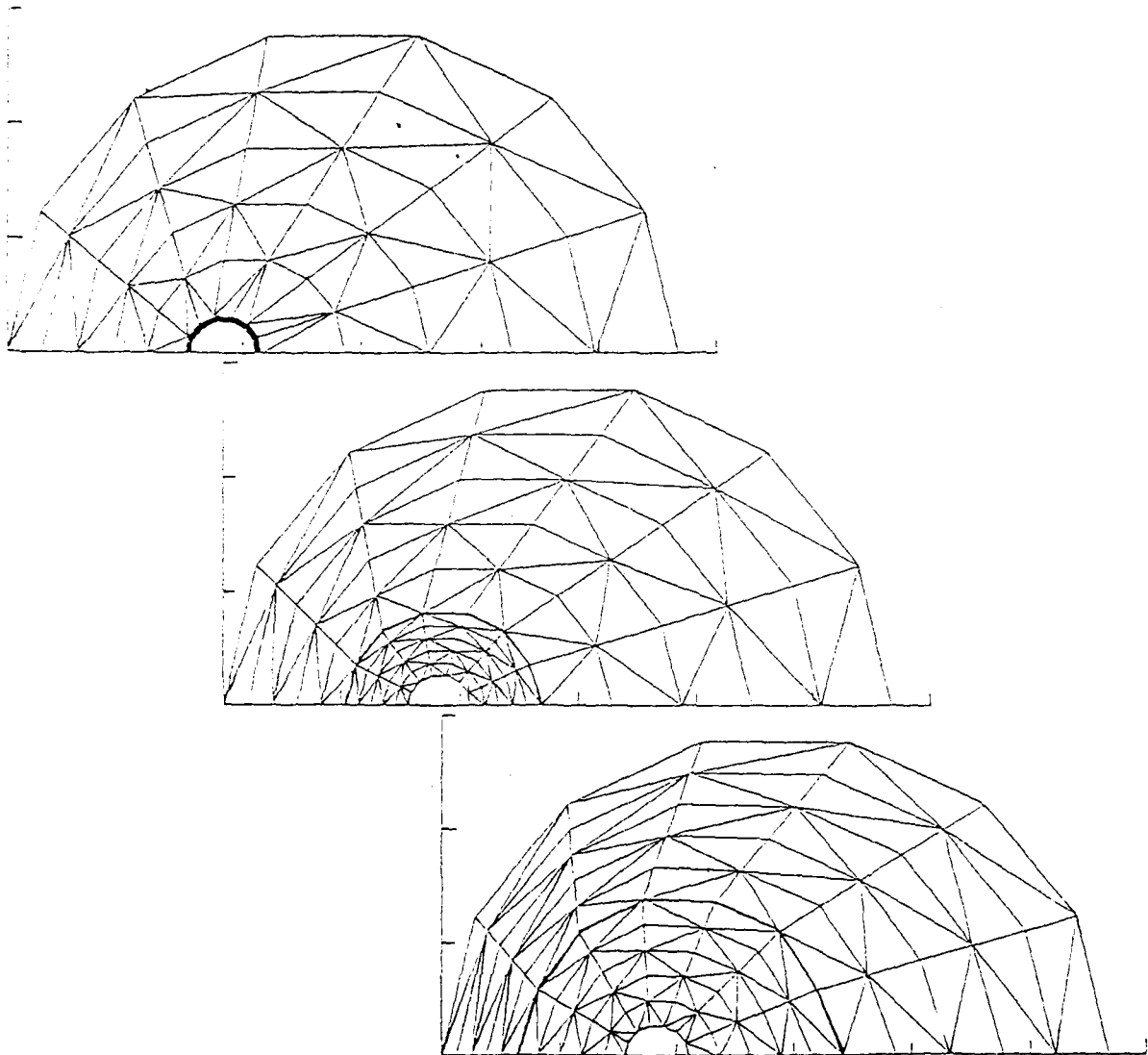


Figure 10: Stages in finite element mesh movement during simulation of freezing around a pipe. The dark line shows the phase front. Albert (1984).

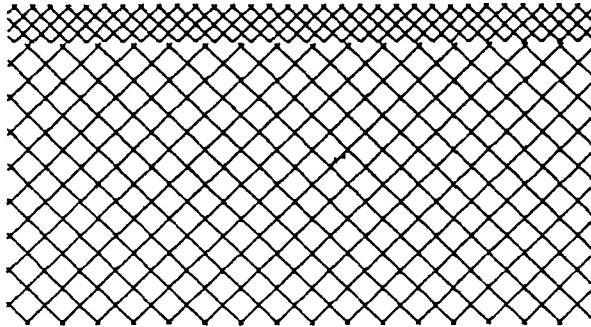
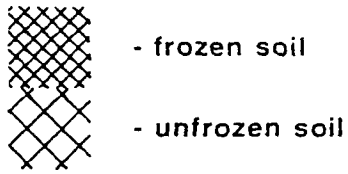
saturated sand in which freezing was instigated by the passage of cold ethylene glycol through a vertical copper pipe situated off-centre within the drum. It may be possible that with the continued development of finite element programming and automatic mesh generation techniques it will eventually be possible to obtain standard deforming grid packages to which the model can be linked. Alternatively, the connections established at CRREL may provide the basis for the development of a scheme to directly link in with the current model.

Secondly, as mentioned in the previous section is the desire to enhance the model by further interaction with observed data sources. Harris *et al* (1993) are continuing the project based at CAEN with a further series of freeze/thaw experiments planned in 1994. Additional data on soil moisture status during freezing will be collected then and the aim is to collect a more comprehensive data set particularly with regard to internal profile conditions during freezing and thawing. They aim to concentrate on two different soil types to highlight the differences between coarse grained sandy soils and finer grained silty soils.

In addition to the use of the model to simulate specific experimental conditions it should be possible to extend the use of the model as a tool in investigating controlling factors on slope movement and development in cold climates. This could be to investigate movements in current cold regions or to try to unravel climatic conditions which could have led to the formation of certain geomorphological features.

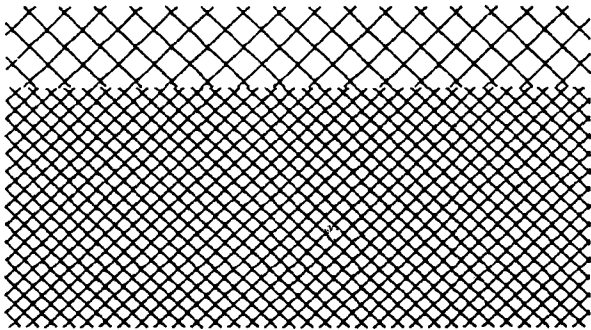
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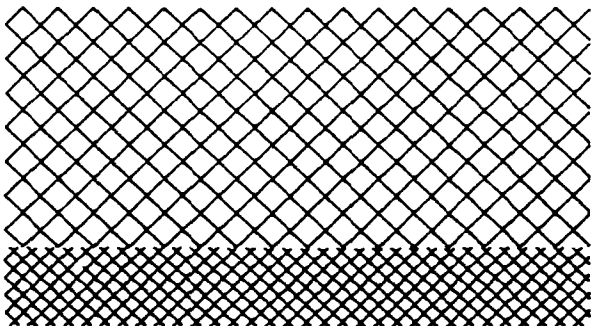
CONDITION (1)

critical depth of frozen soil which will support chosen vehicle



CONDITION (2)

critical depth of unfrozen ('slippery') layer where traction is too low. If tires reach frozen layer then zone may be trafficable.



CONDITION (3)

soil moisture status is critical for determining trafficability. The frozen layer is too far down the profile to give support and acts to impede drainage.

Figure 11: Summary diagram of critical soil trafficability conditions in freeze/thaw environments.

APPENDIX 1

Input files required to run PERIMOVE and parameter definitions.

GUIDE TO THE INPUT FILES REQUIRED TO RUN PERIMOVE

The input files are listed in alphabetical order. All inputs are in free format unless otherwise stated.

INDEX:

Ambient1.data	initial ambient temperature and number of temperature changes in simulation.
Ambt.data	array of ambient temperatures for simulation.
Cvin.data	soil freezing parameter data.
Hydro.data	soil hydrological property data.
InitialH.data	soil hydrological condition at start of simulation.
InitialT.data	soil temperature condition at start of simulation.
Kcurve.data	ice characteristic curve data.
Out.data	output option controls.
SM.data	data for suction/moisture curve.
Soli.data	data for solifluction potential.
Struc.data	cell configuration and definition of computational points.
TW.data	data for temperature/water curve.
Temporal.data	controls for length of simulation.
Ttsec.data	time of each temperature change.

Ambient1.data (unit 49)

L1	NOTC, STAMBT, BTEMP
	NOTC number of temperature changes in run.
	STAMBT initial ambient temperature (K).
	BTEMP basal boundary temperature (K).

Ambt.data (unit 48)

L1	AMBT(NOTC)
	AMBT ambient temperature (K).

* L1 is repeated NOTC times.

Cvin.data (unit 59)

L1		UDENS, FDENS, SHCSOIL, CW, CI
	UDENS	unfrozen soil bulk density g m^{-3} .
	FDENS	frozen dry bulk density g m^{-3} .
	SHCSOIL	specific heat capacity of soil $\text{J g}^{-1} \text{K}^{-1}$.
	CW	specific heat capacity of soil water $\text{J g}^{-1} \text{K}^{-1}$.
	CI	specific heat capacity of ice $\text{J g}^{-1} \text{K}^{-1}$.

Hydro.data (unit 57)

L1		STCON1, STCON2
	STCON1	saturated soil hydraulic conductivity for soil layer 1 m s^{-1} .
	STCON2	saturated soil hydraulic conductivity for soil layer 2 m s^{-1} .
L2		ASR1, ASR2
	ASR1	saturated soil moisture content for soil layer 1 $\text{m}^3 \text{m}^{-3}$.
	ASR2	saturated soil moisture content for soil layer 2 $\text{m}^3 \text{m}^{-3}$.
L3		(HYDG(I), I=1,NC)
	HYDG	hydraulic gradient under saturated flow - based on slope angle.
	NC	number of columns in simulation.
L4		(ORES1, I=1,NC)
	ORES1	residual soil moisture content for layer 1.
L5		(ORES2, I=1,NC)
	ORES2	residual soil moisture content for layer 2.

InitialH.data (unit 54)

L1		(ATHESP(I,J), J=1,M)
	ATHESP	initial soil moisture contents $\text{m}^3 \text{m}^{-3}$.
	M	number of cells in column.
	* L1 is repeated NC times.	

L2 (VATHET(I,J), J=1,M)
 VATHET ice content or potential volume of moisture available for release
 on thaw $\text{m}^3 \text{m}^{-3}$.

InitialT.data (unit 53)

L1 (sltp(I,J), J=1,M)
 sltp initial soil temperature (K) for each cell.
 M number of cells in column.
 * L1 is repeated NC times (I=1,NC).

Kcurve.data (unit 58)

L1 NKI
 NKI number of values on thermal conductivity of ice with temperature
 curve (Ktl/T).

L2 (KI(I), I=1,NKI)
 KI thermal conductivity of ice on Ktl/T curve $\text{W m}^{-1} \text{K}^{-1}$.

L3 (TKI(I), I=1,NKI)
 TKI temperature values on Ktl/T curve K.

L4 NKO
 NKO number of values on thermal conductivity of water with temperature
 curve (Ktθ/T).

L5 (KO(I), I=1,NKO)
 KO thermal conductivity of water on Ktθ/T curve $\text{W m}^{-1} \text{K}^{-1}$.

L6 (TKO(I), I=1,NKO)
 TKO temperature values on Ktθ/T curve °K.

L7 TDEP temperature at which ice starts to melt

L8 (KS(I,J), J=1,NNL(I))
 KS intrinsic thermal conductivity of soil particles $\text{W m}^{-1} \text{K}^{-1}$.

NNL number of cells in column
 * L8 repeated for each column (l=1,NC).

Out.data (unit 61)

L1 outmin, nops
 outsec iteration period for output of data (minutes).
 nops number of output data options.

L2 (option(l), l=1,nops)
 option output data option codes. (see key below)

L3 nslices
 nslices number of columns for which graphics data are to be written out.

L4 (slice(K), K=1,nslices)
 slice column for which graphics data is to be written out.

L5 NOCOP number of cells for which output is required

L6 (CELLOP(l), l=1,NOCOP)
 cell numbers for which output is required
 *L5 and L6 are repeated for each column (NC times)

L7 outstart minute from which output is produced

Key to PERIMO output options.

Numerical data files with captions.

OPTION	FILE NAME	FORTTRAN UNIT NO.	DATA
1	sltp.data	81	sltp
2	ice.data	82	VATHET
3	water.data	83	ATHESP
4	nhflux.data	84	NHFLUX
5	hflux.data	85	HFLUX
6	heatcap.data	86	VHC
7	thermcon.data	87	APKT
8	hydcon.data	88	KU
9	suction.data	89	sct
10	fcum.data	90	FCUM
11	fmax.data	91	FMAX
12	fmin.data	92	FMIN
13	fs.data	93	FS

14	ath1.data	94	ATH1
15	vath1.data	95	VATH1
16	atemp.data	96	ATEMP
17	btflux.data	97	BTFLUX
40	move.data	40	GDISP+HDISP

Graphical data files

OPTION	FILE NAME	FORTTRAN UNIT NO.	DATA
20	Gsttp.data	70	sttp
21	Gice.data	72	VATHET
22	Gwater.data	73	ATHESP
23	Gsuction.data	73	sct
24	Ghycond.data	74	KP
25	gtemp.data	75	ATEMP
26	Gkt.data	76	KT
27	Gapkt.data	77	APKT
28	Gfs.data	78	FS
29	Gnhflux.data	79	NHFLUX
30	Ghflux.data	80	HFLUX
33	Gexp.data	33	COLHT, COLHT-DEPTH
34	Gcelexp.data	34	EXP
35	fort.300-400	(isotherm position)	ISO, KKK
	e.g. fort.302 contains ISO and KKK data for column 2		
36	Gsoilp.data	36	SOILP
37	Gexwat.data	37	EXWAT
38	Gdisplace.data	38	HDISP, GDISP

SM.data (unit 56)

L1	NQ1, NQ2, KQ
	NQ1 number of observations on suction/moisture curve for hydrological layer 1.
	NQ2 number of observations on suction/moisture curve for hydrological layer 2.
	KQ number of observations on hydraulic conductivity/moisture curve.
L2	(X1(l), l=1,NQ1)
	X1 moisture values on suction/moisture curve for layer 1 m ³ m ⁻³ .
L3	(Y1(l), l=1,NQ1)
	Y1 suction values on suction/moisture curve for layer 1 m.

* lines L2 and L3 are repeated for layer 2 for X2 and Y2.

Soli.data (unit 60)

L1		GAMMAS, GAMMAU
	GAMMAS	saturated unit weight of soil g m^{-3} .
	GAMMAU	unsaturated unit weight of soil g m^{-3} .
L2		(BETA(I), I=1,NC)
	BETA	slope angle in degrees
L3		(C(I,J), J=1,M)
	C	cohesion KPa for each cell (J=1,M).
		* L3 is repeated for each column.
L4		(PHI(I,J), J=1,M)
	PHI	internal friction angle in radians.
		* L4 is repeated for each column.
L5	GELRAT	heave:gelifluction ratio for soil type

Structural.data (unit 51)

L1		NC, NOC, LENGTH, BDT
	NC	number of columns in simulation
	NOC	number of cells in first column of the profile
	LENGTH	cell length (m).
	BDT	breadth of cell (m)
L2		(NNL(I), I=1,NC)
	NNL	number of cells in each column of the profile.
L3		(CDEPTH(I,J), dcp(I,J), I=1,NC, J=1,NNL(I))
	CDEPTH	cell depth (m)
	dcp	depth to computation point in cell(m)
L4		(DEPTH(I), I=1,NC)
	DEPTH	depth of profile in each column (m).
L5		(HT(I,J), J=1,M)

HT height of each computation point above the profile base.
 * L3 is repeated for each column

L6 (LC1(I), I=1,NC)
 LC1 position (j) of last cell in hydrological layer 1.

L7 (FC2(I), I=1,NC)
 FC2 position (j) of first cell in hydrological layer 2.

L8 (TTCOM(I,J), J=1,M)
 depth of cell computation point from surface (m).
 * line L8 is repeated for each column.

TW.data (unit 55)

L1 TF
 TF number of observations on characteristic freezing curve.

L2 (TEMCV(I), I=1,TF)
 TEMCV temperature values on characteristic freezing curve, K (highest first).

L3 (WATECV(I), I=1,TF)
 WATECV moisture values on characteristic freezing curve ($m^3 m^{-3}$).

Temporal.data (unit=50)

L1 IP, NOSECR
 IP iteration period (seconds).
 NOMINR number of minutes in simulation.

Ttsec.data (unit=52)

L1 TTSEC(JJJ)
 TTSEC time from start of simulation to temperature change input to PERIMO model (minutes).
 JJJ number of temperature changes during simulation (NOTC).

DEFINITIONS OF VARIABLES FOR PERIMOVE

AMBT	ambient temperature K.
APKT	apparent thermal conductivity $W m^{-1} K^{-1}$.
ASR1	saturated moisture content for soil type 1 $m^3 m^{-3}$.
ASR2	saturated moisture content for soil type 2 $m^3 m^{-3}$.
ATEMP	Ambient temperature (K).
ATH1	volume of water in cell.
ATHESP	soil water content $m^3 m^{-3}$.
ATHESP1	soil water volume m^3 .
ATHET	volumetric water content derived from T/0 curve.
ATHET1	water content derived from T/0 curve, m^3 .
AVK	average hydraulic conductivity $m s^{-1}$.
AVKT	Average thermal conductivity across cell boundary.
BDT	cell breadth m.
BETA	slope angle.
BFLUX	volume of flux through base of profile.
BISO	depth of basal isotherm.
BWFLUX	basal flux.
C	cohesion.
CDEPTH	Depth (size) of cells (m).
CELVOL	cell volume.
CH	Temperature change.
CHPERC	Temperature change expected given net input or loss of energy on phase change. Calculated per cell.
COLHT	height of each soil column.
cmdt	Thermal component of energy budget for dT.
CVGR	curve gradient on T/0 curve.
dcp	depth to computation point in cell.
deltaT	Change in temperature during iteration period.
deltaw	Change in moisture during iteration period.
DEPTHP	depth of profile in each column.
DISO	depth from surface at which 273.15K isotherm occurs.
EXP	expansion of cell in vertical from original volume.
EXSAT	change in moisture content due to saturated flow.
EXWAT	excess water content expelled from cell on thawing.
FC2	position (j) of first cell in layer 2.
FCUM	cumulative factor of safety per column.
FDENS	frozen dry bulk density ($g m^{-3}$).
FDMAX	depth of max. FS.
FDMIN	depth of min. FS.
FLUX	soil water flux between cells, volume.
FMAX	maximum factor of safety per column.
FMIN	minimum factor of safety per column.
FS	factor of safety.
G1, G2	average incremental gradients on suction/moisture curve, layers 1 and 2.
GAMMAS	saturated unit weight of soil.
GAMMAU	unsaturated weight of soil.
GZ1, GZ2	average incremental gradients on K/0 curve.
HFLUX	Heat flux across cell boundary.
HPOT	hydraulic potential.
HT	height of each computation point above profile base.
HYDG	hydraulic gradient sat. flow, ie. slope angle.
IKG	incremental gradient on Kt/T curve.
IP	iteration period secs.
ISO	depth from previous compn. point at which 273.15K isotherm occurs.

ITT	count for iterations.
KI	thermal conductivity of ice, KtI/T curve.
KIK	thermal conductivity of ice at given temp.
KO	thermal conductivity of water, KtO/T curve.
KOK	thermal conductivity of water at given temp.
KP	cell hydraulic conductivity, $m s^{-1}$.
KQ	number of observations on K/O curve.
KS	intrinsic thermal conductivity of soil solids.
LA	Latent heat of fusion.
LAO	hydrological component of energy release for dT.
LC1	position (j) of last cell in layer 1.
LENGTH	cell length, m.
MORES	Mean residual moisture contents.
NC	no. of columns.
NEWVOL	updated cell volume following expansion or resettlement.
NFLUX	net water flux of cell.
NHFLUX	net heat flux $W m^{-3} s^{-1}$.
NKI	number of values on KtI/T curve.
NKO	number of values on KtO/T curve.
NNL	number of cells in each column of profile.
NNLMAX	maximum number of cells in any of the columns.
NOC	number of cells in first column of the profile.
NOIT	number of iterations in simulation.
NOMINR	number of minutes in run.
NOSECR	number of seconds in run.
NOTC	number of temperature changes.
NQ1, NQ2	number of observations on suction/moisture curve, layers 1 and 2.
NSEIT	number of seconds passed by end of iteration.
nslices	number of columns for which graphics data are written.
OKG	incremental gradient of KtO/T curve.
option	output data options.
ORES1, ORES2	residual soil moisture content, layer 1 & 2.
ORIVOL	cell volume at start of simulation.
OUTMIN	iteration period for output data, minutes.
OUTSEC	iteration period for output, minutes.
OUTSTART	control for start time for output, minutes.
PHI	internal friction angle of soil.
sct	soil suction in each cell.
SEC	difference between: nseit and tsec input.
SFI	shape factor ice.
SFO	shape factor water.
SFS	shape factor solids.
SHCSOIL	specific heat capacity of soil ($J m^{-3} K^{-1}$).
NC	number of columns.
slice	column for which graphics data is to be output.
sltp	soil temp. (K).
sltp2	soil temp calculated by heattrans.f on basis of changing ambient temp.
SOILP	percentage volume of soil particles per cell.
SOILV	volume of soil solids, m^{-3} .
STAMBT	starting ambient temperature.
STCH	ambient temperature.
STCON1 STCON2	saturated hydraulic conductivity $m s^{-1}$.
TEMCV	temperature values on T/O curve.
TEMCV	Temperature values on T/O curve (K).
TF	number of observations on T/O curve.
tgrad	temperature gradient across frozen fringe.
TKI	temperature values on KtI/T curve.
TKO	temperature values on KtO/T curve.

TQ number of observations on Ts/0 curve.
TTCOM depth of cell computation point from surface.
TTMIN time(minutes) from start of run to temperature change.
TTSEC time from start of run for temperature change.
UDENS unfrozen dry bulk density (g m⁻³).
VADFAC volume adjustment factor for cell expansion/contraction.
VATH1 volume of ice in cell.
VATHET ice content m³ m⁻³.
VOLCH ratio of volume to original cell volume.
VHC volumetric heat capacity of cell (J m⁻³ K⁻¹).
WATECV Moisture values on T/0 curve m³ m⁻³.
X1, X2 moisture values on suction moisture curve, layers 1 & 2.
Y1, Y2 suction values on suction moisture curve, layers 1 & 2.
Z1, Z2 conductivity values on conductivity moisture curve, layers 1 & 2.

APPENDIX 2

PERIMOVE model code.

PERIMOVE FORTRAN CODE

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```

c      Program PERIMOVE.f
c
c      *****
c      PERIMOVE.F
c      *****
c      Program to predict the volume of ice present in a soil profile during
c      freezing and thawing. The program also predicts heave and resettlement
c      of the soil profile and volumetric water excesses following thawing.
c
c      The program calculates the soil temperature at depth, through
c      time, modifies the temperature for transfer of latent heat
c      and updates soil moisture content accordingly.
c
c      *****
c      program PERI
c      program perimove
c      *****
c      DECLARATION OF VARIABLES
c      *****
c      DOUBLE PRECISION TTCOM(5,60), HT(5,60),COLHT(5)
c      DOUBLE PRECISION LENGTH, CDEPTH(5,60), BDT,CELVOL(5,60)
c      DOUBLE PRECISION dcp(5,60),ORIVOL(5,60),VADFAC(5,60)
c      DOUBLE PRECISION DEPTH(5),SOILP(5,60),EXP(5,60),CDEP(5,60)
c      DOUBLE PRECISION ATHET(5,60), VATHET(5,60),ATHESP(5,60)
c      DOUBLE PRECISION ORES1(50), ORES2(50), X1(50), X2(50)
c      DOUBLE PRECISION Y1(50), Y2(50),STCON1, STCON2
c      DOUBLE PRECISION ASR1, ASR2, sltp(5,60), sltpt2(5,60)
c      DOUBLE PRECISION KP(5,60),sct(5,60),EXWAT(5,60)
c      DOUBLE PRECISION LA, VHC(5,60), G1(50), G2(50)
c      DOUBLE PRECISION Z1(50), Z2(50), GZ1(50), GZ2(50)
c      DOUBLE PRECISION FLUX(5,60), NFLUX(5,60), HPOT(5,60)
c      DOUBLE PRECISION BWFLUX(50), STATH(5,60), STVATH(5,60)
c      DOUBLE PRECISION AVK(5,60), HYDG(50)
c      DOUBLE PRECISION TEMCV(50), WATECV(50), CVGR(50),TDEP
c      DOUBLE PRECISION APKT(5,60), KT(5,60)
c      DOUBLE PRECISION HFLUX(5,60), NHFLUX(5,60)
c      DOUBLE PRECISION AMBT(525600), ATEMP(1468800)
c      DOUBLE PRECISION STAMBT,BTEMP
c      DOUBLE PRECISION FS(5,60), FMIN(50), FMAX(50)
c      DOUBLE PRECISION FDMAX(50), FDMIN(50), FCUM(50)
c      DOUBLE PRECISION EXSAT(50),ISO(5,60),SOIL1(5,60)
c      DOUBLE PRECISION SHATHET(5,60), SHVATHET(5,60)
c      DOUBLE PRECISION HTHETAC(50),KI(20),TKI(20),KS(5,60)
c      DOUBLE PRECISION KO(20),TKO(20),OKG(20),IKG(20)
c      DOUBLE PRECISION C(5,60), BETA(5),MAXCOLHT(5),HDISP(5)
c      DOUBLE PRECISION GAMMAS, GAMMAU,GELRAT,GDISP(5)
c      DOUBLE PRECISION PHI(5,60)
c
c      Integers
c      INTEGER NC,>NNL(5), NOC
c      INTEGER>NNLMAX, col(50), ncol
c      INTEGER SEC(525600)
c      INTEGER outsec
c      real outmin
c      INTEGER NQ1, NQ2, KQ, TF
c      INTEGER LC1(50), FC2(50)
c      INTEGER option(50)
c      INTEGER IP

```

```

INTEGER NOTC
INTEGER NSEIT(1468800), ITT, NOIT
INTEGER NOSECR,NOMINR,TTSEC(525600),TTMIN(525600)
INTEGER slice(50), nsllices
INTEGER isocol(50), isoncol
INTEGER NK0, NKI, KKK
INTEGER NOCOPV,NNLMA,CELL
INTEGER NOCOP(30),CELLOP(5,60)
INTEGER OUTSTART,OUTTIM

```

```
COMMON DUMMY(5,60)
```

```

C
C *****
C OPEN STATEMENTS
C *****
open(unit=20, file='DATAIN/Mb.data', status='old')
rewind 20

open(unit=21, file='DATAIN/Iso.data', status='old')
rewind 21

c open(unit=32, file='DATAIN/massbal.data', status='unknown')
c rewind 32

open(unit=48, file='DATAIN/Ambt.data', status='old')

open(unit=49, file='DATAIN/Ambient1.data', status='old')
rewind 49

open(unit=50, file='DATAIN/Temporal.data', status='old')
rewind 50

open(unit=51, file='DATAIN/Struc.data', status='old')
rewind 51

open(unit=52, file='DATAIN/Ttsec.data', status='old')

open(unit=53, file='DATAIN/InitialT.data', status='old')
rewind 53

open(unit=54, file='DATAIN/InitialH.data', status='old')
rewind 54

open(unit=55, file='DATAIN/TW.data', status='old')
rewind 55

open(unit=56, file='DATAIN/SM.data', status='old')
rewind 56

open(unit=57, file='DATAIN/Hydro.data', status='old')
rewind 57

open(unit=58, file='DATAIN/Kcurve.data', status='old')
rewind 58

open(unit=59, file='DATAIN/Cv.data', status='unknown')
rewind 59

```

```

open(unit=60, file='DATAIN/Soli.data', status='old')
rewind 60

open(unit=61, file='DATAIN/Out.data', status='old')
rewind 61

OUTTIM=1
*****
c
c  TEMPORAL.data  unit 50
c  *****
c  Read iteration time and no. of minutes in run.
c  read(50,*)IP, NOMINR
c  convert no. of minutes in run to no of seconds
c  NOSECR=NOMINR*60

c  *****
c  STRUCTURAL.data  unit 51
c  *****
c  Space dimensions.
c  READ(51,*)NC, NOC, LENGTH, BDT

c  Number of cells in each column
c  READ(51,*)(NNL(I),I=1,NC)

c  calculate volume of each cell and store in ORIVOL for volume
c  change calculations.  Initialise volume adjustment factors.
DO 16 I=1,NC
  DO 17 J=1,NNL(I)
    READ(51,*) CDEPTH(I,J),dcp(I,J)
    CELVOL(I,J)=LENGTH*BDT*CDEPTH(I,J)
    ORIVOL(I,J)=CELVOL(I,J)
    VADFAC(I,J)=1.0
    EXP(I,J)=0.0
    CDEP(I,J)=CDEPTH(I,J)
17  CONTINUE
16  CONTINUE
c

c
c >NNLMAX=0
c  DO 1 I=1,NC
c  IF(NNL(I).GT.NNLMAX)THEN
c >NNLMAX=NNL(I)
c  ENDIF
1  CONTINUE
c

c  Depth of profile in each column.
c  READ(51,*)(DEPTHP(I),I=1,NC)
c

c  Height of computation point above profile base.
DO 2 I=1,NC
  M=NNL(I)
  READ(51,*)(HT(I,J),J=1,M)
2  CONTINUE
c

c  Define LC1 and FC2.
c  READ(51,*)(LC1(I),I=1,NC)
c  READ(51,*)(FC2(I),I=1,NC)

```



```

c
c   Depth of each computation point from surface.
DO 3 I=1,NC
M=NNL(I)
READ(51,*)(TTCOM(I,J),J=1,M)
3   CONTINUE
do 3012 I=1,NC
write(39,*)depthp(i)
MAXCOLHT(I)=DEPTH(I)
do 3013 J=1,M
WRITE(39,*)I,J,HT(I,J),TTCOM(I,J),CELVOL(I,J)
3013 CONTINUE
3012 CONTINUE

c
c   *****
c   INITIALT.data
c   *****
c
c   Starting soil temperatures.
DO 4 I=1, NC
M=NNL(I)
READ(53,*)(sltp(I,J),J=1,M)
4   CONTINUE
c
c   *****
c   INITIALH.data   unit 54
c   *****
c
c   Starting moisture conditions
DO 5 I=1,NC
M=NNL(I)
read(54,*)(ATHESP(I,J),J=1,M)
5   continue

c
c   Starting ice content
DO 60 I=1,NC
M=NNL(I)
read(54,*)(VATHET(I,J),J=1,M)
60  continue
c
c   *****
c   TW.data   unit 55
c   *****
c
c   Number of values on soil water characteristic curve for temps
c   below 273.15K. (Temperature and moisture values)
c   Ambient thermal conductivity
READ(55,*)TF

c
c   Values on soil water characteristic curve.
READ(55,*)(TEMCV(I),I=1,TF)
READ(55,*)(WATECV(I),I=1,TF)

c
c   *****
c   Ktcurve.data   unit
c   *****
c
c   Read no. of values on KI/T curve
read(58,*)NKI

c
c   Read Ki and T values on KI/T curve
read(58,*)(KI(I),I=1,NKI)
read(58,*)(TKI(I),I=1,NKI)

```

```

c   Read no. of values on KO/T curve
    read(58,*)NKO

c   Read KO and T values on KO/T curve
    read(58,*)(KO(I),I=1,NKO)
    read(58,*)(TKO(I),I=1,NKO)

c   Read thaw initiation temperature TDEP
    read(58,*)TDEP

c   Read values for intrinsic thermal conductivity of soil particles
    do 25 I=1,NC
      read(58,*)(KS(I,J),J=1,NNL(I))
25  continue

```

```

c   Establish thermal conductivity of water
c   Incremental gradients on K/O curve
    do 26 I=1,NKO-1
      OKG(I)=(KO(I+1)-KO(I))/(TKO(I+1)-TKO(I))
26  continue

```

```

c   Incremental gradients on K/I curve
    do 27 I=1,NKI-1
      IKG(I)=(KI(I+1)-KI(I))/(TKI(I+1)-TKI(I))
27  continue

```

```

c   *****
c   SM.data  unit 56
c   *****
c   Number of observations on suction moisture curve, Ku/moisture
c   curve.
    READ(56,*)NQ1, NQ2, KQ

```

```

c   Values of moisture and suction on suction moisture curve.
    READ(56,*)(X1(I),I=1,NQ1)

    READ(56,*)(Y1(I),I=1,NQ1)

    READ(56,*)(X2(I),I=1,NQ2)

    READ(56,*)(Y2(I),I=1,NQ2)

```

```

c   *****
c   HYDRO.data  unit 57
c   *****
c   Saturated hydraulic conductivity
    READ(57,*)STCON1, STCON2

```

```

c   Saturated moisture content for layers 1 and 2, and hydraulic gradient
c   (slope angle)
    READ(57,*)ASR1, ASR2
    READ(57,*)(HYDG(I),I=1,NC)

c   Initialise starting volumetric contents of soil solids
    DO 591 I=1,NC

```

```

      M=NNL(I)
      DO 592 J=1,M
        if(J.LE.LC1(I))SOILP(I,J)=1.0-ASR1
        if(J.GE.FC2(I))SOILP(I,J)=1.0-ASR2
        SOIL1(I,J)=SOILP(I,J)*CELVOL(I,J)
592   continue
591   continue
c
c   Residual soil moisture content for layers 1 and 2.
READ(57,*)(ORES1(I),I=1,NC)
READ(57,*)(ORES2(I),I=1,NC)
c
c   .....
c   READ IN STABILITY DATA
c   .....
      READ(60,*)GAMMAS, GAMMAU

      READ(60,*)(BETA(I),I=1,NC)

      DO 40 I=1,NC
        M=NNL(I)
        READ(60,*)(C(I,J),J=1,M)
40    CONTINUE

      DO 41 I=1,NC
        M=NNL(I)
        READ(60,*)(PHI(I,J),J=1,M)
41    CONTINUE

      DO 42 I=1,NC
        M=NNL(I)
        READ(60,*)GELRAT
42    CONTINUE

c   .....
c   OUT.data   unit 61
c   .....
c   Iteration period of output, number of data sets.
read(61,*)outmin,nops
c   convert iteration period to seconds
outsec=outmin*60
c   Data set definitions
read(61,*)(option(I),I=1,nops)
c   Graphics, column data to sent to graphics files.
read(61,*)nslices
read(61,*)(slice(K),K=1,nslices)

do 598 I=1,NC
  READ(61,*)NOCOP(I)
  NOCOPV=NOCOP(I)
  READ(61,*)(CELLOP(I,M),M=1,NOCOPV)
598  CONTINUE

do 601 I=1,NC
  M=1
  CELL = CELLOP(I,M)
 >NNLMA=NNL(I)
  DO 602 J=1,>NNLMA
    IF(CELL.EQ.J)THEN

```

```

        DUMMY(I,J)=1
        M=M+1
        CELL=CELLP(I,M)
        ELSE
        DUMMY(I,J)=0
        ENDIF
602    CONTINUE

601    CONTINUE
c      read time in minutes at which output is to start
      READ(61,*) OUTSTART
      OUTSTART=OUTSTART*60

c      *****
c      Mb.data
c      *****
c      Number of columns for which mass balance data is to be written
      read(20,*)ncol

c      Column numbers for which mass balance data is to be written
      do 350 l=1,ncol
      read(20,*)col(l)
350    continue
c      *****
c      iso.data
c      *****
c      Number of columns for which isotherm is to be identified
      read(21,*)isoncol

c      Column numbers for which isotherm is to be identified
      do 351 l=1,isoncol
      read(21,*)isocol(l)
351    continue
c      *****
c      INITIAL SECTION
c      *****
c      MILLINGTON and QUIRK method is used to calculate
c      hydraulic conductivity for each layer from a given
c      soil moisture characteristic curve.
c
      NQJ=NQ1

c      Layer 1
      DO 6 l=1,NQJ
      llJ=NQJ-l+1
      XII=X1(llJ)
      TOPS=0.0
      BOTS=0.0
      DO 7 J=1,NQJ
      JF=NQJ-J+1
      YJJ=Y1(JF)
7      BOTS=((2*J-1)*YJJ**(-2))+BOTS
      ll=l
      DO 8 J=ll,NQJ
      JF=NQJ-J+1
      YJJ=Y1(JF)
8      TOPS=((2*J+1-2*ll)*YJJ**(-2))+TOPS

```

```

JT=NQJ-I+1

Z1(JT)=STCON1*(XII/ASR1)*TOPS/BOTS
open(unit=2,file='zval',status='unknown')
6  write(2,*)JT,Z1(JT),XII

c  Layer 2
   NQJ=NQ2
   DO 9 I=1,NQJ
     IJ=NQJ-I+1
     XII=X2(IJ)
     TOPS=0.0
     BOTS=0.0
     DO 10 J=1,NQJ
       JF=NQJ-J+1
       YJJ=Y2(JF)
10  BOTS=((2*J-1)*YJJ**(-2))+BOTS
     II=I
     DO 11 J=II,NQJ
       JF=NQJ-J+1
       YJJ=Y2(JF)
11  TOPS=((2*J+1-2*I)*YJJ**(-2))+TOPS
     JT=NQJ-I+1

9  Z2(JT)=STCON2*(XII/ASR2)*TOPS/BOTS

c  Calculate gradients on suction/moisture curves, K/moisture
c  curve, and T/moisture characteristic curve.

c  Calculate average incremental gradients of suction-moisture
c  and K-moisture curves.
c  Layer 1
   DO 12 I=1,NQ1-1
     G1(I)=(Y1(I+1)-Y1(I))/(X1(I+1)-X1(I))
     GZ1(I)=(Z1(I+1)-Z1(I))/(X1(I+1)-X1(I))
12  CONTINUE

c  Layer 2
   DO 13 I=1,NQ2-1
     G2(I)=(Y2(I+1)-Y2(I))/(X2(I+1)-X2(I))
     GZ2(I)=(Z2(I+1)-Z2(I))/(X2(I+1)-X2(I))
13  CONTINUE

c  T/0 curve.
   DO 14 I=1,TF-1
     CVGR(I)=(WATECV(I+1)-WATECV(I))/(TEMCV(I+1)-TEMCV(I))
14  CONTINUE
c  *****
c  SET PARAMETERS
c  *****
c  Latent heat of fusion, units J g-1
   LA=333.2852
c  *****
c  DYNAMIC SECTION
c  *****
c  Set number of iterations
c  Iterations = no. of seconds in run/ iteration period

```

```

NOIT=NOSECR/IP

read(49,*)NOTC,STAMBT,BTEMP

DO 600 KKK=1,NOIT
ITT=KKK

    DO 650 I=1,NC
    M=NNL(I)
    DO 651 J=1,M
    STATH(I,J)=ATHESP(I,J)
    STVATH(I,J)=VATHET(I,J)
651     CONTINUE
650     CONTINUE

c   Second marking the end of iteration
NSEIT(KKK)=ITT*IP

JJJ=0

c   Set up loop for the number of temp. changes in data set.
DO 500 JJJ=1,NOTC
    if(JJJ.eq.1)then
        rewind 52
        rewind 48
    endif

c   read in time in minutes of next temperature change
read(52,*)TTMIN(JJJ)

c   convert this value to seconds
TTSEC(JJJ)=TTMIN(JJJ)*60
read(48,*)AMBT(JJJ)

c   -----
if(TTSEC(JJJ).gt.NSEIT(KKK))then

    if(KKK.eq.1)then
        SEC(JJJ)=TTSEC(JJJ)
        ATEMP(KKK)=STAMBT+(((AMBT(JJJ)-STAMBT)/SEC(JJJ))
& *IP)
    endif

    if(KKK.gt.1)then
        ATEMP(KKK)=ATEMP(KKK-1)+(((AMBT(JJJ)-ATEMP(KKK-1))
& /(TTSEC(JJJ)-NSEIT(KKK-1)))*IP)
    endif

    do 900 NN=1,nops
        if(option(NN).eq.31)then
            open(unit=8, file='atemp.data', status='unknown')
            rewind 8
            open(unit=9, file='Gatemp.data', status='unknown')
            rewind 9
            write(8,*)ATEMP=,ATEMP(KKK),'KKK=',KKK
            write(9,*)ATEMP(KKK),KKK
        endif
900     continue

```

```

        goto 700
    endif

c -----
    if(TTSEC(JJJ).eq.NSEIT(KKK))then
        ATEMP(KKK)=AMBT(JJJ)
        do 1000 NN=1,nops
            if(option(NN).eq.31)then
                open(unit=8,file='atemp.data',status='unknown')
                rewind 8
                open(unit=9, file='Gatemp.data', status='unknown')
                rewind 9
                write(8,*)ATEMP=,ATEMP(KKK),'KKK=',KKK
                write(9,*)ATEMP(KKK),KKK
            endif
1000    continue
        goto 700
    endif

    if(TTSEC(JJJ).lt.NSEIT(KKK))goto 500

500    continue

700    call ku(NC, LC1, FC2, KQ, ATHESP,
    & X1, X2, Z1, Z2, GZ1, GZ2, KP)

    call kta(VATHET, ATHESP, CELVOL,
    & sltp,>NNLMAX, NC,>NNL, ASR1, ASR2, LA,
    & APKT, KT,KI,NKI,KO,NKO,TKI,TKO,KS,OKG,IKG,SOILP)

    call heatrans(ATEMP, CDEPTH, sltp, sltp2,
    & VATHET, NC,>NNL, VHC, LENGTH, option,
    & ATHESP,>NNLMAX, BDT, TEMCV, WATECV,
    & TF, LA, ORES1, ORES2, CVGR, ATHET, HFLUX, NHFLUX,
    & APKT, STAMBT, IP, KKK, LC1, FC2, ASR1, ASR2,
    & SHVATHET, CELVOL,ORIVOL, VADFAC,SOILP,SOIL1,TDEP,BTEMP,EXWAT)

    call masstrans(NC,LC1,NQ1,ATHESP,X1,Y1,G1,FC2,>NNL,
    & Y2,G2,HT,KP,KQ,Z1,GZ1,Z2,GZ2,CDEPTH,LENGTH,
    & sct,HPOT,AVK,FLUX,NFLUX,HYDG,IP,DEPTHP,TTCOM,dcp,
    & BDT, X2, NQ2, STCON1, STCON2, ASR1, ASR2,
    & EXSAT, BWFLUX, HTHETAC, SHATHET,VATHET,COLHT,
    & CDEP,EXP,vadfac,MAXCOLHT)

c    call massbal(ATHESP, VATHET, NC,>NNL, STATH, STVATH,
c    & NFLUX, KKK, EXSAT, NSEIT, BWFLUX, HTHETAC,
c    & SHATHET, SHVATHET, col, ncol, nops, option)

    call soli(NC,>NNL,>NNLMAX, TTCOM, sct, FS, FDMAX, FDMIN,
    & FMAX, FMIN, FCUM,GAMMAS,GAMMAU,BETA,C,PHI,maxcolht,
    & DEPTHP,HDISP,GELRAT,GDISP)

    if(mod(NSEIT(KKK),outsec).eq.0.and.NSEIT(KKK).ge.OUTSTART)
    & call doutput(outsec,nops,
    & option,NC,sltp,VATHET,ATHESP,sct,KP,NHFLUX,HFLUX,

```

```
1  
  
& VHC,FCUM,FMIN,FDMIN,FMAX,FDMAX,FS,APKT,KKK,NNLMAX,  
& NOTC,ATH1,VATH1,TTSEC,ATEMP,BTFLUX,KT,nslices,  
& slice,OUTTIM,DEPTH,COLHT,EXP,SOILP,HDISP,EXWAT,GDISP)  
  
do 888 NN=1,nops  
  if(option(NN).eq.35)then  
    if(mod(NSEIT(KKK),outsec).eq.0)call isotherm(NC, sltp,  
  & TTCOM,dcp,KKK,NNL,ATEMP,ISO,outsec,isocol,isoncol,  
  & nops,option)  
  endif  
888 continue  
  
600 continue  
  
DO l=1,NC  
  write(40,*)'total downslope displacement at surface'  
  write(40,*)'column ',l, (GDISP(l)+HDISP(l))*100,' cm'  
enddo  
c  
c  
c  
  stop  
end
```



```

c   Subroutine Cv.f
c
c   Calculates volumetric heat capacity of a soil cell given
c   intrinsic thermal properties of soil mineralogy, and relative
c   proportions of soil ice and water.
c
c   DEFINITIONS
c   Volumetric heat capacity (Cv) is the amount of heat required to change
c   the temperature of a unit volume of a substance by 1K.
c   Units = KJ / m-3 . K
c
c   For unfrozen soil
c    $C_{vu} = d(C_s + C_w W/100)$ 
c
c   For frozen soil
c    $C_{vf} = df(C_s + C_w W/100 + C_i I/100)$ 
c
c    $C_{vu}$  = volumetric heat capacity of unfrozen soil (J m-3 K-1)
c    $C_{vf}$  = volumetric heat capacity of frozen soil (m-3 K-1)
c    $d$  = dry bulk density of unfrozen soil (g m-3)
c    $df$  = dry bulk density of frozen soil (g m-3)
c    $C_s$  = specific heat capacity of dry mineral matter (J g-1 K-1)
c    $C_w$  = " " " " water (J g-1 K-1)
c    $C_i$  = " " " " ice (J g-1 K-1)
c    $W$  = water (volumetric %) and  $W/100$  is unit % ie. ATHESP
c    $I$  = ice (volumetric %) and  $I/100$  is unit % ie. VATHET
c
c
c   *****
c   subroutine Cv(NC,NNL,VHC,ATHESP,VATHET,TT,NNLMAX)
c   *****
c   DECLARATION OF VARIABLES
c   *****
c   DOUBLE PRECISION UDENS, FDENS, SHCSOIL, VHC(5,60)
c   DOUBLE PRECISION CW, CI, ATHESP(5,60), VATHET(5,60)
c
c   INTEGER NC, NNL(5), TT, NNLMAX
c
c   *****
c   OPEN STATEMENTS
c   *****
c   open(unit=59, file='DATAIN/Cvin.data', status='unknown')
c   rewind 59
c   *****
c   READ DATA INPUT FILE
c   *****
c   READ(59,*)UDENS, FDENS, SHCSOIL., CW, CI
c   *****
c
c   DYNAMIC SECTION
c   *****
c
c   Values of ice and water transferred from main program are already in
c   % volumetric terms.
c   Determine whether soil cell contains ice.
c   DO 1 I=1,NC
c   M=NNL(I)

```

DO 2 J=1,M

c Frozen soil. VHC units = J m⁻³ K⁻¹.

IF(VATHET(I,J).GT.0.0)VHC(I,J)=FDENS*
& (SHCSOIL+(CW*ATHESP(I,J))+(CI*VATHET(I,J)))

c Unfrozen soil. VHC units = J m⁻³ K⁻¹.

IF(VATHET(I,J).LE.0.0)VHC(I,J)=UDENS*(SHCSOIL+
& (CW*ATHESP(I,J)))

2 CONTINUE

1 CONTINUE

return

end

c **Subroutine freeze.f**

c *****
 c Program freeze.f
 c *****

c freeze is entered if the net heat flux for a cell is negative
 c as calculated in heatrans.f

subroutine freeze(ATHESP, TEMCV, WATECV, NC,
 & TF, sltp, sltpt2, CVGR, VHC, LA, TT, LENGTH, CDEPTH,
 & VATHET, BDT, ORES1, ORES2, ATHET, option, I, J, KKK,
 & ASR1, ASR2, CELVOL, VADFAC, SOILP)

c DECLARATION OF VARIABLES

c *****
 DOUBLE PRECISION ATHESP(5,60), ATHET(5,60)
 DOUBLE PRECISION ATHESP1(5,60), ATHET1(5,60)
 DOUBLE PRECISION VATHET(5,60), VATADD(5,60)
 DOUBLE PRECISION TEMCV(50), WATECV(50)
 DOUBLE PRECISION sltpt2(5,60), sltp(5,60), CVGR(50)
 DOUBLE PRECISION deltaT(5,60), deltaW(5,60), cmdt(5,60)
 DOUBLE PRECISION LAO(5,60), VHC(5,60), LA
 DOUBLE PRECISION CH(5,60), VADFAC(5,60), NEWVOL(5,60)
 DOUBLE PRECISION LENGTH, CDEPTH(5,60), BDT
 DOUBLE PRECISION ORES1(50), ORES2(50), MORES(50)
 DOUBLE PRECISION CHPERC(5,60), TEMPCH(5,60), SOILP(5,60)
 DOUBLE PRECISION ASR1, ASR2, ADFAC(5,60), CELVOL(5,60)

INTEGER NC, TF, TFMAX, option, I, J, KKK

ADFAC(I,J)=1.0

c *****
 c DYNAMIC SECTION
 c *****

c Freezing conditions in cell?
 IF((sltpt2(I,J)-273.15).LT.1.0e-10)GOTO 12
 IF(sltpt2(I,J).LE.273.15)GOTO 12
 IF((sltpt2(I,J)-273.15).GT.1.0e-10)THEN
 sltp(I,J)=sltpt2(I,J)
 GOTO 2
 ENDIF

c New soil temperature at end of iteration has been calculated by
 c heatrans.f, (sltpt2), using the O/T characteristic curve.
 c Establish corresponding soil moisture content for temperature
 c sltpt2.
 c The value returned from the O/T curve is adjusted by the volume
 c change factor for the cell. This means that the absolute volume of
 c water at a given temperature is the same whilst the volumetric
 c percentage will be decreased if expansion has taken place.

c Calculate dt deltaT, for iteration period. (K)
 12 IF(sltpt(I,J).GT.273.15)deltaT(I,J)=273.15-sltpt2(I,J)
 IF(sltpt(I,J).LE.273.15)deltaT(I,J)=sltp(I,J)-sltpt2(I,J)

```

IF(deltaT(I,J).EQ.0.0D-10)THEN
  GOTO 2
ENDIF
IF(sltpt2(I,J).GE.TEMCV(1))ATHET(I,J)=WATECV(1)
TFMAX=TF
IF(sltpt2(I,J).LE.TEMCV(TFMAX))ATHET(I,J)=WATECV(TFMAX)

```

```

DO 4 K=1,TF-1
IF(sltpt2(I,J).LE.TEMCV(K).AND.sltpt2(I,J).GT.TEMCV(K+1))
  & ATHET(I,J)=WATECV(K)+CVGR(K)*(sltpt2(I,J)-TEMCV(K))
4 CONTINUE
ATHET(I,J)=ATHET(I,J)/VADFAC(I,J)

```

c -----

```

c Calculate dW deltaW.
c Calculate mean residual moisture content and adjust for any volume
c change that has occurs based on the absolute residual moisture
c content being unchanged.
MORES(I)=((ORES1(I)+ORES2(I))/2)/VADFAC(I,J)

```

```

c Balance checks, ATHESP must not fall below residual moisture content
c adjusted for cell volume. The assumption is made that the residual
c moisture content for a cell is an absolute volume and not a unit
c percentage so it is adjusted for any change in volume.

```

```

IF(DABS(ATHESP(I,J)-MORES(I)).LT.1.0D-10)ATHESP(I,J)=MORES(I)
c .....no further freezing unless water migrates to cell.

```

```

IF(ATHESP(I,J).LT.MORES(I))ATHESP(I,J)=MORES(I)
c .....no further freezing unless water migrates to cell.

```

```

c Convert ATHESP from %volume to volume m3.
ATHESP1(I,J)=CELVOL(I,J)*ATHESP(I,J)

```

```

c Balance checks , ATHET must not fall below residual moisture content.
IF(DABS(ATHET(I,J)-MORES(I)).LT.1.0D-10)ATHET(I,J)=MORES(I)

```

```

IF(ATHET(I,J).LT.MORES(I))ATHET(I,J)=MORES(I)

```

```

c Convert ATHET from %volume to volume m3.
ATHET1(I,J)=CELVOL(I,J)*ATHET(I,J)

```

```

c d0, deltaW m3
deltaW(I,J)=ATHESP1(I,J)-ATHET1(I,J)

```

```

c *****
c cmdt, thermal component of energy budget. J (m-3 K-1).
cmdt(I,J)=((VHC(I,J)*CELVOL(I,J))*deltaT(I,J))

```

```

c LAO, hydrological component of energy budget. J (m-3).
LAO(I,J)=(LA*1000000.0)*deltaW(I,J)
c *****

```

c PHASE CHANGE LOGIC

```

c dQ = La d0 dt - cmdt
c dt

```

c *****

c If change in moisture content is negative, (given the predicted
 c temperature change and isothermal nature of phase change in previous
 c iteration), phase change cannot occur. Energy is therefore not
 c released as a hydrological response to transient thermal conditions
 c and thus has no effect on the temperature of the cell.

```
IF(deltaW(I,J).LE.0.0)THEN
c .....no further change in 0 or 1
  sltp(I,J)=sltp2(I,J)
  GOTO 2
ENDIF
```

```
.....
c
c LAO = cmdt
```

```
IF(DABS(LAO(I,J)-cmdt(I,J)).LT.1.0D-10)THEN
c .....isothermal phase change occurs.
```

```
c Temperature
c .....reset temperature to that before heat transfer, i.e. temp
c processes.
```

```
c Ice
c VATHET(I,J)=VATHET(I,J)+(ATHESP(I,J)-ATHET(I,J))
  VATADD(I,J)=(ATHESP(I,J)-ATHET(I,J))*1.09
  VATHET(I,J)=VATHET(I,J)+VATADD(I,J)
```

```
c Moisture
  ATHESP(I,J)=ATHET(I,J)
c .....where ATHET is defined by the characteristic curve for
c sltp2.
```

```
c determine if new volumetric total exceeds unity
  NEWVOL(I,J)=VATHET(I,J)+ATHESP(I,J)+SOILP(I,J)
  IF(NEWVOL(I,J).GT.1.0)THEN
    ADFAC(I,J)=NEWVOL(I,J)
    VATHET(I,J)=VATHET(I,J)/ADFAC(I,J)
    ATHESP(I,J)=ATHESP(I,J)/ADFAC(I,J)
    SOILP(I,J)=SOILP(I,J)/ADFAC(I,J)
    CDEPTH(I,J)=CDEPTH(I,J)*ADFAC(I,J)
    CELVOL(I,J)=CELVOL(I,J)*ADFAC(I,J)
  endif
```

```
GOTO 2
ENDIF
```

```
c COMMENT
c Moisture content is set to ATHET on characteristic curve as defined
c by sltp2, however after phase change release of energy, the temp.
c returns to sltp. Therefore at the beginning of the next time step,
c soil water and temperature are not in accordance with the
c characteristic curve. ie. the soil moisture content is less than
c normally expected given the cell temperature. 0 in next iteration
c is set to ATHET->ATHESP.
```

```
.....
c
c LAO<cmdt
c NEW IF STATEMENT
  IF(deltaW(I,J).GE.0.0)THEN
```

```

c   IF(LAO(I,J).LT.cmdt(I,J))THEN
c
c   .....no isothermal phase change, cell cools. Energy released by
c   cooling of cell (thermal component), is partially compensated by
c   phase change energy release. Temperature is not maintained at
c   273.15 K, rather the cell cools by an amount equivalent to LAO given
c   the volumetric heat capacity of the soil mixture.
c   Temperature
c   Calculate the temperature change expected by an input of 1J.
c   TEMPCH(I,J)=1.0/VHC(I,J)
c
c   Temperature change given the cell volume.
c   CHPERC(I,J)=TEMPCH(I,J)/CELVOL(I,J)
c   CH(I,J)=LAO(I,J)*CHPERC(I,J)
c   sltp(I,J)=sltp2(I,J)+CH(I,J)
c
c   Isothermal ???
c   if(sltp(I,J).gt.273.15)sltp(I,J)=273.15
c
c   VATADD(I,J)=(ATHESP(I,J)-ATHET(I,J))*1.09
c   VATHET(I,J)=VATHET(I,J)+VATADD(I,J)
c   VATHET(I,J)=VATHET(I,J)+(ATHESP(I,J)-ATHET(I,J))
c   ATHESP(I,J)=ATHET(I,J)
c
c   determine if new volumetric total exceeds unity
c   NEWVOL(I,J)=VATHET(I,J)+ATHESP(I,J)+SOILP(I,J)
c   if(NEWVOL(I,J).GT.1.0)THEN
c   ADFAC(i,j)=NEWVOL(I,J)
c   VATHET(I,J)=VATHET(I,J)/ADFAC(I,J)
c   ATHESP(I,J)=ATHESP(I,J)/ADFAC(I,J)
c   SOILP(I,J)=SOILP(I,J)/ADFAC(I,J)
c   CDEPTH(I,J)=CDEPTH(I,J)*ADFAC(I,J)
c   CELVOL(I,J)=CELVOL(I,J)*ADFAC(I,J)
c   endif
c
c   ENDIF
c   .....
2   return
c   end

```

c **Subroutine heatrans.f**

c Subroutine calculates flux of heat energy in one dimension on basis
 c of thermal conductivity of medium and temperature gradient.
 c Temperature of cells are updated given the volumetric heat capacity
 c of the medium, modified on each iteration for changing ice and water
 c content.

$$C \frac{dT}{dt} = -d \left(k \frac{dT}{dx} \right)$$

c where

c C = volumetric heat capacity J m⁻³ K⁻¹
 c x = depth m
 c t = time sec
 c k = thermal conductivity Jsec m⁻¹ K⁻¹
 c T = temperature K

c and

$$1 \text{ W} = 1 \text{ J sec}$$

c Subroutine contains time element for thermal conductivity
 c Kt. W m⁻¹ K⁻¹ where 1W = 1 J sec
 c i.e. multiply thermal conductivity entered in input file
 c by iteration period in seconds.

c *****
 c subroutine heatrans(ATEMP, CDEPTH, sltp, sltpt2,
 c & VATHET, NC, NNL, VHC, LENGTH, option,
 c & ATHESP, NNLMAX, BDT, TEMCV, WATECV, TF, LA,
 c & ORES1, ORES2, CVGR, ATHET, HFLUX, NHFLUX, APKT,
 c & STAMBT, IP, KKK, LC1, FC2, ASR1, ASR2,
 c & SHVATHET, CELVOL, ORIVOL, VADFAC, SOILP, SOIL1, TDEP, BTEMP, EXWAT)

c *****
 c DECLARATION OF VARIABLES
 c *****

DOUBLE PRECISION ATEMP(1468800), sltp(5,60), sltpt2(5,60)
 DOUBLE PRECISION CDEPTH(5,60), VATHET(5,60), AVKT(5,60)
 DOUBLE PRECISION NHFLUX(5,60), BTFLUX(50), HFLUX(5,60)
 DOUBLE PRECISION LENGTH, TCHA(5,60), CELVOL(5,60)
 DOUBLE PRECISION ATHESP(5,60), ATHET(5,60), BDT
 DOUBLE PRECISION VHC(5,60), WATTS(5,60)
 DOUBLE PRECISION TEMCV(50), WATECV(50), LA, CVGR(50), TDEP
 DOUBLE PRECISION ORES1(50), ORES2(50), EXWAT(5,60)
 DOUBLE PRECISION APKT(5,60), STAMBT, BTEMP
 DOUBLE PRECISION ASR1, ASR2, SOILP(5,60), SOIL1(5,60)
 DOUBLE PRECISION SHVATHET(5,60), VADFAC(5,60), ORIVOL(5,60)

INTEGER NC, NNL(5), option, NNLMAX
 INTEGER LC1(50), FC2(50)
 INTEGER TF, I, J, IP, KKK

c *****

```

c   DYNAMIC SECTION
c   .....

c   Call subroutine Cv.f for calculation of volumetric heat capacities
c   in each cell of matrix.
c   call Cv(NC,NNL,VHC,ATHESP,VATHET,TT,NNLMAX)

      DO 666 I=1,NC
      M= NNL(I)
      DO 667 J=1,M
      SHVATHET(I,J)=VATHET(I,J)
c   calculate volume adjustment factor (VADFAC) before going through
c   heatrans and on to thaw or freeze
      VADFAC(I,J)=CELVOL(I,J)/ORIVOL(I,J)
667   CONTINUE
666   CONTINUE

c   SURFACE BOUNDARY CONDITIONS

c   Mean thermal conductivity across surface interface.
      DO 1 I=1,NC
      J=1
      AVKT(I,J)=APKT(I,J)

c   Surface boundary heat flux.
c   Divide by 0.5 x cell depth.

      HFLUX(I,J)=((AVKT(I,J)*IP)*((ATEMP(KKK)-sltp(I,J))/
& (0.5*CDEPTH(I,J))))*(LENGTH*BDT)
1     CONTINUE

c   PROFILE HEAT FLUX
c   Mean thermal conductivities across cell boundaries.
      DO 2 I=1,NC
      M=NNL(I)
      DO 3 J=2,M
      AVKT(I,J)=((APKT(I,J)*CDEPTH(I,J)+(APKT(I,J-1)*
& CDEPTH(I,J-1)))/(CDEPTH(I,J)+CDEPTH(I,J-1)))

c   Heat flux across cell boundaries. Jsec cm-3 -K.
      HFLUX(I,J)=((AVKT(I,J)*IP)*((sltp(I,J-1)-sltp(I,J))/
& ((CDEPTH(I,J)+CDEPTH(I,J-1))/2)))*(LENGTH*BDT)

3     CONTINUE
2     CONTINUE

c   Basal heat flux. Jsec m-3 -K
      DO 4 I=1,NC
      J=NNL(I)

c   btemp set as input basal temperature
c   also possible to set bflux to zero if do not want basal heat supply
      bflux(i)=((avkt(i,j)*ip)*((sltp(i,j)-btemp)/cdepth(i,j)))*
& (length*bdt)
4     CONTINUE

c   Net heat flux across all cell boundaries. J cm-3 -K.

```



```

DO 5 I=1,NC
M=NNL(I)
DO 6 J=1,M

c   Net flux of heat across all boudaries
IF(J.NE.NNL(I))NHFLUX(I,J)=HFLUX(I,J)-HFLUX(I,J+1)

c   Basal cell net heat flux.
IF(J.EQ.NNL(I))NHFLUX(I,J)=HFLUX(I,J)-BTFLUX(I)

6   CONTINUE
5   CONTINUE

c   Temperature change of all cells in matrix
DO 7 I=1,NC
M=NNL(I)
DO 8 J=1,M

c   Calculate Joules required to change temperature by 1K
WATTS(I,J)=1.0/VHC(I,J)

c   Temperature change.
TCHA(I,J)=(WATTS(I,J)/(LENGTH*CDEPTH(I,J)*BDT))*NHFLUX(I,J)

c   New soil temperature.
sltp2(I,J)=sltp(I,J)+TCHA(I,J)

8   CONTINUE
7   CONTINUE

c   *****
208 DO 888 I=1,NC
M=NNL(I)
DO 999 J=1,M

IF(NHFLUX(I,J).GT.0.0)THEN
call thaw(sltpt2, TF, LENGTH,
& CDEPTH, BDT, ATHESP, VATHET, VHC, sltp, NC, LA,
& NHFLUX, I, J, KKK, CELVOL, VADFAC, SOILP, ORIVOL, SOIL1,
& ASR1, ASR2, LC1, FC2, TDEP, EXWAT)
ENDIF

c
IF(NHFLUX(I,J).LE.0.0)THEN
call freeze(ATHESP, TEMCV, WATECV, NC, TF, sltp,
& sltp2, CVGR, VHC, LA, TT, LENGTH, CDEPTH, VATHET,
& BDT, ORES1, ORES2, ATHET, option, I, J, KKK,
& ASR1, ASR2, CELVOL, VADFAC, SOILP)
ENDIF

999 CONTINUE
888 CONTINUE

c   *****
return
end

```

```

c      subroutine isotherm.f
c
c      .....
c      This subroutine defines and writes out the depth of the 273.15K
c      isotherm to a file which may then be converted into a graphical
c      plot.
c      .....
c      subroutine isotherm(NC, sltp, TTCOM, dcp,
&      KKK,>NNL, ATEMP,ISO,outsec,isocol,isoncol,
&      nops,option)
c
c      DECLARATION OF VARIABLES
c
c      DOUBLE PRECISION sltp(5,60), ATEMP(1468800)
c      DOUBLE PRECISION TTCOM(5,60), dcp(5,60)
c      DOUBLE PRECISION ISO(5,60), DISO(5,60)
c      DOUBLE PRECISION SISO(5,60),TCH(5,60),TG(5,60)
c      DOUBLE PRECISION TISO(525600), BISO(525600)
c
c
c      INTEGER KKK, NC,>NNL(5)
c      INTEGER outsec, MM
c      INTEGER isoncol, isocol(50)
c      INTEGER nops, option(50)
c
c      .....
c      SURFACE CELLS
c      Does 273.15K isotherm lie between surface and first compn. point?
c      DO 1 I=1,isoncol
c
c      M>NNL(I)
c      DO 2 J=1,M
c
c          II=isocol(I)
c
c          if(J.eq.1)then
c              if(sltp(II,J).le.273.15.and.ATEMP(KKK).gt.273.15)then
c
c      Temperature gradient
c      TG(II,J)=(ATEMP(KKK)-sltp(II,J))/dcp(II,j)
c
c      What is the temperature difference between surface and 273.15K?
c      TCH(II,J)=ATEMP(KKK)-273.15
c
c      At what depth from the surface does the isotherm occur?
c      ISO(II,J)=TCH(II,J)/TG(II,J)
c
c          if(ISO(II,J).le.0.0)goto 140
c
c      Write isotherm data to files with suffix 300-399
c
c      do 506 NN=1,nops
c          if(option(NN).eq.35)then
c      MM=(II+300)

```

```

write(MM,*)ISO(I,J),KKK
endif
506 continue
140 endif

c -----
if(sltp(I,J).gt.273.15.and.ATEMP(KKK).le.273.15)then
TG(I,J)=(sltp(I,J)-ATEMP(KKK))/dcp(I,J)
TCH(I,J)=273.15-ATEMP(KKK)
SISO(I,J)=TCH(I,J)/TG(I,J)
ISO(I,J)=SISO(I,J)

if(ISO(I,J).le.0.0)goto 150

c Write isotherm data to files with suffix 300-400
do 501 NN=1,nops
if(option(NN).eq.35)then
MM=(I+300)
write(MM,*)ISO(I,J),KKK
endif
501 continue
150 endif

endif

c .....
c PROFILE CELLS

if(J.gt.1)then

c Does 273.15K isotherm lie between computation points of cells
c J-1 and J.
if(sltp(I,J).le.273.15.and.sltp(I,J-1).gt.273.15)then

c Temperature gradient
TG(I,J)=(sltp(I,J-1)-sltp(I,J))/(dcp(I,J)+dcp(I,J-1))

c What is the temperature difference between cell J-1 and 273.15?
TCH(I,J)=sltp(I,J-1)-273.15

c At what depth from compn. point of cell J-1 does 273.15 isotherm
c occur?
DISO(I,J)=TCH(I,J)/TG(I,J)

c Depth of isotherm from surface.
ISO(I,J)=TTCOM(I,J-1)+DISO(I,J)

if(ISO(I,J).ne.0.0)goto 160

do 502 NN=1,nops

```

```
        if(option(NN).eq.35)then
          MM=(II+300)
          write(MM,*)ISO(II,J),KKK
        endif
502      continue

160      endif
c      -----
      if(sltp(II,J).gt.273.15.and.sltp(II,J-1).le.273.15)then
        TG(II,J)=(sltp(II,J)-sltp(II,J-1))/(dcp(II,J)+dcp(II,J-1))
        TCH(II,J)=273.15-sltp(II,J-1)
        DISO(II,J)=TCH(II,J)/TG(II,J)
        ISO(II,J)=TTCOM(II,J)+DISO(II,J)

        if(ISO(II,J).le.0.0)goto 170

        do 503 NN=1,nops
          if(option(NN).eq.35)then
            MM=(II+300)
            write(MM,*)ISO(II,J),KKK
          endif
503      continue

170      endif

      endif

2      continue
1      continue

      return
      end
```

c **Subroutine kta.f**

c Sub routine kta.f calculates temperature dependent thermal
c conductivity ($W\ m^{-1}\ K^{-1}$) for each cell of profile. The system is
c assumed to comprise three phases, soil solids, water and ice.

```
subroutine kta(VATHET, ATHESP, CELVOL,
&  sltp, NNLMAX, NC, NNL, ASR1, ASR2, LA,
&  APKT, KT, KI, NKI, KO, NKO, TKI, TKO, KS, OKG, IKG, SOILP)
```

```
c     *****
c     DECLARATION OF VARIABLES
c     *****
DOUBLE PRECISION KI(20), TKI(20), KO(20), TKO(20)
DOUBLE PRECISION VATHET(5,60), ATHESP(5,60), CELVOL(5,60)
DOUBLE PRECISION KS(5,60), sltp(5,60), SOILV(5,60)
DOUBLE PRECISION KOK(5,60), KIK(5,60), ASR1, ASR2
DOUBLE PRECISION VATH1(5,60), ATH1(5,60), SOILP(5,60)
DOUBLE PRECISION SFI(5,60), SFO, SFS(5,60)
DOUBLE PRECISION KT(5,60), APKT(5,60), LA, IKG(20), OKG(20)
```

```
INTEGER NKI, NKO, NC, NNL(5), NNLMAX, NKOMAX, NKIMAX
```

```
c     *****
c     DYNAMIC SECTION
c     *****
c     Derive volume of soil solids per cell
DO 10 I=1,NC
  M=NNL(I)
  DO 20 J=1,M
    SOILV(I,J)=CELVOL(I,J)*SOILP(I,J)

c     Derive volume of ice per cell
  VATH1(I,J)=CELVOL(I,J)*VATHET(I,J)

c     Derive volume of water per cell
  ATH1(I,J)=CELVOL(I,J)*ATHESP(I,J)

20     continue
10     continue
```

```
c     Extrapolation from curve
c     Water
DO 40 I=1,NC
  M=NNL(I)
  DO 41 J=1,M
    if(sltp(I,J).GE.TKO(1))KOK(I,J)=KO(1)
    NKOMAX=NKO
    if(sltp(I,J).LE.TKO(NKOMAX))KOK(I,J)=KO(NKOMAX)
    do 4 K=1,NKO-1
      if(sltp(I,J).LE.TKO(K).AND.sltp(I,J).GT.TKO(K+1))
        & KOK(I,J)=KO(K)+OKG(K)*(sltp(I,J)-TKO(K))
4     continue
41     continue
40     continue
```

```

c   Ice
DO 60 I=1,NC
M=NNL(I)
DO 70 J=1,M
  if(sltp(I,J).GE.TKI(1))KIK(I,J)=KI(1)
  NKIMAX=NKI
  if(sltp(I,J).LE.TKI(NKIMAX))KIK(I,J)=KI(NKIMAX)

  do 5 K=1,NKI-1
    if(sltp(I,J).LE.TKI(K).AND.sltp(I,J).GT.TKI(K+1))
      & KIK(I,J)=KI(K)+IKG(K)*(sltp(I,J)-TKI(K))
5   continue

70  continue
60  continue
c   Derive shape factor for water
SFO=1.0D00

c   Derive shape factor for ice
DO 50 I=1,NC
M=NNL(I)
DO 51 J=1,M
  SFI(I,J)=(1.0D00/3.0D00)
  & *((1.0D00/(1.0D00+((KIK(I,J)/KOK(I,J))-1.0D00)*0.125))
  & +(1.0D00/(1.0D00+((KIK(I,J)/KOK(I,J))-1.0D00)*0.125))
  & +(1.0D00/(1.0D00+((KIK(I,J)/KOK(I,J))-1.0D00)*0.750)))

c   Derive shape factor for soil solids
SFS(I,J)=(1.0D00/3.0D00)
  & *((1.0D00/(1.0D00+((KS(I,J)/KOK(I,J))-1.0D00)*0.125))
  & +(1.0D00/(1.0D00+((KS(I,J)/KOK(I,J))-1.0D00)*0.125))
  & +(1.0D00/(1.0D00+((KS(I,J)/KOK(I,J))-1.0D00)*0.750)))

c   Calculate thermal conductivity of the cell W m-1 K-1
KT(I,J)=((ATH1(I,J)*KOK(I,J)*SFO)
  & +(VATH1(I,J)*KIK(I,J)*SFI(I,J))
  & +(SOILV(I,J)*KS(I,J)*SFS(I,J)))
  & /((ATH1(I,J)*SFO)+(VATH1(I,J)*SFI(I,J))
  & +(SOILV(I,J)*SFS(I,J)))

  APKT(I,J)=KT(I,J)

51  continue
50  continue

  return
  end

```

c **Subroutine ku.f**

c Subroutine ku.f calculates hydraulic conductivity
c for all cells in the profile

```
subroutine ku(NC, LC1, FC2, KQ, ATHESP,
& X1, X2,Z1, Z2, GZ1, GZ2, KP)
```

c Z1 Z2.....curve gradients on suction/moisture curve.

```
DOUBLE PRECISION ATHESP(5,60),X1(50),X2(50)
DOUBLE PRECISION Z1(50),Z2(50),GZ1(50),GZ2(50),KP(5,60)
```

```
INTEGER NC, KQ, LC1(50), FC2(50)
```

```
c *****
c DYNAMIC SECTION
c *****
```

```
DO 1 I=1,NC
M=LC1(I)
DO 2 J=1,M
if(KQ.gt.20)write(428,*)KQ,athesp(i,j),i,j,kp(i,j)
DO 3 K=1,KQ-1
IF(ATHESP(I,J).GT.X1(K).AND.ATHESP(I,J).LE.X1(K+1))
& KP(I,J)=Z1(K)+GZ1(K)*(ATHESP(I,J)-X1(K))
IF(ATHESP(I,J).LE.X1(1))KP(I,J)=Z1(1)
IF(ATHESP(I,J).GT.X1(20))KP(I,J)=Z1(20)
3 CONTINUE
2 CONTINUE
1 CONTINUE
```

```
DO 4 I=1,NC
M=LC1(I)
N=FC2(I)
DO 5 J=M,N
DO 6 K=1,KQ-1
IF(ATHESP(I,J).GT.X2(K).AND.ATHESP(I,J).LE.X2(K+1))
& KP(I,J)=Z2(K)+GZ2(K)*(ATHESP(I,J)-X2(K))
IF(ATHESP(I,J).LE.X2(1))KP(I,J)=Z2(1)
IF(ATHESP(I,J).GT.X2(20))KP(I,J)=Z2(20)
6 CONTINUE
5 CONTINUE
4 CONTINUE
```

```
return
end
```

c **Subroutine masstrans.f**

c Subroutine masstrans.f, is used to reasses the distribution of
c water in the soil profile after freezing or thawing have occurred,
c on the basis of Darcy's Law.

c subroutine masstrans(NC,LC1,NQ1,ATHESP,X1,Y1,G1,FC2,NNL,
& Y2,G2,HT,KP,KQ,Z1,GZ1,Z2,GZ2,CDEPTH,LENGTH,
& sct,HPOT,AVK,FLUX,NFLUX,HYDG,IP,BDT,STCON,dcp,
& BDT,X2,NQ2,STCON1,STCON2,EXSAT,
& BWFUX,HTHETAC,SHATHET,VATHET,COLHT,CDEP,EXP,
& VADFAC,MAXCOLHT)

c DOUBLE PRECISION SHATHET(5,60),ASR1,ASR2,ATHESP(5,60)
c DOUBLE PRECISION ATHESPT(5,60),VATHET(5,60)
c DOUBLE PRECISION STCON1,STCON2
c DOUBLE PRECISION Y1(50),Y2(50),G1(50),G2(50),X1(50),X2(50)
c DOUBLE PRECISION sct(5,60),HPOT(5,60),CDEP(5,60),EXP(5,60)
c DOUBLE PRECISION HT(5,60),CDEPTH(5,60),LENGTH,BDT
c DOUBLE PRECISION AVK(5,60),KP(5,60),GZ1(50),GZ2(50)
c DOUBLE PRECISION FLUX(5,60),NFLUX(5,60),BWFUX(5,
c DOUBLE PRECISION COLHT(5),CUMHT,MAXCOLHT(5)
c DOUBLE PRECISION Z2(50),Z1(50),vadfac(5,60)

c INTEGER NC,LC1(50),FC2(50),NQ1,NQ2,KQ,NNL(5),IP

c UNSATURATED FLOW

c Soil water pressure (suction)

c Layer 1

DO 20 I=1,NC

M=LC1(I)

DO 30 J=1,M

c is saturated set suction to zero

IF (ATHESP(I,J).GE.ASR1)THEN

SCT(I,J)=0.0

ELSEIF(ATHESP(I,J).LT.X1(1))THEN

SCT(I,J)=Y1(1)

ELSE

K=NQ1-1

DO 40 L=1,K

IF(ATHESP(I,J).GE.X1(L).AND.ATHESP(I,J).LT.X1(L+1))THEN

sct(I,J)=Y1(L)+G1(L)*(ATHESP(I,J)-X1(L))

ENDIF

40 CONTINUE

ENDIF

30 CONTINUE

20 CONTINUE

c

c

c

c

Layer 2

DO 50 I=1,NC

M=FC2(I)

N=NNL(I)

DO 60 J=M,N

IF (ATHESP(I,J).GE.ASR2)THEN

SCT(I,J)=0.0


```

ELSEIF(ATHESP(I,J).LT.X2(1))THEN
SCT(I,J)=Y2(1)
ELSE
K=NQ2-1
DO 70 L=1,K
IF(ATHESP(I,J).GE.X2(L).AND.ATHESP(I,J).LT.X2(L+1))THEN
sct(I,J)=Y2(L)+G2(L)*(ATHESP(I,J)-X2(L))
ENDIF
70 CONTINUE
ENDIF
60 CONTINUE
50 CONTINUE
c
c
c calculate cell mid-point heights, total column height
c and volume expansion of each cell
DO 151 I=1,NC
M=NNL(I)
CUMHT=0.0
COLHT(I)=0.0
DO 152 J=M,1,-1
c find height of midpoint of cell
CUMHT=CUMHT+(CDEPTH(I,J)/2.0)
c cell expansion calculated as value in mm
EXP(I,J)=(CDEPTH(I,J)-CDEP(I,J))*1000
HT(I,J)=CUMHT
c convert CUMHT to height of column so far
CUMHT=CUMHT+(CDEPTH(I,J)/2.0)
COLHT(I)=COLHT(I)+CDEPTH(I,J)
152 CONTINUE
IF(COLHT(I).GT.MAXCOLHT(I))MAXCOLHT(I)=COLHT(I)
151 CONTINUE

c Hydraulic potential
DO 80 I=1,NC
DO 90 J=1,NNL(I)
HPOT(I,J)=sct(I,J)+HT(I,J)
90 CONTINUE
80 CONTINUE
c
c
c Hydraulic conductivity
c Layer 1
DO 100 I=1,NC
M=LC1(I)
DO 110 J=1,M
IF(ATHESP(I,J).GE.ASR1) THEN
KP(I,J)=STCON1
ELSEIF(ATHESP(I,J).LT.X1(1))THEN
KP(I,J)=Z1(1)
ELSE
DO 120 K=1,KQ-1
IF(ATHESP(I,J).GE.X1(K).AND.ATHESP(I,J).LT.X1(K+1))
& KP(I,J)=Z1(K)+GZ1(K)*(ATHESP(I,J)-X1(K))
120 CONTINUE
ENDIF
110 CONTINUE
100 CONTINUE

```

```

c
c
c   Layer 2
      DO 130 I=1,NC
      M=FC2(I)
      N=NNL(I)
      DO 140 J=M,N
      IF(ATHESP(I,J).GE.ASR2)THEN
      KP(I,J)=STCON2
      ELSEIF(ATHESP(I,J).LT.X2(1))THEN
      KP(I,J)=Z2(1)
      ELSE
      DO 150 K=1,KQ-1
      IF(ATHESP(I,J).GE.X2(K).AND.ATHESP(I,J).LT.X2(K+1))
      & KP(I,J)=Z2(K)+GZ2(K)*(ATHESP(I,J)-X2(K))
150   CONTINUE
      ENDIF
140   CONTINUE
130   CONTINUE
c
c
c   Mean hydraulic conductivity for surface cells
      DO 1401 I=1,NC
      J=1
      AVK(I,J)=KP(I,J)
c   write(6,*) 'kp',KP(I,J),'avk', AVK(I,J)
1401  CONTINUE

c   Mean hydraulic conductivity for profile cells
      DO 160 I=1,NC
      DO 170 J=2,NNL(I)
c   if freezing has started use upper cell conductivity as average
c   conductivity as a guide to ability to draw up water
      IF(VATHET(I,J-1).GT.0.0)THEN
      AVK(I,J)=KP(I J-1)
      ELSE
      AVK(I,J)=((KP(I,J-1)*CDEPTH(I,J-1))+KP(I,J)*CDEPTH(I,J))/
      & (CDEPTH(I,J-1)+CDEPTH(I,J))
      ENDIF
170   CONTINUE
160   CONTINUE
c
c
c   Convert hydraulic conductivity, m sec-1 into m IP-1
      DO 600 I=1,NC
      M=NNL(I)
      DO 601 J=1,M
      AVK(I,J)=AVK(I,J)*IP
601   CONTINUE
600   CONTINUE

c   Calculate flux across surface boundary
      DO 302 I=1,NC
      J=1
      FLUX(I,J)=0.0
302   CONTINUE

```

```

c   Calculate flux between cells, (vol m-3 hr)
DO 180 I=1,NC
M=NNL(I)
DO 190 J=2,M
FLUX(I,J)=(((HPOT(I,J-1)-HPOT(I,J))*AVK(I,J))/((CDEPTH(I,J-1)
& +CDEPTH(I,J)/2))*(LENGTH*BDT)
c   approach to stemming downward flux if cell below is
c   saturated with water or water+ice.
if(flux(i,j).gt.0.0.and.vadfac(i,j+1).gt.1.0)flux(i,j)=0.0
190  CONTINUE
180  CONTINUE
c
c
c   Basal water flux
DO 303 I=1,NC
J=NNL(I)
BWFLUX(I)=(AVK(I,J)*(0.5*CDEPTH(I,J)))*(LENGTH*BDT)
303  CONTINUE

c   Calculate net flux for each cell
DO 200 I=1,NC
M=NNL(I)
DO 210 J=1,M

c   Profile net water flux.
if(J.NE.NNL(I))NFLUX(I,J)=FLUX(I,J)-FLUX(I,J+1)
c   Basal net water flux.
if(J.EQ.NNL(I))NFLUX(I,J)=FLUX(I,J)-BWFLUX(I)
210  CONTINUE
200  CONTINUE
c
c
c   Calculate new soil moisture conditions
DO 220 I=1,NC
DO 230 J=1,NNL(I)
ATHESP(I,J)=ATHESP(I,J)+(NFLUX(I,J)/(CDEPTH(I,J)*LENGTH
& *BDT))
ATHESP(I,J)=ATHESP(I,J)
230  CONTINUE
220  CONTINUE
c
c
DO 240 I=1,NC
M=NNL(I)
DO 241 J=1,M
SHATHET(I,J)=ATHESP(I,J)
241  CONTINUE
240  CONTINUE

return
end

```

c **Subroutine output.f**

c Subroutine OUTPUT.F writes data to files according to
c selection of option(I).

```

subroutine doutput(outsec, nops, option, NC,
&  sltp, VATHET, ATHESP, sct, KP, NHFLUX, HFLUX, VHC,
&  FCUM, FMIN, FDMIN, FMAX, FDMAX, FS, APKT, KKK,NNLMAX,
&  NOTC, ATH1, VATH1, TTSEC, ATEMP, BTFLUX, KT, nslices,
&  slice, OUTTIM, DEPTH, COLHT, EXP, SOILP, HDISP, EXWAT, GDISP)

```

c
c **DECLARATION OF VARIABLES**
c

```

DOUBLE PRECISION sltp(5,60), VATHET(5,60), ATHESP(5,60)
DOUBLE PRECISION sct(5,60), KP(5,60), NHFLUX(5,60)
DOUBLE PRECISION HFLUX(5,60), VHC(5,60), SOILP(5,60)
DOUBLE PRECISION FCUM(50), FMIN(50), FDMIN(50)
DOUBLE PRECISION FMAX(50), FDMAX(50), FS(5,60)
DOUBLE PRECISION APKT(5,60)
DOUBLE PRECISION ATH1(5,60), VATH1(5,60), KT(5,60)
DOUBLE PRECISION BTFLUX(50), DEPTH(5), COLHT(5), EXP(5,60)
DOUBLE PRECISION ATEMP(1468800), HDISP(5), GDISP(5), EXWAT(5,60)

```

```

INTEGER NC, outsec, nops, option(50)
INTEGER KKK,>NNLMAX
INTEGER TTSEC(525600)
INTEGER nslices, slice(50), OUTTIM
COMMON DUMMY(5,60)

```

c

```

IF(OUTTIM.EQ.1)THEN
do 100 NN=1, nops
if(option(NN).eq.1)then
open(unit=81, file='sltp.data', status='unknown')
rewind 81
endif

if(option(NN).eq.2)then
open(unit=82, file='ice.data', status='unknown')
rewind 82
endif

if(option(NN).eq.3)then
open(unit=83, file='water.data', status='unknown')
rewind 83
endif

if(option(NN).eq.4)then
open(unit=84, file='nhflux.data', status='unknown')
rewind 84
endif

if(option(NN).eq.5)then
open(unit=85, file='hflux.data', status='unknown')
rewind 85
endif

```

```
if(option(NN).eq.6)then
open(unit=86, file='heatcap.data', status='unknown')
rewind 86
endif
```

```
if(option(NN).eq.7)then
open(unit=87, file='thermcon.data', status='unknown')
rewind 87
endif
```

```
if(option(NN).eq.8)then
open(unit=88, file='hydcon.data', status='unknown')
rewind 88
endif
```

```
if(option(NN).eq.9)then
open(unit=89, file='suction.data', status='unknown')
rewind 89
endif
```

```
if(option(NN).eq.10)then
open(unit=90, file='fcum.data', status='unknown')
rewind 90
endif
```

```
if(option(NN).eq.11)then
open(unit=91, file='fmax.data', status='unknown')
rewind 91
endif
```

```
if(option(NN).eq.12)then
open(unit=92, file='fmin.data', status='unknown')
rewind 92
endif
```

```
if(option(NN).eq.13)then
open(unit=93, file='fs.data', status='unknown')
rewind 93
endif
```

```
if(option(NN).eq.14)then
open(unit=94, file='ath1.data', status='unknown')
rewind 94
endif
```

```
if(option(NN).eq.15)then
open(unit=95, file='vath1.data', status='unknown')
rewind 95
endif
```

```
if(option(NN).eq.16)then
open(unit=96, file='atemp.data', status='unknown')
rewind 96
endif
```

```
if(option(NN).eq.17)then
open(unit=97, file='btflux.data', status='unknown')
```

```
rewind 97
endif

if(option(NN).eq.20)then
open(unit=70, file='Gsltp.data', status='unknown')
rewind 70
endif

if(option(NN).eq.21)then
open(unit=71, file='Gice.data', status='unknown')
rewind 71
endif

if(option(NN).eq.22)then
open(unit=72, file='Gwater.data', status='unknown')
rewind 72
endif

if(option(NN).eq.23)then
open(unit=73, file='Gsuction.data', status='unknown')
rewind 73
endif

if(option(NN).eq.24)then
open(unit=74, file='Ghycond.data', status='unknown')
rewind 74
endif

if(option(NN).eq.25)then
open(unit=75, file='gtemp.data', status='unknown')
rewind 75
endif

if(option(NN).eq.26)then
open(unit=76, file='Gkt.data', status='unknown')
rewind 76
endif

if(option(NN).eq.27)then
open(unit=77, file='Gapkt.data', status='unknown')
rewind 77
endif

if(option(NN).eq.28)then
open(unit=78, file='Gfs.data', status='unknown')
rewind 78
endif

if(option(NN).eq.29)then
open(unit=79, file='Gnhflux.data', status='unknown')
rewind 79
endif

if(option(NN).eq.30)then
open(unit=80, file='Ghflux.data', status='unknown')
rewind 80
endif
```

```

if(option(NN).eq.33)then
open(unit=33,file='Gexp.data',status='unknown')
rewind 33
endif

if(option(NN).eq.34)then
open(unit=34,file='Gcelexp.data',status='unknown')
rewind 34
endif

if(option(NN).eq.36)then
open(unit=36,file='Gsoilp.data',status='unknown')
rewind 36
endif

if(option(NN).eq.37)then
open(unit=37,file='Gdisplace.data',status='unknown')
rewind 37
endif

if(option(NN).eq.38)then
open(unit=38,file='Gexwat.data',status='unknown')
rewind 38
endif

100  continue
      OUTTIM=OUTTIM+1
      endif

      do 103 NN=1,NOPS

c      OUTPUT OPTIONS

c      *** SLTP ***
      if(option(NN).eq.1)then
11      write(81,11)KKK
      &  format('//Soil temperature, sltp. K',/,
      &  'Iteration = ',i8)

      call sieve(sltp,81,NC,NNLMAX)

      endif

c      -----
c      *** ICE ***
c      if(option(NN).eq.2)then

13      write(82,13)KKK
      &  format('//Soil ice content m3 m-3',/,
      &  'Iteration = ',i8)

      call sieve(VATHET,82,nc,nnlmax)

      endif

c      -----
c      *** WATER ***
c      if(option(NN).eq.3)then

      write(83,14)KKK

```

```

14   format(//Soil moisture content m3 m-3',/,
    &   'Iteration = ',i8)

      call sieve(ATHESP,83,NC,nnlmax)

      endif
c   -----
c   *** NET HEAT FLUX ***
      if(option(NN).eq.4)then

        write(84,16)KKK
16   format(//Net heat flux W m-3 K-1',/,
    &   'Iteration = ',i8)

        call sieve(NHFLUX,84,nc,nnlmax)

        endif
c   -----
c   *** HEAT FLUX ***
      if(option(NN).eq.5)then

        write(85,18)KKK
18   format(//Heat flux W m-3 K-1',/,
    &   'Iteration = ',i8)

        call sieve(HFLUX,85,nc,nnlmax)

        endif
c   -----
c   *** VOLUMETRIC HEAT CAPACITY ***
      if(option(NN).eq.6)then

        write(86,20)KKK
20   format(//Volumetric heat capacity J m-3 K-1',/,
    &   'Iteration = ',i8)

        call sieve(VHC,86,nc,nnlmax)

        endif
c   -----
c   *** THERMAL CONDUCTIVITY ***
      if(option(NN).eq.7)then

        write(87,22)KKK
22   format(//Thermal conductivity W m-3 sec-1 K-1',/,
    &   'Iteration = ',i8)

        call sieve(APKT,87,nc,nnlmax)

        endif
c   -----
c   *** HYDRAULIC CONDUCTIVITY ***
      if(option(NN).eq.8)then

        write(88,24)KKK
24   format(//Hydraulic conductivity m sec-1',/,
    &   'Iteration = ',i8)

```



```

call sieve(KP,88,nc,nnlmax)

endif
c -----
c *** SUCTION ***
c if(option(NN).eq.9)then

write(89,26)KKK
26 format(//Soil suction m',/,
& 'Iteration = ',i8)

call sieve2(sct,89,nc,nnlmax)

endif
c -----
c *** CUMULATIVE "F" RATIO ***
c if(option(NN).eq.10)then

write(90,28)KKK
28 format(//Cumulative F ratio per column',/,
& 'Iteration = ',i8)

write(90,*)(FCUM(I),I=1,NC)
endif
c -----
c *** MAXIMUM "F" RATIO ***
c if(option(NN).eq.11)then

write(91,29)KKK
29 format(//Maximum F ratio per column',/,
& 'Iteration = ',i8)
write(91,*)(FMAX(I),I=1,NC)
write(91,30)
30 format(//Depth of maximum F ratio from surface m'//)
write(91,*)(FDMAX(I),I=1,NC)
endif
c -----
c *** MINIMUM "F" RATIO ***
c if(option(NN).eq.12)then

write(92,31)KKK
31 format(//Minimum F ratio per column',/,
& 'Iteration = ',i8)
write(92,*)(FMIN(I),I=1,NC)

write(92,32)
32 format(//Depth of minimum F ratio from surface m'//)
write(92,*)(FDMIN(I),I=1,NC)
endif
c -----
c *** FACTOR OF SAFETY ***
c if(option(NN).eq.13)then

write(93,33)KKK
33 format(//Factor of Safety',/,
& 'Iteration = ',i8)

```

```

        call sieve(FS,93,nc,nnlmax)

        endif
c -----
c *** ATH1 ***
c if(option(NN).eq.14)then

        write(94,200)KKK
200   format('//Soil water content, volume m3',/,
      & 'Iteration=', i8)

        call sieve(ATH1,94,nc,nnlmax)
        endif
c -----
c *** VATH1 ***
c if(option(NN).eq.15)then

        write(95,202)KKK
202   format('//Soil ice content, volume m3',/,
      & 'Iteration=', i8)

        call sieve(VATH1,95,nc,nnlmax)
        endif
c -----
c *** ATEMP ***
c if(option(NN).eq.16)then

        write(96,204)KKK
204   format('//Ambient temperature change and time',/,
      & 'Iteration=', i8)

        write(96,*)ATEMP(I),TTSEC(I)

        endif
c -----
c *** BTFLUX ***
c if(option(NN).eq.17)then
        write(97,206)KKK
206   format('Basal heat flux W m-3 s-1',/,
      & 'Iteration =', i8)

        write (97,*)(BTFLUX(I),I=1,NC)
        endif
c *****
c GRAPHICS
c *****
c *** Gsltp ***

c ***** change this for correct output *****
do 800 K=1,nslices
if(option(NN).eq.20)then

c   do 35 J=1,NNLMAX
c   l=slice(K)
c   call sieve(70,*) KKK, sltp(I,J)
c   call sieve(sltp,70,nc,nnlmax)
c   call sieve(112,*) atemp(kkk)
c   continue
35

```

```

endif
C -----
C *** Gice ***
if(option(NN).eq.21)then

call sieve(VATHET,71,nc,nnlmax)
endif
C -----
C *** Gwater ***

if(option(NN).eq.22)then
call sieve(ATHESP,72,nc,nnlmax)
endif

C -----
C *** Gsuction ***
if(option(NN).eq.23)then

call sieve2(sct.73,nc,nnlmax)
endif
C -----
C *** Ghycond ***
if(option(NN).eq.24)then

call sieve2(KP.74,nc,nnlmax)
endif
C -----
C *** gATEMP ***
if(option(NN).eq.25)then

do 220 l=1,NOTC
C call sieve(75,*)TTSEC(l),ATEMP(l)
220 continue
endif
C -----
C *** gKT ***
if(option(NN).eq.26)then

call sieve2(KT,76,nc,nnlmax)
endif
C -----
C *** gAPKT ***
if(option(NN).eq.27)then

do 251 J=1,NNLMAX
l=slice(K)
C call sieve(77,*)APKT(l,J)
251 continue
endif
C -----
C *** FS ***
if(option(NN).eq.28)then

do 252 J=1,NNLMAX
l=slice(K)
C call sieve(78,*)FS(l,J)
252 continue

```

```

endif
c -----
c *** NHFLUX ***
c if(option(NN).eq.29)then
c   do 253 J=1,NNLMAX
c     l=slice(K)
c     call sieve(NHFLUX,79,nc,nnlmax)
c 253   continue
c   endif
c -----
c *** HFLUX ***
c if(option(NN).eq.30)then
c   do 254 J=1,NNLMAX
c     l=slice(K)
c     call sieve(80,*)HFLUX(I,J)
c 254   continue
c   endif
c -----
c *** COLUMN HEIGHT (CM) ***
c if(option(NN).eq.33)then
c   do 255 l=1,NC
c     write(33,*)COLHT(l)*100.0,(COLHT(l)-DEPTH(l))*100.0
c 255   continue
c   endif
c -----
c *** CELL EXPANSION (MM) ***
c if (option(NN).eq.34)then
c   call sieve(exp,34,nc,nnlmax)
c   endif
c -----
c *** SOIL VOLUMETRIC PERCENTAGE ***
c if(option(nn).eq.36)then
c   call sieve(soilp,36,nc,nnlmax)
c   endif
c -----
c *** SOIL DOWNSLOPE DISPLACEMENT POTENTIAL (cm)***
c if(option(nn).eq.37)then
c   do 256 l=1,NC
c     write(37,*)HDISP(l)*100.,GDISP(l)*100.,(GDISP(l)+HDISP(l))
c     & *100.
c 256   continue
c   endif
c -----
c *** EXPELLED EXCESS WATER VOLUMES ****
c if(option(nn).eq.38)then
c   call sieve(exwat,38,nc,nnlmax)
c   endif
c -----
c *** ISOTHERM ***
800   continue
103   continue

```

```

return
end

```

```

c      subroutine sieve
c      subroutine sieve sifts the desired arrays and outputs the cells
c      specified in the Out.data file thereby enabling selective output
c      rather than the entire column to be printed.

```

```

subroutine sieve(array, outid,nc,nnlmax)
integer nc,nnlmax,outid, counter
COMMON DUMMY(5,60)
double precision buffer(60), array(5,60)

```

```

do 1013 i=1,nc
counter=0
do 2013 j=1,nnlmax
if(DUMMY(i,j).eq.1)then
counter=counter+1
buffer(counter)=array(i,j)
endif
2013 continue
write(outid,'(30f9.4,x)')(buffer(col), col=1,counter)
1013 continue
return
end

```

```

c      subroutine sieve2

```

```

subroutine sieve2(array, outid,nc,nnlmax)
integer nc,nnlmax,outid, counter
COMMON DUMMY(5,60)
double precision buffer(60), array(5,60)

```

```

do 1013 i=1,nc
counter=0
do 2013 j=1,nnlmax
if(DUMMY(i,j).eq.1)then
counter=counter+1
buffer(counter)=array(i,j)
endif
2013 continue
c write(outid,'(30f9.4,x)')(buffer(col), col=1,counter)
write(outid,'*)(buffer(col), col=1,counter)
1013 continue
return
end

```

```

c      Subroutine soli.f

c      Subroutine soli.f calculates solifluction movement potential for column cells
c
c
c      subroutine soli(NC, NNL, NNLMAX, TTCOM,sct,FS,rDMAX,
&      FDMIN, FMAX, FMIN, FCUM,GAMMAS,GAMMAU,BETA,C,PHI,MAXCOLHT
&      ,DEPTH,HDISP,GELRAT,GDISP)

c      DECLARATION OF VARIABLES

      DOUBLE PRECISION C(5,60), BETA(5)
      DOUBLE PRECISION GAMMAS, GAMMAU,GELRAT,GDISP(5)
      DOUBLE PRECISION sct(5,60), PHI(5,60), FS(5,60)
      DOUBLE PRECISION TTCOM(5,60), FMAX(50)
      DOUBLE PRECISION FMIN(50), SMAX(50)
      DOUBLE PRECISION FDMAX(50), FDMIN(50),DEPTH(5)
      DOUBLE PRECISION FCUM(50),MAXCOLHT(5),HDISP(5)

      INTEGER NC, NNL(5), NNLMAX

c      *****
c      calculate maximum downslope displacement due to heave
c      resettlement using maximum displacement and slope angle

      DO 5 I=1,NC
      HDISP(I)=(MAXCOLHT(I)-DEPTH(I)) * TAND(BETA(I))
      GDISP(I)=HDISP(I)*GELRAT
5      CONTINUE

c      *****
c      Set FMAX and FMIN limits
      DO 1 I=1,NC
      FMAX(I)=-20000
      FMIN(I)=20000
      SMAX(I)=-20000
1      CONTINUE

c      Calculation of factor of safety.
      DO 3 I=1,NC
      M=NNL(I)
      DO 4 J=1,M
      IF(sct(I,J).GE.0.0)THEN
&      FS(I,J)=(C(I,J)+(GAMMAS*TTCOM(I,J)*(DCOS(BETA(I))))**2
&      -sct(I,J)*9.8)*DTAN(PHI(I,J))/(GAMMAS*TTCOM(I,J)
&      *DSIN(BETA(I))*DCOS(BETA(I)))
      ENDIF

      IF(sct(I,J).LT.0.0)THEN
&      FS(I,J)=(C(I,J)+(GAMMAU*TTCOM(I,J)*(DCOS(BETA(I))))**2
&      -sct(I,J)*9.8)*DTAN(PHI(I,J))/(GAMMAU*TTCOM(I,J)
&      *DSIN(BETA(I))*DCOS(BETA(I)))
      ENDIF

      IF(FS(I,J).GT.FMAX(I))FMAX(I)=FS(I,J)
      FDMAX(I)=TTCOM(I,J)
      IF(FS(I,J).LT.FMIN(I))FMIN(I)=FS(I,J)
      FDMIN(I)=TTCOM(I,.)

```

```
IF(sct(I,J).GT.SMAX(I))SMAX(I)=sct(I,J)
FCUM(I)=FCUM(I)+FS(I,J)
```

```
4   CONTINUE
3   CONTINUE
```

```
return
end
```

c **Subroutine thaw.f**

c Subroutine thaw.f calculates the partition of energy between
c two processes:-

- c (i) phase change
c (ii) thermal change
c when NHFLUX is +ve and thaw of soil moisture
c is a likely result.

subroutine thaw(sltpt2, TF, LENGTH,
& CDEPTH, BDT, ATHESP, VATHET, VHC, sltp, NC, LA,
& NHFLUX, I, J, KKK, CELVOL, VADFAC, SOILP, ORIVOL, SOIL1, ASR1, ASR2,
& LC1, FC2, TDEP, EXWAT)
c *****

c DECLARATION OF VARIABLES

DOUBLE PRECISION sltpt2(5,60), sltp(5,60), ASR1, ASR2
DOUBLE PRECISION ATHESP(5,60), ATHESP1(5,60)
DOUBLE PRECISION VATHET(5,60), VATH1(5,60)
DOUBLE PRECISION LENGTH, CDEPTH(5,60), BDT
DOUBLE PRECISION deltaW(5,60), deltaT(5,60), Et(5,60)
DOUBLE PRECISION LA, VHC(5,60), CELVOL(5,60), SOIL1(5,60)
DOUBLE PRECISION SOILP(5,60), volch(5,60), ORIVOL(5,60)
DOUBLE PRECISION VADFAC(5,60), TDEP
DOUBLE PRECISION EXWAT(5,60), WAT(5,60)

INTEGER LC1(50), FC2(50), TF, NC, I, J, KKK
c *****

c Is energy partition between phase change and thermal heating
c required, i.e. is ice present.
IF(VATHET(I,J).LE.0.0D-06.OR.SLTPT2(I,J).LT.TDEP)THEN
 sltp(I,J)=sltpt2(I,J)
ELSEIF(SLTPT2(I,J).GE.TDEP.AND.VATHET(I,J).GT.0.0)THEN
 sltp(i,j)=tdep
c *****

c Change in temperature for iteration period predicted by heattrans.f
c this is heat energy that will be used to melt ice
deltaT(I,J)=sltpt2(I,J)-TDEP

c calculate thermal energy available for melting ice
Et(I,J) = ((VHC(I,J)*CELVOL(I,J))*DeltaT(I,J))

c -----
c HYDROLOGY

c Convert Athesp and Athet from % volume to volume m3
ATHESP1(I,J)=CELVOL(I,J)*ATHESP(I,J)

c Volumetric change in moisture content - ice melted (wat equiv) m3
deltaW(I,J)=Et(i,j)/(LA*1000000.0)

c Convert VATHET ice content % into a volume m3
VATH1(I,J)=CELVOL(I,J)*VATHET(I,J)


```

c      check that ice is available for melting
      IF(VATH1(I,J).GT.(deltaw(I,J)*1.09))THEN
        VATH1(I,J)=VATH1(I,J)-(deltaW(I,J)*1.09)
      ELSE
        DELTAW(I,J)=VATH1(I,J)/1.09
        VATH1(I,J)=0.0
      ENDIF

      ATHESP1(I,J)=ATHESP1(I,J)+DELTAW(I,J)
      ATHESP(I,J)=ATHESP1(I,J)/CELVOL(I,J)
      VATHET(I,J)=VATH1(I,J)/CELVOL(I,J)

c      calculate if any volume adjustment is required
      IF(VADFAC(I,J).GT.1.0)THEN
        IF(J.LE.LC1(I))WAT(I,J)=ASR1*ORIVOL(I,J)
        IF(J.GE.FC2(I))WAT(I,J)=ASR2*ORIVOL(I,J)
c      include lines to expel water content that is above the original
c      porosity of the soil
        IF(ATHESP1(I,J).GT.WAT(I,J))THEN
          EXWAT(I,J)=EXWAT(I,J)+((ATHESP1(I,J)-WAT(I,J))
&          /ORIVOL(I,J))
          ATHESP1(I,J)=WAT(I,J)
        ENDIF

      VOLCH(I,J)=CELVOL(I,J)-(ATHESP1(I,J)+VATH1(I,J)+SOIL1(I,J))
      CELVOL(I,J)=CELVOL(I,J)-VOLCH(I,J)
      VADFAC(I,J)=CELVOL(I,J)/ORIVOL(I,J)
      IF(VADFAC(I,J).LT.1.0)THEN
        VADFAC(I,J)=1.0
        VOLCH(I,J)=CELVOL(I,J)-ORIVOL(I,J)
        CELVOL(I,J)=ORIVOL(I,J)
      ENDIF
      ATHESP(I,J)=ATHESP1(I,J)/CELVOL(I,J)
      VATHET(I,J)=VATH1(I,J)/CELVOL(I,J)
      SOILP(I,J)=SOIL1(I,J)/CELVOL(I,J)
      CDEPTH(I,J)=CDEPTH(I,J)-(VOLCH(I,J)/(BDT*LENGTH))

      endif
      else
      write(6,*)'slipped out in thaw'

      ENDIF

      return
      end

```