Poyal Norwegian Council for Scientific and Industrial Research NTNF)



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NORSAR Scientific Report No. 2-92/93

Semiannual Technical Summary

1 October 1992 — 31 March 1993

Kjeller, July 1993





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Abstract (cont.)

This Semiannual Report also presents statistics from operation of the Intelligent Monitoring System (IMS). The IMS has been operated in an experimental mode, and the performance has been very satisfactory. Since October 1991, a new version of the IMS that accepts data from an arbitrary number of arrays and single 3-component stations has been operated.

The NORSAR Detection Processing system has been operated throughout the period with an average uptime of 97.8% as compared to 96.7% for the previous reporting period. A total of 1733 seismic events have been reported in the NORSAR monthly seismic bulletin. The performance of the continuous alarm system and the automatic bulletin transfer by telex to AFTAC has been satisfactory. The system for direct retrieval of NORSAR waveform data through an X.25 connection has been tested successfully for acquiring such data by AFTAC. Processing of requests for full NORSAR and regional array data on magnetic tapes has progressed according to established schedules. There have been no modifications made to the NORSAR data acquisition system.

On-line detection processing and data recording at the NORSAR Data Processing Center (NDPC) of NORESS, ARCESS, FINESA and GERESS data have been conducted throughout the period. Data from two experimental small-aperture arrays at sites in Spitsbergen and Apatity, Kola Peninsula, have been recorded and processed in an experimental mode. Monthly processing statistics for the arrays as well as results of the IMS analysis for the reporting period are given.

Maintenance activities in the period comprise preventive/corrective maintenance in connection with all the NORSAR subarrays, NORESS and ARCESS. In addition, the maintenance center has been involved with occasional maintenance of equipment for FINESA. Other activities have involved testing of the NORSAR communications systems and work in connection with the experimental small-aperture arrays in Spitsbergen and the Kola Peninsula.

Starting 1 October 1991, an effort began to carry out a complete technical refurbishment of the NORSAR array. This project is funded jointly by AFTAC, ARPA and NTNF. During the reporting period, efforts have focused upon evaluation and laboratory testing of technical options for field instrumentation, in particular state-of-the-art A/D converters, data acquisition and synchronization devices. During the reporting period, we have also been testing several such systems under realistic operating conditions in the field. When these studies have been completed, a recommendation for a system to be installed will be presented to the funding agencies.

Summaries of six scientific contributions are presented in Chapter 7 of this report.

Section 7.1-7.3 describe the basic principles and give examples of application of a promising new approach to automatic analysis of events from sites with recurring seismic activity. This method, which we have called "intelligent post-processing of seismic events" is particularly well suited to implement as a post-event processing technique in the Intelligent Monitoring System.

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Section 7.1 describes the basic principles of this approach. From experience with analyst review of events automatically defined by the Intelligent Monitoring System (IMS), we have realized that the quality of the automatic event locations can be significantly improved if the event intervals are reprocessed with signal processing parameters tuned to phases from events in the given region. The tuned processing parameters are obtained from off-line analysis of events located in the region of interest. The primary goal of such intelligent post-processing is to provide event definitions of a quality that minimizes the need for subsequent analysis and review.

The first step in this post-processing is to subdivide the area to be monitored in order to identify sites of interest. Clearly, calibration will be easiest and potential savings in manpower are largest for areas with high, recurring seismic activity. We have identified 8 mining sites in Fennoscandia/W. Russia/Estonia and noted that 65.6 per cent of the events of $M_L > 2.0$ in this region can be associated with one of these sites. This result is based on 1 1/2 years of data.

Section 7.2 describes the second step in the post-processing approach. This second step is to refine the phase arrival time and azimuth estimates using frequency filters and processing parameters that are tuned to the initial event location provided by the IMS. We have studied, as a first example, a set of 58 events from the Khibiny Massif in the Kola peninsula. Very accurate locations of these events has been provided by the Kola Regional Seismological Centre. Our refinement of phase arrival times for these events, as recorded in Apatity and at ARCESS, have given quite remarkable results. Using the autoregressive likelihood technique of Pisarenko et al (1987) we have been able to estimate onset times automatically to an accuracy (standard deviation) of about 0.05 s for P-phases and 0.15-0.20 s for S-phases. These accuracies are as good as for analyst picks, and are considerably better than in the current SigPro analysis.

Section 7.3 describes the third step in the post-processing approach. This third step in the post-processing is a recomputation of the location estimate, using the refined arrival times and broad-band azimuths associated with a fixed frequency band (tied to the initial IMS location). Again using the set of 58 known Khibiny Massif events as a reference, we have investigated the accuracy of the LocSat location procedure using varying amounts of regional corrections. Using a depth constrained to 0, we find that the epicenter can be estimated to 2-3 km accuracy (median value), with "worst case" error less than 10 km, even without any regional corrections. It is noted that the method needs to be tested for additional target sites before firm conclusions can be drawn.

Section 7.4 describes a four-month monitoring experiment, using the Continuous Threshold Monitoring technique applied to the Novaya Zemlya test site. Data from the NORESS, ARCESS and FINESA regional arrays were used. Starting 1 December 1992, we have been compiling daily statistics of all peaks on the threshold diagram exceeding $m_b = 2.75$, and associating these peaks to regional or teleseismic events whenever possible. In addition, we are analyzing smaller peaks (below $m_b = 2.75$) that can possibly be associated with Novaya Zemlya epicenters.

We have conducted extensive analysis of the $m_b = 2.5$ seismic event at Novaya Zemlya at 09.29.25 GMT on 31 December 1992. This event was detected by ARCESS (P and S), Spitsbergen (P and S), NORESS (P) and Apatity (S). In addition, the Kola Regional Seis-

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mological Centre provided readings for the station Amderma (Pg and Sn) made from analog recordings. Our results indicate that the epicenter was slightly to the north of the northern Novaya Zemlya test site. However, there are some uncertainties in the travel time tables, and an on-site location cannot entirely be ruled out.

Section 7.5 contains some initial processing results from the Spitsbergen array. Preliminary analysis of data from this new small-aperture (1km) array on the Arctic island of Spitsbergen has revealed a large variation in signal and noise amplitudes across the array. Further investigations are needed in order to determine the reason for this variation, and it may prove necessary to redeploy one or more of the array sensors. Site B2 in this array exhibits a particularly favorable signal focusing for nearly all arrival azimuths. Initial studies of the noise characteristics show that the noise is quite stable over time, especially at high frequencies. Also, the noise amplitudes at high frequencies seem to be similar across the array, in contrast to the large variation seen at low frequencies.

Section 7.6 contains results from an evaluation of the most recent version of the Intelligent Monitoring System (IMS). In order to assess the performance of IMS, the automatic results for a one-week test period were carefully and thoroughly reviewed. The conclusion is that the overall performance of IMS has improved considerably from previous versions and can now be rated as very satisfactory, but also that there is still room for some improvements. Nearly 80% of the events automatically declared by IMS were judged to be acceptable in the sense that they correspond to real seismic events for which only relatively minor modifications were made during the analyst review process. The majority of the remaining 20% were judged to be false events, and some modifications to IMS are proposed that may help in reducing the number of such events. Using various reference bulletins, it was found that some events were missed by IMS, even if data were available to IMS that should allow it to associate phases and form events. An evaluation of the manual corrections made to the automatically determined onset times shows that there is a potential for improving these by, e.g., implementing in IMS new procedures for onset time estimation like those described in section 7.2.

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threshold monitoring of the northern Novaya Zemlya test site

Initial processing results from the Spitsbergen small-aperture array

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1 Summary

This Semiannual Technical Summary describes the operation, maintenance and research activities at the Norwegian Seismic Array (NORSAR), the Norwegian Regional Seismic Array (NORESS, the Arctic Regional Seismic Array (ARCESS) and the experimental Spitsbergen regional array for the period 1 October 1992 -31 March 1993. Statistics are also presented for additional seismic stations, which through cooperative agreements with institutions in the host countries provide continuous data to the NORSAR Data Processing Center (NPDC). These stations comprise the Finnish Experimental Seismic Array (FINESA), the German Experimental Seismic Array (GERESS), and an experimental regional seismic array in Apatity, Russia.

This Semiannual Report also presents statistics from operation of the Intelligent Monitoring System (IMS). The IMS has been operated in an experimental mode, and the performance has been very satisfactory. Since October 1991, a new version of the IMS that accepts data from an arbitrary number of arrays and single 3-component stations has been operated.

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Summaries of six scientific contributions are presented in Chapter 7 of this report.

Section 7.1-7.3 describe the basic principles and give examples of application of a promising new approach to automatic analysis of events from sites with recurring seismic activity. This method, which we have called "intelligent post-processing of seismic events" is particularly well suited to implement as a post-event processing technique in the Intelligent Monitoring System.

Section 7.1 describes the basic principles of this approach. From experience with analyst review of events automatically defined by the Intelligent Monitoring System (IMS), we have realized that the quality of the automatic event locations can be significantly improved if the event intervals are reprocessed with signal processing parameters tuned to phases from events in the given region. The tuned processing parameters are obtained from off-line analysis of events located in the region of interest. The primary goal of such intelligent post-processing is to provide event definitions of a quality that minimizes the need for subsequent analysis and review.

The first step in this post-processing is to subdivide the area to be monitored in order to identify sites of interest. Clearly, calibration will be easiest and potential savings in manpower are largest for areas with high, recurring seismic activity. We have identified 8 mining sites in Fennoscandia/W. Russia/Estonia and noted that 65.6 per cent of the events of $M_L > 2.0$ in this region can be associated with one of these sites. This result is based on 1 1/2 years of data.

Section 7.2 describes the second step in the post-processing approach. This second step is to refine the phase arrival time and azimuth estimates using frequency filters and processing parameters that are tuned to the initial event location provided by the IMS. We have studied, as a first example, a set of 58 events from the Khibiny Massif in the Kola peninsula. Very accurate locations of these events has been provided by the Kola Regional Seismological Centre. Our refinement of phase arrival times for these events, as recorded in Apatity and at ARCESS, have given quite remarkable results. Using the autoregressive likelihood technique of Pisarenko et al (1987) we have been able to estimate onset times automatically to an accuracy (standard deviation) of about 0.05 s for P-phases and 0.15-0.20 s for S-phases. These accuracies are as good as for analyst picks, and are considerably better than in the current SigPro analysis.

Section 7.3 describes the third step in the post-processing approach. This third step in the post-processing is a recomputation of the location estimate, using the refined arrival times and broad-band azimuths associated with a fixed frequency band (tied to the initial IMS location). Again using the set of 58 known Khibiny Massif events as a reference, we have investigated the accuracy of the LocSat location procedure using varying amounts of regional corrections. Using a depth constrained to 0, we find that the epicenter can be estimated to 2-3 km accuracy (median value), with "worst case" error less than 10 km, even without any regional corrections. It is noted that the method needs to be tested for additional target sites before firm conclusions can be drawn.

Section 7.4 describes a four-month monitoring experiment, using the Continuous Threshold Monitoring technique applied to the Novaya Zemlya test site. Data from the NORESS, ARCESS and FINESA regional arrays were used. Starting 1 December 1992, we have been compiling daily statistics of all peaks on the threshold diagram exceeding $m_b = 2.75$, and associating these peaks to regional or teleseismic events whenever possible. In addition, we are analyzing smaller peaks (below $m_b = 2.75$) that can possibly be associated with Novaya Zemlya epicenters.

We have conducted extensive analysis of the $m_b = 2.5$ seismic event at Novaya Zemlya at 09.29.25 GMT on 31 December 1992. This event was detected by ARCESS (P and S), Spitsbergen (P and S), NORESS (P) and Apatity (S). In addition, the Kola Regional Seismologicai Centre provided readings for the station Amderma (Pg and Sn) made from analog recordings. Our results indicate that the epicenter was slightly to the north of the northern Novaya Zemlya test site. However, there are some uncertainties in the travel time tables, and an on-site location cannot entirely be ruled out.

Section 7.5 contains some initial processing results from the Spitsbergen array. Preliminary analysis of data from this new small-aperture (1km) array on the Arctic island of Spitsbergen has revealed a large variation in signal and noise amplitudes across the array. Further investigations are needed in order to determine the reason for this variation, and it may prove necessary to redeploy one or more of the array sensors. Site B2 in this array exhibits a particularly favorable signal focusing for nearly all arrival azimuths. Initial studies of the noise characteristics show that the noise is quite stable over time, especially at high frequencies. Also, the noise amplitudes at high frequencies seem to be similar across the array, in contrast to the large variation seen at low frequencies.

Section 7.6 contains results from an evaluation of the most recent version of the Intelligent Monitoring System (IMS). In order to assess the performance of IMS, the automatic results for a one-week test period were carefully and thoroughly reviewed. The conclusion is that the overall performance of IMS has improved considerably from previous versions and can now be rated as very satisfactory, but also that there is still room for some improvements. Nearly 80% of the events automatically declared by IMS were judged to be acceptable in the sense that they correspond to real seismic events for which only relatively minor modifications were made during the analyst review process. The majority of the remaining 20% were judged to be false events, and some modifications to IMS are proposed that may help in reducing the number of such events. Using various reference bulletins, it was found that some events were missed by IMS, even if data were available to IMS that should allow it to associate phases and form events. An evaluation of the manual corrections made to the automatically determined onset times shows that there is a potential for improving these by, e.g., implementing in IMS new procedures for onset time estimation like those described in section 7.2.

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2 NORSAR Operation

2.1 Detection Processor (DP) operation

There have been 58 breaks in the otherwise continuous operation of the NORSAR online system within the current 6-month reporting interval. The uptime percentage for the period is 97.8 as compared to 96.7 for the previous period.

Fig. 2.1.1 and the accompanying Table 2.1.1 both show the daily DP downtime for the days between 1 October 1992 and 31 March 1993. The monthly recording times and percentages are given in Table 2.1.2.

The breaks can be grouped as follows:

a)	Hardware failure	3
b)	Stops related to program work or error	2
c)	Hardware maintenance stops	0
d)	Power jumps and breaks	0
e)	TOD error correction	5
f)	Communication lines	48

The total downtime for the period was 119 hours and 3 minutes. The mean-time-betweenfailures (MTBF) was 3.6 days, as compared to 2.0 for the previous period.

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Fig. 2.1.1. Detection Processor uptime for January (top), February (middle) and March (bottom) 1993.

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Date	Time	Cause
12 Oct	0939 - 1140	Software work
21 Oct	0720 - 1315	Line failure
02 Nov	1605 -	Line failure
03 Nov	- 0555	
07 Dec	1801 -	Line failure
08 Dec	- 0606	
21 Dec	0423 - 0802	Line failure
01 Jan	1007 - 1131	Software work
31 Jan	0906 - 1448	Hardware failure
31 Jan	1808 -	Hardware failure
01 Feb	-	Hardware failure
02 Feb	- 1455	Hardware failure
03 Mar	0759 -	Line failure
04 Mar	- 0747	
15 Mar	1248 - 1314	Hardware failure

 Table 2.1.1. The major downtimes in the period 1 October 1992 - 31 March 1993.

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Month	DP Uptime Hours	DP Uptime %	No. of DP Breaks	No. of Days with Breaks	DP MTBF* (days)
Oct 92	735.24	98.84	11	9	2.6
Nov 92	705.09	97.94	14	11	2.0
Dec 92	728.10	97.87	4	4	6.1
Jan 93	723.57	97.31	8	7	3.4
Feb 93	724.51	97.43	4	4	6.0
Mar 93	700.45	97.34	17	13	1.6
		97.79	58	48	3.6

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*Mean-time-between-failures = total uptime/no. of up intervals.

Table 2.1.2. Online system performance, 1 October 1992 - 31 March 1993.

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2.2 Array communications

General

Table 2.2.1 reflects the performance of the communications system throughout the reporting period.

The most prominent events which have affected the different systems have been: loss of synch, bad cable, spikes on data, PCM irregularities, Digital Access Cross Connection System (DACCS), SLEM, line outage and 2701 data adapter.

Detailed Summary

October (weeks 40-44), 28.9-1.11.92

NTA/Lillestrøm replaced the carrier frequency system pertaining to 01A-06C (-02C), with pulse code modulation equipment (PCM) between Lillestrøm and Hamar 1 Oct 93. An attempt to replace the 02C system was also made, but without success.

01A, which had been down since 18 June, was back in operation 23 Oct with satisfactory performance.

8 Oct 01B started with deteriorated data (spikes) and excessive error rate, after the original SLEM was reinstalled. Between 9 and 12 Oct the subarray was out of operation. During a subarray visit 15 Oct a SLEM failure occurred and after this the system was down throughout October.

02C continued to cause us trouble after NTA's first attempt to replace the carrier system with PCM equipment (1 Oct). Although it was (according to NTA/Lillehammer) changed back to the original system the same day. Between 2 and 7 Oct the subarray was affected by different irregularities related to the communication path, and was more or less out of operation until 12 Oct, when it was demasked. On 28 Oct 02C was again affected in connection with NTA work in Oslo.

Other subarrays which were affected in October were 03C and 06C; 03C between 4 and 6 Oct and 06C between 22 and 26 Oct.

All systems were affected 21 Oct between 0720 and 1316 GMT.

Average outages for October, individual weeks 40-44

Week 40 (-01A,02C)	:	0.003%
Week 41 (-01A,02C)	:	0.001%
Week 42 (-01A,01B,02C)	:	0.002%
Week 43 (N/A)	:	
Week 44-01B,02C,06C)	:	0.003%

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November (weeks 45-48), 2-29.11.92

01B resumed operation 11 Nov (week 46) after having been down since 8 Oct, due to a SLEM failure.

In the same period 02C and 04C were also affected. 02C had synch problems, and the 04C irregularity was caused by NTA/Hamar.

02B was affected between 12 and 16 Nov and from 17 through 30 Nov due to a bad communications cable. Also the 02C synch problem continued week 46 (39.5% down), and week 47 (17.4% down).

06C was affected week 48 (38.1% down), probably by a line outage. A Modcomp restart caused 06C reinitialization.

Average outages in November, individual weeks:

Week 45 (-01B,02C,04Cl)	:	0.002%
Week 46 (-01B,02B,02Cl)	:	0.0006%
Week 47 (-02B,02C)	:	0.003%
Week 48(-02B,06C)	:	0.002%

December (weeks 49-53), 30.11.92-3.1.93

02B went down week 46 due to a bad communications cable located in a swampy area, with snow depths exceeding 1.5 meters. A cable repair was therefore difficult according to NTA/Hamar. They suggested, however, a start date mid-January 93. 21 December, after 1 a.m. all systems were affected 4 hours caused by an NTA system irregularity.

Average outage in December, individual weeks:

Week 49 (-02B)	:	0.0008%
Week 50 (-02B)	:	0.0003%
Week 51 (-02B)	:	0.001%
Week 52 (-02B)	:	2.380%
Week 53 (-02B)	:	0.0008%

January (weeks 1-4), 4-31.1.93

NTA/Hamar finally finished the repair of the 02B communications cable on 29 Jan 93. The subarray had been out of operation since week 46/92.

After a few days' observation of the performance, the subarray communications were accepted. Communications errors were not discovered between 29 Jan 0945 GMT and 31 Jan approx 1800 GMT, which was the approximate hour that the communication adapter 2701 stopped, and the data transfer between the Modcomp and the IBM 4381 machine was suspended. The 2701 remained down throughout 31 Jan.

Average outages in January, individual weeks:

Week 1 (-02B)	:	0.002%
Week 2 (-02B	:	0.002%
Week 3 (-02B)	:	0.001%
Week 4 (-,02B)	:	0.001%

February (weeks 5-8), 1-28.2.93

The 2701 communications adapter which failed 31 Jan 93 approx. 1800 GMT was repaired 3 Feb, and the NORSAR system started again 1455 GMT. After restart all systems performed satisfactorily.

Average outages in February, individual weeks:

Week 5	: 0.0	02%
Week 6	: 0.0	02%
Week 7	: 0.0	02%
Week 8	: 0.0	02%

March (weeks 9-12), 1-28.3.92

3 March a DACCS (Data Access Cross Connection System) failed at the NTA/Lillestrøm premises and affected all NORSAR subarray communications systems (01A-06C). According to NTA/Lillestrøm all the systems (-02C) were in operation approx 2100 hrs GMT. On 4 March at 0544 GMT, an attempt to start NORSAR was made, but without success. The TOD (Time Code Generator), which was replaced 3 March, was reinstalled. 0747 GMT a new ONLINE start was successful, but 02C was still inoperative.

During the period misalignment of received data and TOD generator clock pulses have resulted in several Modcomp restarts. 4 March 02C, which remained down after the NTA/DACCS failure in Lillestrøm, was loop-tested between Kjeller and Sjusjøen. No errors were observed.

15 March a NORSAR ONLINE stop occurred at 1248 hrs GMT, up 25 min 42 sec later.

16 March a new stop occurred at 1256 GMT, and the system was down for about 5 min 42 seconds.

18 March NTA/Hamar carried out scheduled work on the 02C PCM system between the Nes peninsula and Sjusjøen.

22 March deteriorated data (spikes) were observed on 01A, 02B and 02C. At the same time a 2-second misalignment of received NORSAR data and TOD clock pulses were detected. In order to find out if the seriously deteriorated data had created a situation in the Modcomp which resulted in the misalignment, the three subarrays were omitted. In response to a request, NTA/Hamar informed us that an attenuator had been removed in connection with changing the path between Løten and 06C from PCM cable to radiolink. After having reinstalled the attenuator, NORSAR performance was significantly improved.

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Average outages in March, individual weeks:

Week 9 (-02C)	:	0.002%
Week 10 (-02C)	:	0.001%
Week 11 (all)	:	0.022%
Week 12 (all)	:	0.522%

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	Oct (5)	Nov (4)	Dec (5)	Jan (4)	Feb (4)	Mar (4	Average
Subarrays	28.9- 1.11.93	2-29.11.93	30.11.92- 3.1.93	4-31.1.93	1-28.2.93	1-28.3.93	1/2 year
01A	0.010 ¹⁾	0.004	0.001 ¹³⁾	0.008	0.001	0.225	0.047
01B	0.030 ²⁾	0.0005 ⁸⁾	0.0009 ¹⁴⁾	0.0007	0.0006	0.002	0.006
02B	0.0007 ³⁾	0.0019)	100.0 ^{N/A}	92.755 ^{N/A}	0.002	0.329	0.003 ²⁰⁾
02C	0.005 ⁴)	0.002 ¹⁰⁾	0.001 ¹⁵⁾	0.002	0.002	0.010 ¹⁹⁾	0.004
03C	0.002 ⁵⁾	0.0004	0.001 ¹⁶⁾	0.001	0.0004	0.011	0.003
04C	0.001 ⁶⁾	0.000111)	0.001 ¹⁷⁾	0.001	0.004	0.009	0.003
06C	0.0027)	0.002 ¹²⁾	0.003 ¹⁸⁾	0.003	0.002	0.369	0.063
AVER	0.007	0.001	0.001 ²¹⁾	0.016 ²²⁾	0.002	0.136	0.018

Figures representing error rate (in per cent) followed by number 1), 2), etc., are related to legend below.

Table 2.2.1. Communications performance. The numbers represent error rates in per cent based on total transmitted frames/week (28 Sep 92 - 28 Mar 93).

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1), 4)	one week (43/44)
9), 10)	one week (45/48)
2), 8), 19)	average 2 weeks (40,41/47,48) (11,12/93)
7), 11),12)	average 3 weeks (40,41,42/46,47,48/45,46,47)
3), 5), 6)	average 4 weeks (40,41,42,44)
13), 14), 15), 16)	average 4 weeks (49,50,51,53)
17), 18)	average 4 weeks (49,50,51,53)
20)	average 4 months (Oct, Nov 92/ Feb, Mar 93)
21)	average 6 subarrays (01A,02B, 02C-06C)
22)	average 6 subarrays (01A,02B, 02C-06C)

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2.3 NORSAR Event Detection operation

In Table 2.3.1 some monthly statistics of the Detection and Event Processor operation are given. The table lists the total number of detections (DPX) triggered by the on-line detector, the total number of detections processed by the automatic event processor (EPX) and the total number of events accepted after analyst review (teleseismic phases, core phases and total).

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	Tatal	Tatal	Accepte	d events			
	DPX	EPX	P-phases	Core Phases	Sum	Daily	
Oct 92	13450	1812	241	131	372	12.0	
Nov 92	12175	1324	189	50	239	8.0	
Dec 92	13752	1512	189	61	250	8.1	
Jan 93	11324	1265	148	74	222	7.2	
Feb 93	11450	1241	225	61	286	10.2	
Mar 93	13000	1505	257	107	364	11.7	
			1249	484	1733	9.5	

Table 2.3.1. Detection and Event Processor statistics, 1 October 1992 - 31 March 1993.

NORSAR Detections

The number of detections (phases) reported by the NORSAR detector during day 275, 1992, through day 090, 1993, was 74,237, giving an average of 412 detections per processed day (180 days processed). Table 2.3.2 shows daily and hourly distribution of detections for NORSAR.

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Table 2.3.2 (Page 1 of 4)

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Table 2.3.2. (Page 2 of 4)

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Table 2.3.2. (Page 3 of 4)

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hour and the total sum of detections during the period. The averages show number of processed days, hourly distribution and average per processed day. (Page 4 of 4)

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3 **Operation of regional arrays**

3.1 **Recording of NORESS data at NDPC, Kjeller**

Table 3.1.1 lists the main outage times and reasons.

The average recording time was 99.69% as compared to 98.55% during the previous reporting period.

Date	Time	Cause				
11 Nov	1925 - 1941	Transmission line failure				
31 Dec	2310 - 0000	Software work				
08 Mar	1200 - 1216	Hub maintenance				
15 Mar	0934 - 0952	Transmission line failure				

Table 3.1.1.	Interruptions in recording of NORESS data at NDPC, 1 October 1992 - 3	31
March	1993.	

Monthly uptimes for the NORESS on-line data recording task, taking into account all factors (field installations, transmissions line, data center operation) affecting this task were as follows:

October	:	99.97
November	:	99.96
December	:	99.89
January	:	100.00
February	:	100.00
March	:	99.9 1

Fig. 3.1.1 shows the uptime for the data recording task, or equivalently, the availability of NORESS data in our tape archive, on a day-by-day basis, for the reporting period.

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December (bottom) 1992.



Fig. 3.1.1. (cont.) NORESS data recording uptime for January (top), February (middle) and March (bottom) 1993.

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3.2 Recording of ARCESS data at NDPC, Kjeller

Table 3.2.1 lists the main outage times and reasons.

The average recording time was 99.61% as compared to 99.25% for the previous reporting period.

Date	Time	Cause				
27 Nov	2123 - 2214	Hardware failure NDPC				
28 Nov	2153 - 2230	Hardware failure NDPC				
16 Dec	1057 - 1147	Hardware failure NDPC				
18 Dec	0641 - 0912	Software work				
31 Dec	2311 - 0000	Software work				
01 Feb	1821 - 1927	Satellite link failure				
01 Feb	2054 -	Satellite link failure				
02 Feb	- 0516					

Table 3.2.1. The main interruptions in recording of ARCESS data at NDPC, 1 October1992 - 31 March 1993.

Monthly uptimes for the ARCESS on-line data recording task, taking into account all factors (field installations, transmissions line, data center operation) affecting this task were as follows:

October	:	99.99%
November	:	99.78%
December	:	99.4 1%
January	:	100.00%
February	:	98.49%
March	:	99.98%

Fig. 3.2.1. shows the uptime for the data recording task, or equivalently, the availability of ARCESS data in our tape archive, on a day-by-day basis, for the reporting period.

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Fig. 3.2.1. ARCESS data recording uptime for October (top), November (middle) and December (bottom) 1992.

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Fig. 3.2.1. ARCESS data recording uptime for January (top), February (middle) and March (bottom) 1993.

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3.3 Recording of FINESA data at NDPC, Kjeller

The average recording time was 98.78% as compared to 93.75% for the previous period. As can be seen from Table 3.3.1 below, the main reason for the downtime is transmission line failure.

Date	Time	Cause
23 Oct	0434 - 0935	Transmission line failure
10 Dec	0547 - 1314	Transmission line failure
18 Dec	1337 - 1422	Transmission line failure
18 Dec	1427 - 1447	Transmission line failure
21 Dec	1021 - 1035	Transmission line failure
26 Jan	1636 - 1651	Transmission line failure
26 Jan	2104 - 2137	Transmission line failure
27 Jan	1255 -	Software work in Helsinki
28 Jan	- 1440	
29 Jan	2316 -	Hardware failure NDPC
30 Jan	- 0030	
31 Jan	2213 -	Transmission line failure
01 Feb	- 0712	Transmission line failure
10 Mar	1043 - 1057	Transmission line failure
10 Mar	1334 - 1349	Transmission line failure
10 Mar	1525 - 1544	Transmission line failure

Table 3.3.1. The main interruptions in recording of FINESA data at NDPC, 1 October1992 - 31 March 1993.

Monthly uptimes for the FINESA on-line data recording task, taking into account all factors (field installations, transmission lines, data center operation) affecting this task were as follows:

October	:	99.13%
November	:	99.98%
December	:	98.79%
January	:	95.99%
February	:	98.91%
March	:	99.87%

Fig. 3.3.1 shows the uptime for the data recording task, or equivalently, the availability of FINESA data in our tape archive, on a day-by-day basis, for the reporting period.

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Fig. 3.3.1. FINESA data recording uptime for October (top), November (middle) and December (bottom) 1992.



Fig. 3.3.1. FINESA data recording uptime for January (top), February (middle) and March (bottom) 1993.

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3.4 Event detection operation

This section reports results from one-array automatic processing using signal processing recipes and "ronapp" recipes for the ep program (NORSAR Sci. Rep. No 2-88/89).

Three systems are in parallel operation to associate detected phases and locate events:

- 1. The ep program with "ronapp" recipes is operated independently on each array to obtain simple one-array automatic solutions.
- 2. The Generalized Beamforming method (GBF) (see F. Ringdal and T. Kværna (1989), A mulitchannel processing approach to real time network detection, phase association and threshold monitoring, BSSA Vol 79, no 6, 1927-1940) processes the four arrays jointly and presents locations of regional events.
- 3. The IMS system is operated on the same set of arrivals as ep and GBF and reports also teleseismic events in addition to regional ones.

IMS results are reported in section 3.5 and GBF results in section 3.6.

In addition to these three event association processes, we are running test versions of the so-called Threshold Monitoring (TM) process. This is a process that monitors the seismic amplitude level at the four regional arrays continuously in time to estimate the upper magnitude limit of an event that might go undetected by the network. The current TM process is beamed to several sites of interest, including the Novaya Zemlya test site. Simple displays of so-called threshold curves reveal instants of particular interest; i.e., instants when events above a certain magnitude threshold may have occurred in the target region. Results from the three processes described above are used to help resolve what actually happened during these instances.

NORESS detections

The number of detections (phases) reported from day 275, 1992, through day 090, 1993, was 36,783 giving an average of 202 detections per processed day (182 days processed).

Table 3.4.1 shows daily and hourly distribution of detections for NORESS.

Events automatically located by NORESS

During days 275, 1992, through 090, 1993, 2074 local and regional events were located by NORESS, based on automatic association of P- and S-type arrivals. This gives an average of 11.4 events per processed day (183 days processed). 63% of these events are within 300 km, and 88% of these events are within 1000 km.

ARCESS detections

The number of detections (phases) reported during day 275, 1992, through day 090, 1993, was 91,474, giving an average of 503 detections per processed day (182 days processed).

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Table 3.4.2 shows daily and hourly distribution of detections for ARCESS.

Events automatically located by ARCESS

During days 275, 1992, through 090, 1993, 3359 local and regional events were located by ARCESS, based on automatic association of P- and S-type arrivals. This gives an average 18.5 events per processed day (182 days processed). 43% of these events are within 300 km, and 84% of these events are within 1000 km.

FINESA detections

The number of detections (phases) reported during day 275, 1992, through day 090, 1993, was 58,399, giving an average of 321 detections per processed day (182 days processed).

Table 3.4.3 shows daily and hourly distribution of detections for FINESA.

Events automatically located by FINESA

During days 275, 1992, through 090, 1993, 2835 local and regional events were located by FINESA, based on automatic association of P- and S-type arrivals. This gives an average of 15.6 events per processed day (182 days processed). 65% of these events are within 300 km, and 87% of these events are within 1000 km.

GERESS detections

The number of detections (phases) reported from day 275, 1992, through day 090, 1993, was 7,456, giving an average of 182 detections per processed day (41 days processed).

Table 3.4.4 shows daily and hourly distribution of detections for GERESS.

Events automatically located by GERESS

During days 275, 1992, through 090, 1993, 717 local and regional events were located by GERESS, based on automatic association of P- and S-type arrivals. This gives an average of 17.5 events per processed day (41 days processed). 72% of these events are within 300 km, and 90% of these events are within 1000 km.

Apatity detections

The number of detections (phases) reported from day 275, 1992, through day 090, 1993, was 90,712, giving an average of 501 detections per processed day (181 days processed).

As described in earlier reports, the data from Apatity is transferred by one-way (simplex) radio links to Apatity city. The transmission suffers from radio disturbances that result in a
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large number of small data gaps and spikes in the data. Although the communication protocol may correct such errors by requesting retransmission of data, this cannot be done at Apatity. For such error corrections, a two-way radio link is needed (duplex radio). Some of the detections are consequently due to bad data spikes. However, it should be noted that noise from cultural activities and from the nearby lakes cause most of the unwanted detections. These unwanted detections are "filtered" in the signal processing, as they give seismic velocities that are outside accepted limits for regional and teleseismic phase velocities.

Table 3.4.5 shows daily and hourly distribution of detections for Apatity.

Events automatically located by Apatity

During days 275, 1992, through 090, 1993, 2691 local and regional events were located by Apatity, based on automatic association of P- and S-type arrivals. This gives an average of 15.0 events per processed day (180 days processed). 31% of these events are within 300 km, and 66% of these events are within 1000 km.

Spitsbergen detections

The number of detections (phases) reported from day 275, 1992, through day 090, 1993, was 19,359, giving an average of 150 detections per processed day (129 days processed).

Table 3.5.6 shows daily and hourly distribution of detections for Spitsbergen.

Events automatically located by Spitsbergen

The installation at Spitsbergen is not completed, and automatic processing for location of events has only been done for experimental purposes. See also section 7.5.

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hour of the day, and number of detections for that day. The end statistics give total number of detections distributed for each hour and the total sum of detections during the period. the averages show number of processed days, hourly distribution and average Table 3.4.3. Daily and hourly distributions of FINESA detections. For each day is shown number of detections within each 128 14 15 15 16 12 9 8 9 11 13 14 19 22 18 12 11 12 11 13 13 14 16 15 320 Average workdays 313 Average weekends Thuraday Priday Saturday Sunday Nonday 182 14 15 15 16 13 11 9 9 11 12 13 16 18 15 12 12 13 12 13 14 15 16 15 321 Total average 58399 Total sum ****** Date 33 EEEE à **J** *53322-656 8 6 14 14 15 17 14 13 11 10 11 10 10 9 10 9 10 14 15 13 15 14 15 17 16 16 202 8 49 2700 2768 2460 2914 33 8 24454545454545 20 21 52°55555°C 7 19 30 62.6 vendidides 19 2524288454 . 2657 2928 1930 1698 2192 2888 2780 2156 2170 2449 2534 2700 2370 1630 2031 2325 3359 2152 2343 2265 24 08 09 10 11 12 13 14 15 16 17 18 19 18 85555577755558 855555777555558 -17 E. 82878 é N **مە**ق 16 848425466689 รูอุท 15 a~a 16°611°° 1 9 11 1282848 3°821 07 08 09 10 11 12 21.96 88681951-723 88611951-723 detections 312 49166798065 © **₹** 5 221 400-04443 σ distribution of 101 1 ~ 75 5 510 ູ່ມານກາວກິດຜິ 05 06 4 (1 0 04 05 06 งน้องนดงบบดอยี่ย 01101 68701 **1**30 per processed day. (page 4 of 4) 20 5 3066 10 90 81 a 60 02 03 .FKX Hourly 3 ၿရှိဖ 516% C3953865 5 5 23225 2522 • 8 8 PIN 5 NI Qay Sum

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Figs. 7.1.2-7.1.4 show the event distribution in Fennoscandia for all magnitudes, $M_L > 1$

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3.5 IMS operation

The Intelligent Monitoring System (IMS) was installed at NORSAR in December 1989 and was operated at NORSAR from 1 January 1990 for automatic processing of data from ARCESS and NORESS. A second version of IMS that accepts data from an arbitrary number of arrays and single 3-component stations was installed at NORSAR in October 1991, and regular operation of the system comprising analysis of data from the 4 arrays ARCESS, NORESS, FINESA and GERESS started on 15 October 1991. As opposed to the first version of IMS, the one in current operation also locates events at teleseismic distance.

During this reporting period, two new data sources were made available to IMS: the array near Apatity on the Kola peninsula of Russia, and the array on Spitsbergen. The analysts had waveforms available from these two arrays from 11 December 1992. On 14 December 1992, phase detections from the Apatity array were included in the automatic phase association. The phase detections from the Spitsbergen array were made available to the analysts on 5 February 1993. These detections are not used in the automatic phase association, but can be added manually during analysis.

The operational stability of IMS has been very good during the reporting period. In fact the IMS event processor (pipeline) has had no downtime of its own; i.e., all data available to IMS have been processed by IMS.

Phase and event statistics

Table 3.5.1 gives a sumary of phase detections and events declared by IMS. From top to bottom the table gives the total number of detections by the IMS, the number of detections that are associated with events automatically declared by the IMS, the number of detections that are not associated with any events, the number of events automatically declared by the IMS, the total number of events defined by the analyst, and finally the number of events automatically declared by the analyst without any changes (i.e., from the set of events automatically declared by the IMS).

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	Oct 92	Nov 92	Dec 92	Jan 93	Feb 93	Mar 93	Total
Phase detections	38026	26150	39885	40148	49504	50710	206397
-Associated phases	5835	5439	5936	5838	6127	8923	32263
-Unassociated phases	32191	20711	33949	34310	43377	41787	174134
Events automatically declared by IMS	1742	1671	1785	1755	1765	2554	9530
No. of events defined by the analyst	1275	1359	1136	1141	1430	2125	7191
No. of events accepted without modifications	1018	880	951	954	1105	1623	5513

Table 3.5.1. IMS phase detections and event summary

U. Baadshaug B. Ferstad B.Kr. Hokland L.B. Loughran

3.6 GBF operation

Events automatically located by GBF

During days 275, 1992, through 090, 1993, 9280 local and regional events were located by GBF. This gives an average of 51.0 events per processed day (182 days processed). 70% of these events are within 300 km of nearest station, and 85% of these events are within 1000 km of the nearest station.

70.8% of these events were defined by 2 regional phases. Teleseismic phases are currently not used by GBF. 86.7% of all events had 3 defining phases or less.

T. Kværna

4 Improvements and Modifications

4.1 NORSAR

NORSAR data acquisition

No modification has been made to the NORSAR data acquisition system.

The data are recorded on a 30-hour circular disk buffer on the IBM system, and archived onto 1/2 inch magnetic tapes. In addition to this, the data are now regularly transmitted to a SUN system for recording on a 48-hour circular disk buffer.

NORSAR detection processing

The NORSAR detection processor has been running satisfactorily on the IBM 4381 computer during this reporting period.

Detection statistics are given in section 2. In addition to the detection processing done on IBM, the dp program is doing regular detection processing on a SUN system, using the unix-based circular disk buffer (see below). A detection SNR threshold of 20.0 triggers automatic saving of waveforms into CSS 3.0 data files.

NORSAR event processing

There have been no changes in the routine processing of NORSAR events, using the IBM system.

In parallel with the IBM processing, routine event processing is also done on a SUN computer using the "old" IBM time delay correction data base that has been converted to SUN/UNIX. The automatic solutions produced are equal to or better than the old system with a lower false alarm rate. Alert messages are sent to USGS for events above magnitude 5.5.

NORSAR refurbishment

As reported earlier, the main problem in this refurbishment is to find 24-bit A/D converters with low power consumption. We have received evaluation units from Refraction Technology, Teledyne Geotech and Science Horizons. During the reporting period, the available units have been extensively tested, and they appear to perform satisfactorily. Details on the refurbishment effort will be reported on separately.

4.2 Regional Arrays

Detection processing

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The routine detection processing of the arrays is running satisfactorily on each of the array's SUN-3/280 acquisition systems. The same program is used for NORSAR, NOR-

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ESS, ARCESS, FINESA, GERESS, Apatity and Spitsbergen, but with different "recipes". The beam table for NORESS and ARCESS is found in NORSAR Sci. Rep. No. 1-89/90. The beam table for FINESA and GERESS is found in NORSAR Sci. Rep. No. 1-90/91. The beam table for Apatity is found in NORSAR Sci. Rep. No. 1-92/93, and that for Spitsbergen is given in the present report (section 7.5).

Detection statistics are summarized in section 3.

Signal processing. Phase estimation

This process performs f-k and polarization analysis for each detection to determine phase velocity, azimuth and type of phase, and the results are put into the ORACLE detection data base for use by the IMS.

Event Processing. Plot and epicenter determination

A description of single-array event processing is found in NORSAR Sci. Rep. No. 2-88/ 89, and NORSAR Sci. Rep. No. 2-89/90.

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5 Maintenance Activities

5.1 Activities in the field and at the Maintenance Center

This section summarizes the activities at the Maintenance Center (NMC) Hamar, and NDPC activities related to monitoring and control of NORSAR, including monitoring of NORESS, ARCESS, FINESA, GERESS and Apatity.

Activities involve preventive and corrective maintenance, planning and installation of the Spitsbergen array (28 Oct - 7 Nov 92) and completion of the Apatity installation (22 Sep - 3 Oct 92).

NORSAR

Visits to subarrays in connection with:

- Test of SLEM equipment in order to locate a failure
- Reinstallation of SLEM used for test purposes
- Adjustment of gain SP/LP channels
- Adjustment of MP/FP
- Adjustment of DC offset SP channels
- Replacement of MP/FP motors incl. adjustments
- SLEM rest
- Replacement of BE-card
- Comm. cable test
- Modem and line test

NMC

- Preparations for the Spitsbergen installation
- NORSAR refurbishment preparations
- Other tasks related to repair of SLEM and related subarray equipment

NORESS

• Repair of C6, D6

Spitsbergen

- Installation of the Spitsbergen array (Oct/Nov 92)
- Installation of a second windmill and extra batteries (Feb 93)

Apatity

Installation continued through 30 Sep 92 and was completed 3 Oct 92

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Subarray/ area	Task	Date
NORSAR		
01 B	Installation of 01A analog unit in 01B SLEM for test purposes (possible because 01A was still down)	6 Oct
01 B	Spikes in data still required CTV engagement. A test with a data scope at the NDPC and the CTV modem (in C-loop) proved that the communications system was OK. A Digital Unit from the 01B SLEM was installed in the 01A SLEM for continued testing	12 Oct
01 B	Visit to the SA in connection with installation of original SLEM with a new power unit. Power unit failed (lack of +15 Volt). No spare unit was available.	15 Oct
01 A	Reinstalled the SLEM used for test purposes on 01B. Attempts to adjust the gain SP/LP resulted in improper data.	15 Oct
01 A	A new visit to the subarray resulted in proper data and demasking of the array.	22 Oct
NPDC	Daily checks of the following arrays have been carried out, i.e., NORSAR, NORESS, ARCESS, FINESA and Apatity. SP/LP instruments have been calibrated. Free Period (FP) and Mass Pos. (MP) were measured. Those outside specifications adjusted (when feasible from NDPC).	Οα
NORSAR		
01 B	SLEM brought to NMC for repair	9 Nov
01 B	SLEM installed	11 Nov
02B	Found communication inoperative	20 Nov
04C	Replaced BE-card channel 1. Adjusted gain all SP/LP chan- nels. Adjusted FP/MP all LP instruments	30 Nov
NDPC	Daily checks of the following arrays have been carried out, i.e., NORSAR, NORESS, ARCESS, FINESA, Apatity and Spitsbergen (partly). SP/LP instruments were calibrated weeks 45, 47 and 48. FP/ MP were measured. Those outside specifications adjusted (when feasible from NDPC).	Nov

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Subarray/ area	Task	Date
NMC	The staff have been engaged in preparations in connection with the refurbishment of the NORSAR array. In addition, installation of the Spitsbergen array (28 Oct - 7 Nov has been carried out.	Nov
NORSAR		
03C	Adjusted channel gain all SP/LP channels. Adjusted MP/FP V, NS, EW seismometers. Adjusted DC offset all SP chan- nels.	2 Dec
04C	LP seismometer NS repaired	4 Dec
02B(tel)	Replaced the telemetry receiver station batteries.	4 Dec
NMC	The NMC staff continued with preparations for the refurbish- ment of the NORSAR array.	Dec
NDPC	Daily checks of the following arrays have been carried out, i.e., NORSAR, NORESS, ARCESS, FINESA, Apatity and Spitsbergen. SP/LP instruments were calibrated weeks 45, 49 - 52. FP/ MP were measured. Those outside specifications adjusted (when feasible from NDPC).	Dec
NORSAR		1993
02B	A SLEM reset carried out.	11 Jan
NMC	NORSAR refurbishment preparations continued.	Jan
NDPC	Daily checks of the following arrays have been carried out, i.e., NORSAR, NORESS, ARCESS, FINESA, Apatity and Spitsbergen. SP/LP instruments were calibrated (NORSAR). FP/ MP were measured. Those outside specifications adjusted (when feasi- ble from NDPC).	Jan
Spitsbergen	A second windmill and extra batteries were mounted.	19 Feb
NMC	NORSAR refurbishment preparations continued.	Feb

Subarray/ area	Task	Date
NDPC	Daily checks of the following arrays have been carried out, i.e., NORSAR, NORESS, ARCESS, FINESA, Apatity and Spitsbergen. SP/LP instruments were calibrated (NORSAR). FP/ MP were measured. Those outside specifications adjusted (when feasi- ble from NDPC).	
NORSAR		
02C	Replaced NS MP motor. Adjusted EW,NS Mass Position and Free Period (MP/FP).	11 Mar
02C	Replaced EW MP/FP motor and adjusted.	29 Mar
NORESS	Replaced C6.	23 Mar
	Final repair of D6. NORSAR refurbishment	31 Mar
NPDC	Daily checks of the following arrays have been carried out, i.e., NORSAR, NORESS, ARCESS, FINESA, Apatity and Spitsbergen. SP/LP instruments were calibrated (NORSAR). FP/ MP were measured. Those outside specifications adjusted (when feasi- ble from NDPC).	

Table 5.1. Activities in the field and the NORSAR Maintenance Center, including NDPC activities related to NORSAR, NORESS, ARCESS, FINESA, GERESS, Apatity and Spitsbergen 1 Oct 92 - 31 Mar 93.

5.2 Array status

As of 31 March 1992 the following NORSAR channels deviated from tolerances:

01A 01	8 Hz filter
02	8 Hz filter
04	30 dB attenuation
02B 01	
08	
09	
02C 01	
08	
03C 04	

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7 Summary of Technical Reports / Papers Published

7.1 Intelligent post-processing of seismic events -- Part 1: Basic approach

Introduction

This is the first in a series of three contributions in this report addressing the topic of intelligent post-processing of seismic events. In this first contribution we discuss how to subdivide the area to be monitored in order to identify sites of particularly high seismic activity. We further introduce the basic idea behind this post-processing technique, which is to use as a starting point the initial event location provided by the Intelligent Monitoring System (IMS) and then use region-specific information to refine the solution. By applying this technique to areas with significant recurring seismic activity, such as mining sites, a considerable part of the analyst work can be eliminated.

Since 15 October 1991, the Intelligent Monitoring System (IMS, Bache et al, 1993) has been processing seismic data from four high-frequency arrays in northern Europe. These are NORESS and ARCESS in Norway, FINESA in Finland and GERESS in Germany. During October 1992 a small-aperture array was installed near Apatity on the Kola peninsula, and during late October/early November 1992 another small-aperture array on the island of Spitsbergen became operational. The data from these installations are now included in the IMS processing for the production of the event bulletin.

Since four of the arrays providing data to the IMS are located in Fennoscandia, see Fig. 7.1.1, the IMS event bulletin shows an excellent event detection capability for this region. Ringdal (1991) found that for Fennoscandia/NW Russia, a network consisting of NOR-ESS, ARCESS and FINESA has a 90% detection capability close to M_L 2.0. Near the individual arrays, the detection capability is considerably better, and consequently a large number of events less than M_L 1.0 are detected.

IMS event statistics

The basic principle of the post-processing method is to start by subdividing the area to be monitored into smaller areas, and subsequently apply region-specific analysis to each such area. As an example, we will consider in some detail the statistics of events in Fennoscandia and NW Russia for the 18-month time period 10/15/91 - 04/15/93. We will only consider "well-defined" events; thus we ignore events with author identification "yes/no" and "ESAL/Poor_Loc" in the origin table.

For the time period 10/15/91 to 04/15/93 the IMS bulletin contains 19503 well-defined events. 65.6% (12799) of these events are located in the Fennoscandian/NW Russian region defined by the map of Fig. 7.1.2, 15.8% (3089) are located within 5 degrees of the GERESS array, and the remaining 18.5% (3615) are distributed around the rest of the world, mostly at teleseismic distances from the regional array network.

Figs. 7.1.2-7.1.4 show the event distribution in Fennoscandia for all magnitudes, $M_L > 1$ and $M_L > 2$, respectively. In each figure, we have marked the approximate geographical extent of 8 main mining areas. Table 7.1.1 lists these mining sites and gives details on the number and percentages of events associated to the sites at various magnitudes.

From the three figures and Table 7.1.1 we can make the following general observations:

- Out of the total 12799 events, 6317 (49.4 per cent) are above $M_L = 1$, and only 1131 (8.8 per cent) are above $M_L = 2$.
- The total percentage of events associated with the 8 mining sites is 47.88% (all magnitudes), 56.66% ($M_L > 1.0$) and 65.61% ($M_L > 2.0$). Thus, these sites become more dominant for the largest events, in terms of relative number of events reported.
- Some mining sites have a relatively high proportion of large events ($M_L > 2.0$). This is particularly noticeable for the mining areas in Western Russia/Estonia. On the other hand, the Kiruna mine has the largest number of events altogether, but almost none of these are above $M_L 2$.

Being based on about 1 1/2 year of data, the statistics discussed here should be reasonably representative for the situation in the Fennoscandian/NW Russia region. Thus, analysis of recurring events from these mining areas is a significant workload for the analyst. An automatic method to improve the automatic analysis so as to obtain location precisions comparable with the analysts' results would be a significant development. In this and the next two sections of this report, we will show that such an improvement is possible for a well-calibrated mining area (the Khibiny Massif).

General outline of the method

Most automatic detection processor algorithms work without any *a priori* assumptions as to when and where a seismic event occurred. This is, of course, quite reasonable, and to some extent inevitable. The detector (SigPro) associated with the IMS works in this way. As a result, some of the SigPro output parameters, which later will be used by the IMS ESAL system, are less than optimum.

However, once an initial event location is given by the IMS, it is possible to use this initial location successively in an automatic iteration scheme. Each iteration gives a more precise location, which in turn allows the automatic program to place successively stronger constraints on the processing parameters.

As a first example (see Section 7.2), we can consider the estimation of signal arrival time. Given that an event has occurred in a certain area, the automatic program can select a set of optimum filter bands and beam parameters for this area, prior to reassessing the arrival time estimate. As shown in Section 7.2, this can lead to a remarkable improvement in timing precision. The examples given in Section 7.2 make use of an autoregressive likelihood technique (Pisarenko et al, 1987; Kushnir et al, 1990). It is notewortary that this method seems to require that the search be limited to a relatively short time window in order to work well. If an initial location and origin time is known, we can obtain the required short

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data in a low passband, e.g. 1 - 2 Hz, where S has a low SNR, and the first S-onset is in many cases missed. With this problem in mind we have found it difficult to justify a com-

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time window for the search. The method is therefore well suited to a post-processing application.

Another example is the estimation of azimuth from either arrays or three-component stations. The advantage of using a fixed frequency interval for broadband F-k analysis was convincingly demonstrated by Kværna and Ringdal (1986). Again, a prerequisite was the knowledge that the event in question was located in a certain known area.

Section 7.3 demonstrates that the approach of doing post-processing based on IMS initial solutions has the potential of providing an order-of-magnitude improvement in location precision, at least in certain cases such as the Khibiny Massif near the Apatity array. The improvement may be less if no network station is located close to the source, but it should still be significant. For example, the data from Kværna and Ringdal (1986) indicate that a single array (NORESS) would be capable of locating the Blåsjø explosions to within an accuracy of 10 km or better at a distance of 300 km. This is compared to the typical uncertainty of about 30 km in traditional single-array location estimates at this distance (Mykkeltveit and Ringdal, 1981).

In general, it is true that regional corrections are required in order to compute an optimum location. Again, the post-processing analysis is well suited toward this end, because the corrections can be tied to the general area, to which the initial IMS processing assigns the event.

In this context, it is important to note that no regional travel-time tables need to be involved as long as an adequate set of calibration events for the general area are available. The corrections for systematic bias may be made both to the phase arrival times and to the estimated azimuths. Again, this subject will be discussed in detail in Section 7.3, in connection with an application of the method to the Khibiny Massif area in the Kola peninsula.

T. Kværna

U. Baadshaug

F. Ringdal

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Dester	All Mag		Mag > 1.0		Mag > 2.0	
Kegion	Number	Percent	Number	Percent	Number	Percent
Fennoscandia/NW Russia	12799	100.00	6317	49.36	1131	8.84
Estonia	1487	11.62	1159	18.36	225	19.89
Karelia	379	2.96	212	3.36	70	6.19
Khibiny	1374	10.74	1106	17.51	233	20.60
Kiruna	1953	15.26	634	10.04	11	0.97
Kostomuksha	69	0.54	69	1.09	47	4.16
Kovdor	112	0.88	99	1.57	34	3.01
Nikel	620	4.84	181	2.87	104	9.20
Siilinjaervi	134	1.05	119	1.88	18	1.59
Total for 8 mines	6128	47.88	3579	56.66	742	65.61

Table 7.1.1. Distribution of events in mining regions of Fennoscandia and NW Russia. Events with author identificantion "yes/no" and "ESAL/Poor_Loc" in the origin table are not included in the statistics. 6

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Seismic stations



Fig. 7.1.1. Map showing the location of the regional arrays currently used by the Intelligent Monitoring System in operation at the NORSAR processing center.



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Fig. 7.1.2. Location of 12,799 events (all magnitudes) processed by the IMS for an 18month period. Only event solutions of satisfactory quality have been included (see text for details). Note the concentration of events in selected mining areas.





Fig. 7.1.3. Same as Fig. 7.1.2, but showing only events of $M_L > 1.0$. The total number of events is 6317 for the 18-month period.

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Fig. 7.1.4. Same as Fig. 7.1.2, but showing only events of $M_L > 2.0$. The total number of events is 1131 for the 18-month period.

7.2 Intelligent post-processing of seismic events -- Part 2: Accurate determination of phase arrival times using autoregessive likelihood estimation

Introduction

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A precise estimation of the onset time of seismic phases is necessary to obtain an accurate event location. In the context of automatic processing of the regional arrays operated by NORSAR, a two-step procedure has been in use since the first regional array, NORESS, was established in 1985 (Mykkeltveit and Bungum, 1984). This procedure consist of first applying a series of STA/LTA detectors in parallel to a set of filtered beams. When one or more of the STA/LTA detectors exceed a predefined threshold, a phase detection is declared and a detection time is found. Subsequently, a time domain phase timing algorithm is applied to the filtered beam with the highest SNR, using the detection time as the starting value. A detailed description of this algorithm is found in Mykkeltveit and Bungum (1984).

These estimates of the onset times are used by the automatic phase association and event location procedure (ESAL) of the Intelligent Monitoring System (IMS) (Bache et al, 1993), and the fully automatic processing results are finally reviewed and corrected by the analyst using the Analyst Review Station (ARS) of the IMS. Through the analyst review we have experienced that in most cases the phase onsets need to be adjusted. In section 7.6 of this report, comparisons between the automatic and manual onset time estimates of various seismic phases are presented. For P-type phases the standard deviations of automatic - manual onset time generally varies between 0.5 and 0.7 seconds. It should be noticed that these estimates represent averages for all seismic phases for a one-week period, and thus include phases with very different SNR's, frequency characteristics and signatures. For phases with an high SNR and an instantaneous signature, the performance of the phase timing algorithm is significantly better. Nevertheless, one of the conclusions from the study presented in section 7.6 is that in order to reduce the analyst's workload, there is a strong need to improve the precision of the automatic onset time estimates.

In this study we will investigate the potential automatic use of an onset picker based on autoregressive likelihood estimation (Pisarenko et al, 1987, Kushnir et al, 1990). Both a single-component version (ESTON1) and a three-component version (ESTON3) of this method will be tested on data from events located in the Khibiny Massif on the Kola Peninsula of Russia, recorded at the Apatity array, the Apatity three-component station and the ARCESS array, see Fig. 7.2.1.

The Khibiny Massif events

Six apatite mines are located within an area of about 10 km² in the Khibiny Massif (see Fig. 7.2.1). A detailed description of these mines and the mining activity is found in Mykkeltveit (1992). The coordinates of the mines are given in Table 7.2.1. Notice that Mine I consists of both an underground mine and an open-pit mine. Although we have no

explicit information on the exact sizes of these mines, interpretation of various maps suggest that the typical size is about 1 km^2 . The Kola Regional Seismological Centre has since the beginning of 1991 provided NORSAR with information on mining blasts in the six Khibiny mines. The information provided contains an assignment of the relevant mine (I-VI), P (and normally also S) arrival times at the analog APA (Apatity) station (co-located with the three-component station), the amplitude and period of the signal, and the total charge size in tons. Detailed information on the 58 events used in this study is given in Table 7.2.2.

At the Apatity array (APA0) and the Apatity three-component station (APZ9) which are located within a distance range of 18 - 49 km from the different Khibiny mines, the seismograms show clear P (Pg), S (Sg or Lg) and Rg arrivals (see Figs. 7.2.2 and 7.2.3). At ARCESS which is located about 400 km from the mines, a clear Pn and emergent Pg, Sn and Lg phases are observed (see Fig. 7.2.4). Table 7.2.3 gives detailed information on the distances and azimuths from APZ9, APA0 and ARCESS to the six Khibiny mines.

Estimation of P arrival time

We will not go into any detail concerning the theory of the autoregressive likelihood technique for onset time estimation, but refer to Pisarenko et al (1987) and Kushnir et al (1990) for details. We will, however, concentrate on the practical aspects of implementing the method as part of an automatic processing sequence. The onset time procedure requires that we have available an approximate timing of the phase arrival, and the search for the exact onset is limited to an interval around the initial arrival time. The initial arrival time can be obtained in several ways:

- Pisarenko et al (1987) and Kushnir et al (1990) suggest that an optimal event detector based on an Bayesian approach using autoregressive modelling of the data should be used for phase detection and approximate timing.
- Another alternative, which we will use for P-phases at APA0 and ARCESS, is to take the onset time provided by the IMS as the initial estimate. This onset time is calculated using the algorithm described in the *Introduction*.
- A third alternative is to use the predicted phase arrival time, e.g., derived from an initial event location and origin time or from the expected pattern of phase arrivals from events in a given region. This is the approach to be used for both P- and S-phases at APZ9 and for S-phases at APA0.

The autoregressive likelihood estimator is based on regarding the signal onset time as the moment in time when the statistical features of the observed time series (single-component or three-component) are abruptly changed. The single-component method thus takes into account changes in both power and frequency content, and it is therefore important that the broadband signal waveforms are retained (no narrow bandpass filtering). In addition to taking into account changes in power and frequency content, the three-component method is also sensitive to changes in the polarization characteristics of the three-component data. From experimenting with both ESTON1 and ESTON3 we have found that in order to obtain stable estimates, the data should first be lowpass filtered and decimated

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according to the highest frequencies of the signals. For P-phases from the Khibiny events recorded at the Apatity stations and at ARCESS, signal frequencies up to the Nyquist frequency of 20 Hz are in almost all cases observed. Consequently, no lowpass filtering or decimation is applied to the data. However, a low-order prediction error filter, designed from a 25 second noise sample preceding the P-phase, is applied.

For the arrays APA0 and ARCESS, the slowness and azimuth of the P-phases are computed using broad-band f-k analysis (Kværna and Doornbos, 1986). Kværna and Ringdal (1986) showed that the key to obtain stable slowness and azimuth estimates of phases from a given region is to process the data in a fixed frequency band, using a fixed time window positioning. The time window positioning for the f-k analysis and the fact that the onset time is computed from a steered beam (with steering delays found from f-k analysis) make these two processes to be closely integrated. Through extensive testing we have found that for P-phases from the Khibiny mines recorded at APA0, the most stable slowness and azimuth estimates are obtained if the f-k spectrum is computed in the frequency band 4.0 - 10.0 Hz using the 9 SPZ-component sensors of the array. The time window has a length of 2 seconds and starts 0.3 seconds ahead of the P-arrival. For ARCESS the most stable estimates are obtained if all 25 SPZ-component sensors are processed in the frequency band 3.0 - 6.0 Hz using a 3 second time window starting 0.3 seconds ahead of the P-arrival. Further refinement can be made by including a procedure that checks the SNR in the predefined frequency band, and adjusts the frequency cutoffs in the case of low SNR.

For APA0 and ARCESS, the automatic algorithm used for estimating the P-onsets from events from the Khibiny Massif can be summarized in the following way:

- Use the P onset from the IMS as the initial arrival time estimate.
- Generate a beam with steering delays corresponding to the IMS f-k results and prewhiten the beam with a low order prediction error filter.
- Apply ESTON1 to the prewhitened beam in an interval of ±2 seconds around the initial onset estimate. A sliding window length of 1 second is used by ESTON1, and the autoregressive modelling is of order 3.
- Position the time window according to the onset time from ESTON1 and run the f-k analysis as outlined above.
- Generate a new beam with steering delays corresponding to the new f-k results and prewhiten the beam with a low order prediction error filter.
- Apply ESTON1 to the new prewhitened beam in an interval of ±2 seconds around the previous ESTON1 onset estimate.
- Position the time window according to the last onset time estimate from ESTON1 and run the f-k analysis once more as outlined above.

For the three-component station APZ9 the situation is somewhat different as this station is not part of the automatic IMS processing. APZ9 is, however, located only 18 km from APA0, and we therefore choose to use the onset time at APA0 as the initial arrival time estimate. The processing can be summarized as follows:

• Use the P onset at APAO as the initial arrival time estimate.

- Prewhiten the Z-component APZ9 with a low order prediction error filter.
- Apply ESTON1 to the prewhitened Z-component APZ9 in an interval of ±5 seconds around the initial onset estimate from APA0. The sliding window length used by ESTON1 is 1 second and the autoregressive modelling is of order 3.
- Reestimate the onset time applying ESTON1 to the prewhitened z-component APZ9 in an interval of ±2 seconds around the previous ESTON1 onset estimate.

Precision of automatic P-onsets using the single-component estimator (ESTON1)

The precision of the automatic P-onsets using ESTON1 has been assessed by two different methods. The first method is to compare the ESTON1 estimates to the best manual pick. The purpose of this approach is to obtain information on any bias in ESTON1 estimates, and also to check the consistency between the automatic and manual onset estimates. In the second method we estimate the standard deviation of the phase picks by looking at the consistency of the arrival time differences between phases from events located in the same mine, see Sereno (1990). In this way, an unbiased estimate of the standard deviation is found for both the automatic and manual picks.

In Fig. 7.2.5.a the time differences between the automatic onsets from IMS and the manual pick at APA0 are presented as a function of signal-to-noise ratio (SNR) on the prewhitened beam. The standard deviation of these differences is as high as 0.43 seconds. In comparison, the time difference between the automatic ESTON1 onsets and the manually picked onsets, given in Fig. 7.2.5.b has a standard deviation as low as 0.02 seconds, and has a systematic positive bias of 0.05 seconds, i.e. the ESTON1 onsets are consistently picked a bit late.

In Fig. 7.2.6 similar plots are presented for Pn phases recorded at ARCESS. The automatic onsets from IMS shows significantly less scatter ($\sigma = 0.13$ s) than for APA0. On the other hand, the automatic ESTON1 onsets has a somewhat larger scatter ($\sigma = 0.03$ s) and a somewhat larger positive bias (0.08 s) than at APA0. Nevertheless, the accuracy of the ESTON1 time picks appears to be significantly better than those of the current processing system.

To address possible dependency of ESTON1 on the signal-to-noise ratio, we have in Fig. 7.2.7 plotted the time difference between the automatic ESTON1 onsets and the manually picked onsets for both APA0 and APZ9. In this way, we get an overview of the scatter for SNR's ranging from about 3 to about 500. We find from this figure that the time differences seem to be almost independent of SNR, which again indicate that the ESTON1 onset estimator works well for quite low signal-to-noise ratios. For direct comparison, we have in Fig. 7.2.8 plotted the ARCESS data on the same scale, but for ARCESS no low SNR phases are observed, as all SNR's exceed 6. By SNR, we mean the maximum of the linear ratio STA/LTA (i.e., short term average divided by long term average.)

The results presented above show that an automatic onset time procedure for P-phases using ESTON1 can give a remarkable improvement in precision compared to the current algorithm used in the IMS. The ESTON1 onset estimator has a bias that is dependent on

the dominant frequency of the signal. At the Apatity stations the average dominant P frequency is 13 Hz and the average bias is 0.05 s, whereas at ARCESS the average dominant P frequency is 6 Hz and the average bias is 0.08 s. The difference in bias can not be due to differences in sampling rate, since both APAO, APZ9 and ARCESS have a sampling frequency of 40 Hz. For the high frequency signals at the Apatity stations, the bias shows no clear dependency on SNR, at least not for SNR above 3. As expected, the results also suggest that the precision of ESTON1 increases with increasing signal frequencies. This can be inferred from the fact that the automatic P-onset estimates of the high-frequency signals at the Apatity stations are more consistent with the manual picks than the P-onsets at ARCESS, which generally have a lower dominant frequency. According to Pisarenko et al (1987) there are no analytical expressions for the theoretical biases and variances of ESTON1 and ESTON3, and it is therefore necessary to obtain empirical values. In any case, the biases and variances are less than 0.1 s for the P-phases considered in this study.

An unbiased estimate of the measurement variance is determined from the arrival time difference between two phase observations for repeated events in the same mine. Specifically:

$$\sigma_{1, pick}^{2} + \sigma_{2, pick}^{2} = \frac{\sum_{k=1}^{N_{mines}} \sum_{i=1}^{N_{obs}} [\Delta T_{obs_{ik}} - \langle \Delta T_{obs} \rangle_{k}]^{2}}{(N_{obs} - N_{mines})}$$

where σ_1^2 and σ_2^2 are the picking variances of each phase, $\Delta T_{obs_{a}}$ is the *ith* observation of the arrival-time difference for the *kth* mine. $\langle \Delta T_{obs} \rangle_k$ is the mean arrival time difference for the *kth* mine. N_{obs} is the total number of observations (at all mines), and N_{mines} is the number of mines.

By computing the arrival-time differences between the P observations at the three stations APAO(1), APZ9(2) and ARCESS(3), we get three equations of the type above, with altogether three unknowns. These equations can easily be solved to obtain estimates of each individual variance value. For example, for the automatic picks we obtain:

$$\sigma_1^2 + \sigma_2^2 = (0.065)^2$$

$$\sigma_1^2 + \sigma_3^2 = (0.065)^2$$

$$\sigma_2^2 + \sigma_3^2 = (0.074)^2$$

which gives $\sigma_1 = 0.04$, $\sigma_2 = 0.05$ and $\sigma_3 = 0.05$ (see Table 7.2.4).

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The standard deviations of both the manual picks and the automatic picks from ESTON1 are given in Table 7.2.4. All standard deviations are less than or equal to 0.06 s, and a part of this variability is likely due to the fact that the events of each mine are not located at the same spot, but are distributed within the mine. Under the assumption that each mine has an extent of 1 x 1 km (which is reasonable from interpretation from various maps), we have computed the maximum theoretical P travel-time difference between APA0 and APZ9 for the six Khibiny mines. The largest value is obtained for Mine I, where a 0.09 s travel-time difference is possible. The smallest values are found for Mine V and Mine VI, where a 0.04 s travel-time difference is possible. This clearly suggest that location variability within each mine can have a significant impact on the estimates of the picking precision. If we, however, assume that the distribution of the events within each mine increases all picking error estimates with a similar amount, a likely interpretation of the differences in picking precision between the three stations is as follows:

- The P picks at APZ9 are less precise than at APA0 due to generally lower SNR.
- The P picks at ARCESS are less precise than at APA0 due to generally lower dominant frequency of the signals.

The results of Table 7.2.4 show that the automatic ESTON1 method matches the human precision, and that for the relatively high SNR P arrivals analyzed in this study, the standard deviation of the automatic picks is well below 0.1 s.

Estimation of S arrival time at the Apatity stations

As seen from Fig. 7.2.2 and 7.2.3, the seismograms of the Khibiny events show clear P, S and Rg arrivals at the Apatity stations APAO and APZ9. The S-onsets do, however, become more emergent with increasing source-receiver distance. The S-phases typically have the largest SNR on the transverse component, but they are also clearly observed on the radial and vertical components. For automatic estimation of S-onsets, we have experimented with both the single-component ESTON1 onset estimator applied to the transverse component data, and found that the ESTON3 method gave the best results. The most stable results were obtained if the data were first filtered in a relatively wide band between 2.0 and 8.0 Hz and then decimated to a sampling rate of 20 Hz.

The Rg arrivals occur very close in time to the S-onset. Moreover, the Rg phase is dispersive. We therefore did not succeed in estimating the Rg-onset in a reliable way with the autoregressive likelihood estimation technique. In order to design an automatic processing sequence for the Khibiny events at the Apatity stations APA0 and APZ9, we have utilized the expected pattern of phase observations from events in this region, and the procedure is as follows:

- Estimate P-onset as previously outlined
- Identify the peak of the Rg-phase from an STA envelope created from the filtered zcomponent (0.8 - 2.0 Hz). The search interval for the Rg maximum is currently limited to 20 seconds after the P-onset.

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- For the APAO array, the slowness and azimuth of the Rg phase is estimated using data in a 5 s window starting 3 s ahead of the Rg peak. A frequency band of 0.8 2.0 Hz is used in the f-k analysis.
- For estimation of S-onsets, the three-component data are first filtered and decimated as outlined above, and the ESTON3 estimator is then applied within a time interval that starts 2 s after the P-onset and stops at the time of the Rg peak. The sliding window length used by ESTON3 is 2 seconds and the autoregressive modelling is of order 3.
- For the APAO array, the slowness and azimuth of the S phase is estimated using data in a 2 s window starting 0.3 s ahead of the S-onset. A frequency band of 2-5 Hz is used in the f-k analysis.

Illustrations of the automatic processing sequence are shown if Figs. 7.2.9 and 7.2.10. Notice the clear peak of the ESTON3 likelihood function at the S-onset for both events.

Precision of automatic S-onsets using the three-component estimator (ESTO? 3)

When comparing the automatic ESTON3 onsets to the manually picked S-onsets at APA0, we find that the standard deviation of the time differences is 0.18 s and that the bias is 0.01 s. For APZ9 the standard deviation is 0.12 s and the bias is 0.0^{4} s. Due to the relatively large scatter in the observations, this bias of 0.04 is not significantly different from 0.

To compute the measurement variance of the S-onset estimates, we have again investigated the consistency of the arrival time difference between two phase observations for repeated events in the same mine, according to equation 7.2.1. By using the P-onsets at both APAO, APZ9 and ARCESS as references, for which the measurement variances are known (see. Table 7.2.4), we can reliably estimate both the manual and the automatic Sonset variances at both APAO and APZ9. The results are summarized in Table 7.4.5

We see from Table 7.4.5 that the picking uncertainty estimates based on all three reference P-phases are very consistent, indicating that the method for estimating uncertainty works well. The fact that the precision of the manual and automatic picks are about equal, indicates that the automatic S-onset algorithm using ESTON3 matches the human precision. Table 7.4.5 also show that the S-onsets at APZ9 ($\sigma = 0.13$ s) are generally more precise than at APA0 ($\sigma = 0.19$ s). We believe that this is due to the fact that the S-phases become more emergent with increasing source-receiver distance, as illustrated in Figs. 7.2.2 and 7.2.3. The six Khibiny mines are located within a distance range of 18 - 33 km from APZ9, and within a distance range of 32 - 49 km from APA0, see Table 7.2.3.

It would have been interesting to compare the precision of the automatic S-onsets from ESTON3 to the precision of the automatic S-onsets used by the IMS. We have, however, experienced that the continuous processing of the Apatity array data (APA0) has problems in detecting and estimating the onsets, slowness and azimuths of secondary phases that have little time separation and large differences in frequency content. The first S-detection of the Khibiny events recorded at APA0 is in almost all cases declared in a low frequency band, which is typical for detection of Rg. Consequently the onset routine prefilters the

data in a low passband, e.g. 1 - 2 Hz, where S has a low SNR, and the first S-onset is in many cases missed. With this problem in mind we have found it difficult to justify a comparison between the automatic S-onsets from ESTON3 and the automatic S-onsets from the procedure providing data to the IMS.

Precision of azimuth estimates from broad-band f-k analysis

As described above, the estimation of onset time and the estimation of slowness and azimuth by f-k analysis are closely integrated processes. In one-array location of seismic events, the azimuth estimates are necessary to be able to compute an event location, and in the event location procedure of the IMS (Bratt and Bache, 1988) the phase azimuth estimates are required to be accompanied with an uncertainty estimate.

In Table 7.5.6, the mean and the standard deviation of the azimuth residuals relative to the Khibiny mine locations of the phases recorded at the arrays APA0 and ARCESS are presented. We see from the table that the P and Rg azimuths at APAO have about the same standard deviation (3.9 and 3.4 degrees, respectively), but that the P azimuths have a systematic bias of 8.2 degrees. The Rg azimuth bias is as low as -1.8 degrees. The ARCESS P azimuths have a very low standard deviation (0.9 degrees), but a systematic bias of 4.6 is consistently observed. The S-azimuths at APAO show a very large scatter. This show that if the Khibiny events are located without introducing corrections for the azimuth biases, the Rg azimuths should be given the smallest a priori uncertainty. If the systematic biases are removed, the a priori uncertainty of P and Rg at APA0 become comparable. With the systematic bias removed, the a priori uncertainty of ARCESS P is very small, but it should be noticed that in the event location procedure (Bratt and Bache, 1988) the a priori azimuth uncertainty is scaled by the source-receiver distance, such that an azimuth observation at 400 km distance with an uncertainty of 0.9 degrees (ARCESS P) is given less weight than an azimuth observation at 40 km distance with an uncertainty of 3.9 degrees (APA0 P).

Conclusions

The results presented in this study show that very precise automatic estimates of phase onsets from events in the Khibiny Massif can be obtained with the autoregressive likelihood estimation method. Implementation of the method requires that we have available approximate estimates of the phase arrival, and we have shown that such approximate estimates can be obtained from the IMS event definitions (phase association and event location) or from analysis of previous events in the region. In this way, the autoregressive likelihood estimation method can provide phase onsets that match the human precision. The uncertainties and biases of the automatic onset estimates of various phases at the Apatity stations and at ARCESS have been quantified, and the precision of the automatic phase picks shows very large improvement in comparison to the automatic phase onsets from the continuous processing providing input to the IMS.

In addition to the automatic onset estimation, we have estimated the phase azimuths and slowness using broad-band f-k analysis. This has been done using data in a fixed frequency band, using a fixed window positioning, as suggested by Kværna and Ringdal

(1986) for the purpose of obtaining increased stability. For P-phases recorded at ARCESS, the results are in accordance with those of Kværna and Ringdal (1986), where a scatter of only ± 1 degree were observed at NORESS for Pn phases from repeated events at a distance of 300 km. Comparing to the overall azimuth uncertainty of phases recorded at ARCESS and NORESS (Sereno, 1990), we find that if the systematic biases are removed from our azimuth estimates, the event location precision can be significantly improved. Without introducing azimuth corrections, the Rg azimuths at APAO show to be quite reliable.

We realize that in order to obtain accurate event locations, precise onset time and azimuth estimates are necessary, but not sufficient. If the theoretical travel-time model used in the event location deviates from the true travel-times, the precision of the event locations will be reduced. To overcome this, we will in Section 7.3 of this report discuss the introduction of travel-time corrections.

A natural extension to this study will be to investigate the performance of ESTON1 and ESTON3 applied to phases at the other stations of the network recording the Khibiny events. At NORESS and FINESA the detected Pn phases have quite low SNR, and it would be interesting to quantify the precision of the automatic onset estimates of these phases. It also remains to test and implement the automatic picking of Sn and Lg phases at ARCESS, NORESS and FINESA.

During the work with the autoregressive likelihood estimation method, we have experienced that the display of the likelihood functions, as illustrated in Figs. 7.2.9 and 7.2.10 can assist the analyst in picking correct phase onsets. In the context of interactive analysis of seismic data, we believe that the idea of making such likelihood functions available to the analyst should be pursued.

T. Kværna

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Mine	Location	
I, Underground	67.6702°N	33.7285°E
I, Open-pit	67.665°N	33.744°E
Ш	67.647°N	33.761°E
III	67.631°N	33.835°E
IV	67.624°N	33.896°E
V	67.632°N	34.011°E
VI	67.665°N	34.146°E

Table 7.2.1. The table gives the location of the six mines in the Khibiny Massif shown in Fig. 7.2.1.

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Event	Mine	ESTON1 P-arrival (APA0)	Size	
1	I (U)	1992-312:06.31.49.425	185 tons	
2	I (U)	1992-355:07.10.51.931	190 tons	
3	I (O)	1992-366:04.35.22.825	14 tons	
4	I (O)	1993-017:06.51.15.831	4 tons	
5	I (U)	1993-031:04.14.11.263	140 tons	
6	I (O)	1993-038:07.13.16.881	7 tons	
7	I (U)	1993-045:07.39.45.200	185 tons, double	
8	I (O)	1993-052:07.24.49.706	14 tons, double	
9	I (O)	1993-059:04.24.01.850	30 tons	
10	I (O)	1993-066:06.58.54.356	12 tons	
11	I (O)	1993-073:11.28.09.731	10 tons	
12	I (O)	1993-080:11.26.36.106	10 tons	
13	II	1992-334:04.08.51.725	130 tons	
14	II	1992-366:09.41.50.156	100 tons	
15	II	1993-024:08.04.48.656	24 tons	
16	II	1993-066:04.14.37.475	60 tons	
17	III	1992-354:08.38.32.806	15 tons	
18	III	1992-361:09.53.07.550	320 tons	
19	ш	1993-016:09.20.47.756	20 tons	
20	ш	1993-030:10.16.00.581	20 tons	
21	III	1993-030:13.07.15.650	265 tons	
22	III	1993-044:12.05.41.725	300 tons	
23	III	1993-051:07.28.29.931	16 tons	
24	III	1993-058:12.01.17.356	17 tons	
25	III	1993-065:08.51.42.356	18 tons	
26	III	1993-065:11.00.36.931	130 tons	

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Event	Mine	ESTON1 P-arrival (APA0)	Size
27	II	1993-065:11.29.46.150	Induced earth- quake, M _L 2.3
28	Ш	1993-086:10.35.05.925	14 tons
29	IV	1992-353:12.29.32.681	340 tons
30	IV	1992-360:12.37.40.556	380 tons
31	IV	1992-365:11.39.53.300	214 tons
32	IV	1993-006:11.01.24.331	310 tons
33	IV	1993-012:12.04.00.406	130 tons
34	IV	1993-022:12.50.56.806	310 tons
35	IV	1993-036:12.35.08.831	230 tons
36	IV	1993-043:13.37.27.181	350 tons
37	IV	1993-057:12.31.59.475	380 tons
38	IV	1993-064:12.32.40.931	300 tons, double
39	IV	1993-075:12.29.49.650	70 tons
40	IV	1993-082:15.08.34.250	95 tons
41	IV	1993-089:14.44.12.256	100 tons
42	v	1992-346:08.56.58.906	183 tons
43	v	1992-360:08.03.37.581	138 tons
44	v	1993-006:08.18.41.606	101 tons
45	v	1993-015:14.08.51.075	188 tons
46	V	1993-029:10.07.17.831	195 tons
47	v	1993-036:08.05.42.006	unknown size
48	v	1993-050:12.34.53.506	203 tons
49	v	1993-057:10.33.31.006	149 tons
50	v	1993-078:09.45.13.906	260 tons, general
51	v	1993-085:08.59.22.900	146 tons
52	VI	1992-318:06.41.16.250	60 tons
53	VI	1992-332:07.05.26.950	unknown size

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Event	Mine	ESTON1 P-arrival (APA0)	Size	
54	VI	1992-353:07.35.26.956	73 tons	
55	VI	1992-365:07.23.22.631	80 tons	
56	VI	1993-022:08.20.09.231	285 tons	
57	VI	1993-043:08.13.29.381	140 tons	
58	VI	1993-071:07.33.22.906	140 tons, general	

Table 7.2.2. This table contains information on the 58 Khibiny Massif events analyzed in this study. For each event, the assigned mine, the automatic P-onset at the Apatity array and the event size are given. I(U) means the underground Mine I, whereas I(O) means the open-pit mine. When the event size is given as "double", the reported charge is distributed between two explosions that are closely separated in time. If the term "general" is used, the reported charge is distributed among several explosions that are closely separated in time.

	АР	'A0	AP	Z9	ARC	CESS
Mine	Delta (km)	Az (deg)	Delta (km)	Az (deg)	Delta (km)	Az (deg)
I(U)	32.0	76.8	17.8	50.4	393.8	118.0
I (O)	32.6	78.0	18.0	53.1	394.7	118.0
Π	33.0	81.7	17.5	59.8	396.4	118.2
III	35.9	85.2	19.6	69.0	400.0	118.1
IV	38.4	86.6	21.8	73.3	402.6	118.0
v	43.4	85.7	26.7	74.4	406.0	117.5
VI	49.4	81.8	33.3	70.8	408.6	116.6

Table 7.2.3. Distance and azimuths from the three stations APA0, APZ9 and ARCESS to the mines considered in this study.

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	omanual	σ _{automatic}
P, Apatity array	0.04 s	0.04 s
P, Apatity 3-comp.	0.06 s	0.05 s
Pn, ARCESS	0.06 s	0.05 s

Table 7.2.4.	. Estimated stan	idard deviations	of manual and	automatic (E	STON1) P-onsets
at the	Apatity array, t	he Apatity three	e-component st	ation and at A	RCESS.

S-phase	Reference phase	σ _{manual}	^o automatic
APA0	P at APA0	0.20	0.19
	P at APZ9	0.20	0.19
	Pn at ARCESS	0.18	0.20
APAO	Average	0.19	0.19
APZ9	P at APA0	0.13	0.15
	P at APZ9	0.12	0.14
	Pn at ARCESS	0.13	0.16
AP29	Average	0.13	0.15

Table 7.2.5. Estimated standard deviations of manual and automatic (ESTON3) S-onsets at APAO and APZ9. In addition to the average uncertainty, we give for each phase the uncertainty estimates computed using the three different reference phases.

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	Mean (°)	σ (⁰)
P, Apatity array	8.2	3.9
Pn, ARCESS	4.6	0.9
S, Apatity array	-1.0	19.8
Rg, Apatity array	-1.8	3.4

 Table 7.2.6. Mean and standard deviation of azimuth residuals relative to Khibiny mine locations.



Fig. 7.2.1. In the upper part, a large reference area is shown. The location of the ARCESS array is given by a filled circle, and the location of the Khibiny Massif region is shown. The lower part shows a detailed picture of the Khibiny Massif region. The locations of the six mining sites are given by large numbers 1-6. The Apatity array is shown as a filled circle and the three-component station in the town of Apatity is shown as a large triangle.

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Fig. 7.2.2. Seismograms from an event at the underground Mine I recorded at the Apatity three-component station. The source-receiver distance is 18 km. The upper trace, emphasizing the P-phase, is the z-component prewhitened with a low-order prediction error filter designed from a 25 s noise sample preceding the P-phase. The three-component data rotated with the azimuth to the mine can be seen at traces 2-4. The data are filtered in a relatively wide passband of 2 - 8 Hz, and these three traces clearly show an instantaneous S-phase. The lower trace is bandpass filtered between 0.8 and 2.0 Hz, and clearly illustrate the low-frequency Rg phase.

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Fig. 7.2.3. Seismograms from an event at mine VI recorded at the Apatity array. The source-receiver distance is 49 km. The only difference from the traces of Fig. 7.2.2 is that the upper trace is the array beam steered with steering delays corresponding to the slowness and azimuth of the P-arrival. Notice that in this case the S-phase is much more emergent than the S-phase of Fig. 7.2.2.

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Fig. 7.2.4. ARCESS recording of the event also shown in Fig. 7.2.2. The source-receiver distance is 394 km. The upper trace is the array beam steered with the slowness and azimuth of the P-phase. The beam is prewhitened with a low-order prediction error filter. The lower three traces are the rotated three-component data at the central element (ARA0) of the ARCESS array. The data are bandpass filtered between 2 and 5 Hz. Notice that both Pn, Pg, Sn and Lg are clearly seen, and that Sn and Lg are most prominent on the transverse component.



Fig. 7.2.5. These two plots show the time differences between the automatic P-onsets and the manual pick at APAO as a function of signal-to-noise ratio (SNR) on the prewhitened beam. All 58 P-observations are included. Figure a) shows the time differences between the automatic P-onsets from the IMS and the manual picks, whereas figure b) shows the time differences between the automatic P-onsets from ESTON1 and the manual picks. The two vertical dotted lines represent ±σ2.



Fig. 7.2.6. Plots similar to those of Fig. 7.2.5, but in this case for Pn-onsets at ARCESS. For one of the Khibiny Massif events, the ARCESS array was not operational, and consequently only 57 P-phases were analyzed.



Fig. 7.2.7. This figure show the time differences between the automatic P-onsets from ESTON1 and the manual pick at both Apatity stations (APA0 and APZ9) as a function of signal-to-noise ratio (SNR). Compared to Fig. 7.2.5 the x-axis is strongly expanded. Notice that P-phases at APZ9 (triangles) generally have lower SNR. The average dominant frequency of these P-phases is 13 Hz.



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Fig. 7.2.8. Plot similar to Fig. 7.2.7, but in this case for ARCESS P. The average dominant frequency is 6 Hz.



Fig. 7.2.9. The seismograms of this figure are Apatity array (APA0) recordings of the Mine I event also given in Fig. 7.2.2. The source-receiver distance is 32 km. Trace no. 2 is the prewhitened P-beam used for onset-time estimation using ESTON1. The upper trace gives the likelihood function from ESTON1 after processing an interval of ± 2 s around the initial P-onset estimate, and the peak of this likelihood function correspond to the estimated onset. The lower trace is the vertical component APAO_sz filtered in a low passband (0.8 - 2.0 Hz) to enhance the Rg phase, and the trace above it is the STA envelope. The peak of this envelope is defined as the peak of the Rg phase. After the P-onset and the Rg maximum is found, we are searching for the S-onset using the three-component ESTON3 estimator. The search interval starts 2 s after the P-onset and stops at the Rg peak, as seen from the ESTON3 likelihood function of trace no. 3. The three-component data processed by ESTON3 are given in traces 4-7. Notice that there is no need to rotate the three-component data before using ESTON3, but to visualize that the S-phase has the largest SNR on the transverse component, we have in this figure rotated the data. Notice that both the ESTON1 and the ESTON3 likelihood functions show clear peaks at the P- and Sonsets.

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Fig. 7.2.10. The traces of this plot are similar to those of Fig. 7.2.9. The seismograms are Apatity array (APA0) recordings of the Mine VI event also shown in Fig. 7.2.3. The source-receiver distance is 49 km. Notice that the S-phase is much more emergent than the S-phase of Fig. 7.2.9, but the ESTON3 likelihood function still shows a clear peak at the S-onset.

7.3 Intelligent post-processing of seismic events -- Part 3: Precise relocation of events in a known target region

Introduction

From experience with analyst review of events automatically defined by the Intelligent Monitoring System (IMS) (Bache et al, 1993), we have realized that the quality of the automatic event locations can be significantly improved if the event intervals are reprocessed with signal processing parameters tuned to phases from events in the given region. The tuned processing parameters are obtained from off-line analysis of events located in the region of interest. The primary goal of such intelligent post-processing is to provide event definitions of a quality that minimizes the need for subsequent analysis and review.

In this third paper on intelligent post-processing of seismic events we test an event location method that makes use of accurate retiming of seismic phases and accurate re-estimates of azimuth using broad-band analysis of array data. The key to obtaining such accurate estimates is the knowledge, provided by the IMS, of an approximate event location in a known area with good calibration information. Various aspects of generalizing the method and incorporating it into the IMS are also discussed.

Location of regional events

Seismic event location has traditionally relied upon P-wave arrival times from a large number of stations. For small events observed only at regional distances, it has been necessary to include arrival times of later phases as well as arrival azimuths in order to obtain acceptable location accuracy (Bratt and Bache, 1988; Thurber et al, 1989). With a sparse network, the number of arrival time observations may often be so low that azimuth information is given considerable weight in the location scheme. However, as the number of observations increases, the relative importance of the azimuths decreases sharply, in view of the generally large uncertainties (10 degrees or more) that are associated with azimuth estimates (e.g., Suteau-Henson, 1990; Bame et al, 1990).

Kværna and Ringdal (1986) demonstrated that the uncertainty of azimuth estimates from regional arrays can be greatly reduced by applying a broad-band estimation scheme in which the frequency band is kept constant for a given epicentral region. Thus, for NOR-ESS observations from Blåsjø explosions (distance 300 km) with high SNR, they found an azimuth standard deviation of only 1.0 degrees for Pn phases and 1.5 degrees for Sn and Lg phases.

This observation suggests that given an initial (approximate) event epicenter, it should be feasible to determine an optimum frequency band (based upon previous observations) and apply broad-band F-k analysis in order to obtain very accurate azimuth estimates, which can then be used to refine the event location. The basic idea of this paper is to establish and test a method for location refinement which takes into account both the stable azimuth estimates obtained by fixed frequency broad-band analysis and the accurate timing observations described in Section 7.2. We use as an example IMS-reported events in the Khibiny Massif recorded by the nearby Apatity array and the more distant ARCESS array. We

show that a significant improvement can be obtained relative to the current IMS processing. Moreover, the method is completely automatic and can easily be incorporated into the IMS, as illustrated in Fig. 7.3.1.

As shown in Section 7.1, the majority of seismic events in Fennoscandia can be associated with mining activity at a few sites. The Khibiny Massif is one of the mining areas with the most events. Obtaining accurate automatic solutions for events from this and other mining areas will lead to a considerably lower workload for the analyst, and will consequently allow more time to be spent analyzing events of particular interest.

Method

The automatic method applied in this study consists of the following steps:

- 1. The IMS detects and locates a seismic event within a given distance from the target region.
- 2. For each phase detection by a station in the network, the estimated arrival time is refined using the method of Pisarenko et al (1987) (see Section 7.2). A standard deviation is assigned to each refined estimate.
- 3. For each phase detection by one of the arrays, an "optimum" frequency band for F-k analysis is extracted from a data base which we assume has been previously established and calibrated.
- 4. Broad-band F-k analysis is then applied to each phase. The resulting values are corrected for systematic bias and assigned a standard deviation, based upon SNR, phase type and previously observed calibration information.
- 5. The LocSat program (Bratt and Bache, 1988) is then applied to the revised data set, and a new event location estimate is obtained. In practice, zero depth is assumed a priori in the automatic process.

In the case study discussed in detail in this paper, we will analyze events in the Khibiny Massif recorded by IMS, for which we have independent location information. We will primarily make use of the Apatity array, using arrival times for the P and S phases and azimuths derived from the Rg phase in order to refine the location estimates as described above. In addition, we will consider possible improvement in the location accuracy when using additional available stations.

Data

The data for this study comprises 58 events at 6 mines in the Khibiny Massif. The mines are only a few kilometers apart (Mykkeltveit, 1992), and they each have a dimension ranging from a few hundred meters to about 1 km. These dimensions are small enough so that we ignore the areal extent of each mine. A list of the event parameters is given in Section 7.2.

To obtain an initial location, we have used the automatic IMS epicenter solution for each event. The automatic IMS solution relative to the "true" epicenter for each of the events is

displayed graphically in Fig. 7.3.2. Although many of the events are very accurately located (to within a few km), there is a fair amount of scatter in the location estimates. The median "error" is 10.6 km. In a few cases, the location is wrong by several tens of km; this is due to occasional erroneous phase identification by the automatic process.

The Apatity array has been described by Mykkeltveit et al (1992), and initial data analysis results have been presented by Ringdal and Fyen (1992). The latter paper discusses in particular the high stability of azimuth estimates using the Rg phase for shallow events at local distances.

In Section 7.2, some examples of Apatity array recordings of events in the Khibiny Massif are shown. The filtered recordings (0.8-2.0 Hz) are dominated by the Rg phase. In contrast, the P and S phases are far more high-frequent, as is to be expected at such close distances. Because of the low frequency of the Rg phase, this phase has a high coherency across the array and thus provides more stable slowness estimates than are obtained from the other phases.

For the present study, we have used Apatity array recordings for all 58 events. For all but six events, high-frequency three-component data from the array site have been available. Furthermore, all events have associated 3-component broadband recordings from the station installed in the town of Apatity, 15 km from the array site. ARCESS array data have been available for all but one of the events.

Results

In the following, we present results from analyzing the 58 events in the data base under various scenarios. Each scenario assumes that we have an operational situation in which a certain amount of calibration information has been assembled.

a) Use of Apatity array only; P and S arrival times, Rg azimuth, no calibration.

In this scenario, we assume no prior knowledge except for the "optimum" frequency bands for timing and azimuth estimation. The travel-time tables are those that the IMS uses for Fennoscandia in general, with no regional corrections.

The results are displayed in Figs. 7.3.3. We note that there is a small systematic bias, as there seems to be a trend of events being located to the north and east of the sites. This indicates that some improvement could be obtained by introducing a regional velocity model for the Khibiny area, together with regional azimuthal corrections.

Nevertheless, the scatter relative to the 6 mines in this plot (using uncalibrated Apatity array data only) is significantly reduced compared to the IMS results plotted in Fig. 7.3.2.

The improvement can be quantified by considering that the median "error" is only 3.1 km, with a worst case "error" of 6.7 km. This can be compared to the IMS results of 10.6 km and 75.3 km, respectively. Thus the reprocessing, even without regional correction, could significantly improve the precision in the epicenter solutions if incorporated in the automatic IMS processing.

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b) Use of Apatity array data only, with a regional P-S travel-time bias correction and a correction for systematic Rg azimuth bias.

The resulting location plot is shown in Fig. 7.3.4. The median error is reduced to 2.1 km, and the "worst case" error is 6.1 km.

c) Use of Apatity array data plus P and S times from the 3-component station in Apatity. Regional corrections have been applied both for the array and the 3-comp. station.

The resulting location plot (Fig. 7.3.5) shows that the median error is 1.4 km, and the maximum error is 5.4 km.

d) This final scenario includes both Apatity array data, Apatity 3-comp. data and ARCESS P arrival times. Regional corrections have been applied for all the data.

The resulting location plot (Fig. 7.3.6) shows a median error of 1.6 km, and a maximum error of 7.1 km. It is interesting to note that this is not quite as good as case c). Thus, it is not necessarily an improvement to add data, even if (as in this case) the additional data are extremely accurate (error $\leq \pm 0.05$ seconds on P-recordings).

We should note that the travel-time corrections used in this paper are *relative* and based upon the same data set that we have evaluated. In view of the large number of events, the possible bias introduced by this procedure should be negligible. As the origin time of the Khibiny events are unknown, we have fixed the travel-time correction of P at the Apatity three-component station APZ9 to 0. The P phase at APZ0 has travelled a shorter distance and spent less time in the earth than the other phases used in the event locations. Thus, the influence of an erroneous travel-time model is likely to be the least for this phase.

Travel-time corrections:

Ρ	APZ9	0 s (fixed)
P	APA0	-0.10 s
Pn	ARCESS	-0.31 s
S	APZ9	0.22 s
S	APA0	0.09 s

Conclusions

In this study we have conducted extensive off-line analysis of 58 mining explosions in the Khibiny Massif recorded at the Apatity array. Independent locations of these explosions are provided by seismologists from the Kola Science Centre. Most of these events show clear P, S and Rg phases at the nearby Apatity array located 30-50 km away from the mining areas and the events have also been detected by the ARCESS array. By using the onset-time estimator of Pisarenko et al (1987) and comparing to the best manual pick, we have found that P-onsets can be automatically estimated with an accuracy of better than 0.1 seconds, and S-onsets with an accuracy better than 0.5 seconds. In addition, the azimuth estimates from F-k analysis of the P and Rg phases show to be accurate well within ± 5 degrees after removing the biases. The key to achieving stable azimuth estimates is to

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process the data in a fixed frequency band, using a fixed time window positioning, as demonstrated by Kværna and Ringdal (1986) for Blåsjø recordings at NORESS.

Our observations suggest that based on data from the Apatity array alone, we are able to locate these events (assuming 0 depth) with a median error of about 3 km relative to the true location. Even better accuracy can be achieved using calibration information, i.e., correcting the azimuth and arrival time observations for systematic bias. The excellent precision of the automatic phase onsets and azimuth estimates also indicate that the need for subsequent manual analysis of these events may be eliminated.

After we have established the tuned processing parameters for events in this region, a natural next step will be to automatically activate such intelligent post-processing every time the IMS locates an event in the Khibiny region. Extensions of the method to other mining areas will also be considered.

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New processing flow:



Fig. 7.3.1. Schematic view of the principle behind intelligent post-processing of seismic events.

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Fig. 7.3.2. Location of the six mining sites in the Khibiny Massif (large numbers 1-6) and the locations of the 58 reference events (small numbers 1-6) as given by the automatic IMS processing. In the **upper part**, a large reference area is shown, with the mines plotted as filled squares. The **lower part** shows a detailed picture for the area near the mines. The small number (1-6) associated with each event represents the mine in which the event actually occurred. The Apatity array is shown as a filled circle and the three-comp. station in the town of Apatity is shown as a filled triangle.

Apatity array location errors (uncalibrated) Median: 3.1 km 90% quantile: 4.9 km Maximum: 6.7 km 31 32° 33° 34° 35° 36° 68 68° Array 3comp 50 km 67' ± 67° 33. 36° 31 32° 34° 35° 33° 30' 34° 00' 67' 42' 67° 42' 66 **6**⁶ 55 67' 36' 67° 36' 5 km 33° 30' 34° 00'

Fig. 7.3.3. Same as Fig. 7.3.2, but the event locations (small numbers) have now been taken from the post-processing results using uncalibrated Apatity array data only.

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Fig. 7.3.4. Same as Fig. 7.3.2, but with the event locations resulting from post-processing of calibrated Apatity array data (see text for details).

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tity array data and 3-comp. data (both calibrated) in the post-processing.

Fig. 7.3.5. Same as Fig. 7.3.2, but with the event locations resulting from combining Apa-

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Fig. 7.3.6. Same as Fig. 7.3.2, but with the event locations resulting from combining Apatity array, Apatity 3-comp. and ARCESS array data (all calibrated) in the post-processing.

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7.4 Monitoring a moratorium: An experiment in continuous seismic threshold monitoring of the northern Novaya Zemlya test site

Introduction

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After 1 October 1992 a moratorium on nuclear testing has been in effect for the US, UK, Russia and France. On this background, we have applied the continuous threshold monitoring technique to the northern Novaya Zemlya test site for a full six-month period (1 October - 31 March 1993), using the NORESS, ARCESS and FINESA regional arrays. Starting 1 December 1992, we have compiled daily statistics of all peaks on the threshold diagram exceeding $m_b = 2.75$, and associated these peaks to regional or teleseismic events whenever possible. In addition, we have analyzed smaller peaks (below $m_b = 2.75$) that can possibly be associated with Novaya Zemlya epicenters.

The theoretical background for and applications of the continuous seismic threshold monitoring metod (CSTM) have been described in several articles. The approach was introduced by Ringdal and Kværna (1989), who showed that by continuously monitoring the seismic amplitude level at several seismic stations or arrays, one can at any time obtain an instant network-based magnitude threshold for a given target region. The magnitude threshold can be interpreted as the maximum magnitude of a possible clandestine explosion, given a predefined level of confidence. In the context of a comprehensive or threshold test ban treaty, the continuous assessment of the magnitude thresholds makes it possible to focus attention upon those specific time intervals when realistic evasion opportunities exist, while retaining confidence that no treaty violation has occurred at other ti.nes.

Previous results from experimental monitoring

Kværna (1992) presented results from a one-month experiment of continuously monitoring the northern Novaya Zemlya test site. Data from the Fennoscandian regional array network (ARCESS, FINESA, and NORESS), see Fig. 7.4.1, were used to calculate the magnitude thresholds. It was found that the test site could be consistently monitored at a very low magnitude level (typically $m_b = 2.5$). In fact, every occurrence of the threshold exceeding $m_b = 2.6$ could be explained as resulting from an identified interfering event signal either at teleseismic or regional distance, except for three instances when a short gap in ARCESS recording caused the network threshold to increase.

The excellent capability of the Fennoscandian regional array network to monitor the northern Novaya Zemlya test site was further confirmed by an experiment where recordings of the Novaya Zemlya nuclear test of October 24, 1990 were downscaled to $m_b = 2.6$ and superimposed on different noise intervals (Kværna, 1991).

In the context of using CSTM as a tool in routine monitoring, it is important to determine how the method will work under different conditions. Variability in the seismic noise level, occurrences of large earthquakes and aftershock sequences, station downtimes and data quality problems are all factors that will influence the performance of CSTM.

Analysis of network threshold peaks

Our monitoring experiment was conducted in the same way and with the same parameter settings as used by Kværna (1992). In Kværna (1992) the monitoring results were presented in terms of plots covering one data day each. In Figs. A-1 to A-29 of the Appendix of that report, each covering one day of February, 1992, all time periods where the network magnitude thresholds at the 90% confidence level exceeded $m_b = 2.6$ were identified. A similar approach was used for this experiment.

For the remainder of this paper, the term magnitude threshold implies the magnitude threshold at the 90% confidence level.

Figs. 7.4.2 shows a typical example of a one-day plot. The upper three traces of each figure represent the magnitude thresholds obtained from the three indivitual arrays, whereas the bottom trace illustrates the network threshold. Typically, the individual array traces have a number of significant peaks for each 24-hour period, due to signals from interfering events (regional or teleseismic). On the network trace, the number and sizes of these peaks are significantly reduced, because an interfering event usually will not provide matching signals at all stations. From probabilistic considerations, it can in such cases be inferred that the actual network threshold is lower than these individual peaks might indicate.

The arrows on the one-day threshold plots indicate peaks with network magnitude threshold exceeding $m_b = 2.75$. A T at the arrow indicates that the peak is caused by signals from a teleseismic event, whereas an **R** indicates signals from a regional or local event. Fig. 7.4.3 shows summary statistics for one data day, with an explanation of all threshold peaks exceeding 2.75. Such summary statistics were generated for each day during the four-month period, 1 December 1992 - 31 March 1993.

The peaks in the threshold traces are caused by either large teleseismic events or by regional events. The regional events were all processed and located by the Intelligent Monitoring System (IMS) (Bache et al., 1993), and the teleseismic events were located either by the IMS or by the QED service of the USGS.

Fig. 7.4.4 shows a histogram of the number of peaks exceeding given magnitude thresholds. During the entire four-month period, there were only 40 peaks exceeding $m_b = 3.0$. Each of these peaks could be unambiguously associated with either a regional or teleseismic event. Consequently, at the specified confidence level, we can state that no seismic event of $m_b > 3.0$ occurred at the test site during this four-month period.

The event at Novaya Zemlya on 31 December 1992

Figs. 7.4.5 and 7.4.6 show the threshold plots and peak statistics for 31 December 1992. At 0929 GMT that day, a peak occurred corresponding to an event at Novaya Zemlya, located by the IMS. This peak (which is the only one associated to a Novaya Zemlya event during the six-month interval) had an upper magnitude limit of 2.6. The actual event magnitude, according to Carter et al (1993) was 2.5.

This event was detected by ARCESS (P and S), Spitsbergen (P and S), NORESS (P) and Apatity (S). In addition, the Kola Science Centre provided readings for the station Amderma (Pg and Sn) made from analog recordings. Tables 7.4.1 and 7.4.2 summarize the available observations for this event. Table 7.4.3 gives the results of applying the LOCSAT program (Bratt and Bache, 1988) to this parameter set.

Our results indicate that the epicenter was slightly to the north of the test site. However, there are some uncertainties in the travel time tables, and an on-site location cannot entirely be ruled out. We are hesitant to introduce travel-time corrections in this case, since no good reference event is available for some of the key stations.

A plot of the IMS-processed traces is shown in Fig. 7.4.7. Notice in particular the high SNR at the Spitsbergen B2 single sensor (filter band 8-16 Hz) and on the ARCESS array beam. Fig. 7.4.8 shows the IMS solution for this event. The Amderma station is not included in the IMS processing, but adding its data does not cause any significant change in the event location.

Conclusions

This work has documented the practical capability of the Continuous Seismic Threshold Monitoring method to monitor a specific nuclear test site at a very low threshold over an extended time period.

Specifically, we have used the Fennoscandian array network (NORESS, ARCESS and FINESA) to monitor the northern Novaya Zemlya test site for a full four-month period at a threshold of $m_b = 2.75$. We have identified only one instance where an event close to this threshold has occurred near the test site. In fact, the event magnitude (2.5) was below our target threshold, but the peak was still easily identified on the threshold trace.

Recently, additional array stations have been installed or are planned for installation in the Arctic region. These stations would contribute to further improving the CSTM capability, both for Novaya Zemlya and on a general regional basis. This will be the subject for additional studies in the future.

T. Kværna F. Ringdal

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Station	Lat (^o N)	Lon (°E)	Type of Station
AMD (Amderma)	69.742	61.655	3-comp analog
APA (Apatity)	67.600	33.000	Аттау
ARAO (ARCESS)	69.535	25.506	Аттау
FIAO (FINESA)	61.444	26.079	Array
NRA0 (NORESS)	60.735	11.541	Аттау
SVA (Spitsbergen)	78.180	16.350	Агтау

Table 7.4.1. Location of stations used in this study.

Station	Phase	Type ¹⁾	Arrival time	Azimuth	St. dev. ²⁾
AMD	Pg	t	09.30.43.7	-	2.0
AMD	Sn	t	09.31.20.7	-	2.0
APA	Sn	t	09.33.22.0	-	2.0
ARC	Pn	t	09.31.48.7	-	1.0
ARC	Pn	a	-	62.5	15.0
ARC	Sn	t	09.33.37.2	-	2.0
ARC	Sn	a	-	58.4	15.0
NRS	Р	t	09.34.04.3	-	1.0
NRS	Р	a	-	24.0	15.0
SVA	Pn	t	09.31.50.7	-	1.0
SVA	Sn	t	09.33.41.7	•	2.0

Table 7.4.2. Observed arrivals for the 31 Dec 92 event.

1) t = time, a = azimuth

2) A priori standard deviation in seconds (time) or degrees (azimuth)

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Final location estimate (+/- S.D.): Latitude: 73.620 deg. +/-6.673 km. Longitude: 55.196 deg. +/-7.177 km. 0.000 km. +/-0.000 km. (Fixed) Depth: Relative O.T.: -79.301 sec. +/-Absolute O.T.: -79.301 sec. +/-0.849 sec. 0.849 sec. : 1969 12 31 23:58:40.70 Confidence region at 0.90 level: Semi-major axis: 18.1 km. = 0.16 deg. Semi-minor axis: 10.6 km. = 0.10 dag. Major-axis strike: 49.3 deg. clockwise from North Orig. time error: 1.4 sec. Standard errors (sigma): Prior: 1.00 (9999 deg. of freedom) Posterior: 1.00 (10007 deg. of freedom) Posterior: 0.65 (Normalized sample S.D.) Azimuthal weighting: 1.00 Effective rank of matrix: 2.00 Maximum azimuthal GAP: 195 deg. - No damping required ! -----Data Residuals Distance Azimuth Data Ariv ID Statn Phase Type at True Normalized (deg.) (deg.) Import Err ____ ____ --245771 APA -0.044 9.417 241.62 Sn -0.087 0.314 ŧ d 4.371 149.27 245771 ND 0.900 0.450 0.319 Pa £ d 245771 MD Sn t d -1.495 1.068 20.495 254.54 -0.652 20.495 254.54 -0.317 10.096 26 -0.748 4.371 149.27 0.455 245782 NRS 1.068 0.016 ₽ t d 245782 NRS P d -9.782 0.000 a 245782 ARC -0.317 0.055 Pn t d -0.317 245782 ARC 9.672 0.645 10.096 261.08 0.000 Pn a d 245782 ARC Sn t d -1.009 -0.504 10.096 261.08 0.220 245782 ARC 10.096 261.08 5.572 0.371 0.000 Sn a d 245782 SVA Pn đ -0.320 -0.320 10.244 313.75 0.408 t 245782 SVA Sn t d -0.033 -0.016 10.244 313.75 0.211

Table 7.4.3. Epicenter solution for the 31 Dec 92 event at Novaya Zemlya using the LOC-SAT program (Bratt and Bache, 1988). The depth has been constrained to 0. See Table 7.4.1 for station locations and Table 7.4.2 for observed arrival data and assumed a priori standard deviations.



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Fig. 7.4.1. Location of the target area (the northern Novaya Zemlya test site) for the monitoring experiment. The locations of the three arrays NORESS ($\Delta = 2280$ km), ARCESS ($\Delta = 1100$ km) and FINESA ($\Delta = 1780$ km) are indicated.



Fig. 7.4.2. Example of continuous threshold monitoring of Novaya Zemlya for one day (13 February 1993). A "threshold trace" is shown for each of the 3 arrays, and the combined network threshold trace is shown at the bottom. Note that the network threshold is well below magnitude 2.5 almost all the time.

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Date: February 13, 1993 Day_of_year: 1993-044

	# peaks	# seconds	% of time
Mag >2.50	4	335	0.39
Mag >2.75	2	67	0.078
Mag >3.00	0	0	0
Mag >3.50	0	0	0

Individual Peaks > 2.75

Explanation	Kola Peninsula, Russia, probable explosion	Rat Islands, Aleutian Islands, earthquake			
Agency	IMS	SMI			
Mag	2.4	5.4			
Depth	0	33F			
Lon	32.9E	176.4E			
Lat	68.1N	51.1N			
Or.time	10.37.16	21.19.35			
Reg Tele	R	Т			
# sec > 2.75	27	4			
TM _{time}	10.38.10	21.26.57			
TM _{max}	2.89	2.99			

Fig. 7.4.3. Characterization of individual peaks in the threshold plot for 13 February 1993 (Fig. 7.4.2). The two peaks exceeding m_b = 2.75 are due to an event in the Kola Peninsula and an earthquake in the Aleutian Islands.

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Fig. 7.4.4. Histogram showing the distribution of peaks on the network threshold trace for the four-month monitoring experiment.





Fig. 7.4.5. Threshold plot for 31 Dec 92 (see Fig. 7.4.3 for explanation). The small Novaya Zemlya event at 09.29 GMT is indicated.

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Date: December 31, 1992 Day_of_year: 1992-366

	# peaks	# seconds	% of time
Mag >2.50	14	648	0.75
Mag >2.75	ø	107	0.12
Mag >3.00	2	19	0.022
Mag >3.50	1	2	0.0081

Individual Peaks > 2.75

Explanation		Bonin Islands Region	Novaya Zemlya	North Pacific Ocean	Gap in ARCESS and NORESS recording. Local noise at FINESA.	Gap in ARCESS and NORESS recording. Local noise at FINESA.	Local event at FINESA.	Local event at FINESA.
Agency		SMI	SMI	IMS				
Mag		4.9	2.3	4.7				
Depth		33F	OF	33F				
Lon		138.8E	55.2E	146.9E				
Lat		27.4N	73.6N	22.3N				
Ortime		07.25.10	09.29.24	32.01.06				
Reg	Tele	ч	R	т			R	ĸ
#sec > 2.75		80	0	38	31	11	5	14
TM _{time}		07.34.10	09.29.23	23.09.49	23.27.45	23.34.38	23.42.23	23.53.35
TMmax		2.80	2.64	3.89	3.02	2.88	2.79	2.92

Fig. 7.4.6. Threshold monitoring statistics for 31 Dec 92. Note that there were some data problems just before midnight, causing a rise in the threshold.

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Waveforms of Novaya Zemiya Event 31 Dec 1992 - Magnitude 2.5

Fig. 7.4.7. Waveform plots for 5 regional arrays for the Novaya Zemlya event on 31 Dec 92. There are P-phase detections at ARCESS, Svalbard (Spitsbergen) and NORESS, and S-phase detections at Apatity, ARCESS and Svalbard.

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Novaya Zemiya event 31 Dec 1992

Fig. 7.4.8. IMS solution and associated error ellipse for the Novaya Zemlya event ($m_b \sim 2.5$) on 31 Dec 92.

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7.5 Initial processing results from the Spitsbergen small-aperture array

During the fall of 1992, a new regional array was installed near the town of Longyearbyen on the Arctic island of Spitsbergen (see Mykkeltveit et al, 1992). Fig. 7.5.1 shows the location of this array and Fig. 7.5.2 shows the array geometry. As of early 1993, six of the nine array sites had been finalized, and this study is based on these six instruments.

In the following we report on some initial results from analyzing data from the new array. It is emphasized that these results are preliminary, since the available data cover only a few months, and only a partial array has been installed until now. Thus a more comprehensive assessment must await the collection of more complete array data over a longer time span.

Noise spectra

Fig. 7.5.3a shows an example of corrected noise spectra for the 6 vertical elements of the Spitsbergen array, taken at 00.00 hours (GMT) on day 345, 1992. For comparison, the ARCESS spectra for the same number of channels and taken at the same time are shown in Fig. 7.5.3b. From these figures, it is seen that the Spitsbergen array has practically the same noise level as ARCESS, especially for the "best" Spitsbergen sensor (B2). However, the Spitsbergen array has a much larger variation in noise level than ARCESS, especially at low frequencies.

Fig. 7.5.4a shows 41 corrected noise spectra, taken daily at 00.00 GMT during the period 10 December 1992 - 18 January 1993 for the Spitsbergen array (instrument B2). These spectra represent typical winter conditions. For comparison, a similar set of spectra during summer (23 May - 7 June) is shown in Fig. 7.5.4.b. The range is similar for the two time periods, although the highest noise levels at low frequencies are observed during winter.

Noise suppression

Previous studies have shown that regional arrays are very effective in suppressing seismic noise, thus providing gains that are often of the order of \sqrt{N} , or sometimes even in excess of this value (N being the number of sensors). Such better than \sqrt{N} suppression occurs when particular subgeometries are chosen, enhancing the suppression of noise at certain frequencies. As a first check on the capabilities of the Spitsbergen array in this regard, we have computed noise suppression curves for the initial array geometry.

We calculate an uncorrected power density spectrum first by prewhitening 60 seconds of data. Then we estimate the autocorrelation for 6 partially overlapping windows (window length 12 seconds), and compute power density spectra from the average autocorrelation, with compensation for the prewhitening filter.

An average spectrum is obtained by averaging the individual channel spectra for the array. The averaging is done after a logarithmic transform of the spectra, and the standard deviation for each frequency point is calculated. Each spectrum is pointwise compared to the average spectrum. If a value is outside 1.5 standard deviation from the mean value, the point is considered an outlier. If a single channel spectrum has more than 60% outliers, the spectrum is excluded and a new average spectrum is estimated. A beam is calculated using only those channels for which the spectrum was accepted. The suppression is then the beam spectrum divided by the average spectrum.

Fig. 7.5.5 shows noise suppression in the frequency range 0-20 Hz for an infinite-velocity beam (no time delays) for the Spitsbergen array. The figure shows an average of 24 noise samples, taken hourly on day 345, 1992. We note that the noise suppression is between 5 and 10 dB above 2 Hz. This is about the theoretical value (8 dB for 6 sensors). We anticipate further investigations into this subject after the full array is deployed.

Detection processing

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Since the array was installed, the Spitsbergen data have been subjected to continuous online detection processing (DP) at NORSAR using the standard small-array detection algorithm. The initial beam deployment is shown in Table 7.5.1.

As an example of a regional seismic event recorded and processed using Apatity data, we show here results for a low-magnitude earthquake ($m_b \sim 2.0$) west of the array (on the North Atlantic Ridge).

Fig. 7.5.6 shows individual recordings for this event. Note the large variations in signal amplitudes across the array. In particular, the B2 instrument shows excellent signal-to-noise ratio for this event as well as for other events we have processed. The reason for this large amplitude variation is as yet unknown. There are two distinct P-onsets and one clear S-onset on the waveforms.

Frequency-wavenumber analysis

Frequency-wavenumber solutions for P and S are shown in Fig. 7.5.7. In spite of the small aperture of the array and the incomplete deployment, the peaks in the F-K diagram are well-defined. However, there is a tendency to side lobes in the plots.

From Fig. 7.5.7 we note that the phase velocities of each of the two phases are reasonably consistent with the phase type, the S phase having the slowest velocity. The estimated azimuths are quite consistent (256 and 260 degrees, respectively).

In conclusion, our preliminary analysis indicates that the Spitsbergen array will be an important supplement to the seismic array network in Northern Europe. Its inclusion into the Intelligent Monitoring System (IMS) will in particular serve to improve the location precision and source characterization of the large number of events in the Barents Sea and adjacent regions. The excellent recordings of this array from the small Novaya Zemlya event on 31 Dec 1992 (see subsection 7.4) is a further illustration of this point. Further

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evaluation of the array capabilities will be made after the 9-element array has been completed.

J. Fyen F. Ringdal

References

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BEAM	Velocity	Azimuth	Filter	band	Threshold	N	Configuration
S011	99999.9	0.0	0.5 -	1.5	4.4	6	AOB
S021	99999.9	0.0	1.0 -	3.0	4.4	6	AOB
S031	99999.9	0.0	1.5 -	3.5	4.4	6	AOB
S032	11.0	30.0	1.5 -	3.5	4.4	6	AOB
S033	11.0	90.0	1.5 -	3.5	4.4	б	AOB
S034	11.0	150.0	1.5 -	3.5	4.4	6	AOB
S035	11.0	210.0	1.5 -	3.5	4.4	6	AOB
S036	11.0	270.0	1.5 -	3.5	4.4	6	AOB
S037	11.0	330.0	1.5 -	3.5	4.4	6	AOB
S038	15.0	88.0	1.5 -	3.5	3.9	6	AOB
S039	10.0	95.0	1.5 -	3.5	3.9	6	AOB
S041	99999.9	0.0	2.0 -	4.0	4.4	9	AOAB
S042	10.2	30.0	2.0 -	4.0	4.4	9	AOAB
S043	10.2	90.0	2.0 -	4.0	4.4	9	AOAB
S044	10.2	150.0	2.0 -	4.0	4.4	9	AOAB
S045	10.2	210.0	2.0 -	4.0	4.4	9	AOAB
S046	10.2	270.0	2.0 -	4.0	4.4	9	AOAB
S047	10.20	330.0	2.0 -	4.0	4.4	9	AOAB
S048	15.0	88.0	2.0 -	4.0	3.9	9	AOAB
S049	10.0	95.0	2.0 -	4.0	3.9	9	AOAB
S051	99999.9	0.0	2.5 -	4.5	4.4	9	AOAB
S052	8.9	30.0	2.5 -	4.5	4.4	9	AOAB
S053	8.9	90.0	2.5 -	4.5	4.4	9	AOAB
S054	8.9	150.0	2.5 -	4.5	4.4	9	AOAB
S055	8.9	210.0	2.5 -	4.5	4.4	9	AOAB
S056	8.9	270.0	2.5 -	4.5	4.4	9	AOAB
S057	8.9	330.0	2.5 -	4.5	4.4	9	AOAB
S058	15.0	88.0	2.5 -	4.5	3.9	9	AOAB
S059	10.0	95.0	2.5 -	4.5	3.9	9	AOAB
S061	99999.9	0.0	3.0 -	5.0	4.4	9	AOAB
S062	10.5	30.0	3.0 -	5.0	4.4	9	AOAB
S063	10.5	90.0	3.0 -	5.0	4.4	9	AOAB
S064	10.5	150.0	3.0 -	5.0	4.4	9	AOAB
S065	10.5	210.0	3.0 -	5.0	4.4	9	AOAB
S066	10.5	270.0	3.0 -	5.0	4.4	9	AOAB
S067	10.5	330.0	3.0 -	5.0	4.4	9	AOAB
S068	15.0	88.0	3.0 -	5.0	3.9	9	AOAB
S069	10.0	95.0	3.0 -	5.0	3.9	9	AOAB
S071	99999.9	0.0	3.5 -	5.5	4.4	9	AOAB
S072	11.1	30.0	3.5 -	5.5	4.4	9	AOAB
S073	11.1	90.0	3.5 -	5.5	4.4	9	AOAB
S074	11.1	150.0	3.5 -	5.5	4.4	9	AOAB
S075	11.1	210.0	3.5 -	5.5	4.4	9	AOAB
S076	11.1	270.0	3.5 -	5.5	4.4	9	AOAB

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Table 7.5.1. Spitsbergen beam table, valid from 1992/328 (23 November 1992). The table shows the name of the beam, velocity (km/sec), azimuth (degrees), filter band (Hz), STA/LTA threshold, and configuration. The configuration is described with number of sensors and a configuration code. Here, AOAB means center A0 SPZ plus A-ring plus B-ring, and A0B means A0 SPZ plus B-ring. SI01 - SI06 are incoherent beams using SPZ channels only. (Page 1 of 2)

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S077	11.1	330.0	3.5 - 5.5	4.4	9	AOAB	
S081	9999 9.9	0.0	4.0 - 8.0	4.4	9	AOAB	
S082	9.5	30.0	4.0 - 8.0	4.4	9	AOAB	
S083	9.5	90.0	4.0 - 8.0	4.4	9	AOAB	
S084	9.5	150.0	4.0 - 8.0	4.4	9	AOAB	
S085	9.5	210.0	4.0 - 8.0	4.4	9	AOAB	
S086	9.5	270.0	4.0 - 8.0	4.4	9	AOAB	
S087	9.5	330.0	4.0 - 8.0	4.4	9	AOAB	
S091	99999.9	0.0	5.0 - 10.0	4.9	9	AOAB	
S092	10.5	30.0	5.0 - 10.0	4.9	9	AOAB	
S093	10.5	90.0	5.0 - 10.0	4.9	9	AOAB	
S094	10.5	150.0	5.0 - 10.0	4.9	9	AOAB	
S095	10.5	210.0	5.0 - 10.0	4.9	9	AOAB	
S096	10.5	270.0	5.0 - 10.0	4.9	9	AOAB	
S097	10.5	330.0	5.0 - 10.0	4.9	9	AOAB	
S101	99999.9	0.0	8.0 - 9.0	4.9	9	AOAB	
S102	9.9	30.0	8.0 - 9.0	4.9	9	AOAB	
S103	9.9	90.0	8.0 - 9.0	4.9	9	AOAB	
S104	9.9	150.0	8.0 - 9.0	4.9	9	AOAB	
S105	9.9	210.0	8.0 - 9.0	4.9	9	AOAB	
S106	9.9	270.0	8.0 - 9.0	4.9	9	AOAB	
S107	9.9	330.0	8.0 - 9.0	4.9	9	AOAB	
S201	99999.9	0.0	1.0 - 3.0	4.0	6	AOB	
S207	99999 .9	0.0	8.0 - 9.0	4.5	6	AOB	
S254	9999 9.9	0.0	2.0 - 4.0	4.0	6	AOB	
S282	99999.9	0.0	4.0 - 8.0	4.0	6	AOB	
S310	99999.9	0.0	1.0 - 2.0	2.5	6	AOB	
S312	99999.9	0.0	2.0 - 4.0	2.4	6	AOB	
SI01	99999.9	0.0	0.5 - 1.5	3.2	6	AOB	
SI02	99999.9	0.0	1.0 - 2.0	3.2	6	AOB	
SI03	99999.9	0.0	1.5 - 2.5	3.0	6	AOB	
SI04	99999.9	0.0	2.0 - 4.0	2.6	6	AOB	
SI05	99999.9	0.0	3.5 - 5.5	2.6	6	AOB	
SI06	99999.9	0.0	5.0 - 10.0	2.8	6	AOB	

Table 7.5.1. (Page 2 of 2)



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Fig. 7.5.1. This map of the Svalbard archipelago with its main island Spitsbergen shows the location of the array site at Janssonhaugen (J), the location of the array controller at Norwegian Telecom's facility at Longyearbyen (L), and the location of the NOR-SAT B earth station at Isfjord Radio (IR).

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Fig. 7.5.5. Spitsbergen array noise suppression by beamforming for the initial geometry. The plot shows an average of 24 curves. To produce each of these curves, an infinitevelocity beam is formed and the spectrum for this beam is divided by the average of the single sensor spectra. The 24 curves result from one minute of data taken hourly between 00.00 and 23.00 hours GMT on day 345, 1992. The horizontal line at -8 dB represents \sqrt{N} suppression for 6 sensors.

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Fig. 7.5.6. Plot of individual Spitsbergen SPZ channels for the event discussed in the text. Note the very prominent amplitude variation across the array, as seen by the scale factors to the left of each trace.

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a) P phase





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b) S phase



Fig. 7.5.7. (Page 2 of 2)

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7.6 An evaluation of the performance of the Intelligent Monitoring System

Introduction

The Intelligent Monitoring System (IMS) is a computer hardware and software system which analyzes data from a network of seismic stations (arrays and three-component stations) to automatically detect and locate seismic events. The events automatically declared by IMS are reviewed by an analyst and modified as appropriate.

The version in current operation at NORSAR is described in Bache et al (1993). This version of IMS provides for joint processing of data from the NORESS and ARCESS arrays in Norway, the FINESA array in Finland, the GERESS array in Germany, and the new (fall of 1992) array new Apatity on the Kola peninsula of Russia (see Fig. 7.6.1). In addition, data from the array installed at Spitsbergen in November 1992 (Mykkeltveit et al, 1992) are made available to IMS for interactive e analysis; i.e., detections from this array are not yet automatically associated by IMS. Based on these data sources, IMS automatically produces a bulletin of local, regional and teleseismic events. The key module used by IMS in doing this is ESAL (Expert System for Association and Location). ESAL uses the detection information resulting from the signal processing (the SigPro module, containing automatic signal detection and characterization) performed separately for each array, and forms and locates events using artificial intelligence technologies, notably in the form of rule-based reasoning.

One of the basic objectives of IMS is to provide automatic event definitions of a quality that significantly reduces the burden on the seismic analyst. Operational experience with IMS is thus taken into account to produce enhancements to the system that result in performance improvements. NORSAR has tried to assist in this process by undertaking evaluations of IMS performance following the release of new versions by SAIC, the system developer. The results of such studies are discussed with SAIC, and taken into account in the next set of enhancements to the IMS.

The subject of this study is the outcome of another such evaluation of the IMS. The performance of the system was checked carefully during a recent one-week test period, and the focus was on the performance of ESAL. The performance statistics presented in the following thus basically reflect the quality of the automatic results produced by ESAL, and not the quality of the final event definitions after analyst intervention. In addition to this, we also point to some potential for improvements in SigPro that would enhance the overall performance of IMS.

Evlauation procedure

The evaluation was conducted by NORSAR analyst and seismologist staffs, who analyzed carefully and thoroughly complete IMS data for the 7-day period 26 April - 2 May 1993. The task was threefold, namely, (i) to characterize the automatic IMS event definitions, (ii) to modify these as deemed appropriate, and (iii) to check for missed events, i.e., events

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that were not defined by IMS, but where detections were available from SigPro that should allow ESAL to associate phases and form events. Task (ii) is important in providing information on the performance of various IMS algorithms. Examples here are statistics on retiming of phase onsets and renaming of regional phases. Task (iii) above was performed through various means: The bulletin of the GBF (Generalized Beamforming) processing, which uses the detection information produced by SigPro to locate regional events, was checked to see whether it contained events not declared by IMS. Likewise, the results from the single-array "ep"-processing performed at NORSAR based on the SigPro output was checked. In addition, completely independent bulletins like the NORSAR array bulletin and the PDE bulletin of the USGS were checked, basically to identify teleseismic events that IMS might have failed to catch, even if appropriate detections were available from SigPro.

IMS automatically produces "event plots" like the one shown in Fig. 7.6.2. These plots give a good starting point for the review process, as the inferences made by ESAL in forming and locating the events can be judged from these plots. The actual assessment of all events reviewed for this one-week evaluation period was made on the basis of using the interactive analysis tools offered by the Analyst Review Station.

Each event automatically declared by IMS was assigned to one of the following five categories:

Acceptable events: In addition to events accepted without any modifications, this category includes events where the anlayst made relatively minor changes. The character of these changes was such that the original IMS location was not strongly affected, i.e., the resulting change in location should be less than 50-100 km for regional events. For teleseismic events the requirements on "acceptable" events were more relaxed. The modifications in this category usually amounted to retiming of phases, association and disassociation of phases, as well as renaming of phases, especially for single-array events (e.g., from Sn to Lg or vice versa, from Pg to Px, from Rg to Lg, etc.).

Seriously mislocated events: This category includes events that are real events, but where the event location is too far away from the true location to qualify for the "Acceptable events" category. For these cases, either the phase assignments were wrong, or phases not belonging to the same event were associated.

False events: This category consists of those events declared by the IMS that the analyst rejected, believing they were not real events.

Inconclusive events: For this category, the analyst was not able to reach a definite conclusion whether the events were real or not.

Missed events: This category includes events that were *not* declared by IMS, but where detections were available from SigPro, that should permit ESAL to associate phases and form events. In addition to such events, the analyst occasionally produces events by manually adding signal arrivals that were not detected by SigPro. An example here would be a regional event for which SigPro has detected a P phase but where the S phase, though vis-

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ible to the analyst during interactive waveform analysis, has gone undetected. Such events are *not* counted as missed in the statistics presented in this report.

All events in these five categories are divided into regional and teleseismic events, based on their epicentral distance from the network:

Regional events: Events where the closest station is within 20 degrees of the epicenter.

Teleseismic events: Events where the distance to the closest detecting station exceeds 20 degrees.

Results and discussion

Table 7.6.1 summarizes the results of the characterization of the events automatically declared by IMS during the one-week test period, and in addition includes statistics on events missed by IMS. As seen from this table, nearly 80% of the events declared are considered to be acceptable, whereas the majority of the remaining ones are considered to be false.

All events in the category "acceptable" were considered by the analyst to be real seismic events with automatic location estimates that were either not changed during the subsequent interactive analysis, or that were slightly modified during this process through minor changes as described in the previous paragraph. Renaming of regional phases is one example of such modifications, and Table 7.6.2 offers statistics in this regard. Only phases for which the original arrival time was changed by less than 2.0 seconds are included in this table. This is done to exclude cases where a change in phase name was accompanied by a substantial change in the phase arrival time. It is noteworthy that only for one single case was a phase changed from P- to S-type by the analyst, and there were no cases where an Stype phase was changed to P. The extent of renaming can thus be characterized as being relatively modest, meaning that the automatic phase identification in ESAL now works very well.

As seen in Table 7.6.1, altogether 105 events automatically declared by IMS were rejected by the analyst as false, based on various kinds of evidence. We have taken a closer look at these events to see if there is parametric or other kind of information available that might permit ESAL to automatically reject these events. The following observations are made:

- For 42 (all regional) of these 105 events, there is, in the judgement of the analyst, parametric information available that might be used by ESAL to preclude the formation of an event. Examples of such parametric evidence are high or low Pn velocities, and frequency content of Pn and Lg phases well outside the expected range for these phases. The formation of these events might be precluded through the addition of new rules in ESAL in the form of consistency checks on the parameters pertaining to phases used in forming these events.
- 15 (also all regional) of the false events were rejected by the analyst because they were located close to an array in the network which showed no sign of any signal arrivals from this event. (It was checked that the closest array was operating nor-

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mally in these cases.) This is an expample where contextual information could be used to automatically reject an event hypothesis, and an appropriate consistency check in ESAL might rectify the situation. For several of the events in this category, also parametric information indicated that the events were false.

Altogether 18 events (both regional and teleseismic in this category) were judged by the analyst to be so-called "split" events, i.e., phase arrivals (often coda detections) belonging to a real event were used to define an additional, false event. Sometimes the arrival azimuth estimates for the phases used to define the split event deviated by 20° or so from those of the phases used to define the real event, and it would be difficult or maybe even inadvisable to preclude formation of the second event.

We find that 27 of the false events originate from detections resulting from bad data (spikes and gaps in the data). For 5 of the events, it is not possible for the analyst to see any signal at all (not even after beamforming) in the traces, and there may be some malfunction of the detector in these cases. Possible remedies for these 32 events would amount to changes in SigPro, but it should be noted that considerable work has already been invested in preventing SigPro from declaring detections when bad data are recorded.

There are a few false regional events at distances around 10^o that fall into none of the categories dealt with above. For these events, the only reason for rejecting them is the impression left after close inspection of the waveforms. The S wavetrain is impulsive and of very short duration, and thus does not match the expected shape (emergent onset and a coda of some length). It is of course very difficult to reject such events automatically, at least until the AI technology is able to match the trained eye of a skillful analyst.

As seen in Table 7.6.1, there are 15 regional events that were missed by IMS in the sense that IMS had available detection information that should allow ESAL to associate phases and form and locate events. Only one of these events had detections on more than one array, and this is a small (local magnitude 1.2) Khibiny Massif event recorded at ARCESS and the Apatity array. Three of the missed one-array events were associated with double events (two mine blasts at the same site, 10 seconds or so apart), where IMS only defined one event. For 4 of the remaining 11 missed one-array events, the arrival azimuth estimates for P and S phases differed by more than 20° (they were in the range 22.8-25.9°). This may be the reason why ESAL did not form these 4 events, but it appears that the remaining 7 events (with azimuth differences between P and S phases of less than 16.4°) should all have been formed by ESAL.

Table 7.6.1 shows that 14 teleseismic events were missed by IMS. These were events defined in the reference bulletins, and for which ESAL had detections available (from 2 or more arrays) that apparently should have been associated. It will be necessary to have a closer look at all of these to determine why they were not formed by ESAL.

Conclusions and recommendations

Our main impression after having carefully analyzed one week of data is that the overall performance of IMS is now very satisfactory. For example, the rules used for automatic phase identification in ESAL appear to work very well. It is also observed that problems
associated with earlier versions of IMS have now to a large extent been solved. There is, however, still room for some improvement, as discussed above and summarized in the following.

The number of false events may appear a bit high, and we have suggested some possible remedies that might help the IMS to automatically reject some of these. Such changes must, however, be tested very carefully to make sure they do not have unintended effects, like throwing out real events. In fact, in order to make sure that the system does not miss real events, we will just have to cope with a certain number of false events. This number can probably be reduced quite a bit from today's about 20% of the total number of events automatically declared by IMS, and it must remain a goal for IMS to minimize the burden on the analyst by defining as few false events as possible.

Events missed by IMS represent a more serious problem than that of the false events. Due to the large amounts of data processed by the IMS and the limited manpower resources available for interactive analysis of the automatic IMS results, it is unlikely that the analyst review process will pick up the events missed by ESAL. Keeping in mind the basic purpose of IMS, it is therefore of utmost importance that ESAL captures all real events, for which there is a solid basis for phase association and event formation. We have seen in the aforegoing that some of the missed regional events were associated with double mining blasts. Although this is a situation where a trained analyst could pick up the missed event when inspecting the waveform traces, it will certainly present a challenge to capture all such events missed were very small one-array events. Still, it is necessary to have a close look at ESAL to rectify this problem, as well as the problem of the missed teleseismic events.

As we have already seen, IMS performance could also be enhanced through certain modifications to SigPro. We will here touch upon another aspect where changes to SigPro might be beneficial: Figs. 7.6.3 and 7.6.4 provide statistics on retiming of the phases Pn and Pg, respectively, during the course of analyst review of the data for the one-week evaluation period. The figures show the differences between the arrival time as automatically determined by SigPro and the arrival time as determined by the analyst during the review process, plotted versus the SNR of the Pn or Pg phase, calculated from the detecting beam. Only phases that after analyst review retained their original, automatic phase assignment (Pn or Pg) and in addition correspond to real seismic events, are included in Figs. 7.6.3 and 7.6.4. The figures show standard deviations of the order of half a second for the differences in the arrival time estimates, and there are appreciable differences even for high SNR phases. This indicates that there should be quite some potential for improvement in the automatic estimation of phase arrival times. Section 7.2 of this Semiannual Technical Summary presents an approach that might be implemented in SigPro and that holds considerable promise to improve the automatic onset times.

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The current version of IMS makes use of some region-specific knowledge. Further enhancements to IMS performance are likely to be obtained through introduction of additional such knowledge, and section 7.3 of this Semiannual Technical Summary demonstrates how event locations may be significantly improved using region-specific knowledge.

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References

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	Regional	Tele- seismic	Total
Acceptable events	384	67	451
Seriously mislocated events	5	4	9
False events	97	8	105
Inconclusive	0	4	4
Total number of events declared			569
Missed events	15	14	29

Table 7.6.1 Characterization of all events automatically declared by IMS during the seven-day period 26 April - 2 May 1993 (see text for explanation of the various categories). The table also includes statistics on events missed by the IMS.

		Pn	Pg	Рх	Sn	Lg	Rg	Sx
ESAL	Pn	213	19	2	1			
	Pg	11	203	24	1			
	Px	2	2	60		1		
	Sn	1			34	5		9
	Lg			1	7	325	7	24
	Rg					6	30	7
	Sx				1	20	1	75

Analyst

Table 7.6.2 The table shows the ESAL automatic phase assignments versus analyst assignments made during interactive analysis, for all regional phases that were used by IMS to define regional events during the time period 26 April - 2 May 1993.





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Fig. 7.6.1 The network of six regional arrays in northern Europe.



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Fig. 7.6.2. A typical example of a plot automatically produced by IMS, basically showing the judgements and inferences made by ESAL in associating phases at the various arrays and forming this event in Estonia.



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Fig. 7.6.3. The figure shows differences between arrival times as determined by SigPro and arrival times determined by the analyst during review, plotted versus SNR, for Pn phases used by IMS to form events during the evaluation period 26 April - 2 May 1993. The mean value of the arrival time differences is marked by a solid vertical line, whereas the two strippled lines denote ± 2 standard deviations.



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Fig. 7.6.4. Same as Fig. 7.6.3, but for the Pg phase.