

PL-TR-92-2006

CLUSTER ANALYSIS OF CLOSELY SPACED MINING BLASTS AS A METHOD OF EVENT LOCATION

Florence Riviere-Barbier Lori Teresa Grant

Science Applications International Corporation Center for Seismic Studies 1300 N. 17th Street, Suite 1450 Arlington, VA 22209

30 January 1992

Final Technical Report 1 October 1990 - 1 October 1991

Approved for public release; distribution unlimited

4 08 127





99

PHILLIPS LABORATORY AIR FORCE SYSTEMS COMMAND HANSCOM AIR FORCE BASE, MASSACHUSETTS 01731-5000

SPONSORED BY Defense Advanced Research Projects Agency Nuclear Monitoring Research Office ARPA ORDER NO. 5307

MONITORED BY Phillips Laboratory Contract F19628-88-C-0159

The views and conclusions contained in this document are those of the authors and should not be interpreted as representing the official policies, either expressed or implied, of the Defense Advanced Research Projects Agency or the U.S. Government.

This technical report has been reviewed and is approved for publication.

JAMES F. LEWKOWICZ Contract Manager Solid Earth Geophysics Branch Earth Sciences Division

JAMES F. LEWROWICZ

Branch Chief Solid Earth Geophysics Branch Earth Sciences Division

malel teller. DONALD H. ECKHARDT, Director

Earth Sciences Division

This report has been reviewed by the ESD Public Affairs Office (PA) and is releasable to the National Technical Information Service (NTIS).

Qualified requestors may obtain additional copies from the Defense Technical Information Center. All others should apply to the National Technical Information Service.

If your address has changed, or if you wish to be removed from the mailing list, or if the addressee is no longer employed by your organization, please notify PL/IMA, Hanscom AFB, MA 01731-5000. This will assist us in maintaining a current mailing list.

Do not return copies of this report unless contractual obligations or notices on a specific document requires that it be returned.

REPORT D	OCUMENTATION P	AGE	Form Approved OMB No. 0704-01	
Public reporting burden for this collection of in gathering and maintaining the data needed, an collection of information, including suggestion Davis Highway, Suite 1204, Arlington, VA 2220	formation is estimated to average 1 hour pe id completing and reviewing the collection of s for reducing this burden, to Washington He 2-302, and to the Office of Management an	r response, including the time for r information. Send comments rega adquarters Services, Directorate fo Budget, Paperwork Reduction Pro	eviewing instructions, searching existin rding this burden estimate or any other r information Operations and Reports, ject (0704-0188), Washington, DC 2050	g data sources. If aspect of this 1215 Jefferson 3.
1. AGENCY USE ONLY (Leave black	nk) 2. REPORT DATE	3. REPORT TYPE AN Final (1 Oct	D DATES COVERED	
4. TITLE AND SUBTITLE	Jo Sandary 1992		5. FUNDING NUMBERS	
Cluster Analysis of (Method of Event Locat	Closely Spaced Mining	Blasts as a	PE 62714E PR 8A10 TA DA W	U AT
6. AUTHOR(S)	··· <u>·</u> ································		Contract F19628-8	8C-0159
Florence Riviere-Barb Lori Teresa Grant	Dier			
7. PERFORMING ORGANIZATION N	AME(S) AND ADDRESS(ES)		8. PERFORMING ORGANIZA	TION
Science Applications Center for Seismic St 1300 N. 17th Street, Arlington, VA 22209	International Corpora udies Suite 1450	ation	REPORT NUMBER	
9. SPONSORING/MONITORING AG	ENCY NAME(S) AND ADDRESS(E	5)	10. SPONSORING / MONITO AGENCY REPORT NUME	RING BER
Phillips Laboratory Hanscom AFB, MA 01731	-5000		PL-TR-92-2006	
Contract Manager: Ja	mes Lewkowicz/GPEH			
12a. DISTRIBUTION/AVAILABILITY Approved for public r	STATEMENT release;		125. DISTRIBUTION CODE	
13. ABSTRACT (Maximum 200 word	d ds)	. <u></u>		
This is the Final Re Seismic Data. The a characterize the hig of discriminating be frequency seismic da Israelsson and Carte contract and covered teleseismic waveform NRDC high frequency	port of contract F196 im of the research co h frequency content of tween mine blasts and ta. Two previous rep r, 1991) describe muc the high frequency co as as recorded at the stations in the Sovie	28-88-C-0159, An inducted under th if noise and sign other seismic s orts (Israelsson in of the researc haracteristics o short period Sca t Union.	alysis of High Fred is contract was to als, and to develop ources using high et al., 1990; and h conducted under to f local, regional, ndinavian arrays ar	uency methods this and and the
The final six months characterization for used to group events Karelian, northwest visual classificatio (continued on revers	of the contract were location and discrim with similar charact of St. Petersburg. T n of the same data an e)	devoted to a sta ination purposes eristics from the he groupings com d were hypothesis	udy of mining event . Cluster analysis e mining district c pared well with a c zed to be associate	s was of areful ed with
14. SUBJECT TERMS			15. NUMBER OF P	AGES
FINESA Cluster ARCESS f-k ana	analysis IM lysis	S	54 16. PRICE CODE	
NORESS Cross-c 17. SECURITY CLASSIFICATION T	Orrelation 18. SECURITY CLASSIFICATION	19. SECURITY CLASSIFIC	CATION 20. LIMITATION O	FARSTRACT
OF REPORT Unclassified	OF THIS PAGE Unclassified	OF ABSTRACT Unclassified	SAR	

NSN 7540-01-280-5500

¥

Standard Form 298 (Rev. 2-89)

13. Abstract (continued)

specific mines. The results of this study are very promising and warrant expanding to other mining districts. Future studies based on this work, however, should verify the event sources using non-seismic data before an absolute event identification is assigned.

TABLE OF CONTENTS

LIST OF FIGURES iv	
LIST OF TABLES vii	
FOREWORD ix	
INTRODUCTION 1	
DATA	
MINE LOCATIONS	
ANALYSIS METHODS	
FINESA RESULTS	
RESULTS OF THE CLUSTER ANALYSIS USING ARCESS AND NOR- ESS DATA	
ASSOCIATION OF EVENTS TO MINE LOCATION	
SUMMARY AND CONCLUSIONS	
ACKNOWLEDGEMENTS	
REFERENCES	

	AC	uion Fer
	NTT3 D740 Uusiin Just1	GRAAI 2 TAB
\bigcirc	By Distr	toutiou/
(HEED)	Diat A-1	Aveil and/or Special

LIST OF FIGURES

Figure 1:	Location of ARCESS, FINESA and NORESS arrays collecting the data processed by the Intelligent Monitoring System. ARCESS and NORESS are located in Norway while FINESA is located in Fin- land. Three important mining districts have been identified: in Esto- nia, in Karelian, and on the Kola Peninsula.	2
Figure 2:	Magnitudes reported in the Helsinki bulletin for 89 of the studied events. Many of the events for which the IMS magnitude was above 1.8 were located using an automatic process. Only a few events with an IMS magnitude below 1.2 were reported in the Helsinki bulletin. For the remaining events, a "manual location" was reported that cor- responded to a mine location. The Finnish analysts routinely recog- nize and "manually locate" recurring mining blasts close to their stations.	5
Figure 3:	Comparison between mine locations determined using Helsinki seismic data (HC1 to HC14) and the locations determined on SPOT photos by Fox (1991). Helsinki locations have been determined seis- mically and an attempt has been made to provide an equivalent SPOT location for each of them. A SPOT location can include sev- eral small mines within a few kilometers of each other.	7
Figure 4:	One hundred-forty-four events recorded at FINESA were studied and classified. Their IMS locations are shown on this map using la- bels corresponding to each group. Their IMS local magnitudes range from 0.22 to 2.6. Even though IMS locations gave a good idea of the event locations, they were not accurate enough to distinguish be- tween events from two mines 5 km apart.	10
Figure 5:	Eighty-nine of the events studied were reported in the Helsinki bul- letin; 31 events were located automatically and 55 events had "man- ual locations". On this plot, IMS locations are compared with Helsinki locations, both manual and automatic. The largest discrep- ancies in location occur for events with a low signal-to-noise ratio. Two of them were clearly mislocated by the IMS. An interactive f - k analysis further confirmed that the azimuths determined by the IMS were erroneous.	15
Figure 6:	Reference events for 18 groups determined by visual analysis of the data. Both unfiltered and filtered data were used. Groups are labeled from A through W.	16

ж.

Figure 7:	Events from group K and S are plotted to show the difference be- tween two groups as well as the repeatability of the signal in each group. A difference is clearly seen in the Lg -P time and in the pres- ence of an Rg phase. For other groups, a difference is evident only in the shape of the first arrival, the Lg -P time being the same
Figure 8:	This figure shows the first arrival of reference events for groups C through M. Their Lg -P times and Rg -P times vary by less than one second. Differences between events can best be seen using unfiltered data to look at the shape of the first arrival. Although only four mines have been reported in this area in the Helsinki Bulletin, visual classification and cluster analysis determined 8 groups with more than one event and two with only one event. A careful SPOT photo analysis should be performed to determine whether or not this sub-division corresponds to a real distribution of the mines
Figure 9:	Two events from group R and group S have been plotted. They have a magnitude of 2.05 and 2.12, respectively. Groups R and S are lo- cated at the same mine according to the Helsinki bulletin (HC13). In the absence of the SPOT photo location for this mine, two assump- tions can be made: either these events are from two different parts of the same large mine or they come from the same mine and have dif- ferent source parameters. 20
Figure 10:	Apparent velocities of the first arrival have been computed for each event using an <i>f-k</i> analysis method. This map shows the spatial dis- tribution of the average apparent velocity for each mine group. North of Lake Ladoga, the velocity is around 6.45 km/s while South of the lake, the velocity is around 7.35 km/s and keeps increasing up to 8.77 km/s for the events located on an island in the Gulf of Fin- land. Strong variations in the thickness of the crust as well as sharp lateral boundaries in the crust can explain these changes in the ap- parent velocity
Figure 11:	Contour map of crustal thickness (km) and schematic map of Pn ve- locity for the Baltic Shield and adjacent areas based on DSS data. Lines of equal Moho depth are represented by thick solid lines for reliable data and dashed lines for unreliable data. Values of Pn ve- locity are: 1, 7.8 to 8.0; 2, 8.1 to 8.3; 3, 8.3 to 8.5. Thin solid lines denote DSS profiles. The crustal thickness varies from as little as 30 to 35 km near the coast to 50 to 55 km within the interior areas. Pn wave velocity varies from 7.8 to 8.0 km/s up to 8.3 to 8.5 km/s (the most frequently observed values are 8.1 to 8.2 km/s). There is no di- rect relationship between variations of the crustal thickness and Pn

	velocity (Ryaboy, 1990).	24
Figure 12:	The tree resulting from cluster analysis using a "complete linkage method". Envelopes of filtered data (1-15 Hz) recorded on the ver- tical channel of the FIA1 sensor were computed. A cross-correlation value was calculated for each pair of events. These cross-correlation values were used as similarity measurements in the cluster analy- sis.	26
Figure 13:	Fifty-three events belonging to groups C through O have been reprocessed using the frequency band 1-15 Hz. The resulting clustering shows a good agreement with the visual classification. This set of data is especially difficult and in addition to the labels C O, others labels were used. Lower case letters were used to label events that were similar but not identical to the group labeled with the upper case letter, the difference residing mostly in the Lg -P time. The "a" label was used for an event that was visually not close to any other group.	27
Figure 14:	Signal-to-noise ratios are compared for 39 events recorded at the three arrays: ARCESS, NORESS and FINESA. ARCESS and NOR-ESS have very similar values. These two arrays are located at about the same distance from the mining district. The signals recorded at FINESA have a signal-to-noise ratio roughly 9 times higher than the signals recorded at NORESS and ARCESS.	29
Figure 15:	Result of the cluster analysis performed on 31 events recorded at ARCESS. Data were filtered between 2 to 5 Hz. Events with the highest signal-to-noise ratio give reasonable results.	31
Figure 16:	Result of the cluster analysis performed on 31 events recorded at NORESS. The same frequency band applied to the ARCESS data was applied to these signals.	31

LIST OF TABLES

Table 1:	Mine Locations from the Helsinki bulletin (Hels.) and SPOT photos (Fox, 1990)	. 4
Table 2:	One-hundred-forty-four events recorded at FINESA	11
Table 3:	Azimuth and velocity values computed automatically by the IMS (Azim2, Ap.vel2) and the same values computed using f-k analysis (Azim1, Ap.vel1).	23

FOREWORD

This is the Final Report of contract F19628-88-C-0159, Analysis of High Frequency Seismic Data. The aim of the research conducted under this contract was to characterize the high frequency content of noise and signals, and to develop methods of discriminating between mine blasts and other seismic sources using high frequency seismic data. Two previous reports (Israelsson *et al.*, 1990; and Israelsson and Carter, 1991) describe much of the research conducted under this contract and covered the high frequency characteristics of local, regional, and teleseismic waveforms as recorded at the short period Scandinavian arrays and the NRDC high frequency stations in the Soviet Union.

Much of the work presented in the previous reports contributed to the mine characterization study described herein. The material covered in Israelsson *et al.* (1990) showed evidence that shooting practices at the Kola Peninsula mines is different than at the Scandinavian mines. Studies were also made of event characterization using spectrograms (in an attempt to characterize events with spectral scalloping) and waveform correlation of closely spaced events. The spectral scalloping study showed that while spectral scalloping was evident for some mine blasts, it was not a consistent feature. Waveform correlation, on the other hand, showed promise as a method for grouping events from a specific mine and was adopted in a modified form for this final study. Israelsson and Carter, (1991) contains studies on the high frequency content of teleseismic P-waves; estimating the characteristics of ripple-fired explosions; and slowness estimation with interpolated NORESS data. The slowness estimation study was able to discern differences in the wave-fronts of the first arrivals from the northeast and southwest sections of the Balapan test site using interpolated NORESS data.

The final six months of the contract were devoted to a study of mining event characterization for location and discrimination purposes. Cluster analysis was used to group events with similar characteristics from the mining district of Karelian, northwest of St. Petersburg. The groupings compared well with a careful visual classification of the same data and were hypothesized to be associated with specific mines. The results of this study are very promising and warrant expanding to other mining districts. Future studies based on this work, however, should verify the event sources using non-seismic data before an absolute event identification is assigned.

INTRODUCTION

Numerous discrimination studies have been performed to distinguish between events of different source type (Blandford, 1982; Pomeroy *et al.*, 1982). The initial studies used modeling in an attempt to reproduce the effects of the source on the signal or spectrum. More recent studies used a case-based approach and tried to parameterize the signal and spectrum. Studies of frequency content in order to detect "ripple firing" (Baumgardt and Ziegler, 1988; Smith, 1989) showed that spectral modulations could be used to distinguish mine blasts from earthquakes and nuclear explosions. Dysart and Pulli (1989) characterized chemical explosions and earthquakes recorded at the NORESS array using amplitude ratios and spectral complexity. Israelsson (1990), in an application of cluster analysis to the discrimination problem, was able to distinguish closely spaced events located within an area of 20 km by 75 km by taking advantage of the repeatability of signals from c⁺ nilar sources. In general, these studies showed that methods used to discriminate events should be applied at regional distances so that the tectonic environment is integrated (Bennett *et al.*, 1989).

The purpose of this study is to characterize and distinguish between mines located within a few kilometers of each other. This characterization may then be implemented in an automatic detection, location, and discrimination system such as the Intelligent Monitoring System (IMS) (Bache *et al.*, 1990). A discrimination method that can associate events with a particular mine will reduce the number of events that need to be investigated in more detail.

The IMS processes data from four arrays (ARCESS, NORESS, FINESA and GERESS) providing a large data set of local and regional events from the baltic shield and the western part of the east-european platform (*Figure 1*). An automatic location is computed for each event which is later reviewed by an analyst. Even when an event location is constrained by data from three different arrays, the error ellipse can be large enough to include several mines. The average estimated error of the IMS locations is 20 km (Bratt *et* al., 1990).

Three different areas with a high concentration of mines in the local to regional distance range from the IMS arrays were identified in the Russian territories closest to Finland. Two areas are located near St. Petersburg and the third one includes mines on the Kola Peninsula. This paper is devoted to the results obtained for the mining district located north of St. Petersburg, in Karelian.

The discrimination technique used for this study is a cluster analysis method (Everitt, 1986) based on waveform similarity. It does not require any pre-classification of events into groups.



Figure 1: Location of ARCESS, FINESA and NORESS arrays collecting the data processed by the Intelligent Monitoring System. ARCESS and NORESS are located in Norway while FINESA is located in Finland. Three important mining districts have been identified: in Estonia, in Karelian, and on the Kola Peninsula.

DATA

The data set covers the period of time from November 4, 1990 to June 28, 1991. Waveforms recorded at FINESA, ARCESS and NORESS as well as phase parameters computed and saved during the automatic detection and location of events by the IMS were utilized in this study. The set of phase parameters extracted from the IMS database was similar to the set used by Baumgardt (1987).

Waveforms

Waveforms used in this study were automatically saved on optical disk by the IMS. Since November 4, 1990, all waveforms with at least one associated phase after analyst review have been archived, unless a software failure caused a loss of data.

Database Parameters

Only events with at least one Pn or Pg and one Lg phase were considered. For each phase, the following parameters were extracted from the IMS parameter database:

- detection time
- azimuth computed from *f*-*k* analysis
- velocity computed from *f*-*k* analysis
- center period of the phase
- short term average measured on the incoherent vertical beam filtered between 2 and 4 Hz.

The detection time was used to compute a "relative time". This relative time was defined as the time difference between the earliest phase detected at any of the stations and the detection time of the other phases (Bache *et al*,1990). The short term average (sta) values were used in a ratio: sta(Pn)/sta(Lg) computed in decibels. Additional polarization parameters such as azimuth and rectilinearity for the *P*-type phase, horizontal-to-vertical ratio for *P*-type and *S*-type phases, and planarity for the *S*-type phases were added to the set of parameters described above.

Event locations

In addition to the IMS bulletin, a source of accurate locations for small events in this area is the bulletin published by the University of Helsinki (Uski *et al.*, 1990). Monthly bulletins are released several months after the events occur. A weekly bulletin is released earlier.

Fewer events were reported in the Helsinki bulletin than in the IMS bulletin as the Hels-

inki bulletin did not report most of the events with an IMS magnitude below 1.2 (Figure 2). Of the 55 events not reported in the Helsinki bulletin, only seven had an IMS magnitude greater than 1.2. Most of the events with a magnitude between 1.2 and 1.8 in the Helsinki bulletin had a "manual location" which means that the Finnish analyst assigned them to a particular mine by visual inspection. An automatic location was reported for the events with an IMS magnitude above 1.8 and included information from other stations in the Nordic countries. The IMS location included information from only the FINESA array for small events or from all three arrays (ARCESS, NORESS, FINESA) for larger events (usually with a local magnitude above 1.5).

MINE LOCATIONS

Hels.	lat.	lon.	azim.	dist.	SPOT	lat.	lon.	azim.	dist.
HC1	60.7°	28.7°	119.2°	1.48°	SC1	60.749°	28.836°	116.3°	1.51°
HC2	60.7	29.0	116.47	1.60	SC2	60.700	29.181	114.99	1.68
HC3	60.6	29.2	117.78	1.74	SC3	60.581	29.065	119.48	1.69
HC4	60.8	29.3	111.06	1.69	SC4	60.846	28.99	111.76	1.53
HC5	60.9	29.3	107.88	1.65	SC5	61.008	29.038	105.71	1.50
HC6	60.9	29.4	107.29	1.70	SC6	60.953	29.176	106.84	1.58
HC7	60.8	29.5	109.78	1.78	SC7	60.902	29.348	107.53	1.67
HC10	61.1	29.9	98.93	1.88					
HC11	61.1	30.2	98.04	2.00	SC11	61.142	29.870	97.75	1.85
HC12	61.5	30.4	86.56	2.07					
HC13	61.9	30.6	76.05	2.20					
HC14	61.4	31.6	88.54	2.65	SC14	61.605	31.424	84.06	2.57
HB15	60.0	29.9	125.99	2.37	SB15	60.019	29.742	126.87	2.30
N114	61.03	28.18	111.34	1.10					
N117	61.9	29.0	70.55	1.47					

Table 1: Mine Locations from the Helsinki bulletin (Hels.) and SPOT photos (Fox,1990).



Figure 2: Magnitudes reported in the Helsinki bulletin for 89 of the studied events. Many of the events for which the IMS magnitude was above 1.8 were located using an automatic process. Only a few events with an IMS magnitude below 1.2 were reported in the Helsinki bulletin. For the remaining events, a "manual location" was reported that corresponded to a mine location. The Finnish analysts routinely recognize and "manually locate" recurring mining blasts close to their stations.

Table 1 is a list of the mine locations available for the area. The Helsinki bulletin made reference to 15 mines in this area during this time period. The left half of the table corresponds to the locations found in the Helsinki bulletin. The right half of the table gives information compiled by Fox (1990) using satellite photos. Helsinki mine locations have been determined seismically by averaging repeated events. The accuracy of these locations is only given to one-tenth of a degree. Some locations clearly include several small mines about 2.5 km apart. According to the scientists at the University of Helsinki, locations of events outside of the Finnish network are not considered accurate to better than 5 km. An attempt was made to provide an equivalent SPOT location for each Helsinki mine. SPOT locations can be an average location that includes several small mines. Most SPOT locations have a corresponding Helsinki location; however, a few Helsinki mine locations do not have a corresponding SPOT location (*Figure 3*)

ANALYSIS METHODS

Cluster analysis

In order to characterize and classify the mining events from a small area, a method was required that could distinguish the subtle differences between closely spaced events. The ability of the method to work in an automated system was also considered. Cluster analysis was chosen for the task as it provided a method of grouping events based on an comparison of event pairs. The result of each comparison was reduced to a single number, the collection of which constituted the elements of a matrix used by the cluster analysis.

Cluster analysis can be based on the comparison of either "distance" or "similarity" between data (Everitt, 1986). Distance values are greater than or equal to zero. Similarity values range from zero to one. In a distance scheme, data that are similar are assigned a small distance value and data that are dissimilar are assigned large distance values. The opposite is true for similarity measurements. A similarity measurement was applied to the waveforms while a distance measurement was more appropriate for parameter data. The results of either measurement are represented by a cluster tree or dendogram.

Among the several methods used to build cluster trees, the most popular are the *complete linkage* method and the *single linkage* method. The results of these clusters can be quite different. In the *single linkage* method, distance between groups is defined as the distance between the closest members while in the *complete linkage* method, this distance is defined as the distance between the furthest members. The results presented in this paper were better represented by the *complete linkage* method because it handles problematic data better (i.e. multiple events). Several similarity measurements were tried and are



Figure 3: Comparison between mine locations determined using Helsinki seismic data (HC1 to HC14) and the locations determined on SPOT photos by Fox (1991). Helsinki locations have been determined seismically and an attempt has been made to provide an equivalent SPOT location for each of them. A SPOT location can include several small mines within a few kilometers of each other.

7

described below.

Cross-correlation

By assuming that events from the same mine should look very similar, the computation of cross-correlation values provides an easy way to build a numerical link between similar events. A high cross-correlation value between two events would imply a high probability that the events were from the same mine. A cross-correlation function w_i s computed between each pair of events using signals recorded on the sz channel (FIA1 sensor) at FINESA. The input to the cluster analysis was a symmetric $n \times n$ matrix where n was the number of events. The elements of the matrix were the maximum values of the cross-correlation function.

Waveforms recorded at FINESA from the Karelian mines were studied because the signalto-noise ratio was higher than for the data recorded at ARCESS or NORESS. Different data processing techniques were tested before computing the cross-correlation function. Raw data, filtered data, and the envelope of either unfiltered or filtered data were tried using sz, sn and se components either separately or in combination. Different signal lengths were tested including the entire signal, only P, and only Lg. The best results were obtained for the cross-correlation between the envelope (from the Hilbert transform) of the entire signal on the sz component. When the signal-to-noise ratio was low, best results were obtained with signals filtered between 1 and 15 Hz.

Other similarity measurements

The method described by Israelsson (1990) was based on a covariance matrix which was built for each component pair and each event. The six covariance matrices were stacked for each event and a cross-correlation was computed for each pair of stacked traces. This method was tested but did not provide more accurate results than the cross-correlation between envelopes.

Another similarity measurement was made using the maximum value of the coherence between pairs of events. Two different methods were used to pick this maximum value. In one method, the maximum of the entire coherence function was used; and in the other method, the maximum was picked in the frequency band that showed the best signal-tonoise ratio. Neither method was successful because there was little coherent energy above the noise between event pairs and it was difficult to find a consistent frequency band from event to event that provided useful results.

Ł

Distance measurements

In an attempt to use data parameters, relations between events were better represented by a

distance measurement. The set of phase parameters characterizing each event was used to compute a "distance" value between each pair of events. The data were first normalized as the parameters did not have the same units. Then, a principal component analysis was performed on the data to eliminate any redundant information. Finally, an euclidean distance (Everitt, 1986) was computed for each pair of events that were used in the cluster analysis.

Visual classification

In order to better interpret and verify the results of the cluster analysis, a visual classification was performed using recordings of the vertical channel. The visual classification was based on the following characteristics listed in order of importance:

- *Lg P* time,
- Rg P time,
- similarity in the shapes of the P arrivals for the first three seconds,
- similarity in the shapes of the Lg and/or the Rg phases,
- superposition of the waveforms for each phase,
- frequency content.

The visual classification using filtered and unfiltered data sometimes resulted in different groupings for events from closely spaced mines. The classification based on unfiltered data (when the signal-to-noise ratio was high enough) was preferably used.

Interactive f-k analysis

An f-k analysis was performed interactively on most of the events using all eighteen sz channels of the FINESA array. The aperture of the array is 2 km. An apparent velocity (from a slowness) and an azimuth were obtained for each of the studied events. In this way, a coherent beam was calculated for each event. The purpose of this analysis was to see if these f-k results could provide event locations that were accurate enough to associate events to specific mines. Azimuth and slowness values obtained by the IMS using automatically picked phases did not provide sufficient accuracy.

FINESA RESULTS

Seven months of FINESSA data were searched for events that occurred within the area between 60 and 62° latitude, and 28 and 32° longitude. One hundred and forty-four events met these criteria and were used in the cross-correlation and cluster analysis computations. The IMS locations of the events are plotted in *Figure 4* and listed in Table 2.. The IMS local magnitudes ranged between 0.22 and 2.66. The distance from FINESA was between



Figure 4: One hundred-forty-four events recorded at FINESA were studied and classified. Their IMS locations are shown on this map using labels corresponding to each group. Their IMS local magnitudes range from 0.22 to 2.6. Even though IMS locations gave a good idea of the event locations, they were not accurate enough to distinguish between events from two mines 5 km apart.

Ref.	V .	C.	H.C.	Date	Time	ml	Latl	Lon1	Lat2	Lon2
c93	#	Z2	AU	04/24/91	10:00:33	1.49	61.81	30.78	59.27	27.69
c7	#	Z1	HB9	11/21/90	12:17:07	1.26	60,71	28.50	59.2	27.6
c79	#	W	NR	04/09/91	12:57:51	1.08	60.10	28.98		
c104	#	Z2	NR	04/29/91	10:59:55	1.49	60.07	30.80		
c17	?	B1	NR	12/16/90	02:40:17	1.06	61.92	29.55		
c37	?	Z2	NR	01/08/91	19:40:37	0.47	61.87	29.79		
c40	?	Z2	NR	01/15/91	10:03:54	0.38	60.89	29.96		
c41	?	Z2	NR	01/16/91	00:07:25	0.59	61.87	30.36		
c101	?	Z1	NR	04/26/91	11:33:56	0.61	61.34	28.70		
c112	?	Z2	NR	05/12/91	12:45:19	0.37	60.11	28.97		
c113	?	Z2	NR	05/12/91	13:06:58	0.20	60.07	28.90		
c136	?	B1	NR	06/06/91	20:41:54	0.74	61.81	29.63		
c144	?	B1	NR	06/28/91	07:31:24	0.59	61.02	29.18		
c10	A	A	AU	11/26/90	12:01:35	1.27	60.76	28.76	61.04	28.31
c3	A	A	NR	11/15/90	08:02:47	0.83	60.59	28.83		
c30	A	A	NR	01/03/91	12:50:42	0.96	60.68	28.72		
c33	A	A	NR	01/04/91	12:37:10	0.94	60.77	28.41		
c36	A	A	NR	01/07/91	10:31:23	0.91	60.68	28.80		
c45	A	A	NR	01/30/91	12:37:27	0.90	60.57	28.81		
c47	A	Z1	NR	02/04/91	12:10:58	0.76	60.77	28.77		
c48	A	Z2	NR	02/06/91	12:06:49	0.40	60.79	28.69		
c106	A	A	NR	04/30/91	11:03:02	1.00	60.74	28.61		
c111	A	A	NR	05/10/91	11:03:20	1.07	60.98	28.31		
c34	B	B2	N117	01/04/91	14:17:51	1.37	61.80	29.63	61.9	29.0
c29	B	BI	NR	01/02/91	11:52:31	0.50	61.76	29.81		
c31	B	B1	NR	01/03/91	19:37:23	0.66	61.92	29.43		
c32	В	B1	NR	01/04/91	11:39:42	0.88	61.74	29.63		
c38	B	B1	NR	01/14/91	11:42:44	0.70	61.92	29.80		
c39	В	B1	NR	01/14/91	19:50:54	0.71	61.63	29.81		
c44	B	B2	NR	01/28/91	14:23:25	1.28	61.68	28.55		
c52	В	B1	NR	03/11/91	11:46:43	0.54	61.72	29.52		
c54	B	B1	NR	03/11/91	19:40:21	0.56	61.94	29.43		
c56	В	B1	NR	03/14/91	11:47:30	1.04	61.73	29.50		
c62	В	B1	NR	03/21/91	19:38:48	0.70	61.75	29.68		
c67	В	B1	NR	03/26/91	19:39:11	0.68	61.73	29.47		
c78	В	B1	NR	04/08/91	21:56:35	0.60	61.76	28.82		
c86	В	Z2	NR	04/17/91	18:43:58	0.92	61.86	28.88		
c88	В	Z2	NR	04/18/91	18:50:56	0.37	61.72	29.66		
c92	В	B1	NR	04/22/91	18:50:23	0.54	61.57	29.67		
c109	В	B1	NR	05/08/91	18:37:51	0.25	61.76	29.26		
c116	В	B1	NR	05/14/91	18:40:40	0.37	61.78	29.51		
c122	В	B1	NR	05/23/91	19:33:23	0.47	61.44	29.61	*	
c126	В	Z2	NR	05/29/91	20:51:57	0.38	61.75	29.69		
c133	B	BI	NR	06/04/91	18:55:28	0.22	61.76	29.48		

Table 2: One-hundred-forty-four events recorded at FINESA.

Ref.	V .	C.	H.C.	Date	Time	ml	Lat1	Lon1	Lat2	Lon2
c18	С	С	HC4	12/17/90	12:51:58	1.96	60.76	29.09	60.8	29.3
c81	C	С	HC5	04/11/91	13:21:38	2.14	60.93	28.94	60.9	29.3
c140	C	C	HC5	06/26/91	12:31:59	2.09	60.93	29.01	60.9	29.3
c12	С	C	HC6	11/29/90	12:29:11	2.66	60.91	29.13	60.9	29.4
c69	C	С	HC6	03/28/91	14:49:15	2.32	60.90	28.96	60.9	29.4
c123	D	D	AU	05/24/91	12:02:00	1.99	60.80	28.96	60.92	29.03
c49	D	D	HC6	12/14/90	11:52:05	2.22	61.11	29.15	60.9	29.4
c77	D	D	HC6	04/08/91	11:41:08	1.76	60.98	28.97	60.9	29.4
c53	E	E	AU	03/11/91	12:29:45	2.12	60.91	29.02	60.95	29.03
c94	E	E	AU	04/24/91	12:07:05	2.25	60.98	29.08	60.96	29.30
c98	E	E	AU	04/25/91	12:10:27	2.18	60.89	29.13	61.00	29.17
c129	E	E	AU	05/31/91	13:55:23	2.18	61.03	29.11	60.89	29.12
c22	E	E	HC5	12/25/90	12:28:48	2.16	60.78	29.08	60.9	29.3
c28	Ē	E	HC5	12/28/90	12:27:34	2.08	60.95	29.10	60.9	29.3
c89	F	F	HC4	04/19/91	12:43:01	1.76	60.93	29.12	60.8	29.3
c21	F	F	HC5	12/22/90	16:48:34	1.99	60.75	29.08	60.9	29.3
c134	F	F	HC5	06/05/91	11:43:12	1.79	60.77	29.01	60.9	29.3
c2	G	G	HC2	11/05/90	12:00:21	2.13	60.85	29.05	60.7	29.0
c50	G	G	HC2	03/07/91	11:11:24	2.29	60.64	28.72	60.7	29.0
c97	Н	•	AU	04/25/91	10:43:00	2.05	60.75	29.03	60.67	29.07
c117	H	Н	AU	05/21/91	12:33:30	1.86	60.78	29.05	60.74	28.99
c125	H	Н	AU	05/29/91	14:06:55	1.85	60.74	28.73	60.64	28.96
c57	Н	Н	HC1	03/14/91	11:08:08	1.73	60.72	28.91	60.7	28.7
c11	Н	Н	HC2	11/27/90	12:17:01	1.42	60.67	28.55	60.7	29.0
c75	Н	Н	HC2	04/04/91	12:53:16	1.50	60.75	28.82	60.7	29.0
c16	K	K	AU	12/06/90	14:22:38	2.22	60.85	29.19	60.89	29.09
c46	K	K	AU	01/30/91	12:59:33	2.24	60.85	29.15	60.89	29.11
c74	K	K	AU	04/03/91	12:25:23	1.99	60.94	29.11	60.91	29.16
c76	K	K	AU	04/05/91	12:54:23	2.14	60.91	29.13	60.92	29.13
c95	K	K	AU	04/24/91	13:10:52	2.32	60.84	29.04	60.90	29.21
c27	K	K	HC4	12/27/90	12:30:36	2.23	60.83	29.23	60.8	29.3
c51	K	К	HC6	03/07/91	12:37:41	2.17	60.78	29.24	60.9	29.4
c142	K	K	HC6	06/27/91	13:08:04	2.40	60.87	29.12	60.9	29.4
c35	L	L	AU	01/05/91	11:41:24	1.56	60.91	29.64	60.90	29.31
c25	M	К	AU	12/26/90	10:07:28	2.25	60.83	29.31	60.83	29.21
c9	Μ	Μ	HC5	11/24/90	07:48:19	1.65	60.89	29.43	60.9	29.3
c85	М	M	HC6	04/17/91	09:48:32	1.94	60.97	29.17	60.9	29.4
c60	М	K	HC7	03/20/91	11:59:51	1.93	60.96	29.19	60.8	29.5
c82	М	М	HC7	04/12/91	14:23:49	1.68	60.97	29.15	60.8	29.5
c132	М	М	NR	06/04/91	12:00:02	1.11	61.07	29.35	•	
c108	0	01	AU	05/06/91	16:07:34	2.06	60.55	29.40	60.48	29.19
c124	0	O2	AU	05/28/91	16:13:45	1.87	60.53	29.25	60.51	29.09
c23	0	01	HC1	12/25/90	14:50:11	1.25	60.44	29.31	60.7	28.7
c73	0	<u>O2</u>	HC3	04/02/91	14:14:26	2.01	60.64	28.95	60.6	29.2

Table 2: One-hundred-forty-four events recorded at FINESA.

۰.

-

Ref.	V.	C.	H.C.	Date	Time	ml	Lat1	Lon1	Lat2	Lon2
c4	σ	01	NR	11/16/90	12:51:27	1.54	60.56	29.03		
c100	0	02	NR	04/25/91	13:08:04	1.67	60.54	29.24		
c121	Р	Р	AU	05/23/91	13:29:58	1.68	61.21	29.97	61.07	30.03
c128	P	Р	AU	05/31/91	13:33:41	1.43	61.15	29.88	61.10	29.88
c8	Р	Р	HC10	11/22/90	11:37:10	1.88	61.24	29.95	61.1	29.9
c14	P	Р	HC10	11/30/90	11:49:14	1.48	61.03	29.94	61.1	29.9
c59	Р	Р	HC10	03/19/91	16:16:40	1.67	61.17	29.69	61.1	29.9
c66	Р	Р	HC10	03/25/91	09:23:51	2.06	61.21	29.78	61.1	29.9
c72	Р	Р	HC10	04/02/91	09:48:18	1.68	61.13	30.11	61.1	29.9
c90	Р	Р	HC10	04/19/91	12:58:09	1.93	61.17	29.85	61.1	29.9
c87	P	P	HC11	04/18/91	13:14:26	1.59	61.40	29.79	61.1	30.2
c135	Р	Р	HC11	06/05/91	14:26:01	1.27	60.91	29.54	61.1	30.2
c68	P	Р	NR	03/28/91	09:19:48	1.53	61.27	29.82		
c105	Р	Р	NR	04/30/91	10:57:16	1.59	61.30	29.83		
c99	R	R	AU	04/25/91	12:42:54	1.89	61.85	30.40	61.86	30.66
c5	R	R	HC13	11/16/90	13:01:22	2.31	61.92	30.44	61.9	30.6
c13	R	R	HC13	11/30/90	09:50:22	1.43	61.93	30.50	61.9	30.6
c19	R	R	HC13	12/18/90	08:59:28	1.28	61.88	30.43	61.9	30.6
c20	R	R	HC13	12/19/90	11:01:31	2.05	62.00	30.38	61.9	30.6
c42	R	R	HC13	01/21/91	10:00:47	1.46	61.78	30.38	61.9	30.6
c71	R	R	HC13	03/29/91	13:41:28	1.89	61.96	30.28	61.9	30.6
c80	R	R	HC13	04/10/91	10:10:09	1.04	61.88	30.31	61.9	30.6
c84	R	R	HC13	04/17/91	08:46:47	1.26	61.86	30.45	61.9	30.6
c114	R	R	HC13	05/13/91	08:52:24	1.08	61.69	30.35	61.9	30.6
c61	R	R	NR	03/21/91	08:08:54	0.85	61.86	30.17		
c26	S	S	AU	12/27/90	10:36:59	2.54	61.80	30.59	61.75	30.79
c43	S	S	AU	01/25/91	12:20:30	2.12	61.94	30.44	61.86	30.66
c83	S	S	AU	04/13/91	11:52:34	2.55	61.81	30.66	61.71	30.92
c65	S	S	HC13	03/23/91	12:07:26	2.45	61.84	30.60	61.9	30.6
c110	S	S	NR	05/10/91	07:10:04	1.23	61.91	30.39		
c137	Т	W	AU	06/07/91	05:39:36	2.32	60.06	29.59	60.11	29.69
c143	Т	•	HB15	06/27/91	15:22:31	2.40	60.10	29.61	60.0	29.9
c118	U	Z2	NR	05/23/91	12:09:53	0.78	60.66	30.54		
c120	U	Z2	NR	05/23/91	13:27:15	0.75	60.62	30.57		
c130	V	S	NR	06/01/91	12:46:19	1.05	60.74	30.49		
c131	V	W	NR	06/01/91	13:41:11	1.12	60.82	30.53		
c103	W	W	AU	04/26/91	13:28:10	1.56	61.71	31.20	61.31	31.56
c127	W	W	AU	05/31/91	12:55:57	2.31	61.69	31.42	61.52	31.74
c138	W	Ŵ	AU	06/07/91	17:47:06	1.47	61.48	31.01	61.66	31.95
c1	W	W	HC14	11/05/90	09:46:23	1.74	61.42	31.55	61.4	31.6
c15	W	W	HC14	11/30/90	15:07:45	2.42	61.61	31.44	61.4	31.6
c24	W	W	HC14	12/26/90	08:47:57	1.82	61.70	31.42	61.4	31.6
ئ58	W	W	HC14	03/19/91	11:27:01	1.62	61.74	30.86	61.4	31.6
c91	W	W	HC14	04/19/91	13:08:41	1.80	61.67	31.23	61.4	31.6

Table 2: One-hundred-forty-four events recorded at FINESA.

Ref.	V.	C.	H.C.	Date	Time	ml	Lat1	Lon1	Lat2	Lon2
c141	W	W	HC14	06/26/91	13:47:53	2.32	61.84	31.02	61.4	31.6
c55	W	Ŵ	NR	03/13/91	10:11:56	0.87	61.88	31.05		
c96	W	W	NR	04/25/91	09:38:12	1.09	61.54	30.81		
c63	а	K	HC7	03/22/91	13:22:27	1.71	60.80	29.10	60.8	29.5
c70	b	D	HC5	03/29/91	13:38:51	2.37	60.97	28.86	60.9	29.3
c102	d	C	AU	04/26/91	12:13:15	2.04	61.02	28.99	60.99	29.02
c64	h	F	HC5	03/22/91	12:34:16	1.55	60.99	29.07	60.9	29.3
c119	k	K	AU	05/23/91	13:03:18	2.13	60.89	29.22	60.90	29.17
c6	k	0	HC6	11/20/90	13:13:52	1.31	60.79	28.51	60.9	29.4
c107	k	E	NR	05/04/91	14:27:50	0.50	60.86	29.52		
c139	m	M	HC6	06/26/91	12:04:19	1.93	60.90	29.17	60.9	29.4
c115	m	D	NR	05/14/91	05:51:45	0.83	61.05	29.14		

Table 2: One-hundred-forty-four events recorded at FINESA.

Ref.: Reference number

V. Visual classification results

C.:Cluster analysis results

H.C.: Helsinki mine location, (AU:automatic; NR: Not reported in the bulletin)

ml: IMS local magnitude

Lat1, Lon1: IMS coordinates

Lat2, Lon2: Helsinki coordinates

#: events that do not belong to the Karelian mine district

?: events not classified visually because of a low signal-to-noise ratio

.: Multiple event

 1.48° (165 km) and 2.65° (295 km) and the azimuthal coverage was 50°. For 89 of the events, a location was reported in the bulletin from Helsinki (monthly or weekly bulletins). In *Figure 5*, locations from the IMS are compared with the locations from the Helsinki bulletin. The average difference in location between events in these two bulletins was 23.01 ± 0.23 km for the manual locations and 24.45 ± 8.30 km for the automatic locations. The large discrepancy between the standard errors of these mislocations is due to four events that were grossly mislocated by the IMS

٤

Results of the visual classification

To do the visual classification, an upper limit on the number of groups was set above which signals were considered to belong to the same group even though small differences could be observed. The number of groups was defined by the results of cluster analysis. All of the events were classified into 18 basic groups. *Figure* δ shows the reference events



Figure 5: Eighty-nine of the events studied were reported in the Helsinki bulletin; 31 events were located automatically and 55 events had "manual locations". On this plot, IMS locations are compared with Helsinki locations, both manual and automatic. The largest discrepancies in location occur for events with a low signal-to-noise ratio. Two of them were clearly mislocated by the IMS. An interactive *f-k* analysis further confirmed that the azimuths determined by the IMS were erroneous.



Figure 6: Reference events for 18 groups determined by visual analysis of the data. Both unfiltered and filtered data were used. Groups are labeled from A through W.

that were identified as representing distinct waveforms. Each group has been labeled with a capital letter that will be used in the text to refer to these groups: A being the nearest group to FINESA and W being the farthest group. To illustrate the repeatability of these mining events, all the events belonging to groups K and S are displayed in *Figure 7*.

The events from group A were unique: the Lg phase was barely observable and the Rg phase was strong. For the ten events included in this group, the automatic phase identification of IMS identified only one event with an Lg phase. For the remaining nine events, the analyst reviewing the IMS solutions either added an Lg or renamed Rg to Lg.

B was the largest group with 18 events. In addition to having a small Lg-P time, their poor signal-to-noise ratio made them unique in the data set. These events were mostly located in Finland according to the IMS and their locations spread over a large area (100 km²). Because most of the **B** events had a magnitude less than 1.0, only the largest event (c111)

with a magnitude of 1.07 was reported in the Helsinki bulletin

Groups C through M included events related to a cluster of mines, all within an area of about 16 km². This subgroup was the most challenging to the clustering technique. The Lg-P time as well as the Rg-P time were not characteristic enough to allow a classification based on these parameters. However, the shape of the first P arrival along with the shape of the *PMP* phase were used in the visual classification. The first five seconds of the signal for representatives of the eight groups C through M are plotted in *Figure 8*. The interpretation of results from the cluster analysis made it necessary to distribute these events into ten main groups.

Events from group O were easily identified based on their unique Lg-P and Rg-P times.

The events from group P occurred at mines HC10 or HC11 according to the Helsinki bulletin. A thorough visual analysis along with the results of the cluster analysis, showed that their Lg-P times were within one second of each other. Only one mine has been located on SPOT photos for this area (SC10).

Events from groups **R** and **S**, which were all located at mine HC13 according to the Helsinki bulletin, were separated into two groups based on the shape of the first arrival. The difference in the shape of the first arrival could be due either to a difference in the source or to a path effect. **R** events exhibited an impulsive *P* arrival while the **S** events showed an emergent *P* followed by a strong second arrival (probably a *PMP* wave). Even though **R** events had smaller magnitudes, an *f*-k analysis showed that no emergent *P*-wave could be distinguished before the impulsive *P*. Figure 9 shows one event from each group with approximately the same magnitude. The Lg-P time was slightly larger for the events from **S**. This observation implies the presence of two mines so close to each other that a regular



Figure 7: Events from group K and S are plotted to show the difference between two groups as well as the repeatability of the signal in each group. A difference is clearly seen in the Lg-P time and in the presence of an Rg phase. For other groups, a difference is evident only in the shape of the first arrival, the Lg-P time being the same.



Figure 8: This figure shows the first arrival of reference events for groups C through M. Their Lg-P times and Rg-P times vary by less than one second. Differences between events can best be seen using unfiltered data to look at the shape of the first arrival. Although only four mines have been reported in this area in the Helsinki Bulletin, visual classification and cluster analysis determined 8 groups with more than one event and two with only one event. A careful SPOT photo analysis should be performed to determine whether or not this subdivision corresponds to a real distribution of the mines.





location routine could not distinguish between them. The quality of the SPOT photo presently available for this area did not allow a confirmation or a denial of the presence of two different mines.

Groups U and V contained two events each and were located within 22 km of each other by the IMS but far from any reported mine. They were the only events located in this particular area during the 7-month period. Two events (U) occurred on one day and the other two (V) one week later at about the same time of the day.

The events from group W showed a large scatter in their location despite their relatively high magnitudes. The IMS located most of them close to the SPOT location HC14. Their Lg-P time was unique and the Rg phase had very small amplitudes for most of the events.

Four events did not fit any of the 18 groups described above. The visual analysis showed that these four events were clearly mislocated by IMS; one event had a Lg-P time too large

for it to be located within the studied area, and the other three events had very poor IMS azimuth estimates and were actually located in Estonia. These observations were later confirmed by the interactive f-k analysis as discussed below.

Some conclusions based on the visual classification are given below:

- Most of the groups were easily distinguished using their Lg-P and Rg-P time.
- For groups C through M, the shape of the first arrival was a determining factor.
- The visual review showed that four events were mislocated in the IMS bulletin.
- The Rg phase was an important feature among the characteristics used to separate different groups. It has been shown that Rg waves are strongly site dependent (Murphy and Shah, 1988).
- The resulting visual classification agreed partially with the Helsinki "manual location" but more groups were found than the number of mines reported in the Helsinki bulletin.

Results of the interactive f-k analysis

A f-k analysis was performed on each event using a 0.5 s window starting at the beginning of the first arrival. Table 3 shows the average values and standard deviations of the appar-

ent velocity and the azimuth (Azim1, Ap. vel1) computed for each group. The average values included only the events for which the visual classification was in agreement with the cluster analysis. Only eight events belonging to group **B** were studied due to low signal-to-noise ratios. Although these events also had very low signal-to-noise ratios, the values computed with the *f*-*k* analysis were stable (small standard deviation). The velocities derived from the *f*-*k* analysis for the first arrival are plotted on a map (*Figure 10*) to show



Figure 10: Apparent velocities of the first arrival have been computed for each event using an *f-k* analysis method. This map shows the spatial distribution of the average apparent velocity for each mine group. North of Lake Ladoga, the velocity is around 6.45 km/s while South of the lake, the velocity is around 7.35 km/s and keeps increasing up to 8.77 km/s for the events located on an island in the Gulf of Finland. Strong variations in the thickness of the crust as well as sharp lateral boundaries in the crust can explain these changes in the apparent velocity.

group	Azim1	Ap.vel1	Azim2	Ap.vel2
A	121.27 0.23	6.45 0.04	122.24 3.49	6.77 0.69
В	79.32.0.53	7.42 0.07	78.93 8.55	8.10 0.70
C	120.71 0.36	6.47 0.06	112.99 1.45	7.02 0.09
D	120.68 0.35	6.44 0.09	117.59 2.06	6.76 0.04
E	117.58 7.87	6.65 0.40	111.18 2.14	7.11 0.08
F	120.70 0.86	6.44 0.02	111.73 1.28	6.94 0.15
G	121.05 0.00	6.41 0.00	120.04 0.47	7.12 0.12
Н	120.79 0.38	6.43 0.02	119.33 2.64	6.97 0.15
K	120.81 0.55	6.46 0.08	114.16 1.05	7.26 0.09
М	120.72 0.50	6.45 0.04	112.49 1.23	7.10 0.45
0	120.63 0.47	6.43 0.02	122.46 2.65	7.24 0.41
Р	102.02 0.68	7.35 0.08	106.20 6.05	7.01 0.28
R	79.43 0.50	7.36 0.06	76.22 1 1.62	6.94 1.06
S	93.52 1.95	7.17 0.38	81.74 10.63	7.39 0.21
U	101.54 0.00	7.46 0.00	106.97 3.47	6.97 0.01
W	101.93 0.97	7.30 0.00	95.01 8.52	7.28 0.42
c7	125.22	6.70	148.24	7.30
c79	134.77	7.19	134.13	8.92
c93	74.44	6.52	168.51	7.46
c104	117.11	8.83	31.39	6.49

Table 3: Azimuth and velocity values computed automatically by the IMS (Azim2, Ap.vel2) and the same values computed using *f*-*k* analysis (Azim1, Ap.vel1).

their spacial distribution. These events were located at distances from FINESA that correspond to the cross-over point of the travel-time curves for this area. Events located near the northern part of Lake Ladoga showed an apparent velocity ranging between 7.27 to 7.41 km/s; events located south of the lake exhibited an apparent velocity between 6.39 and 6.54 km/s; and events located further south in the Gulf of Finland (Island not plotted on the map) showed an apparent velocity close to 9.0 km/s. The apparent velocity of the second P type arrival was about the same for all of the events: 6.5 km/s. The events with the lowest velocity were also the closest to the array and the first arrival was a Pg phase. The two other sets of events were at about the same distance from FINESA and the difference in the apparent velocity could be explained by a difference in the travel path. In both cases, the first arrival was a Pn phase. However, the crust thickens rapidly from the Gulf of Finland to the Baltic Shield (*Figure 11*), thus increasing the apparent velocity.

A second P phase was seen, more or less clearly, within the two seconds following the first arrival on most of the signals. The second arrival was a PMP phase (reflected from the Moho). The presence of this phase was not related to the size of the shot. It did not appear on the signals from group **P** whose magnitudes were between 1.27 and 2.06 while the



Figure 11: Contour map of crustal thickness (km) and schematic map of Pn velocity for the Baltic Shield and adjacent areas based on DSS data. Lines of equal Moho depth are represented by thick solid lines for reliable data and dashed lines for unreliable data. Values of Pn velocity are: 1, 7.8 to 8.0; 2, 8.1 to 8.3; 3, 8.3 to 8.5. Thin solid lines denote DSS profiles. The crustal thickness varies from as little as 30 to 35 km near the coast to 50 to 55 km within the interior areas. Pn wave velocity varies from 7.8 to 8.0 km/s up to 8.3 to 8.5 km/s (the most frequently observed values are 8.1 to 8.2 km/s). There is no direct relationship between variations of the crustal thickness and Pn velocity (Ryaboy, 1990).

events from group **B** with smaller magnitudes (0.22 to 1.37), showed a sharp second arrival. The distance to the array also did not influence it's presence. For instance, events from group **R** did not show the *PMP* phase, but it was seen on signals from groups located both closer to and farther from the array. Lateral heterogeneities, in addition to differences in crustal thickness, may explain the presence of the PMP phase. Its interpretation will require further study.

F-k analysis was performed automatically in the IMS based on a computed arrival time. The window length was 3.0 s and began 1.0 s before the arrival time. Our results were based on a 0.5 s window length, but tests showed that there were no significant differences in the results for window lengths up to 5.0 s. Average values of the apparent velocity and the azimuth computed from the IMS values are reported in Table 3. Only the IMS results for the first arrival are shown and compared to the interactive f-k analysis. IMS also stores the results for secondary phases.

A careful f-k analysis provides important information about events. Even though computed values for the same group of events are very stable, the azimuth can vary by up to

one degree around this value. For our purpose, this is too large of a location error (about 5 km) and the visual classification, if associated with a known mine, provides more accurate results.

Cluster analysis results using waveforms

Figure 12 shows the cluster tree resulting from the analysis of filtered data (1-15 Hz). The lowest level of the tree (cross-correlation value of 0.4) shows the separation between events located north of Lake Ladoga and events located south of the lake. For the branch on the left, there is a dichotomy in the tree separating events with a large Rg phase from events with little to no Rg phase. At the level of the tree corresponding to a cross-correlation value of 0.57, the cluster separates the events into 10 groups. At this level, groups A, M, P, R, S, V and W are defined while events from group B are split into two subgroups (B1 and B2) and events from groups C through M belong to one large group.

The events from group **B** had the smallest signal-to-noise ratios. **B1** corresponds to the largest events with a magnitude greater than 1.0 while the other events group in **B2**. If a narrower frequency band is used (1.0 to 8.0 Hz), no subgroups are obtained.

The Z branch of the cluster tree does not define a group of events that look alike. These events cannot be classified in any group because of low signal-to-noise ratio, uniqueness, or the fact that they are multiple events.

Figure 13 shows the cluster tree using the 55 events from groups C through O. Seven









small groups can be distinguished from one another at a cross-correlation value of 0.67, but a level of 0.75 must be used to distinguish between groups F, G, and H. The high level of correlation is related to the stability of the shape of the signal in this tight grouping of events. Note that the M events, previously grouped together by the visual classification, have been split into two groups here.

The visual classification differed from the cluster analysis for ten of the 144 events with-

out including the events that were labeled with lower case letters. These ten events were characterized either by a low signal-to-noise ratio or were multiple events. Nine events had such a low signal-to-noise ratio that a visual classification could not be performed. The events labeled with lower case letters were unique and could not be classified in any group, thus their visual classification was questionable.

Four events were erroneously included in the data set due to large azimuth errors in the IMS and actually belong to mining districts in other areas. This error was discovered through interactive f-k analysis. The P-Lg time for these three events was nearly identical to the P-Lg time for mines in the study area. One of these events erroneously clustered with events from group W while the three other events were classified with either group

Z1 or Z2. These events clustered at low levels with events most similar to them.

Cluster analysis results using phase parameters data

The use of *f*-*k* parameters in addition to other measurements (SNR, period, azimuth, slowness, etc.) computed during the automatic processing of the IMS in a cluster analysis gives results comparable to those that would be obtained by grouping events based on IMS locations. The addition of polarization parameters from three-component polarization analysis did not improve the results of this cluster analysis because polarization parameters are more characteristic of the receiver than of the source.

RESULTS OF THE CLUSTER ANALYSIS USING ARCESS AND NORESS DATA

Among the 144 events, 39 were recorded at ARCESS and 31 at NORESS. Both arrays are located at a distance of about 9° from the mining district. Despite their relatively high magnitudes (1.56 to 2.66 for ARCESS data and 2.01 to 2.66 for NORESS data), the data had low signal-to-noise ratios (*Figure 14*). Envelopes were computed using data filtered between 2.0 and 5.0 Hz. The signal length used in the cross-correlation process was set to 175 s including 15 s before the arrival. The cluster analysis based on these data gave poor





quality results (*Figure 15* and 16) probably due to poor signal-to-noise ratios and the lack of an Rg phase in the NORESS and ARCESS data.

ASSOCIATION OF EVENTS TO MINE LOCATION

Having defined clusters of events both visually and through cluster analysis, the task

becomes associating a mine with each group of events. Twenty-one groups of events were identified above, but only fifteen mines have been listed in the Helsinki Bulletin and even fewer were located on SPOT photos by Fox (1990). Because ground truth was not available for the events used at the time this study was made, seismic locations reported in the IMS and Helsinki Bulletins and the results of the interactive f-k analysis was the only information used to make the association. Other information (possibly provided by mining authorities) may be used in future analysis efforts.

Events from group A and from group I were all located close to mine SC1 (HC1) by the IMS. The visual analysis showed a clear difference in Lg-P times between the two groups, however: events from group I have Lg-P times 8 s greater than the events from group A. The Helsinki bulletin reported an automatic location for event c10 from group A that was

close to mine N114. The Helsinki bulletin also reported a manual location for event c57 from group I at mine HC1. This discrepancy between bulletins may be explained by a misidentification of the phases for group A in the IMS. As noted in the visual description of events from group A, the analyst often renamed the Rg as an Lg resulting in a location too far from the array. The distance between HC1 and N114 is 43 km, which corresponds to a difference in travel-time of about 8.25 s. Event c111 with the only Lg correctly identified by the IMS for this group of events was located near mine N114. Thus, events from group A are associated with mine N114, and events from group I are associated with mine SC1 (HC1)

Only one event from group **B** was manually located by the Helsinki analysts. In Table 1, this mine is labeled N117. The interactive f-k analysis gave stable values for the events in this group and most of them could be located at this mine. The large scatter observed in the IMS locations for these events is explained by their small magnitudes.

Events from group O were located by the IMS close to mine SC3. There was little scatter in the locations. One of these events was manually located at mine HC3 by the Helsinki analysts.

Events from group P were located at one of two different mines, HC10 or HC11, accord-



Figure 15: Result of the cluster analysis performed on 31 events recorded at ARCESS. Data were filtered between 2 to 5 Hz. Events with the highest signal-to-noise ratio give reasonable results.



Figure 15: Result of the cluster analysis performed on 31 events recorded at NORESS. The same frequency band applied to the ARCESS data was applied to these signals.

ing to the Helsinki bulletin. For this area, only one mine was reported on SPOT photos by Fox (1990). Lg-P times for this group vary by 1.0 s, which is not enough to justify the distance between HC10 and HC11, but could correspond to events from different parts of the same big mine.

Seven events from group \mathbf{R} and one event from group \mathbf{S} were manually associated with mine HC13 according to the Helsinki bulletin. Three other events from group \mathbf{S} were automatically located around this mine by the same bulletin. Any difference in location

between the two groups was not clearly seen using the IMS location. No reliable SPOT photo was available for this area. Events from groups \mathbf{R} and \mathbf{S} may have been from the same large mine or from two different mines.

Events from group T were located on an island in the Gulf of Finland where a mine has been identified on a SPOT photo (SB15).

Events from group U could not be associated with any mine as none were reported on this part of the southern coast of Lake Ladoga. Although two of the events had an IMS magnitude greater than 1.0, they were not reported in the weekly bulletin from Helsinki. These events may have been related to construction activity.

Events from group V were located far from each other by the IMS, probably because of their small magnitude, and they were not reported in the Helsinki bulletin. No mine was

reported close to these events.

The Helsinki bulletin reported five group W events and located them all at mine HC14. The IMS locations for the set of events were closer to mine SC14.

Associating the events from groups C through M with specific mines was not possible with the limited information available. These groups contained events that were located in a small area where only five mines were identified in the Helsinki bulletin and on SPOT photos (SPOT locations can include several small mines). In this particular case, the limitations of the current seismic location procedures are shown. Even when careful f-k analysis was used to obtain more stable values of the azimuth, the results were not accurate enough to allow a perfect match between event groups and mines. Possible flaws in the Helsinki manual locations were also observed. In some cases, several events that cluster together and were identified as belonging to the same group visually, were manually located by the Helsinki analysts as originating from different mines.

SUMMARY AND CONCLUSIONS

Events recorded at the FINESA, ARCESS and NORESS arrays were studied with the aim of identifying and characterizing mines located north of St. Petersburg. Both waveforms and parameters computed during the automatic process were used in this study.

Cross-correlation values computed between the envelopes of the vertical component seismograms were found to be the best similarity measurements for use in a cluster analysis. This method was verified by a visual classification of the waveforms along with an interactive f-k analysis. The f-k analysis provided explanations for the questionable results of

the cluster analysis and showed that the assumption of identical travel paths for all of these events was not valid. Eighteen different groups (with more than one event per group) were identified from the 144 events by both cluster analysis and visual classification. Four events were not classified as they were originally mislocated and do not really belong to this mining district.

The Helsinki and IMS bulletins were used as references for the locations of these events. Only 89 events were reported in the Helsinki bulletin for the covered period of time. Fiftyeight events had been "manually located" by the Helsinki analysts at 15 different mines. Their classification agreed with our grouping for events located in areas where only one mine had been reported: HC3, HC10, HC13, HC14 and HB15.

SPOT photos from this area are currently being analyzed. They provide information about the size and the actual number of active mines. This information allows a better classifica-

tion of the events as the number of groups (mines) is predetermined. But, in order to improve the matching between groups of events and mines, information concerning the origin time, the "true" location of the shot, as well as the shooting parameters and the yield of the shot need to be gathered directly from the mines.

Two other mining districts showing a high concentration of mines are available to test this method. The implementation of this method in an automatic system such IMS would required a selection of master events for each mine. A "pre-location" of a given event by the automatic routine would determine what mining district was relevant. The event would be compared to the reference events built for this particular district. The maximum value of the cross-correlation of the envelope functions would be compared to the thresholds previously determined for each mine by cluster analysis. As seen above, a threshold needs to be set in order to delimit the different groups. If this threshold is too low to separate several groups, a second comparison should be performed using a smaller set of reference events. For mislocated events, the result of the cross-correlation with the master events should give a value below the threshold. In such cases, these events would be re-analyzed carefully.

ACKNOWLEDGEMENTS

The program (massproc) used to compute cross-correlation values directly from the database was written by John Coyne, SAIC. The program (geotool) used for all waveform comparisons, stacking, and f-k analysis was written and customized by Ivan Henson, Teledyne Geotech with assistance from John Coyne of SAIC. Information about SPOT photos was provided by Warren Fox, SAIC. The authors would like to thank Jerry Carter for his constant encouragement.

REFERENCES

- Bache T.C., J.T. Anderson, D. Baumgardt, S.R. Bratt, W.E. Farrell, R.F. Fung, J.W. Given, A.S. Henson, C. Kobryn, H.J. Swanger, J. Wang (1990). Intelligent Array System. Final Technical Report, October 1990. SAIC, San Diego.
- Baumgardt, D.R. and K. A. Ziegler (1988). Spectral Evidence for Source Multiplicity in Explosions: Application to Regional Discrimination of Earthquakes and Explosions. Bull. Seism. Soc. Am. 78, 5, 1773-1795.
- Baumgardt, D.R. (1987). Case-based Reasoning Applied to Regional Seismic Event Characterization, DARPA/AFGL Seismic Research Symposium, Harbor House, Nantucket, MA, 15-18 June 1987, 173-178, GL-TR-90-0300, ADA229025.
- Bennett, T.J., B.W. Barker, K.L. McLaughlin, J.R. Murphy (1989). Regional Discrimination of Quarry Blasts, Earthquakes and Underground Nuclear Explosions. S-Cubed Final Report SSS-TR-89-1039, GL-TR-89-0114. ADA223148
- Blandford, R.R. (1982). Seismic Event Discrimination. Bull. Seism. Soc. Am. 72, S69-S87.
- Bratt S.R., H.J. Swanger, R.J. Stead, F. Ryall, T.C. Bache (1990). Initial Results from the Intelligent Monitoring System. Bull. Seism. Soc. Am. 80b, 1852-1873.
- Dysart, P.S. and J.J. Pulli (1988). Waveform and Spectral Characteristics of Regional Earthquakes and Chemical Explosions Recorded at the NORESS Array. *Technical Report C88-01*. Center for Seismic Studies, Internal Report.
- Everitt, B. (1986). Cluster Analysis. Halsted Press, Division of John Wiley & Sons, New York (2nd Edition).
- Fox, W. (1990). Satellite Imagery of Areas in Northeast Europe. Internal Report. SAIC, San Diego.
- Israelsson H. (1990). Analysis of High Frequency Data. Scientific report #1, GL-TR-90-0299. Center for Seismic Studies, ADA231804.
- Israelsson H., and J. Carter (1991). Analysis of High Frequency Seismic Data. Scientific report #2, PL-TR-91-2032. Center for Seismic Studies, ADA235579

- Murphy, J.R. and H.K. Shah (1988). An Analysis of the Effects of Site Geology on the Characteristics of Near-field Rayleigh Waves.Bull.Seism.Soc.Am. 78, 1, 64-82.
- Pomeroy, P.W., W.J. Best, T.W. McEvilly (1982). Test Ban Treaty Verification with Regional Data A Review. Bull. Seism. Soc. Am. 72, S89-S129.
- Ryaboy, V. (1990). Upper Mantle Structure along a Profile from Oslo (NORESS) to Helsinki to Leningrad, based on Explosion Seismology. Bull. Seism. Soc. Am. 80, 6, 2194-2213.
- Smith, A.T. (1989). High-Frequency Seismic Observations and Models of Chemical Explosions: Implications for the Discrimation of Ripple-fired Mining Blasts. Bull. Seism. Soc. Am. 79, 1089-1110.
- Uski, M., E. Polkonen, M. Franssila, M.Raime (1990-1991). Seismic Events in Northern Europe. Ed. H. Korhonen.Institute of Seismology, University of Helsinki, Helsinki, Finland.

Prof. Thomas Ahrens Seismological Lab, 252-21 Division of Geological & Planetary Sciences California Institute of Technology Pasadena, CA 91125

Prof. Keiiti Aki Center for Earth Sciences University of Southern California University Park Los Angeles, CA 90089-0741

Prof. Shelton Alexander Geosciences Department 403 Deike Building The Pennsylvania State University University Park, PA 16802

Dr. Ralph Alewine, III DARPA/NMRO 3701 North Fairfax Drive Arlington, VA 22203-1714

Prof. Charles B. Archambeau CIRES University of Colorado Boulder, CO 80309

Dr. Thomas C. Bache, Jr. Science Applications Int'l Corp. 10260 Campus Point Drive San Diego, CA 92121 (2 copies)

Prof. Muawia Barazangi Institute for the Study of the Continent Cornell University Ithaca, NY 14853

Dr. Jeff Barker Department of Geological Sciences State University of New York at Binghamton Vestal, NY 13901

Dr. Douglas R. Baumgardt ENSCO, Inc 5400 Port Royal Road Springfield, VA 22151-2388

Dr. Susan Beck Department of Geosciences Building #77 University of Arizona Tuscon, AZ 85721 Dr. T.J. Bennett S-CUBED A Division of Maxwell Laboratories 11800 Sunrise Valley Drive, Suite 1212 Reston, VA 22091

Dr. Robert Blandford AFTAC/TT, Center for Seismic Studies 1300 North 17th Street Suite 1450 Arlington, VA 22209-2308

Dr. G.A. Bollinger Department of Geological Sciences Virginia Polytechnical Institute 21044 Derring Hall Blacksburg, VA 24061

Dr. Stephen Bratt Center for Seismic Studies 1300 North 17th Street Suite 1450 Arlington, VA 22209-2308

Dr. Lawrence Burdick Woodward-Clyde Consultants 566 El Dorado Street Pasadena, CA 91109-3245

Dr. Robert Burridge Schlumberger-Doll Research Center Old Quarry Road Ridgefield, CT 06877

Dr. Jerry Carter Center for Seismic Studies 1300 North 17th Street Suite 1450 Arlington, VA 22209-2308

Dr. Eric Chael Division 9241 Sandia Laboratory Albuquerque, NM 87185

Prof. Vernon F. Cormier Department of Geology & Geophysics U-45, Room 207 University of Connecticut Storrs, CT 06268

Prof. Steven Day Department of Geological Sciences San Diego State University San Diego, CA 92182

1

Marvin Denny U.S. Department of Energy Office of Arms Control Washington, DC 20585

Dr. Zoltan Der ENSCO, Inc. 5400 Port Royal Road Springfield, VA 22151-2388

Prof. Adam Dziewonski Hoffman Laboratory, Harvard University Dept. of Earth Atmos. & Planetary Sciences 20 Oxford Street Cambridge, MA 02138

Prof. John Ebel Department of Geology & Geophysics Boston College Chestnut Hill, MA 02167

Eric Fielding SNEE Hall INSTOC Cornell University Ithaca, NY 14853

Dr. Mark D. Fisk Mission Research Corporation 735 State Street P.O. Drawer 719 Santa Barbara, CA 93102

Prof Stanley Flatte Applied Sciences Building University of California, Santa Cruz Santa Cruz, CA 95064

Dr. John Foley NER-Geo Sciences 1100 Crown Colony Drive Quincy, MA 02169

Prof. Donald Forsyth Department of Geological Sciences Brown University Providence, RI 02912

Dr. Art Frankel U.S. Geological Survey 922 National Center Reston, VA 22092 Dr. Cliff Frolich Institute of Geophysics 8701 North Mopac Austin, TX 78759

Dr. Holly Given IGPP, A-025 Scripps Institute of Oceanography University of California, San Diego La Jolla, CA 92093

Dr. Jeffrey W. Given SAIC 10260 Campus Point Drive San Diego, CA 92121

Dr. Dale Glover Defense Intelligence Agency ATTN: ODT-1B Washington, DC 20301

Dr. Indra Gupta Teledyne Geotech 314 Montgomery Street Alexanderia, VA 22314

Dan N. Hagedon Pacific Northwest Laboratories Battelle Boulevard Richland, WA 99352

Dr. James Hannon Lawrence Livermore National Laboratory P.O. Box 808 L-205 Livermore, CA 94550

Dr. Roger Hansen HQ AFTAC/TTR Patrick AFB, FL 32925-6001

Prof. David G. Harkrider Seismological Laboratory Division of Geological & Planetary Sciences California Institute of Technology Pasadena, CA 91125

Prof. Danny Harvey CIRES University of Colorado Boulder, CO 80309 Prof. Donald V. Helmberger Seismological Laboratory Division of Geological & Planetary Sciences California Institute of Technology Pasadena, CA 91125

Prof. Eugene Herrin Institute for the Study of Earth and Man Geophysical Laboratory Southern Methodist University Dallas, TX 75275

Prof. Robert B. Herrmann Department of Earth & Atmospheric Sciences St. Louis University St. Louis, MO 63156

Prof. Lane R. Johnson Seismographic Station University of California Berkeley, CA 94720

Prof. Thomas H. Jordan Department of Earth, Atmospheric & Planetary Sciences Massachusetts Institute of Technology Cambridge, MA 02139

Prof. Alan Kafka Department of Geology & Geophysics Boston College Chestnut Hill, MA 02167

Robert C. Kemerait ENSCO, Inc. 445 Pineda Court Melbourne, FL 32940

Dr. Max Koontz U.S. Dept. of Energy/DP 5 Forrestal Building 1000 Independence Avenue Washington, DC 20585

Dr. Richard LaCoss MIT Lincoln Laboratory, M-200B P.O. Box 73 Lexington, MA 02173-0073

Dr. Fred K. Lamb University of Illinois at Urbana-Champaign Department of Physics 1110 West Green Street Urbana, IL 61801 Prof. Charles A. Langston Geosciences Department 403 Deike Building The Pennsylvania State University University Park, PA 16802

Jim Lawson, Chief Geophysicist Oklahoma Geological Survey Oklahoma Geophysical Observatory P.O. Box 8 Leonard, OK 74043-0008

Prof. Thorne Lay Institute of Tectonics Earth Science Board University of California, Santa Cruz Santa Cruz, CA 95064

Dr. William Leith U.S. Geological Survey Mail Stop 928 Reston, VA 22092

Mr. James F. Lewkowicz Phillips Laboratory/GPEH Hanscom AFB, MA 01731-5000(2 copies)

Mr. Alfred Lieberman ACDA/VI-OA State Department Building Room 5726 320-21st Street, NW Washington, DC 20451

Prof. L. Timothy Long School of Geophysical Sciences Georgia Institute of Technology Atlanta, GA 30332

Dr. Randolph Martin, III New England Research, Inc. 76 Olcott Drive White River Junction, VT 05001

Dr. Robert Masse Denver Federal Building Box 25046, Mail Stop 967 Denver, CO 80225

Dr. Gary McCartor Department of Physics Southern Methodist University Dallas, TX 75275 Prof. Thomas V. McEvilly Seismographic Station University of California Berkeley, CA 94720

Dr. Art McGarr U.S. Geological Survey Mail Stop 977 U.S. Geological Survey Menlo Park, CA 94025

Dr. Keith L. McLaughlin S-CUBED A Division of Maxwell Laboratory P.O. Box 1620 La Jolla, CA 92038-1620

Stephen Miller & Dr. Alexander Florence SRI International 333 Ravenswood Avenue Box AF 116 Menlo Park, CA 94025-3493

Prof. Bernard Minster IGPP, A-025 Scripps Institute of Oceanography University of California, San Diego La Jolla, CA 92093

Prof. Brian J. Mitchell Department of Earth & Atmospheric Sciences St. Louis University St. Louis, MO 63156

Mr. Jack Murphy S-CUBED A Division of Maxwell Laboratory 11800 Sunrise Valley Drive, Suite 1212 Reston, VA 22091 (2 Copies)

Dr. Keith K. Nakanishi Lawrence Livermore National Laboratory L-025 P.O. Box 808 Livermore, CA 94550

Dr. Carl Newton Los Alamos National Laboratory P.O. Box 1663 Mail Stop C335, Group ESS-3 Los Alamos, NM 87545

Dr. Bao Nguyen HQ AFTAC/TTR Patrick AFB, FL 32925-6001 Prof. John A. Orcutt IGPP, A-025 Scripps Institute of Oceanography University of California, San Diego La Jolla, CA 92093

Prof. Jeffrey Park Kline Geology Laboratory P.O. Box 6666 New Haven, CT 06511-8130

Dr. Howard Patton Lawrence Livermore National Laboratory L-025 P.O. Box 808 Livermore, CA 94550

Dr. Frank Pilotte HQ AFTAC/TT Patrick AFB, FL 32925-6001

Dr. Jay J. Pulli Radix Systems, Inc. 2 Taft Court, Suite 203 Rockville, MD 20850

Dr. Robert Reinke ATTN: FCTVTD Field Command Defense Nuclear Agency Kirtland AFB, NM 87115

Prof. Paul G. Richards Lamont-Doherty Geological Observatory of Columbia University Palisades, NY 10964

Mr. Wilmer Rivers Teledyne Geotech 314 Montgomery Street Alexandria, VA 22314

Dr. George Rothe HQ AFTAC/TTR Patrick AFB, FL 32925-6001

Dr. Alan S. Ryall, Jr. DARPA/NMRO 3701 North Fairfax Drive Arlington, VA 22209-1714 Dr. Richard Sailor TASC, Inc. 55 Walkers Brook Drive Reading, MA 01867

Prof. Charles G. Sammis Center for Earth Sciences University of Southern California University Park Los Angeles, CA 90089-0741

> Prof. Christopher H. Scholz Lamont-Doherty Geological Observatory of Columbia University Palisades, CA 10964

Dr. Susan Schwartz Institute of Tectonics 1156 High Street Santa Cruz, CA 95064

Secretary of the Air Force (SAFRD) Washington, DC 20330

Office of the Secretary of Defense DDR&E Washington, DC 20330

Thomas J. Sereno, Jr. Science Application Int'l Corp. 10260 Campus Point Drive San Diego, CA 92121

Dr. Michael Shore Defense Nuclear Agency/SPSS 6801 Telegraph Road Alexandria, VA 22310

Dr. Matthew Sibol Virginia Tech Seismological Observatory 4044 Derring Hall Blacksburg, VA 24061-0420

Prof. David G. Simpson IRIS, Inc. 1616 North Fort Myer Drive Suite 1440 Arlington, VA 22209 Donald L. Springer Lawrence Livermore National Laboratory L-025 P.O. Box 808 Livermore, CA 94550

Dr. Jeffrey Stevens S-CUBED A Division of Maxwell Laboratory P.O. Box 1620 La Jolla, CA 92038-1620

Lt. Col. Jim Stobie ATTN: AFOSR/NL Bolling AFB Washington, DC 20332-6448

Prof. Brian Stump Institute for the Study of Earth & Man Geophysical Laboratory Southern Methodist University Dallas, TX 75275

Prof. Jeremiah Sullivan University of Illinois at Urbana-Champaign Department of Physics 1110 West Green Street Urbana, IL 61801

Prof. L. Sykes Lamont-Doherty Geological Observatory of Columbia University Palisades, NY 10964

Dr. David Taylor ENSCO, Inc. 445 Pineda Court Melbourne, FL 32940

Dr. Steven R. Taylor Los Alamos National Laboratory P.O. Box 1663 Mail Stop C335 Los Alamos, NM 87545

Prof. Clifford Thurber University of Wisconsin-Madison Department of Geology & Geophysics 1215 West Dayton Street Madison, WS 53706

Prof. M. Nafi Toksoz Earth Resources Lab Massachusetts Institute of Technology 42 Carleton Street Cambridge, MA 02142 Dr. Larry Turnbull CIA-OSWR/NED Washington, DC 20505

Dr. Gregory van der Vink IRIS, Inc. 1616 North Fort Myer Drive Suite 1440 Arlington, VA 22209

Dr. Karl Veith EG&G 5211 Auth Road Suite 240 Suitland, MD 20746

Prof. Terry C. Wallace Department of Geosciences Building #77 University of Arizona Tuscon, AZ 85721

Dr. Thomas Weaver Los Alamos National Laboratory P.O. Box 1663 Mail Stop C335 Los Alamos, NM 87545

Dr. William Wortman Mission Research Corporation 8560 Cinderbed Road Suite 700 Newington, VA 22122

Prof. Francis T. Wu Department of Geological Sciences State University of New York at Binghamton Vestal, NY 13901

AFTAC/CA (STINFO) Patrick AFB, FL 32925-6001

DARPA/PM 3701 North Fairfax Drive Arlington, VA 22203-1714

DARPA/RMO/RETRIEVAL 3701 North Fairfax Drive Arlington, VA 22203-1714 DARPA/RMO/SECURITY OFFICE 3701 North Fairfax Drive Arlington, VA 22203-1714

HQ DNA ATTN: Technical Library Washington, DC 20305

Defense Intelligence Agency Directorate for Scientific & Technical Intelligence ATTN: DTIB Washington, DC 20340-6158

Defense Technical Information Center Cameron Station Alexandria, VA 22314 (2 Copies)

TACTEC Battelle Memorial Institute 505 King Avenue Columbus, OH 43201 (Final Report)

Phillips Laboratory ATTN: XPG Hanscom AFB, MA 01731-5000

Phillips Laboratory ATTN: GPE Hanscom AFB, MA 01731-5000

Phillips Laboratory ATTN: TSML Hanscom AFB, MA 01731-5000

2

Phillips Laboratory ATTN: SUL Kirtland, NM 87117 (2 copies)

Dr. Michel Bouchon I.R.I.G.M.-B.P. 68 38402 St. Martin D'Heres Cedex, FRANCE Dr. Michel Campillo Observatoire de Grenoble I.R.I.G.M.-B.P. 53 38041 Grenoble, FRANCE

Dr. Kin Yip Chun Geophysics Division Physics Department University of Toronto Ontario, CANADA

Prof. Hans-Peter Harjes Institute for Geophysic Ruhr University/Bochum P.O. Box 102148 4630 Bochum 1, GERMANY

Prof. Eystein Husebye NTNF/NORSAR P.O. Box 51 N-2007 Kjeller, NORWAY

David Jepsen Acting Head, Nuclear Monitoring Section Bureau of Mineral Resources Geology and Geophysics G.P.O. Box 378, Canberra, AUSTRALIA

Ms. Eva Johannisson Senior Research Officer National Defense Research Inst. P.O. Box 27322 S-102 54 Stockholm, SWEDEN

Dr. Peter Marshall Procurement Executive Ministry of Defense Blacknest, Brimpton Reading FG7-FRS, UNITED KINGDOM

Dr. Bernard Massinon, Dr. Pierre Mechler Societe Radiomana 27 rue Claude Bernard 75005 Paris, FRANCE (2 Copies)

Dr. Svein Mykkeltveit NTNT/NORSAR P.O. Box 51 N-2007 Kjeller, NORWAY (3 Copies)

2

Prof. Keith Priestley University of Cambridge Bullard Labs, Dept. of Earth Sciences Madingley Rise, Madingley Road Cambridge CB3 OEZ, ENGLAND Dr. Jorg Schlittenhardt Federal Institute for Geosciences & Nat'l Res. Postfach 510153 D-3000 Hannover 51, GERMANY

Dr. Johannes Schweitzer Institute of Geophysics Ruhr University/Bochum P.O. Box 1102148 4360 Bochum 1, GERMANY