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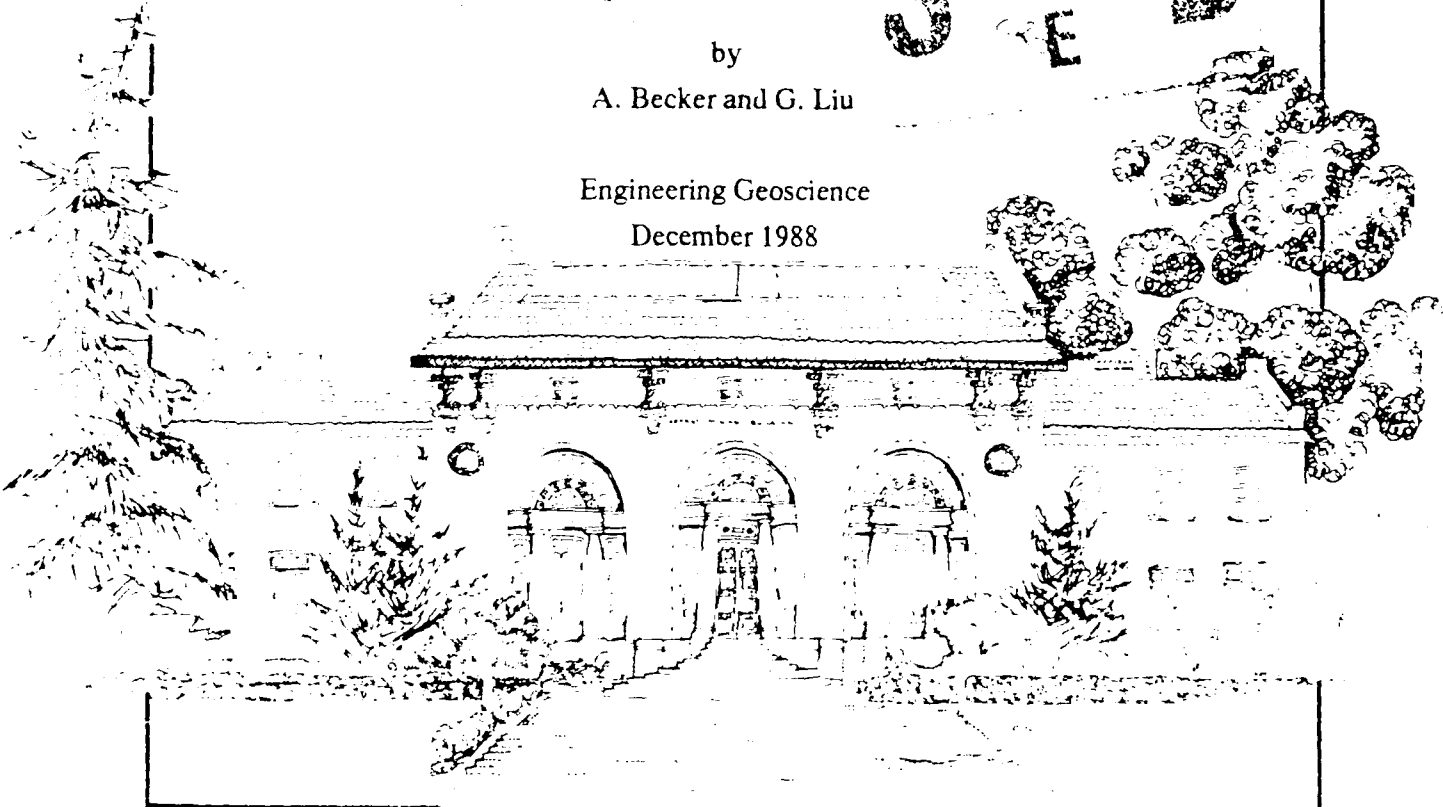
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by

A. Becker and G. Liu

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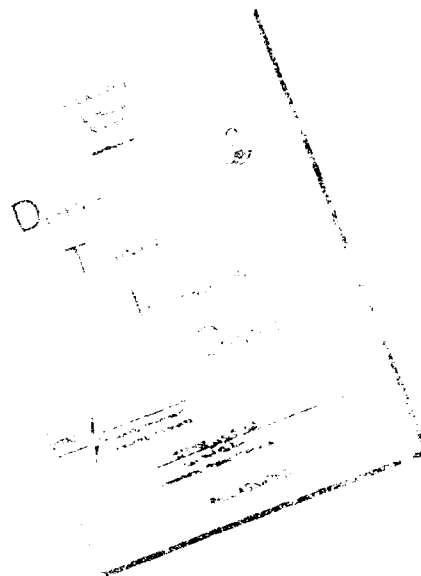
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Summary

A conventional frequency domain helicopter-borne electromagnetic (HEM) system can be used to map sea ice keels with a reasonable degree of accuracy. A preliminary interpretation of the acquired data can be made manually with the help of a nomogram or automated with the use of a table look-up routine on a small computer. Such data may also be more accurately interpreted with the use of an adaptation of Occam's inversion. This scheme allows for the practical non-uniqueness of the inverse solution but selects the smoothest keel shape that is consistent with the field data. The inversion method is much more computationally intensive than the table look-up technique. While the latter can be implemented on a small computer to form an interactive in-flight interpretation system, the inversion technique involves many forward computations and, for the present, is best reserved for post flight data analysis. It is possible that this difficulty can be resolved with the use of specialized computing equipment.

In the strict sense both the proposed interpretation techniques are only suitable for use on data acquired over two dimensional features whose strike length (measured in a direction perpendicular to the flight line) is much greater than the flight height. An examination of the anomalies for three-dimensional keels however, reveals that good data interpretation is possible whenever the keel strike length exceeds the system height by a factor of three.

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Airborne Electromagnetic Mapping of Sea Ice Keels

Alex Becker and Guimin Liu

1.0 Introduction

Sea ice is a unique environment encountered in most Arctic work. This includes the transportation of vehicles through ice-covered waters, the construction of offshore drilling structures, and the safe operation of submarines. In such circumstances, knowledge of the thickness and properties of sea ice is important. This is true especially for ice keels, which constitute a local downward indentation of the ice-water interface.

Recently, Canpolar consultants (1985) reviewed the possible techniques for remotely measuring sea ice thickness, among these they included airborne impulse radar and airborne electromagnetics. The impulse radar technique has been very successful in some areas (Kovacs and Morey, 1985). It may however, occasionally yield erroneous data when saline moisture zones exist within the ice cover. Becker et al. (1983), on the other hand, examined the feasibility of using frequency-domain airborne electromagnetics to determine sea ice thickness. At low frequencies, sea ice is practically transparent to electromagnetic waves and the observed secondary magnetic field can be used to estimate the distance from the EM system boom to the ice-water interface. At the same time, a laser altimeter installed on the boom measures the distance to the surface of the ice, so that the difference of these two distances is the sea ice thickness. This technique was successfully used in mapping the average sea ice thickness in Prudhoe Bay, Alaska, but it failed in an area of a multi-year pressure keel because of the inappropriate one-dimensional (1-D) techniques used to interpret data (Kovacs, et al., 1986; Becker, et al., 1987). In order to recover

the geometry of the keel, two- or three-dimensional interpretation techniques are required.

In this report we primarily concern ourselves with the development of techniques for interpreting AEM data for two-dimensional keels which are assumed to be infinitely long in a direction perpendicular to the flight line. The computational methods used to construct the necessary forward solutions are outlined in the Appendix. These serve either for the construction of interpretation nomograms or in the implementation of a numerical inversion scheme. In order to establish the validity of the two-dimensional interpretation scheme we also examine the effects of finite keel length on the observable anomalies.

In terms of computer time, the modeling technique we employ is particularly well suited for the interpretation of the AEM data collected over sea ice keels. Using the conventional finite element method, 30 minutes CRAY II CPU time is required to compute the AEM system response along a traverse line over an ice keel. However, for a similar problem, using an approach rooted in potential theory, the computation takes about 10 seconds on IBM 3090 (equivalent to about 3 seconds on CRAY) with a gain more than two orders of magnitude in speed. We will show that the assumption of a perfectly conducting sea water model is not a significant drawback since it can be used for any system operation frequency greater than 30 KHz.

2.0 AEM Anomalies for Ice Keels

Let us first briefly examine the airborne electromagnetic response to two-dimensional sea ice keels as a function of their size and shape (cf. Fig.1). For these calculations, the electromagnetic system has a coil separation of 6.5m and is "flown" 25m above the upper ice surface which is flat. With the exception of the zone containing the keel, the sea ice is 5m thick and is assumed to have negligible electrical conductivity. The data is calculated for both the vertical-coaxial coil configuration (HX system) and the horizontal-coplanar one (HZ system). The keel is assumed to have

the shape of a Gaussian curve and its parameters and geometry are shown in Figure 2.

Figure 3 shows the HX system response for two different keels with parameters $A = 12\text{m}$, and $W = 28\text{m}$ (solid line) or 14m (dashed line). The observed secondary magnetic field is expressed in parts per million (ppm) of the primary field at the receiver. There is a big anomaly in the system response due to either of these ice keels. For a keel width $W = 28\text{m}$, the maximum anomaly is about 35% of the background response and we define this to be the "percent anomaly". In order to characterize the shape of the observed anomaly, we define the "anomaly width", q , to be the width of the anomaly at one half of its maximum value. For a keel width $W = 28\text{m}$, the anomaly width q is 45m . When the keel width W is halved and other conditions are kept unchanged, the HX system anomaly is significantly reduced. In this case, the maximum anomaly (or percent anomaly) decreases by 37%, and the anomaly width decreases by 33%.

The situation for the HZ system is similar. Figure 4 shows the HZ system response due to the same two ice keels that were previously examined. For a keel defined by $A = 12\text{m}$ and $W = 28\text{m}$, the percent anomaly for the HZ system response is 32%. The anomaly width q is 66m in this case, which is much larger than that seen for the HX system. When the keel width is halved, the HZ system anomaly also decreases greatly. In this case, the anomaly amplitude (or percent anomaly) drops by 32% and the anomaly width decreases by 20%. From the comparison of Figures 3 and 4, we can see that the HX system response tracks the shape of the keel closely, while the HZ system anomaly is much wider.

Next, the keel drawdown A is halved while the keel width is kept constant at $W = 28\text{m}$. The corresponding HX system response is shown in Figure 5 (dashed line), and HZ system response is shown in Figure 6 (dashed line). The anomaly amplitude contracts by 29% for the HX system and by 37% for the HZ system. However, the

corresponding decrease in the anomaly width is small, 8% for the HX system, and 3% for the HZ system.

From the above model studies, we find that the amplitude of the anomaly (or accordingly the percent anomaly) is sensitive to both the thickness and the width of the keel. Thus, it is essentially a function of the area of the cross section of the keel. In contrast, the anomaly width is primarily related to the keel width. It is not sensitive to the keel thickness as long as the shape of the keel does not change. It is also noticed that the HX system response is more sensitive to keel shape than the HZ system response. The curve of HX system response resembles the keel shape and the change in the anomaly width due to the change of the keel width is much larger than that for the HZ system (33% compared to 20%).

It is worthwhile to compare the result calculated with the above technique to that obtained by the finite element method (Lee and Morrison, 1985) that allows a finite conductivity for both the sea ice and the sea water as well as a low operation frequency. This is done in Figure 7 for a triangular sea ice keel. The ice is uniformly 5m thick except for the keel area and it has a conductivity of 0.002 S/m. The conductivity of the sea water is 4 S/m and the HX system flies at 30m above the surface of the ice and operates at 2500 Hz. The dashed line in Figure 7 (upper part) represents the in-phase component of the response, and the solid line corresponds to the system response that would be observed if the sea water conductivity were infinite and sea ice conductivity were zero. Notice the similarity of these two curves and that the percent anomalies and anomaly widths are almost identical for these two curves. Thus it appears appropriate to use the perfect conductor model to interpret low frequency data observed in practice.

3.0 Charts for Data Interpretation

From the above analysis of numerical data, it is seen that the percent anomaly is a function of the cross area of a sea ice keel and the anomaly width is primarily a function of the keel width. To

interpret field data, we need a strategy to relate the observed electromagnetic anomaly to the keel geometry. In terms of the smooth keel used in this analysis (see Figure 2), we need to estimate the two keel parameters A and W from the observed anomaly of the system response. In the following, we are going to concentrate on the HX system configuration which appears superior in terms of sensitivity of the keel geometry. Parallel analysis can be carried out for the HZ system. .

From the computation and analysis of the HX system response for model keels with systematic values of A and W , one can construct an interpretation chart as shown in Figure 8. The vertical axis of the chart is the percent anomaly, while the horizontal axis represents the ratio of the anomaly width q to the flight height h above the flat part of the ice-water interface. Here, the two sets of parametric curves intersect each other clearly and are well separated. The solid lines are for constant values of the drawdown A , and the dashed lines are for constant keel width W . The charted values of a and w are the keel drawdown "A" and width "W" normalized by h , the system height above the flat seawater surface. As shown, the percent anomaly decreases with the decrease of the keel drawdown and keel width. In fact, the line $a = 0.8$ tends to zero fast and intersects the lines $a = 0.4$, $a = 0.2$, etc.. For purpose of clarity, this is not shown in Figure 8. Nonetheless, this does not pose a serious problem in practice since such narrow and sharp keels are highly improbable. Although the interpretation chart is constructed for $h = 30\text{m}$, it can be used for the range $h = 25\text{m} - 50\text{m}$ with an error less than $\pm 4\%$. .

As pointed out at the end of last section, the anomaly shape for low frequency data is almost identical to that obtained in case of the perfect conductor model. This makes the interpretation chart shown in Figure 8 applicable for a wide range of frequencies. We expect that the interpretation chart is useful for the frequency range between 1 - 100 kHz. A similar interpretation chart can be constructed for the HZ system response. It is shown in Figure 9.

Now consider the 2500 Hz theoretical data shown in Figure 7 (dashed line) for an asymmetric triangular keel. From the in-phase component of the AEM response, the percent anomaly and anomaly width are calculated to be 7% and 27m respectively. Thus $q/h = 27/35 = 0.72$. We draw this point in the interpretation chart (point C in Figure 8) and find the corresponding values of a and w to be 0.1 and 0.35 respectively. Hence $A = a \times h = 3.5\text{m}$, and $W = w \times h = 12.3\text{m}$. The estimated keel is drawn (symmetrically about the maximum anomaly of the data) in the lower portion of the illustration (dashed line in Figure 7). We can see that the size of the interpreted keel is close to that of the model keel. But since we have assumed that the keel is symmetric in the interpretation, the position of the maximum of the interpreted keel is misplaced 3.5m to the right.

The field data collected over an ice keel in the Prudhoe Bay by Geotech Ltd. for CRREL (Kovacs et al., 1986) are interpreted next. The AEM system used consists of two pairs of vertical coaxial coils (HX system) and two pairs of horizontal coplanar coils (HZ system). The former operate at 930 Hz and 4158 Hz respectively, and the latter operate at 530 Hz and 16290 Hz respectively. The transmitter-receiver separation of each coil pair is about 6.5m.

A part of the 930 Hz and 16290 Hz data for line F6L3 is shown in Figure 10. The altitude is the distance from the system boom to the ice surface measured by a laser altimeter. Note that the quadrature component of the 16290 Hz data is of bad quality. As we can see in the figure, the data are highly correlated with the altitude as expected. First, we interpreted the data using a 1-D technique (Becker et al, 1987). The result shows an average of 3m thick ice (see Figure 11) but no ice keel is apparent in the interpretation. However, we notice that there is an anomaly in the system response from fiducial numbers 2668 - 2675, which can not be related to the small altitude variation in that area. Moreover, the 1-D interpreted results also show thicker ice in that region. For the 930 Hz data, the anomaly width q is found to be 31.4m and the percent anomaly is 6.5%. Since h is about 38.5m (altitude + average

ice thickness), $q/h = 0.83$. The corresponding point is found in the interpretation chart (point D in Figure 8), which gives $a = 0.08$ and $w = 0.42$. Hence $A = a \times h = 3.08\text{m}$, $W = w \times h = 16.2\text{m}$. The interpreted keel is plotted symmetrically about the maximum anomaly in Figure 11 (dashed line). The solid line in the figure is the average of the drill hole measurements along three parallel lines 11.5 meters apart. As we can see, the interpreted keel is a good approximation of the real feature. Note, however, that the effect of the variation of the altitude of the system boom has been neglected in the interpretation.

The interpretation of 4158 Hz data yields similar result to the above. But difficulty was encountered in interpreting the HZ system data which were acquired at 530 and 16290 Hz. As shown in Figure 10, there is also an anomaly in the 16290 Hz data from fiducial numbers 2668 - 2675. If we calculate the percent anomaly and anomaly width in this case, the corresponding point will fall off the interpretation chart shown in Figure 9 and can not be interpreted.

Generally speaking, the chart interpretation method gives an overall estimate of the size and shape of an ice keel. In order to recover the detail shape of an ice keel, we consider in the next section an iterative technique for data inversion. The chart interpretation can be used to obtain an initial model for the iterative technique where the variation of the system altitude can be easily accounted for.

4.0 inversion of AEM Data

Now consider the determination of sea ice keel shape by data inversion. In order to overcome the problem associated with 1-D data interpretation, we present a two-dimensional (2-D) AEM data inversion scheme where ice thickness may vary along the flight direction of the AEM system. In the perpendicular direction, however, the geometry of the ice/water interface is assumed to be invariant. Although the actual ice bottom topography is three dimensional (3-D), most ice keels do exhibit a dominant extension in

one direction which is called the strike because they are formed by the interaction of two ice plates (Canpolar, 1985). Indeed, as will be shown in the next section, if the strike length of an ice keel of Gaussian shape is greater than three times the AEM system flight height 2-D interpretation may be quite appropriate.

The interpretation scheme that we have developed is based on the principle of the Occam's inversion presented by Constable et al. (1987). Due to practical non-uniqueness of the inverse problem, there may be many sea ice keels that may fit the observed data within a specified error. This scheme yields the smoothest among these. The smooth keel shows the basic features of the actual keel although small bumps on its surface may not appear in the inversion results. Because the process of the data acquisition however constitutes a low-pass filter, it is also possible that the information needed to define the keel in detail is absent from the acquired data.

We first examine the theory of constrained smooth inversion and then use it to interpret synthetic field data which are generated with our numerical modeling program (cf. Appendix). Both noise free and artificially contaminated data are used to test the scheme. Finally we propose a procedure to apply this inversion scheme to low frequency data. In addition to this field data sets (collected in Prudhoe Bay by Geotech, Ltd. in 1985) are also interpreted using this procedure.

4.1 Theory

Let x denote the traverse direction of the AEM system, y the strike direction of the ice keel and z the downward pointing vertical. The interface of the ice and the water is described by $z = h(x)$. Our purpose here is to reconstruct this interface from the airborne electromagnetic data $d(x)$ while the upper surface of the ice is mapped by a laser device. The difference in elevation between the bottom and the top surface of the ice is the required ice thickness. The data $d(x)$ can be either the horizontal or/and the vertical secondary magnetic field recorded during the flight.

Now consider $h(x)$ as the only unknown. This is true when the two assumptions used in our previous modeling algorithm (see the Appendix) are valid. These are: 1) the sea ice is transparent to the electromagnetic wave, and 2) the sea water can be treated as a perfect electric conductor. With these considerations, the AEM data can be written as

$$d(x) = F[h(x)] \quad \dots (1)$$

where $F[]$ is the functional for forward modeling. Equation (1) immediately reveals that the inversion problem is actually one-dimensional (1-D) with the direction of variation along x instead of along z . Fortunately there is a large amount of geophysical literature dealing with 1-D inversion. One particular scheme that well suits our needs is Occam's inversion presented by Constable et al. (1987).

The basis for Occam's inversion is the search for the smoothest solution that fits the observed data within a specified tolerance. This principle applies particularly well to our work because electromagnetic wave propagation in sea water is a diffusion process, so that the resolution of sharp edges on the ice/water interface can not be expected from the data. Furthermore, any information on the sharp edges is contained in the high wave-number range of the data, which is contaminated by noise. Because inversion implies downward continuation (Parker, 1977), the attempts to reconstruct the fine structures of the interface will amplify any noise and result in unstable solutions. Hence a stable solution is necessarily smooth.

In order to find the smoothest solution, let us first define the roughness of the ice/water interface. Physically the rougher the interface is, the larger the magnitude of its derivatives. Thus we define

$$R = \int_{p_1}^{p_2} \left| \frac{dh(x)}{dx} \right|^2 dx$$

to be the roughness of the interface $h(x)$; (p_1, p_2) define the lateral keel extent outside of which the interface is assumed to be flat. In the discrete sense,

$$R = \sum_{i=2}^N (h_i - h_{i-1})^2 = \|\partial \mathbf{h}\|^2 \quad \dots (2)$$

where $h_i = h(x_i)$ and $\|\cdot\|$ denotes the l_2 norm, and

$$\partial = \begin{vmatrix} 0 & 0 & 0 & \dots & 0 \\ -1 & 1 & 0 & \dots & 0 \\ 0 & -1 & 1 & \dots & 0 \\ \vdots & \vdots & \vdots & \ddots & \vdots \\ 0 & 0 & \dots & -1 & 1 \end{vmatrix}$$

$$\mathbf{h} = (h_1, h_2, \dots, h_N)^T$$

Note that $p_1 = x_1$, $p_2 = x_N$ and that the points x_1, x_2, \dots, x_N are usually equally spaced.

Suppose that there are M data points recorded over an ice keel, $d_j = d(X_j)$, $j=1, 2, \dots, M$. The corresponding computed predictions from the discrete model \mathbf{h} are $F_j(\mathbf{h})$. The goodness of fit of the predictions to the actual data can be evaluated using the least-squares criterion

$$E = \sum_{j=1}^M [d_j - F_j(\mathbf{h})]^2 = \|\mathbf{d} - \mathbf{F}(\mathbf{h})\|^2 \quad \dots (3)$$

where

$$\mathbf{d} = (d_1, d_2, \dots, d_M)^T$$

$$\mathbf{F}(\mathbf{h}) = (F_1(\mathbf{h}), F_2(\mathbf{h}), \dots, F_M(\mathbf{h}))^T$$

The ice/water interface \mathbf{h} can only vary within some physical bounds. If we assume that the $z = 0$ plane is chosen such that it coincides with the flat part of that interface then a reasonable lower bound is:

$$h_i \geq 0, \quad i = 1, 2, \dots, N \quad \dots (4)$$

since the ice keel protrudes downward. An upper bound can also be set for each individual case

$$h_i \leq T_i, \quad i = 1, 2, \dots, N \quad \dots (5)$$

In the Arctic, small ice keels may protrude several meters into the water, whereas large keels can protrude tens of meters (Lowry and Wadhams, 1979). The values of T_i should be estimated for each specific ice keel encountered. Note that $h_1 = h_N = 0$ should be included in the constraints in order for the solution to be smooth at the two end points. This condition can be met by simply letting $T_1 = T_N = 0$.

Now the mathematical problem to be solved can be stated as follows: find a solution h which minimizes the roughness R and brings the misfit E within an acceptable tolerance, while the bound constraints (4) and (5) are satisfied. Without the bound conditions (4) and (5), this problem is exactly identical to the one solved by Constable et al. (1987).

The condition for the data misfit is

$$\|d - F(h)\|^2 \leq E_* \quad \dots (6)$$

where E_* is the tolerance. If we treat this inequality as an equality and apply the method of Lagrange multipliers, the above problem can be reduced to the minimization of

$$U = \|\partial h\|^2 + \mu^{-1} (\|d - F(h)\|^2 - E_*) \quad \dots (7)$$

with constraints (4) and (5). Here μ^{-1} is the Lagrange multiplier. As interpreted by Constable et al, μ is a smoothing parameter. The larger the μ is, the less the solution is affected by the misfit. On the contrary, if μ is small, the data misfit is minimized with little influence from the roughness term.

The original problem can now be solved with the following procedures: Solve the above minimization problem for a series of μ

values to obtain a set of solutions of the ice/water interface. Among these solutions, choose the one which satisfies the tolerance condition (6). If more than one solution satisfies (6), choose the one with the largest μ value for this corresponds to the smoothest solution desired.

Such solutions can not however be easily obtained by the direct minimization of the objective function U (equation (7)) since the minimization is non-linear. It is first necessary to transform the non-linear problem into a problem of quadratic programming, for which existing mathematical tools can be used.

Let us linearize the $F(h)$ about an initial model h^0

$$F(h) = F(h^0) + J \Delta \quad \dots (8)$$

Here $h = h^0 + \Delta$, and the Jacobian matrix is

$$J = \begin{pmatrix} \frac{\partial F_1}{\partial h_1} & \frac{\partial F_1}{\partial h_2} & \dots & \frac{\partial F_1}{\partial h_N} \\ \frac{\partial F_2}{\partial h_1} & \frac{\partial F_2}{\partial h_2} & \dots & \frac{\partial F_2}{\partial h_N} \\ \dots & \dots & \dots & \dots \\ \frac{\partial F_M}{\partial h_1} & \frac{\partial F_M}{\partial h_2} & \dots & \frac{\partial F_M}{\partial h_N} \end{pmatrix} \quad \dots (9)$$

Substituting (8) into (7) and dropping the constant term $\mu^{-1}E$, we obtain

$$U = \|\partial h\|^2 + \mu^{-1} \|\tilde{d} - Jh\|^2$$

where $\tilde{d} = d - F(h^0) + Jh^0$ is the modified data. Rearrangement of the above equation gives

$$V = \frac{1}{2} \mu U - \tilde{d}^T \tilde{d} = \frac{1}{2} h^T H h - C^T h \quad \dots (10)$$

Here V is the new objective function and

$$\mathbf{H} = \mu \partial^T \partial + \mathbf{J}^T \mathbf{J}$$

$$\mathbf{C} = \mathbf{J}^T \tilde{\mathbf{d}}$$

Note that \mathbf{H} is a symmetric positive-definite matrix. The minimization of the new objective function V is equivalent to the minimization of U for a fixed value of μ .

Now that the new objective function is in quadratic form the problem of optimization with bound constraints (4) and (5) can be solved using quadratic programming (Gill et al., 1981). There are subroutines available in existing mathematical software libraries which can be used for this purpose. These are E04AF in the NAG Fortran Library and VE04A in the Harwell Fortran Library. We selected VE04A because of its simplicity.

The smoothest solution can now be actually obtained in the following way. Starting from an initial model \mathbf{h}^0 , solve the minimization problem for different μ values. From these solutions choose the one that minimizes the actual misfit E instead of V . (Minimizing V may result in divergence of the solution (Constable et al., 1987)). Use this solution for the next initial model and iterate until the solution for the ice/water interface that brings the misfit below a specified tolerance is found.

The initial model \mathbf{h}^0 can be chosen arbitrarily since it does not affect the convergence of the inversion to the smoothest solution. This is one of the beauties of the smooth inversion scheme which sets out to seek a unique solution. In our problem we set the initial model to be a flat ice cake, i.e., $\mathbf{h}^0 = (0, 0, \dots, 0)^T$.

4.2 Inversion of synthetic data

In this section we will apply the above scheme to invert some synthetic data which are generated with the fast modeling algorithm (see Appendix). To test the stability of the inversion scheme white

noise will also be added to the numerical data. Most inversion schemes require the knowledge of solutions to the forward problem in the computation of the Jacobian matrix. Here these are obtained with the fast forward modeling algorithm. The partial derivatives in the Jacobian are computed using the forward finite-difference of two numerical solutions. Thus one iteration of the inversion process requires $N+1$ forward calculations for M data points.

Before considering the inversion results let us first describe the geometry of the problem for all the synthetic models. Unless otherwise specified, the transmitter and the receiver are both small circular loops separated by 6.5 meters, and both are maintained at 25 meters above the upper ice surface. For convenience the vertical co-axial coil system is referred to as the HX system because the axes of the coils are in the x direction while the horizontal co-planar coil system is referred to as the HZ system. The receiver measures the secondary magnetic field which is expressed in parts per million of the received primary field. The inversion is performed for the HX and the HZ system data independently although the joint inversion can be easily accomplished. Except in the keel area the ice thickness is assumed to be five meters. Note that for the synthetic data generation we assume the induction limit to hold so that the system frequency is not involved. Applications to low frequency data, however, will be shown in the next section.

The position of the end points p_1 and p_2 of the ice keel must now be chosen. These can be arbitrary as long as they are located outside the range of the keel. They should not be too far apart however, because the computation of the Jacobian matrix can be quite expensive and convergence of the inversion may be slow. From our experience, locating these two points at the one quarter of the peak HX system anomaly points, and at the two points at half of the peak HZ system anomaly yields good results. Note that this condition may need to be relaxed for field data because they may be acquired at fluctuating flight heights, which will distort the shape of the observed anomaly.

The sampling interval of $h(x)$, Δx , is taken to be 3 meters for all the examples shown in this report. The data $d(x)$ is sampled at an interval of 3.5 meters unless otherwise specified, which at a helicopter speed of 68 knots corresponds to a recording rate of 10 samples/second.

Figure 12(a) shows the inversion results of the HX system data for a triangular model keel (Model 1), which is symmetric and is located between 60 and 90 meters along the profile. The keel drawdown at the peak is 5 meters and its two sides slope at 18 degrees from the horizontal. The solid line in the upper graph is the synthetic data corresponding to the ice keel shown in solid line in the lower graph, while the dashed line is the system response from the inverted keel shown in dashed line in the lower graph. Similar results for the HZ system data are shown in Figure 12(b). In the inversion the two end points p_1 and p_2 are taken at 45 and 105 meters respectively. The tolerance criterion for convergence is set at 1 ppm for the HX system data and 2 ppm for the HZ system data. Three iterations in the inversion yield the convergent solutions shown in Figure 12 from an initial guess of a flat ice cake that is 5 meters thick. As we can see the interpreted keels have smooth vertices as can be expected for the constrained smooth inversion. The interpreted keels are about 1.5 meters too shallow, and the independent results for the HX and the HZ system data are almost identical.

In the next example we invert the synthetic data for an irregular trapezoidal ice keel (Model 2). The steep side of the keel is 34 degrees from the horizontal while the slope of the other side is 16 degrees (Figure 13). The keel protrudes 6 meters down into the water with a flat bottom of 12 meters. Again both the HX and the HZ system data are inverted independently and the results shown were obtained after three iterations. Misfit for the HX system data is again held below 1ppm, and that for the HZ system data is below 2ppm. For the HX system the inverted maximum keel thickness is correct although the two vertices on the steep side are smoothed out. The

other side however is imaged correctly. For the HZ system, the results are similar except that the maximum keel thickness is a half meter too small.

Now white noise of 5% to 10% of the peak anomaly is added to the Model 1 and Model 2 theoretical data. Here system noise may be small but "geological noise" is estimated to be about 5 percent of the anomaly amplitude. The tolerance criterion for the inversion is set at the RMS noise level and the convergence is usually achieved within 4-5 iterations. The inversion results are shown in Figures 14 - 16. As demonstrated in these figures the inversion still yields good results except at the 10% noise level where the keel interpreted from the HZ system data is flat at the bottom (Figure 15). Our experience with noisy data shows that the inversion results obtained from the HX system data are usually better than those obtained from the HZ system data. In practice 5% noise is probably excessive.

We next examine the case where the height of the system varies during a flight, as is usually the case in practice. The model keel is a symmetric trapezoid in shape with a maximum drawdown of 3 meters (Model 3). The flat bottom is 12 meters wide and the two sides of the keel slope at 12 degrees from the horizontal (Figure 17). The system height over the flat part of the ice/water interface, shown in Figure 17(a), varies between 37 and 41 meters. This height variation was taken from a test flight in Prudhoe Bay. Here it is assumed that the transmitter and the receiver are always at the same level although in practice the instrument pod may tilt in space. The inversion results for the HX and the HZ system response are shown in Figures 17(b) and (c). As can be seen the inverted keels agree very well with the model keel in this case. The final data misfit is 0.3ppm for the HX system and 0.8ppm for the HZ system.

4.3 Application to low frequency data

The inversion procedure can also be extended to the interpretation of low frequency field data. As mentioned earlier, the synthetic data used above were generated in the induction limit where the secondary field only exhibits an in-phase component. In practice however, this can not be realized because the sea water is not a superconductor and the operating frequency of the system must be low enough for the EM wave to penetrate the ice freely. Hence the measured secondary magnetic field has both an in-phase and a quadrature component. Ideally the in-phase and quadrature components can be inverted simultaneously using the above scheme if a fast modeling algorithm for a general conductivity distribution exists. The finite-element and finite-difference methods have been very successful in solving such problems (Lee and Morrison, 1985; Stoyer and Greenfield, 1983). The computational costs associated with these methods however, are prohibitive so that they can not be used in solving this problem as too many forward computations must to be carried out even for one iteration.

Although the fast modeling algorithm that we developed generates theoretical data at the induction limit it can be used to invert low-frequency data. For data collected at 30 kHz the arithmetic sum of the in-phase and quadrature components gives a good approximation of the data that would be obtained at the induction limit (Becker et al, 1983). This sum can be directly interpreted using the previous scheme, but it must first be multiplied by a scale factor prior to interpretation. The same scheme can be used since the anomaly shape changes but little with frequency (Becker and Liu, 1987). The scale factor can be computed as the ratio of the response at the induction limit to the sum of the in-phase and quadrature components at that frequency for a 1-D model. Note that this factor varies with the system altitude. If the altitude does not change much in a flight over the ice keel this factor can be taken as a constant.

Let us now consider an example of synthetic data of the HX system at the frequency of 2500 Hz which was generated by the finite-element method (Lee and Morrison, 1985). The conductivities of the ice and water are 0.002 S/m and 4 S/m respectively. The ice is uniformly 5 meters thick except at the keel area where it protrudes 5 meters into the sea water. The shape of the keel is step-triangular as shown in Figure 18 and the HX system is flown at 30 meters above the ice surface. The scaled sum of the in-phase and quadrature components of the theoretical data and the inversion results are also shown in Figure 18. There are ten unevenly spaced data points and the scale factor is 1.044 in this case. As we can see the inverted ice keel is close to the model although it is somewhat shallower. We consider this an encouraging success of the application of the above proposed procedure.

We have not yet had the opportunity to experiment with large quantities of low frequency data. Thus it is not clear in which frequency range the above procedure can be used to yield reasonable results. Furthermore the role of the water conductivity has not yet been investigated although we suspect that its main effect is to decrease the resolution of the inversion results expected from the superconductor assumption. It is worthwhile however to attempt an inversion of the data collected in the Prudhoe Bay by Geotech Ltd. in 1985 (Kovacs, et al., 1987). The data is for line F6L3 collected with the HZ system at the frequency of 16290 Hz. Because the quadrature component is of very poor quality only the in-phase component was scaled by a factor of 1.104 and interpreted. The system altitude, the scaled in-phase component and the inversion results are all shown in Figure 19. The inversion was performed with the two end points of the ice keel fixed at $p_1 = 54$ meters and $p_2 = 102$ meters. The background ice thickness was fixed at 3.3 meters which is the average from the 1-D interpretation.

The interpreted ice thickness is 1.5 meters too shallow as compared to the drill-hole measurements (solid line in the bottom of Figure 19) in the keel area. However the keel is clearly visible in the

inversion result, which is encouraging since it is much better than the 1-D inversion that shows but little of the keel (Becker et al., 1987). The inversion was stopped at an RMS data misfit of 15 ppm since more iterations could not reduce the misfit. There may be several causes for such a large misfit:

- 1) The effects of the pitch and roll of the system during the flight was not accounted for in this inversion.
- 2) The time constant of the instrument may be too large for such a small keel to be fully represented in the data. The inversion treats the data as being recorded with a zero time constant, which ignores the integral effect due to the instrument. This effect has been carefully studied by Becker and Cheng (1987).
- 3) The top surface of the ice was treated as a flat plane and the effect of the associated ridge was not considered.
- 4) The assumed 2-D structure of the ice may be incorrect.
- 5) For this frequency channel the operational system noise was large. The quadrature component was quite erratic and was not used in the inversion.

Errors due to 1) can be diminished by improving the computer code. Errors due to 2) and 5) can be reduced to a large extent by improving the instrument. But errors due to 4) can not be reduced in the present 2-D inversion scheme. To do this a 3-D inversion algorithm needs to be developed which can be an extension of the present scheme. The process will be computer intensive and furthermore, the data acquisition will necessarily cover an area instead of a line over the keel. This will be very costly and may not be worthwhile.

Errors due to 3) need to be further investigated. Treating the ice top as a flat plane has an effect of pressing the ridge down into the sea water. The ice thickness probably may not be largely

affected. But it remains to be seen whether a topographic correction is necessary before a 2-D inversion is attempted.

Applying the constrained smooth inversion to the scaled sum of the in-phase and quadrature components of the real data appears to be an attractive simple scheme. If the phase of the response does not change, this is identical to scaling the in-phase component only. The performance, mainly the resolution, of this scheme will deteriorate with lowering the operation frequency. The range of the frequency, in which the above procedure can be applied, needs to be defined by further investigation. Effects of the conductivities of the sea ice and water will also reduce the resolution of the keel geometry expected from the assumption of transparent ice and superconducting sea water, which need be studied in the future.

All the computations in the above inversion have been done on a IBM RT PC computer, which has a comparable speed as the Microvax. An typical case of inversion of $M = N = 20$ takes about 4 hours of the CPU time. At the present the computer code is not optimized and it can be shown that its optimization may sharply reduce the CPU time by a factor of 5. It appears that in flight data interpretation will require some highly specialized computing equipment.

In all our examples the ice keel position was assumed known. In practice the position of the ice keel may be obtained from the videocamera record of the associated ridge. If this is not possible the AEM data itself can be used to estimate the position of the ice keel prior to the inversion. To do this the influence of the altitude should first be removed from the data so that the remaining anomaly will indicate the position of the keel.

4.4 Discussion

The 2-D inversion scheme can successfully recover the ice keel information, which is usually lost in 1-D inversion. Because the electromagnetic induction is a diffusion process, sharp edges can not

be resolved from the data and we set out to seek a smooth ice keel, which retains the major features of the true ice keel. Although the scheme is designed to work for data collected in the induction limit with transparent ice and superconducting sea water, it can be applied to low frequency data, where the sum of the in-phase and quadrature components is scaled to approximate the required data. At the present the range of the frequency, in which this procedure can be applied, remains unclear and needs further investigation. With refinements and optimization of the computer program however, there is no doubt that this algorithm can be used routinely to process field data.

5.0 The Three-Dimensional Keel

Under usual circumstances the ice-water interface in arctic seas exhibits the same irregular topography that is normally associated with land forms. In areas covered by young ice however, the topographic features are minimal and a one-dimensional data interpretation technique for the average thickness of the ice sheet from airborne electromagnetic data (Kovacs, et. al. , 1987) can be used. This technique fails in areas of even moderate ice-water interface topography (Kovacs, et al, 1987). In cases where the axial length of the keel is very large (compared to the AEM system altitude) it is possible to remedy this situation with the use of interpretation charts that allow for a two-dimensional description of the keel. If necessary, a 2-D data inversion outlined above may be carried out to delineate ice keels in more details. Because the two-dimensional data interpretation method appears to produce adequately accurate results it appears worthwhile to find the minimal keel axis length that is necessary for its proper application. This is done below where theoretical data for three-dimensional (3-D) keels is presented. In all 14 different keel models are considered.

5.1 The keel model

Cartesian coordinates are used and the XY plane is chosen to coincide with the flat part of the ice-water interface so that the Z axis points vertically down into the sea water. The ice is assumed to be transparent to electromagnetic waves and sea water is assumed to be a perfect electric conductor. Both the co-axial (HX) and co-planar (HZ) coil systems are considered.

The 3-D Gaussian keel model is an extension of the 2-D model described above. Its shape is given by

$$t(x,y) = A e^{-\frac{y^2}{0.361 s^2} - \frac{x^2}{0.361 w^2}} \quad (1)$$

where

$t(x,y)$ is the protruding depth of the ice keel into the sea water,
 A is the keel drawdown or maximum keel thickness,
 w is the keel width at $t = A/2$ on the cross section $y = 0$,
 s is the keel width at $t = A/2$ on the cross section $x = 0$.

Here we define the y direction as the strike direction of the ice keel and hence designate s as the strike length of the keel. An infinite value of s corresponds to a 2-D ice keel. When $s = w$, the keel is a body of revolution. The numerical calculations were carried out as outlined in the Appendix.

5.2 Airborne electromagnetic profiles

In this section, we examine AEM profiles over three sets of model ice keels. Except in the keel area, the sea ice is assumed to have a uniform thickness of 5 meters. The keel parameters A , w , and s are varied to demonstrate the corresponding changes in the AEM system response. For each model set, the keel drawdown A and keel width w are fixed, while the keel strike length is varied. Model Set 1 represents a shallow keel, where $A = 3\text{m}$, $w = 24\text{m}$, and keel

strike lengths of 12, 24, 48, 96, and inf. meters respectively (Figure 20). The corresponding theoretical AEM response for the HX and HZ system over the center of these features along a line perpendicular to the keel strike is shown in Figure 21, where curves 1, 2, 3, 4, and 5 correspond to $s = 12, 24, 48, 96,$ and inf. meters respectively. Both the transmitter and the receiver, which are separated by 6.5m, are "flown" 25m above the upper ice surface. The system response is plotted at a point located mid-way between the transmitter and receiver, and is expressed in parts per million of the primary field at the receiver. Figure 21 (c) shows the the cross section of the ice directly below the flight line.

It is evident in Figure 21(a) that as the strike length s decreases the HX system anomaly becomes smaller but the general shape of the anomaly curve changes little. If these system anomalies for 3-D ice keels are interpreted using the 2-D chart given in our previous report, the interpreted cross section of the keel will be smaller than that in the 3-D model. The significant effect here is a reduction of the interpreted keel drawdown. The keel width however is less affected since the anomaly width does not contract significantly with the keel strike.

As shown in Figure 21(a), the HX system response for a keel strike length s of 96m (curve 4) closely resembles the response for an infinitely long keel (curve 5). In fact the maximum difference in the relative amplitude or percent anomaly (c.f. Becker and Liu) is of the order of 1.5% and would result in an under-estimation of the keel depth of about 10%. In fact the relative error in the percent anomaly size is also of this order and suggests that a 10% value for this quantity be used as a threshold value for the definition of a two-dimensional keel. Thus, If the maximum relative difference in the relative anomaly is less than 10%, the finite strike keel can be effectively considered as two-dimensional, and hence 2-D interpretation techniques can be applied.

Using this criterion, any ice keel whose strike length exceeds 96m can be effectively considered as two-dimensional for the purpose of data interpretation. In general, the ratio of the strike length to the system height over the ice-water interface must be larger than 96/30. This value is much greater than the value expected from the footprint for the HX system (see Becker, et al., 1987), which is 1.4. At the first sight, this may seem puzzling. A careful look at the current distribution on the water surface, however, will yield the explanation.

The current pattern for a horizontal-axis transmitter is shown in Figure 22 (from Becker, et al., 1987). It consists of two circular current rings while most of the current flows in a direction parallel to the strike of the ice keel. For a keel of finite-strike, this major current flow is distorted. The footprint however was computed on the basis of the unaltered current flow so that it is not surprising to find that the strike/height ratio needed for 2-D interpretation is much larger than the footprint of the HX system.

Now let us look at the HZ system response and see how it changes as the keel strike length is varied. Curves 1, 2, 3, 4, and 5 in Figure 21(b) correspond to $s = 12, 24, 48, 96,$ and inf. meters respectively. As expected, the HZ system response also decreases when the strike length of the keel is reduced. The major change is the decrease of the anomaly amplitude. The anomaly width, in contrast, stays almost constant. Nonetheless, as the keel strike is further decreased, the anomaly shape also changes and the lower part of the anomaly flattens out.

The HZ system response for $s = 96\text{m}$ is also close to that for the 2-D keel. The maximum relative difference in the relative anomaly in this case is about 12%. Thus taking 10% as the threshold, the keel strike should be somewhat longer than 96m in order to use the 2-D interpretation. Roughly speaking, we may again take the strike/height ratio to be approximately $96/30 = 3.2$ for the 2-D interpretation to be valid. This is the same as that for the HX system

and is very close to the value obtained from the footprint of HZ system which is 3.7 . As a matter of fact, the current pattern on a flat water surface for a vertical-axis transmitter consists of concentric circles so that a 3-D ice keel will distort this pattern as much as a 2-D ice keel does.

Now let us increase the keel drawdown A to 6m and keep the keel width at $w = 24$ m. Figure 23 shows the deeper ice keels with different strike lengths (Model Set 2). The corresponding HX and HZ system responses are shown in Figure 24 (a) and (b) respectively. Now both the HX and HZ system responses are much larger than those for Model Set 1 due to the larger cross section of these keels. But the characteristics of these anomalies are similar to those discussed for Model Set 1. Therefore, the analysis for Model Set 1 still holds in this case.

Figure 25 shows the narrower keels that make up Model Set 3, where $A = 3$ m and $w = 12$ m. The keel strike, again, is 12, 24, 48, 96, and inf. meters respectively. Figure 26 (a) and (b) displays the HX and HZ system response and curves 1, 2, 3, 4, and 5 correspond to $s = 12, 24, 48, 96,$ and inf. meters. Figure 26 (c) shows the cross section of the keel below the flight line.

Now the HX system response is much smaller than those shown in Figure 21 (a) because of the smaller size of the ice keels. But the relative characteristics remain unaltered and the analysis for Model Set 1 still applies here.

The HZ system anomaly also reduces with the decrease of the strike length of the keel. Furthermore, it is observed that a minor feature begins to appear at the center of the anomaly when the keel strike is small. This feature, however, may not be real as it may be caused by numerical errors in the calculation. At present, we are unable to make this point clear due to computational limitations.

5.3 Discussion

- 1) The strike length of an ice keel must be at least three times the flight height of the AEM system in order to successfully use the 2-D techniques to interpret both HX and HZ systems response.
- 2) For the HX system, the anomaly amplitude decreases as the keel strike shortens, but the anomaly width changes little.
- 3) For the HZ system, the anomaly amplitude also decreases as the keel strike contracts. At the same time, the lower part of the anomaly tends to become flat.

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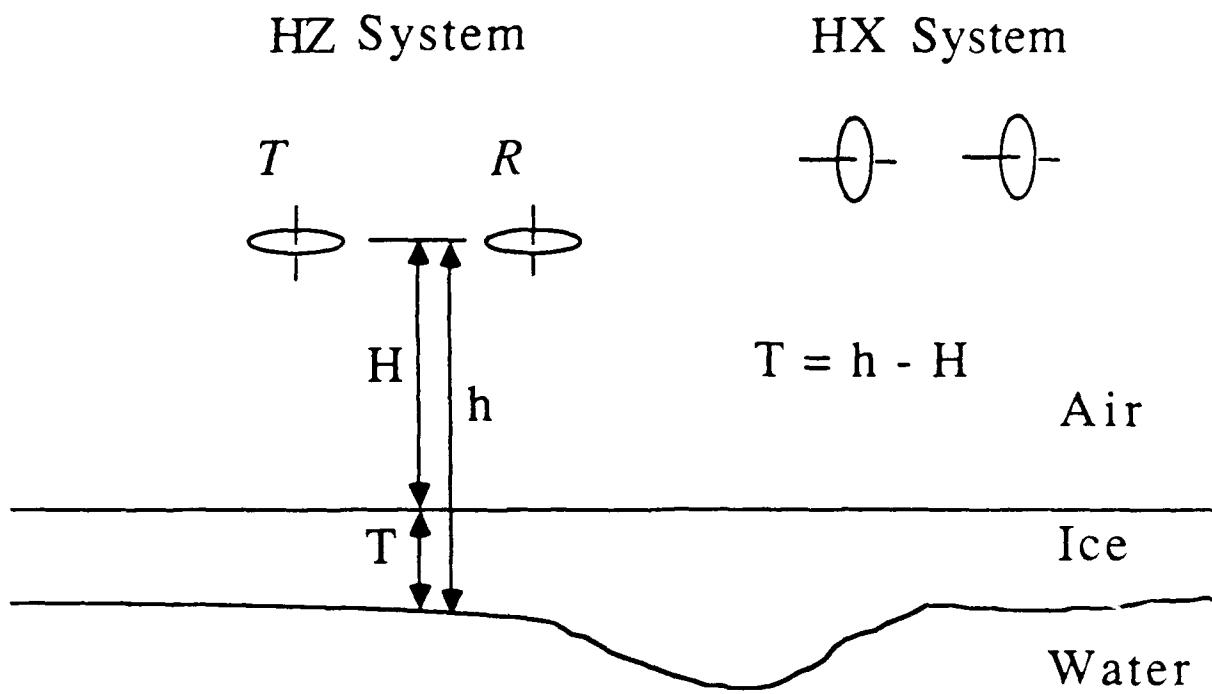


Fig.1 HX and HZ systems over sea ice

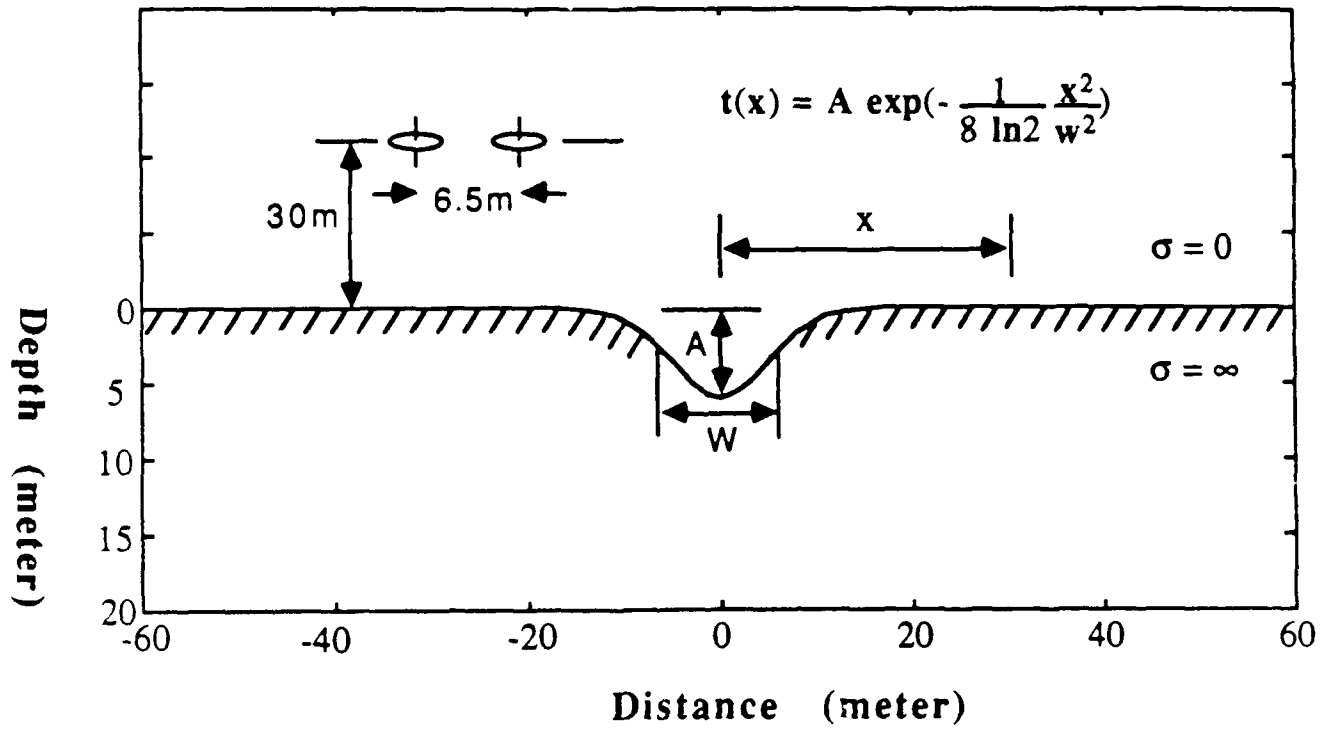


Fig.2 Cross section of a smooth keel

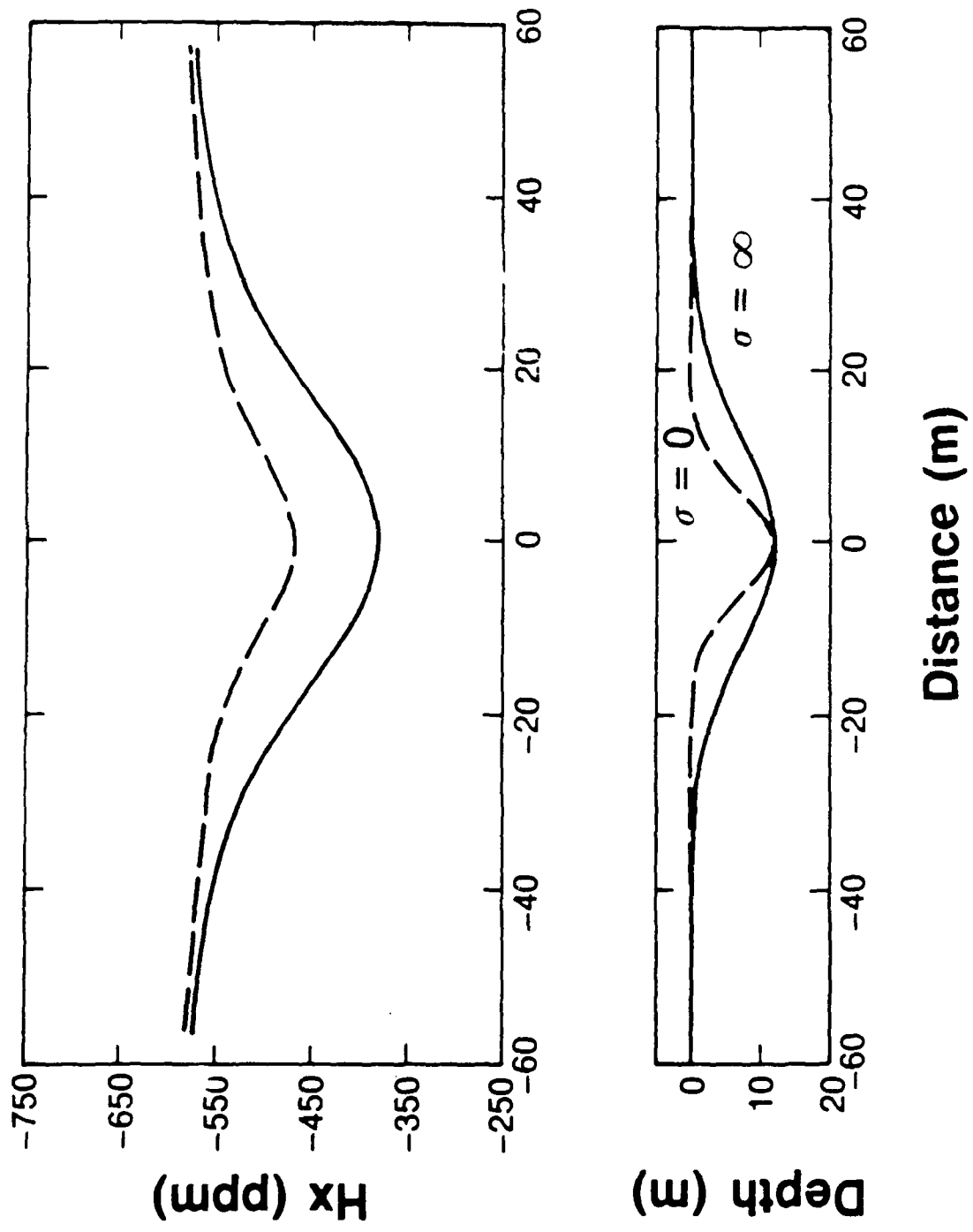
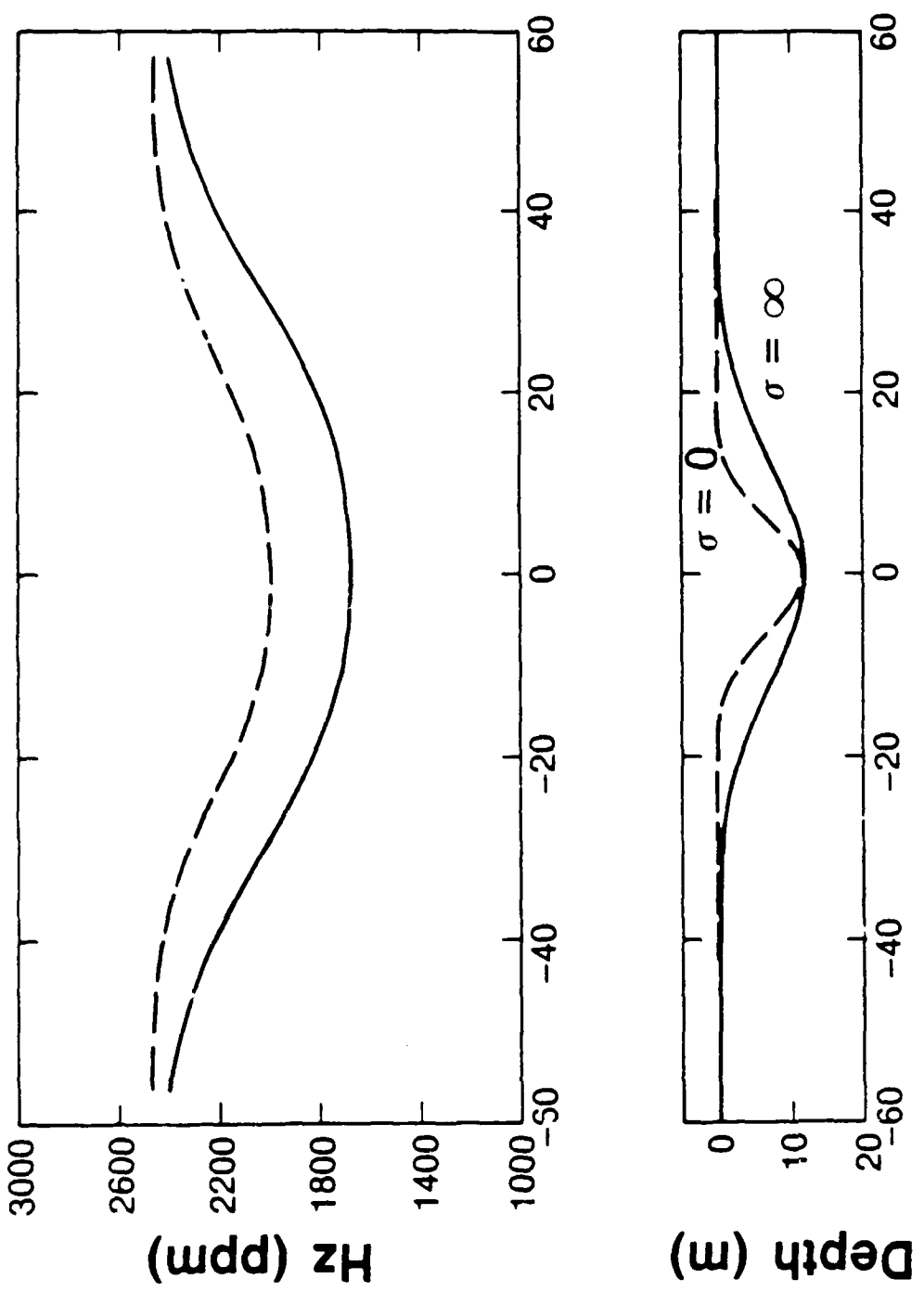
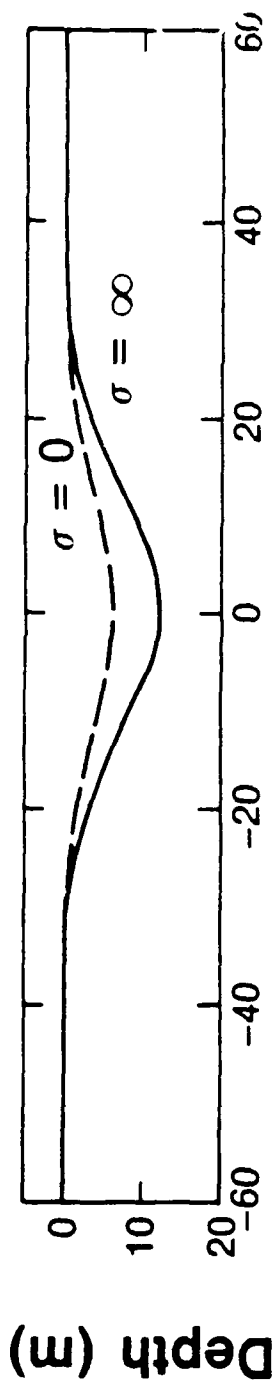
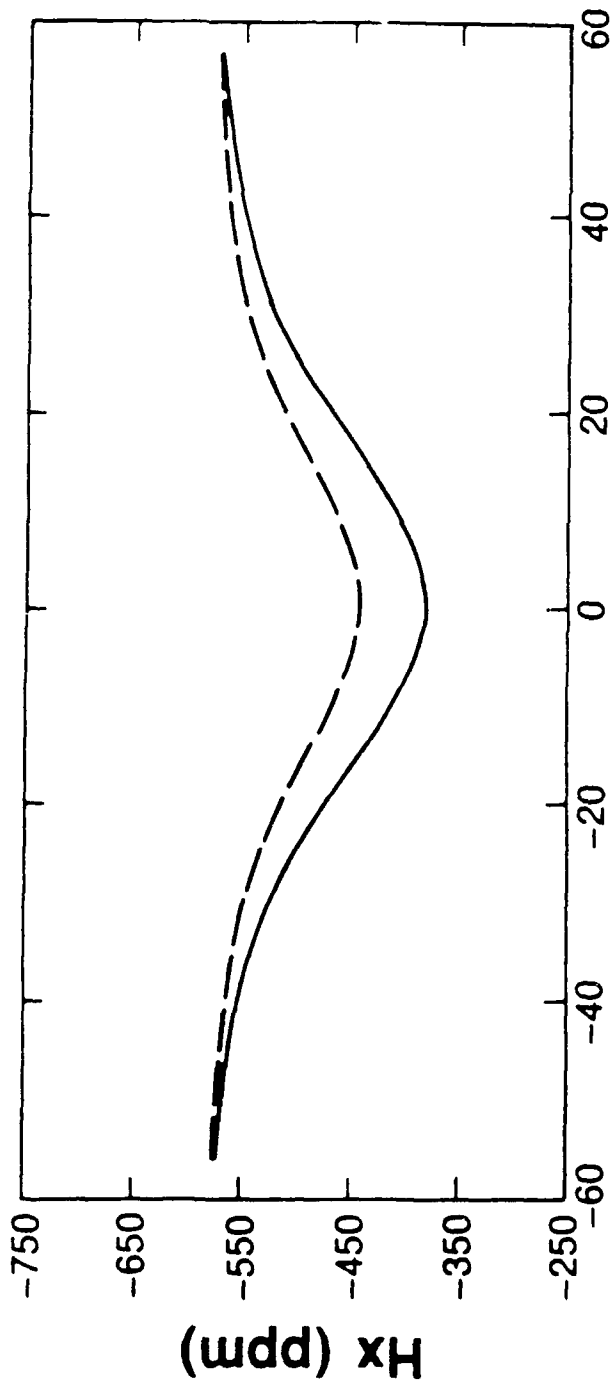


Fig.3 Effect of Keel Width on IIX System Reponse



Distance (m)

Fig.4 Effect of Keel Width on HZ System Response



Distance (m)

Fig.5 Effect of Keel Drawdown on HX System Response

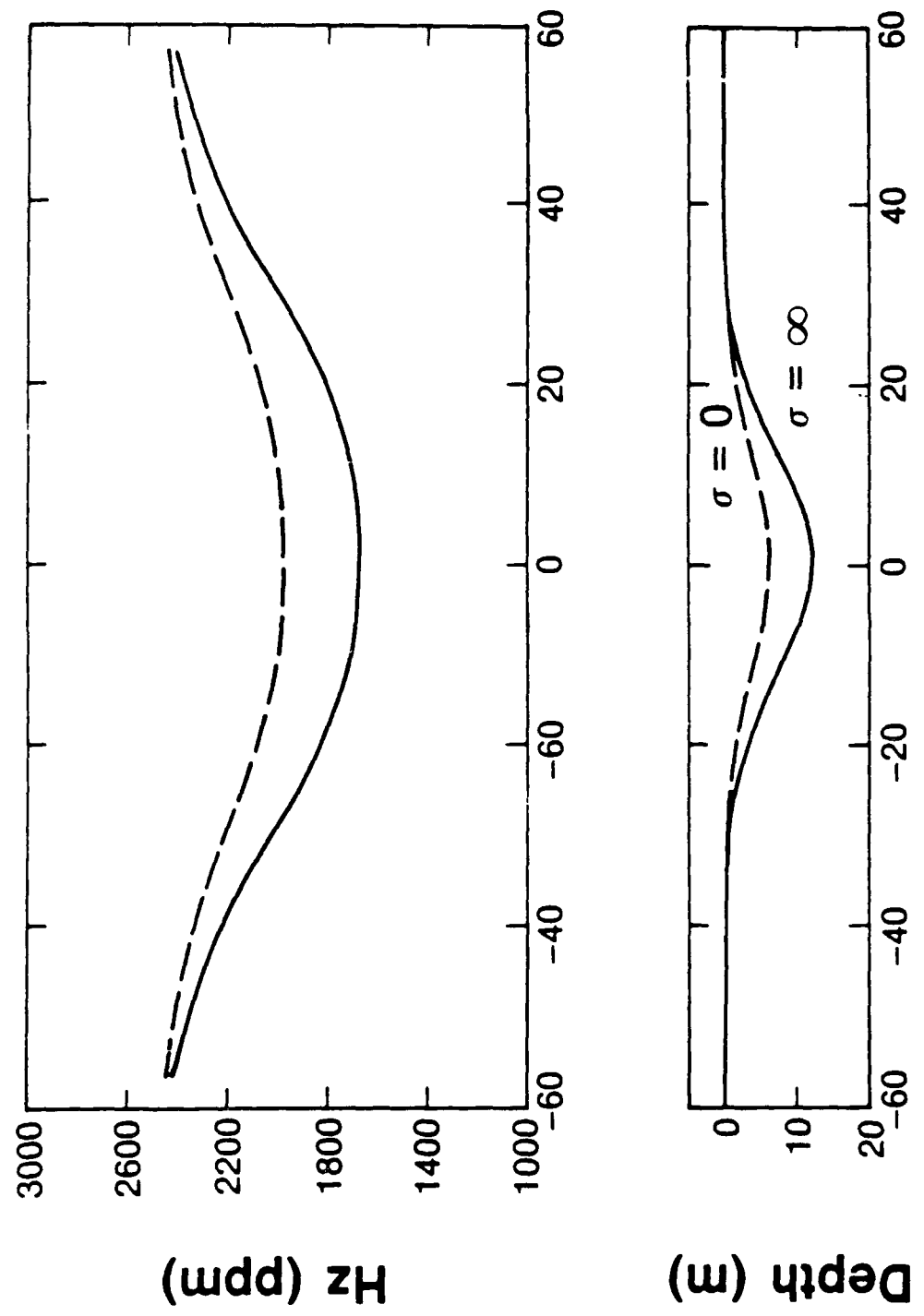
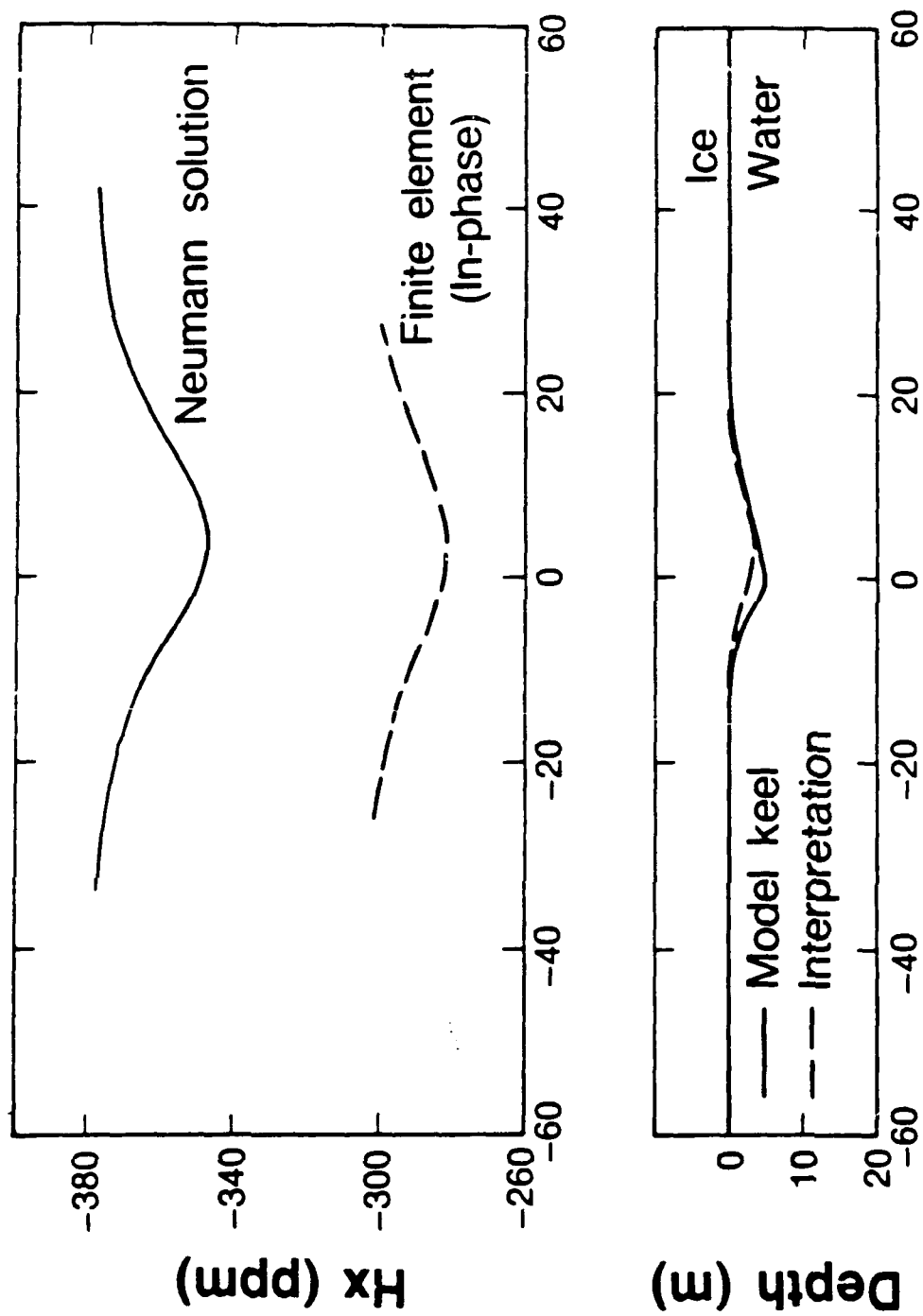


Fig.6 Effect of Keel Drawdown on HZ System Response



Distance (m)

Fig.7 Model Response

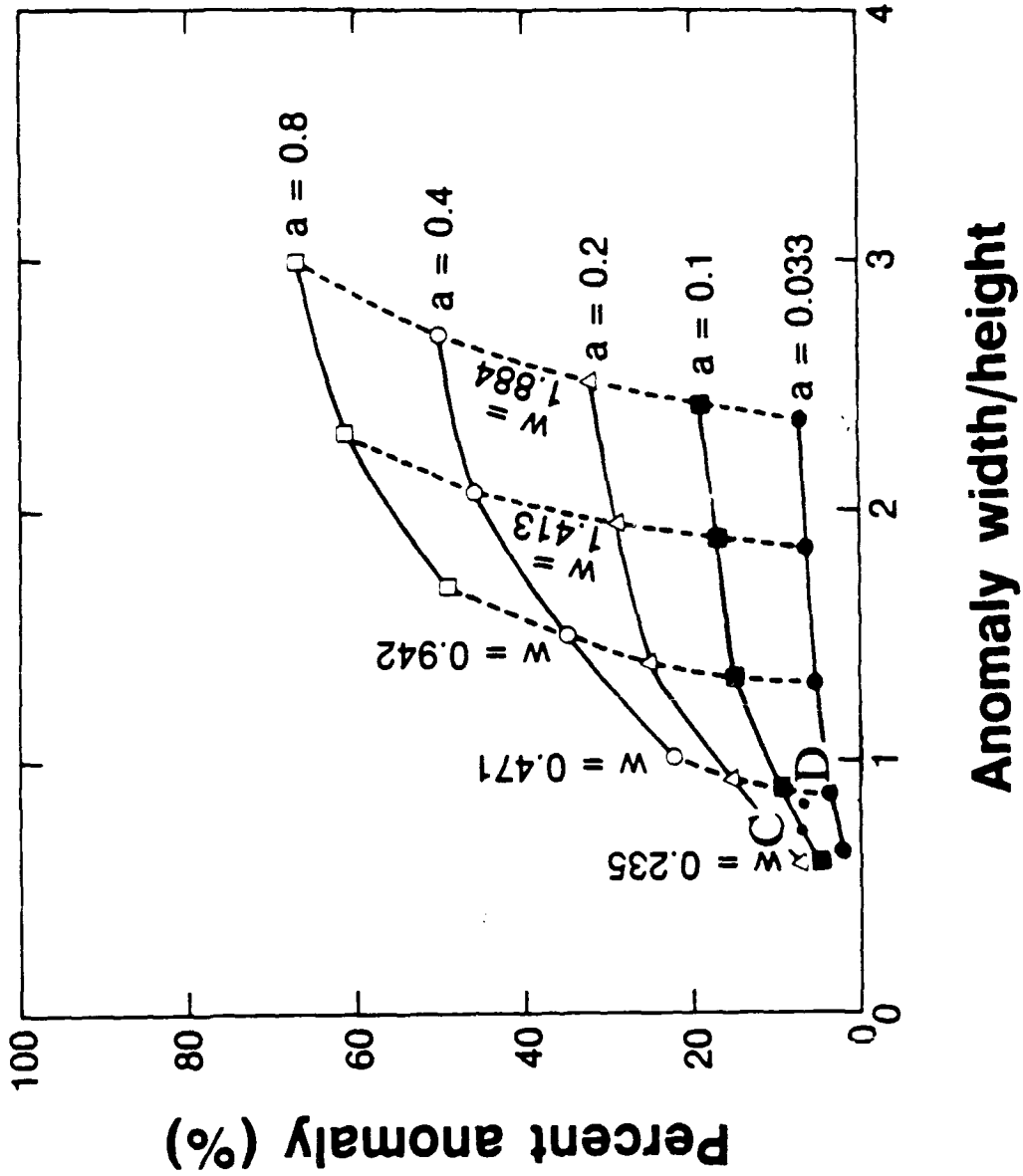
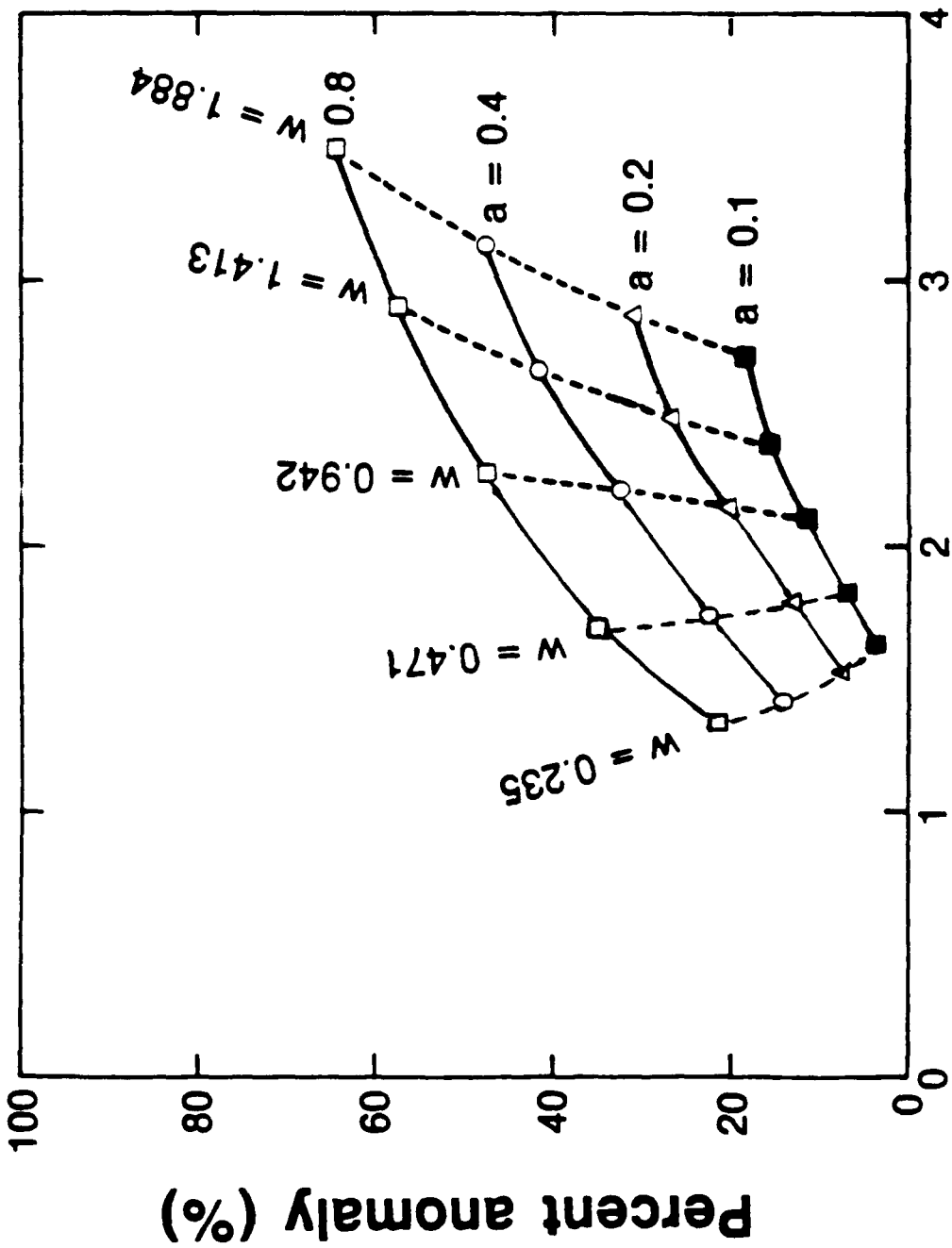


Fig.8 Interpretation Chart for HX System

Anomaly width/height

Percent anomaly (%)



Anomaly width/height

Fig.9 Interpretation Chart for HZ system

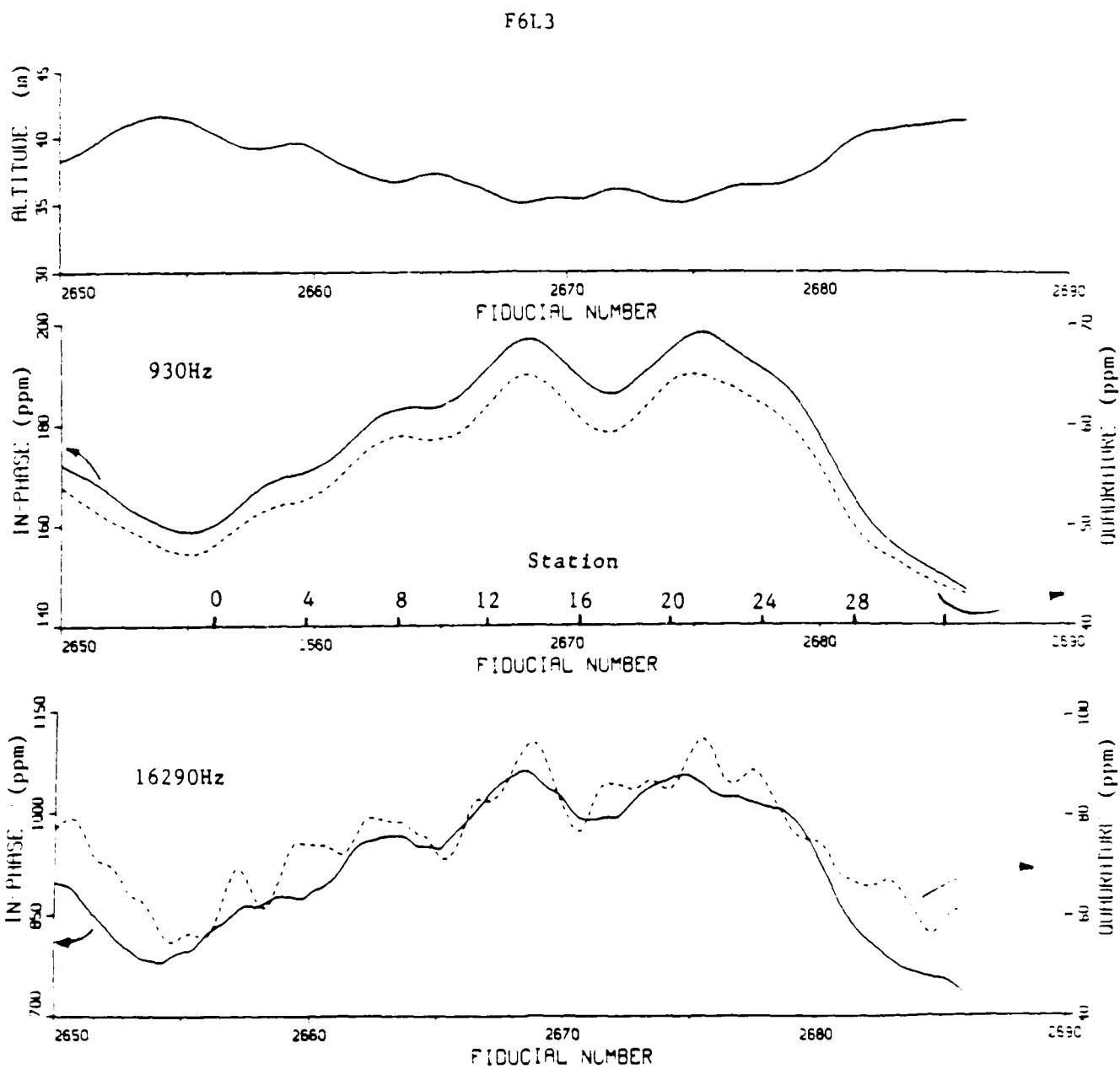


Fig.10 Field data. Line F6L3.

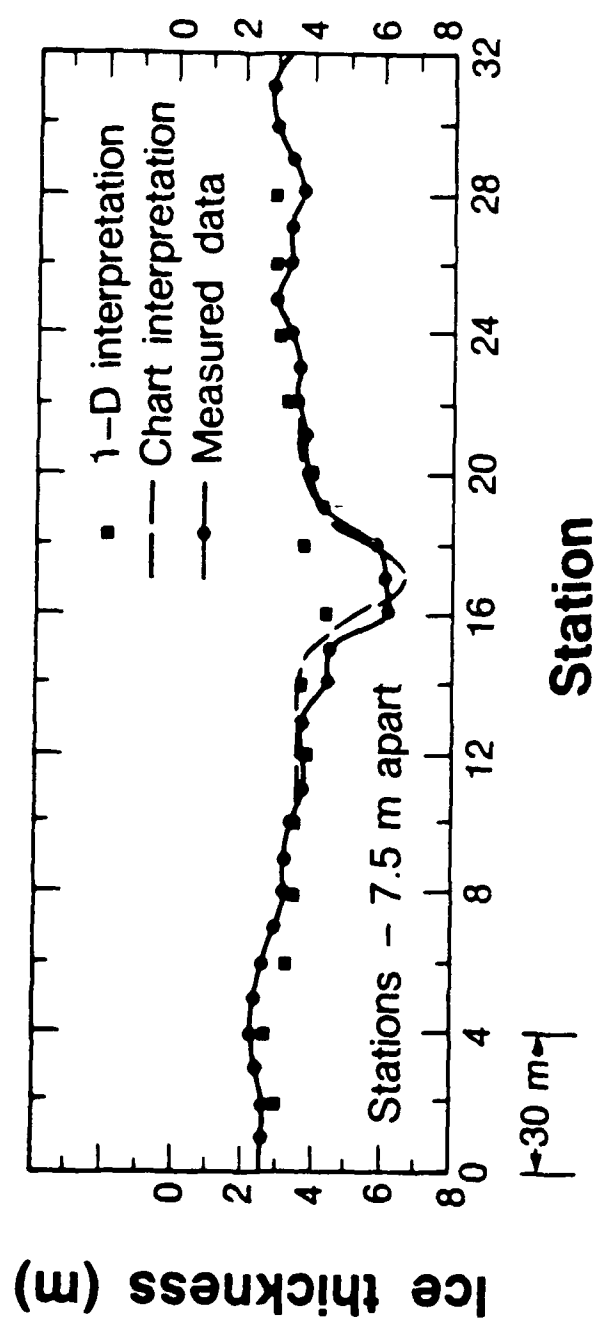


Fig.11 Interpretation of Field Data

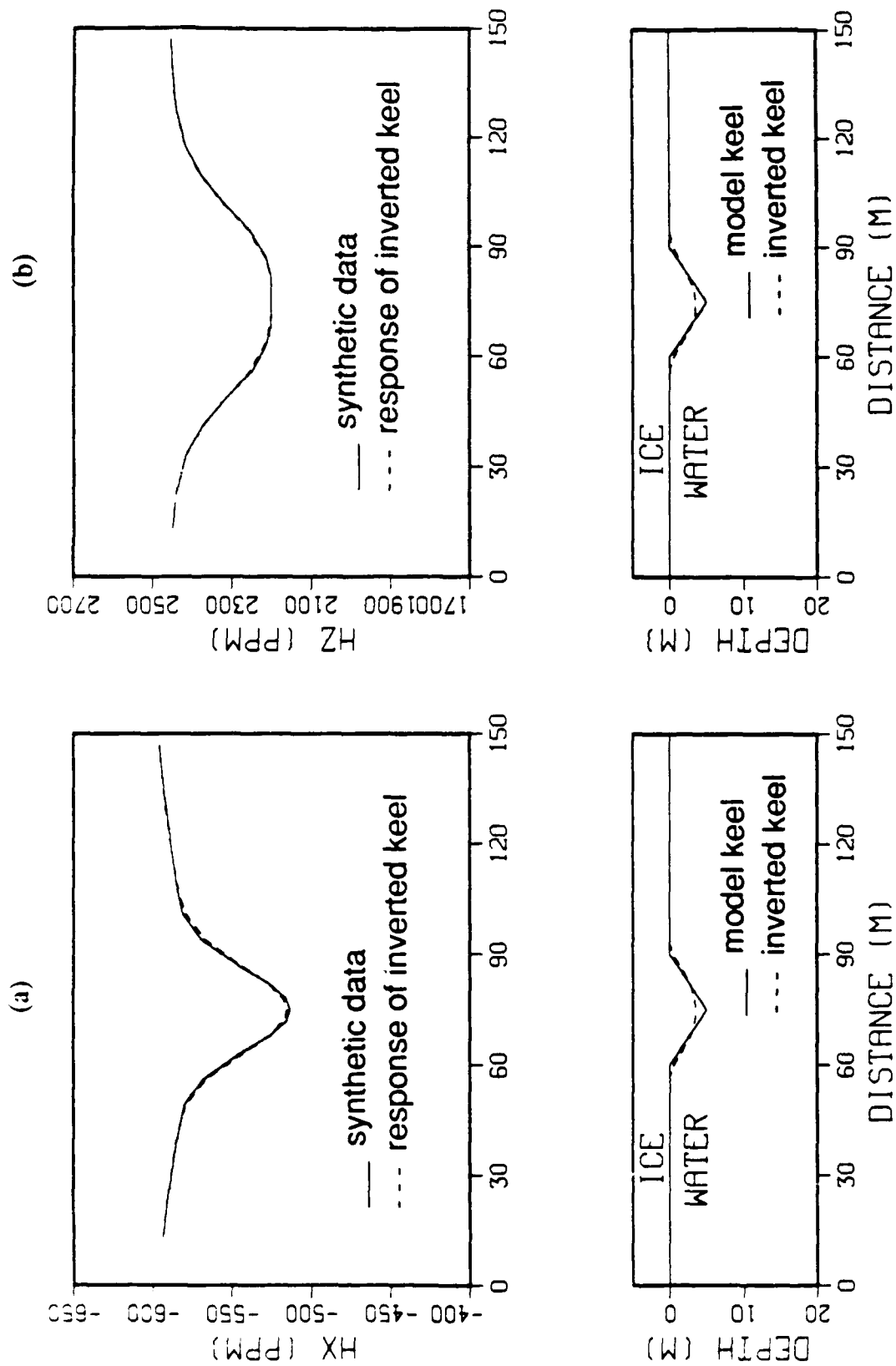


Figure 12 (a) Synthetic data for the HX system and inversion results, Model 1.

(b) Synthetic data for the HZ system and inversion results, Model 1.

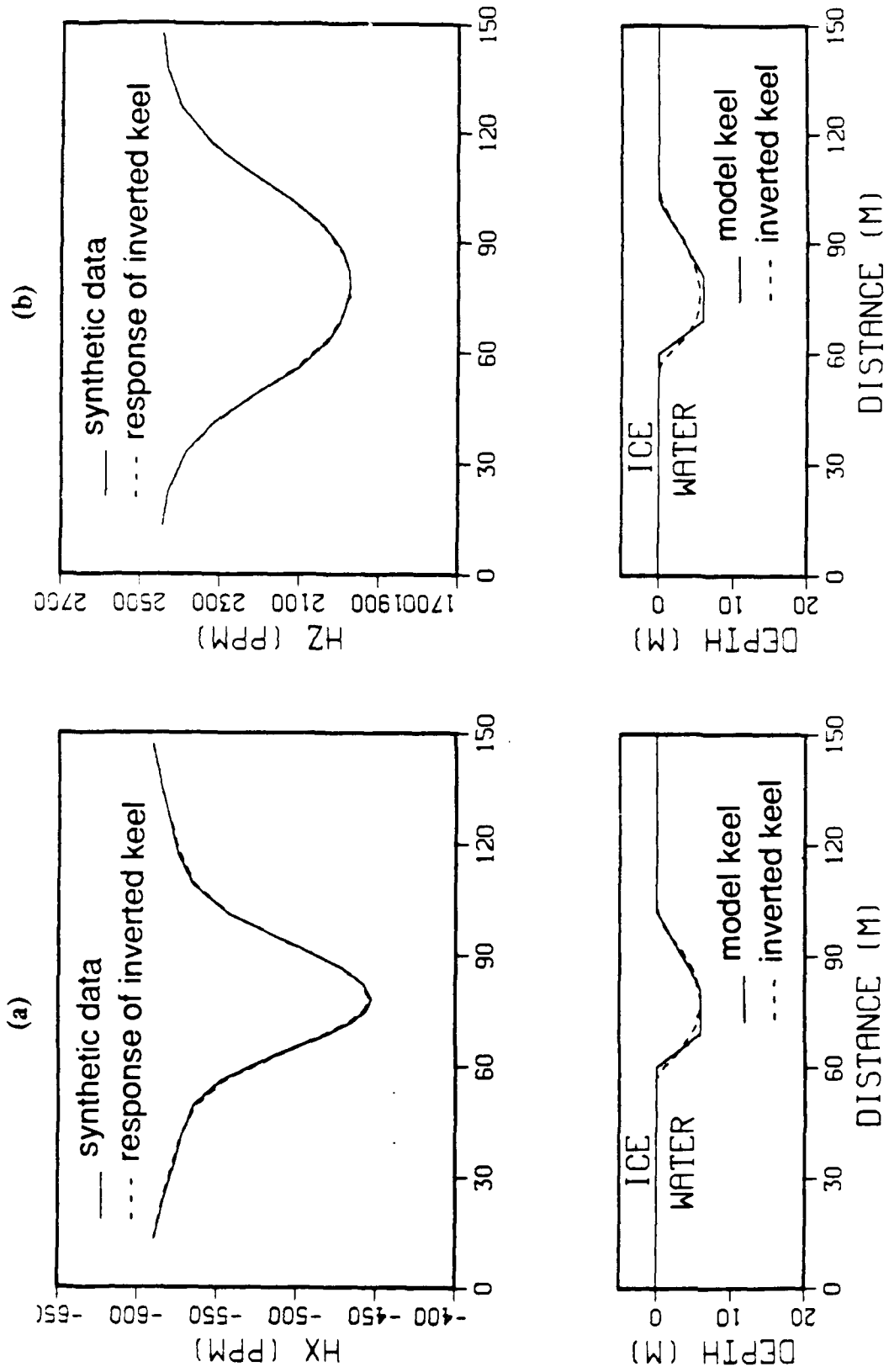


Figure 13 (a) Synthetic data for the IX system and inversion results, Model 2.
 (b) Synthetic data for the IIZ system and inversion results, Model 2.

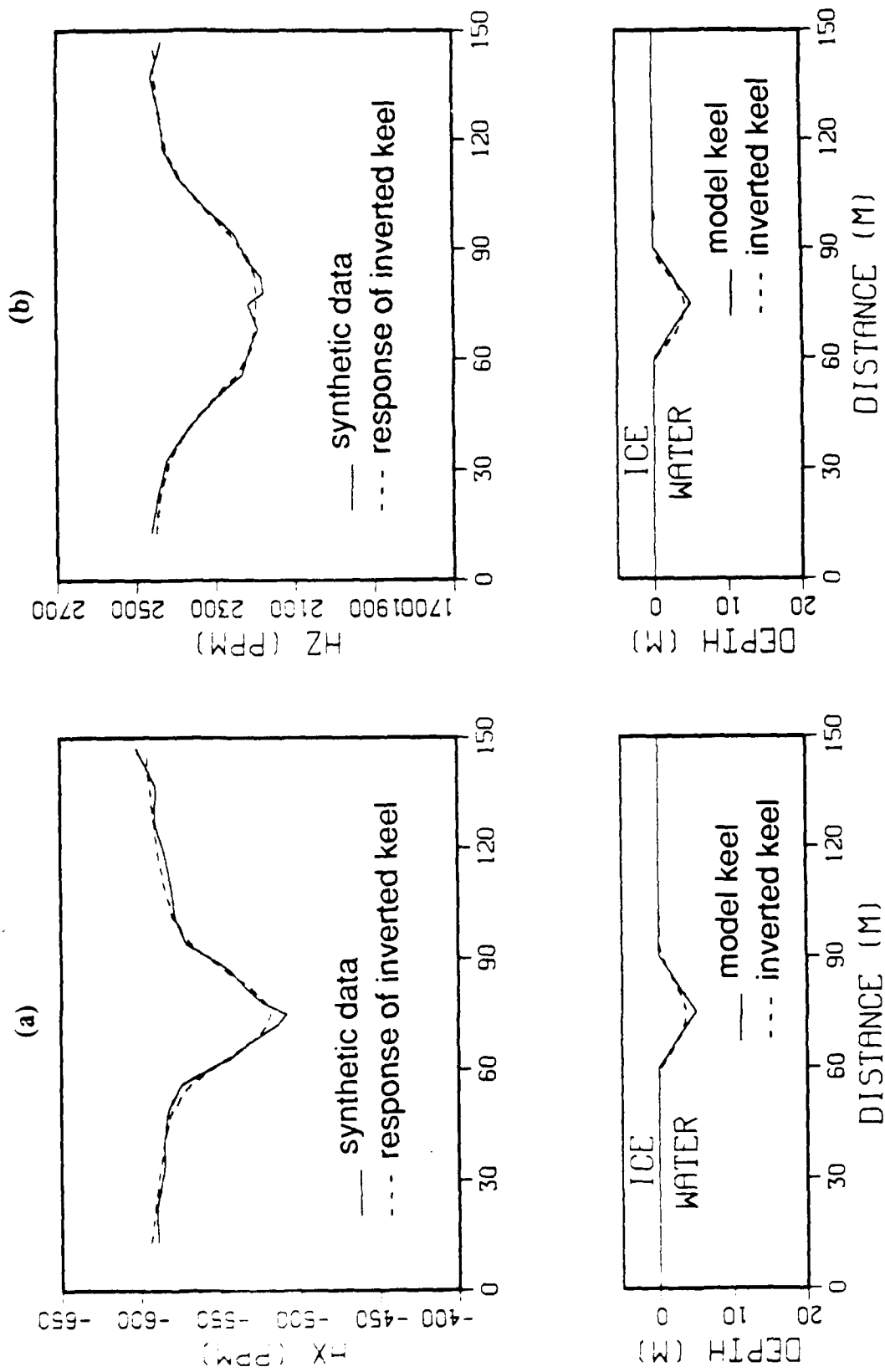


Figure 14 (a) Synthetic data with 5% added noise and inversion results for the HX system, Model 1. (b) Synthetic data with 5% added noise and inversion results for the HZ system, Model 1.

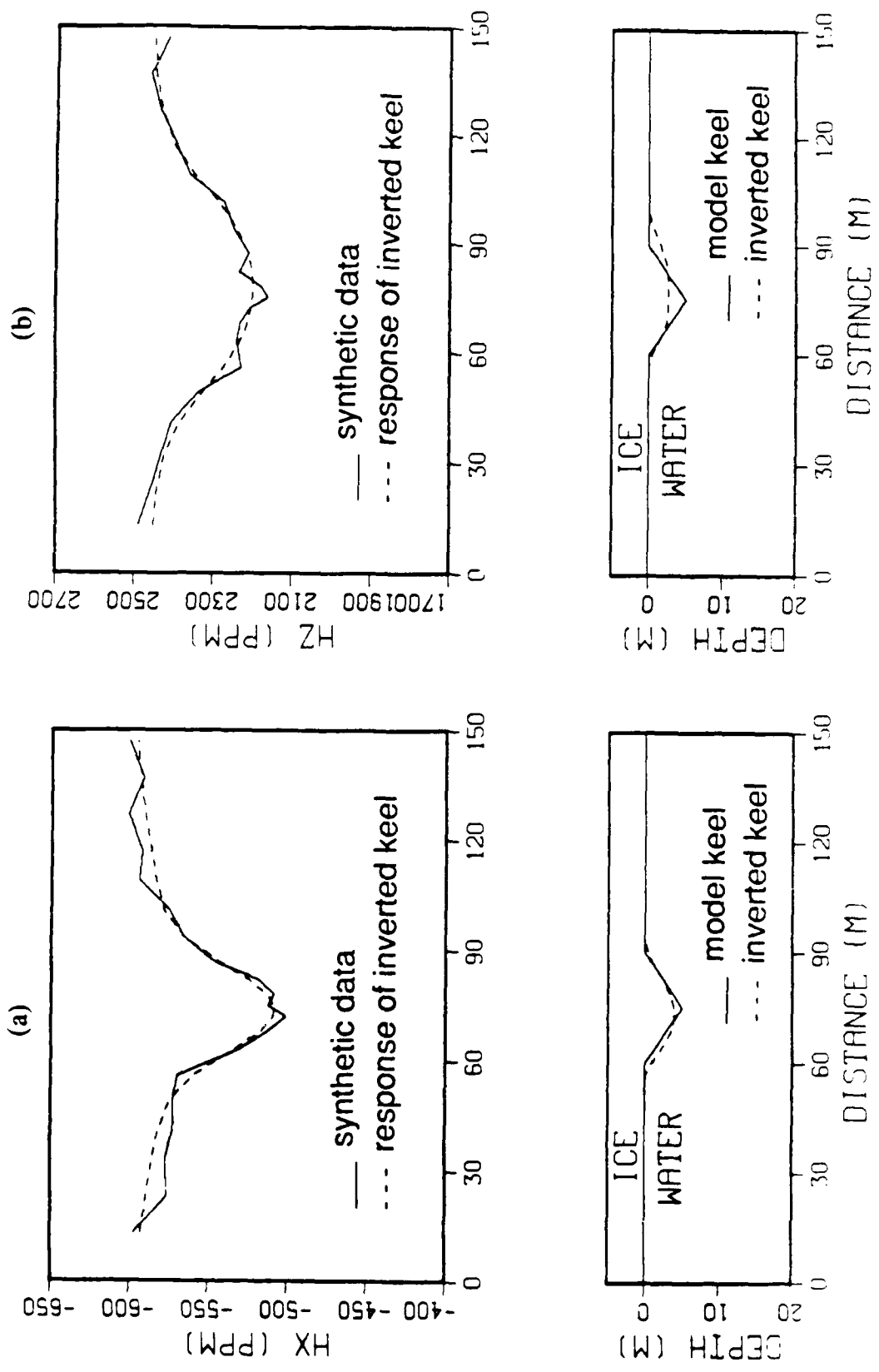


Figure 15 (a) Synthetic data with 10% added noise and inversion results for the HX system, Model 1. (b) Synthetic data with 10% added noise and inversion results for the HZ system, Model 1.

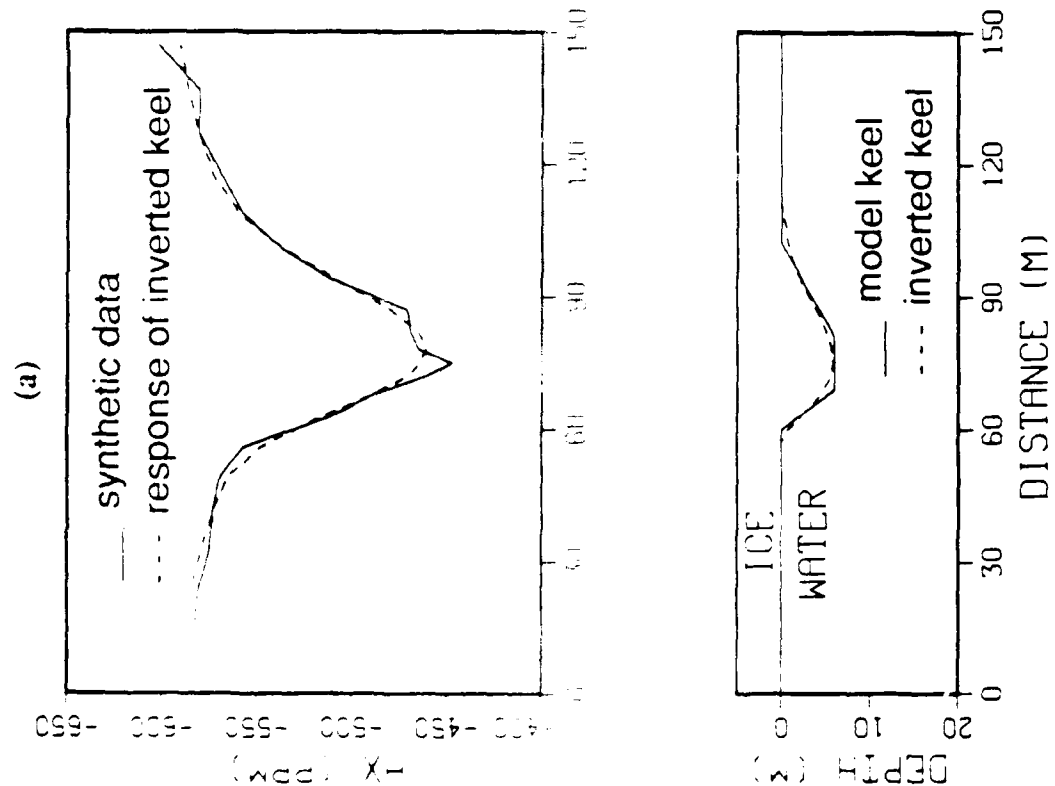
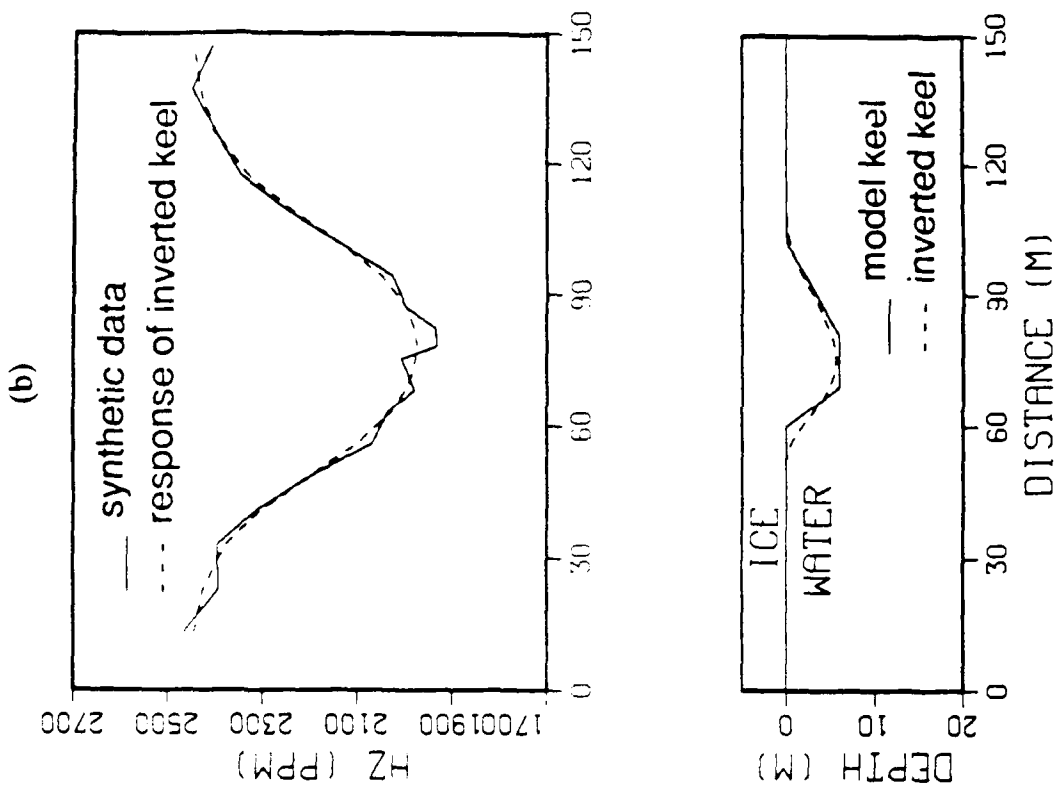


Figure 16 (a) Synthetic data with 5% added noise and inversion results for the HX system, Model 2. (b) Synthetic data with 5% added noise and inversion results for the IIZ system, Model 2.

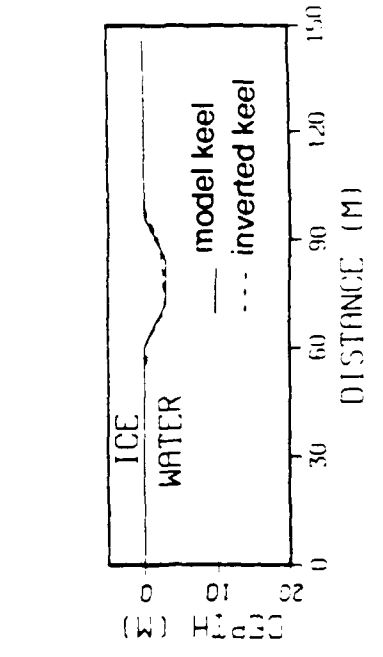
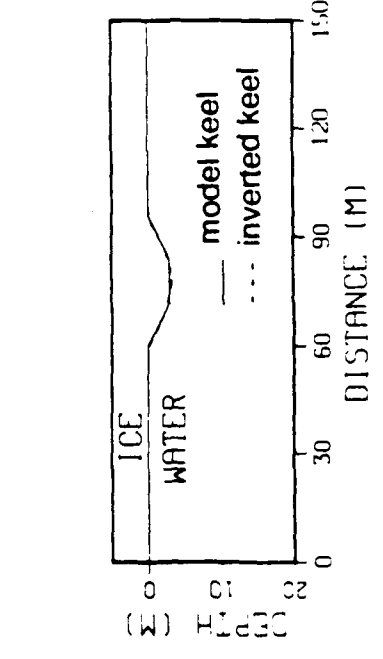
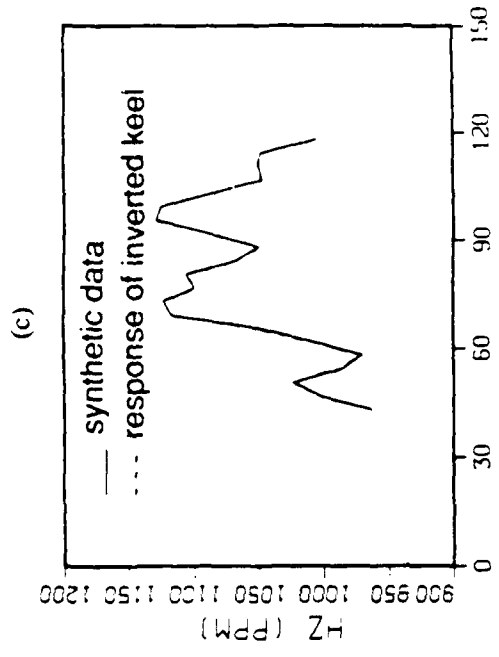
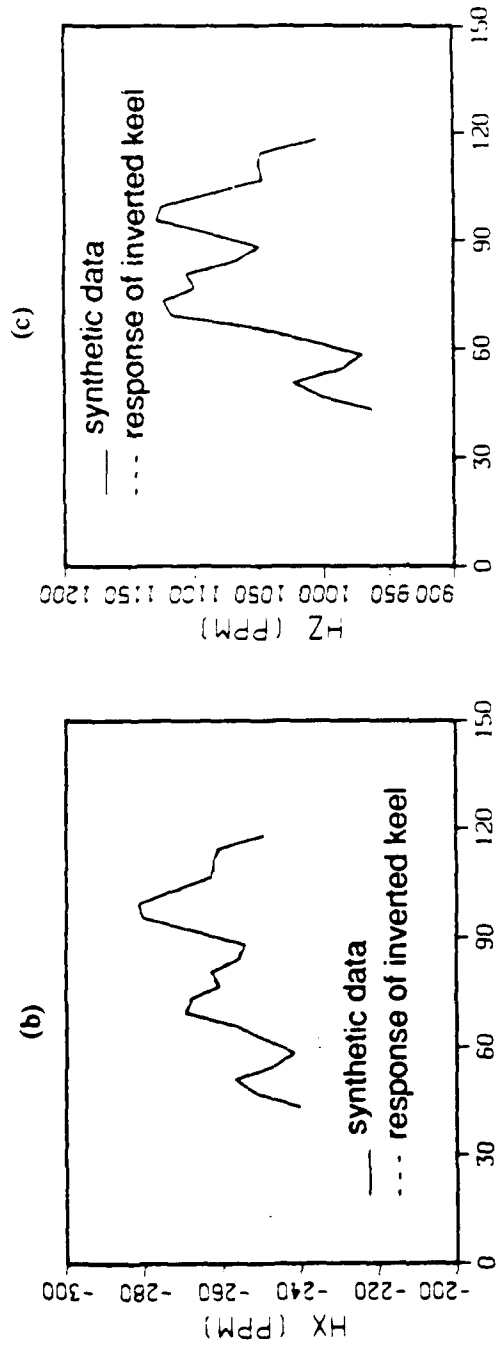
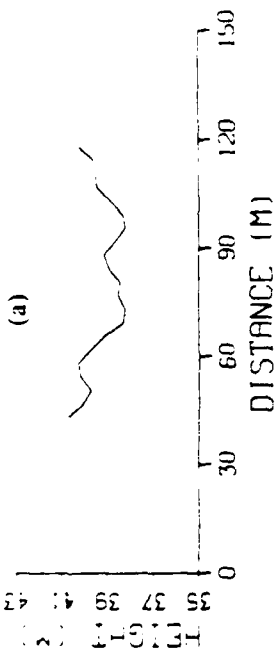


Figure 17 (a) Height of the system over the ice/water interface during the flight. (b) Synthetic data and inversion results for the IIX system, Model 3. (c) Synthetic data and inversion results for the IIZ system, Model 3.

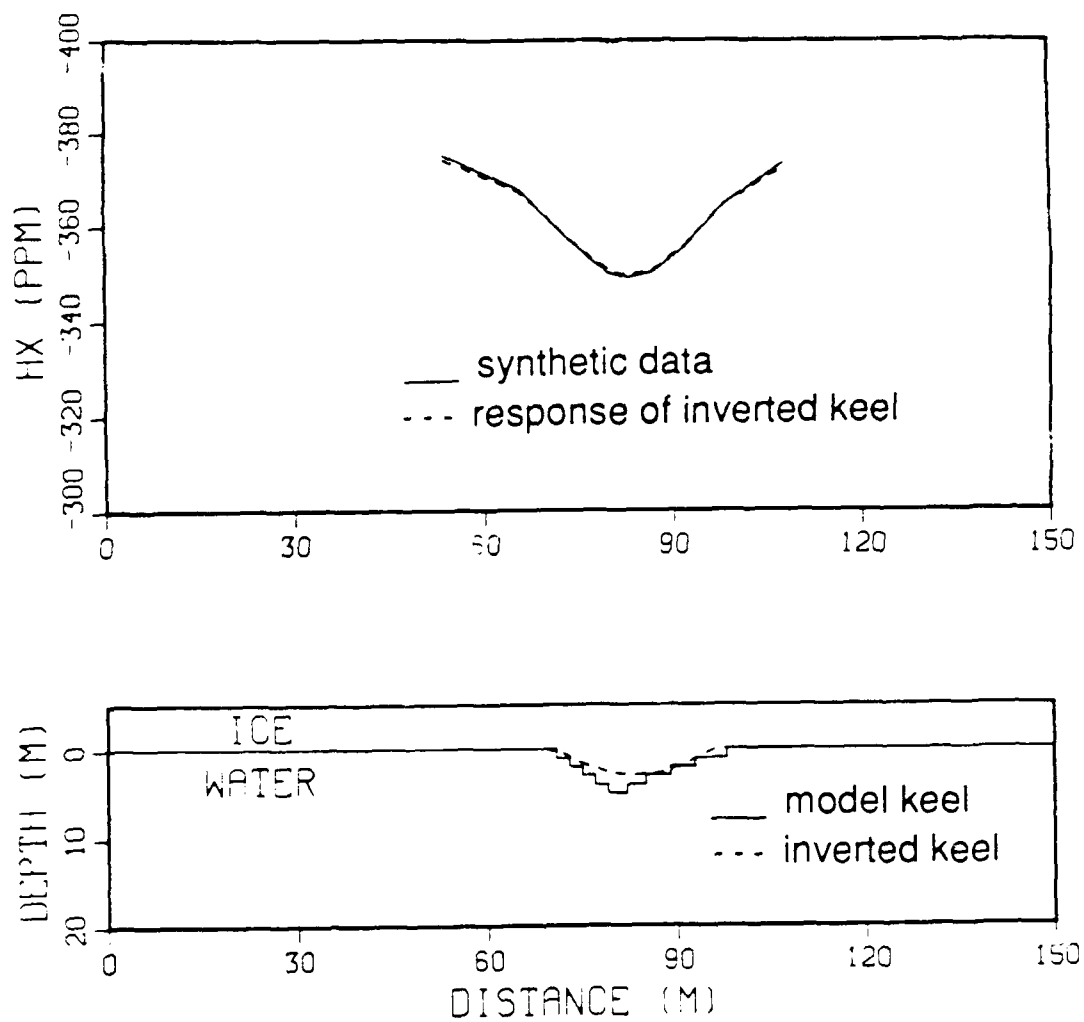


Figure 18 Synthetic data at 2500Hz and inversion results for the HX system, Model 4. The data shown are the scaled sum of the in-phase and quadrature components of the secondary magnetic field.

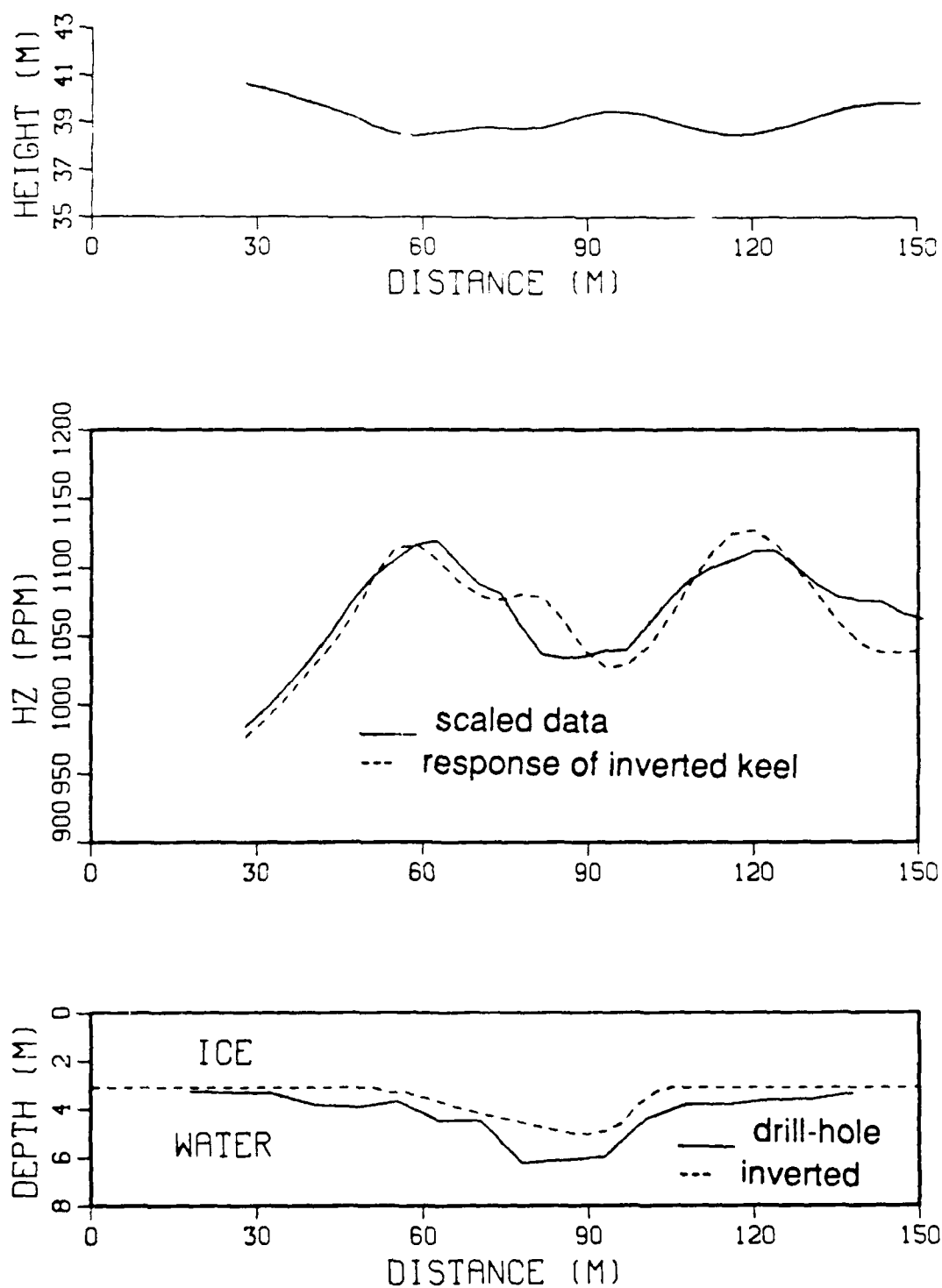
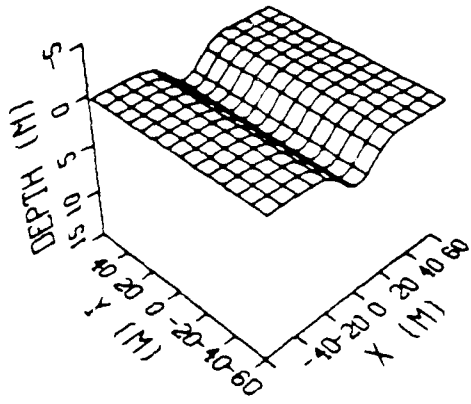
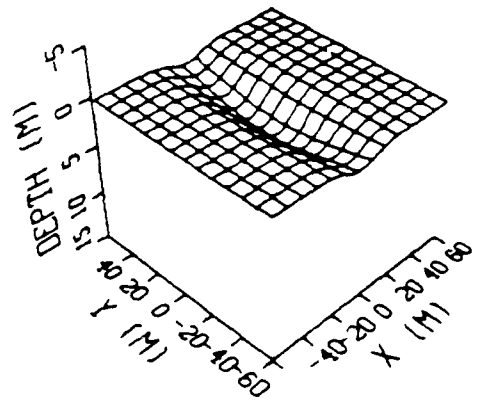


Figure 19 Scaled in-phase data at 16,290Hz and the inversion results for the HZ system. The data were collected at the Prudhoe Bay by Geotech Ltd. in 1985 on line F6L3.

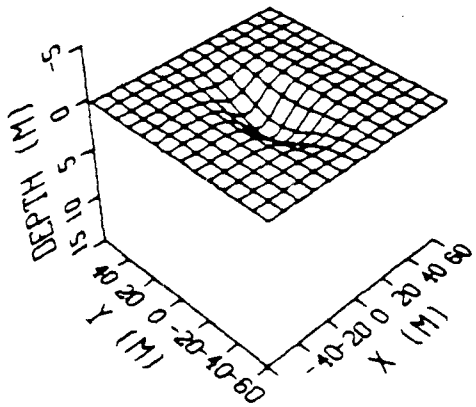
(a) $s = \text{inf.}$



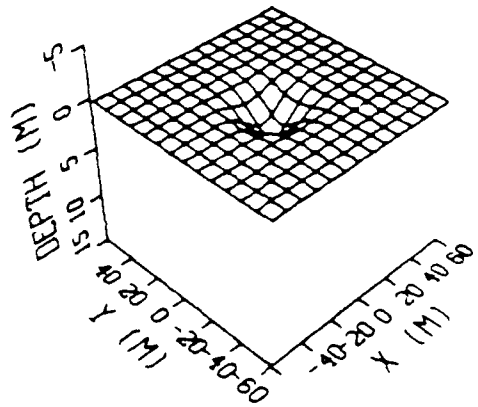
(b) $s = 96\text{m}$



(c) $s = 48\text{m}$



(d) $s = 24\text{m}$



(e) $s = 12\text{m}$

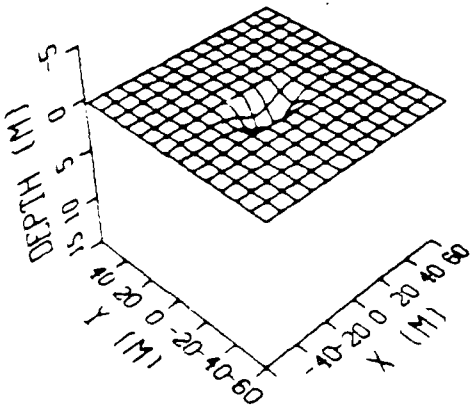


Figure 20 Sea ice keels of Model Set 1.
 $A = 3\text{m}$, $w = 24\text{m}$.

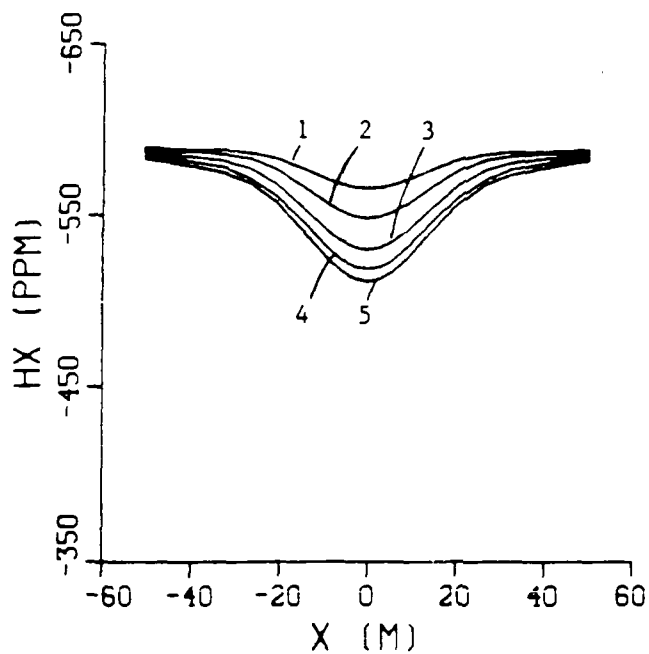


Figure 21(a) HX system response. Curves 1, 2, 3, 4, and 5 correspond to $s = 12, 24, 48, 96,$ and infinity meters respectively.

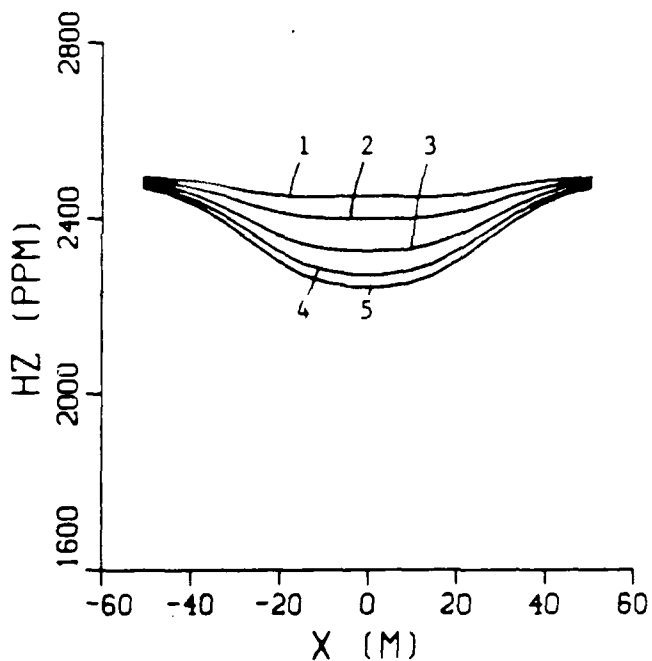


Figure 21(b) HZ system response. Curves 1, 2, 3, 4, and 5 correspond to $s = 12, 24, 48, 96,$ and infinity meters respectively.

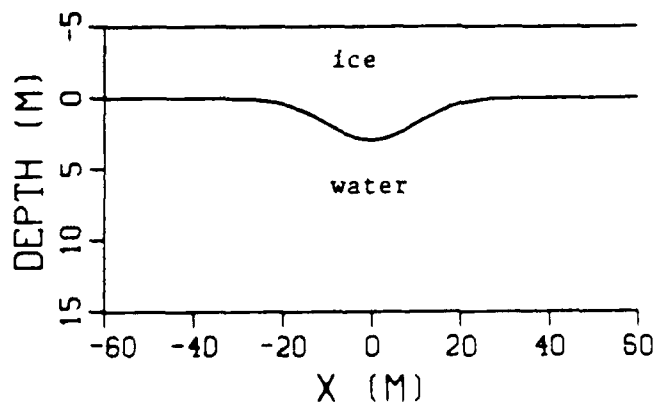


Figure 21(c) Cross section of the ice below the flight line.

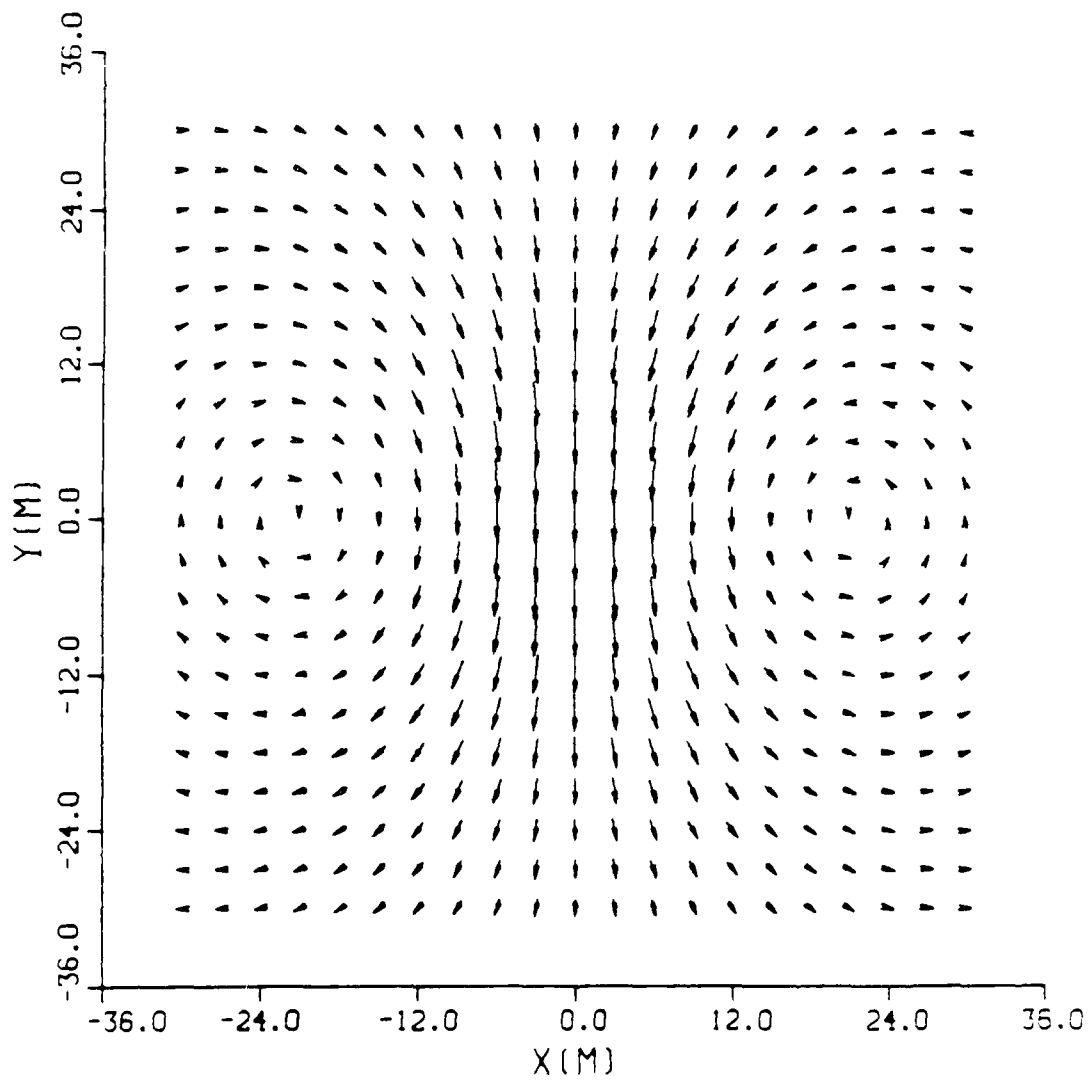


Figure 22 Surficial currents for a horizontal-axis transmitter, which is 30 meters above the center (0, 0).

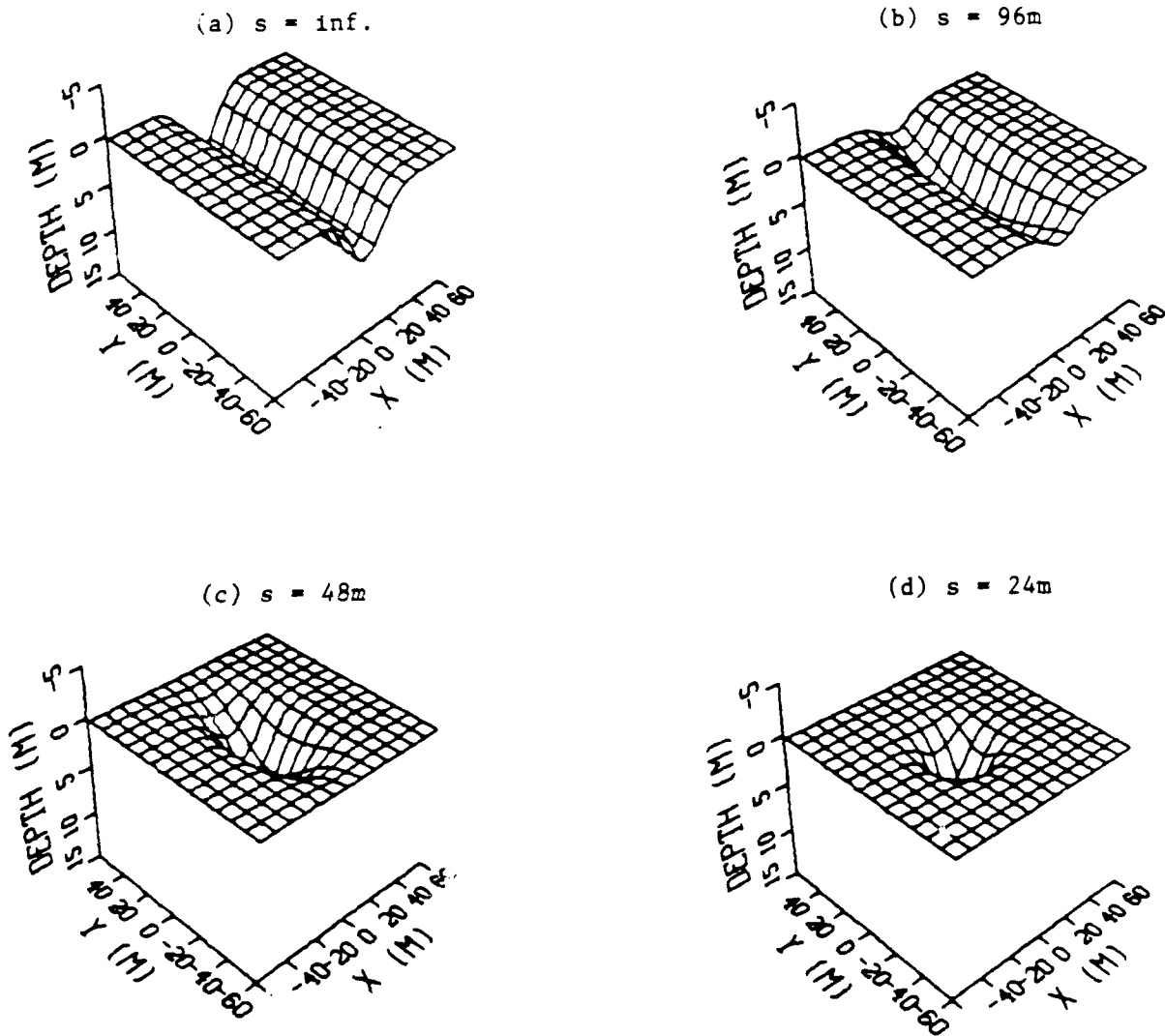


Figure 23 Sea ice keels of Model Set 2.
 $A = 6\text{m}$, $w = 24\text{m}$.

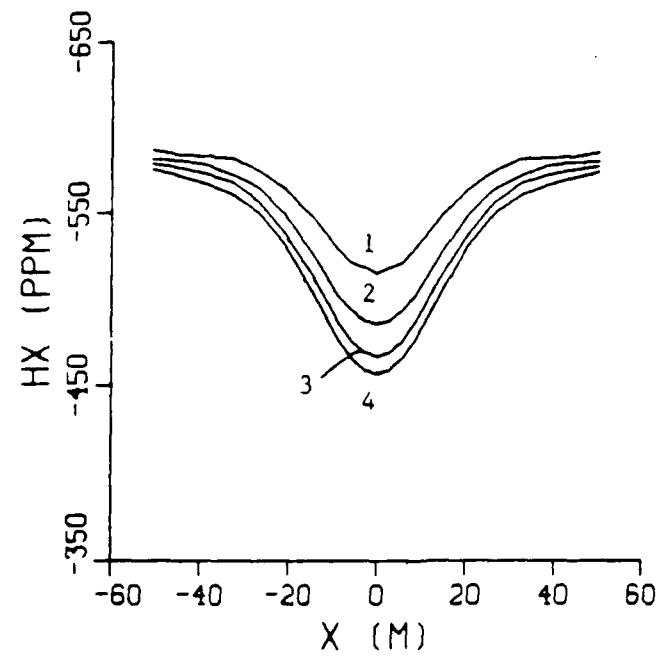


Figure 24(a) HX system response
Curves 1, 2, 3, 4 correspond to
 $s = 24, 48, 96,$ and inf. meters
respectively.

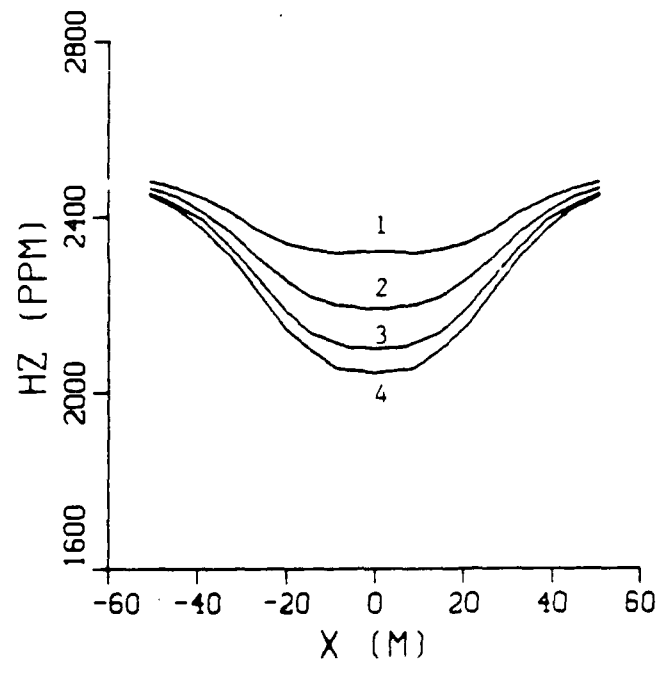


Figure 24(b) HZ system response
Curves 1, 2, 3, 4 correspond to
 $s = 24, 48, 96,$ and inf. meters
respectively.

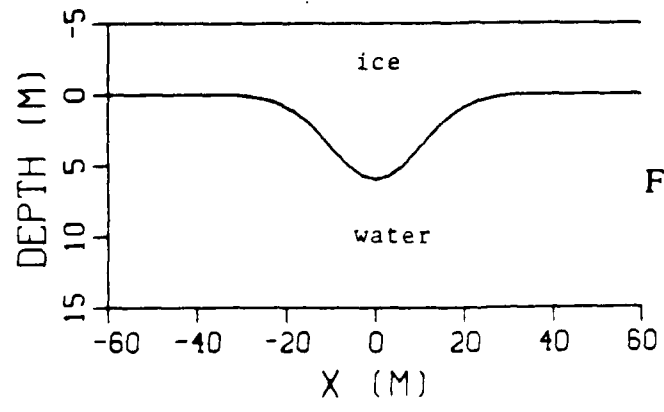


Figure 24(c) Cross section of the ice
below the flight line.

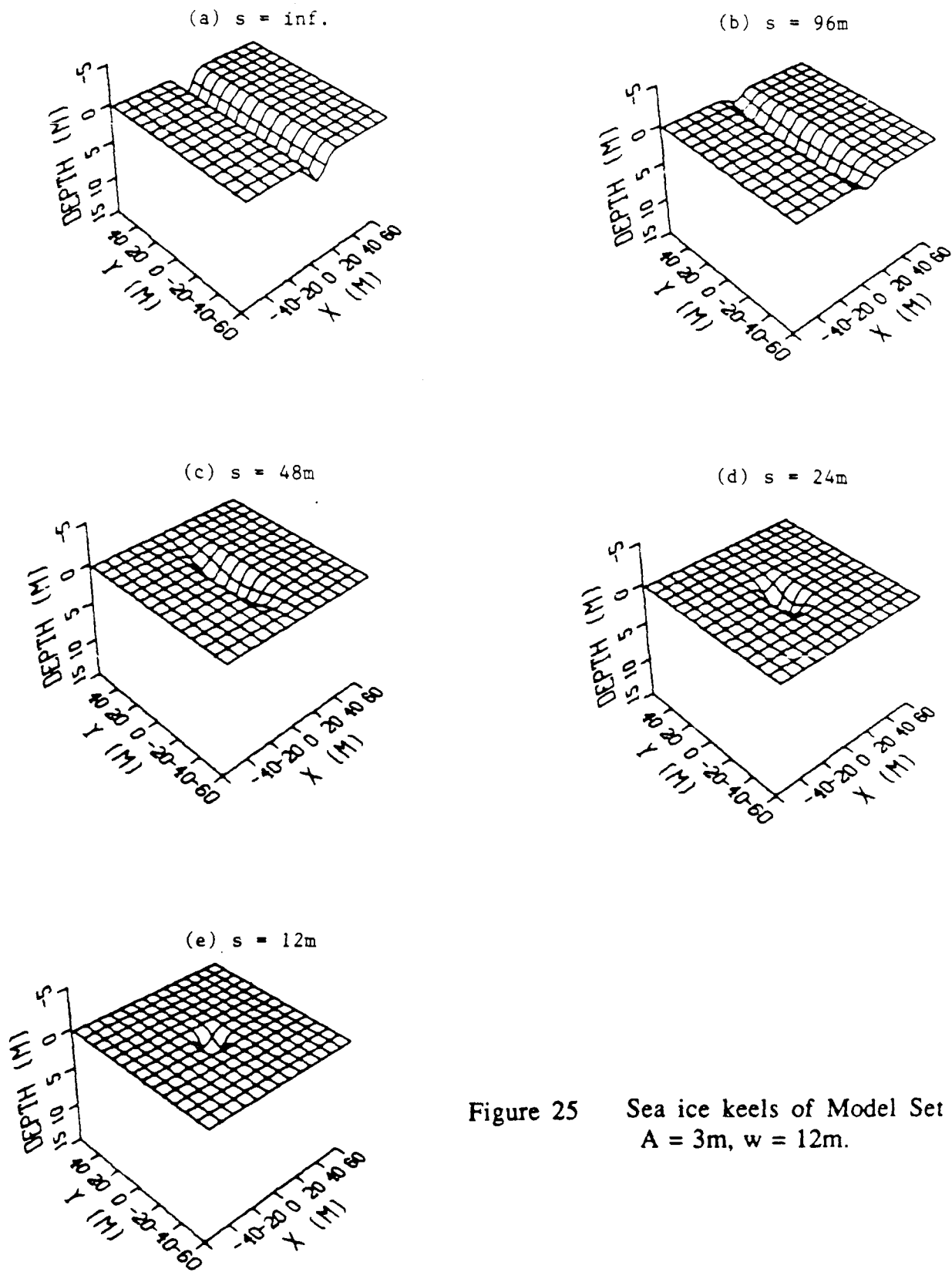


Figure 25 Sea ice keels of Model Set 3.
 $A = 3\text{m}$, $w = 12\text{m}$.

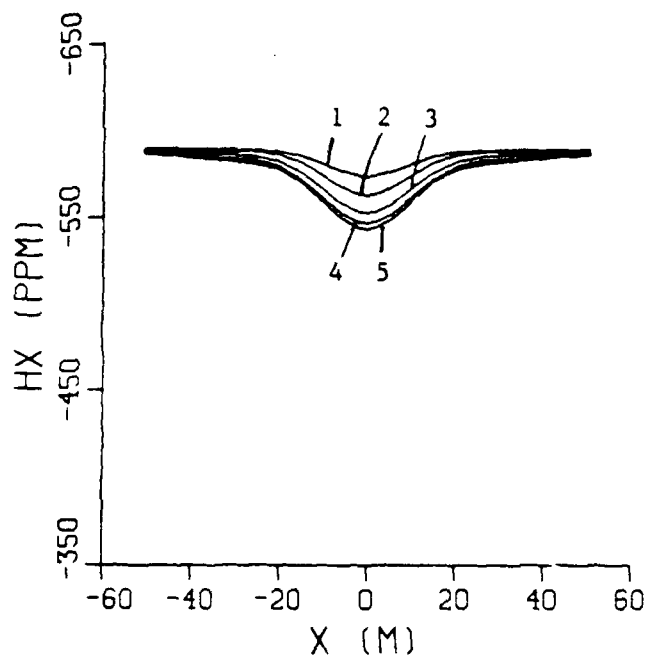


Figure 26(a) HX system response
Curves 1, 2, 3, 4 and 5 correspond to $s = 12, 24, 48, 96,$ and i meters respectively.

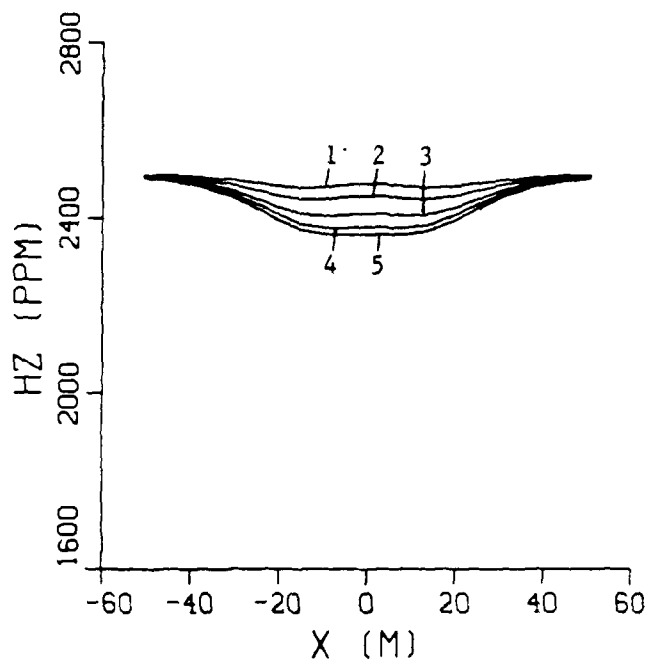


Figure 26(b) HZ system response
Curves 1, 2, 3, 4 and 5 correspond to $s = 12, 24, 48, 96,$ and i meters respectively.

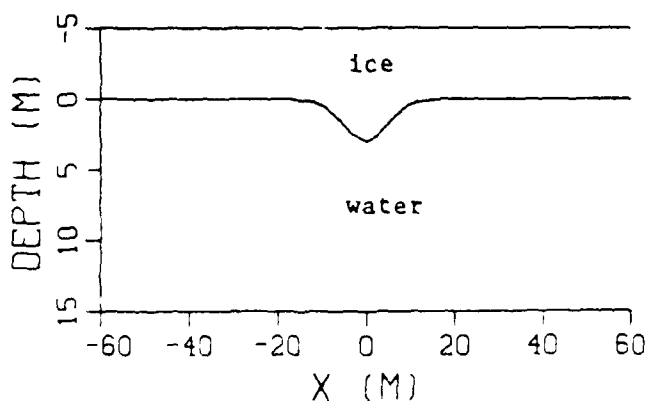


Figure 26(c) Cross section of the ice
below the flight line.

Appendix A

Computational Methods

A.1 The General Case

Consider an alternating magnetic dipole (current loop) source T, located in free space as shown in Figure A1. Its orientation is arbitrary and it is positioned over a homogeneous, perfectly conductive medium with three-dimensional surface relief. Due to this source, electric currents are induced on the surface of the medium and they give rise to the secondary magnetic field \mathbf{H}_S in free space. It is our objective to calculate this quantity at any point above the surface S.

Here the height of the source above the medium is assumed to be small compared to the wavelength. The observation point is also assumed to be close to the source so that the electromagnetic field is quasi-static. In free space, $\nabla \times \mathbf{H}_S = 0$, and we may relate \mathbf{H}_S to a scalar magnetic potential ϕ

$$\mathbf{H}_S = -\nabla \times \phi \quad (1)$$

We also have $\nabla \cdot \mathbf{H}_S = 0$ and correspondingly

$$\nabla^2 \phi = 0 \quad (2)$$

Since the lower medium is assumed to be infinitely conductive, the normal component of the total magnetic field must vanish on the surface S, i.e.

$$H_{pn} + H_{sn} = 0 \quad \text{on S} \quad (3)$$

Here, H_{pn} and H_{sn} are the normal components of the secondary and primary magnetic fields respectively. Hence,

$$\frac{\partial \phi}{\partial n} \Big|_S = -H_{sn} \Big|_S = H_{pn} \Big|_S \quad (4)$$

The Laplace equation (2) and the boundary condition (4) constitute the Neumann boundary value problem. That is, given the normal derivative of the potential on a surface S , we wish to calculate the potential itself in the free space. Once ϕ is found, H_S may be calculated from equation (1).

The solution of the Neumann problem outside a closed surface can be expressed as the potential of a surface charge layer (Graham, 1980)

$$\phi(O) = \int_S \frac{\xi(P)}{r_{PO}} ds \quad (5)$$

where $\xi(P)$ is a fictitious charge density function, r_{OP} is the distance between points P and O (see Figure 1). The charge density satisfies a Fredholm integral equation of the second kind

$$\xi(M) = - \frac{1}{2\pi} \frac{\partial \phi}{\partial n} \Big|_S - \frac{1}{2\pi} \int_S \xi(P) \frac{\cos(\mathbf{r}_{PM}, \mathbf{N})}{r_{PM}^2} ds \quad \text{on } S \quad (6)$$

where $(\mathbf{r}_{PM}, \mathbf{N})$ is the angle between \mathbf{r}_{PM} (the vector connecting P to M) and \mathbf{N} (the unit normal vector at M). In our problem, the surface S extends to infinity and equations (5) and (6) are still valid.

To solve the integral equation (6), we use the successive approximation method (Mikhlin, 1964). The initial solution (first iteration) is assumed to be the first term on the right hand side of equation (6), i.e.

$$\xi(M) = - \frac{1}{2\pi} \frac{\partial \phi}{\partial n} \Big|_S = - \frac{1}{2\pi} H_{pn}$$

We then use this value in the integral in the equation to compute an improved value of $\xi(M)$ and so on. Once the charge density is known,

the secondary magnetic field can be calculated from the following equation, which is obtained by combining equations (1) and (5)

$$\mathbf{H}_s = \int_S \frac{\xi(\mathbf{p})}{r_{PO}^3} \mathbf{r}_{PO} ds \quad (7)$$

Here \mathbf{r}_{PO} is the vector connecting P to O.

In the above context, the electrodynamic problem is reduced to a potential problem under the quasi-static field assumption. The charge distribution on the surface of the conductive medium is computed first. The secondary magnetic field in the free space is then found by summing up the contributions from the individual charges. This is analogous to the integral equation approach for solving electromagnetic scattering problem (Parry and Ward, 1970), where the equivalent electric and magnetic currents are first sought; the electromagnetic field are then obtained by integrating the contributions from the current distribution.

In the special case where the surface S forms a plane, the solution given by equation (7) for an oscillating magnetic dipole source is identical to that obtained by the method of images (Jackson, 1975). To illustrate our computational method, we now show this to be true for a vertical magnetic dipole source. The coordinate system is chosen such that the z axis is pointing vertically downward. The XOY plane is on the surface of the conductor which occupies the half space $z > 0$. The dipole source is h meters above the plane on the z axis and points in the positive z direction.

Since \mathbf{r}_{PM} is perpendicular to N (cf. Figure 1), equation (6) reduces to

$$\xi(M) = - \frac{1}{2\pi} \frac{\partial \phi}{\partial n} \Big|_S = - \frac{1}{2\pi} H_{pm} \quad (8)$$

On the conductor surface, the normal component of the primary magnetic field is

$$H_{pn} = -\frac{1}{4\pi} \frac{2h^2 - x^2 - y^2}{(x^2 + y^2 + h^2)^{5/2}} \quad (9)$$

Substituting equations (8) and (9) into (5), we obtain

$$\phi(x, y, z) = \int_{-\infty}^{+\infty} \int_{-\infty}^{+\infty} \frac{\xi(x', y', 0)}{[(x-x')^2 + (y-y')^2 + z^2]^{1/2}} dx' dy' \quad (z < 0)$$

$$\phi(0, 0, z) = \frac{1}{8\pi^2} \int_{-\infty}^{+\infty} \int_{-\infty}^{+\infty} \frac{2h^2 - x'^2 - y'^2}{(x'^2 + y'^2 + h^2)^{5/2}} \frac{dx' dy'}{(x'^2 + y'^2 + z^2)^{1/2}}$$

Let $x' = R\cos\theta$, $y' = R\sin\theta$, $dx' dy' = Rd\theta dR$, then

$$\phi(0, 0, z) = \frac{1}{8\pi^2} \int_0^{+\infty} \int_0^{2\pi} \frac{2h^2 - R^2}{(R^2 + h^2)^{5/2}} \frac{Rd\theta dR}{(R^2 + z^2)^{1/2}}$$

Furthermore, let $r^2 = R^2 + h^2$, $RdR = rdr$, then

$$\phi(0, 0, z) = \frac{1}{4\pi} \int_h^{\infty} \frac{3h^2 - r^2}{r^4 (r^2 - h^2 + z^2)^{1/2}} dr$$

The above integration can be carried out and the result is as follows

$$\phi(0, 0, z) = \frac{1}{4\pi} \frac{1}{(h-z)^2} \quad (z < 0)$$

The vertical component of the secondary magnetic field on z axis is

$$H_{sz}(0, 0, z) = -\frac{\partial\phi}{\partial z} = \frac{1}{2\pi} \frac{1}{(h-z)^3} \quad (10)$$

and as required is identical to the image field.

A.2 The 2-D keel

In the 2-D case, the geometry of the ice-water interface does not change in the strike direction (y -direction here). The relief of the interface is only a function of x , i.e. $h(x,y) = h(x)$. In this case, the potential of the scattered magnetic field is given by,

$$\phi(x, y, z) = \iint_{-\infty}^{+\infty} \frac{\xi(x', y') \sqrt{1 + [dh(x')/dx']^2}}{\sqrt{(x-x')^2 + (y-y')^2 + (z-h(x'))^2}} dx' dy' \quad (11)$$

and the surface charge density $\xi(x, y)$ satisfies

$$\begin{aligned} \xi(x, y) = & -\frac{1}{2\pi} H_{pn}(x, y) - \frac{1}{2\pi} \iint_{-\infty}^{+\infty} \xi(x', y') \frac{(x-x')dh(x')/dx' - (h(x)-h(x'))}{[(x-x')^2 + (y-y')^2 + (h(x)-h(x'))^2]^{\frac{3}{2}}} \\ & \cdot \left[\frac{1 + (dh(x')/dx')^2}{1 + (dh(x)/dx)^2} \right]^{\frac{1}{2}} dx' dy' \end{aligned} \quad (12)$$

Here $H_{pn}(x, y)$ is the normal component of the primary magnetic field at the ice-water interface.

Notice that in equations 11 and 12 the integral with regard to y' is a convolution. Taking the Fourier transform of both sides we get

$$\begin{aligned} \phi(x, k_y, z) = & \int_{-\infty}^{+\infty} \phi(x, y, z) e^{-ik_y y} dy \\ = & 2 \int_{-\infty}^{+\infty} \xi(x', k_y) \sqrt{1 + [dh(x')/dx']^2} K_0(\rho |k_y|) dx' \end{aligned} \quad (13)$$

and

$$\xi(x, k_y) = -\frac{1}{2\pi} H_{pn}(x, k_y) - \frac{1}{2\pi} \int_{-\infty}^{+\infty} \xi(x', k_y) f(x, x', k_y) dx' \quad (14)$$

where the kernel of the integration is

$$f(x, x', k_y) = 2 \left[\frac{1 + (dh(x')/dx')^2}{1 + (dh(x)/dx)^2} \right]^{\frac{1}{2}} \frac{(x-x')dh(x')/dx' - (h(x)-h(x'))}{[(x-x')^2 + (h(x)-h(x'))^2]^{\frac{1}{2}}} \cdot |k_y| K_1(\rho' |k_y|) \quad (15)$$

In the above equations

k_y = angular wave number in the y direction

$$\rho = \sqrt{(x-x')^2 + (z-h(x'))^2}$$

$$\rho' = \sqrt{(x-x')^2 + (h(x)-h(x'))^2}$$

$K_0()$ = modified Bessel function of the zeroth order. Second kind.

$K_1()$ = modified Bessel function of the first order. Second kind.

We have thus simplified the problem by decomposing a two dimensional integral equation into a number of integral equations in one dimension.

For the case where the source is a horizontal or vertical magnetic dipole, the normal component of the primary magnetic field can be analytically transformed into the wave-number domain (i.e. $H_{pn}(x, k_y)$) as follows.

The normal component of the primary magnetic field at the ice-water interface is

$$H_{pn}(x, y) = H_p(x, y) \cdot n(x, y) \quad (16)$$

where $n(x, y)$ is the outward normal at point (x, y) on the interface. It is given by

$$n = \left(\frac{h'(x)}{\sqrt{1 + [h'(x)]^2}}, 0, -\frac{1}{\sqrt{1 + [h'(x)]^2}} \right) \quad (17)$$

where $h'(x) = dh(x)/dx$. Thus

$$H_{pn}(x, y) = \frac{h'(x)}{\sqrt{1 + [h'(x)]^2}} H_{px} - \frac{1}{\sqrt{1 + [h'(x)]^2}} H_{pz} \quad (18)$$

(a) Horizontal magnetic dipole source located at $(x_s, y_s=0, z_s)$

$$H_{px}(x, y) = \frac{2(x - x_s)^2 - y^2 - (h(x) - z_s)^2}{4\pi r^5}$$

$$H_{pz}(x, y) = \frac{3(x - x_s)(h(x) - z_s)}{4\pi r^5}$$

Here

$$r = \sqrt{(x - x_s)^2 + y^2 + (h(x) - z_s)^2}$$

Substituting the above equations into (18) and taking the Fourier transform give

$$H_{pn}(x, k_y) = \frac{1}{2\pi \rho^2 \sqrt{1 + [h'(x)]^2}} \left\{ [h'(x)(x - x_s)^2 - (x - x_s)(h(x) - z_s)] |k_y|^2 K_0(\rho |k_y|) \right. \\ \left. - \{ [(h(x) - z_s)^2 - (x - x_s)^2] h'(x) + 2(x - x_s)(h(x) - z_s) \} \frac{|k_y|}{\rho} K_1(\rho |k_y|) \right\} \quad (19)$$

where

$$\rho = \sqrt{(x - x_s)^2 + (h(x) - z_s)^2}$$

As $k_y \rightarrow 0$, the asymptotic forms of the modified Bessel functions are

$$K_0(\rho |k_y|) \rightarrow -\ln \rho |k_y|$$

$$K_1(\rho |k_y|) \rightarrow \frac{1}{\rho |k_y|}$$

Therefore

$$H_{pn}(x, k_y=0) = \frac{-1}{2\pi \rho^4 \sqrt{1 + [h'(x)]^2}} \{ [(h(x) - z_s)^2 - (x - x_s)^2] h'(x) + 2(x - x_s)(h(x) - z_s) \} \quad (20)$$

(b) Vertical magnetic dipole source located at $(x_s, y_s=0, z_s)$

$$H_{px}(x, y) = \frac{3(x - x_s)(h(x) - z_s)}{4\pi r^5}$$

$$H_{pz}(x, y) = \frac{2(h(x) - z_s)^2 - y^2 - (x - x_s)^2}{4\pi r^5}$$

$$H_{pn}(x, k_y) = \frac{1}{2\pi\rho^2\sqrt{1+[h'(x)]^2}} \left\{ [-(h(x)-z_s)^2 + h'(x)(x-x_s)(h(x)-z_s)] |k_y|^2 K_0(\rho|k_y|) \right. \\ \left. - [(h(x)-z_s)^2 - (x-x_s)^2 - 2h'(x)(x-x_s)(h(x)-z_s)] \frac{|k_y|}{\rho} K_1(\rho|k_y|) \right\} \quad (21)$$

$$H_{pn}(x, k_y=0) = \frac{-1}{2\pi\rho^4\sqrt{1+[h'(x)]^2}} \{(h(x)-z_s)^2 - (x-x_s)^2 - 2h'(x)(x-x_s)(h(x)-z_s)\} \quad (22)$$

Note here that the kernel $f(x, x', k_y)$ is independent of the source and the receiver positions. Therefore it may be calculated once for each k_y value and stored for the computation of a complete profile of AEM system response. This is quite economical but impossible in the 3-D case where the matrix is too large to store. At $x = x'$, the kernel has a singularity. But this presents no difficulty for the numerical computation because it is integrable in the sense of Cauchy principal value.

The integral equation (14) can be solved for $\xi(x, k_y)$ using successive approximation method identical to the one suggested above for the general case. This needs to be done at a number of positive k_y harmonics (including $k_y = 0$). The values of $\xi(x, k_y)$ at the negative k_y harmonics may be easily obtained by its property of symmetry. As usual, the sampling in the k_y space is done on a logarithmic scale.

Once the charge density is known, the x- and z- components of the scattered magnetic field may be directly computed from the following equations

$$\begin{aligned} H_{sx}(x, k_y, z) &= -\frac{\partial \phi(x, k_y, z)}{\partial x} \\ &= 2 \int_{-\infty}^{+\infty} (x - x') \xi(x', k_y) \sqrt{1 + [dh(x')/dx']^2} \frac{|k_y|}{\rho} K_1(\rho |k_y|) dx' \end{aligned} \quad (23)$$

and

$$\begin{aligned} H_{sz}(x, k_y, z) &= -\frac{\partial \phi(x, k_y, z)}{\partial z} \\ &= 2 \int_{-\infty}^{+\infty} (z - h(x')) \xi(x', k_y) \sqrt{1 + [dh(x')/dx']^2} \frac{|k_y|}{\rho} K_1(\rho |k_y|) dx' \end{aligned} \quad (24)$$

Now that the scattered magnetic field is obtained in the wave-number domain its inverse Fourier transform will yield the desired result in the space domain. But prior to performing the inverse Fourier transform, the field values at the logarithmically-spaced points in the k_y space need to be interpolated for uniformly spaced values. This is accomplished using cubic spline interpolation.

The above algorithm has been successfully implemented and its speed is more than 20 times faster than that of the general 3-D algorithm. The computation of a 20 point profile of the AEM system response takes about only 10 seconds CPU time on the IBM 3090.

A.3 Computational Check

A scale model experiment was also performed to check the numerical solution. Aluminum was used to simulate the infinitely conductive medium at a scale of 1:250. A model airborne electromagnetic system was built at the same scale and was "flown" at a field height of 10m. The system consisted of a coplanar, vertical axis transmitter and receiver, which were operated at 6 kHz. Details of the model are shown in Figure A2 which also exhibits the traversed feature.

The cross section of the indented surface is a Gaussian curve that simulates a smooth ice keel. Its relief is given by

$$t(x) = A \exp\left(-\frac{x^2}{2\tau^2}\right) \quad (25)$$

and

$$\tau = 0.425W$$

Here

x = distance from the keel center line

t = keel thickness

A = drawdown or maximum keel thickness

W = keel width at half drawdown

The shape of the keel does not change along its strike direction. For this scale model, A and W were taken to be 3.4m and 21m respectively. We chose this type of surface because its simulation on the computer is simple as there are only two input parameters. Furthermore, it is easy to adjust these two parameters to simulate any real sea ice keel. The measurements and numerical calculation results (two iterations) for this model are displayed in Figure A3, which shows an excellent agreement between the numerical and experimental data.

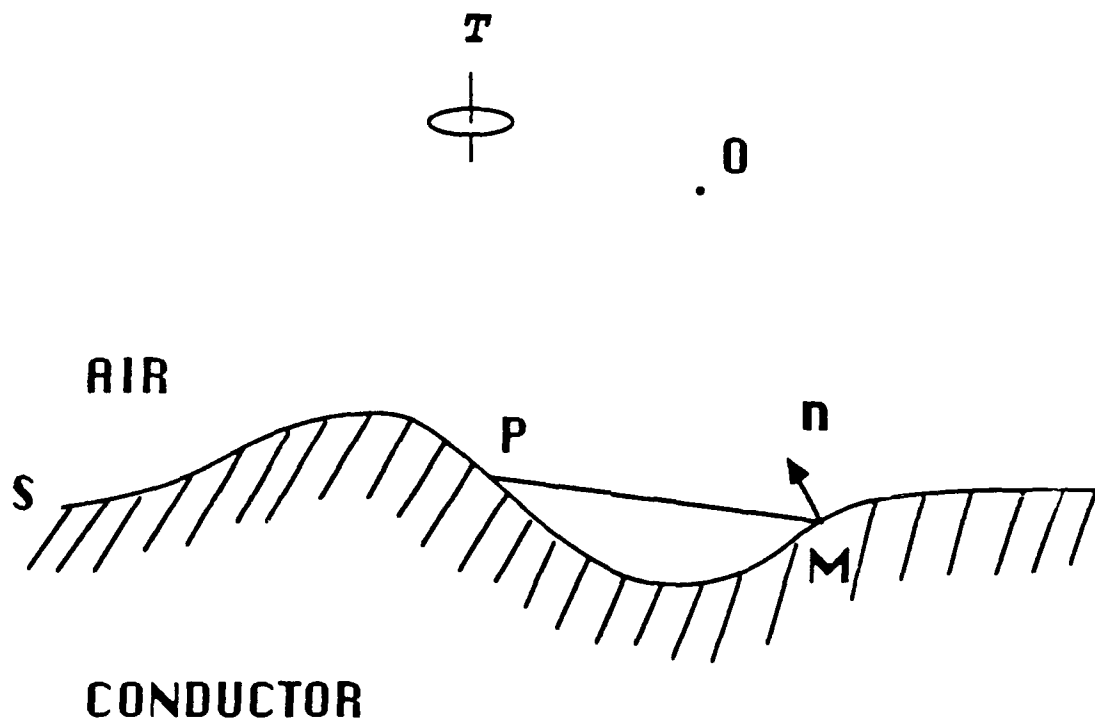


Fig.A1 The Neumann Problem

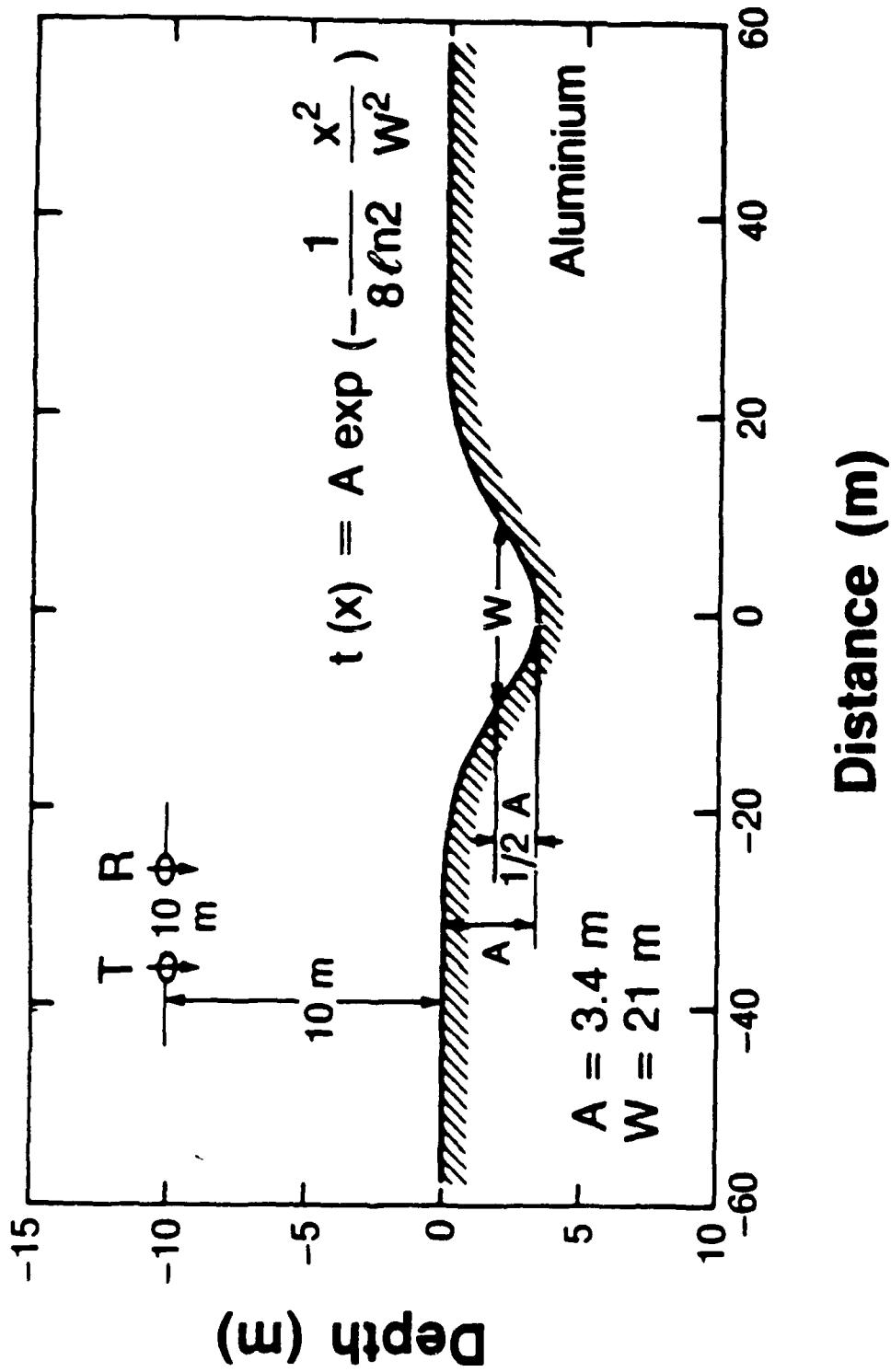


Fig.A2 Scale Model Geometry

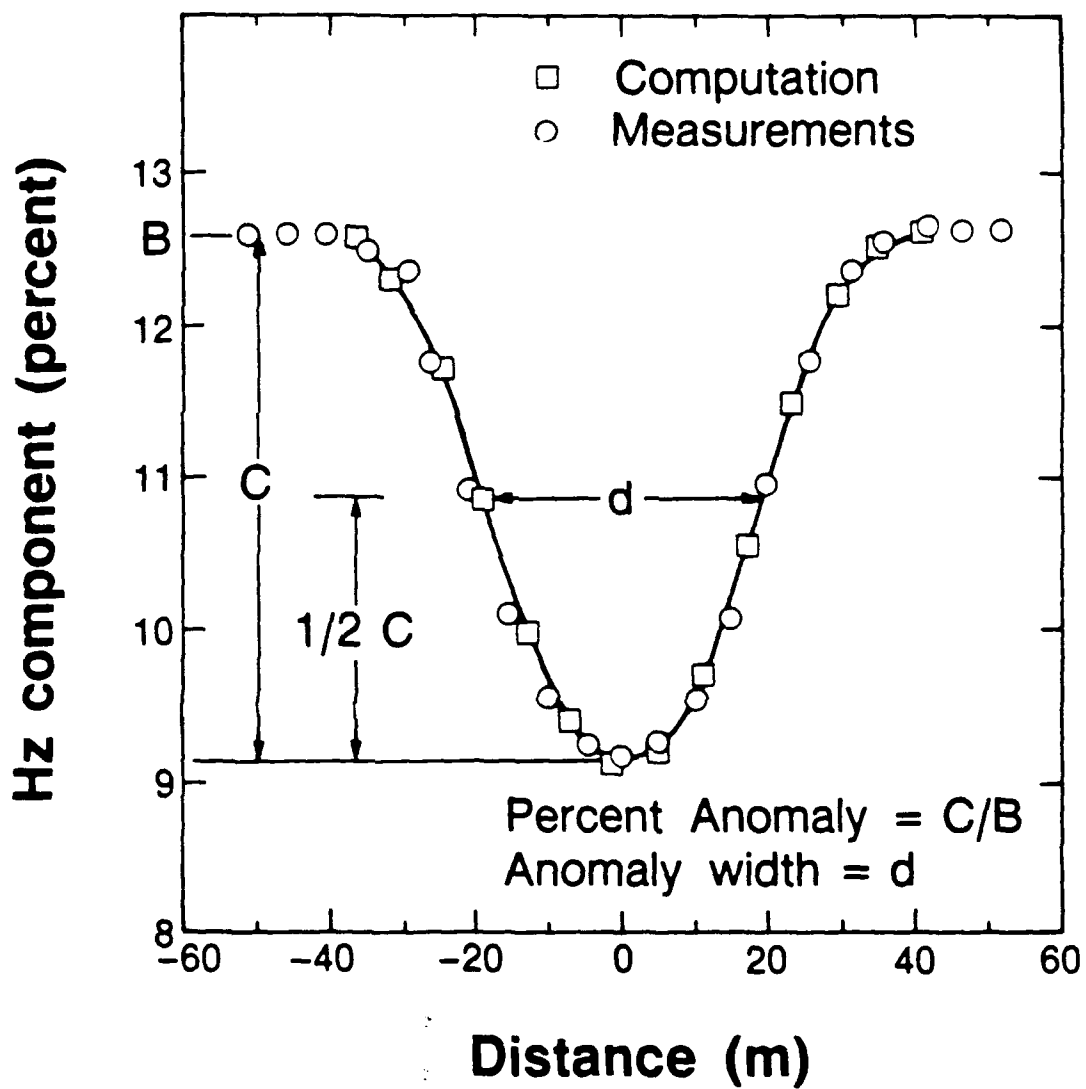


Fig.A3 Comparison of Measurements and Computation

Appendix B

This Appendix includes a listing of three Fortran programs: SURF3D, ICE2D1, and INVSM.

SURF3D is a program for computing the AEM system response for a general three-dimensional ice keel. It may also be used to calculate the surface current distribution on the ice/water interface.

ICE2D1 is a fast modeling program for the computation of AEM system response over a two-dimensional (2-D) ice keel. It is generally more than 20 times faster than using SURF3D for an identical problem.

INVSM is an inversion program for obtaining a 2-D smooth ice keel from the AEM data. It uses ICE2D1 for modeling AEM system response. Here it is acknowledged that part of this program is adapted from program OCCAM1 by Constable.


```

PROGRAM SURF3D
implicit real*8 (a-h, o-z)
*****
c c Purpose...
c c to compute the secondary electromagnetic response due to a mag-
c c netic dipole which sits above a perfectly conductive medium.
c c The secondary field is expressed in parts per million of the
c c received primary field. The orientation of the dipole is arbi-
c c trary and the surface of the conductor is arbitrary. When ipattn
c c = 1, the program outputs the total surface current density.
c c Program not optimized and can be further improved.
c c Program not documented.
c c Gaussian quadrature is used for the integration of the kernel.
c c For fast speed, ignore the Gaussian quadrature by setting 0 in
c c the IF statement.
c c Author ... Guimin Liu, Engineering Geoscience, U.C.Berkeley.
c c (415)642-3809 1988
c c *****
c c DIMENSION ZH(121,121),XDER(121,121),YDER(121,121),HN(121,121),
c c * WZ(121,121)
c c DIMENSION AMPDER(121,121),
c c COMMON /AA/ZH/BB/XDER/CC,YDER/DD/HH/EE/W/AMP/AMPDER
c c COMMON /TR/XS,YS,ZS,XR,YR,ZR
c c COMMON NX,NY,NX2,NY2,XDELTA,YDELTA
c c open (unit=3,file='surf3d.dat')
c c open (unit=4, file ='surf3d.out')

c c WHEN NPL0T=1 , THE OUTPUT IS ARRANGED FOR PLOTTING PURPOSE
c c READ(3,*) NPL0T,ipattn
c c ANGV = VERTICAL ANGLE. ANGH = HORIZONTAL ANGLE FROM X DIRECTION
c c INORM = NORMALIZATION COMPONENT. 1 = HX, 2 = HY, 3 = HZ
c c READ(3,*) ANGV, ANGH, INORM
c c READ IN NUMBER OF SOURCES
c c READ(3,*) NSOR
c c READ IN GRID NUMBER , MAXIMUM NUMBER OF ITERATIONS
c c READ(3,*) NX,NY,NLOOP
c c READ IN SURFACE PARAMETERS
c c GAUSSIAN DISTRIBUTION SHAPE
c c READ(3,*) HEIGHT,XWIDTH,YWIDTH
c c READ IN GRID SIZE
c c READ(3,*) XDELTA,YDELTA
c c NS=0
c c ANGV=ANGV/180*3.14159
c c ANGH=ANGH/180*3.14159
c c IF (INORM.EQ.1) WRITE(4,41)
c c IF (INORM.EQ.3) WRITE(4,42)
c c FORMAT(4X,'HORIZONTAL MAGNETIC DIPOLE SOURCE')
c c FORMAT(4X,'VERTICAL MAGNETIC DIPOLE SOURCE')
c c IF (NPL0T.EQ.1) GOTO 501
c c WRITE(4,112) ANGV,ANGH,INORM
c c WRITE(4,111) HEIGHT,XWIDTH,YWIDTH
c c WRITE(4,200)NX,NY,XDELTA,YDELTA,NLOOP
c c FORMAT(7X,'ANGLE OF THE DIPOLE FROM VERTICAL =',F8.2,/,
c c * ,7X,'ANGLE OF THE DIPOLE FROM X DIRECTION =',F8.2,/,
c c * ,7X,'NORMALIZATION COMPONENT 1 = HX 2 = HY 3 = HZ',I5,/,
c c * ,7X,'GRID NX =',I5,3X,' NY =',I5,
c c * ,7X,'ELEMENT SIZE DX =',F10.5,' DY =',F10.5,
c c * ,7X,'MAXIMUM ITERATION NUMBER',I8,/)
c c GOTO 666
c c 501 CONTINUE

```

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SUR00160
SUR00170
SUR00180
SUR00190
SUR00200
SUR00210
SUR00220
SUR00230
SUR00240
SUR00250
SUR00260
SUR00270
SUR00280
SUR00290
SUR00300
SUR00310
SUR00320
SUR00330
SUR00340
SUR00350
SUR00360
SUR00370
SUR00380
SUR00390
SUR00400
SUR00410
SUR00420
SUR00430
SUR00440
SUR00450
SUR00460
SUR00470
SUR00480
SUR00490

```

```

C      WRITE(4,113) ANGV, ANGH, INORM
      WRITE(4,101) HEIGHT, XWID, H, YWIDTH
      WRITE(4,201) NX, NY, XDELTA, YDELTA
101    FORMAT(5X, 'H = ', F6.2, 4X, 'XW = ', F6.2, 4X, 'YW = ', F6.2)
113    FORMAT(5X, 2F18.2, I8)
201    FORMAT(5X, 2I8, 5X, 2F18.2, //)
866    CONTINUE
C      READ IN NUMBER OF RECEIVERS
      READ(3,*) NREC
C      READ IN SOURCE POSITION
      READ(3,*) XS, YS, ZS
      NR=0
111    FORMAT(7X, 'SURFACE PARAEETER   HEIGHT = ', F7.1, 1X, ' XWIDTH = ', 1X,
      * F7.1, ' YWIDTH = ', F7.1)
C      GENERATING THE SURFACE OF THE WATER
      CALL SURF(HEIGHT, XWIDTH, YWIDTH)
C
C      CALCULATE THE PRIMARY FIELD AT THE SURFACE. GENERATE THE BOUNDARY CO
      FOR NEUMANN PROBLEM.
C      CALL HNPRIM ( ANGV, ANGH )
C
C      CALCULATE THE SURFACE CHARGE DENSITY BY SOLVING THE FREDHOLM INTEGR
      EQUATION OF THE SECOND KIND.
C      CALL SIGMA(NPLOT, NLOOP)
C
C      compute the surface current pattern and then stop
C      if(ipattn.eq.1) call currnt(angv, anvh)
C
C      CALCULATE THE MAGNETIC FIELD AT A RECEIVER POINT.
555    CONTINUE
C      READ IN RECEIVER POSITION
      READ(3,*) XR, YR, ZR
      CALL FIELD(NPLOT, ANGV, ANGH, INORM)
      NR=NR+1
      IF(NR.LT.NREC) GOTO 555
      NS=N+1
      IF(NS.LT.NSOR) GOTO 666
      STOP
      END
C
C      SUBROUTINE SURF(Z0, XW, YW)
      implicit real*8 (a-h, o-z)
      DIMENSION ZH(121,121), XDER(121,121), YDER(121,121), AMPDER(121,121)
      COMMON /AA/ZH/BB/XDER/CC/YDER/AMP/AMPER
      COMMON /TR/XS, YS, ZS, XR, YR, ZR
      COMMON NX, NY, NX2, NY2, XDELTA, YDELTA
      NX2= (NX+1) /2
      NY2= (NY+1) /2
      TAUx=0 42466* XW

```

```

SUR00500
SUR00510
SUR00520
SUR00530
SUR00540
SUR00550
SUR00560
SUR00570
SUR00580
SUR00590
SUR00600
SUR00610
SUR00620
SUR00630
SUR00640
SUR00650
SUR00660
SUR00670
SUR00680
SUR00690
SUR00700
SUR00710
SUR00720
SUR00730
SUR00740
SUR00750
SUR00760
SUR00770
SUR00780
SUR00790
SUR00800

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SUR00810
SUR00820
SUR00830
SUR00840
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SUR00860
SUR00870
SUR00880
SUR00890
SUR00900
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SUR00930
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SUR00950
SUR00960
SUR00970
SUR00980
SUR00990
SUR01000
SUR01010
SUR01020
SUR01030
SUR01040

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SUR01050
 SUR01060
 SUR01070
 SUR01080
 SUR01090
 SUR01100
 SUR01110
 SUR01120
 SUR01130
 SUR01140
 SUR01150
 SUR01160
 SUR01170
 SUR01180
 SUR01190
 SUR01200
 SUR01210
 SUR01220
 SUR01230
 SUR01240
 SUR01250
 SUR01260
 SUR01270
 SUR01280
 SUR01290
 SUR01300
 SUR01310
 SUR01320

 SUR01330
 SUR01340
 SUR01350
 SUR01360
 SUR01370
 SUR01380
 SUR01390
 SUR01400
 SUR01410
 SUR01420
 SUR01430
 SUR01440
 SUR01450
 SUR01460
 SUR01470
 SUR01480
 SUR01490
 SUR01500
 SUR01510
 SUR01520
 SUR01530
 SUR01540
 SUR01550
 SUR01560
 SUR01570
 SUR01580
 SUR01590
 SUR01600
 SUR01610
 SUR01620
 SUR01630

```

TAUY=0.42466*YW
XDEV=TAUX*TAUX*2.0
YDEV=TAUY*TAUY*2.0
DO 10 I=1,NX
DO 20 J=1,NY
X=(I-NX2)*XDELTA+XS
Y=(J-NY2)*YDELTA+YS
EX=-X*X/XDEV+Y*Y/YDEV
IF (EX.LT.(-10)) GOTO 333
RR=EXP(EX)
GOTO 444
RR=0.0
ZH(I,J)=Z0*RR
XDER(I,J)=-ZH(I,J)*X /XDEV*2.0
YDER(I,J)=-ZH(I,J)*Y /YDEV*2.0
AMPDR(I,J)=SQRT(1.0+XDER(I,J)*XDER(I,J)+YDER(I,J)*YDER(I,J))
CONTINUE
WRITE(4,100)Z0,WIDTH,XDELTA,YDELTA,RR
WRITE(4,100)((ZH(I,J),J=1,NY),I=1,NX)
WRITE(4,100)((XDER(I,J),J=1,NY),I=1,NX)
WRITE(4,100)((YDER(I,J),J=1,NY),I=1,NX)
FORMAT(3X,/,,(5E12.5))
RETURN
END
C
C
SUBROUTINE HNPTRIM (ANGV, ANGH)
implicit real*8 (a-h, o-z)
DIMENSION HN(121,121),ZH(121,121),XDER(121,121),YDER(121,121)
DIMENSION AMPDR(121,121)
COMMON /AA/ZH/BB/XDER/CC/YDER/DD/HN/AMP/AMPDR
COMMON NX,NY,NX2,NY2,XDELTA,YDELTA
COMMON /TR/XS,YS,ZS,XR,YR,ZR
COSV=COS(ANGV)
SINVS=SIN(ANGV)*COS(ANGH)
SINVSN=SIN(ANGV)*SIN(ANGH)
DO 10 I=1,NX
DO 20 J=1,NY
XD=(I-NX2)*XDELTA
YD=(J-NY2)*YDELTA
ZD=ZH(I,J)-ZS
XD2=XD*XD
YD2=YD*YD
ZD2=ZD*ZD
XYD=XD*YD
YZD=YD*ZD
ZXD=ZD*XD
R5=(XD2+YD2+ZD2)**2.5
HX1=3.0*ZXD
HY1=3.0*YZD
HZ1=(2*ZD2 -XD2 -YD2 )
HX2=(2*XD2 -YD2 -ZD2 )
HY2=3.0*XYD
HZ2=3.0*ZXD
HX3=3.0*XYD
HY3=(2*YD2 -XD2 -ZD2 )
HZ3=3.0*YZD
HX=(HX1+COSV*HX2+SINVS*HX3+SINVSN)/R5
  
```

333
 444
 20
 10
 C
 C
 C
 C
 100
 C
 C
 C

SUR01640
SUR01650
SUR01660
SUR01670
SUR01680
SUR01690
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SUR01770
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SUR01800
SUR01810
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SUR01880
SUR01890
SUR01910
SUR01920
SUR01930
SUR01940
SUR01950
SUR01960
SUR01970
SUR01980
SUR01990
SUR02000
SUR02010
SUR02020
SUR02030
SUR02040
SUR02050
SUR02060
SUR02070
SUR02080
SUR02100

20 HY=(HY1+COSV*HY2+SINVC3*HY3+SINVS4)/R5
10 HZ=(HZ1+COSV*HZ2+SINVC3*HZ3+SINVS4)/R5
C HN(I,J)=(HX*YDER(I,J)+HY*YDER(I,J)-HZ)/AMPDER(I,J)
CONTINUE
10 WRITE(4,100)((HN(I,J),J=1,NY),I=1,NX)
C FORMAT(3x,/,/, (5E12.5))
RETURN
END

C
C

SUBROUTINE SIGMA (NPL0T,NL00P)
implicit real*8 (a-h, o-z)
DIMENSION W1(121,121),ZH(121,121),YDER(121,121),YDER(121,121)
DIMENSION W2(121,121),AMPDER(121,121),HN(121,121)
DIMENSION XD(200),YD(200),XD2(200),YD2(200)
COMMON /AA/ZH/BB/XDER/CC/YDER/DD/HN/EE/W2/AMP/AMPDER
COMMON /TR/XS,YS,ZS,XR,YR,ZR
COMMON NX,NY,NX2,NY2,XDELTA,YDELTA
common /in/a,b,c,xm,ym,y1,y2,hdX,hdY

C

DXY=XDELTA*YDELTA
FAC=1.0/(2*3.1415926)
D0 10 I=1,NX
D0 10 J=1,NY
W2(I,J)=0.0
W1(I,J)=-HN(I,J)*FAC
NUM=0
N2X=NX*2
N2Y=NY*2

10

D0 13 I=1,N2X
XD(I)=(I-NX)*XDELTA
XD2(I)=XD(I)*XD(I)
D0 14 I=1,N2Y
YD(I)=(I-NY)*YDELTA
YD2(I)=YD(I)*YD(I)
CONTINUE
D0 15 I=1,NX
D0 20 J=1,NY
WS=0.000

13

14
999

D0 40 I1=1,NX
D0 30 J1=1,NY
K1= I1-I+NX
K2= J1-J+NY
ZD=ZH(I1,J1)-ZH(I,J)
IF(I.EQ.I1 .AND. J.EQ.J1) GOTO 30
IF (ABS(I1-I).LE.4 .AND. ABS(J1-J).LE.4) THEN
XM=(I-NX2)*XDELTA
YM=(J-NY2)*YDELTA
X1=(I1-NX2-0.5)*XDELTA
X2=(I1-NX2+0.5)*XDELTA
Y1=(J1-NY2-0.5)*YDELTA
Y2=(J1-NY2+0.5)*YDELTA
HDX=XDER(I,J)
HDY=YDER(I,J)
a=-XDER(I1,J1)
b=-YDER(I1,J1)
C=-ZD-a*(I1-NX2)*XDELTA-b*(J1-NY2)*YDELTA
CALL QGAUS(X1,X2,SS4)

```

WS = WS+SS4*AMPDER(I1,J1)/AMPDER(I,J)*W1(I1,J1)
else
  RPM2=XD2(K1)+YD2(K2)+ZD*ZD
  BT=(ZD*XDER(I,J)+XD(K1)-YDER(I,J)+YD(K2))*d+y
  WS=WS+BT*(I1,J1)+BT*AMPDER(I1,J1)/(RPM2*SQR(RPM2))*AMPDER(I,J)
endif
30 CONTINUE
40 CONTINUE
W2(I,J)=FAC*(-WS-HN(I,J))
20 CONTINUE
15 CONTINUE
SUM=0.0
IF(NPLOT.EQ.1) GOTO 501
D0 50 I=1,NX
D0 50 J=1,NY
  WW=W2(I,J)-W1(I,J)
50 SUM=SUM+WW*WW
  AVER=SQR(SUM)/(NX*NY)
  WRITE(4,100) AVER
100 FORMAT(20X, 'AVERAGE DIFFERENCE BETWEEN TWO SUCCESSIVE ITERATIONS',
  +,/,E10.5)
  NUM=NUM+1
  T1=ABS(AVER/W2(NX2,NY2))
  WRITE(4,122) NUM
122 FORMAT(//,/,20X, 'NUM=',I5)
  IF (.LT.0.1E-4 .OR. NUM.GT.NLOOP) GOTO 1000
  GOTO 502
501 NUM=NUM+1
502 IF ( NUM.GT.NLOOP) GOTO 1000
CONTINUE
D0 44 I=1,NX
D0 44 J=1,NY
  W1(I,J)=W2(I,J)
44 GOTO 999
1000 CONTINUE
C D0 111 I=1,NX
C111 WRITE(4,110)(W2(I,J),J=1,NY)
C110 FORMAT(1X, //,(5E12.5))
RETURN
END
C
C
SUBROUTINE FIELD(NPLOT,ANGV,ANGH,INORM)
implicit real*8 (a-h,o-z)
DIMENSION W(121,121),ZH(121,121),XDER(121,121),YDER(121,121)
DIMENSION AMPDER(121,121)
COMMON /AA/ZH,BB/XDER/CC/YDER/AMP/AMPDER
+ /EE/W
COMMON NX,NY,NX2,NY2,XDELTA,YDELTA
COMMON /TR/XS,YS,ZS,XR,YR,ZR
COSV=COS(ANGV)
SINVC=SIN(ANGV)*COS(ANGH)
SINYSN=SIN(ANGV)*SIN(ANGH)
XD=XR-XC
YD=YR-YS
ZD=ZR-ZS
XD2=XD*XD
YL2=YD*YD
SUR02090
SUR02110
SUR02120
SUR02130
SUR02140
SUR02150
SUR02160
SUR02170
SUR02180
SUR02190
SUR02200
SUR02210
SUR02220
SUR02230
SUR02240
SUR02250
SUR02260
SUR02270
SUR02280
SUR02290
SUR02300
SUR02310
SUR02340
SUR02350
SUR02360
SUR02370
SUR02380
SUR02390
SUR02400
SUR02410
SUR02420
SUR02430
SUR02440
SUR02320
SUR02330
SUR02450
SUR02460
SUR02470
SUR02700
SUR02710
SUR02390
SUR02600
SUR02610
SUR02620
SUR02630
SUR02540
SUR02760
SUR02770
SUR02100
SUR02790
SUR02800
SUR02810
SUR02820
SUR02830
SUR02840
SUR02850

```

```

ZD2=ZD*ZD
XD=XD*XD
YD=YD*YD
ZD=ZD*ZD
RS=(X2*Y2+Y2*Z2)*.5
HX1=3*ZD*XD
HY1=3*ZD*YD
HZ1=(2*ZD*XD+YD*YD)
HZ2=(2*ZD*YD+YD*ZD)
HZ3=3*ZD*XD
HZ4=3*ZD*YD
HZ5=(2*YD2+XD2-ZD2)
HZ6=3*ZD*YD
HXP=(HX1+COSV*HX2+SINVC*HX3+SINVS)/RS
HYP=(HY1+COSV*HY2+SINVC*HY3+SINVS)/RS
HZP=(HZ1+COSV*HZ2+SINVC*HZ3+SINVS)/RS
HXS=0.0
HYS=0.0
HZZ=0.0
DO 10 I=1,NX
DO 20 J=1,NY
XD=(I-NX2)*XDELTA+XS-XR
YD=(J-NY2)*YDELTA+YS-YR
ZD=ZH(I,J)-ZR
RPM2=XD*XD+YD*YD+ZD*ZD
RPM=SQR(RPM2)
HXS=HXS+W(I,J)*XD*AMPDER(I,J)/(RPM*RPM2)
HYS=HYS+W(I,J)*YD*AMPDER(I,J)/(RPM*RPM2)
HZZ=HZZ+W(I,J)*ZD*AMPDER(I,J)/(RPM*RPM2)
CONTINUE
10 CONTINUE
HXS=-HXS*XDELTA*YDELTA
HYS=-HYS*XDELTA*YDELTA
HZZ=-HZZ*XDELTA*YDELTA
IF (INDRM.EQ.1) GOTO 606
IF (INDRM.EQ.2) GOTO 607
HX=HXS/HZP*1000000
HY=HYS/HZP*1000000
HZ=HZZ/HZP*1000000
GOTO 608
606 HX=HXS/HXP*1000000
HY=HYS/HXP*1000000
HZ=HZZ/HXP*1000000
GOTO 608
607 HX=HXS/H ' *1000000
HY=HYS/HYP*1000000
HZ=HZZ/HYP*1000000
CONTINUE
IF (NPLOT.EQ.1) GOTO 501
WRITE(4,200) XS,YS,ZS
FORMAT (
* /, 7X, 'SOURCE POSITION (',F10.5,1X,F10.5,1X,F10.5,')' )
WRITE(4,100)XR,YR,ZR,HX,HY,HZ
FORMAT( 7X,'RECEIVER POSITION (',F10.5,1X,F10.5,1X,F10.5,')',/,
* 7X,'SECONDARY MAGNETIC FIELD IN PPM',/, 7X,'STATIC FIELD
* APPROXIMATION',/, 5X,
* 'HX = ',F5.8, ' HY = ',F5.8, ' HZ = ',F5.8)
GOTO 502
501 WRITE(4,101) XS,YS,ZS
SUR02860
SUR02870
SUR02880
SUR02890
SUR02900
SUR02910
SUR02920
SUR02930
SUR02940
SUR02950
SUR02960
SUR02970
SUR02980
SUR02990
SUR03000
SUR03010
SUR03020
SUR02700
SUR02710
SUR02720
SUR02730
SUR02740
SUR02750
SUR02760
SUR02770
SUR02780
SUR02790
SUR02800
SUR02810
SUR02820
SUR02830
SUR02840
SUR02850
SUR02860
SUR02870
SUR03060
SUR03070
SUR03080
SUR03090
SUR03100
SUR03110
SUR03120
SUR03130
SUR03140
SUR03150
SUR03160
SUR03170
SUR03180
SUR03190
SUR03200
SUR03210
SUR03220
SUR03230
SUR03240
SUR03250
SUR03260
SUR03270
SUR03280
SUR03290
SUR03300

```

```

101 WRITE(4,201)XR,YR,ZR,HX,HY,HZ
201 CUMAT(10X,3F18.2)
301 FORMAT(10X,3F18.2,/,10X,3E18.5,/)
502 RETURN
END

```

```

5
6
7 the following routine compute the surface currents.
SUBROUTINE currnt(angv,angh)

```

```

implicit real*8 (a-h, c-z)
DIMENSION ZH(121,121),XDER(121,121),YDER(121,121)
DIMENSION W2(121,121),AMPDER(121,121)
dimension curamp(121,121),curdir(121,121)
DIMENSION XD(200),YD(200),XD2(200),YD2(200)
COMMON /AA/ZH/BB/XDER/CC/YDER/EE/W2/AMP/AMPDER
COMMON /TR/XS,YS,ZS,XR,YR,ZR
COMMON NX,NY,NX2,NY2,XDELTA,YDELTA
common /in/a,b,c,xm,ym,y1,y2,hdx,hdy

```

```

N2X=NX*2
N2Y=NY*2

```

```

DXY=XDELTA*YDELTA
DO 13 I=1,N2X

```

```

XD(I)=(I-NX)*XDELTA
XD2(I)=XD(I)*XD(I)
DO 14 I=1,N2Y
YD(I)=(I-NY)*YDELTA
YD2(I)=YD(I)*YD(I)

```

```

DO 20 J=1,NY
curx=0.0
cury=0.0
curz=0.0

```

```

DO 40 I1=1,NX
DO 30 J1=1,NY
K1= I1-I+NX
K2= J1-J+NY

```

```

ZD=ZH(I1,J1)-ZH(I,J)
IF(I.EQ.I1 .AND. J.EQ.J1) GOTO 30
IF(ABS(I1-I).LE.4 .AND. ABS(J1-J).LE.4) THEN
XM=(I-NX2)*XDELTA
YM=(J-NY2)*YDELTA

```

```

X1=(I1-NX2-0.5)*XDELTA
X2=(I1-NX2+0.5)*XDELTA
Y1=(J1-NY2-0.5)*YDELTA
Y2=(J1-NY2+0.5)*YDELTA
HDX=XDER(I,J)
HDY=YDER(I,J)

```

```

b=-XDER(I1,J1)
b=-YDER(I1,J1)

```

```

C=-ZD-A*(I1-NX2)*XDEL B*(J1-NY2)*YDEL
CALL QCAUS2(X1,X2,SS1,SS2,SS3)

```

```

WS =W2(I1,J1)*AMPDER(I1,J1)/ampder(i,j)
CURX=WS*SS1*CURY
CURY=WS*SS2*CURX
CURZ=WS*SS3*CURZ

```

```

GOTO 30

```

```

ENDIF

```

```

RPM2=XD2(K1)+YD2(K2)+ZD*ZD
IF (RPM2.LT.0.1 ) GOTO 30

```

```

SUR03310
SUR03320
SUR03330
SUR03340
SUR03350

```

```

SUR01770
SUR01780

```

```

SUR01790
SUR01800
SUR01810
SUR01820

```

```

SUR01830
SUR01920
SUR01930

```

```

SUR01940
SUR01950
SUR01960
SUR01970
SUR01980
SUR01990
SUR02010
SUR02020

```

```

SUR02040
SUR02050
SUR02060
SUR02070
SUR02080
SUR02100

```

```

SUR02090
SUR02100

```

SUR02110
SUR02110
SUR02110

BTX=-ZD*YDER(I,J)-YD(K2)
RTY=XD(K1)*XDER(I,J)+ZD
RTZ=-YDER(I,J)+YD(K2)+XD(K1)*YDER(I,J)
WS=W2(I1,J1)*AMPDER(I1,J1)/(RPM2*SQR1(RPM2)*AMPDER(I,J))*DXY
CURX=CURX+WS*BTX
CURY=CURY+WS*BTY
CURZ=CURZ+WS*BTZ

SUR02130
SUR02140

30
40
CONTINUE
CONTINUE

C primary field calculation

SUR01390
SUR01400
SUR01410
SUR01440
SUR01450
SUR01460
SUR01470
SUR01480
SUR01490
SUR01500
SUR01510
SUR01520
SUR01530
SUR01540
SUR01550
SUR01560
SUR01570
SUR01580
SUR01590
SUR01600
SUR01610
SUR01620
SUR01630
SUR01640
SUR01650

COSV=COS(ANGV)
SINVC5=SIN(ANGV)*COS(AHIGH)
SINVS=SIN(ANGV)*SIN(ANGH)
X1=(I-NX2)*XDELTA
Y1=(J-NY2)*YDELTA
Z1=ZH(I,J)-ZS
X2=X1*X1
Y2=Y1*Y1
Z2=Z1*Z1
XYD=X1*Y1
YZD=Y1*Z1
ZXD=Z1*X1
R5=(X2+Y2+Z2)**2.5
HX1=3.0*ZXD
HY1=3.0*YZD
HZ1=(2*Z2-X2-Y2)
HX2=(2*X2-Y2-Z2)
HY2=3.0*XYD
HZ2=3.0*7XD
HX3=3.0*YD
HY3=(2*Y2-X2-Z2)
HZ3=3.0*YZD
HX=(HX1*COSV+HX2*SINVC5+HX3*SINVS)/R5
HY=(HY1*COSV+HY2*SINVC5+HY3*SINVS)/R5
HZ=(HZ1*COSV+HZ2*SINVC5+HZ3*SINVS)/R5
CURX=CURX+(HZ*YDER(I,J)+HY)/AMPDER(I,J)
CURY=CURY+(-HX-HZ*XDER(I,J))/AMPDER(I,J)
CURZ=CURZ+(HY*XDER(I,J)-HX*YDER(I,J))/AMPDER(I,J)
CURAMP(I,J)=SQR(CURX**2+CURY**2+CURZ**2)
IF(ABS(CURX) LT 1.0E-30 .AND. CURY GE 0.0) THEN
CURDIR(I,J)=90.

else
if(abs(curx).lt.1.0e-30 .and. cury.lt.0.0) then
CURDIR(I,J)=-90.

else
CURDIR(I,J)=atan2(cury,curx)*180./3.1415928
endif

SUR02160
SUR02170

20
15
CONTINUE
CONTINUE
DO 100 I=1,NX

write(4,101)I,(curamp(i,j),j=1,ny)
continue
DO 110 I=1,NX

write(4,102)I,(curdir(i,j),j=1,ny)
continue

101
102
FORMAT(4X,'AMPLITUDE OF SURFACE CURRENT I =',I5,'/',5(1PE12.3))))
FORMAT(4X,'DIRECTION OF SURFACE CURRENT I =',I5,'/',5(1PE12.3))))
stop


```

return
end

c
c
subroutine ggaus(a,b,ss4)
implicit real*8 (a-h, o-z)
c nine point Gauss quadrature
dimension x(5),w(5)
data x/0.0,0.324283,0.613371,0.836031,0.968160/
data w/0.330239,0.312347,0.260611,0.180648,0.081274/
xm=0.5*(b+a)
xr=0.5*(b-a)
ss4=0.
do 11 j=2,5
dx=xr*x(j)
ss4=ss4+w(j)*(fun4(xm+dx)+fun4(xm-dx))
11 continue
return
END

c
c
function fun4(xp)
implicit real*8 (a-h, o-z)
common /in/a,b,c,xm,ym,y1,y2,hdx,hdy
d=a*xp+c
e=(xm-xp)**2+d*d+ym*ym
f=2.*e*(d*b-ym)
g=1.+b*b
del=4*e*g-f*f
if(abs(del).lt.1.e-20) then
sg=1./sqrt(g)
fg=f/(g+g)
yg1=1./(y1+fg)
yg2=1./(y2+fg)
aix1=abs(0.5*sg*(yg2*yg2-yg1*yg1))
aix2=sg*(yg2*(-1.+0.5*fg*yg2)-yg1*(-1.+0.5*fg*yg1))
if(yg1.lt.0) aix2=-aix2
else
r2=1./sqrt(e*(f+g*y2)*y2)
r1=1./sqrt(e*(f+g*y1)*y1)
aix1=2.*((2.*g*y2*f)*r2-(2.*g*y1*f)*r1)
aix2=-2.*((2.*e+f*y2)*r2-(2.*e+f*y1)*r1)
endif
fun4=((xm-xp)*hdx+ym*hdy-d)*aix1-(hdy+b)*aix2
return
end

```

ICE00010
ICE00020
ICE00030
ICE00040
ICE00050
ICE00060
ICE00070

PROGRAM ICE2D1
implicit real*8 (a-h,o-z)
C ... PURPOSE ...
C COMPUTE THE AEM SYSTEM RESPONSE OVER A 2-D ICE-WATER PROFILE
DIMENSION ZSURF(200), XDER(200), AMPOER(200), HPN(120,20)
DIMENSION AKERN(120,120,20), CHAR(120,20), XSYS(50), ZSYS(50)
COMMON /KY/ AKY(20)/KER/ AKERN/CH/CHAR/HPN/HPN
COMMON KYE, NSYS, NST, NPT, NXT, NTRUNK, XDEL, TRDIS
open (unit=3, file='ice2d1.dat')
open (unit=4, file='ice2d1.out')

ICE00090

READ(3,*)
READ(3,*) NSYS, NST, NPT, NXT, NRESP, NTRUNK, NLOOP
READ(3,*)
READ(3,*) XDEL, TRDIS
READ(3,*)
READ(3,*) (ZSURF(I), I=1, NPT)
READ(3,*)
READ(3,*) (XSYS(I), ZSYS(I), I=1, NRESP)
IF (NSYS.EQ.1) WRITE(4,101)
IF (NSYS.EQ.2) WRITE(4,102)
IF (NSYS.NE.1 .AND. NSYS.NE.2) PAUSE 'BAD INPUT NSYS'
WRITE(4,104) TRDIS, XDEL
WRITE(4,103) NST, NPT, NXT, NRESP, NTRUNK, NLOOP
WRITE(4,105) (ZSURF(I), I=1, NPT)
WRITE(4,106) (XSYS(I), I=1, NRESP)
WRITE(4,107) (ZSYS(I), I=1, NRESP)
CALL SURF(NST, NPT, NXT, XDEL, ZSURF, XDER, AMPDER)
K=1

ICE00100

C ... ASSIGN KY HARMONICS. THE FOLLOWING IS OK FOR ZR=10 - 100 METERS.
C KYE= 15
NM=7
AKYMIN=1.0/(2.0*1000)
AKDEL=0.0006
AKLOG=0.17
AKYEX=LOG10(AKYMIN)
AKY(1)=0.0
DO 20 I=1, KYE-1
EX=AKYEX*(I-1)*AKLOG
AKY(I+1)=10**EX
20 CONTINUE
WRITE(4,201) (AKY(I), I=1, KYE)
C ...
WRITE(4,202)
1001 CONTINUE
KS=(XSYS(K)-TRDIS/2.)/XDEL+2-NTRUNK/2
IF (KS.LE.0 .OR. KS.GE.NXT-NTRUNK/2) PAUSE 'BAD INPUT XSYS'
CALL HPRIM(K, KS, ZSURF, XDER, AMPDER, XSYS, ZSYS)
CALL KERNEL(K, KS-1, KS1, ZSURF, XDER, AMPDER)
CALL CHARGE(NTRUNK, KYE, XDEL, NLOOP)
CALL FIELD(K, KS, NM, AKDEL, HXS, HZS, XSYS, ZSYS, ZSURF, AMPDER)
WRITE(4,203) XSYS(K), ZSYS(K), HXS, HZS
K=K+1
KS1=KS
IF (K.LE.NRESP) GOTO 1001
101 FORMAT(4X, 'HORIZONTAL AXIS DIPOLE SYSTEM', /)
102 FORMAT(4X, 'VERTICAL AXIS DIPOLE SYSTEM', /)
103 FORMAT(4X, 'NST =', I5, 2X, 'NPT =', I5, 2X, 'NXT =', I5, /,
+ 4X, 'NRESP =', I5, 2X, 'NTRUNK =', I5, 2X, 'NLOOP =', I5)
104 FORMAT(4X, 'T-R SEPARATION =', F7.2, ' METERS', /, 4X, 'XDEL =', F7.2,

ICE00110

ICE00120

ICE00130

ICE00140

ICE00150

ICE00160

ICE00170

ICE00180

ICE00190

ICE00200

ICE00210

ICE00220

ICE00230

ICE00240

ICE00250

ICE00260

ICE00270

ICE00280

ICE00290

ICE00300

ICE00310

ICE00320

ICE00330

ICE00340

ICE00350

ICE00360

ICE00370

ICE00380

ICE00390

ICE00400

ICE00410

ICE00420

ICE00430

ICE00440

ICE00450

ICE00460

ICE00470

ICE00480

ICE00490

ICE00500

ICE00510

ICE00520

ICE00530

ICE00540

ICE00550

ICE00560

ICE00570

ICE00580


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ICE01170
ICE01180
ICE01190
ICE01200
ICE01210
ICE01220
ICE01230
ICE01240
ICE01250
ICE01260
ICE01270
ICE01280
ICE01290
ICE01300
ICE01310
ICE01320
ICE01330
ICE01340
ICE01350
ICE01360
ICE01370
ICE01380
ICE01390
ICE01400
ICE01410

ICE01420
ICE01430
ICE01440
ICE01450
ICE01460
ICE01470
ICE01480
ICE01490
ICE01500
ICE01510
ICE01520
ICE01530
ICE01540
ICE01550
ICE01560
ICE01570
ICE01580
ICE01590
ICE01600
ICE01610
ICE01620
ICE01630
ICE01640
ICE01650
ICE01660
ICE01670
ICE01680
ICE01690
ICE01700
ICE01710
ICE01720
ICE01730
ICE01740
ICE01750

IF(NSYS.EQ.1) THEN
  GEOM1=(-XDER(J)*XD*ZD)*XD/D
  GEOM2=(-(XD*XD-ZD*ZD)*XDER(J)+2.*0*(XD*ZD))/D
ELSE
  GEOM1=(-XDER(J)*XD*ZD)*ZD/D
  GEOM2=((ZD*ZD-XD*XD)-XDER(J)*2.*0*(XD*ZD))/D
ENDIF
DO 20 I=1,KYE
  WK=AKY(I)*PI2
  RKY=RHO*WK
  IF(I.GT.1) THEN
    HPN(M,I)=(GEOM1*WK+BESSK0(RKY)+GEOM2*BESSK1(RKY)
              /RHO)*WK
  ELSE
    HPN(M,I)=GEOM2/(RHO*RHO)
  ENDIF
20 CONTINUE
10 CONTINUE
30 WRITE(4,30)(HPN(J,1),J=1,NTR)
  FORMAT(4X,'HPN',/,5(1PE12.4))
RETURN
END

SUBROUTINE KERNEL(K,KS,KS1,ZSURF,XDER,AMPDER)
implicit real*8(a-h,o-z)
C COMPUTE THE KERNEL IN THE INTEGRAL EQUATION
C AND SAVE IT
C .....
C .....
C .....
PARAMETER (PI=3.1415926)
DIMENSION ZSURF(1),XDER(1),AMPDER(1)
DIMENSION AKERN(120,120,20)
COMMON /KER/AKERN/KY/AKY(20)
COMMON KYE,MSYS,NST,NPT,NXT,NTR,XDEL,TRDIS
PI2=PI*2.
DO 1 J=1,NTR
  DO 1 I=1,KYE
    AKERN(J,J,I)=0.0
1 CONTINUE
IF(K.GT.1) GOTO 30
C COMPUTE THE AKERNEL
1001 CONTINUE
DO 20 J=1,NTR
  DO 20 L=1,NTR
    IF(J.EQ.L) GOTO 20
    XD=(J-L)*XDEL
    ZD=ZSURF(J+KS)-ZSURF(L+KS)
    RHO=SQRT(XD*XD+ZD*ZD)
    GEOM=2.*(AMPDER(L+KS)/AMPDER(J+KS))*(XD*XDER(J+KS)-ZD)/RHO
    DO 10 I=1,KYE
      WK=AKY(I)*PI2
      IF(I.EQ.1) THEN
        AKERN(J,L,I)= GEOM/RHO
      ELSE
        AKERN(J,L,I)= GEOM*WK+BESSK1(RHO*WK)
      ENDIF
10 CONTINUE
20 CONTINUE
RETURN

```

```

* METERS')
105 FORMAT(4X,'RELIEF OF ICE-WATER INTERFACE',/, (10F7.2))
106 FORMAT(4X,'X POSITION OF THE SYSTEM',/, (10F7.2))
107 FORMAT(4X,'Z POSITION OF THE SYSTEM',/, (10F7.2))
201 FORMAT(4X,'KY HOMOINICS =',/, (5E14.4))
202 FORMAT(/,4X,' XSYS (M)  ZSYS (M)  HXS (PPM)  HZS (PPM) ')
203 FORMAT(2X,2F11.2,3X,2(1PE12.4))
STOP
END
C
C
C
SUBROUTINE SURF(NST,NPT,NXT,XDEL,XSURF,ZSURF,XDER,AMPDER)
implicit real*8 (a-h,o-z)
C
C  GENERATE THE INFORMATION OF THE ICE-WATER INTERFACE, I.E.,
C ZSURF, XDER, AMPDER
C
DIMENSION ZSURF(NXT),XDER(NXT),AMPDER(NXT)
NE =NST+NPT
CALL CUBOER(NPT+2,XDEL,ZSURF,XDER)
DO 10 I=NXT,1,-1
IF (I.GE.NE .OR. I.LT.NST) THEN
ZSURF(I)=0.0
XDER(I)=0.
AMPDER(I)=1.
GOTO 10
ENDIF
ZSURF(I)=ZSURF(I-NST+1)
XDER(I)=XDER(I-NST+1)
AMPDER(I)=SQRT(1.0+XDER(I)*XDER(I))
CONTINUE
WRITE(4,50) (ZSURF(I),I=1,NXT-1)
WRITE(4,60) (XDER(I),I=1,NXT-1)
WRITE(4,70) (AMPDER(I),I=1,NXT-1)
50 FORMAT(4X,' ZSURF',/,4X,(10F7.3))
60 FORMAT(4X,' XDER',/,4X,(10F7.3))
70 FORMAT(4X,' AMPDER',/,4X,(10F7.3))
RETURN
END
C
C
SUBROUTINE HNPRIM(K,KS,ZSURF,XDER,AMPDER,XSYS,ZSYS)
C
C  COMPUTE THE NORMAL COMPONENT OF THE MAGNETIC FIELD
C
implicit real*8 (a-h,o-z)
PARAMETER (PI=3.1415926)
DIMENSION ZSURF(1),XDER(1),AMPDER(1),XSYS(1),ZSYS(1)
DIMENSION HPN(120,20)
COMMON /KY/ AKY(20)/HN/HPN
COMMON KYE,NSYS,NST,NPT,NXT,NTR, XDEL, TROIS
PI2=PI*2.
DO 10 J=KS,NTR+KS-1
M=J-KS+1
XD=(J-1)*XDEL-XSYS(K)+TROIS/2.
ZD=ZSURF(J)-ZSYS(K)
RHO=SQRT(XD*XD+ZD*ZD)
D=RHO*RHO*AMPDER(J)*PI2
IF (D.EQ.0.) PAUSE 'BAD HPN'

```

```

ICE00590
ICE00600
ICE00610
ICE00620
ICE00630
ICE00640
ICE00650
ICE00660
ICE00670
ICE00680
ICE00690
ICE00700
ICE00710
ICE00720
ICE00730
ICE00740
ICE00750
ICE00760
ICE00770
ICE00780
ICE00790
ICE00800
ICE00810
ICE00820
ICE00830
ICE00840
ICE00850
ICE00860
ICE00870
ICE00880
ICE00890
ICE00900
ICE00910
ICE00920
ICE00930
ICE00940
ICE00950
ICE00960
ICE00970
ICE00980
ICE00990
ICE01000
ICE01010
ICE01020
ICE01030
ICE01040
ICE01050
ICE01060
ICE01070
ICE01080
ICE01090
ICE01100
ICE01110
ICE01120
ICE01130
ICE01140
ICE01150
ICE01160

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ICE02350
ICE02360
ICE02370
ICE02380
ICE02390
ICE02400
ICE02410
ICE02420
ICE02430
ICE02440
ICE02450
ICE02460
ICE02470
ICE02480
ICE02490
ICE02500
ICE02510
ICE02520
ICE02530
ICE02540
ICE02550
ICE02560
ICE02570
ICE02580
ICE02590
ICE02600
ICE02610
ICE02620
ICE02630

```

```

I=1
KK=0
1001 CONTINUE
DO 10 J=1,NTR
W1(J)=HPN(J,I)*FAC
999 CONTINUE
DO 20 J=1,NTR
SUM=0.0
DO 30 L=1,NTR
SUM=SUM+W1(L)*AKERN(J,L,I)
CONTINUE
W2(J)=FAC*(-SUM*XDEL+HPN(J,I))
20 CONTINUE
DO 40 J=1,NTR
W1(J)=W2(J)
KK=KK+1
IF(KK.LT.NLOOP) GOTO 999
DO 50 J=1,NTR
CHAR(J,I)=W1(J)
50 I=I+1
IF(I.LE.KYE) GOTO 1001
WRITE(4,22)(CHAR(J,I),J=1,NTR)
22 FORMAT(4X,'CHARGE',/, (E14.4))
RETURN
END
C
C
C SUBROUTINE FIELD(K,KS,NM,AKDEL,HXS,HZS,XSYS,ZSYS,ZSURF,AMPDER)
implicit real*8 (a-h,o-z)
C
C COMPUTE THE SECONDARY MAGNETIC FIELD, BOTH
C X AND Z COMPONENTS, EXPRESSED IN PPM OF THE
C RECEIVED PRIMARY FIELD.
C
C
C
PARAMETER (PI=3.1415926)
DIMENSION AMPDER(1),CHAR(120,20),ZSURF(1),XSYS(1),ZSYS(1),AA(513)
DIMENSION TEMPI(20),TEMP2(20)
COMPLEX*16 CF(1024)
INTEGER JWK(20)
COMMON /CH/CHAR/KY/AKY(20)
COMMON KYE,NSYS,NST,NPT,NXT,NTR,XDEL,TRDIS
R=TRDIS
XF=R/2
IF(NSYS.EQ.1) THEN
HNORM=1./(2.*PI*R**3)
ELSE
HNORM=-1./(4.*PI*R**3)
ENDIF
DO 10 I=1,KYE
WK=AKY(I)*PI*2.
AZ=0.0
AX=0.0
DO 20 J=KS,NTR+KS-2
M=J-KS+1
RHO=SQRT((XSYS(K)-(J-1)*XDEL+XF)**2+(ZSYS(K)-ZSURF(J))**2)
IF(I.GT.1) THEN
AXZ=2.0*CHAR(M,I)*AMPDER(J)*WK*BESSKI(RHO*WK)/RHO
ELSE

```

```

ICE02640
ICE02650
ICE02660
ICE02670
ICE02680
ICE02690
ICE02700
ICE02710
ICE02720
ICE02730
ICE02740
ICE02750
ICE02760
ICE02770
ICE02780
ICE02790
ICE02800
ICE02810
ICE02820
ICE02830
ICE02840
ICE02850
ICE02860
ICE02870
ICE02880
ICE02890
ICE02900
ICE02910
ICE02920
ICE02930

```

```

C
C
C
30 COMPUTE THE UNKNOWN PART OF THE KERNEL
CONTINUE
M=KS-KSI+1
KE2=NTR-M
IF(M.LE.0) PAUSE 'BAD INPUT XSYS'
IF(M.GE.NTR) GOTO 1001
DO 40 J=1,KE2
  DO 40 L=1,KE2
    DO 40 I=1,KYE
      AKERN(J,L,I)=AKERN(J+M,L+M,I)
40 CONTINUE
DO 60 J=1,KE2
  DO 60 L=KE2+1,NTR
    XD=(J-L)*XDEL
    ZD=ZSURF(J+KS)-ZSURF(L+KS)
    RHO=SQRT(XD*XD+ZD*ZD)
    GEOM=2.*(AMPDER(L+KS)/AMPDER(J+KS))*(XD*XDER(J+KS)-ZD)/RHO
    DO 50 I=1,KYE
      WK=AKY(I)*PI2
      IF(I.EQ.1) THEN
        AKERN(J,L,I)= GEOM/RHO
      ELSE
        AKERN(J,L,I)= GEOM*WK*BESSKI(RHO*WK)
      ENDIF
50 CONTINUE
60 CONTINUE
DO 80 J=KE2+1,NTR
  DO 80 L=1,NTR
    IF(J.EQ.L) GOTO 80
    XD=(J-L)*XDEL
    ZD=ZSURF(J+KS)-ZSURF(L+KS)
    RHO=SQRT(XD*XD+ZD*ZD)
    GEOM=2.*(AMPDER(L+KS)/AMPDER(J+KS))*(XD*XDER(J+KS)-ZD)/RHO
    DO 70 I=1,KYE
      WK=AKY(I)*PI2
      IF(I.EQ.1) THEN
        AKERN(J,L,I)= GEOM/RHO
      ELSE
        AKERN(J,L,I)= GEOM*WK*BESSKI(RHO*WK)
      ENDIF
70 CONTINUE
80 CONTINUE
WRITE(4,90)(AKERN(40,10,I),I=1,KYE)
90 FORMAT(4X,'KERNEL',/,5(1PE12.4))
RETURN
END
C
C
SUBROUTINE CHARGE(NTR,KYE,XDEL,NL00P)
implicit real*8 (a-h,o-z)
C
C
C
C COMPUTE THE CHARGE DISTRIBUTION ON THE SURFACE OF
C THE SEA WATER
C
C DIMENSION CHAR(120,20),HPN(120,20),AKERN(120,100,100),WK(120)
C DIMENSION W2(120)
C COMMON /SER/AKERN/HPN/CHAR
C FAC=1.0/(2*3.1415926)

```

```

ICE01760
ICE01770
ICE01780
ICE01790
ICE01800
ICE01810
ICE01820
ICE01830
ICE01840
ICE01850
ICE01860
ICE01870
ICE01880
ICE01890
ICE01900
ICE01910
ICE01920
ICE01930
ICE01940
ICE01950
ICE01960
ICE01970
ICE01980
ICE01990
ICE02000
ICE02010
ICE02020
ICE02030
ICE02040
ICE02050
ICE02060
ICE02070
ICE02080
ICE02090
ICE02100
ICE02110
ICE02120
ICE02130
ICE02140
ICE02150
ICE02160
ICE02170
ICE02180
ICE02190
ICE02200
ICE02210
ICE02220
ICE02230
ICE02240
ICE02250
ICE02260
ICE02270
ICE02280
ICE02290
ICE02300
ICE02310
ICE02320
ICE02330
ICE02340
ICE02340

```

```

C .....
C CUBIC SPLINE INTERPOLATION BEFORE TAKING THE
C INVERSE FOURIER TRANSFORM.
C .....
C
C DIMENSION AKY(KYE), O(NK1), YA(KYE), Y2(20)
C CALL SPLINE(AKY, YA, KYE, 1.E30, 1.E30, 2)
C O(1)=YA(1)
C DO 10 I=2, NK1
C   XX=(I-1)*AKDEL
C   KHI=KYE
C   IF (KHI-KLO.GT.1) THEN
C     K=(KHI+KLO)/2
C     IF (AKY(K).GT.XX) THEN
C       KHI=K
C     ELSE
C       KLO=K
C     ENDIF
C   GOTO 1
C ENDIF
C H=AKY(KHI)-AKY(KLO)
C A=(AKY(KHI)-XX)/H
C B=(XX-AKY(KLO))/H
C O(I)=A*YA(KLO)+B*YA(KHI)+(A*(A-1.))*Y2(KLO)+
C   B*(B-1.)*Y2(KHI)/(H*H)/6.
C
* CONTINUE
C RETURN
C END
C
C SUBROUTINE SPLINE(X, Y, N, YP1, YPN, Y2)
C implicit real*8 (a-h, o-z)
C
C GIVEN ARRAYS X AND Y OF LENGTH N CONTAINING A TABULATED FUNCTION,
C I.E., YJ=F(XJ), WITH X1 < X2 < X3 < ... < XN, THIS ROUTINE RETURNS AN ARRAY
C Y2 OF LENGTH N WHICH CONTAINS THE SECOND DERIVATIVES OF THE INTER-
C POLATING FUNCTION AT THE TABULATED POINTS XJ. NATURAL SPLINE FROM
C NUMERICAL RECIPES, PRESS AT AL.
C .....
C
C PARAMETER (NMAX=1024)
C DIMENSION X(N), Y(N), Y2(N), U(NMAX)
C IF (YP1.GT..99E30) THEN
C   Y2(1)=0.
C   U(1)=0.
C ELSE
C   Y2(1)=-0.5
C   U(1)=(3./(X(2)-X(1)))*((Y(2)-Y(1))/(X(2)-X(1))-YP1)
C ENDIF
C DO 11 I=2, N-1
C   SIG=(X(I)-X(I-1))/(X(I+1)-X(I-1))
C   P=SIG*Y2(I-1)+2.
C   Y2(I)=(SIG-1.)/P
C   U(I)=(0.)*((Y(I+1)-Y(I))/(X(I+1)-X(I))-
C     (Y(I)-Y(I-1))/(X(I)-X(I-1)))/(X(I+1)-X(I-1))-SIG*U(I-1))/P
C
* CONTINUE
C IF (YPN.GT..99E30) THEN
C   QN=0.
C   UN=0.
C ELSE

```

```

ICE03520
ICE03530
ICE03540
ICE03550
ICE03570
ICE03580
ICE03590
ICE03600
ICE03610
ICE03620
ICE03630
ICE03640
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ICE03660
ICE03670
ICE03680
ICE03690
ICE03700
ICE03710
ICE03720
ICE03730
ICE03740
ICE03750
ICE03760
ICE03770
ICE03780
ICE03790
ICE03800
ICE03810
ICE03820

```

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ICE03830
ICE03840
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ICE03890
ICE03900
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ICE03930
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ICE03960
ICE03970
ICE03980
ICE03990
ICE04000
ICE04010
ICE04020
ICE04030
ICE04040
ICE04050
ICE04060
ICE04070
ICE04080
ICE04090
ICE04100

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```

      AXZ=2.0*CHAR(M,I)*AMPDER(J)/(RHO*RHO)
      E*DIFF
      AZ=(ZSYS(K)-ZSURF(J))*AXZ+AZ
      AX=(XSYS(K)-(J-1)*XDEL+XF)*AXZ+AX
30  CONTINUE
      TEMP1(I)=AZ*XDEL
      TEMP2(I)=AX*XDEL
10  CONTINUE
      NK=2.0*NM
      IF(NK*AKDEL.GT.AKY(KYE)) PAUSE 'BAD AKDEL OR AKY'
      CALL CURSPL(KYE,NK,AKDEL,AKY,TEMP1,AA)
      SUM=0.0
      DO 11 I=2,NK
          SUM=SUM+AA(I)
11  HZS=(SUM+AA(1))/2)*AKDEL*2.
      CALL CURSPL(KYE,NK,AKDEL,AKY,TEMP2,AA)
      SUM=0.0
      DO 12 I=2,NK
          SUM=SUM+AA(I)
12  HXS=(SUM+AA(1))/2)*AKDEL*2.
      HXS=HXS/HNORM*1000000
      HZS=HZS/HNORM*1000000
      RETURN
      END
C
C SUBROUTINE CUBDER(N,XDEL,YA,YD)
      implicit real*8 (a-h,o-z)
C
C COMPUTE THE DERIVATIVES AT EACH NODE USING
C CUBIC SPLINE.
C NOTE: N=NPT+2 ONE POINT ADDED TO EACH END.
C
C
C DIMENSION YA(50),YA(1),Y2(50),YD(N)
      DO 11 I=1,N
          YA(I)=(I-1)*XDEL
      DO 12 I=N-1,2,-1
          YA(I)=YA(I-1)
      YA(1)=0.
      YA(N)=0.
      CALL SPLINE(XA,YA,N,0.,0.,Y2)
      DO 10 I=2,N-1
          XX=XA(I)
          KLO=I
          KHI=I+1
          H=XA(KHI)-XA(KLO)
          A=(XA(KHI)-XX)/H
          B=(XX-XA(KLO))/H
          XD=XA(KHI)-XA(KLO)
          YD(I-1)=(YA(KHI)-YA(KLO))/XD-XD*((3.*AA+1.)*Y2(KLO)
              -(3.*0B-1.)*Y2(KHI))/6.
10  CONTINUE
      DO 14 I=1,N-2
14  YA(I)=YA(I+1)
      RETURN
      END
C
C SUBROUTINE CURSPL(KYE,NK1,AKDEL,AKY,YA,0)
      implicit real*8 (a-h,o-z)

```

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ICE02940
ICE02950
ICE02960
ICE02970
ICE02980
ICE02990
ICE03000
ICE03010
ICE03020
ICE03030
ICE03040
ICE03050
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ICE03070
ICE03080
ICE03090
ICE03100
ICE03110
ICE03120
ICE03130
ICE03140
ICE03150
ICE03160
ICE03170
ICE03180
ICE03190
ICE03200
ICE03210
ICE03220
ICE03230
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ICE03250
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ICE03290
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ICE03400
ICE03410
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ICE03450
ICE03460
ICE03470
ICE03480
ICE03490
ICE03500
ICE03510

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ICE04690
ICE04700
ICE04710

ICE04720
ICE04730
ICE04740
ICE04750
ICE04760
ICE04770
ICE04780
ICE04790
ICE04800
ICE04810
ICE04820
ICE04830
ICE04840
ICE04850
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ICE04880
ICE04890
ICE04900
ICE04910
ICE04920
ICE04930
ICE04940

ICE04950
ICE04960
ICE04970
ICE04980
ICE04990
ICE05000
ICE05010
ICE05020
ICE05030
ICE05040
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ICE05060
ICE05070
ICE05080
ICE05090
ICE05100
ICE05110
ICE05120
ICE05130
ICE05140
ICE05150
ICE05160
ICE05170

Function BESSII(x)

implicit real*8 (a-h,o-z)

..... Returns the modified Bessel function I1(x) for any real x

.....
real*8 y,p1,p2,p3,p4,p5,p6,p7,q1,q2,q3,q4,q5,q6,q7,q8,q9
data p1,p2,p3,p4,p5,p6,p7/0.5d0,0.87890594d0,0.51498869d0,
* 0.15084934d0,0.2658733d-1,0.301532d-2,0.32411d-3/
data q1,q2,q3,q4,q5,q6,q7,q8,q9/0.39894228d0,-0.3988024d-1,
* -0.362018d-2,0.163801d-2,-0.1031555d-1,0.2282967d-1,
* -0.289631d-1,0.1787654d-1,-0.420059d-2/
if(abs(x).lt.3.75) then

.....
y=(x/3.75)**2
BESSII=x*(p1+y*(p2+y*(p3+y*(p4+y*(p5+y*(p6+y*p7))))))

.....
else

.....
ax=abs(x)

.....
y=3.75/ax

.....
BESSII=(exp(ax)/sqrt(ax))*(q1+y*(q2+y*(q3+y*(q4+y*(q5+y*(q6+y*(q7+y*(q8+y*q9))))))

.....
endif

.....
return

.....
end

Function BESSKI(x)

implicit real*8 (a-h,o-z)

..... Returns the modified Bessel function K1(x) for positive x

.....
real*8 y,p1,p2,p3,p4,p5,p6,p7,q1,q2,q3,q4,q5,q6,q7
data p1,p2,p3,p4,p5,p6,p7/1.0d0,0.15443144d0,-0.67278579d0,
* -0.18156897d0,-0.1919402d-1,-0.110404d-2,-0.4686d-4/
data q1,q2,q3,q4,q5,q6,q7/1.25331414d0,0.23498619d0,-0.3655620d-1,
* 0.1604268d-1,-0.780353d-2,0.325614d-2,-0.68245d-3/
if(x.le.2.0) then

.....
y=x*x/4.0

.....
BESSKI=(log(x/2.0)*BESSII(x)+(1.0/x)*(p1+y*(p2+y*(p3+y*(p4+y*(p5+y*(p6+y*p7))))))

.....
else

.....
if(x.gt.40) then

.....
BESSKI=0.0

.....
else

.....
y=(2.0/x)

.....
BESSKI=(exp(-x)/sqrt(x))*(q1+y*(q2+y*(q3+y*(q4+y*(q5+y*(q6+y*(q7+y*(q8+y*q7))))))

.....
endif

.....
return

.....
end


```

QN=0.5
UN=(3./(X(N)-X(N-1)))*(Y(N)-(Y(N-1)))/(X(N)-X(N-1))
ENDIF
Y2(N)=(UN-QN*U(N-1))/(QN*Y2(N-1)+1.)
DO 12 K=N-1,1,-1
  Y2(K)=Y2(K)*Y2(K+1)+U(K)
CONTINUE
RETURN
END

12
C
C
C
C
Function BESSI0(x)
implicit real*8 (a-h,o-z)
.....
RETURNS THE MODIFIED BESSEL FUNCTION I0(X) FOR ANY REAL X
.....
real*8 y,p1,p2,p3,p4,p5,p6,p7,q1,q2,q3,q4,q5,q6,q7,q8,q9
data p1,p2,p3,p4,p5,p6,p7/1.0d0,3.5156229d0,3.0899424d0,
+ 1.2067492d0,0.2659732d0,0.360788d-1,0.45813d-2/
data q1,q2,q3,q4,q5,q6,q7,q8,q9/0.39894228d0,0.1328592d-1,
+ 0.225319d-2,-0.167565d-2,0.916281d-2,-0.2057706d-1,
+ 0.2635537d-1,-0.1647833d-1,0.392377d-2/
if(abs(x).lt.3.75) then
  y=(x/3.75)**2
  BESSI0=p1+y*(p2+y*(p3+y*(p4+y*(p5+y*(p6+y*p7))))
else
  ax=abs(x)
  y=3.75/ax
  BESSI0=(exp(ax)/sqrt(ax))*(q1+y*(q2+y*(q3+y*(q4+y*(q5+
+ y*(q6+y*(q7+y*(q8+y*q9)))))))
endif
return
end

C
C
Function BESSK0(x)
implicit real*8 (a-h,o-z)
.....
Returns the modified Bessel function K0(x) for positive x
.....
real*8 y,p1,p2,p3,p4,p5,p6,p7,q1,q2,q3,q4,q5,q6,q7
data p1,p2,p3,p4,p5,p6,p7/-0.57721566d0,0.42278420d0,0.23069756d0,
+ 0.3488590d-1,0.262698d-2,0.10760d-3,0.74d-5/
data q1,q2,q3,q4,q5,q6,q7/1.26331414d0,-0.7832358d-1,0.2189568d-1,
+ -0.1062446d-1,0.587872d-2,-0.251540d-2,0.53208d-3/
if(x.le.2.0) then
  y=x*x/4.0
  BESSK0=(-log(x/2.0)+BESSI0(x))*(p1+y*(p2+y*(p3+
+ y*(p4+y*(p5+y*(p6+y*p7))))))
else
  if(x.gt.40.) then
    BESSK0=0.0
  else
    y=(2.0/r)
    BESSK0=(exp(-x)/sqrt(x))*(q1+y*(q2+y*(q3+y*(q4+y*(q5+
+ y*(q6+y*q7))))))
  endif
endif
return
end

```

ICE04110
ICE04120
ICE04130
ICE04140
ICE04150
ICE04160
ICE04170
ICE04180
ICE04190
ICE04200
ICE04210
ICE04220

ICE04230
ICE04240
ICE04250
ICE04260
ICE04270
ICE04280
ICE04290
ICE04300
ICE04310
ICE04320
ICE04330
ICE04340
ICE04350
ICE04360
ICE04370
ICE04380
ICE04390
ICE04400
ICE04410
ICE04420
ICE04430
ICE04440
ICE04450

ICE04460
ICE04470
ICE04480
ICE04490
ICE04500
ICE04510
ICE04520
ICE04530
ICE04540
ICE04550
ICE04560
ICE04570
ICE04580
ICE04590
ICE04600
ICE04610
ICE04620
ICE04630
ICE04640
ICE04650
ICE04660
ICE04670
ICE04680


```

IF (KK.GT.NITER) THEN
  WRITE(4,17)
  STOP
ENDIF
ENDIF
11 FORMAT(4X,'COMPUTED ICE-WATER INTERFACE',/, (5(1PE14.2)))
12 FORMAT(4X,'COMPUTED DATA',/, (5(1PE14.4)))
16 FORMAT(4X,'FIRST CONVERGENCE CRITERION SATISFIED')
17 FORMAT(4X,'SECOND CONVERGENCE CRITERION SATISFIED')
101 FORMAT(4X,'HORIZONTAL AXIS DIPOLE TRANSMITTER',/)
102 FORMAT(4X,'VERTICAL AXIS DIPOLE TRANSMITTER',/)
103 * 4X,'NRESP',,15,2X,'NPT',,16,2X,'NXT',,16,/,
  * 15,2X,'NITER',,16)
104 * 4X,'T-R SEPARATION',,F7.2,' METERS',/,4X,'XDEL',,F7.2,
  * 4X,'FIRST CONVERGENCE CRITERIA RMS ERR',,1PE12.2)
105 FORMAT(4X,'INITIAL ICE-WATER INTERFACE',/, (10F7.2))
106 FORMAT(11X,'ZSYS (M) DATA IN PPM',/, (3F17.2))
GOTO 1001
END

C
C
C SUBROUTINE HARMOC(KYE,NM,AKDEL,AKY)
C
C .....
C ASSIGN KY HARMONICS. THE FOLLOWING IS OK FOR ZR=10 - 100 METERS.
C IMPLICIT DOUBLE PRECISION (A-H, O-Z)
C DIMENSION AKY(KYE)
C KYE=15
C NM=7
C AKYMIN=1.0/(2.0*1000)
C AKDEL=0.00006
C AKLOG=0.17
C AKYEX=DLOG10(AKYMIN)
C AKY(1)=0.0
C DO 20 I=1,KYE-1
  EX=AKYEX*(I-1)*AKLOG
  AKY(I+1)=10**EX
C CONTINUE
20 WRITE(4,201)(AKY(I),I=1,KYE)
201 FORMAT(4X,'KY HARMONICS =',/, (5E14.4))
C .....
C RETURN
C END
C
C
C SUBROUTINE FORDIV(NPT,NRESP,ZH1,CALC,WJ)
C
C COMPUTE THE APPROXIMATE JACOBIAN MATRIX
C
C IMPLICIT DOUBLE PRECISION (A-H, O-Z)
C PARAMETER (NDD=40,NPP=30,NXT=200)
C PARAMETER (PI2=3.1415926*2.,PI=3.1415926,KM=250)
C DIMENSION ZH0(NXT),XSYS(NDD),ZSYS(NDD),WJ(NDD,NPP),ZH1(1)
C DIMENSION AMPDER(NXT),TEMP(20),AA(KM),ZH2(NPP),CALC2(NDD),CALC(1)
C COMMON /FCM/ZH0,XSYS,ZSYS,AMPDER,NM,IREC,AKDEL
C COMMON /KY/AKY(20)
C COMMON KYE,NSYS,NST,XDEL,TRDIS
C CALL FORMOD(NPT,NRESP,ZH1,CALC)
C K=1
1001 CONTINUE
DO 10 I=1, NPT

```

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INV00600
INV00610
INV00620
INV00630
INV00640
INV00650
INV00660
INV00670
INV00680
INV00690
INV00700
INV00710
INV00720
INV00730
INV00740
INV00750
INV00760
INV00770
INV00780
INV00790
INV00800
INV00810
INV00820
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INV00960
INV00970
INV00980
INV00990
INV01000
INV01010
INV01020
INV01030
INV01040
INV01050
INV01060
INV01070
INV01080
INV01090
INV01100
INV01110
INV01120
INV01130
INV01140
INV01150
INV01160
INV01170
INV01180
INV01190

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INV00010
INV00020
INV00030
INV00040
INV00050
INV00060
INV00070
INV00080
INV00090
INV00100
INV00110
INV00120
INV00130
INV00140
INV00150
INV00160

INV00180
INV00190
INV00200
INV00210
INV00220
INV00230
INV00240
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INV00500
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INV00570
INV00580
INV00590

PROGRAM INVSM
C INVERSION PROGRAM FOR IMAGING THE ICE-WATER INTERFACE FROM THE
C AIRBORNE ELECTROMAGNETIC DATA.
C QUADRATIC PROGRAMMING APPROACH.
C SMOOTH INVERSION

IMPLICIT DOUBLE PRECISION (A-H, O-Z)
PARAMETER (NXT=200, NTRUNK=100, NLOOP=8, NDD=40, NPP=30)
DIMENSION FKERN(NDD, NPP), DELZH(NDD), AMPDER(NXT), ZH0(NXT)
DIMENSION XSYS(NDD), ZSYS(NDD), CALC(NDD)
COMMON /DATA/NRESP, DAT(NDD), SD(NDD)
COMMON /MODEL/NPT, ZH1(NPP), DM(NDD), L_JMP(NPP)
COMMON /KY/ AKY(20)

COMMON /FCM/ZH0, XSYS, ZSYS, AMPDER, NM, IREC, AKDEL
COMMON KYE, NSYS, NST, XDEL, TRDIS
open (unit=3, file='invhz.dat')
open (unit=4, file='invhz.out')

READ(3,*)
READ(3,*) NSYS, IREC, NST, NPT, NRESP, NITER
READ(3,*)
READ(3,*) XDEL, TRDIS, ERR, PMU
READ(3,*)
READ(3,*) (ZH1(I), I=1, NPT)
READ(3,*)
READ(3,*) (XSYS(I), ZSYS(I), DAT(I), CALC(I), I=1, NRESP)
READ(3,*)
READ(3,*) (SD(I), I=1, NRESP)
READ(3,*)
READ(3,*) (L_JMP(I), I=1, NPT)
IF (IREC.EQ.2) THEN
C O 3 I=1, NRESP
DAT(I)=CALC(I)
ENDIF

3 OUTPUT THE INPUT DATA

IF (NSYS.EQ.1) WRITE(4,101)
IF (NSYS.EQ.2) WRITE(4,102)
IF (NSYS.NE.1 .AND. NSYS.NE.2) PAUSE 'BAD INPUT NSYS'
IF (IREC.NE.1 .AND. IREC.NE.2) PAUSE 'BAD INPUT IREC'
WRITE(4,104) TRDIS, XDEL, ERR
WRITE(4,103) NST, NPT, NPT, NRESP, NTRUNK, NLOOP, IREC, NITER
WRITE(4,105) (ZH1(I), I=1, NPT)
WRITE(4,106) (XSYS(I), ZSYS(I), DAT(I), I=1, NRESP)

C ASSIGN THE KY HARMONICS

C CALL HARMOC(KYE, NM, AKDEL, AKY)

KK-1

1001 CONTINUE

KK=KK+1

C CALL THE SMOOTH INVERSION ROUTINE

CALL OCCAM1(1, ERR, TOBT, 0, NIT, STEPSZ, KONV, 4, PMU)

WRITE(4,11) (ZH1(I), I=1, NPT)

WRITE(4,12) (DM(I), I=1, NRESP)

IF (TOBT.LT.ERR) THEN

WRITE(4,10)

STOP

ELSE

```

C TOBT = RMS MISFIT OF NEW MODEL RESPONSE.
C NIT = NUMBER OF PREVIOUS CALLS TO OCCAMI DURING THIS INVERSION
C STEPSZ = SUM OF SQUARES OF THE CHANGES IN THE MODEL PARAMETERS
C KONV = STATUS FLAG:
C 0 = NORMAL EXIT FOR A SUCCESSFUL ITERATION.
C 1 = A PERFECTLY SMOOTH MODEL HAS BEEN FOUND FOR THE REQUIRED TOLERANCE
C 2 = CONVERGENCE PROBLEMS. UNABLE TO FIND A MODEL WITH AN R.M.S. MISFIT
C LESS THAN OR EQUAL TO THAT OF THE INITIAL MODEL.
C UM = THE MU VALUE USED AT THE MINIMUM OF THIS ITERATION
C
C SUBROUTINES REQUIRED AND SUPPLIED:
C MAKDEL, A SUBROUTINE WHICH CREATES THE ROUGHENING MATRIX
C TOFMU(AMU). A FUNCTION WHICH RETURNS THE RMS MISFIT OF THE RESPONSE
C OF THE MODEL CONSTRUCTED USING THE LAGRANGE MULTIPLIER AMU.
C CALLS CHOLIN, CHOLSL, FORMOD, ANORM.
C INITM, TRMLT, MULT AND ANORM, SUBROUTINES FOR SIMPLE MATRIX OPERATIONS
C FMIN, A SUBROUTINE WHICH MINIMISES A UNIVARIATE FUNCTION
C FRONT, A SUBROUTINE WHICH FINDS THE ROOT OF A UNIVARIATE FUNCTION
C MIN, RK, A SUBROUTINE WHICH BRACKET'S A UNIVARIATE MINIMUM
C
C SUBROUTINES WHICH MUST BE SUPPLIED BY THE USER:
C FORMOD(NP,ND,PM,DP,DM), COMPUTES THE FORWARD FUNCTION FOR MODEL PM()
C THE DATA PARAMETERS DP() AND RETURNS IT IN DM().
C FORDIV(NP,ND,PM,DP,DM,WJ), COMPUTES THE FORWARD FUNCTION AS DOES FORMOD
C BUT ALSO RETURNS THE MATRIX OF PARTIALS, WJ(ND,NP).
C
C PARAMETERS NDD AND NPP SHOULD BE SET TO THE MAXIMUM DIMENSIONS OF THE
C DATA AND PARAMETER VECTORS, RESPECTFULLY. DON'T FORGET TO ADJUST
C THEM IN TOFMU, MAKDEL, BLOCKDATA AND ANY I/O SUBROUTINES.
C
C NOTE THAT DOUBLE PRECISION IS USED THROUGHOUT. SOME COMPILERS WILL
C EXTERNAL TOFMU TO BE DECLARED IN OCCAMI'S CALLING PROGRAM.
C
C IMPLICIT DOUBLE PRECISION (A-H, O-Z)
C PARAMETER (NDD = 40, NPP = 30)
C COMMON /DATA/ ND, D(NDD), SD(NDD)
C COMMON /MODEL/ NP, PM(NPP), DM(NDD), L, JMP(NPP)
C /RESULT/ IS USED TO CARRY INFO IN AND OUT OF TOFMU AND STORE LAST
C VALUE OF MU USED.
C COMMON /RESULT/ PTP(NPP,NPP), WJTWJ(NPP,NPP), WJTWJ(NPP), PM2(NPP),
C • PMU,NFOR,NITER,FRAC,RLAST,IFFTOL
C DIMENSION WJ(NDD,NPP),DD(NDD),DHAT(NDD),WK(NPP)
C EXTERNAL TOFMU
C
C ON FIRST ENTRY CREATE TRANS(P), WHERE P IS THE ROUGHENING MATRIX OR
C ROUGHENING MATRIX SQUARED, AND SET MU TO AN ARBITRARY (LARGE) VALUE.
C
C IF (NITER .EQ. 0) THEN
C CALL MAKDEL(IRUF)
C PMU = 1.0000.
C PMU=UM
C RLAST = 1.0000.
C KONV = 0
C END IF
C
C FRAC CONTROLS THE STEP SIZE; NORMALLY WILL REMAIN AT 1.0 UNLESS WE
C CONVERGENCE PROBLEMS
C FRAC = 1.0
C CALL FORDIV(NP,ND,PM,DM,WJ)
C WRITE(IOUT,11) (DM(1),I=1,ND)

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INV01800
INV01810
INV01820
INV01830
INV01840
INV01850
INV01860
INV01870
INV01880
INV01890
INV01900
INV01910
INV01920
INV01930
INV01940
INV01950
INV01960
INV01970
INV01980
INV01990
INV02000
INV02010
INV02020
INV02030
INV02040
INV02050
INV02060
INV02070
INV02080
INV02090
INV02100
INV02110
INV02120
INV02130
INV02140
INV02150
INV02160
INV02170
INV02180
INV02190
INV02200
INV02210
INV02220
INV02230
INV02240
INV02250
INV02260
INV02270
INV02280
INV02290
INV02300
INV02310
INV02320
INV02330
INV02340
INV02350
INV02360
INV02370
INV02380
INV02390

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```

10      ZH2(I) = ZH1(I)
      ZH2(K) = ZH1(K)*0.2
      CALL FORMOD(NPT,NRESP,ZH2,CALC2)
      DO 30 I=1, NRESP
      WJ(I,K) = (CALC2(I)-CALC(I))/0.2
30      CONTINUE
      K=K+1
      IF(K.LE.NPT) GOTO 1001
      DO 51 I=1,NRESP
      WRITE(4,50)(WJ(I,J),J=1,NPT)
C61      CONTINUE
50      FORMAT(4X,'INVERSION KERNEL',/, (6(1PE14.4)))
      RETURN
      END
C.....
      SUBROUTINE OCCAM1(IRUF,TOLREQ,TOBT,IBUG,NIT,STEPSZ,KONV,IOUT,UM)
C
C   MODIFIED VERSION OF OCCAM1
C   OCCAM1 EXECUTES ONE ITERATION OF A ONE DIMENSIONAL SMOOTH MODEL FINDER
C   REFERENCE: CONSTABLE, PARKER & CONSTABLE, 1987: GEOPHYSICS 52, 289-300
C
C   S. CONSTABLE, MARCH 1986, S.I.O., LA JOLLA, CA 92093, U.S.A. (VERSION
C   CONSTABLE, OCTOBER 1987, (VERSION
C
C   IF YOU OBTAIN THIS CODE FROM A THIRD PERSON, PLEASE SEND YOUR NAME AND
C   TO S. CONSTABLE. YOU WILL THEN RECEIVE UPDATES, NEWS ON BUGS, ETC.
C
C   ON INPUT:
C   /DATA/ CONTAINS
C   ND = NUMBER OF DATA
C   D(NDD) = VECTOR OF DATA VALUES
C   SD(NDD) = VECTOR OF ABSOLUTE STANDARD ERRORS IN THE DATA
C   DP(NDD,4) = MATRIX CONTAINING UP TO FOUR PARAMETERS ASSOCIATED WITH
C   EACH DATUM (RANGE, FREQUENCY, DATA TYPE ETC)
C   NPD = NUMBER OF PARAMETERS ASSOCIATED WITH EACH DATUM (E.G. FOR FR
C   NPD = 1, FOR FREQ. AND RANGE NPD = 2)
C
C   /MODEL/ CONTAINS
C   NP = NUMBER OF MODEL PARAMETERS, USUALLY THE NUMBER OF LAYERS
C   PM(NPP) = VECTOR OF INITIAL MODEL PARAMETERS, USUALLY LOG10(LAYER
C   LJMP(NPP) = A VECTOR DESCRIBING PLACES WITHIN THE MODEL TO RELAX TH
C   SMOOTHNESS CONSTRAINT, IN MONOTONIC ORDER FOLLOWED BY A ZERO.
C   THUS LJMP = 4,10,0
C   WOULD ALLOW A DISCONTINUITY BETWEEN LAYER 3 AND 4 AND LAYERS
C
C   IRUF = 1 FOR FIRST DERIVATIVE MINIMIZATION
C   = 2 FOR SECOND DERIVATIVE MINIMIZATION
C   TOLREQ = REQUIRED RMS MISFIT
C   IRUG = 1 IF A TABLE OF RMS MISFIT VERSUS LAGRANGE MULTIPLIER IS REQUI
C   (USEFUL FOR TRACKING DOWN CONVERGENCE PROBLEMS)
C   0 OTHERWISE (LOTS FASTER)
C   IOUT = FORTRAN UNIT NUMBER FOR OUTPUT, CAN EITHER BE TERMINAL OR A FI
C   UM = THE MU VALUE USED AT THE MINIMUM OF THE MISFIT IN THE LAST CALL
C   TO OCCAM1
C
C   ON OUTPUT:
C   /MODEL/ CONTAINS
C   PM(NPP) = VECTOR OF UPDATED MODEL PARAMETERS
C   DM(NDD) = VECTOR CONTAINING RESPONSE OF NEW MODEL

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INV01200
INV01210
INV01220
INV01230
INV01240
INV01250
INV01260
INV01270
INV01280
INV01290
INV01300
INV01310
INV01320
INV01330
INV01340
INV01350
INV01360
INV01370
INV01380
INV01390
INV01400
INV01410
INV01420
INV01430
INV01440
INV01450
INV01460
INV01470
INV01480
INV01490
INV01500
INV01510
INV01520
INV01530
INV01540
INV01550
INV01560
INV01570
INV01580
INV01590
INV01600
INV01610
INV01620
INV01630
INV01640
INV01650
INV01660
INV01670
INV01680
INV01690
INV01700
INV01710
INV01720
INV01730
INV01740
INV01750
INV01760
INV01770
INV01780
INV01790

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165      WK(I) = PM2(I) - PM2(I-1)
      ELSE
        DO 170 I = 2, NP-1
          WK(I) = PM2(I+1) - 2.*PM2(I) + PM2(I-1)
          WK(NP) = 0.0
        END IF
      C REMOVE ENTRIES IN ROUGHNESS VECTOR WHERE JUMPS ARE ALLOWED
      I = 1
      IF (LJMP(I) .NE. 0) THEN
        WK(LJMP(I)) = 0.0
        I = I+1
        GOTO 176
      END IF
      RUF = ANORM(NP, WK)
      C IF WE HAVE ATTAINED THE INTERCEPT BUT THE MODEL IS GETTING ROUGHER
      C WE HAVE PROBLEMS WITH CONVERGENCE.
      C SAVE NEW MODEL AND COMPUTE STEP SIZE
      DO 180 I = 1, NP
        WK(I) = PM2(I) - PM(I)
        PM(I) = PM2(I)
        STEPSZ = ANORM(NP, WK)
        STEPSZ = SQRT(STEPSZ/NP)
        WRITE(IOUT, *) 'STEPSIZE IS = ', STEPSZ
        WRITE(IOUT, *) 'ROUGHNESS IS = ', RUF
        NITER = NITER + 1
      RETURN
      END
180
C .....
C SUBROUTINE MAKDEL(IRUF)
C
C MAKDEL CONSTRUCTS TRANS(DEL).DEL FOR 1ST OR 2ND DERIVATIVE ROUGHNESS
C
C ON INPUT:
C   IRUF = 1 FOR 1ST DERIVATIVE ROUGHNESS
C         2 FOR 2ND DERIVATIVE ROUGHNESS
C
C /MODEL/ CONTAINS
C   LJMP = ARRAY OF LAYER LOCATIONS WHERE A JUMP IN RESISTIVITY IS ALLOWED
C         TERMINATED BY A ZERO
C
C IMPLICIT DOUBLE PRECISION (A-H, O-Z)
C PARAMETER (MDD = 40, NPP = 30)
C COMMON /MODEL/ NP, PM(NPP), DM(NDD), LJMP(NPP)
C COMMON /RESULT/ PTP(NPP, NPP), WJTWJ(NPP, NPP), WJTWJ(NPP), PM2(NPP),
C * PMJ, NFOR, NITER, FRAC, RLAST, IFFTOL
C DIMENSION DEL(NPP, NPP)
C
C MAKE PARTIALS MATRIX
C CALL INITM(NPP, NP, NP, DEL)
C DO 10 I = 2, NP
  DEL(I, I-1) = -1.
  DEL(I, I) = 1.
10
C REMOVE ANY ENTRIES WHERE MODEL IS ALLOWED TO JUMP
  I = 0
  I = I+1
  N = LJMP(I)
C CAN NOT REMOVE AIR/EARTH INTERFACE (N=1):
  IF (N .EQ. 1) GOTO 20
  IF (I .NE. NP+1 .AND. N.NE.0) THEN

```

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INV03000
INV03010
INV03020
INV03030
INV03040
INV03050
INV03060
INV03070
INV03080
INV03090
INV03100
INV03110
INV03120
INV03130
INV03140
INV03150
INV03160
INV03170
INV03180
INV03190
INV03200
INV03210
INV03220
INV03230
INV03240
INV03250
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INV03270
INV03280
INV03290
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INV03540
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INV03570
INV03580
INV03590

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11 FORMAT(4X,'COMPUTED DATA',/, (5(1PE14.4)))
C CALC MISFIT VECTOR AND MISFIT
DO 60 I = 1,ND
  DD(I) = (D(I) - DM(I))/SD(I)
  CHI0 = ANORM(ND,DD)
  TOL0 = SQRT(CHI0/ND)
  IF (NITER.EQ.0) THEN
    WRITE(IOUT,*) 'STARTING R.M.S. = ',TOL0
  END IF
C WEIGHT THE PARTIALS MATRIX
DO 70 I = 1,ND
  DO 70 J = 1,NP
    WJ(I,J) = WJ(I,J)/SD(I)
C FORM W.J.TRANS(W.J)
CALL TRMULT(ND,ND,NP,NP,WJ,WJ,WJTWJ)
C FORM THE WEIGHTED, TRANSLATED DATA AND PREMULTIPLY BY TRANS(W.J)
CALL MULT(ND,ND,NP,NP,1,WJ,PM,DHAT)
DO 80 I = 1,ND
  DHAT(I) = DD(I) + DHAT(I)
  CALL TRMULT(ND,ND,NP,NP,1,WJ,DHAT,WJTD)
C PRODUCE THE MISFIT FUNCTION IF REQUIRED
  IF (IBUG.EQ.1) THEN
    WRITE(IOUT,*) 'MISFIT AS A FUNCTION OF MU'
    DO 100 K = 18,1,-1
      AMU = 10.**((FLOAT(K)/2. - 1.))
      T = TOFMU(AMU)
      IF (T.GE.1.0E+09) GOTO 110
    WRITE(IOUT,*) AMU,T
  END IF
100 CONTINUE
  WRITE(IOUT,*) ' '
  WRITE(IOUT,*) ' ** ITERATION ',NITER+1,' **'
C THE NEXT BLOCK OF CODE CONTROLS THE SELECTION OF THE LAGRANGE MULTIPLIER
C PMU
120 TOLM = 0.2
  AMU1 = 0.9*PMU
  AMU2 = PMU
  NFOR = 0
  IFFTOL = 0
C FIND MINIMUM
  CALL MINBRK(AMU1,AMU2,AMU3,TAMU1,TAMU2,TAMU3,TOFMU)
  WRITE(IOUT,*) 'AMU1',AMU1,'MISFIT',TAMU1
  WRITE(IOUT,*) 'AMU2',AMU2,'MISFIT',TAMU2
  WRITE(IOUT,*) 'AMU3',AMU3,'MISFIT',TAMU3
  TOBT = FMIN(AMU1,AMU2,AMU3,TAMU2,TOFMU,TOLM,PMU)
  PMU = AMU2
  TOBT = TAMU2
C COMPUTE THE MODEL PARAMETERS (AGAIN) AT WHICH MINIMUM MISFIT WAS
  C OBTAINED. TORBDM IS DUMMY
  TORBDM=TOFMU(PMU)
  WRITE(IOUT,*) 'MINIMUM TOL IS AT MU = ',PMU
  WRITE(IOUT,*) 'AND IS = ,TOBT
  WRITE(IOUT,*) 'USING ',NFOR,' EVALUATIONS OF FUNCTION'
C IF THE NEW MINIMUM TOLERANCE IS GREATER THAN THE TOLERANCE FROM THE
C PREVIOUS MODEL, WE ARE HAVING CONVERGENCE PROBLEMS.
C **FIND LAGRANGE MULTIPLIER SELECTION
C COMPUTE ROUGHNESS
  WK(1) = 0.0
  IF (IRUF.EQ.1) THEN
    DO 165 I = 2,NP

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INV02400
INV02410
INV02420
INV02430
INV02440
INV02450
INV02460
INV02470
INV02480
INV02490
INV02500
INV02510
INV02520
INV02530
INV02540
INV02550
INV02560
INV02570
INV02580
INV02590
INV02600
INV02610
INV02620
INV02630
INV02640
INV02650
INV02660
INV02670
INV02680
INV02690
INV02700
INV02710
INV02720
INV02730
INV02740
INV02750
INV02760
INV02770
INV02780
INV02790
INV02800
INV02810
INV02820
INV02830
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INV02850
INV02860
INV02870
INV02880
INV02890
INV02900
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INV02920
INV02930
INV02940
INV02950
INV02960
INV02970
INV02980
INV02990

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PD (I) = PM(I)
CVEC(I) = WJTD(I)
CONTINUE
CALL QUAD(PD,NPP,NPP, NP,HESS,CVEC)
DO 15 I=1,NP
PM2(I) = PD(I)
C CUT SIZE IF NECESSARY
IF (FRAC.LT. 0.999) THEN
DO 30 I = 1,NP
PM2(I) = PM(I) + FRAC*(PM2(I)-PM(I))
END IF
C CALC. MODEL RESPONSE AND MISFIT
CALL FORMOD(NP,ND,PM2,DM)
DO 40 I = 1,ND
WK(I) = (D(I) - DM(I))/SD(I)
CHI = ANORM(ND,WK)
TOFMU = Sqrt(CHI/ND)
RETURN
END
C
SUBROUTINE QUAD(PD,NROWH,NCOLH,N,HESS,CVEC)
IMPLICIT REAL*8 (A-H,O-Z)
PARAMETER (ITMAX=50,MSGLVL=1,NROWA=1,
1 BIGND=1.E+10,LIWORK=60,LWORK=1980,NCLIN=0,NPP=50)
REAL*8 A(1,NPP),BL(NPP),BU(NPP),CVEC(N),FEATOL(NPP)
HESS(NROWH,NCOLH),CLAMDA(NPP),WORK(LWORK),IWORK(LIWORK),
2 ISTATE(NPP),PD(N)
LOGICAL COLD,LP,ORTHOG
EXTERNAL QPHESS
DATA NOUT/7/
CALL X04ABF(1,NOUT)
COLD=.TRUE.
LP=.FALSE.
ORTHOG=.FALSE.
NCTOTL=N+NCLIN
DO 10 I=1,N
BL(I) = 0.
BU(I) = 20.
FEATOL(I) = 0.02
10 CONTINUE
BU(1) = 0.
BU(N) = 0.
IFAIL = 0
C QUADRATIC PROGRAMMING
C CALL E04NAF(ITMAX,MSGLVL,N,NCLIN,NCTOTL,NROWA,NROWH,NCOLH,
C 1 BIGND,A,BL,BU,CVEC,FEATOL,HESS,QPHESS,
C 2 COLD,LP,ORTHOG,PD,ISTATE,ITER,DEJ,
C 3 CLAMDA,IWORK,LIWORK,WORK,LWORK,IFAIL)
C CALL VE04A(N,HESS,NROWH,CVEC,BL,BU,PD,Q,IWORK,K,WORK)
IF (IFAIL.GT.0) THEN
WRITE(4,20) IFAIL
STOP
ENDIF
20 FORMAT('E04NAF TERMINATED WITH IFAIL = ',I3)
RETURN
END
SUBROUTINE QPHESS(N,NROWH,NCOLH,JTHCOL,HESS,X,HX)
IMPLICIT REAL*8 (A-H,O-Z)
INTEGER JTHCOL,N,NCOLH,NROWH
REAL*8 HESS(NROWH,NCOLH),HX(N),X(N)

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INV04200
INV04210
INV04220
INV04230
INV04240
INV04250
INV04260
INV04270
INV04280
INV04290
INV04300
INV04310
INV04320
INV04330
INV04340
INV04350
INV04360
INV04370
INV04380
INV04390
INV04400
INV04410
INV04420
INV04430
INV04440
INV04450
INV04460
INV04470
INV04480
INV04490
INV04500
INV04510
INV04520
INV04530
INV04540
INV04550
INV04560
INV04570
INV04580
INV04590
INV04600
INV04610

INV04620
INV04630
INV04640
INV04650
INV04660
INV04670
INV04680
INV04690
INV04700
INV04710
INV04720
INV04730
INV04740
INV04750
INV04760
INV04770
INV04780

```

DEL(N,N-1) = 0.
DEL(N,N) = 0.
GOTO 20
END IF
C SQUARE MATRIX AND COPY BACK INTO DEL IF SECOND DERIV. SMOOTHING REQUIRED
CALL MULT(NPP,NP,NPP,NP,DEL,DEL,PTP)
DO 30 I = 1,NP
  DO 30 J = 1,NP
    DEL(J,I) = PTP(J,I)
  END IF
END IF
C FORM TRANS(DEL).DEL
CALL TRMULT(NPP,NP,NPP,NP,DEL,DEL,PTP)
RETURN
END
C
BLOCKDATA BLOCK
IMPLICIT DOUBLE PRECISION (A-H, O-Z)
PARAMETER (NDD = 40, NPP = 30)
COMMON /RESULT/ PTP(NPP,NPP), WJTWJ(NPP,NPP), WJTWJ(NPP), PM2(NPP),
* PMU,NFOR,NITER,FRAC,RLAST,IFFTOL
DATA NITER /0/
END
C
C*****
FUNCTION TOFMU(AMU)
C
C FUNCTION TOFMU RETURNED THE RMS MISFIT OF THE RESPONSE OF A MODEL CONSTANT
C USING THE GIVEN VALUE OF LAGRANGE MULTIPLIER
C S. CONSTABLE, MARCH 1988, S.I.O., LA JULLA, CA 92093, U.S.A.
C
C ON INPUT:
C AMU = LAGRANGE MULTIPLIER
C THE MODEL IS ALSO A FUNCTION OF THE ARRAYS PASSED IN COMMON BLOCK /RES/
C ON OUTPUT
C TOFMU = R.M.S. MISFIT OR 1.0E+10 IF CHOLESKY DECOMPOSITION FAILED
C /MODEL/ CONTAINS THE MODEL WHICH PRODUCES THE MISFIT TOFMU, AS WELL
C AS ITS RESPONSE
C SUBROUTINES USED:
C CHOLIN, CHOLSL ARE R.L. PARKER'S SUBROUTINES TO PERFORM CHOLESKY DE
C AND THEN SOLVE A LINEAR SYSTEM BY BACK SUBSTITUTION.
C ANORN, FORMOD ARE EXPLAINED IN OCCAM1
C
IMPLICIT DOUBLE PRECISION (A-H, O-Z)
PARAMETER (NDD = 40, NPP = 30)
COMMON /DATA/ ND,D(NDD),SD(NDD)
COMMON /MODEL/ NP,PM(NPP),DM(NDD),LJMP(NPP)
COMMON /RESULT/ PTP(NPP,NPP),WJTW(NPP,NPP),WJTW(NPP),PM2(NPP),
* PMU,NFOR,NITER,FRAC,RLAST,IFFTOL
DIMENSION WK(NDD)
REAL*8 PD(NPP),HESS(NPP,NPP),CVEC(NPP)
NFOR = NFOR + 1
C ADD AMU,TRANS(PARTIAL).PARTIAL TO TRANS(W.J).W.J
DO 20 I = 1,NP
  DO 20 J = 1,NP
    HESS(I,J) = AMU*PTP(I,J) + WJTWJ(I,J)
  END DO
DO 10 I = 1,NP

```

INV03600
INV03610
INV03620
INV03630
INV03640
INV03650
INV03660
INV03670
INV03680
INV03690
INV03700
INV03710
INV03720
INV03730
INV03740
INV03750
INV03760
INV03770
INV03780
INV03790
INV03800
INV03810
INV03820
INV03830
INV03840
INV03850
INV03860
INV03870

INV03880
INV03890
INV03900
INV03910
INV03920
INV03930
INV03940
INV03950
INV03960
INV03970
INV03980
INV03990
INV04000
INV04010
INV04020
INV04030
INV04040
INV04050
INV04060
INV04070
INV04080
INV04090
INV04100
INV04110
INV04120
INV04130
INV04140
INV04150
INV04160
INV04170
INV04180
INV04190

```

12 CONTINUE
RETURN
END
C
C
C
C.....
SUBROUTINE TRMULT(MAD,MA,NAD,NA,NB,A,B,C)
C MULTIPLIES THE TRANSPOSE OF A 2D MATRIX BY ANOTHER MATRIX
IMPLICIT DOUBLE PRECISION (A-H, O-Z)
DIMENSION A(MAD,NA),B(MAD,NB),C(NAD,NB)
DO 12 I = 1,NA
DO 12 J = 1,NB
C(IJ) = 0.0
DO 10 K = 1,MA
C(IJ) = A(K,I)*B(K,J) + C(IJ)
END
RETURN
END
C
C
C.....
SUBROUTINE MINBRK(AX,BX,CX,FA,FB,FC,FUNC)
C MINBRK BRACKETS A UNIVARIATE MINIMUM OF A FUNCTION. TO BE USED PRIOR
C A UNIVARIATE MINIMISATION ROUTINE.
C
C ON INPUT:
AX, BX = TWO DISTINCT ESTIMATES OF THE MINIMUM'S WHEREABOUTS
FUNC = THE FUNCTION IN QUESTION
C ON OUTPUT:
AX, BX, CX = THREE NEW POINTS WHICH BRACKET THE MINIMUM
FA, FB, FC = FUNC(AX), FUNC(BX) ETC.
C
C
C.....
IMPLICIT DOUBLE PRECISION (A - H, O - Z)
PARAMETER (GOLD=1.618034, GLIMIT=100., TINY=1.E-21)
FA = FUNC(AX)
FB = FUNC(BX)
IF (FR.GT.FA) THEN
DUM = AX
AX = BX
BX = DUM
DUM = FB
FB = FA
FA = DUM
ENDIF
CX = BX + GOLD*(BX - AX)
FC = FUNC(CX)
IF (FB.GE.FC) THEN
R = (BX - AX)*(FB - FC)
Q = (BX - CX)*(FB - FA)
U = BX - ((BX - CX)*Q - (BX - AX)*R) / (2.*SIGN(MAX(ABS(Q - R), TINY), Q - R))
ULIM = BX + GLIMIT*(CX - BX)
IF ((BX - U)*(U - CX).GT.0.) THEN
FU = FUNC(U)
IF (FU.LT.FC) THEN
AX = BX
FA = FB
BX = U

```

```

INV05390
INV05400
INV05410
INV05420
INV05430
INV05440
INV05450
INV05460
INV05470
INV05480
INV05490
INV05500
INV05510
INV05520
INV05530
INV05540
INV05550
INV05560
INV05570
INV05580
INV05590
INV05600
INV05610
INV05620
INV05630
INV05640
INV05650
INV05660
INV05670
INV05680
INV05690
INV05700
INV05710
INV05720
INV05730
INV05740
INV05750
INV05760
INV05770
INV05780
INV05790
INV05800
INV05810
INV05820
INV05830
INV05840
INV05850
INV05860
INV05870
INV05880
INV05890
INV05900
INV05910
INV05920
INV05930
INV05940
INV05950
INV05960
INV05970
INV05980

```

```

INV04790
INV04800
INV04810
INV04820
INV04830
INV04840
INV04850
INV04860
INV04870
INV04880
INV04890
INV04900
INV04910
INV04920
INV04930
INV04940
INV04950
INV04960
INV04970
INV04980
INV04990
INV05000
INV05010
INV05020
INV05030
INV05040
INV05050
INV05060
INV05070
INV05080
INV05090
INV05100
INV05110
INV05120
INV05130
INV05140
INV05150
INV05160
INV05170
INV05180
INV05190
INV05200
INV05210
INV05220
INV05230
INV05240
INV05250
INV05260
INV05270
INV05280
INV05290
INV05300
INV05310
INV05320
INV05330
INV05340
INV05350
INV05360
INV05370
INV05380

```

```

        C.....
        FUNCTION ANORM(N,D)
        C RETURNS THE SQUARE OF THE EUCLIDEAN NORM OF A VECTOR
        IMPLICIT DOUBLE PRECISION (A-H, O-Z)
        DIMENSION D(N)
        ANORM = 0.f
        DO 10 I = 1,N
            ANORM = ANORM + D(I)*D(I)
        CONTINUE
        RETURN
    END

```

```

10      C.....
        SUBROUTINE INITM(MD,M,N,A)
        C ZEROS A 2D MATRIX
        IMPLICIT DOUBLE PRECISION (A-H, O-Z)
        DIMENSION A(MD,N)
        DO 10 I = 1,M
            DO 10 J = 1,N
                A(I,J) = 0.0
            CONTINUE
        END
    END

```

```

        C.....
        SUBROUTINE MULT(MAD,MA,NAD,NA,NB,A,B,C)
        C MULTIPLIES TWO 2D MATRICES
        IMPLICIT DOUBLE PRECISION (A-H, O-Z)
        DIMENSION A(MAD,NA),B(NAD,NB),C(MAD,NB)
        DO 12 I = 1,MA
            DO 12 J = 1,NB
                C(I,J) = 0.0
                DO 10 K = 1,NA
                    C(I,J) = A(I,K)*B(K,J) + C(I,J)
                CONTINUE
            CONTINUE
        END
    END

```

```

        C.....
        SUBROUTINE DIMULT(MAD,MA,NA,DIAG,A,B)
        C MULTIPLIES A 2D MATRIX BY A DIAGONAL MATRIX
        IMPLICIT DOUBLE PRECISION (A-H, O-Z)
        DIMENSION DIAG(MA),A(MAD,NA),B(MAD,NA)
        DO 12 J = 1,NA
            DO 11 I = 1,MA
                B(I,J) = DIAG(I)*A(I,J)
            CONTINUE
        END
    END

```

INV066590
 INV066600
 INV066610
 INV066620
 INV066630
 INV066640
 INV066650
 INV066660
 INV066670
 INV066680
 INV066690
 INV066700
 INV066710
 INV066720
 INV066730
 INV066740
 INV066750
 INV066760
 INV066770
 INV066780
 INV066790
 INV066800
 INV066810
 INV066820
 INV066830
 INV066840
 INV066850
 INV066860
 INV066870
 INV066880
 INV066890
 INV066900
 INV066910
 INV066920
 INV066930
 INV066940
 INV066950
 INV066960
 INV066970
 INV066980
 INV066990
 INV070000
 INV070010
 INV070020
 INV070030
 INV070040
 INV070050
 INV070060
 INV070070
 INV070080
 INV070090
 INV071000
 INV071010
 INV071020
 INV071030
 INV071040
 INV071050
 INV071060
 INV071070
 INV071080

```

V = BX
W = V
X = V
E = 0
FX = FBX
FV = FX
FW = FX
DO 11 ITER = 1, ITMAX
  XM = 0.5*(A + B)
  TOL1 = TOL*ABS(X) + ZEPS
  TOL2 = 2.*TOL1
  IF(ABS(X - XM).LE.(TOL2 - .5*(B - A))) GOTO 3
  IF(ABS(E).GT.TOL1) THEN
    R = (X - W)*(FX - FV)
    Q = (X - V)*(FX - FW)
    P = (X - V)*Q - (X - W)*R
    Q = 2.*(Q - R)
    IF(Q.GT.0.) P = - P
    Q = ABS(Q)
    ETEMP = E
    E = D
  IF(ABS(P).GE.ABS(.5*Q*ETEMP).OR.P.LE.Q*(A - X).OR.
    P.GE.Q*(B - X)) GOTO 1
  D = P/Q
  U = X + D
  IF(U - A.LT.TOL2 .OR. B - U.LT.TOL2) D = SIGN(TOL1,XM - X)
  GOTO 2
ENDIF
IF(X.GE.XM) THEN
  E = A - X
ELSE
  E = B - X
ENDIF
D = CGOLD*E
IF(ABS(D).GE.TOL1) THEN
  U = X + D
ELSE
  U = X + SIGN(TOL1,D)
ENDIF
FU = F(U)
IF(FU.LE.FX) THEN
  IF(U.GE.X) THEN
    A = X
  ELSE
    B = X
  ENDIF
  V = W
  FV = FW
  W = X
  FW = FX
  X = U
  FX = FU
ELSE
  IF(U.LT.X) THEN
    A = U
  ELSE
    B = U
  ENDIF
  IF(FU.IE.FW .OR. W.EQ.X) THEN
    V = W
  
```

```

FB = FU
GO TO 1
ELSE IF (FU.GT.FB) THEN
CX = U
FC = FU
GO TO 1
ENDIF
U = CX + GOLD*(CX - BX)
FU = FUNC(U)
ELSE IF ((CX - U)*(U - ULM) .GT. 0.) THEN
FU = FUNC(U)
IF (FU.LT.FC) THEN
BX = CX
CX = U
U = CX + GOLD*(CX - BX)
FB = FC
FC = FU
FU = FUNC(U)
ENDIF
ELSE IF ((U - ULM)*(ULM - CX) .GE. 0.) THEN
U = ULM
FU = FUNC(U)
ELSE
U = CX + GOLD*(CX - BX)
FU = FUNC(U)
ENDIF
AX = BX
BX = CX
CX = U
FA = FB
FB = FC
FC = FU
GO TO 1
ENDIF
IF (AX.LT.0.) AX=0.
IF (CX.LT.0.) CX=0.
RETURN
END

```

```

INV065990
INV066000
INV066010
INV066020
INV066030
INV066040
INV066050
INV066060
INV066070
INV066080
INV066090
INV066100
INV066110
INV066120
INV066130
INV066140
INV066150
INV066160
INV066170
INV066180
INV066190
INV066200
INV066210
INV066220
INV066230
INV066240
INV066250
INV066260
INV066270
INV066280
INV066290
INV066300
INV066310
INV066320
INV066330
INV066340
INV066350
INV066360
INV066370
INV066380
INV066390
INV066400
INV066410
INV066420
INV066430
INV066440
INV066450
INV066460
INV066470
INV066480
INV066490
INV066500
INV066510
INV066520
INV066530
INV066540
INV066550
INV066560
INV066570
INV066580

```

```

C .....
C FUNCTION FMIN(AX,BX,CX,FBX,F,TOL,XMIN)
C FMIN RETURNS THE MINIMUM VALUE OF A FUNCTION
C
C ON INPUT:
C AX,BX,CX = INDEPENDENT VARIABLE WHICH BRACKET THE MINIMUM
C FBX = F(BX) (USUALLY AVAILABLE FROM THE BRACKETING PROCEDURE)
C F = FUNCTION IN QUESTION
C TOL = FRACTIONAL TOLERANCE REQUIRED IN THE INDEPENDENT VARIABLE
C ON OUTPUT:
C XMIN = ABSCISSA OF MINIMUM
C FMIN = F(XMIN)
C
C
C IMPLICIT DOUBLE PRECISION (A - H, O - Z)
C PARAMETER (ITMAX=100,CGOLD=.3819888,ZEPS=1.0E - 10)
C A = MIN(AX,CX)
C B = MAX(AX,CX)

```

```

CALL CUBDER(NPT+2,XDEL,ZSURF,XDER)
DO 10 I=NXT,1,-1
IF(I.GE.NE.OR.I.LT.NST) THEN
ZSURF(I)=0.0
XDER(I)=0.
AMPDER(I)=1.
GOTO 10
ENDIF
ZSURF(I)=ZSURF(I-NST+1)
XDER(I)=XDER(I-NST+1)
AMPDER(I)=SQRT(1.0+XDER(I)*XDER(I))
CONTINUE
10  WRITE(4,60)(ZSURF(I),I=1,NXT)
20  WRITE(4,60)(XDER(I),I=1,NXT)
30  WRITE(4,70)(AMPDER(I),I=1,NXT)
50  FORMAT(4X,'ZSURF',/,4X,(10F7.3))
60  FORMAT(4X,'XDER',/,4X,(10F7.3))
70  FORMAT(4X,'AMPDER',/,4X,(10F7.3))
RETURN
END
C
C
C SUBROUTINE HNPRIM(K,KS,ZSURF,XDER,AMPDER,XSYS,ZSYS,NTR)
C
C COMPUTE THE NORMAL COMPONENT OF THE MAGNETIC FIELD
C
C .....
C IMPLICIT DOUBLE PRECISION (A-H, O-Z)
C PARAMETER (PI=3.1415926)
C DIMENSION ZSURF(1),XDER(1),AMPDER(1),XSYS(1),ZSYS(1)
C DIMENSION HPN(120,20)
C COMMON /KY/ AKY(20)/HN/HPN
C COMMON /KYE/ NSYS,NST,XDEL,TRDIS
C PI2=PI*2.
DO 10 J=KS,NTR+KS-1
M=J-KS+1
XD=(J-1)*XDEL-XSYS(K)+TRDIS/2.
ZD=ZSURF(J)-ZSYS(K)
RHO=SQRT(XD*XD+ZD*ZD)
D=RHO*RHO*AMPDER(J)*PI2
IF(D.EQ.0.) PAUSE 'BAD HPN'
IF(NSYS.EQ.1) THEN
GEOM1=(-XDER(J)*XD+ZD)*XD/D
GEOM2=(-(XD*XD-ZD*ZD)*XDER(J)+2.*XD*ZD)/D
ELSE
GEOM1=(-XDER(J)*XD+ZD)*ZD/D
GEOM2=((ZD*ZD-XD*XD)-XDER(J)*2.*XD*ZD)/D
ENDIF
DO 20 I=1,KYE
WK=AKY(I)*PI2
RKY=RHO*WK
IF(I.GT.1) THEN
HPN(M,I)=(GEOM1*WK*BESSK0(RKY)+GEOM2*BESSK1(RKY)
/RHO)*WK
ELSE
HPN(M,I)=GEOM2/(RHO*RHO)
ENDIF
20 CONTINUE
10 CONTINUE
WRITE(4,30)(HPN(J,1),J=1,NTR)
C 30 FORMAT(4X,'HPN',/,5(1PE12.4))

```

```

INV07790
INV07800
INV07810
INV07820
INV07830
INV07840
INV07850
INV07860
INV07870
INV07880
INV07890
INV07900
INV07910
INV07920
INV07930
INV07940
INV07950
INV07960
INV07970
INV07980
INV07990
INV08000
INV08010
INV08020
INV08030
INV08040
INV08050
INV08060
INV08070
INV08080
INV08090
INV08100
INV08110
INV08120
INV08130
INV08140
INV08150
INV08160
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INV08180
INV08190
INV08200
INV08210
INV08220
INV08230
INV08240
INV08250
INV08260
INV08270
INV08280
INV08290
INV08300
INV08310
INV08320
INV08330
INV08340
INV08350
INV08360
INV08370
INV08380

```

```

11  FV = FW
    W = U
    FW = FU
    ELSE IF(FU.LE.FV .OR. V.EQ.X .OR. V.EQ.W) THEN
      V = U
      FV = FU
    ENDIF
    ENDIF
    CONTINUE
3  WRITE(*,*) 'MAXIMUM ITERATIONS EXCEEDED IN FMIN'
    XMIN = X
    FMIN = FX
    RETURN
    END
C
C
C .....
SUBROUTINE FORMOD(NPT,NRESP,ZHI,CALC)
C
C ... PURPOSE ...
C ... COMPUTE THE AEM SYSTEM RESPONSE OVER A 2-D ICE-WATER PROFILE
C
C IMPLICIT DOUBLE PRECISION (A-H, O-Z)
C PARAMETER (NDD=40,NPP=30,NXT=200,NTRUNK=100,NLOOP=8)
C DIMENSION ZSURF(NXT),XDER(NXT),AMPDER(NXT),HPN(120,20)
C DIMENSION AKERN(120,120,20),CHAR(120,20)
C DIMENSION XSYS(NDD),ZSYS(NDD),ZHI(NPP),CALC(NDD)
C COMMON /KY/ AKY(20)/KER/ AKERN/CH/CHAR/HPN
C COMMON /FCM/ZSURF,XSYS,ZSYS,AMPDER,NM,IREC,AKDEL
C COMMON KYE,NSYS,NST,XDEL,TRDIS
C DO 1 I=1,NPT
1  ZSURF(I)=ZHI(I)
    CALL SURF(NST,NPT,NXT,XDEL,ZSURF,XDER,AMPDER)
    K=1
1001 CONTINUE
    KS=(XSYS(K)-TRDIS/2.)/XDEL*2-NTRUNK/2
    IF(KS.LE.0 .OR. KS.GE.NXT-NTRUNK/2) PAUSE 'BAD INPUT XSYS'
    CALL HNPTRIM(K,KS,ZSURF,XDER,AMPDER,XSYS,ZSYS,NTRUNK)
    CALL KERNEL(K,KS-1,KS1,ZSURF,XDER,AMPDER,NTRUNK)
    CALL CHARGE(NTRUNK,KYE,XDEL,NLOOP)
    CALL FIELD(K,KS,NM,AKDEL,HXS,HZS,XSYS,ZSYS,ZSURF,AMPDER,NTRUNK)
    IF(IREC.EQ.1) CALC(K)=HXS
    IF(IREC.EQ.2) CALC(K)=HZS
    K=K+1
    KS1=KS
    IF(K.LE.NRESP) GOTO 1001
    RETURN
    END
C
C SUBROUTINE SURF(NST,NPT,NXT,XDEL,ZSURF,XDER,AMPDER)
C
C GENERATE THE INFORMATION OF THE ICE-WATER INTERFACE, I.E.,
C ZSURF, XDER, AMPDER
C
C IMPLICIT DOUBLE PRECISION (A-H, O-Z)
C DIMENSION ZSURF(NXT),XDER(NXT),AMPDER(NXT)
C NE=NST+NPT

```

```

INV07190
INV07200
INV07210
INV07220
INV07230
INV07240
INV07250
INV07260
INV07270
INV07280
INV07290
INV07300
INV07310
INV07320
INV07330
INV07340
INV07350
INV07360
INV07370
INV07380
INV07390
INV07400
INV07410
INV07420
INV07430
INV07440
INV07450
INV07460
INV07470
INV07480
INV07490
INV07500
INV07510
INV07520
INV07530
INV07540
INV07550
INV07560
INV07560
INV07570
INV07580
INV07590
INV07600
INV07610
INV07620
INV07630
INV07640
INV07650
INV07660
INV07670
INV07680
INV07690
INV07700
INV07710
INV07720
INV07730
INV07740
INV07750
INV07760
INV07770
INV07780

```



```

RETURN
END

SUBROUTINE KERNEL(K,KS,KSI,ZSURF,XDER,AMPDER,NTR)
.....
COMPUTE THE KERNEL IN THE INTEGRAL EQUATION
AND SAVE IT
.....
IMPLICIT DOUBLE PRECISION (A-H, O-Z)
PARAMETER (PI=3.1415928)
DIMENSION ZSURF(1),XDER(1),AMPDER(1)
DIMENSION AKERN(120,120,20)
COMMON /KER/ AKERN/KY/ AKY(20)
COMMON KYE, NSYS, NST, XDEL, TRDIS
PI2=PI*2.
DO 1 J=1,NTR
DO 1 I=1,KYE
DO 1 I=1,NTR
AKERN(J,J,I)=0.0
1 CONTINUE
IF (K.GT.1) GO TO 30
C COMPUTE THE AKERNEL
1001 CONTINUE
DO 20 J=1,NTR
DO 20 L=1,NTR
IF (J.EQ.L) GO TO 200
XD=(J-L)*XDEL
ZD=ZSURF(J+KS)-ZSURF(L+KS)
RHO=SQRT(XD*XD+ZD*ZD)
GEOM=2.*(AMPDER(L+KS)/AMPDER(J+KS))*(XD*XD+(J+KS)-ZD)/RHO
DO 10 I=1,KYE
WK=AKY(I)*PI2
IF (I.EQ.1) THEN
AKERN(J,L,I)=
GEOM/RHO
ELSE
AKERN(J,L,I)=
GEOM*WK*BESSK1(RHO*WK)
ENDIF
CONTINUE
20 CONTINUE
RETURN
C
C
C COMPUTE THE UNKNOWN PART OF THE KERNEL
CONTINUE
M=KS-KSI+1
KE2=NTR-M
IF (M.LE.0) PAUSE 'BAD INPUT XSYS'
IF (M.GE.NTR) GO TO 1001
DO 40 J=1,KE2
DO 40 L=1,KE2
DO 40 I=1,KYE
AKERN(J,L,I)=AKERN(J+M,L+M,I)
40 CONTINUE
DO 80 J=1,KE2
DO 60 L=KE2+1,NTR
XD=(J-L)*XDEL
ZD=ZSURF(J+KS)-ZSURF(L+KS)
RHO=SQRT(XD*XD+ZD*ZD)
GEOM=2.*(AMPDER(L+KS)/AMPDER(J+KS))*(XD*XD+(J+KS)-ZD)/RHO
DO 50 I=1,KYE

```

```

INV08390
INV08400
INV08410
INV08420
INV08430
INV08440
INV08450
INV08460
INV08470
INV08480
INV08490
INV08500
INV08510
INV08520
INV08530
INV08540
INV08550
INV08560
INV08570
INV08580
INV08590
INV08600
INV08610
INV08620
INV08630
INV08640
INV08650
INV08660
INV08670
INV08680
INV08690
INV08700
INV08710
INV08720
INV08730
INV08740
INV08750
INV08760
INV08770
INV08780
INV08790
INV08800
INV08810
INV08820
INV08830
INV08840
INV08850
INV08860
INV08870
INV08880
INV08890
INV08900
INV08910
INV08920
INV08930
INV08940
INV08950
INV08960
INV08970
INV08980

```

```

RETURN
END
C
C
SUBROUTINE CUBDER(N,XDEL,YA,YD)
C
C . . . COMPUTE THE DERIVATIVES AT EACH NODE USING
C . . . CUBIC SPLINE.
C . . . NOTE: N=NPT*2 . ONE POINT ADDED TO EACH END.
C
C
11     IMPLICIT DOUBLE PRECISION (A-H, O-Z)
11     DIMENSION XA(50), YA(N), Y2(50), YD(N)
11     DO 11 I=1, N
11       XA(I)=(I-1)*XDEL
11       DO 12 I=N-1, 2, -1
11         YA(I)=YA(I-1)
11       YA(1)=0.
11       YA(N)=0.
11     CALL SPLINE(XA, YA, N, 0., 0., Y2)
11     DO 10 I=2, N-1
11       XX=XA(I)
11       KLO=I
11       KHI=I+1
11       H=XA(KHI)-XA(KLO)
11       A=(YA(KHI)-XX)/H
11       B=(XX-XA(KLO))/H
11       XD=XA(KHI)-XA(KLO)
11       YD(I-1)=(YA(KHI)-YA(KLO))/XD-XD*((3.*A-A-1.)*Y2(KLO)
11         + (3.*B*B-1.)*Y2(KHI))/6.
10     CONTINUE
14     DO 14 I=1, N-2
14       YA(I)=YA(I+1)
14     RETURN
14     END
C
SUBROUTINE CUBSPL(KYE, NK1, AKDEL, AKY, YA, 0)
C
C . . . CUBIC SPLINE INTERPOLATION BEFORE TAKING THE
C . . . INVERSE FOURIER TRANSFORM.
C
C
1     IMPLICIT DOUBLE PRECISION (A-H, O-Z)
1     DIMENSION AKY(KYE), 0(NK1), YA(KYE), Y2(20)
1     CALL SPLINE(AKY, YA, KYE, 1.E30, 1.E30, Y2)
1     0(1)=YA(1)
1     DO 10 I=2, NK1
1       XX=(I-1)*AKDEL
1       KLO=I
1       KHI=KYE
1       IF (KHI-KLO.GT.1) THEN
1         K=(KHI+KLO)/2
1         IF (AKY(K).GT.XX) THEN
1           KHI=K
1         ELSE
1           KLO=K
1         ENDIF
1       GOTO 1
1     ENDIF
1     H=AKY(KHI)-AKY(KLO)

```

```

INV10190
INV10200
INV10210
INV10220
INV10230
INV10240
INV10250
INV10260
INV10270
INV10280
INV10290
INV10300
INV10310
INV10320
INV10330
INV10340
INV10350
INV10360
INV10370
INV10380
INV10390
INV10400
INV10410
INV10420
INV10430
INV10440
INV10450
INV10460
INV10470
INV10480
INV10490
INV10500
INV10510
INV10520
INV10530
INV10540
INV10550
INV10560
INV10570
INV10580
INV10590
INV10600
INV10610
INV10620
INV10630
INV10640
INV10650
INV10660
INV10670
INV10680
INV10690
INV10700
INV10710
INV10720
INV10730
INV10740
INV10750
INV10760
INV10770
INV10780

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```

C      22  IF (I.LE.KYE) GOTO 1001
        WRITE(4,22) (CHAR(J,1),J=1,NTR)
        FORMAT(4X,'CHARGE',/, (5E14.4))
        RETURN
        END
C
C
C      SUBROUTINE FIELD(K,KS,NM,AKDEL,HXS,HZS,XSYS,ZSYS,ZSURF,AMPDER,NTR)
C      . . . . .
C      . . . . . COMPUTE THE SECONDARY MAGNETIC FIELD, BOTH
C      . . . . . X AND Z COMPONENTS, EXPRESSED IN PPM OF THE
C      . . . . . RECEIVED PRIMARY FIELD.
C      . . . . .
C      IMPLICIT DOUBLE PRECISION (A-H, 0-Z)
C      PARAMETER (PI=3.1415926)
C      DIMENSION AMPDER(1),CHAR(120,20),ZSURF(1),XSYS(1),ZSYS(1),AA(513)
C      DIMENSION TEMP1(20),TEMP2(20)
C      COMMON /CH/CHAR/KY/AKY(20)
C      COMMON KYE,NSYS,NST,XDEL,TRDIS
C      R = TRDIS
C      XF=R/2.
C      IF (NSYS.EQ.1) THEN
C          HNORM=1./(2.*PI*R**3)
C      ELSE
C          HNORM=-1./(4.*PI*R**3)
C      ENDIF
C      DO 10 I=1,KYE
C          WK=AKY(I)*PI*2.
C          AZ=0.0
C          AX=0.0
C          DO 20 J=KS,NTR+KS-2
C              M=J-KS+1
C              RHO=SQRT((XSYS(K)-(J-1)*XDEL+XF)**2+(ZSYS(K)-ZSURF(J))**2)
C              IF (I.GT.1) THEN
C                  AXZ=2.0*CHAR(M,I)*AMPDER(J)*WK*BESSK1(RHO*WK)/RHO
C              ELSE
C                  AXZ=2.0*CHAR(M,I)*AMPDER(J)/(RHO*RHO)
C              ENDIF
C          AZ=(ZSYS(K)-ZSURF(J))*AXZ+AZ
C          AX=(XSYS(K)-(J-1)*XDEL+XF)*AXZ+AX
C      CONTINUE
C      TEMP1(I)=AZ*XDEL
C      TEMP2(I)=AX*XDEL
C      CONTINUE
C      NK = 2**N'I
C      IF (NK*AKDEL.GT. AKY(KYE)) PAUSE 'BAD AKDEL OR AKY'
C      CALL CUB_PL(KYE,NK,AKDEL,AKY,TEMP1,AA)
C      SUM=0.0
C      DO 11 I=2,NK
C          SUM=SUM+AA(I)
C      HZS=(SUM*AA(1)/2.)*AKDEL*2.
C      CALL CUB_PL(KYE,NK,AKDEL,AKY,TEMP2,AA)
C      SUM=0.0
C      DO 12 I=2,NK
C          SUM=SUM+AA(I)
C      HXS=(SUM*AA(1)/2.)*AKDEL*2.
C      HXS=HXS/HNORM*1.000000
C      HZS=HZS/HNORM*1.000000

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INV099690
INV099800
INV099910
INV099820
INV099930
INV099840
INV099850
INV099860
INV099870
INV099880
INV099890
INV099900
INV099910
INV099920
INV099930
INV099940
INV099950
INV099960
INV099970
INV099980
INV099990
INV100000
INV100010
INV100020
INV100030
INV100040
INV100050
INV100060
INV100070
INV100080
INV100090
INV100100
INV100110
INV100120
INV100130
INV100140
INV100150
INV100160
INV100170
INV100180

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INV11390
INV11400
INV11410
INV11420
INV11430
INV11440
INV11450
INV11460
INV11470
INV11480
INV11490
INV11500
INV11510
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INV11570
INV11580
INV11590
INV11600
INV11610
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INV11630
INV11640
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INV11660
INV11670
INV11680
INV11690
INV11700
INV11710
INV11720
INV11730
INV11740
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INV11770
INV11780
INV11790
INV11800
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INV11870
INV11880
INV11890
INV11900
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INV11920
INV11930
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INV11950
INV11960
INV11970
INV11980

if (abs(x) .lt. 3.75) then
  y=(x/3.75)**2
  BESSIO=p1+y*(p2+y*(p3+y*(p4+y*(p5+y*(p6+y*p7))))
else
  ax=abs(x)
  y=3.75/ax
  BESSIO=(exp(ax)/sqrt(ax))*(q1+y*(q2+y*(q3+y*(q4+y*(q5+
  + y*(q6+y*(q7+y*(q8+y*q9)))))))
endif
return
end

C
C Function BESSK0(x)
C Returns the modified Bessel function K0(x) for positive x.
C
C IMPLICIT DOUBLE PRECISION (A-H, O-Z)
real*8 y,p1,p2,p3,p4,p5,p6,p7,q1,q2,q3,q4,q5,q6,q7
data p1,p2,p3,p4,p5,p6,p7/-0.67721566d0,0.42278420d0,0.23089756d0,
+ 0.348859d-1,0.26289d-2,0.10750d-3,0.74d-6/
data q1,q2,q3,q4,q5,q6,q7/1.25331414d0,-0.7832358d-1,0.2189568d-1,
+ -0.1092448d-1,0.687872d-2,-0.261540d-2,0.53208d-3/
if (x.le.2.0) then
  y=x*x/4.0
  BESSK0=(-log(x/2.0)*BESSIO(x))+p1+y*(p2+y*(p3+
  + y*(p4+y*(p5+y*(p6+y*p7))))
else
  if (x.gt.40.) then
    BESSK0=0.0
  else
    y=(2.0/x)
    BESSK0=(exp(-x)/sqrt(x))*(q1+y*(q2+y*(q3+y*(q4+y*(q5+
    + y*(q6+y*q7))))))
  endif
endif
return
end

C
C Function BESSII(x)
C Returns the modified Bessel function I1(x) for any real x.
C
C IMPLICIT DOUBLE PRECISION (A-H, O-Z)
real*8 y,p1,p2,p3,p4,p5,p6,p7,q1,q2,q3,q4,q5,q6,q7,q8,q9
data p1,p2,p3,p4,p5,p6,p7/0.5d0,0.87890594d0,0.51498869d0,
+ 0.15084934d0,0.2658735d-1,0.301532d-2,0.32411d-3/
data q1,q2,q3,q4,q5,q6,q7,q8,q9/0.39894228d0,-0.3988024d-1,
+ -0.362018d-2,0.163801d-2,-0.1031555d-1,0.2282967d-1,
+ -0.289631d-1,0.1787954d-1,-0.420069d-2/
if (abs(x) .lt. 3.75) then
  y=(x/3.75)**2
  BESSII=x*(p1+y*(p2+y*(p3+y*(p4+y*(p5+y*(p6+y*p7))))))
else
  ax=abs(x)
  y=3.75/ax
  BESSII=(exp(ax)/sqrt(ax))*(q1+y*(q2+y*(q3+y*(q4+y*(q5+
  + y*(q6+y*(q7+y*(q8+y*q9)))))))
endif

```

```

10      A=(AKY(KHI)-XX)/H
        B=(XX-ACY(KLO))/H
        O(I)=A*YA(KLO)+B*YA(KHI)+(A*(A-A-1.)*Y2(KLO)+
        * B*(B-B-1.)*Y2(KHI))/(H*H)/6.
        + CONTINUE
        RETURN
        END
C
C      SUBROUTINE SPLINE(X,Y,N,YP1,YPN,Y2)
C
C      GIVEN ARRAYS X AND Y OF LENGTH N CONTAINING A TABULATED FUNCTION,
C      I.E., YJ=F(XJ), WITH X1<X2<X3...<XN, THIS ROUTINE RETURNS AN ARRAY
C      Y2 OF LENGTH N WHICH CONTAINS THE SECOND DERIVATIVES OF THE INTER-
C      POLATING FUNCTION AT THE TABULATED POINTS XJ. NATURAL SPLINE FROM
C      NUMERICAL RECIPES, PRESS AT AL.
C
C      IMPLICIT DOUBLE PRECISION (A-H, O-Z)
C      PARAMETER (NMAX=1024)
C      DIMENSION X(N),Y(N),Y2(N),U(NMAX)
C      IF (YPI.GT..99E30) THEN
C        Y2(1)=0.
C        U(1)=0.
C      ELSE
C        Y2(1)=-0.5
C        U(1)=(3./(X(2)-X(1)))*((Y(2)-Y(1))/(X(2)-X(1))-YP1)
C      ENDF
C      DO 11 I=2,N-1
C        SIG=(X(I)-X(I-1))/(X(I+1)-X(I-1))
C        F=SIG*Y2(I-1)+2.
C        Y2(I)=(SIG-1.)/P
C        U(I)=(0.*(Y(I+1)-Y(I))/(X(I+1)-X(I-1))-Y(I)-Y(I-1))/P
C      CONTINUE
11      IF (YPN.GT..99E30) THEN
        QN=0.
        UN=0.
      ELSE
        QN=0.5
        UN=(3./(X(N)-X(N-1)))*(YPN-(Y(N)-Y(N-1))/(X(N)-X(N-1)))
      ENDF
      Y2(N)=(UN-QN*U(N-1))/(QN*Y2(N-1)+1.)
      DO 12 K=N-1,1,-1
        Y2(K)=Y2(K)*Y2(K+1)+U(K)
      CONTINUE
      RETURN
      END
C
C      Function BESSI0(x)
C
C      RETURNS THE MODIFIED BESSEL FUNCTION I0(X) FOR ANY REAL X.
C
C      IMPLICIT DOUBLE PRECISION (A-H, O-Z)
      real*8 y,p1,p2,p3,p4,p5,p6,p7,q1,q2,q3,q4,q5,q6,q7,q8,q9
      data p1,p2,p3,p4,p5,p6,p7/1.0d0,3.5156229d0,3.0899424d0,
      + 1.2067492d0,0.2659732d0,0.360788d-1,0.45813d-2/
      data q1,q2,q3,q4,q5,q6,q7,q8,q9/0.39894228d0,0.1328592d-1,
      + 0.225319d-2,-0.167565d-2,0.916281d-2,-0.2057706d-1,
      + 0.2835637d-1,-0.1647833d-1,0.392377d-2/

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INV10790
INV10800
INV10810
INV10820
INV10830
INV10840
INV10850
INV10860
INV10870
INV10880
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INV10900
INV10910
INV10920
INV10930
INV10940
INV10950
INV10960
INV10970
INV10980
INV10990
INV11000
INV11010
INV11020
INV11030
INV11040
INV11050
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INV11070
INV11080
INV11090
INV11100
INV11110
INV11120
INV11130
INV11140
INV11150
INV11160
INV11170
INV11180
INV11190
INV11200
INV11210
INV11220
INV11230
INV11240
INV11250
INV11260
INV11270
INV11280
INV11290
INV11300
INV11310
INV11320
INV11330
INV11340
INV11350
INV11360
INV11370
INV11380

```

```

IOUT=0
DEL=0
DO 21 I=K1,N
LI=LT(I)
IF(X(LI).EQ.BL(LI).AND.G(I).GE.0.)GOTO21
IF(X(LI).EQ.BU(LI).AND.G(I).LE.0.)GOTO21
IF(G(I).LT.0.)GOTO22
Z=X(LI)-BL(LI)
J=1
GOTO23
22 CONTINUE
Z=BU(LI)-X(LI)
J=0
23 CONTINUE
IF(G(ICAC+I).LE.0.)GOTO24
BETA=ABS(G(I))/G(ICAC+I)
YF(BETA.GE.Z)GOTO24
Z=BETA
D=.5*Z*ABS(G(I))
J=-1
GOTO26
24 CONTINUE
D=Z*(ABS(G(I))-5*Z*G(ICAC+I))
26 CONTINUE
IF(D.LT.DEL)GOTO21
DEL=D
ALPHA=Z
IOUT=I
IIN=I
IF(J.LT.0)IIN=0
LB=J
21 CONTINUE
IF(IOUT.NE.0)GOTO29
27 CONTINUE
Q=0
DO 28 I=1,N
LI=LT(I)
Q=Q+X(LI)*(G(I)-B(LI))
Q=Q-Q
RETURN
29 CONTINUE
SIG=1
IF(G(IOUT).GT.0.)SIG=-1.
LIOUT=LT(IOUT)
LIIN=LIOUT
25 CONTINUE
SAS=G(ICAC+IOUT)
IF(K.EQ.0)GOTO31
DO 30 I=1,K
G(IIS+I)=G(ID+I)*A(IOUT,I)
31 CONTINUE
DO 37 I=K1,N
LI=LT(I)
IF(LI-LIOUT)32,37,33
32 Z=A(LI,LIOUT)
GOTO34
33 Z=A(LIOUT,LI)
34 CONTINUE
IF(K.EQ.0)GOTO36
DO 35 J=1,K

```

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INV12690
INV12800
INV12810
INV12820
INV12830
INV12840
INV12850
INV12860
INV12870
INV12880
INV12890
INV12900
INV12910
INV12920
INV12930
INV12940
INV12950
INV12960
INV12970
INV12980
INV12990
INV13000
INV13010
INV13020
INV13030
INV13040
INV13050
INV13060
INV13070
INV13080
INV13090
INV13100
INV13110
INV13120
INV13130
INV13140
INV13150
INV13160
INV13170
INV13180

```

```

return
end
C
C
Function BESSK1(x)
C
C ... Returns the modified Bessel function KI(x) for positive x
C
IMPLICIT DOUBLE PRECISION (A-H, O-Z)
real*8 y,p1,p2,p3,p4,p5,p6,p7,q1,q2,q3,q4,q5,q6,q7
data p1,p2,p3,p4,p5,p6,p7/1.0d0,0.15443144d0,-0.67278679d0,
* -0.1816897d0,-0.191940d-1,-0.110484d-2,-0.4686d-4/
data q1,q2,q3,q4,q5,q6,q7/1.6331414d0,0.23498819d0,-0.3656620d-1,
* 0.1604268d-1,-0.780363d-2,0.326614d-2,-0.88246d-3/
if (x.le.2.0) then
y=x*x/4.0
BESSK1= (log(x/2.0)*BESSI1(x))+(1.0/x)*(p1+y*(p2+y*(p3+
* if (.gt.40) then
y*(p4+y*(p5+y*(p6+y*p7))))))
else
BESSK1=0.0
elseif
y=(2.0/x)
BESSK1=(exp(-x)/sqrt(x))*(q1+y*(q2+y*(q3+y*(q4+y*(q5+
* endif
endif
return
end
SUBROUTINE VE04A(N,A,IA,B,BU,X,Q,L,T,K,G)
IMPLICIT REAL*8 (A-H,O-Z)
C
C
C QUADRATIC PROGRAMMING FROM HARWELL LIBRARY.
C SIMPLE BOUND CONSTRAINTS.
C
DIMENSION A(IA,1),B(1),BL(1),BU(1),X(1),LT(1),G(1)
IS=N
IAS=N
IV=N
ICAC=N+N
ID=ICAC
DO 9 I=1,N
G(I)=-B(I)
DO 10 I=1,N
X(I)=0.
LT(I)=I
G(ICAC+I)=A(I,I)
IF(0..GE.BL(I).AND.0..LE.BU(I))GOTO10
IF(0..LT.BL(I))X(I)=BL(I)
IF(0..GT.BU(I))X(I)=BU(I)
DO 12 J=1,I
G(J)=G(J)+A(J,I)*X(I)
IF(I.EQ.N)GOTO10
II=I+1
DO 11 J=II,N
G(J)=G(J)+A(I,J)*X(I)
10 CONTINUE
K=0
K1=1
20 CONTINUE

```

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INV11900
INV12000
INV12010
INV12020
INV12030
INV12040
INV12050
INV12060
INV12070
INV12080
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INV12100
INV12110
INV12120
INV12130
INV12140
INV12150
INV12160
INV12170
INV12180
INV12190
INV12200
INV12210
INV12220
INV12230
INV12240
INV12250
INV12260
INV12270
INV12280
INV12290
INV12300
INV12310
INV12320
INV12330
INV12340
INV12350
INV12360
INV12370
INV12380
INV12390
INV12400
INV12410
INV12420
INV12430
INV12440
INV12450
INV12460
INV12470
INV12480
INV12490
INV12500
INV12510
INV12520
INV12530
INV12540
INV12550
INV12560
INV12570
INV12580

```



```

VD=V/G(ID+I1)
S1=S0+V*VD
R=S1/S0
G(ID+I)=G(ID+I1)*R
BETA=VD/S1
IF(R.GT.4.)GOTO841
DO 81 J=I2,N
81 G(IV+J)=G(IV+J)-V*A(J,I1)
IF(I1.GT.K2)GOTO83
DO 82 J=I1,K2
J1=J+1
82 A(J,I)=A(J1,I1)+BETA*G(IV+J1)
83 CONTINUE
A(K,I)=BETA
DO 84 J=K1,N
84 A(J,I)=A(J,I1)+BETA*G(IV+J)
GOTO849
841 CONTINUE
IF(I1.GT.K2)GOTO843
DO 842 J=I1,K2
J1=J+1
842 A(J,I)=BETA*G(IV+J1)+A(J1,I1)/R
843 CONTINUE
A(K,I)=BETA
DO 844 J=K1,N
844 A(J,I)=BETA*G(IV+J)+A(J,I1)/R
DO 846 J=I2,N
846 G(IV+J)=G(IV+J)-V*A(J,I1)
849 CONTINUE
LT(I)=LT(I1)
S0=S1
I1=I2
I2=I2+1
SG=1./S1
LT(K)=LIIN
G(ID+K)=SG
IF(IIN.EQ.1)GOTO361
II=IIN-1
DO 862 I=1,II
Z=A(IIN,I)
DO 863 J=IIN,K2
863 A(J,I)=A(J+1,I)
862 A(K,I)=Z
861 CONTINUE
86 CONTINUE
DO 87 I=K1,N
87 G(ICAC+I)=G(ICAC+I)+SG*G(IV+I)**2
K1=K
K=K2
IIN=0
ALPHA=1E76
SAS=G(ICAC+IOUT)
IF(SAS.GT.0.)ALPHA=ABS(G(IOUT))/SAS
IF(G(IOUT).LT.0.)GOTO898
J=1
Z=X(LIOUT)-BL(LIOUT)
GOTO899
898 CONTINUE
J=0
Z=BU(LIOUT)-X(LIOUT)

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INV13790
INV13800
INV13810
INV13820
INV13830
INV13840
INV13850
INV13860
INV13870
INV13880
INV13890
INV13900
INV13910
INV13920
INV13930
INV13940
INV13950
INV13960
INV13970
INV13980
INV13990
INV14000
INV14010
INV14020
INV14030
INV14040
INV14050
INV14060
INV14070
INV14080
INV14090
INV14100
INV14110
INV14120
INV14130
INV14140
INV14150
INV14160
INV14170
INV14180
INV14190
INV14200
INV14210
INV14220
INV14230
INV14240
INV14250
INV14260
INV14270
INV14280
INV14290
INV14300
INV14310
INV14320
INV14330
INV14340
INV14350
INV14360
INV14370
INV14380

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35 Z=Z-A(I,J)*G(IS*J)
36 G(IS*I)=Z
37 CONTINUE
  G(IS*IOUT)=SAS
  IF (K.EQ.0)GOTO42
  G(IS*K)=-A(IOUT,K)
  IF (K.EQ.1)GOTO42
  I=K
  DO 41 II=2,K
  I=I-1
  Z=-A(IOUT,I)
  I1=I+1
  DO 40 J=I1,K
  Z=Z-G(IS*J)*A(J,I)
41 G(IS*I)=Z
42 CONTINUE
  IF (SIG.EQ.1.)GOTO61
  DO 50 I=1,N
  G(IS*I)=-G(IS*I)
51 CONTINUE
  IF (K.EQ.0)GOTO62
  DO 61 I=1,K
  IF (G(IS*I).EQ.0.)GOTO61
  LI=LT(I)
  J=1
  Z=BL(LI)-X(LI)
  IF (G(IS*I).LT.0.)GOTO60
  J=0
  Z=BU(LI)-X(LI)
90 CONTINUE
  Z=Z/G(IS*I)
  IF (Z.GE.ALPHA)GOTO61
  ALPHA=Z
  LB=J
  IIN=I
  LIIN=LI
  LIIN=LI
61 CONTINUE
62 CONTINUE
  X(LIOUT)=X(LIOUT)+SIG*ALPHA
  IF (K.EQ.0)GOTO71
  DO 70 I=1,K
  LI=LT(I)
  X(LI)=X(LI)+ALPHA*G(IS*I)
71 CONTINUE
  DO 72 I=K1,N
  G(I)=G(I)+ALPHA*G(IAS*I)
  IF (IIN.EQ.0)GOTO90
  X(LIIN)=BL(LIIN)
  IF (LB.EQ.0)X(LIIN)=BU(LIIN)
  IF (IIN.EQ.IOUT)GOTO20
  K2=K-1
  SG=G(ID*IIN)
  I1=IIN+1
  DO 80 I=I1,N
  G(IV*I)=A(I,IIN)
  IF (IIN.EQ.K)GOTO86
  I2=IIN+2
  S0=1./SG
  DO 85 I=IIN,K2
  V=G(IV*I1)

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```

INV13190
INV13200
INV13210
INV13220
INV13230
INV13240
INV13250
INV13260
INV13270
INV13280
INV13290
INV13300
INV13310
INV13320
INV13330
INV13340
INV13350
INV13360
INV13370
INV13380
INV13390
INV13400
INV13410
INV13420
INV13430
INV13440
INV13450
INV13460
INV13470
INV13480
INV13490
INV13500
INV13510
INV13520
INV13530
INV13540
INV13550
INV13560
INV13570
INV13580
INV13590
INV13600
INV13610
INV13620
INV13630
INV13640
INV13650
INV13660
INV13670
INV13680
INV13690
INV13700
INV13710
INV13720
INV13730
INV13740
INV13750
INV13760
INV13770
INV13780

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```
DC 100 L=I1,J1
100 Z=Z+A(J,L)*A(L,I)
101 A(J,I)=-Z
102 CONTINUE
AA=1./G(ID*I1)
G(N*I1)=AA
DO 111 J=1,I
Z=A(I1,J)*AA
G(N*J)=G(N*J)+Z*A(I1,J)
IF(I.EQ.1)GO TO 111
J1=J+1
DO 110 L=J1,I
110 A(L,J)=A(L,J)+A(I1,L)*Z
111 A(I1,J)=Z
RETURN
END
```

```
INV14990
INV15000
INV15010
INV15020
INV15030
INV15040
INV15050
INV15060
INV15070
INV15080
INV15090
INV15100
INV15110
INV15120
INV15130
INV15140
```

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899 CONTINUE
IF (Z .GE. ALPHA) GOT026
ALPHA=Z
LB=J
IIN=IOUT
LIIN=LIOUT
GOT026
90 CONTINUE
K2=K1+1
IF (SIG.EQ.1.) GOT091
DO 901 I=K1,N
901 G(IAS+I)=-G(IAS+I)
91 CONTINUE
IF (IOUT.EQ.K1) GOT097
LT(IOUT)=LT(K1)
LT(K1)=LIOUT
G(IAS+IOUT)=G(IAS+K1)
G(ICAC+IOUT)=G(ICAC+K1)
G(ICAC+K1)=SAS
G(IOUT)=G(K1)
IF (K.EQ.0) GOT097
DO 92 I=1,K
Z=A(K1,I)
A(K1,I)=A(IOUT,I)
92 A(IOUT,I)=Z
93 CONTINUE
IF (K2.EQ.IOUT) GOT096
I1=IOUT-1
DO 94 I=K2,I1
94 A(IOUT,I)=A(I,K1)
95 CONTINUE
IF (IOUT.EQ.N) GOT097
I1=IOUT+1
DO 96 I=I1,N
96 A(I,IOUT)=A(I,K1)
97 CONTINUE
G(K1)=0.
K=K1
IF (K.EQ.N) GOT027
DO 98 I=K2,N
Z=G(IAS+I)/SAS
A(I,K1)=Z
98 G(ICAC+I)=G(ICAC+I)-Z*G(IAS+I)
K1=K2
GOT020

ENTRY VE04B(N,A,IA,G,K)
IF (K.EQ.0) RETURN
ID=N+N
G(N+1)=1./G(ID+1)
IF (K.EQ.1) RETURN
N1=K-1
DO 111 I=1,N1
I1=I+1
A(I1,I)=-A(I1,I)
IF (I.EQ.N1) GOT0102
I1=I+2
DO 101 J=I1,K
Z=A(J,I)
J1=J-1

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INV14390
INV14400
INV14410
INV14420
INV14430
INV14440
INV14450
INV14460
INV14470
INV14480
INV14490
INV14500
INV14510
INV14520
INV14530
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INV14690
INV14700
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INV14800
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INV14850
INV14860
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INV14890
INV14900
INV14910
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INV14980

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Introduction

The utilization of downhole current sources in resistivity mapping increases the resolution for detecting and delineating subsurface features. The effects of near surface inhomogeneities are immensely reduced as shown by Asch and Morrison (1988). Being sensitive to changes in resistivities, the surveys with downhole sources are well suited for monitoring surface processes such as injection or leakage of contaminants from a waste site, steam flooding for enhanced oil recovery, or production of geothermal reservoirs.

In most of these applications, the holes are steel-cased and the casing distorts the current flow in the medium. Holladay and West (1984) have shown that surface resistivity surveys are strongly affected by casings. Also, Kauahikaua et al (1980) showed that if the casing itself is used as an electrode, the results are unpredictable because the current leaves the pipe irregularly due to the variability of the contact resistance between the pipe and the formation.

To study the casing effects in more detail, we have recently formulated the problem for a point source of current, either inside or outside the pipe, on the axis of a finite length metal pipe in a conductive half space. The first part of this study (Schenkel and Morrison, 1987) showed that only the region very near the pipe exerts any substantial influence on the potential fields for a point source 100 casing diameters beyond the end of the pipe (Figure 1). For a 5% or less field distortion, the surface measurements must not be closer than 1/2 the pipe's length. In cross-hole surveys, the affected area is greatly reduced; near the pipe's end, measurements as close as 1/6 the length of the pipe for a 5% or less distortion. If the pipe-source separation is sufficiently large, then the resistivity survey can be corrected for the casing effects (assuming that the target are not too close to the pipe).

This study showed that cross-hole and hole-surface studies may be conducted with very little effects: even when the pipe exerts some influence, e.g., when the current source is close to the end, time monitoring experiments could be carried out with very little reduction in the signal strength. Further, this study revealed the intriguing possibility that segments of pipe, separated by insulated couplings, could be used as electrodes.

Proposed Work

The above study was only developed for a very simple case, i.e., a pipe in a homogeneous half-space. A more complex model is required to simulate field situations. Several aims are proposed to create a realistic model and to evaluate field measurements for downhole sources in steel cased wells. These objectives are:

- 1) To determine the effects of contact resistance between the pipe and the host medium. Contact resistance is used to describe the resistance of the pipe-medium contact. If there is a large contact resistance, the results will completely change. The currents in the pipe will only leak out of the pipe in areas where the pipe has made a good contact with the host medium, thus completely changing the potential fields around the pipe. The contact resistance may be found by assuming a very thin layer between the host and pipe for each segment (Figure 2). An equivalent layer can be used to represent the effects of the contact resistance for each segment. The calculated potential field is obtained from the integral equation solution of the pipe variables. With the additional layer included to the model, the effects of pipe coating, corrosion, and cement on the potential fields can be evaluated.
- 2) To use insulated pipe segments as downhole sources and receivers. Downhole current electrodes can be created by energizing isolated segments. Likewise, insulated segments could be used as potential electrodes. By insulating several segments in the well (Figure 3), multiple

downhole source locations could be used to image a target in a hole-surface survey. A pole-pole configuration can be achieved by attaching to different segments current and potential electrodes. The isolated segments in the well can be used for an AC vertical electric dipole. By attaching to adjacent segments a positive and negative AC source, a grounded electric dipole can be produced to study EM properties of the medium. If an additional well is drilled with multiple isolated segments, then cross-hole DC and AC measurements can be acquired. Thus, cross-hole DC tomography (Daily and Morely, 1988 and Shima and Saito, 1988) could be used to reconstruct an image of a target between the two wells.

- 3) To determine the interaction between the source, pipe, and the anomalous body (Figure 4). To monitor the changes in the body, the effects of the metal pipe-body interaction must be investigated. The extent of this interaction will greatly depend on the source separation from the pipe, the distance between the pipe and the body, and the conductivity of the body. The determination of the limiting values of these parameters in which the body and pipe has very little coupling will be the main objective. For this situation, the body can be modeled alone saving computational time.
- 4) To invert for the geometric parameters for a plume-like body. An integral equation solution of an ellipsoidal model will be used to represent the plume. The parameters of the three axial lengths will be obtained by a non-linear least squares inversion which will make use of the integral equation solution. Sensitivity analysis and minimum spatial coverage will be evaluated for various acquisition array configurations.

Computer models will be required to investigate the above proposed tasks. The current algorithm is flexible so that variable source locations, segment lengths, and segment conductivities can be used. The development of an algorithm which includes an anomalous body and an outer layer is

needed. The current integral equation solution can be extended to include an additional layer and a circular disk. The circular disk is a first order approximation to a plume and will give an estimate of the pipe-body interaction.

Various lengths and separations of the insulated segments will be investigated to determine when a point approximation of the segments can be used. The outer layer will be used to evaluate the role of contact resistance on the field distortion and may be included to the pipe-body coupling phenomenon. The spatial separation of the source, pipe, and body to decouple the pipe and the body will be studied.

Field test and model verification are also required. The test would be conducted at U.C. Berkeley, Richmond Field Station where there exist several plastic-cased holes. Four additional 100 foot holes are also needed. Two of the wells will be composed of alternating steel and fiberglass segments. Another would be cased with steel with the last 20 feet perforated. The last well needs to be plastic cased and perforated at the bottom. A surface grid and/or radial arrays of potential electrodes would be installed over the area. Cross-hole and hole-surface measurements would be compared to calculated fields of various pipe models. An injection of the salt water in both the steel-cased and plastic-cased wells would be measured for hole-surface, cross-hole, and pole-pole configurations. These data would be forward modeled and inverted to determine the geometry of the plume. Lastly, other configurations for which models have been published can be field tested with this field setup.

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