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The Influence of Bubbles on Sea Surface Backscatter Measurements

Der Einfluss von Blasen auf die Messung der Rückstreueigenschaften der Meeresoberfläche

> A Paper to be Presented at the Colloquium on Acoustics in Torpedo Technology, Newport, Rhode Island

Paul D. Koenigs Joseph M. Monti Surface Ship Sonar Department Naval Underwater Systems Center

Bernd Nützel Heinz Herwig Forschungsanstalt der Bundeswehr für Wasserschall- und Geophysik

26 October 1986



NAVAL UNDERWATER SYSTEMS CENTER Newport, Rhode Island • New London, Connecticut



FORSCHUNGSANSTALT DER BUNDESWEHR FÜR WASSERSCHALL- UND GEOPHYSIK Kiel

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Preface

This document describes work performed under Data Exchange Agreement MWDDEA-N-67-G-4207, Subproject 2, between the Forschungsanstalt der Bundeswehr fuer Wasserschall- und Geophysik (FWG), Kiel, Federal Republic of Germany (FRG), and the Naval Underwater Systems Center (NUSC), New London, Connecticut, USA. The U.S. effort was accomplished under NUSC Project No. B68201, "Environmental Dependence of Acoustic Surface Scatter," Principal Investigator, P. D. Koenigs (NUSC Code 3331). Funding was provided under Program Element 62759N, by Naval Sea Systems Command (NAVSEA 63R); Dr. R. Farwell (NORDA 113), Manager, and by NUSC Associate Technical Director, Research and Technology Office. The work was performed under the auspices of the Office of the Chief of Naval Research (Code 122).

Vorwort

Die Messungen an der Forschungsplattform NORDSEE wurden im Rahmen des Datenaustauschabkommens MWDDEA-N-67-G-4207, Subproject 2, durchgeführt. Die Abstimmung des jährlichen Forschungsprogrammes der FWG mit dem Inhalt des DEA erfolgte durch RüFo 3 des BMVg der Bundesrepublik Deutschland.

APPROVED: 26 October 1986

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EINVERSTANDEN: 26 Oktober 1986

PROF. Dr. G. Ziehm Director, Defence Research Institute for Underwater Sound and Geophysics Federal Republic of Germany

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This report is the result of a cooperative research project conducted under the auspices of ONR's data exchange agreement MWDDEA-N-67-G-4207, Subproject 2.

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Dr. William A. Von Winkle Naval Underwater Systems Center for the United States Project Officer

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Strength at a 30 grazing angle is caused by the high frequency wavenumber spectrum at low wind speeds and by sub-surface bubbles at high wind speeds. The backscattering strength shows strong fluctuations in the intermediate region caused by both scattering mechanisms. ventionellen Betrieb als Empfänger arbeitete. Rückstreustärken wurden in Abhängigkeit der Frequenz (3 bis 18 kHz) für Windgeschwindigkeiten von 2 bis 45 Knoten gemessen. Der Bericht zeigt, dass die Rückstreustärke bei einem Glanzwinkel von 30° bei geringen Windgeschwindigkeiten vom hoch frequenten Seegangsspektrum und bei hohen Windgeschwin digkeiten von Blasen bestimmt wird. Im Zwischenbereich weist die Rückstreustärke starke Fluktuationen auf, die von beiden Streumechanismen verursacht werden.

THE INFLUENCE OF BUBBLES ON SEA SURFACE BACKSCATTER MEASUREMENTS

Numerous theoretical studies and experimental programs have been instituted to investigate acoustic reverberation originating at or near the sea surface. The results of these studies indicate surface roughness and air bubbles must be considered to explain the environmental dependence of acoustic scattering on sonar engineering parameters such as frequency, grazing angle, and system design characteristics.

A principal reason for conducting a recent sea surface reverberation study was to determine the frequency regime over which different environmental parameters appear to be the governing factor.

--- First viewgraph, please. ---

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VIEWGRAPH 1

The experiment site was the Federal Republic of Germany's research platform NORDSEE, which is approximately 40 nautical miles west of the German and Danish coasts. The platform is located in approximately 30 m of water. A high resolution parametric array was installed atop a 7.5 m tower and emplanted a short distance from the platform. The parametric array was used as an acoustic projector because of its broad bandwidth, narrow beamwidths, and very low sidelobe levels, which help alleviate the problem of multipath separation. This system has a nominal 2.5 degree beamwidth at 18 kHz and could be remotely trained in elevation and azimuth. The system was also used as a conventional receiver for backscattering measurements.

In addition to these measurements, radar backscatter measurements of the sea surface were concurrently obtained. Numerous supporting measurements were also taken such as sound speed, wind velocity, subsurface bubble, and ocean wave characteristics (using photographs, video recordings, wave rider buoy, and capacitance wave staffs) [1].



VIEWGRAPH 2

Shown in this viewgraph are what we currently believe to be the principal features affecting acoustic surface reverberation from high resolution acoustic systems.

Based on the theoretical work of Bass and Fuks [2] a two-scale description of the sea surface is needed to characterize surface scattering. In this description, the scattering surface is represented as a superposition of short and long waves. The surface backscatter is produced by Bragg diffraction from small wavelets of wavenumber (capital K) that are equal to twice the acoustic wavenumber (small k) in the same plane. The projection into the plane is related to the grazing angle, θ_g . The large waves cause the Bragg diffraction grating to tilt, thus modifying the average grazing angle θ .

Based on the work by Thorpe and others [3, 4], bubble density in the ocean is not always horizontally stratified. When the wind is strong enough to form whitecaps, bubble clouds or plumes are generated. These plumes penetrate surprisingly far into the water column and exhibit many different characteristics similar to cloud formations in the atmosphere. The fluctuation periods of plumes are from 1 to 10 minutes.

Thus, it seems appropriate to hypothesize two types of near surface bubble formations. One consisting of uniform layers of microbubbles that form a background bubble density. The other consists of bubble plumes containing much larger bubbles of very high densities, which will exhibit temporal and spatial fluctuations as the plumes disperse and decay.



INCOHERENT MONOSTATIC BACKSCATTER

$$I_{R} = I_{o} \sum_{i} \frac{\Psi_{i}}{TLC_{i}^{2}}$$

SURFACE

$$\Psi_i = \Phi_i A_i$$

VOLUME

$$\psi_{i} = R_{o}V_{i} \sum_{i} N_{ij}\sigma_{ij} = R_{o}V_{i}s_{v_{i}}$$

COHERENT MONOSTATIC BACKSCATTER

 $= \frac{I_o \phi_{coh}}{2 \text{ TLC}}$

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VIEWGRAPH 3

The intensity of an incoherently scattered acoustic wave is equal to the incident intensity times a scattering function. If we only consider monostatic backscatter and modify the incident and scattered intensity by the transmission loss, TLC, we may express the total reverberation at a receiver, I_R , in terms of intensity at the source, I_O , by the upper equation, where ψ_i is the elemental scattering cross section. We thus see the total reverberation is a sum over all contributing scattering elements.

The extent or number of scattering elements may be determined by realizing the reverberation at any instant of time is a convolution of the spatial impulse response of the scatterers and the transmitted signal [5]. When the scattering mechanisms are distributed or located such that their temporal extent is greater than the pulse length, the contributing scatterers are determined by pulse length. When the pulse length is longer than the temporal extent of the scatterers, all elemental scatterers may contribute.

Scattering cross sections are not usually presented in sonar literature. Instead the decibel measure of a scattering function related to target strength is commonly used. For scattering from a surface, the elemental scattering cross section, ψ , is simply related to the scattering function, ϕ , by the second equation, where A is an area. Surface scattering strength is 10 log ϕ . For an admixture of air bubbles and water, the elemental scattering cross section is related to the scattering coefficient, s_{v} , by the third equation. Here we should note the scattering coefficient is a sum over the number j of bubbles in the elemental volume, V_i , and depending on the distribution may not be principally controlled by resonant bubble size. In addition, σ_{ij} is the scattering cross section of a bubble and is strongly dependent on acoustic frequency near and below resonance. Volume scattering strength is 10 log of the volume scattering coefficient times a reference range, normally 1 m.

The intensity of a coherently scattered acoustic wave is found by using the solution of the image reflection problem at a plane interface. For the normal incidence case, the source and receiver are colocated and the solution is given by the fourth equation. The reflection loss in this case is simply 10 log of the scattering function ϕ .



VIEWGRAPH 4

This viewgraph depicts surface backscattering strength versus grazing angle for various wind speeds. The solid lines are the result of the model recommended by the Applied Physics Laboratory at the University of Washington for use at ultrasonic frequencies [6]. The dashed lines indicate the average results of our experiments at a comparable frequency at high wind speeds as measured 47 m above the sea surface. Several results are noteworthy. First, there is good agreement between the model and the data. Secondly, we can observe that at low grazing angles backscattering strength increases with increasing wind speed but appears to reach a maximum or saturation value above 20 knots. We also note scattering strength behaves in the opposite manner at very high grazing angles. In addition, this viewgraph serves as a guide to the specific backscattering parameters we wish to address. That is the dependence of backscattering strength at the two grazing angles indicated by the arrows at 90 and 30 degrees as a function of the environment in the frequency regime of 18 to 3 kHz.



VIEWGRAPH 5

This viewgraph illustrates backscatter strength as a function of wind speed at 3 and 18 kHz for normal incidence.

Backscattering at normal incidence is considered here to be a coherent process. The backscatter strength values (BS) are corrected for source level (SL) and transmission loss only.

At 18 kHz it can be seen backscatter strength decreases as wind speed increases up to about 30 knots. Beyond this wind speed, it appears a saturation region is reached and there is no longer any significant correlation between backscatter and wind speed. This is the same trend predicted by the APL-Washington model and has been observed many times. The data at 3 kHz are significantly different. At very low wind speeds there is a strong dependence on wind speed. This is to be expected because in this region the Rayleigh roughness parameter is small. Had our experimental data at 18 kHz extended to comparable low wind speeds we would expect the same trend. For higher wind speeds the scattering strengths at 3 and 18 kHz are similar, until at wind speeds above 28 knots the scattering strength again shows a dependence on wind speed attaining a saturation in the 38 knot region. Though not shown here, our data at intermediate frequencies exhibit the same trend. It thus appears other phenomena must be playing a significant role in the backscattering mechanism. To examine this more carefully, individual returns obtained at 18 kHz in the saturation region were examined.



VIEWGRAPH 6

This viewgraph shows the envelope of the time series for two different pings at normal incidence. The transmitted frequency was 18 kHz with a pulse width of 4 ms. The significant wave height was 3.3 m and the wind speed was 27 knots. During this time period many breaking waves occurred. The upper plot shows a single ping return 49 s after the measurement set started. As seen from the steep slope at the beginning of the echo there is no obvious bubble reverberation. The lower plot, which shows the return from another ping 80 s later, indicates the presence of bubbles down to about 1 m below the surface. To investigate the influence of bubbles on the sea surface backscatter for normal incidence, the reverberation levels received from the surface and from the bubbles were evaluated for each individual ping. The surface reverberation levels to be presented are the average over the 4 ms past the very steep slope. The bubble reverberation level is the average over 1 ms starting 1.6 ms before this slope.



VIEWGRAPH 7

This viewgraph depicts the reverberation from the surface and the bubbles from each ping using the method just described. The circles indicate the levels that correspond to the two pulses just shown. The time series is 300 s and is a result of 750 pings at a repetition rate of 0.4 s. It can be seen that there is no correlation between surface reverberation and bubble reverberation levels. The reverberation from the surface has a relatively stationary mean. The reverberation level from the bubbles is considerably lower in amplitude and exhibits slowly varying characteristics of much longer time duration than associated with any wave components. This indicates the generation of bubble patches probably caused by breaking waves, which remain below the surface for an extended period of time. This agrees with our visual observations during the experiment.



VIEWGRAPH 8

The envelopes of time series averaged over 750 pings resulting from two different environmental conditions are shown in this viewgraph.

In both cases the transmit frequency was 18 kHz. The arrival time difference of the individual echoes caused by differing wave heights was removed using a thresholding technique. The upper curve shows the average return for a wind speed of 9 knots and a significant wave height of 0.5 m. The nearly constant steep slope of the leading edge indicates there is no apparent scattered energy from a subsurface bubble layer. The lower curve is the average from a high sea state condition. The onset of scattered energy corresponds to a depth of about 3.5 m below the surface. The reverberation level increases until the difference between surface and subsurface scattered energy is about 28 dB. We may thus conclude that at normal incidence and 18 kHz bubbles are not a dominant scattering mechanism for winds up to at least 27 knots.

Using the same analysis technique for the 3 kHz data we found no precursor caused by bubbles in the averaged time series. At lower frequencies, the effects caused by bubbles are even lower at normal incidence, as expected.



VIEWGRAPH 9

This viewgraph illustrates the influence of wind speed on backscattering strength for the low frequency, low wind speed regime and an extended set of grazing angles.

During one measurement set at 3 kHz, the wind speed increased from 2 to 5 knots. This changed the sea surface from glossy to one covered with small ripples. The significant wave height was constant at 0.4 m. The upper curves show backscattering strength as a function of grazing angle for these two conditions. At normal incidence the backscatter strength for the glossy surface is about 6 dB higher than for the case where ripples are present. For lower grazing angles, the ripples cause higher backscattering such that at 60 degrees the difference is about 8 dB. During these measurements, photographs of the sea surface were taken and analyzed using Stilwell's technique as modified by Baur [7]. The results of this analysis, shown in the lower part of the viewgraph, depict the associated directional wave spectra. The analysis bandwidth presented is from 1 to 3 Hz. The contour interval, depicted by the intersection of a black and white stripe is 6.5 dB. The vector U shows the wind direction and ψ indicates the azimuthal orientation of the acoustic axis. The inner circle is drawn for a surface wavelength of 0.25 m. This corresponds to the acoustic wavelength at 3 kHz. The energy of the waves at this frequency in the direction of the acoustic axis is about 10 dB higher when ripples are present. It is expected that the difference between the backscatter values would become greater for lower grazing angles as indicated by the different slopes of the two backscatter curves.

We may thus conclude at low wind speeds and all grazing angles backscatter at 3 kHz is strongly dependent on the high frequency ocean wavenumber spectrum.



VIEWGRAPH 10

This viewgraph illustrates the wind speed dependence of backscatter for low grazing angles near ultrasonic frequencies. The solid line represents APL-University of Washington's recommended model; the dots are a part of the data base upon which the model is founded and the dashed line is the result of our measurements. As with the previous comparison between the recent data set and APL's model, there is no significant difference between averaged results at comparable frequencies. Of concern to sonar engineers dealing with high data rates, high resolution sonars and advanced processing techniques is when one may rely on averaged results. The problem shall be illustrated for lower frequencies where the transition to the saturated region is less rapid.





VIEWGRAPH 11

This viewgraph represents data acquired in attempting to investigate the azimuthal dependence of backscatter relative to wave direction during two environmentally different periods. Each data point represents an average of 500 returns in a 3.3 minute period. The azimuthal transmit direction was then changed by 20 degrees and the next data set obtained. During the measurement periods of about 2 hours, the environmental conditions did not appreciably change.

We found no correlation between azimuthal transmit and wave direction for the given wind speed case, but it is apparent there is substantial variation between 3.3 minute averages. The highest difference is 16 dB for the two consecutive events marked by crosses in the lower curve.

To investigate the high degree of variability, the time history of these two events was plotted on a ping-to-ping basis.



VIEWGRAPH 12

The backscattered energy from 500 pulses at a repetition rate of 0.4 s is shown in each of these figures. The instantaneous energy was obtained at the time when the average maximum backscattered energy occurred. These two time series, though obviously of different character, were obtained 5 minutes apart at the same frequency. We see the backscatter energy suddenly increases by almost 20 dB and exhibits different decay rates comparable to the fluctuation periods of bubble plumes. It is obvious the average backscatter strength is affected by these long term fluctuations, that is the magnitude and duration of the high backscatter periods.



VIEWGRAPH 13

This viewgraph illustrates in a qualitative manner the dependence of 3 kHz backscatter variability on wind speed. The solid line indicates the same level for all plots. At the lowest wind speed of 6 knots, there are almost no fluctuations. As wind speed increases, we note the high scattering periods occur more often and persist for longer periods of time. The amplitude of the high scatter periods does not, however, appear to change appreciably. Comparing the curve at 13 and 40 knots, one notes an almost inverse behavior. At the lower wind speeds, there are short periods of high scatter while at high wind speeds, the backscatter drops down for only short periods of time. At the high wind speeds, the backscatter strength remains high for periods much too long to correlate even with the longest ocean waves. We must, therefore, conclude another mechanism such as bubbles is governing backscattering even at 3 kHz. Correlating the acoustic and bubble measurement is not possible because the bubble diameter for resonant scattering at this frequency is 2.2 mm, much too large to be measured with the optical bubble sensor used during the experiment. Bubbles of this size can only be generated by breaking waves. Video recordings were made of the acoustically insonified area during the experiments and a qualitative correlation on a ping-to-ping basis indicated reasonable agreement between breaking waves and periods of high backscatter. The correlation was particularly good when the backscattering strength increased very sharply with time.

We may, thus, conclude that for a grazing angle of 30 degrees and an acoustic frequency of 3 kHz, there is a transition from one scattering phenomenon to another. All indications lead us to believe the principal mechanism at low wind speeds is Bragg scattering from the sea surface and the principal mechanism at very high winds is volume scattering from localized bubble concentrations in the form of plumes or clouds.

We might expect bubbles to contribute to reverberation for high winds, however, the occurrence of the high scattering value shown in the upper left hand figure for winds of 6 knots is a curiosity. The video recording contained no breaking waves. A possible explanation is given by Middleton and Mellen [8]. They propose wind-generated solitons, moving nondispersively on the wind-driven drift layer are a plausible mechanism for large backscatter returns in the absence of near-surface bubble layers. Because wind stress was not measured during the experiments, we offer no further comment on this aspect of the data.

Curves obtained using the same analysis process at other frequencies exhibit the same behavior when the significant wave height is 1 m or greater.

Based on experience from previous experiments and analyses at lower wind speeds, we assumed backscattering to be a stationary process as long as the sca state remained constant. As can be seen from this viewgraph, the mean and standard deviation of the backscatter energy show strong differences when the time series is relatively short. To increase the integration time, the data from three events, obtained in temporal sequence, were combined and analyzed to obtain the normalized standard error or coefficient of variation.



o 3 kHz

40

30

10 kHz 18 kHz

-- Next viewgraph, please. --

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VIEWGRAPH 15

In an attempt to determine the frequency and environmental regimes over which one or the other mechanism is dominant at 30 degrees a subjective analysis of each time series was made to determine the relative period of time each mechanism appeared to be dominating the total scattering strength. This viewgraph, which is a result of that analysis contains three regions. Region I corresponds to frequencies and wind speeds in which surface scattering is the dominant mechanism at least 80% of the time. Region II is a transition region with the dashed line corresponding to equal time periods of surface and bubble scattering. Region III corresponds to frequencies and wind speeds in which volume scattering from near surface bubbles is the dominant mechanism at least 80% of the time.

The dominating mechanism is very dependent on grazing angle and as indicated earlier a similar figure at 90 degrees would only contain region i. Note the locations of the letter a, b, c. They correspond to three different regimes at 10 kHz.



VIEWGRAPH 16

Because the method used to determine relative time periods is subjective three samples of each time series corresponding to letters a, b, and c are shown in this viewgraph. The left time series corresponds to low wind speed where only two short periods near the beginning and end of the file contain high values of backscatter. In this case an estimate that 90% of the time reverberation was the result of surface scattering was made. The opposite extreme for 10 kHz is shown at the right. Here the wind speed is 43 knots and relatively high values of backscatter are present about 80% of the time. The middle plot illustrates the case when both surface and bubbles contributed to scattering for region II.



VIEWGRAPH 17

In conclusion, the relative contribution of different backscattering mechanisms is dependent on wind speed, grazing angle, and frequency. The wind speed dependence can be described by three regimes.

In region I, at low wind speeds, the scattering is caused by Bragg diffraction. Here, we see a strong dependence of scattering on the high frequency ocean wavenumber, spectrum at all grazing angles. The time history of backscattering shows low temporal variability.

In region II, at intermediate wind speeds, the scattering is caused by the sea surface and subsurface bubbles. The time history of backscattering shows strong fluctuations with target-like returns.

In region III, at high wind speeds, the scattering is caused by bubbles. In this regime, saturation occurs which is frequency and grazing angle dependent. The temporal variability of backscattering decreases again.

Thank you. Are there any questions?

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Institute of Oceanographic Sciences (S. A. Thorpe)	1
Raytheon Co. Submarine Signal Div. (P. Bilazarian)	1
Ferranti O.R.E. Inc. (F. C. Lowell, Jr.)	1
EDO Corp., Western Div. (R. A. Lapentina)	1
University of Miami/RSMAS (H. de Ferrari)	1
DTIC	12



