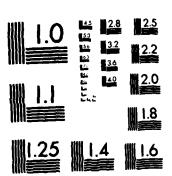
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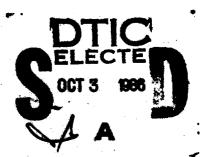
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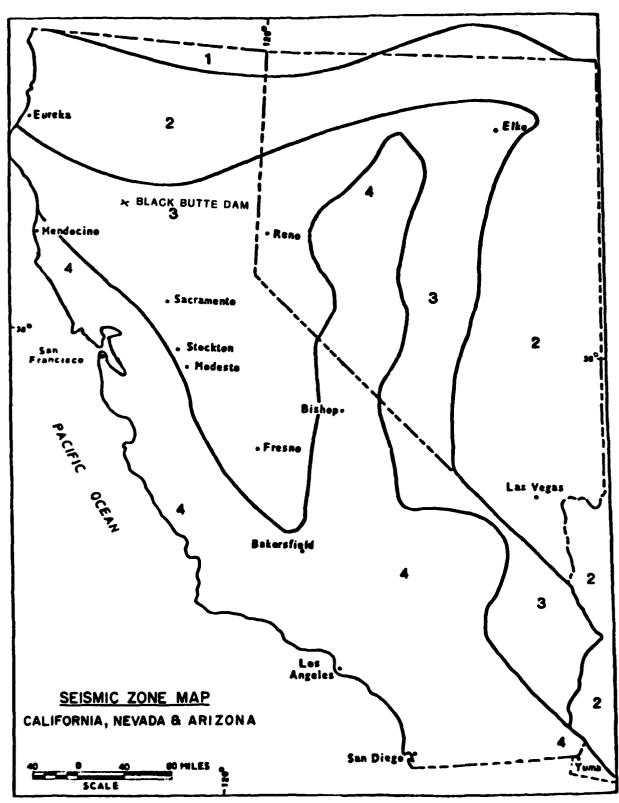
#### EXECUTIVE SUMMARY

Black Butte Dam is located on Stony Creek, a tributary of the Sacramento River, about 9 miles northwest of Orland, California. The Corps of Engineers completed construction of Black Butte Dam in 1963. In accordance with ER 1110-2-1806, "Earthquake Design and Analysis for Corps of Engineers Dams," dated 30 April 1977, an earthquake stability analysis was performed on the dam in 1983. The seismic loading for this analysis was obtained from a geological and seismological study estimating a design earthquake and ground response occurring on a nearby fault 2 to 6 kilometers (km) from the dam. The authors of that study, Drs. Bolt and Seed of the University of California, recommended a site acceleration of 0.45 gravity from a source tectonic structure like the Black Butte Fault or Willows Fault system as defined in literature and on maps of California geology.

Using a dynamic method of analysis it was concluded that the dam may experience significant movements under the given earthquake loadings, resulting in the possibility of slumping in the upstream portion of the dam. The large available freeboard of the dam would, however, retain the reservoir. Because of this slumping, due to this design event, Seattle District, Corps of Engineers, was asked by the Sacramento District, Corps of Engineers, to locate and extensively study nearby faults and determine their capability to produce the design event. The ongoing study was authorized under the Federal Dam Safety Assurance Program.

Black Butte Dam is located within Seismic Zone 3 (see accompanying figure), a zone of expected major earthquake damage. As shown on geologic maps of California, the dam is located approximately 15 miles from the Coast Range Thrust, 14 miles from the Stony Creek Fault, 12 miles from the Paskenta Fault zone, 1 mile from the mapped position of the contested Black Butte Fault, and 15 miles from the northernmost end of the Willows Fault. The contested Willows Fault system is a multibranched fault system extending from the north end of the Willows Fault. A tectonic wedge, with thrust detachments at depth, has also been proposed as existing within the near field area around the dam.

Underlying the dam are folded sediments of the Mesozoic age Great Valley Sequence. Through the interactions of plate tectonics they were folded into a synclinal trough which emerged from marine to subaerial environment. During folding, crustal shortening took place with the development of several deeper level detachments. Formative tectonic stress stopped in the interior valley of California about 3.2 million years ago. The present stress regime is considered an extension, similar to that occurring in the Basin and Range Province to the east, but modified by compression caused by plate interaction between the Pacific and North American plates along the San Andreas Fault zone. The interior valley is marked by a relative seismic quietness when contrasted with the very active western Coast Ranges margin and less active Sierra Nevada eastern border.



EXECUTIVE SUMMARY FIGURE

In past historic time, the site area has been shaken moderately by two earth-quakes. The 1903 Willows event (estimated epicentral Maximum Mercalli (MM) intensity of VII) most probably was not on the nearby Willows Fault but miles to the east of the fault. Historic location for this event is in the central part of the valley, but early settlement diaries indicate that the event was not felt along the North Fork of Stony Creek. The 1906 San Andreas event (epicentral MM intensity of XI) was on the distant San Andreas system. In the site area, this event had an estimated felt intensity of MM V or less.

Faults shown on present day maps resulted from early reports on California geology. For instance, the Black Butte Fault was accepted by the Corps of Engineers (1963) during site investigations for the dam and reservoir based on 1931 vintage reconnaissance work. Prior to this study there was no consensus on detail or activity concerning the Black Butte Fault; substantial investigations of others were divided in opinion about its existence. In addition, numerous faults mapped as short segments were lumped together and proposed to represent a multibranched fault system called the Willows Fault system.

The present study consists of extensive fieldwork and analysis of surface geology, subsurface geology, and geophysical data about the subsurface. Our evaluation of all this data has led to the conclusion that the Black Butte fault and Willows Fault system as proposed do not exist within 10 miles around the dam. Small distinct and separate faults might exist further away.

With elimination of the Black Butte Fault and Willows Fault system near the site, other sources were looked into. The tectonic capability of doubly plunging anticlinal folds in the subsurface near the midvalley axis was assessed. Possibly these folds are evidence for tectonic wedging and thrust faulting at depth, but this evidence is hidden from view. No subsurface evidence of faulting was found to account for a drag folding origin of the anticlines, and they have no seismogenic history.

Based on minor offsets along a few young Holocene terraces on a Stony Creek Fault segment near Thomes Creek, 27 km from the dam, magnitude 6.5 was chosen as the maximum (credible) event. The source is on the youngest segment of the fault, as this is judged to be the only capable segment.

Because of the few number of earthquakes experienced and the limited historic earthquake record in Glenn and Tehama Counties, calculating the exceedance chance for the maximum event during the project life span becomes an academic exercise, as does estimating earthquake recurrence levels. Instead, historic experience in northern California, as a whole, indicates a local magnitude of 5-3/4 as the maximum event in areas away from the San Andreas Fault zone during any 100-year period. Therefore, this level of magnitude is recommended to be a substitute for a statistical event and to serve as the 100-year probable earthquake. The source for this event is either near the Coast Ranges-Great Valley contact or east of the Willows Fault that lies southeast of the dam.

Guidance, opinion, and onsite review of the study were supplied by professional peers Dr. Roy Shlemon and Mr. Alan L. O'Neill. Conclusions arrived at in this study have their concurrence. Based on the surface geology, including two extensive undeformed Holocene terrace sequences and the subsurface stratigraphy, capable faults were not found to exist near the dam. We recommend exclusion of the Black Butte Fault and Willows Fault system from all future design considerations.

#### SECTION 1. INTRODUCTION

- 1.1 Authority. General earthquake risk and seismic hazard studies are authorized under ER 1110-2-1806, dated 16 May 1983. The Sacramento District, Corps of Engineers, requested this study by Memorandum of Understanding between Seattle and Sacramento Districts dated 1 May 1984. This report was prepared by Seattle District, Corps of Engineers, following guidelines of ER 1110-2-1806.
- 1.2 Purpose. All dams constructed by Federal agencies are reviewed for their performance based on current state-of-the-art civil engineering design. This review includes the reaction of the dam to seismic loading. The purpose of this report is to document the results of evaluation of suspected faults capable of generating earthquakes close to Black Butte Dam. The effects of close "capable" faults on the dam are reported in terms of seismic loading, stated as maximum magnitude and distance between the source and the dam.
- 1.3 Previous Studies. The Sacramento District performed various earthquake hazard and engineering studies for Black Butte Dam during 1982. These studies and preliminary results from a dynamic analysis of Black Butte Dam are reported in a series of in-house documents released in October 1983. Emerging from those studies was a single design event with maximum magnitude of  $M_L = 6.0$ . At the time of the study there was no clear consensus concerning activity on nearby faults, thereby generating the prudent procedure of accepting the moderate level earthquake on the closest fault and awaiting further confirmation on that fault's existence. This magnitude 6.0 event was localized on the contested Black Butte Fault producing a site acceleration of 0.45g.
- 1.4 Scope. This study assesses the capability of close faults, those known to the Corps of Engineers as well as those implied by others. The study included a review and summary of seismic history and a data catalogue. It includes the results of a lineament analysis employing a study of remote sensing and photogeologic imagery. It contains a regional structural analysis and fault evaluation. This study relied heavily on relative age of morphostratigraphic units and landforms to evaluate Quaternary faulting near Black Butte Dam (see Harwood and Helley, 1980; Harwood, 1983; Wentworth and Others, 1984a). It includes a reconstruction of Quaternary tectonic history using field analysis of geomorphic-stratigraphic landforms and units. Existing geologic literature and mapping was compiled and reviewed for a study area within a 50-mile radius, centered on the dam.
- 1.5 Investigation and Performance. Field geologic study was accomplished during May through November 1984. This represents a 23-week, two man-day field season. Field mapping of structure, stratigraphy, and morphostratigraphic units was performed at a scale of 1:24,000 (1 inch = 2,000 feet) on six United States Geological Survey (USGS) topographic quadrangle maps of the 7-1/2 minute series. When possible, we employed field review and discussion with geologic authorities. The study of geology and seismology around Black Butte Dam was assisted by the following contracted services:

- o Northwest Geophysical Associates, Inc., 1984, Magnetic Modeling of Black Butte Dam Report and supplemental letter, DACW-67-84-M-1591, NCA.
- o Endacott and Associates, 1984, Discussion and Results of Shallow Seismic Reflection Surveys at Black Butte Dam Technical letter with time sections, DACW-76-84-M-1417.
- o Harlan Miller Tait, 1984, Imagery and Photogeologic Analysis of Major Faults, Black Butte Dam, Orland, California Report, DACW-677-84-0048, HMT.

Guidance, opinion, and onsite review of the study problems were supplied by professional peers Dr. Roy J. Shlemon and Mr. Alan L. O'Neill. We gratefully acknowledge their contribution to the study. Shell Exploration Limited and other oil exploration concerns supplied valuable information and conversation about the acoustical subsurface character in the study area.

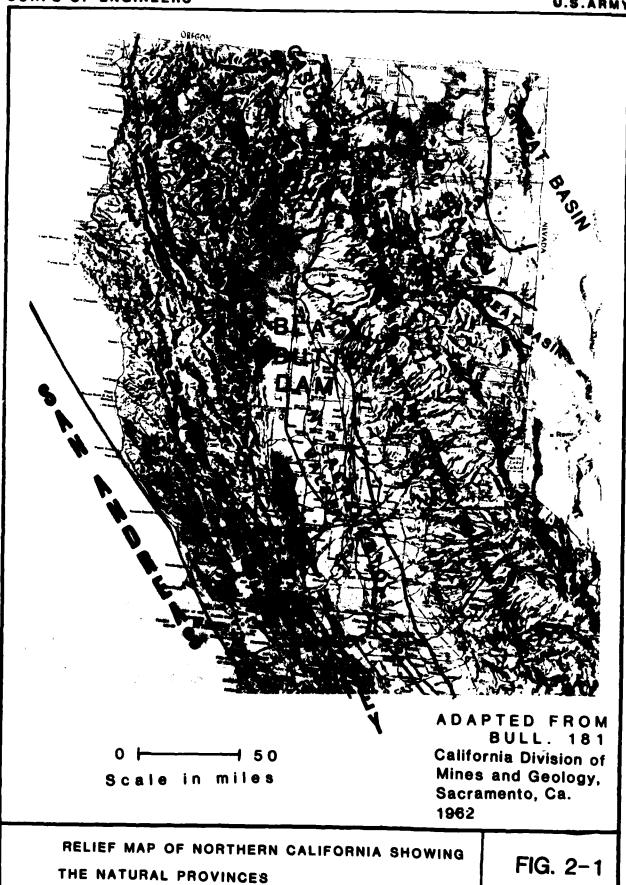
1.6 Project Description (see location and vicinity map on figure 1-1). Black Butte Dam, a flood control and water conservation structure, is located on Stony Creek about 9 miles northwest of Orland, California. The project has a gross pool capacity of 160,000 acre-feet, and was completed in 1963. This dam is a zoned earthfill structure with a crest length of 2,970 feet, a crown width of 20 feet, and a maximum height above streambed of 140 feet. The embankment volume is approximately 2.6 million cubic yards (c.y.) and serves as a closure to a water gap in Orland Buttes. Main project features are shown on figure 1-2. The embankment consists of a central impervious core, transition sections, pervious sections, and random rockfill sections. The embankment spans across streambed alluvium to rock abutments. The core is founded on rock through alluvium. Six earthfill dikes are located at various places around the lake to contain the high pool. The outlet works are founded on and through rock. The spillway is a broad-crested, unlined channel cut into rock.



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#### SECTION 2. GEOLOGY

- 2.1 Geologic Setting. Northern California is divisible into geologic provinces that are clearly marked by differences in morphology, geologic structure, and stratigraphy (see figure 2-1). The project is located in the northern Sacramento Valley portion of the Great Valley (GV) Province. Mountain provinces surround the Black Butte study area. To the north are the Klamath Mountains, to the west and southwest are the Coast Ranges, to the north and east is the Cascade Range, and to the east is the Sierra Nevada Range.
- a. Sacramento Valley. The Sacramento Valley is in the northern part of the Great Valley of California; a nearly flat alluvial plain extending south some 450 miles from the Klamath Mountains on the north. The valley width of 50 miles is seldom broken in its featureless profile. The only elevations of prominence on the valley floor are Marysville (Sutter) Buttes, a Pliocene volcanic plug, and the Dunnigan Hills, a Pliocene sedimentary heap. The Cenozoic sediments and volcanics rests atop the Great Valley Sequence. The subsurface of the valley is a geosynclinal east-west-trending warp with a north-south-trending deep fold axis along the western valley edge.
- b. Coast Ranges. West of the Great Valley are many mountain masses juxtaposed by northwest-trending strike-slip faults. Two principal core complexes are present in these mountains: Jurassic-Cretaceous eugeosynclinal assemblages (Franciscan) and Early Cretaceous granitic and metamorphic rocks. The two are not related. The Coast Ranges are structurally complicated, owing their existence to plate collision and transform movement throughout their history.
- c. Klamath Mountains. The Klamath Mountains consist predominantly of marine arc-related volcanic and sedimentary rocks of Paleozoic and Mesozoic ages. Ultramafic and ophiolitic rocks are also important rock types. Many areas of the Klamath Mountains are intruded by Jurassic granitic plutons. Rocks of the Klamath Mountains are not exposed in the project area; however, they were a source of detritus for the Great Valley Sequence.
- d. <u>Cascade Range</u>. In the northeasternmost part of California and extending into Oregon and Washington is the Cascade Range (including the Modoc Plateau). The dominance of Pleistocene/Holocene constructional volcanic landforms and block faulted basins characterizes this region.
- e. Sierra Nevada. Along the eastern edge of California's Great Valley is a huge batholithic landmass and thick sequence of Upper Paleozoic volcanic strata. Old island arcs thrust onto a Paleozoic continental margin were heavily intruded and plutonized through a Mesozoic Era plate collision. The Sierra Nevada core resulted as an elongate block of crust. Later this block broke free along its eastern edge and rose upwards thousands of feet tilting westward and forming this dominant mountain range on California's landscape.

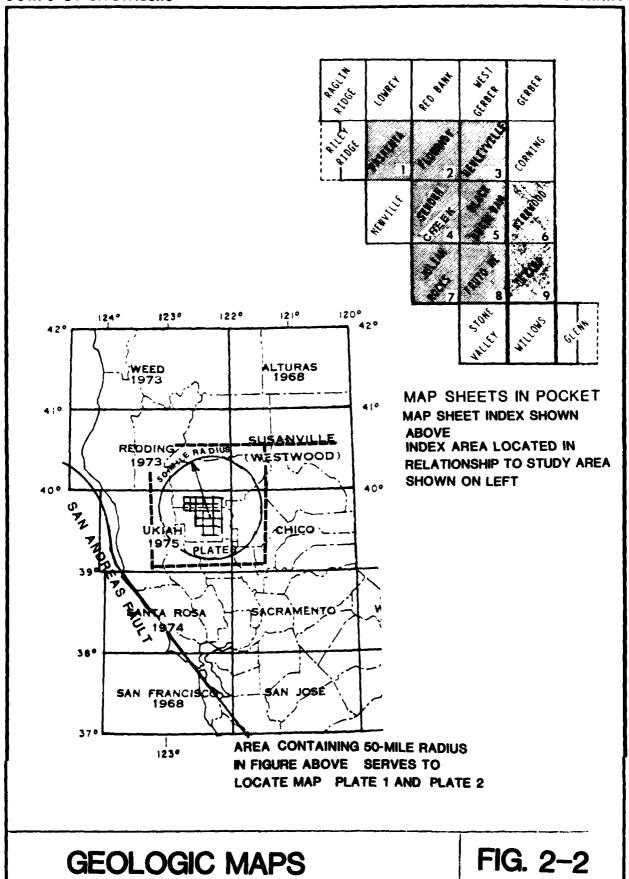


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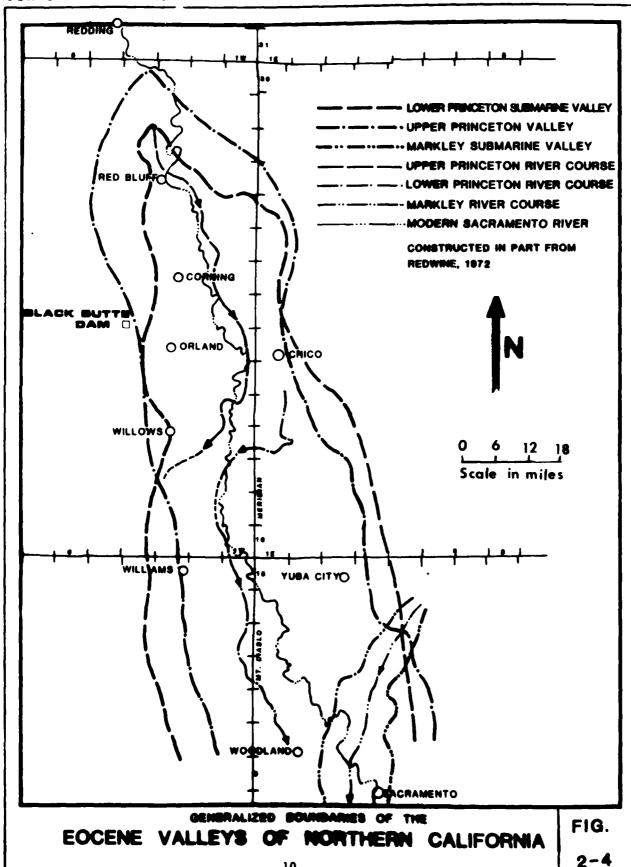
f. San Andreas Fault. In California, faults are numerous, but their hazard varies according to location. With the exception of the Great Valley and the San Andreas (SA) fault zone, provinces surrounding Black Butte Dam have no overwhelming influence on seismic hazard at Black Butte Dam.

The San Andreas Fault is California's most spectacular and best known plate border. The San Andreas Fault zone has had a series of large recurring earthquakes. This fault strikes N350 W along the western flank of the Coast Ranges in nearly a straight line. Extending southward from Cape Mendocino, it has 650 miles of continuous length. The main San Andreas Fault lies within 125 miles of the dam. Inland from the fault is a zone of sympathetic shearing and faulting that mimics the San Andreas trend. The sympathetic faults of the San Andreas lie as close as 57 miles west of Black Butte Dam.

- 2.2 Regional Geology. Quaternary geology and lineaments were compiled on a 1:100,000 scale base map shown as plate 1. The stratigraphy, morphology, and structure within 50 miles of Black Butte Dam were compiled on portions of three State of California 1:250,000 scale geologic maps. Figure 2-2 serves as an index to geologic maps on plates 2, 3, and 4. Figure 2-2 also provides a key to nine maps in the pocket showing more detailed geologic mapping.
- 2.2.1 Pre-Quaternary Stratigraphy. Figure 2-3 accompanies this discussion of stratigraphy. Also refer to plates 2 through 6. Pre-Quaternary stratigraphy is extensively discussed in appendix B, and should be consulted for a better understanding of the stratigraphic section and stratigraphy identified on the figures and plates. Northern Sacramento Valley stratigraphy represents deposition on an emerging crystalline continental margin. Successions of turbidites, shelf deposits, shallow marine, and subaerial deposits are present throughout the valley. Sedimentary deposits overlie a basement consisting of down-thrusted Franciscan Complex and ophiolites in the west and Sierran granodiorites and Paleozoic metasedimentary strata in the east.
- a. The rocks overlying the basement are thick units known as the Great Valley Sequence (up to 24,000 feet thick). Generally the Great Valley Sequence is divided into three series: Knoxville series (Jurassic), Shasta series (Lower Cretaceous), and Chico series (Upper Cretaceous). Each series has several formations. In the project area the depth to basement is unknown. The project and lake are located on top of exposures of the Chico series. Formations present at the damsite include: Forbes mudstone, Dobbins shale, and Guinda sandstone. On the lake shore Funks, Sites, Yolo, Venado, and Boxer (Julian Rocks conglomerate and Clark Valley mudstone) Formations are present in outcrop. Further east along Stony Creek the Lodoga Formation is present (Shasta series).
- b. A major erosional unconformity marks the boundary between the Cretaceous formations and overlying Tertiary deposits. This unconformity represents a period of time in which a huge submarine valley (lower Princeton Submarine Valley) was eroded into the Cretaceous units in the area of the present Sacramento Valley (see figure 2-4).



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- c. Next group in overlying succession is the nonmarine group. An important stratigraphy is the one which makes up the formidable buttes that the dam closes at Stony Creek gap. This strata is a locally occurring series of volcanic sandstone, mudstone, and conglomerate capped with basalt. The Black Butte Formation, as it is found at the buttes in Stony Creek gap, became host to a thin basalt flow or canyon flow about 15 million years ago. The Lovejoy basalt occurs sporadically throughout the valley in wells, presenting in section a sinuous canyon-like nature to its occurrence.
- d. Next in succession is the oldest alluvium, representing a flood of clastic material occurring during the Pliocene that descended from the rising western mountain masses.

Tehama Formation. In the greater project area, the Tehama Formation consists of greenish gray, occasionally pale yellowish gray, sandy, tuffaceous siltstone, lenticular channelized interbeds of conglomerate and sandstone, and locally abundant conglomerate beds. Where near the surface, the conglomerate commonly takes on a bright red-brown color presumably due to weathering on the surface of clasts and the matrix. Where the weathering is incomplete, the conglomerate takes on a mottled appearance. This bright red-brown coloration was also observed locally in conglomerate beds exposed at depth by larger creeks west of the damsite and may represent buried geosoils (see glossary). The Tehama Formation is moderately consolidated and usually uncemented. It lies unconformably on all older formations and was deposited in a fluvial environment of an ancestral (and similar) Sacramento Valley of low relief. The Tehama deposits, derived from the Coast Ranges and Klamath Mountains, were deposited along the west and north sides of this ancestral valley. The Tuscan Formation deposits, derived from the Sierra Nevadas, are coeval with the Tehama, and the two interfinger in the central area of the present Sacramento Valley where they reach a thickness of over 2,000 feet. The Tehama Formation in the general project area takes on the appearance of a large fan from ancestral Stony Creek. The north toe of the fan lies near Black Butte Dam 20-25 miles to the south, the south toe lies about 1 mile east of Willows, and the remnant apex lies aproximately 1 mile southeast of Julian Rocks. The Tehama Formation is late Pliocene in age, bracketed by age dates of 3.4 million years (Evernden and others, 1964) and 1.5 million years (Lydon, 1968) based on fossil contents. The Tehama Formation is relatively thin west of the damsite but thickens rapidly to the east.

Nomlaki Tuff Member. The Nomlaki Tuff Member is 10 to 50 feet thick and found near the base of the Tehama Formation over a large area of the northwestern Sacramento Valley. It is a dacite tuff, locally white, salmon pink, or light gray, massive, and coarse grained. Age dates indicate this tuff is 3.4 million years old. It is exposed consistently near the western margin of the valley and throughout the mapped area of the Tehama Formation. In the course of this study, presumably reworked Nomlaki Tuff was noted in wave cut exposures of the Tehama Formation at Orland Buttes Recreation Area. An additional exposure is noted in the Black Butte Foundation Report (Corps of Engineers, 1983) between the North and South Forks of Stony Creek (sec. 21, T. 22 N, R. 5 E.). These occurrences verify the relatively thin cover of Tehama Formation and the shallow depth to Cretaceous rocks in this area.

### 2.2.2 Quaternary Stratigraphy and Geomorphology.

- a. General. The cycle of uplift, erosion, downwarping, and reestablishment of ancient Sacramento River fluvial deposition, responsible for the deposition of the Neroly and Tehama Formations, was operative at least twice in the Quaternary. The Red Bluff Formation represents a Quaternary cycle and the modern alluvium of the Sacramento River another. Erosion of Tehama and older formations has been occurring along valley margins since early Quaternary time. Through various base level adjustments, numerous morphostratigraphic surfaces have developed along the Quaternary drainages. Modern stream locations are similar in nature to paleo-drainages from the highland areas west of the project. The last major sedimentary cycle (Tehama aggradation) has built an extensive coalescing alluvial fan outward from the Coast Ranges. Erosion is presently moving material down a slope of transportation. About 1.25 million years ago, fluvial deposition ceased, perhaps resulting from a Pleistocene orogeny west of the study area (Earth Sciences Associates (ESA), 1980; Steele, 1979, 1980; Harlan Miller Tait, 1983, 1984). Between about 1.25 million and 0.5 million years ago, extensive "high terraces" were cut across and into the Tehama Formation by ancestral drainages of Stony and Thomes Creeks and the precursor of other east flowing streams.
- b. Red Bluff Formation. The Red Bluff Formation consists of deposits very similar to the Tehama Formation; however, the Red Bluff Formation is generally of uniform deep brick-red color, much coarser, and contains less rounded clasts. These features can usually be used to differentiate the formations where both exist. The principal occurrence of the Red Bluff deposits is north and northeast of the Black Butte Dam area. It reportedly attains a thickness of 100 feet in the vicinity of Redding but rapidly thins to less than 50 feet to the south (Anderson and Russell, 1939). South of the general area of Sehorn Creek only remnant patches of Red Bluff deposits remain, and they become mappably indistinguishable from similarly colored gravels of the Tehama Formation. The Red Bluff deposits unconformably overlie the Tehama Formation and the Cretaceous rocks where the Tehama is absent. The Red Bluff was deposited on an erosional surface in an alluvial environment very similar to that of the Tehama and Neroly Formations.
- c. River Fluvial Terrace Systems. Between about 1.25 and 0.5 million years ago and continuing to the present, an extensive terrace system has been formed along major valley drainages. This system is composed of nested terrace levels which are of regional extent and are mappable. Terrace levels were developed in response to episodic base level lowering resulting mainly from climatic changes (ESA, 1980; Huston, 1973; Steele, 1979; Harlan Miller Tait, 1984; and Shlemon, 1967a, 1984, Per. com.). Terrace levels, either cut or fill surface, and extensive "high" surfaces are correlated along the Sacramento Valley. They have been dated by association with major climatic changes during the Quaternary. Terrace levels are distinguished by morphology and are calibrated by soil stratigraphy to distinct time periods. These morphostratigraphic surfaces provide the best means of evaluating the history of folds and raults mapped across these units. Maps 1 through 10 (in the pocket) indicate that neither faults presently mapped or inferred by others, nor folds, or

lineaments in the study area, have disturbed the units of age Q6 or less (350,000 years before the present).

Morphostratigraphic Units. Ten years ago few engineering geologists understood terms such as morphostratigraphic units, geosoils, or soil-stratigraphic nomenclature; however, these terms have become the geologist's "clue kit" to aid him in solving recent fault history and defining displacement when faults lie across or underneath morphostratigraphic units. These terms are described in the accompanying glossary. Morphostratigraphic units in the Black Butte area are nested terrace levels continuous to downstream channel deposits developed in the youngest prominent fan heads (Steele, 1980; Huston, 1973; Shlemon, 1980; Shlemon and Begg, 1972).

Along the eastern foothills of the northern Coast Ranges the basic geomorphic process has been the development of a slope of transportation by dissection and lowering of the Tehama surface, resulting in the exhumation of underlying folded sedimentary rock. Resurrected strike valleys and thrust zones are exposed in the area of a once existing alluvial plain (i.e., maximum aggradation of Tehama fans) lapping onto the mountain front. On this slope, pedological profiles formed in the last 0.5 million years when given a stable land surface, climate, and vegetation. Where there has been geomorphic stability through time, distinct soil horizons have developed that can be measured and described as stages of profile development. Along Stony Creek, zonal soils develop genetic horizons (such as the argillic horizon) that allow excellent resolution of chronological sequence. Parent material and relief influence zonal soil formation, however, through time past, in California in general, and in this study area, zonal soil formation processes have mostly been dominated by climate and biological forces.

In the foothill region, parent material is of mixed mineralogy and has been inplace for a long time. Materials from the original Coast Ranges source are sedimentary, metasedimentary, and basic volcanic in origin. The oldest materials above bedrock are mixed texture alluvium and gravel channels present as oldest alluvial fans. Remnants of channels are now exposed on high benches, ridge lines, and side-hill terraces as caps (fragipan) containing gravel and cobblestone. These near-level surfaces can be mistaken as "high terraces" suggestive of braided channels on the oldest upland plain. In some instances, as at sec 28, T. 24 N., R. 4 W., these channels are exposed in cuts as exhumed gravel channels of the Tehama. These channels have been weathered long enough in time to develop an abrupt duripan.

Very often, recently exposed nonweathered Tehama materials contain gravel channels that also have duricrust (calcrete and silcrete). Old coarsetextured overwash fans form at the base of secondary drainages dissecting the upland surfaces. The largest ones are seen in poorly sorted material exposed through incision of Thomes and Stony Creeks. Moderately old gravelly alluvium has been deposited on alluvial fans by local intermittent streams draining the old alluvium of the foothills. There, surfaces are moderately to deeply weathered. Moderately old sedimentary alluvium has been deposited from intermittent streams draining the exhumed strike valley and older, fine-grained fan

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materials. Younger alluvium (but older than recent flood plain) has been deposited by Stony Creek as flood plain and coarse fan head gravel channels due to changing base levels because of shifting basins in the lower reaches. Recent alluvium is varied-texture channel sand and gravel and recent flood plain sediments laid down by Stony and Thomes Creeks.

Soil Stratigraphy. Fine tuned division of the Quaternary is most often related to worldwide climatic change and the accompanying alternation of vegetation, sedimentation, and hydraulic regime. These are, in turn, related to interglacial and interstadial epochs. Most fluvial sediments in the Central Valley were laid down during glacial/pluvial time, and most soil formation occurred during preceding and following interglacial and interstadial epochs. On any stable geomorphic surface, exemplified by terraces flanking Stony and Thomes Creeks, soil profile began to form. With the passage of time, distinct horizons slowly formed and, based on physical and chemical characteristics, can be qualitatively distinguished by relative degree of soil profile development: (1) minimal (slight), (2) medial (moderate), and (3) maximal (strong) (Begg 1968, Gowan, 1967). Relative soil profile development can be calibrated to epochs of Quaternary climatic change by association with oxygen isotope chronology. Figure 2-5 shows relative soil profile; schematically mapped soil series particularly applicable to the study area are shown on table 2-1. They are distinguished mainly by relative development of the B2t (argillic) horizon and are divisible into subcategories by differences in parent material and texture.

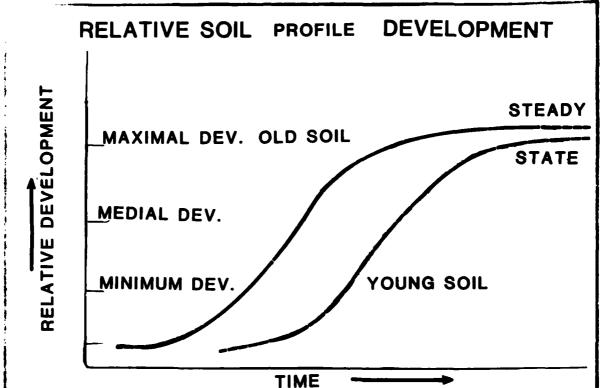
Units. In this study, the major "intermediate" and "low" terrace levels are designated from younger to older as morphostratigraphic units Q1 through Q6. The use and application of the concept of morphostratigraphic units follows the approach discussed in detail in Harlan Miller Tait (1983, pages 37 to 43). The lowermost four or five terraces are correlative to the Modesto and Riverbank Formations of the northeastern San Joaquin Valley (Davis and others, 1959). These formations were extended to the southeastern Sacramento Valley (Shlemon, 1967b; 1972), and later to the western and northern Sacramento Valley (Harwood and others, 1980; Helley and others, 1981; Helley and Jaworowski, 1985). Steele similarly recognized terraces in the study area, naming them after soils series forming on them, e.g., Orland, Yolo, Arbuckle, Perking terraces (Steele, 1980, p. 18).

Many minor discontinuous intermediate units are also preserved in the study area but have been identified only locally by observation in the areas. Detailed field checking of morphostratigraphic units was conducted on Stony Creek and Walker Creek. Reconnaissance checking was performed on Thomes Creek.

Channels. Units Q1 through Q5 have equivalent age channels associated with Stony Creek, downstream from Black Butte Dam. These channels were field mapped in this study using prior mapping by Shlemon, et al. (1976). The channels reflect the migration of Stony Creek across its fan. Soils developed on morphostratigraphic units Q1 through Q5 are of the same pedological development as those developed on equivalent-age channels C1 through C5.



U.S.ARMY



# Typical development of soil groups and profiles in the study area

SUBGROUP	PROFILE	DEVELOPMENT
Typic Xe <b>ro</b> fluvents	Ap-C1-2C2	Minimal
Typic Xerothent	C1-C1011C3	Minimal
Typic Haploxeralfs	Ap-A12-B1t-B2t	Minimal
Typic Haploxeralfs	Ap-Blt-B2lt-B22t-B3tca-Clca	Medial
Typic Palexeralfs	O1-A1-A3-B21t-B22t-B3t-C1	Maxima.
Typic Palexeralfs	Ap-A1-A3-bit-B2t-B31t-B3	Maximal
Abruptic Durixeralfs	A1-A3-B21t-B22qm-Cm1-Cm2	Maximal

#### Subordinate Distinctions

a-accumulation of
carbonates
q-accumulation of silica
t-accumulation of silicate
clay
x-fragipan character
p-plowed
m. Cementation or induration

Classification after Begg (1968) and Gowans (1967).

SOIL DEVELOPMENT IN STUDY AREA

FIG. 2-5

TABLE 2-1 SOILS OF THE SACRAMENTO VALLEY APPLICABLE TO BLACK BUTTE DAM STUDY AREA

Soil (1) Texture	Soil Series	Parent Material	Argillic Development	Pan	Suborder	Youngest (2) State Age
Coarse Textured Soils	Cortina	Coarse Textured Recent Alluvium	Minimal	None	Typic Xerofluvents	Mod e rn
	Arhuckle	Gravelly or Cobbley Alluvium	Minimal	None	Typic Haploxeralfs	10,000
	Perkins	Gravelly or Cobbley Alluvium	Medial	None to Clay	Mollic Haploxeralfs	125,000
	Corning	Gravelly or Cobbley Alluvium Old Terrace	Maximal	Clay	Typic Palexeralfs	125,000
Medium Textures Soils	Orland	Flood Plain Outwash from Sedimentary Rocks	Minimal	None	Typic Xerorthents	Modern
	Wyo	Sedimentary Alluvium	Minimal	None	Mollic Haploxeralfs	з 10,000
	Tehama	Old Alluvial Fans	Medial	None	Typic Haploxeralfs	10,000
	Hillgate	Old Fans and Low Terraces	Maximal	None to Clay Pan	Typic Haploxeralfs	125,000
Fine Textured Soils	Yolo	Recent Alluvium of Sandstone	Minimel	None	Mollic Haploxeralfs	10,000
	Zamora	Mixed Allu- vium on Alluvial Fans		None	Typic Haploxeralfs	10,000
	Kimball	Terraces	Maximal	None	Mollic Haploxeralfs	125,000
Cemented Soils	Moda	Fine Grained Alluvium	Maximal	Iron- Silica	Abruptic Durixeralf	s 125,000
	Redding	Gravelly or Cobbley Alluvium	Maximal	Iron- Silica	Abruptic Durixeralf	s 250,000
	Red Bluff	Old Fans	Maximal	Iron- Silica	Abruptic Durixeralf	s 250,000

 <sup>(1)</sup> Soil series after Begg (1968) and Gowans (1967).
 (2) Estimated minimal ages after ESA (1980); Harlan Miller Tait (1984); Shlemon and Begg (1972); and Steele (1979).

- Q1 This unit consists of the youngest topographically low fill deposit, probably less than 10,000 to 12,000 years old, associated with all the larger drainages in the study area. It includes the active channel and flood plain, the first emergent terrace, which is probably alluvial land and of essentially modern soils, including Orland, Yolo and Cortina.
- Q2 This unit consists of the first preserved "low" terrace remnants above the active channel and flood plain. It is discontinuous but locally present along most major drainages and locally is more extensive than Q3. Unit Q2 is generally used for row-crop agriculture and is characterized by soils less than 10,000 years old, including Zamora, Wyo, and Cortina.
- Q3 This unit consists generally of the most extensive "low" terrace remnants present along most major drainages. It is best preserved on the north side of Thomes Creek. It is locally covered by a thin veneer of young alluvial fan deposits from adjacent sidestreams graded to its surface. Q3 is generally used for row-crop agriculture, but where uncleared is sparsely tree covered. It is characterized by minimally to medially developed soils, approximately 35,000 to 50,000 years old, including Arbuckle, Tehama, and Pleasanton soils.
- Q4 This is locally the most prominent unit in the study area associated with the present drainage system, although it is less prominent in this area than in the northernmost Sacramento Valley. It is generally the most continuous and widest unit, and extends up larger tributary drainages. It may have a distinct hummocky micro-relief, but in the study area it seldom displays this character. It is best preserved on the north side of Thomes Creek. It is locally covered by a thin veneer of young alluvial fan deposits from adjacent sidestreams graded to its surface.

Unit Q4 has moderate relief and is slightly dissected. It characteristically has more minor, discontinuous, intermediate levels than others, particularly on the outside of Q4 age meanders, and locally preserves remnants of now abandoned drainage networks related to its formation. Intermediate levels in these areas include some slightly higher and older and some slightly lower and younger than Q4. Q4 is generally used for pasture, crops, and range, and where uncleared is sparsely to moderately tree covered. It is characterized by medially developed soils approximately 80,000 to 125,000 years old, including the Hillgate, Arbuckle, Tehama, Perkins, and young Kimball soils in the study area.

Q5 - This unit consists of remnants of the first "intermediate" terrace level preserved in the study area. It is generally moderately dissected with moderate relief and is generally narrower and less continuous than Q4. Q5 remnants usually have a distinct hummocky micro-relief on Corning soils; otherwise micro-relief is seldom seen on Q5 in the study area.

Unit Q5 is generally used for pasture and range, with vegetation consisting of annual grasses and forbs and scattered to locally abundant trees. It is characterized by maximally developed soils or equivalent surfaces approximately 200,000 to 250,000 years old developed on sediments greater than 250,000

to 300,000 years old. Commonly associated soils include the Perkins, Red Bluff, Kimball, and Corning in the study area.

Q6 - This unit consists of remnants of a second, topographically higher "intermediate" terrace. It generally is moderately dissected with undisturbed remnants typified by distinct hummocky micro-relief and is somewhat narrower and less continuous than Q5. However, it is readily differentiated from Q5 and occurs along most major drainages in the study area.

Unit Q6 is generally used for pasture and range, with vegetation consisting of annual grasses and forbs with scattered to locally abundant trees. It is characterized by maximally developed soils or equivalent surfaces approximately 300,000 to 350,000 years old. Commonly associated soils are the Redding-Newville and Corning in the area of study.

Red Bluff Pediment/Redding High Flood Plain. The term "Redding High Flood Plain" was coined by Steele (1979) to designate the highest continuous, generally gently eastward-sloping surface preserved in the northern Sacramento Valley west of the Sacramento River. The name "Redding" was derived from the extensive soils series of that name mapped on several high-level geomorphic surfaces in Glenn, Tehama, and Shasta Counties (Begg, 1968; Gowans, 1967; Klassen and Ellison, 1974). These surfaces were inferred by Steele to represent a once continuous, regionally extensive high erosion surface, or pediment, cut across the Plio-Pleistocene Tehama and Pleistocene Red Bluff Formations. Ilelley and others (1981) included the "Redding High Flood Plain" in the Red Bluff Formation, which they interpreted as a large, dissected alluvial fan which truncated the Tehama Formation as a pediment. More recently, Helley and Jaworowski (1985) referred to the surface as "Red Bluff Pediment".

Prior to Steele (1980), Begg (1968) postulated that the oldest fans in the study area were uplifted by tectonic action and subsequently dissected by Stony Creek leaving many disconnected terrace remnants. Steele (1980) attempted to reconstruct the predissected surface of the "Redding High Flood Plain" by drawing generalized contours connecting points of equal elevation on surface remnants. The resultant contour pattern displayed several localized deflections which were postulated by Steele to reflect tectonics within the last 1.25 million years. Helley and Jaworowski produced a Red Bluff age contour map in the study area just south and east of Black Butte Dam (Fruto N.E. quadrangle map) that shows a small east-west-trending dome in the area between Wilson Creek on the south and South Fork Walker Creek on the north. This small structure is one of several east-west-trending folds that their contour mapping suggests along the west side of the valley.

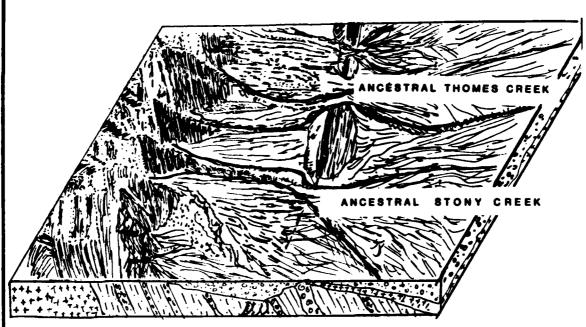
The concept that a once continuous, regionally extensive high erosion surface existed throughout much of the study area and entire Sacramento Valley is intriguing and, if true, would provide one method to evaluate recent tectonics. However, detailed photointerpretation, field mapping, and geomorphic interpretation of published soil surveys conducted as part of the Cottonwood Creek project study (Harlan Miller Tait, 1983) and this study indicate that the

highest surfaces preserved in the study area are not remnants of a once continuous, regionally extensive "flood plain" or pediment. Flood plain may be an incorrect usage of the term, as is pediment, for anything other than a constructional slope. Begg (1968) used alluvial plain; although the Tehama Formation might have fit this concept once, perhaps coalescing in a large semiunified "surface", individual fan surfaces in the formation were probably never regionally continuous. In fact, present remnants of the many high surfaces are representative of distinct fan levels localized by gaps in folded sedimentary bench land. Terraces were cut across and into the Tehama Formation in early Quaternary time between about 1.25 and 0.5 million years ago by ancestral drainages of Thomes Creek and Stony Creek and by precursors of many other east flowing streams in the northwestern Sacramento Valley. They became regionally graded to the ancestral Sacramento River. Figure 2-6 illustrates the maximum aggradation of Tehama fans and the onset of dissection by ancestral Stony and Thomes Creeks.

A further examination of the pediment concept is warranted. A pediment is a gently inclined plane (typically 1 percent in the southwestern United States) at the foot of a mountain front. It is formed by degradation and retreat of the mountain front, the headward source of sediments. The upper erosion surface is thinly veneered with alluvium while at the toe a greatly thickening wedge of clastic is graded to a closed basin or a through-going stream. The outermost zone of aggradation encroaches on the thin veneer or pediment by building of alluvium (Easterbrook, 1969). Pediments are dissected by changes in base levels. The Black Butte area does not fit the description of an undissected pediment as illustrated by figure 2-7.

North of Stony Creek (off of figure 2-7) the Tehama and Red Bluff overlie the bedrock wholly as a zone of aggradation. Thin wedges exist at Table Mountain, elevation 1,100 feet, where the Tehama is presumably the thinnest, but the wedge thickens rapidly eastward to more than 600 feet and persists for 3 miles west. (Table Mountain, sec. 25, T. 26 N., R. 5 W., to Occidental Petroleum Harris 9811 No. 1., sec. 23, T. 26 N., R. 5 W.).

The section sketch in figure 2-7 is south of Stony Creek across T. 21 N., R. 6 to 3 W. down the regional slope. For the section shown, the upper slope is greater than 100 feet per mile; somewhat more than 50 feet per mile is typical of southwest pediments. Since the Tehama dips at 2 degrees (250 feet/mile), this is not the original constructional slope but rather an adjusted slope reflecting degradation and/or slight regional tilting. Field evidence suggests that in places more than 80 feet of thick clastic wedges of Tehama exist on the higher slope areas, not suggestive of a veneer. Many areas in the Fruto and Stone Valley quadrangles contain outcrop and subcrop of Nomlaki Tuff, suggesting that the Tehama has been eroded to near the base of the formation. The lower fan area has intact Q5 and younger units as do the upper fan heads and subsequent valleys, suggesting that this is a 100- to 250-million-year-old surface at the oldest. We found no "high" terraces on Worthington Ranch (Fruto N.E. northern half). The Tehama is 600 feet thick at Worthington Oil Well (sec. 33, T. 24 N., R. 4 W.) indicating that this area could have been a consequent breach in the bedrock bench land probably serving to localize the

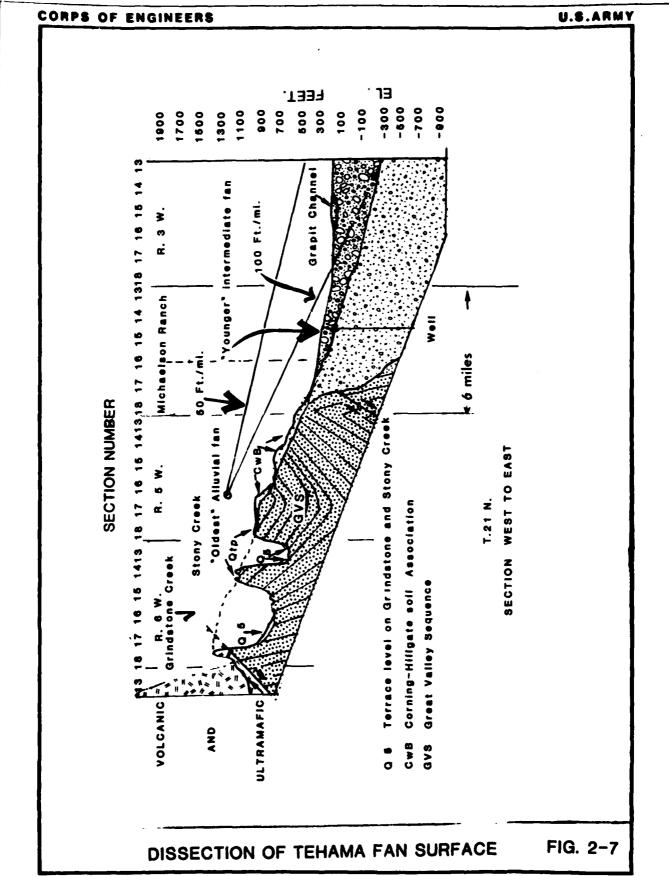


Maximum aggradation of Tehama fans are depicted. Resistant folded sedimentary rock forms ridgelines. Breached by erosion during an older cycle, ridgelines form benches that stand topographically higher than fan surfaces.



As presently occurring near Sheep
Mountain, Wyoming, breached ridgelines
serve to localize new fan heads. Toe
areas of the fans tend to coalesce into
a semiunified alluvial plain.

AGGRADATION OF TEHAMA FAN SURFACE FIG. 2-6



Tehama fan at the gap during its beginning. For at least the last 0.5 million years there have been adjustments in slope profile to a new base level of the Sacramento River. No evidence of the former regional alluvial plain exists. The remnant "high" surfaces mapped by Steele (1980) and Helley and Jaworowski (1985) include morphostratigraphic units Q5, Q6, and older than Q6. Because the remnant highest surfaces in the study area represent at least several distinct terraces formed at differing times and at different elevations, the surfaces cannot be meaningfully contoured. Localized deflections in any resultant contour pattern cannot, therefore, be necessarily inferred to reflect late Quaternary tectonism.

- d. Recent Deposits. Alluvial deposits in the study area, present in nearly all stream valleys, consist principally of material ranging from silt to gravel. Clay and boulders are rare. These materials have been derived locally from the Great Valley Sequence rocks, the Tehama Formation, and the Franciscan Complex. Thickness of the alluvium is quite variable but probably does not exceed 40 feet. Exposure of bedrock in creek bottoms is not common, while thickness of alluvium at the damsite is approximately 20 feet. Colluvial slopewash commonly interfingers with alluvium at the base of slopes and frequently is found as a wedge of material onlapping the morphostratigraphic surfaces. Minor alluvial fans typically are found where small intermittent creeks discharge onto the morphostratigraphic surfaces. Landslides in the study area are prominent on the west side of Orland Buttes as earth flows in the talus and underlying residuum below the steep basalt cliffs. Very few landslides were noted in the Great Valley Sequence rocks, and those were minor.
- 2.2.3 Structure. The geologic structure of the region is dominated by north-to northwest-trending folds and faults resulting from synclinal collapse of the Great Valley in Late Cretaceous time. Development of northeast, north-west, and east-west structures continued throughout Tertiary time. The dominant north-northwest-trending Willows Fault system proposed by Harwood and Helley (1982) is not supported by convincing evidence. Moreover, the absence of such a system is supported by data developed during this study.
- a. General. In the area studied, structures are grouped as to their form (faults or folds). Just after or during the time of accretion of the volcanic arc and forearc basin to the continent, the Coast Range Complex thrust eastward under the Great Valley sequence. Thereafter, additional crustal shortening occurred.

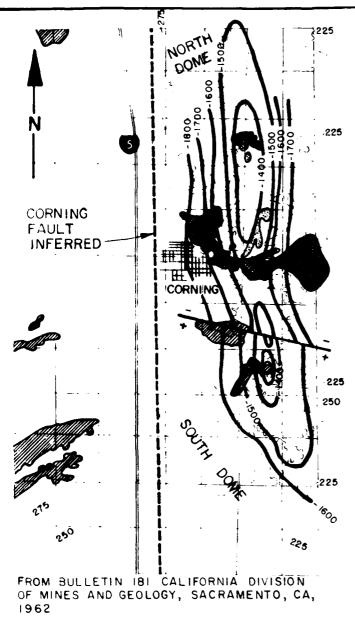
Folds and faults resulting from crustal shortening and the synclinal collapse of the Great Valley are the largest dominant structures. Later northwest, northeast, and east-west oriented faults occurred resulting from the change in stress field during the end of continental accretion and synclinal collapse, approximately 3 million years ago. There is no convincing evidence for a large branching northwest fault system in the area.

### b. Folds, Domes, and Arches (Plate 3).

Red Bluff Arch. This structural feature is a northeast-trending anticline located midway between Cottonwood Creek and the town of Red Bluff. The Red Bluff arch lies about 40 miles northeast of Black Butte Dam. The fold has been the subject of several studies; the most recent by Harlan Miller Tait (1963, 1984). This arch is underlain by the Tehama Formation and is capped by the Red Bluff Formation. Several levels of Q5 terrace on the arch flanks appear to be tilted and the region is believed to have undergone some tectonic instability about 250,000 to 300,000 years ago.

Corning Domes. Corning North and South Domes and Corning Dome are located in Tehama County along a north-south trend 1-mile east of Corning. They are north-trending anticlinal features defined in the subsurface. Prominent hills lie on the crest and western flank of the buried structure on an otherwise flat plain. This topographic and geomorphic expression drew initial geologic investigation to the area for exploration of gas deposits. After gas discovery, the anticlines became better defined and remain as longtime recognized structural features. Some over steepened stream gradients are present over the general area of the anticlines. Additional evidence from Plio-Pleistocene gas stratigraphic markers suggest early Cenozoic tectonic folding. Some investigators suggest that the presence of Red Bluff gravel on the flanks and Tehama-Tuscan Formation at the crest of the topographic hills indicates post-Red Bluff folding (within the last 0.5 million years, Harwood and Helley, 1982). However, the distribution of these deposits does not correspond with the location of the subsurface structural axial fold plane (see figure 2-8). Stratigraphic closure within the Tehama gas sand (unit 3) ranges from 300 feet in the north to 50 feet in the south (California Division Mines and Geological Bulletin 181). This sand is about 200 feet above the Upper Cretaceous Forbes contact with the Kione Formation. The north dome is separated from the south dome by a high angle, northwest-trending fault. Harwood and Helley (1982) cite well log data, earthquake pattern recognized by Marks and Lindh (1978). surficial mapping, and seismic reflection work by Seisdata as evidence that the Corning Dome (and Greenwood anticline, see paragraph below) was formed by east side up drag on a north-trending fault. Based on studies reported in the geophysics section (paragraph 2.4), the evidence cited for drag folding through faulting is far from conclusive. Well location on the domes/folds is orderly and spacially distributed with even density. In the area of the suspected fault, no wells exist. Closest wells are 3 miles west of the dome. earthquake pattern shown by Marks and Lindh (1978) is 10 miles east of Willows (near Glenn) and considerably south of structures discussed. The Seisdata lines examined by this study had considerable data dropout along the Interstate (I) 5 corridor, and very poor geophone coverage adjacent to structures. We were not able to confidently define structure along the I-5 corridor. This area is believed to correspond to the map location of Harwood and Helley's (1982) Corning Fault.

Multiple oil company proprietary seismic lines examined during this study, private discussion with oil companies, and discussions of others' work indicate



## **LEGEND**

# **SURFACE**



CORNING - PERKINS ARBUCKLE SOILS Q5 AGE



REDDING - RED BLUFF SOILS



TUFF (TUSCAN)

€200

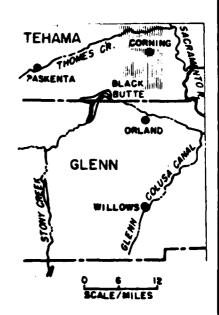
TOPOGRAPHIC CONTOURS IN FEET

# SUBSURFACE STRUCTURE



STRUCTURAL CONTOUR MAP OF CORNING DOMES AT BASE OF LOWER TEHAMA SAND CONTOUR ELEVATION IN FEET

SCALE IN MILES

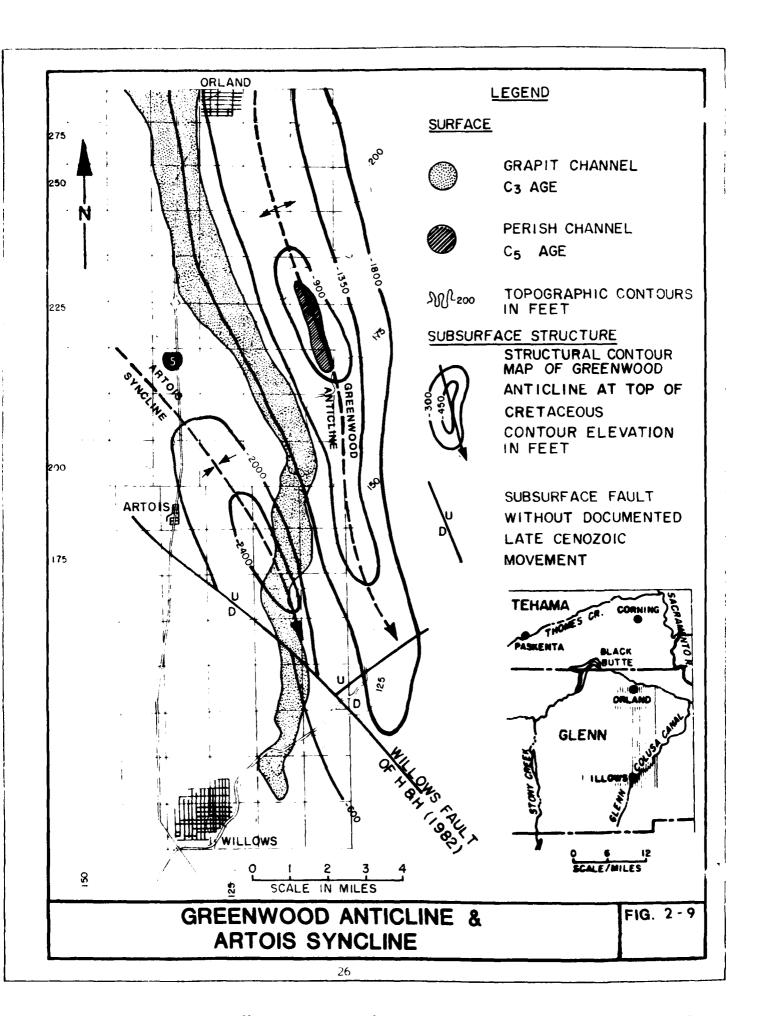


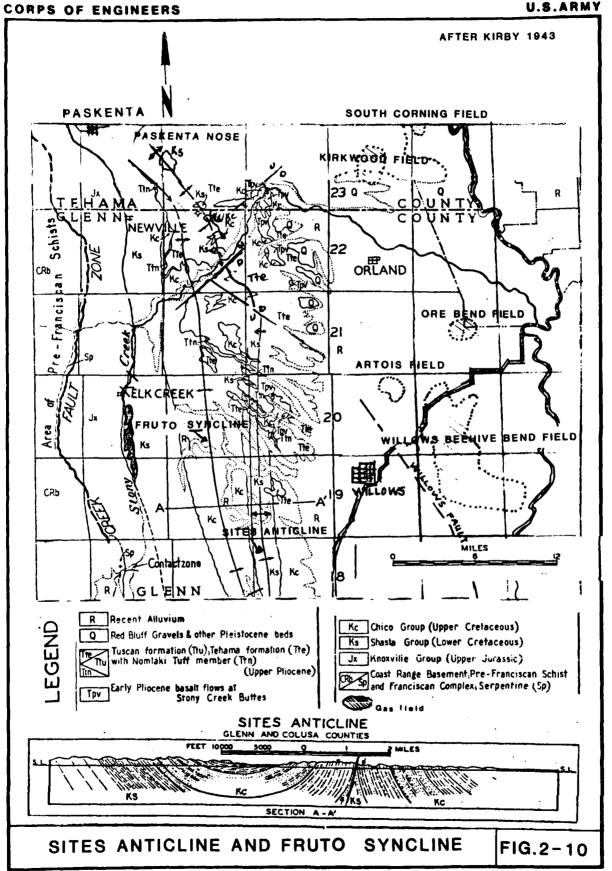
**CORNING DOMES** 

FIG. 2 - 8 faulting along the flanks of folds is equally northeast as well as northwest in trends. North-south trends are generally minor. State of California Bulletin 181 and other oil and gas journals also indicate similar faulting in published maps of the gas fields. Furthermore, faults indicated on time sections are not spacially alined to indicate through-going trends. At best we can speculate on an anastomosing subsurface pattern.

Greenwood and Willows Anticlines. The Greenwood anticline is a minor northtrending fold (Harwood and Helley, 1982). The fold area is shown in figure 2-9. We interpret the fold to be a northwest continuation of the adjacent Willows-Beehive Bend anticline lying to the south and east. The western flank of the Willows anticline is truncated by the Willows Fault. Alkire (1968) suggests this fault terminates at Artois and is one of many cutting the general Willows anticline, both along northwest and northeast trends. Harwood and Helley (1982) do not indicate these faults on their illustrations; however, they continue the Willows Fault beyond Artois and suggest a genetic relationship between the folds and the fault. Stratigraphic producers in the Willows, Artois, and Ore Bend fields are within the Kione or Forbes Formation. A lack of Tehama production is noted because of either the absence of the Tehama gas sand or no deformation in the Tehama. We believe that seismic lines indicate the Tehama to be undeformed throughout most of its section. As shown on figure 2-9, C3 and C5 channels lie across the suspected fault and fold trends in the Greenwood anticline (see also map 9). Both channels are undisturbed in topographic profile across the fold. Thus there has been no deformation in this section for the last 250,000 years.

Sites Anticline. The Great Valley sedimentary bedrock forms a broad synclinal structure with large folds superimposed on the west limb. The Sites anticline and Fruto syncline are two such folds (figure 2-10). The Sites anticline is a flexure named by Kirby (1943). The 35-mile-long fold is a tight, near vertical isoclinal fold structure about 3 miles east of the Fruto syncline. There is evidence to suggest that the Sites anticline is a decollement fold detached from a subregional basement structure by the Sites thrust fault. The section on geophysics (paragraph 2.4) develops the Sites thrust in more detail. The anticline is asymmetric with a steep axial plane, and the east limb is slightly overturned. The west limb rapidly flattens into the Fruto syncline. The axial plane of the fold dips steeply east and generally corresponds with a reverse fault, east side down. This fault appears at depth to be a listric surface that roots into a sole thrust at a depth of about 12,000 feet. This is easily seen in cross sections developed and presented in the geophysics section (2.4). The anticline strikes with excellent exposure up the west side of the valley foothills from Sites for over 10 miles to the vicinity of White Cabin Creek and French Camp (Fruto N.E. quadrangle) where the bedrock disappears beneath the Tehama Formation. Kirby (1943) surmises that either the anticline bent westward to blend into the Paskenta nose or continued north toward the reservoir area of Black Butte. Several seismic reflection lines crossing the north reservoir area and Walker Creek area do not reveal the continuation of the Sites Anticline northward along axial strike. Chuber (1961) mapped multiple southeast-trending, high angle faults that parallel the dominant northwest trend to drainage of creeks in the Fruto N.E. area. These faults sometimes





offset the axial plane of the fold. Limited exposure of beds at the head of White Cabin Creek are warped westward and suggest that the Paskenta nose is a continuation of the Sites anticline, possibly offset eastward by high angle, lateral slip faulting.

Fruto Syncline. The Fruto Syncline is a double plunging fold with a locally steeper west limb. The syncline axis is 3 miles west of the Sites Anticline axis. The fold is approximately 30 miles long. The syncline has several shallow, minor northwest-trending faults disrupting the east limb. High angle northeast and northwest faults are present in individual formation members on the west limb.

Chico Monocline. Along the eastern foothills a remarkable northwest-trending linear feature defines the valley edge. This feature is the Chico Monocline. The structure is a west-dipping monoclinal fold in the Pliocene Tuscan Formation and underlying the Great Valley Sequence. Harwood and Helley (1982) interpret the structure as indicating that beneath the monocline exists a steeply east dipping reverse fault in the basement. Magnetic data, interpreted by Griscom (1973) through modeling, indicate that the fault separates ophiolite from Sierran basement at depth. The existence of earthquake epicenters coinciding with the trace of ground expression suggests a near vertical or slightly west dipping fault (Marks and Lindh, 1978). It appears that the Chico Monocline has been active during the last 1 million years.

#### c. Faults.

Coast Range Thrust. The Coast Range Thrust has been described as a prominent thrust of considerable magnitude. In regional expression the Coast Range Thrust is a sinuous contact between the Great Valley Sequence and the ophiolitic and Franciscan rocks or the sporadic ultramafic sheet that lies between. Bailey and others (1970) considered the thrust to be separate allochthonous tectonic blocks carrying the Great Valley Sequence over the Franciscan assemblage. Maxwell (1974), Raney (1976), and Suppe (1979) have argued that the Coast Range Thrust separates paraautochthons, preferring to have the foreland of Franciscan Complex rocks dive or wedge under the Great Valley Sequence in subduction—like response to a compressional collision. In the present environment of north/south compression, past compressional structures which are presently nonsynchronous with contemporary stress have little bearing for neotectonic response. Harlan Miller Tait (1983) and ESA (1980) reports deal thoroughly with the Coast Range Thrust. Trenching by ESA (1980) indicated no activity along this fault since the Pliocene.

Stony Creek Fault. The Stony Creek Fault is treated in this study as the Great Valley/Ophiolite contact located in the western part of the study area. It extends from the Coastal Range Thrust in the vicinity of the Elder Creek fault zone on the north, southward beyond Glenn and Colusa counties. Multiple studies of this fault have been conducted. 1/ They have considered the

1/Irwin (1966); Kirby (1943); Jennings and Strand (1960); Strand (1962); Chuber (1961); Brown (1964); Fritz (1975); Raney (1976).

contact as being (1) normal depositional in nature, (2) depositional but faulted, or (3) wholly faulted, a large spectrum for any structure. ESA (1980) examined the Stony Creek Fault contact through its extent during a study for the Glenn Reservoir Complex. They concluded that the configuration of the Stony Creek Fault can be characterized as high angle, exhibiting both normal and reverse motion locally, with the west side (ophiolite) moving up to produce the present configuration of the mountain front. The fault represents shearing along a preexisting depositional contact and is composed of multiple segments. ESA and the State of California Department of Water Resources (DWR) investigated terrace deposits located along Stony Creek Fault, although the latter agency's study was more of a review of the former's work. Taken in totality, Quaternary terrace sequences indicate vertical movements on a segment of the Stony Creek Fault near Thomes Creek occurred between 30,000 and 130,000 years ago (ESA, 1980). Southern areas of the fault seem to have been inactive for 250,000 years. DWR does not support younger movement at Thomes Creek.

Battle Creek Fault. This fault has been the subject of analysis by Harlan Miller Tait (1983) for Corps of Engineers, and information is drawn from that study and from Helley, et al. (1980). The Battle Creek Fault zone is the name given by Harwood and others (1980) to the approximately N. 70° E. to easttrending fault zone on the northwest side of Battle Creek. The fault zone extends from east of the Sacramento River near Balls Ferry Bridge northeast toward Lassen Peak for a distance of 35 kilometers (km). It is projected west of the Sacramento River as a lineament. East of the Sacramento River, the Battle Creek Fault is roughly coincident with Battle Creek and forms a prominent escarpment. Vertical offsets progress from 200-foot offsets in broken Quaternary alluvial fans near the Sacramento River to 1,400-foot offsets in Cascade volcanic units. These volcanic units are approximately 0.5 million years old. The Battle Creek Fault is less well defined in the younger sediments of the northern Sacramento Valley west of the Sacramento River. The fault is questionably extended towards Sulphur Springs near the Coast Ranges/ Klamath province on the basis of weak expression by gravity and magnetics. Surface lineaments and some disturbance in Upper Cretaceous basement contours reinforce the fault extension. Harlan Miller Tait (1983) feels there has not been measurable faulting younger than morphostratigraphic unit Q6 southwest of the Sacramento River. This is in conflict with Harwood et al. (1983) who speculate movement more recent than 130,000 years. Seismic reflection data confirms the existence of the Battle Creek Fault zone in the subsurface at least at one point (see Harlan Miller Tait, 1984, page 12). The sense of movement on this line is down to the south as indicated on the records published in the above-cited report. Harlan Miller Tait (1984) believe a 0.5-million-year age for last fault movement is correct because the fault does not penetrate strata 0.5 million years old.

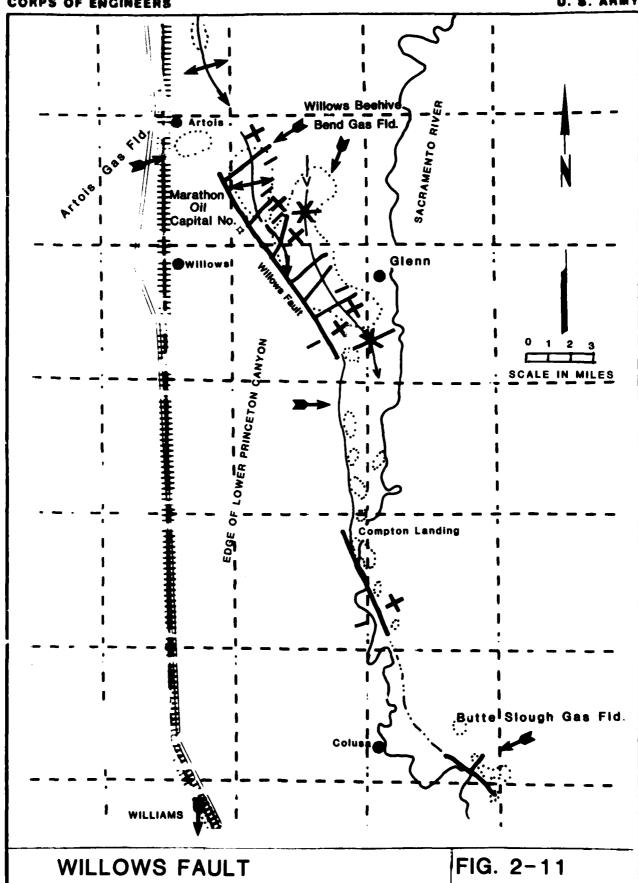
Red Bluff Fault. The Red Bluff Fault is the southernmost of a series of east-northeast alined faults in the northeastern corner of the Sacramento Valley. The Red Bluff Fault is a buried feature corresponding to the linear magnetic gradient that terminates the north end of the Great Valley magnetic high. Jennings (1977) and Steele (1979), each citing Madsen and Johnson (1960), show

movement on the Red Bluff Fault as south-down. Because of the similarity of its relationship to the Great Valley magnetic and gravity high with that shown by the Midland and Willows Faults, it is suggested by the above authors that the Red Bluff Fault may correspond to a basement fault along the east flank of the west side trough of the Great Valley syncline.

Jennings (1975) shows the Red Bluff Fault as a queried feature extending along trend across the west side of the Sacramento Valley to aline with a fault shown offsetting the south end of the Elder Creek Fault. Peppard and Associates (1981) show a fault of similar extent at the horizon of the Dobbins Shale within the Great Valley Sequence. No specific field study of the Red Bluff Fault was made in this study.

Willows Fault. (See figure 2-11) This fault is a north-to-northwest-trending, steeply dipping reverse fault with some slip components (Redwine, 1972, Alkire, 1968). The Willows structure has a suspected length of about 38 miles starting west of Artois and extending southward to a point about 14 miles east of Williams. Harwood and Helley, 1982, extend the Willows Fault south to the area of the Stockton Fault along the eastern side of the Sacramento Valley from the Colusa Dome at Colusa. Their extension is apparently based on a steepening of the crystalline basement west of the -1,500 meter depth contour. However, no clear offset of the subsurface, such as that demonstrated by Redwine, 1972, along the trend of the Willows Fault, north of Colusa, is documented along this trend of Harwood and Helley. Its length and orientation bisects the Willows-Beehive anticline, localizing gas productivity on the east or up-thrown block. Cross faults oriented northeastward also cut the anticline and are mostly truncated by the northwest portion of the Willows Fault. The cross faults cut the anticline into a series of down-to-the-north blocks. These and other faults, as well as the fold, were first defined in seismic reflection and refraction work (Alkire, 1968). Bruce (1958) mapped an unnamed fault west of Compton Landing gas field in Colusa County. In addition, an arcuate northwest-trending fault was mapped southwest of Butte Slough gas field in 1960. Redwine (1972) connected these faults based on the complementary trend of his contours outlining the Princeton submarine valley, the subject of his study. He believed that in this locality the course of the valley was determined by the Willows Fault. This work extended the Alkire (1968) trace of the Willows Fault to the southeast into sec. 1, T. 15N., R. 1 E. Redwine (1972), examining the Marathon Oil Capital-Capital No. 1 well (sec. 30, T.20N., R. 2W.) and other wells, concluded the Willows Fault to have an offset of 850 feet at the top of the Neroly Formation with the offset extending into the Tehama Formation.

From Redwine's structural contouring it appears that most of the movement along the Willows Fault has been dip-slip, northeast-side-up relative to the south-west. In the Artois-Willows area a C3 and C5 age channel deposit crosses the fault without any evidence of disturbance in the channel topographic profile. Even though historic seismicity in the Willows area might be attributed to the Willows Fault, there is no evidence for measurable ground surface deformation since about 120,000 years ago.



Corning Fault. The Corning Fault is envisioned to extend from Kirkwood on the south to nearby Red Bluff on the north (Harlan Miller Tait 1982). In USGS OF 82-737 of Harwood and Helley, speculation on the faults existence is based on unpublished analysis of several seismic lines crossing the fault area. Seismic reflection profiles used are interpreted to show some steep, eastdipping reverse faults with increasing offset with depth. Deformation in the Red Bluff gravel is also used to indicate faulting. In addition, observing the regional slope defined by the bottom of the Cretaceous along the east side of the valley (or top of basement), there appears a possibility of an offset of several hundred meters along the valley medial line which coincides with the Corning fault. Other authors have constructed sections in the area of the fault and have failed to identify fault structures such as the Corning trend described by Harwood and Helley (see Redwine, 1972; Safonov, 1962; Sacramento Petroleum Association, 1962; Calif. Div. of Oil and Gas Vol. 30 No. 2 1944). In this study (see section on Geophysical Study, paragraph 2.4.4), cross faction and topographic profiles indicate the existence of a dome but not the continuous fault shown by Harwood and Helley.

Malton Fault. Harwood and Helley (1982) project the Malton Fault as lying east of Orland Buttes along a north-striking, high angle fault trend. This is based on their interpretation of seismic reflection records and map patterns of Quaternary units. They give the fault a length of 20 miles. The sense of relative movement is similar to the Corning and Willows Faults; that is, the east side is up, and doming of subsurface structure on the east block is due to drag folding. There is, however, no evidence of near surface folding as indicated by a shallow seismic reflection investigation performed during this study and by extensive remnants of Q5 and Q6 terraces which cross fault locations in Thomes Creek and Stony Creek (at Lemon Home Colony). The results of the geophysical survey crossing the fault and ground topographic profiling are reported in the Geophysical Study section (2.4). Extensive analysis of wells containing the Lovejoy Basalt indicates that the flow is essentially unfaulted across the supposed trace of the Malton Fault (figure 2-12).

#### Black Butte Fault. (See figure 2-12)

(1) Previous Observations. The Black Butte Fault as shown by Jennings and Strand (1960) has a map length of 9 miles and trends N. 100 to 200 W. The fault is thought to exhibit west-side-down displacement relative to the east. Russell (1931) was first to show a fault in this position to explain the presence of topographic, high buttes of basalt capping Cretaceous bedrock. He correctly reasoned Orland Butte basalt (referred to as Stony Creek Butte) was the same as basalt present in valley wells beneath the Tehama Formation. Russell evoked the fault mechanism to explain the absence of Tehama materials on top of the Buttes by stripping the uprising block. Later Durrell (1959) and Creely (1954) suggested the origin of the basalt was a Miocene flow that spread into the valley as channel flows from the Honey Lake escarpment. Orland Buttes and Table Mountain were suggested as coeval. Relief on the basalt surface was caused mostly by erosional relief of the underlying sedimentary rocks, the basalt acting as a cast. The basalt on the valley edge was uplifted by continued folding of the valley syncline. However, in the western valley,

WELL LOCATIONS WITH BASALT IN STRATIGRAPHY

FAULTS INFERRED OR SUSPECTED NEAR BLACK BUTTE DAM

FIG. 2-12

Van den Berge (1968) saw tectonic movements caused by folding and uplift as mainly responsible for the relief in structural position of the basalt, with principal tectonic movement occurring at the end of Pliocene time.

The Black Butte Fault was added to the California State Geologic Map by Jennings and Strand (1960). Older versions were first published in Bulletin 181 of the State of California and 1936 vintage maps. Present day map configuration resulted from reports by the Corps of Engineers, Sacramento District. The Black Butte Fault was accepted by the Corps of Engineers (1963) during site investigations for Black Butte Dam and Reservoir based on published maps. In addition to the main Black Butte Fault, four subsidiary faults were mapped during design and construction of the dam. Two north-trending, high angle normal faults with steep easterly dips are present in the right abutment of Black Butte Dam. Both these faults have reported displacements of less than 50 feet (15m) with the east side down-dropped. A third fault trends N. 200 W. and was shown 3/4 mile (1.2 km) west of Black Butte Dam. The presence of this fault is based on an east side-down step in the basalt capped ridge north of Stony Creek. The mapped trace of this fault extends to the southeast about 1-1/2 miles (2-1/2 km) from this basalt capped ridge. The fourth fault has a trend of N. 550 E. and is mapped as displacing the Tehama Formation.

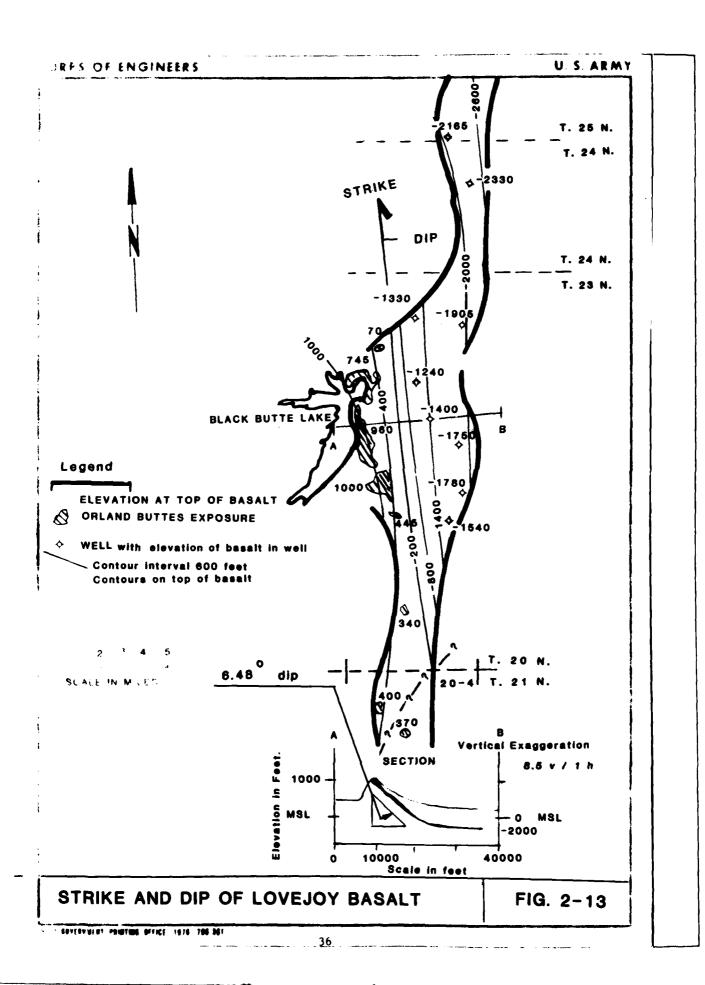
During subsequent work by the California DWR, two additional splay faults were mapped south of Orland Buttes near Walker Creek. These two faults trend N.  $60^{\circ}$  to  $70^{\circ}$  W. and extend to the southeast from the main Black Butte Fault. These faults were mapped based on photolineaments and brief field examination; they are shown as offsetting deposits of the Tehama Formation.

Redwine (1972) examined subsurface oil well data, and suggested that the Willows Fault, to the south, projects into the Black Butte Fault. This combined Black Butte-Willows Fault would have a total fault length of 55 to 60 miles (88 to 96 km).

- (2) Evidence Developed During this Investigation. Field mapping, lineament analysis, stratigraphic reconstruction, gravity and magnetic data analysis, shallow seismic reflection traverses, deep seismic reflection traverses, and trenching developed as part of this investigation leads to the conclusion that the Black Butte Fault does not exist; rather the Buttes are a topographic escarpment and that Orland Buttes was a hogback prior to Tehama deposition. The evidence for this conclusion is summarized as follows:
- o Field mapping of the Black Butte area reveals that Orland Buttes are a topographic expression of an erosion escarpment and that an unfaulted stratigraphic column lies across the inferred location of the fault. Field maps for the area of the inferred fault are Black Butte Dam, Fruto N.E., and Sehorn Creek sheets. Greater understanding of this discussion can be gained by examining figures 2-12 and 2-13. Orland Buttes are capped with up to 80 feet of basalt that strikes in outcrop an average N. 170 W., and dips 4 degrees eastnortheast. Underlying the basalt is Upper Cretaceous Great Valley Sequence that strikes N. 210 W. and dips 20 to 30 degrees east-northeast; hence, the

linear nature of the Buttes. The Tehama Formation was deposited to an elevation of 600 feet on the basalt eastern slope and on the Great Valley Sequence's western slope. The Tehama contact with underlying material is an erosional unconformity in nature; thus the formation draped around the ridge line as it was deposited. The pre-Tehama topography, a hogback consisting of basalt and Great Valley strata, is the original escarpment that was partially buried during Tehama deposition and presently undergoing exhumation.

- o The stratigraphic section of Cretaceous rock at Orland Buttes along the Tehama-Glenn County line was presented in prior paragraphs. As noted in plate 5, the section is not faulted and continuous to the west in normal succession. Gravel benches present at elevation 900 feet on the middle Butte are a pre-Pliocene channel of a west-southwest flowing stream which transported Sierran granite materials as evidenced by granite present in the cobbles. Redwine (1972) suggests, and we agree, that the Lovejoy Basalt marks the rimrock of the lower Princeton gorge. The presence of granite bearing stream gravels yields added evidence to this hypothesis and does not indicate uplifted Tehama formation on the western slope of the escarpment.
- o Lovejoy Basalt lies as a single unfaulted flow on the top and eastern side of the Buttes. Outcrop pattern on the Buttes and subcrop elevations in wells east of the Buttes suggest the basalt is a narrowly confined flow controlled by the existing Great Valley topography at time of deposition. Figure 2-12 shows the limits of the flow. The flow subsequent to deposition appears elevated through continued folding of the valley edge and is tilted eastward. The axis of tilt strikes nearly due north as does the valley axis. As seen in well logs, the flow tilts eastward at 111 feet per 1,000 feet, or approximately 6-1/2 degrees. This is in good agreement with outcrop information and three-point solutions using flow tops. Solutions indicate the average strike of N. 17 W. and 4-1/2 to 6-1/2 degrees eastward dip. Predicted by strike and dip, the strong repetitive presence of basalt at subsurface elevations in wells to the east of the Buttes indicates little if any faulting in the basalt. The geometry of the data presented in figure 2-13 indicates the possibility of only minor faults in the area of T. 21 N. and T. 20 N., R. 4 W. offsetting basalt with less than 200 feet total vectoral displacement.
- o Lineations in the Great Valley Sequence along western slope of the Buttes are not faults but linear side slope benches emerging through the cover of basalt colluvium and Tehama material. The projection of the inferred fault based on photolineament trend south of Orland Buttes, crosses South Fork Walker Creek, and strikes towards an unnamed drainage. The suspected main branch of the lineament crosses Q3 and Q4 (?) terraces on Hambright Creek and Walker Creek without apparent disruption or offset. A cut bank across this lineament on the north side of Walker Creek was cleaned and logged. No Tehama age material is faulted in this exposure (see log of trench on plate 6).
- o The stratigraphic section of Cretaceous rock at Black Butte, along the Tehama-Glenn County line, was presented earlier in this study. No evidence of faulting is revealed in the stratigraphy.



o Gravity and magnetic data were analyzed. Several microgravity traverses were conducted. The Geophysical Study section (2.4) presents the results of the Bouguer gravity profile analysis and Bott 2D analysis. Gravity observations do not support the presence of the Black Butte Fault. Magnetic data in the area indicate that the Lovejoy Basalt is a continuous flow dipping to the east. It is not detected west of Orland Buttes and pinches out to the east. No fault of any magnitude can be modeled using the available magnetic base and traverses made by this study (NGA, 1984).

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- o A shallow seismic reflection survey was run crossing the suspected Black Butte Fault at two locations. No trace of the fault exists in the 1/2-second records (see plates 11 and 12.)
- o Deep seismic reflection records were obtained on three crossings of the fault. No faulting of any significance can be found to correspond with the west side of Orland Buttes (see Geophysical Study section, 2.4).

Two north-trending, high angle faults with steep easterly dips (east down) are present in the right abutment. We analyze these to be block slides in the Butte due to erosion at the toe of the basalt (see Site Geologic Conditions section (2.3)).

The north-trending fault mapped in the basalt cap on North Butte is an apparent step in the flow geometry. Mapping an individual flow across the Butte shows this ridge is not faulted. Plane table mapping of Great Valley Sequence rocks exposed at the base of North Butte and at Eagle Pass did not reveal any offset of strata (communication obtained from Chico State College Geology Department, 1984; see Black Butte map sheet).

The N. 550 E. trending fault mapped on the north side of Black Butte does not appear in the stratigraphy at or above the level of Lovejoy Basalt. A three-point solution to Lovejoy Basalt elevations using North Butte elevation mark 782, Humble-Hall #1 well (sec. 21, T. 23 N., R. 4 W.), and the Grace Petroleum well indicates attitude in good agreement with the north-trending channel of basalt as a whole (strike solution N. 100 W., 4.28 degree dip east). However, this area is southeast of the implied fault and on the indicated up-thrown block. To the north the projection of basalt shows that an outcrop should occur at Butte Mountain-Long Hollow Road, map 5 (elevation 500 feet). Field checking in that area did not reveal any basalt; nor did shallow wells or seismic reflection survey indicate basalt in the subsurface.

We conclude that faulting cannot be confirmed on the basalt occurrence northeast of the supposed fault. Using wells in T. 24 N. and T. 25 N., three-point solutions indicate the basalt flow is unfaulted. The Tehama Formation does not appear faulted in field examination and is so mapped on the Black Butte sheet at Buckhorn area north of the dam. A northeast-trending fault exists in the subsurface Upper Cretaceous stratigraphy. Up to 200 feet of offset exists, with the northwest side up. The fault does not penetrate the upper 1,000 feet of strata. This fault is present only in the subsurface, and the only evidence for its existence is that the fault appears in the Long Hollow-Ham seismic section (see in Geophysical Study section, 2.4).

The Walker Creek Faults identified by ESA (1980) through their lineament analysis do not indicate faulting in the Lovejoy Basalt. A basalt knob lies between the faults (sec. 34, T. 22 N., R. 4 W.). Again, a three-point solution using the basalt contact on Orland Buttes (sec. 21, T. 22 N., R. 4 W.), Worthington Knob, and Franklin Well (S.E. 1/4 sec. 24, T. 22 N., R. 4 W.) basalt at -1,508 feet mean sea level gives a N. 50 W. strike and 4.05 degree dip. This is in excellent agreement with the regional strike and dip on basalt as an unfaulted flow. Field mapping indicates the Tehama is not faulted (Fruto N.E. sheet) in the area of Walker Creek. On individual creeks, Q4 terrace elevations match on each side of the banks that straddle the fault projections. In sum, the Walker Creek Faults have no field evidence to support their existence.

Paskenta Fault Zone. The Paskenta Fault was studied in detail by ESA (1980). That study concluded that the fault zone had no expression in four terrace levels along Thomes Creek (probably Q4, Q5 and Q6 equivalent strata terraces) and had ceased movement by the beginning of Tehama deposition, 3.4 million years ago. Prior to ESA, Jones, et al. (1968, 1969), show the fault zone on a generalized Buchia zone map as 1/2-mile wide, northwest-trending, and curving gently northward to merge with the Stony Creek Fault (Buchia Crassicollis is a prominent faunal marker in Lower Cretaceous rocks.) They conclude that large differences in Buchia zones indicate many miles of stratigraphic separation occurred on the fault and that the apparent 4-1/2 to 6 miles of left-lateral surface offset took place during pre-Tertiary time. The movement probably occurred when the beds were more horizontal. Mapping from Jones and Bailey on the Paskenta 15-minute quadrangle and included in the ESA report (1980) indicates the fault zone to be an anastomizing series of shears that branch toward the south and cannot be followed into the Tehama Formation that caps the fault southeast of Paskenta. Page (1966) sees the Paskenta zone as part of the Sierran/Franciscan boundary where the tear fault becomes a tectonic contact (Stony Creek Fault) between the Franciscan Complex and the Great Valley Sequence.

Kirby (1943) first termed an area of structural disturbance, traceable southward from Paskenta in scarce river and creek bottom outcrops, the Paskenta nose. This synchronous fold/fault zone is mapped northwest of the north arm of Stony Creek trending towards Paskenta (see Sehorn Creek and Newville sheets). The trend disappears beneath the Tehama Formation north of the north branch of Stony Creek along a trend that follows the northwest trend of Kendrick Creek towards Paskenta. The immediate suggestion is that it is related to the Paskenta Fault zone. This study believes the Paskenta nose as that area where the Paskenta tear fault passes into the listric fault associated with the Sites anticline. Geophysical seismic sections developed across the Paskenta nose indicate the nose originates at depth along the same detachment as the Sites thrust fault. (Sections indicating this are shown in the Geophysical Study section 2.4.)

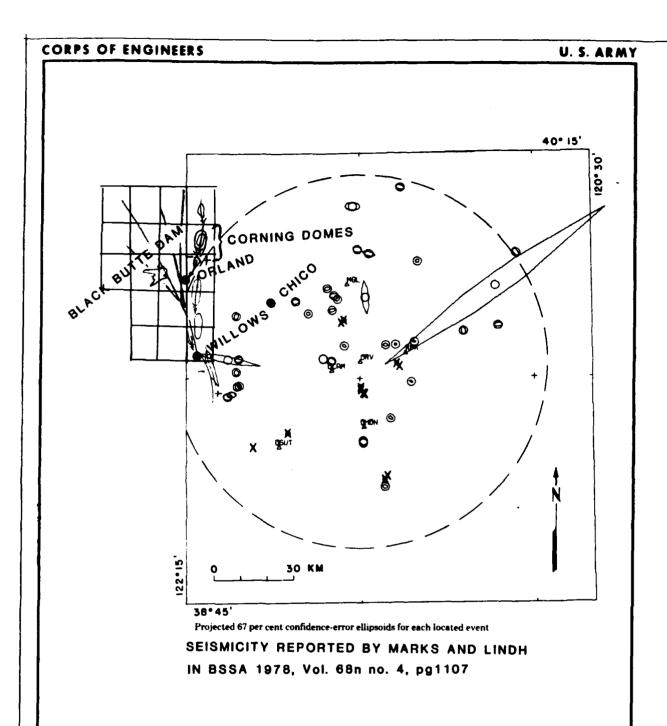
Willows Fault System. The Willows Fault system was defined by Harwood and Helley (1982) based on contouring of stratigraphic units obtained from analysis of well log data, some seismic reflection, and other assumptions. Their

concept of the Willows system is shown on plate 6. As mentioned, the Willows Fault system is reported in USGS Open File 82-737. A large part of our study was devoted to examining the field evidence for the existence of this fault system. We conclude that the Willows Fault system, as so defined, does not exist. Appendix C contains a detailed examination of the assumptions and evidence for the existence of the system presented by Harwood and Helley, including findings of this study that are contrary to statements of evidence given by Harwood and Helley in their text. The existence of the fault system described is controversial and suspect in any hazard analysis for Black Butte Dam. In sum, we do not feel there is any evidence supporting the items in the following list. Quite the contrary, this study presents cogency of evidence against the existence of the items noted below and, therefore, continuity of the fault system proposed:

- o The fault at Orland Buttes.
- o Association and trend of seismic events (shown by Marks and Lindh, 1978, see figure 2-14).
  - o Corning Domes and Greenwood anticline as evidence for drag faulting.
- o Citing the Muculloch Sunray Anchordogry No. 1 well and Occidental Petroleum Harris No. 1 well as evidence for existence of the Malton Fault.
  - o The interpretation of Harwood and Helley concerning the Seisdata lines.
- o The Willows Fault in the Willows gas field and the Paskenta and Elder Creek Faults at three respective zones appear to be well supported. However, the evidence for connecting faults between these zones is unwarranted and contrary to data developed during this investigation.

#### 2.2.4 Lineaments.

- a. General. Recent offset expressed in fault-scarp morphology can often be documented through analysis of low-sun-shadow enhancement of aerial photography (Slemmons, 1977). If the fault has been recently active, segments of known faults are sometimes reflected in disturbance of drainageway and other topographic features. Active or growing fold structure sometimes disrupts drainage patterns, such as producing centripetal drainage nets. Several types of imagery were used to analyze the area (indicated by shaded squares in figure 2-2) for lineaments. The imagery list includes the following:
- o Landsat 1:250,000 scale images, 30x30-inch enlargements, 28 Sept 1979 scene, bands 4, 5 and 6 and a false color composite of bands 4, 5 and 6.
- o Skylab: Three 1:250,000 scale color infrared (CIR) stereophoto enlargements, June 1973 scene of Sacramento Valley.
- o Stereoscopic, Panchromatic, quad-centered 9x9-inch photographs: 1:30,000 scale, 1983 flight, Ames RB47-7 flight lines.



RELATIONSHIP OF SEISMICITY REPORTED BY MARKS
AND LINDH TO THE CORNING STRUCTURES

FIG.

2-14

o Monoscopic Panchromatic low-sun-angle photographs: 1:12,000 scale; 10 flight lines.

desults of independent photographic analysis teams composed of Photographic Interpretation Corporation (PIC) personnel was reported in Harlan Miller Tait (1984). Photoanalysis and ground truth was carried out by Seattle District personnel, the results of which are shown on plate 1.

b. Interpretation of the Final Lineament Maps. The lineaments plotted on the final map are not necessarily faults in bedrock but may represent zones of bedrock structure (joints, fractures, bedding, or faulting) which are expressed as linear features upon the aerial photo or satellite imagery. Those lineaments shown as solid lines merited further investigation. The results of the investigation are given in table 2-2. Table 2-2 is cross referenced to plate 1.

It is difficult to compare lineaments mapped from satellite imagery to those mapped from relatively large scale aerial photographs. Two studies listed on plate 1 demonstrate this difficulty. The difference in scale necessitates that each analysis be based on differing criteria. At the satellite scale, entire river valley systems extending for many miles may aline themselves into one linear element, while at the 1:80,000 or 1:12,000 scale, only straight segments of a river channel within a valley system would be mapped as lineaments. A lineament drawn on a small-scale satellite image will span a width of hundreds of feet, and can be symbolized by a single, straight line on the satellite image. Each image, whether small-scale satellite or large-scale airphoto, provides differing levels of detail and care must be taken in comparing one scale image to another.

In the area of study, careful inspection of the drainage overlays generated as part of the PIC lineament analysis did not reveal obvious deflections in drainage patterns of the present day systems. Thus, no doming or upwarping is suggested. Harlan Miller Tait (1984) does not see any obvious evidence of upwarps associated with the surface projections of major mapped subsurface faults suspected in the Sacramento Valley (particularly those listed by Harwood and Helley, 1982). These faults include the Willows, Corning, Malton, Paskenta, and Black Butte. Drainage associated with the South Fork of Walker Creek makes a large boat-hook bend (see sheet 9, Fruto NE quadrangle map) to the south of Orland Buttes, probably because of the presence of the south basalt butte. Drainage between S. Walker and Wilson Creek north of Michaelson Ranch contains minor deviations from the dominant northwest trend and underlying silt-type dendritic character. In general, there is no obvious disturbance to the drainage net in this area, yet a peculiar pattern does exist and a strong lineament runs between the creeks alined towards the back side of Orland Buttes. This and several other lineaments of sufficient length and closeness to the dam warranted ground investigation. Further details on the above referenced lineaments are in the Geophysical Study section (2.4). Generally few lineaments found to date represent expressions of faults, and the ground investigation revealed that not one fault-lineament correspondence indicated faulting with Quaternary slip.

TARLE 2-2

IMAGERY AND PHOTOGEOLOGIC ANALYSIS

Comments	Photogeologic lineaments related to parallel drainage patterns developed in Tehama silts and areas of consequent fan slone. Do not offset sites artislied	Side slopes develop linear, slope percent abrupt change, banks cut in linear fashion but does not cross pock-mark ordinal.	Trellis pattern. Strong ground truth expression. Consequent drainage pattern in silt.	Consequent drainage pattern in silt. Consequent drainage pattern in silt. Consequent drainage pattern in silt.	Does not offset Nomlaki Tuff.	Does not offset Nomlaki Tuff.	Corresponds to nearby Cretaceous fault.	Resistant hogback ridges in strike-vale topography of Cretaceous outcrop belt.  Trends do not disturb Quaternary surfaces. Siltstone-conglomerate contact in Ku-hogback with subsequent drainage. Julian Rocks conglomerate hogback and Kl. Congolmerate in lower Lodoga Formation (Fm.). Contact between Stony Greek Fm. and Lodoga Fm. Congolmerate. Bedford Greek at Masterson Bridge sandstone-also congolmerate near Buris Greek (hogback). Jurassic-Cretaceous contact (Grassicolis zone). Rocky Ridge conglomerate. Scribner Ranch concretionary mudstone w/resist. Sandstone ridges. Conglomerate ridge unit and pebbly mudstone. Contact between mudstone and basic volcanic units of Coast Ranges.
Photo No.	5484017	10-1, 11A-1 11A-2, 12-1	10-1, 11A-1, 12-1	9-2 94-96 HAP 94-96 HAP	94-96 HAP	94-96 HAP		HAP 83 Strip 39 & 40
Youngest	600,000 ybp	450,000 ybp	450,000 ybp	Tehama 450,000 ybp Tehama	Tehama	Tehama	Cretaceous	10,000 ybp
Type	Drainage	Fabric	Drainage	Drainage Drainage Drainage	Drainage	Drainage	Drainage	Fabric
Location	Lower Wilson Crk. North- west of Willows	White Cabin Crk.	Sheep Coral	Stone Valley Salt Gulch North Hayes	South Hayes	Hoodo Creek	Nye Creek	Strike-Vale topography
Number	1 thru 8	-	2	m 4 v	9	7	80	1 thru 9 1 2 2 4 4 5 5 7 10
Group Letter	<	ra	æ	<b>ខា</b> ខា ខា	æ	Ø	a	w aaa a aaa a

TABLE 2-2 (continued)

Comments	Coast Range Thrust.	North-trending lineaments at front of foothill belt on western side of Great Valley.	No ground trace, possibly mislocated dg - d2.	Trenched at Boathook Bend Walker Creek - no expression.	Corresponds to west side Orland Buttes and strike of Cretaceous bedrock.	Linear drainage in Tehama/no expression in Q4 or 5-45 @ 240' RB not disturbed. (Red Bluff (RB)).	Tehama silt, R.B. cap not cut by linear.	Tehama silt, does not cut Corning age surface or Redding oravel.	See d 4 - continuation of trend.	Combination of drainage and roading system.	Boggs Chapline Road & Rreservation Boundary eastside	tanu use. West Reservation boundary & Table Mountain land use.	Linear drainage in calcrete silt, corresponds to	Old road system.	Does not offset basalt at Worthington Knob with Orland Butte strike and dip.	Does not disrupt or offset Q1-Q5 drainage channels	Sehorn Creek drainage not anomalous.	Linear drainage off Jewett Creek - joint and fault trends in Cretaceous.
Photo No.	HAP 83 Strip 402211	HAP 83-Strip 40213 and 54017	12-3, 2,1	11-A5	11A-1/B-1	12-10,9,8	11B-1	9-15, 14, 8-16,13		118-1, 9, 8	HAP 83 39218	HAP 83	HAP 83	118-4	401710-5	40179-11/ 8-11	392214 HAP 83	40179-21/ 8-22
Youngest Surface	125,000 ybp	10,000 ybp	10,000 ybp	10,000 ybp	10,000 ybp	13×10 <sup>6</sup> ybp	3x106 ybp	600,000 ybp		600,000 ybp	10,000 ybp	10,000 уър	10,000 ybp	3×106 ybp	10,000 ybp	10,000 уьр	Present	10,000 ybp
Type	Fabric	Tone	Tone	Fabric	Drainage	Drainage	Drainage	Drainage	Drainage	Drainage	12ne	Tone	Drainage	Tone	Drainage	Drainage	Drainage	Drainage
Location	West Marg. of Great Valley	Foothill Belt	West Michaelson Ranch	Wilson Creek	Orland Buttes	Buckhorn Crk.	Brannin Crk.	Sourgrass Rd.	Elmore Cem.	Squaw Hollow	Boundary Line	Table Mountain	Reservation	Crumbine Rd.	So. Fork Walker	Stony Creek	Sehorn Creek	Jewett Greek
Number	1	1 thru 15	1	2	3	4	5	9	7	80	6	10	11	12	1	2	3	4
Croup	ပ	د	ס	P	ਚ	P	P 4	3	P	P	ъ	Ţ	P	P	рp	pp	pp	рp

TABLE 2-2 (continued)

Comments	Straight segment in Hambright trellis pattern. No offset in Cretaceous beds at head of lineament. Straight pattern in Masterson Hollow drainage.	Cretaceous is not faulted when exposed at bedrock outcrop. Several ridgelines and creeks north of Elmore. Drainage pattern, Nomlaki Tuff in	subsurface. Fault in Cretaceous, not investigated.	Not investigated.	Reconnaissance did not identify linear on ground. North side limit of Red Bluff gravel pile. Air Reconnaissance failed to identify lineament. Not investigated.	Not investigated, related to Cold Creek fault. Not investigated, related to Cold Creek fault.	not investigated, related to Cold Creek fault. Projects from northeast faults in Cretaceous, Tehama drapes over bedrock fault.	Projects from northeast faults in Cretaceous, Tehama drapes over bedrock fault. Old terrace break (Strath)
Photo No.	401713-5 392213 HAP 83	399212 HAP 83	3992212 HAP 83		402212 Hap 88	402213 Hap 83	402212 Hap 83	402212 Hap 83 392219 Hap 83
Youngest	125,000 ybp 13x10 <sup>6</sup> ybp	Drainage 13x10 <sup>6</sup> ybp	Cretaceous		125,000 y.b.p.			
Type	Drainage Drainage	Drainage	Fabric		Fabric Fabric Fabric Fabric	Fabric Fabric	Fabric	rabric Drainage
Location	Hambright Crk. Masterson Hollow	N. Elmore Crk.	Ranch		Eagle Ford. Henleville Rd. Red Bluff Ranch		Table Mtn. one	McCarty Creek
Number	2 11	m	-		3 3 2 3	v v	ب	4 m
Group	គោ	TU.	60	(a.	യ <b>ഒ</b> രു ഒര	<b>60 60</b>	<b>ж</b>	: <u>.</u> e

- c. Lineament Summary. The photogeologic study and field examination concluded that lineaments in the Black Butte study area are related to drainage character and/or rock fabric. Strongly expressed lineaments in the study area are interpreted as reflecting the megascopic skeletal bedrock fabric, including the following:
  - o North 5 to 25 degree strike of Cretaceous bedding.
  - o Regional eastward dip of Cretaceous bedrock.
- o A regional hinge line running from sec 28, T. 23 N., R. 5 W., through sec 23, T. 23 N., R. 4 W., towards sec 3, T. 19 N., R. 4W.
- o High angle shear faults along N. 600 E, east-west, and N. 450 W. trends.
- o A subdued zone of secondary folding between strike valleys and hinge line.
- 2.3 Site Geologic Conditions. Black Butte Dam was built in a bedrock channel incised through and 420 feet below the upper surface of the sandstone and basalt ridge forming Orland Buttes. Black Butte forms the north abutment while Orland Buttes form the south abutment. Stony Creek passes over Great Valley Sequence rocks with a relatively thin cover of Tehama Formation to the Quaternary alluvium on the floor of the modern Sacramento River flood plain, which is underlain by hundreds of feet of Tehama Formation.
- 2.3.1 Stratigraphy. The stratigraphy at the damsite is somewhat less complex than that for the greater study area and is portrayed on plate 5. The oldest exposed rock is shale of the Cretaceous Great Valley Sequence (Forbes Formation) which dips 20 to 30 degrees to the east and the top of which is marked by an angular unconformity. The volcaniclastic Black Butte Formation overlies the Great Valley Sequence rocks, and in the area of the dam it thickens eastward. Field exploration along the west side of the Buttes indicated that the Black Butte Formation must be thin (less than 20 feet) if it exists at all (see Foundation Report, Black Butte Dam, 1963). Up to 90 feet of Black Butte Formation lies beneath the basalt at the dam abutments and up to 50 feet occurs beneath the main dam embankment. The Lovejoy Basalt unconformably overlies the Black Butte Formation. The basalt varies locally in thickness due to prebasalt topography, being approximately 70 feet thick at the damsite but elsewhere is often about 40 feet thick. The basalt dips 4 degrees to the east. The Tehama Formation laps around the buttes, lying upon and against the Lovejoy Basalt and Great Valley Sequence rocks. Up to 20 feet of recent alluvium underlies the bed of Stony Creek at the dam.
- 2.3.2 Dam Foundation Geology. The channel of Stony Creek has cut through the basalt caprock and into the underlying mudstone, sandstone, and conglomerate of the Black Butte Formation, and in one area it has reached the shale of the Forbes Formation. Plates 7, 8, and 9 depict the foundation geology of the dam and appurtenant structures. The impervious core is founded entirely on the

Black Butte siltstone and sandstone, except from stations 30+50 to 31+20 which is founded on conglomerate and stations 20+10 to 21+40 which is founded on Forbes Shale. The upstream and downstream shells of the dam embankment are founded on recent alluvium of Stony Creek, which is typically 10 to 15 feet thick overlying bedrock. On the left abutment, the main embankment was placed against deposits of the Black Butte Formation up to approximately elevation 415 feet and on Lovejoy Basalt above that elevation. The right abutment was placed against the same materials as the left abutment with the change in rock type at approximately elevation 395 feet. The spillway channel was cut through the Lovejoy Basalt and, in the lower part of the upstream end, siltstone and sandstone of the Black Butte Formation. The concrete structures of the outlet works upstream of the tunnel are founded on Lovejoy Basalt, mudstone and sandstone of the Black Butte Formation, and the Forbes shale. The intake control tower is founded entirely on shale of the Forbes Formation. The intake tower bridge abutment is founded on Lovejoy Basalt. Bridge pier No. 1 is founded on mudstone, and bridge pier No. 3 on conglomerate. The upstream end of the tunnel starts in shale of the Forbes Formation and conglomerate of the Black Butte Formation, and as a result of the downstream dip of the formations, penetrates the shale, conglomerate, mudstone, and lower portion the Lovejoy Basalt. A fault with 25 feet of east side-down displacement was encountered in the tunnel at station 8+02.5 and is discussed in the following section. Concrete structures of the outlet works downstream of the tunnel are founded on Lovejoy Basalt, except for a portion of the north stilling basin wing wall which is founded on Tehama Formation deposits of clay, silty sand, and gravel.

- 2.3.3. Dam Foundation Structure. Structural features of the foundation and attitudes, jointing, and faulting are treated below.
- o Attitudes. As noted above, the rocks of the Great Valley Sequence dip 20 to 30 degrees to the east, a consequence of being part of the limb of the valley syncline. Three-point problem resolution of the upper surfaces of the basalt capped Buttes, projection to subsurface occurrences to the east, and magnetic modeling show the Lovejoy Basalt dipping at 4+ degrees to the east. No bedding was evident in the Black Butte Formation, hence its attitude is unknown. The Tehama Formation generally lacks measurable bedding planes but regional occurrence suggests it dips at less than 2 degrees to the east.
- o Jointing. Two prominent and normal joint systems are present in the basalt at the damsite, and on the right abutment area they extend into the underlying sandstone. They strike N. 40° to 55° W. and N. 0° to 60° E. and both have vertical dips. A less prominent joint set dips gently eastward and was probably horizontal prior to the tilting of the flow. Overall, the joint system defines roughly rectangular blocks whose dimensions average 15 to 20 feet. Additional fracturing in the mudstone and sandstone of the Black Butte Formation is confined to the right abutment area, particularly in the upstream tunnel portal area and the tunnel from stations 10+30 to 11+00 where they are open and closely spaced, with heavily stained surfaces. Conglomerate of the Black Butte Formation does not exhibit jointing. Of the two joint systems in the Forbes Formation, one follows the bedding planes and the other trends slightly east to west of north and has steep dips.

o <u>Faults</u>. Two faults of minor displacement, revealed at the damsite during construction of the outlet tunnel and now inaccessible, were discussed in the foundation report as follows (C.O.E., 1963):

"(Two)...faults are present in the right abutment, one cutting the outlet tunnel at Sta. 8+00 and the other cutting through the tunnel conduit section at Sta. 13+00 Both trend north and dip east at high angles. They are normal faults in which the east blocks have been dropped in relation to the west blocks. Although they have displacements of less than 50 feet, their effects on the rock have been extremely detrimental and almost totally unexpected. ... Multiple gouge zones up to 40 feet in width were uncovered in the basalt of the tunnel conduit section, stilling basin and irrigation diversion channel. These zones cut the basalt mostly at high angles and consist of intensely altered basalt fragments in a matrix of brown, gray or tan clay. ...In the outlet tunnel from Sta. 7+50 to Sta. 9+35 an almost continuous zone of shearing was penetrated. Spacing of the individual shear planes ranged from 3 feet to less than one inch. ... Not infrequently the mudstone has been thoroughly pulverized into a loose, dry silty clay. The individual zones of extreme shearing ranged in width from a few inches to 5 feet or more. These zones, as well as the individual shear planes with sufficient continuity to determine attitude, almost always had strikes in the range of N10°E to N10°W and dips from 75° to vertical. Minor shearing was also noted in the stream channel section of the core trench and in the left abutment upstream from the core trench. ... However, faulting has probably been at least partly responsible for the development in the basalt of a system of near vertical joints in both abutments along which extreme weathering has taken place."

These faults, shown on plates 7 and 9, were considered in the foundation report to be branches of the "Black Butte Fault." However, this report shows that fault to be nonexistent. The basalt apparently flowed across a surface that was fairly level or tilted slightly to the west. Since its deposition, it has been tilted 4+ degrees to the east and possibly more. In addition, after emplacement of the basalt, down cutting of the upper Princeton Valley removed much of the basalt, creating a rimrock topography for the upper Princeton Valley fill, where detached blocks spalled off the valley walls. Such relationships are observed at the terminal end of the spillway where the basalt is in various stages of block detachment and moving downslope over terrain devoid of basalt. Also, underlying the basalt at the crest of the spillway cut where the basalt lies in a normal depositional relationship are "slick" clayey sandstone and mudstones. These factors, taken together, suggest that these two faults are related to gravity sliding and adjustment of blocks of basalt overlying an east-tilted "slick" clay formation.

2.4 Geophysical Study. Plate 6 is a map showing the nature and location of various geophysical studies. Figure 2-15 shows the results of gravity and magnetic modeling.

2.4.1 Gravity Study. Several gravity lines were surveyed along east-west trends perpendicular to regional structure. The lines were read to microgal accuracy and program corrections were made for tides and drift. Free-air, simple Bouguer, terrain, and regional gradient corrections were removed. A mass balance model using modified Bott 2-D routine was calculated (Bott, 1960).

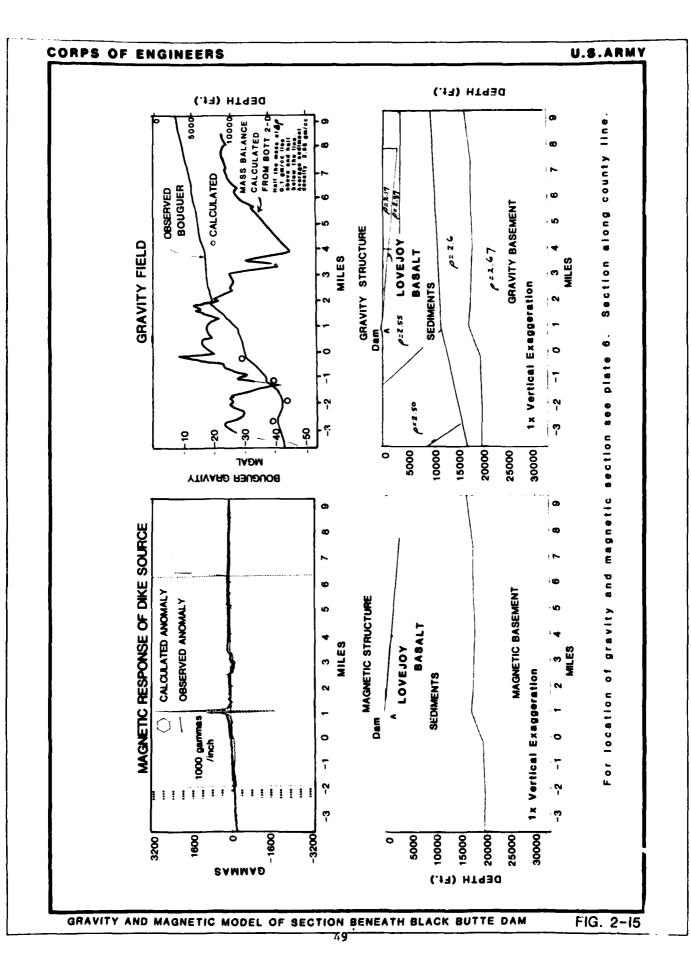
Based on the results indicated for  $\Delta \rho \frac{1}{2} = 0.1$  gm./cc., a two-dimensional section gravity model was attempted. Limited points were calculated and compared to observe Bouguer gravity. Results of this modeling indicate the lack of steep north-trending faults in the study area. Neither the suspected Black Butte nor Malton Fault could be confirmed. Gravity analysis indicates that the principal structure present is at subbasement level in the dam area, about 20,000 feet deep, and is parallel to the regional geologic trend expressed on the surface by visible structure. The gravity model suggests a wedge of less dense material is emplaced at depth near the mapped position of the Paskenta nose. Also, a block of higher density (possibly Franciscan or mafic volcanic) material is piled at depth beneath the dam in section 17 at a depth below 17,500 feet. No high angle faulting is associated with these boundary conditions, but rather the profiles suggest thrust slivers and doming.

2.4.2 Magnetic Survey. Several ground magnetic lines were surveyed along east-west trends perpendicular to the regional topography. The purpose of this work was to determine the most likely geologic structure of the Lovejoy Basalt flows in the area. The geologic structure was obtained by modeling surface magnetic data, using comparisons between observed and calculated magnetic anomalies to refine geologic sections developed from surface and borehole geology. The modeling results were further constrained by geomagnetic and paleomagnetic data. A complete report of modeling is contained in NGA, 1984. The favored computed magnetic model presents the Lovejoy Basalt as a continuous flow dipping to the east at 4.25 degrees. The basalt was not detected west of Orland Buttes nor is there any noticeable displacement in the basalt other than east dip. The only indication of structure other than unfaulted basalt flow is the character of the magnetic basement (as shown in figure 2-15).

There is a noticeable increase in the slope of the basement slightly west of Orland Buttes with the west side down. The inclination of the slope and the offset at depth indicated is dependent upon the magnetic parameters that are used for the basement. While these parameters may vary, the presence of some basement structure is required to model the regional magnetic anomaly. A minor anomaly near the crest of Middle Butte on the east flank corresponds to either the onset of superposition of two flows or a minor fault. Two flows can be mapped on the Butte at some locations. This anomaly does not appear in the South Butte.

2.4.3 Paleomagnetic Analysis and Age. Six sites on Black Butte and one Upper Chico Canyon site were sampled for age and paleomagnetic study. The location

 $1/\Delta \rho$  (Change in density or contrast between density of rock.)



and results of this study are reported in detail in NGA (1984). The Chico Canyon basalt is correlative with the Table Mountain basalt on the east side of the Sacramento Valley. The results of K-Ar radiometric age dates from opposite sides of the valley indicated mid-Miocene age between 13.6 and 15.2 million years ago for the basalts. The Lovejoy Basalt was formed during a reversed magnetic epoch. Some slight variation in declination and inclination exists, possibly due to small relative motion of basalt along Orland Buttes during uplift. However no large scale block rotation at the sample sites is indicated. The paleomagnetic property of the basalt suggests the direction of the flow is from the north in the immediate area.1/

2.4.4 Shallow Reflection Survey. Four shallow seismic reflection survey lines were made across suspected surface locations of fault structures identified by other investigators. The shallow seismic reflection survey, using common offset reflection profiles, was conducted under Corps of Engineers direction by Endacott and Associates. Complete results of the reflection survey are contained in their report (Endacott and Associates, 1984). The interpreted profiles are shown on plates 10 through 13. Results are as follows:

Line 1 (Plate 10). This line was shot across the Paskenta nose and morphostratigraphic unit Q4. This line shows two geological features. The anticlinal fold at 160-foot depth is Cretaceous bedrock configured as the Paskenta nose. The high-angle fault associated with the Sites-Paskenta feature is confined to the Cretaceous strata and does not propagate through the Tehama Formation or the Q4 terrace overlying Tehama.

Line 2 (Plate 11). This line was shot across the inferred Black Butte Fauit at Hambright Creek. The predominant geologic feature seen on this profile is a wavy rise on the time section on reflector R3. This corresponds to the location of buried Cretaceous topography. The east curvature of the reflector is a dip-slope of Cretaceous beds and the west curvature is an eroded slope exposing bed section as a reflector. No fault is present at station 15+00, the inferred position of the fault. Reflectors R4 and R2 are different materials as are R5 and R3. This interpretation is not unique and could also be interpreted to indicate a buried dike or fault. The dike is eliminated through magnetic modeling and no fault is indicated in deeper seismic records. No structure propagates into upper terrace cover (Q3).

Line 3 (Plate 12). This line was shot across the projected Black Butte lineament at Walker Creek. Geologic interpretation of this line indicates very shallow Cretaceous bedrock at the east end where it appears to be a buried topography mimicking monoclinal beds. The presence of the R5 (?) reflector

1/Magnetic grain indicates a general northern channel source. Ultimate source believed to be Honey Lake escarpment.

(diffraction) suggests a fault. A deep seismic record of the same area indicates no faulting of underlying beds. Deep reflection work tends to discount a shallow fault interpretation. A trench dug in the overlying Tehama material shows no evidence of disruption (a bank log of the trench is included on plate 19).

Line 4 (Plate 13). Line four was shot across the inferred position of the Malton Fault. Interpretation of this line indicates a fairly uniform and flat lying geological section of Tertiary sediments. No faulting of the Q5 terrace can be seen. Underlying Tertiary reflectors are continuous without fault interruption. A remnant of Lovejoy Basalt (indicated in well data) is possibly in the subsurface at 345 feet below the surface; however, if this is so, the high speed basalt causes velocity pull ups and makes depth uncertain. Diffraction near the k3 reflector can be the result of a pull up.

2.4.5 <u>Deep Seismic Reflection Work.</u> Four seismic reflection profiles were obtained and examined in detail and were correlated with wells, ground geology, and gravity lines. The results of this analysis are shown on four geologic-seismic sections illustrated on plates 14 through 17. The sections are of the general areas listed below.

Long Hollow Seismic Section (Plate 14). This section is along an east-west line from R. 5 W. to R. 2 W. in the general area of the northern border of T. 23 N. The line crosses the inferred position of Black Butte, Malton, and Corning Faults. The section shows the beginning of the Kirkwood-South Corning dome.

Basement is about 1.4 seconds (about 12,000 feet) in this area. Lying above basement, a series of detachment planes initiate within the Chico Series. The detachments rise in section and cut across east-dipping strata. Within the Chico Series, an orderly, persistent reflector group, characterizing Dobbins-Guinda-Sites Formation, are undisturbed beneath the trace of the Black Butte and Malton Faults. Disrupting this group is a high angle, west dipping fault labeled the Burris Creek Fault with approximately an 800-foot throw. This fault is repeated on strike in section shown on plate 15. In the eastern portion of the record, block and shallow trust faults ramp off the basement detachment below the Corning Dome.

ham Bridge to Black Butte Seismic Section (Plate 15). This section is along an east-northeast oriented line paralleling the general area of the north shore of Black Butte Lake and the North Fork of Stony Creek. This seismic section crosses the Paskenta nose and the inferred position of the Black Butte Fault.

Basement (not shown) is deeper than 3 seconds (greater than 15,000 feet) in this area. In the section plane, the basement dips sharply off to the west. The dip is estimated to be greater than 20 degrees. Approximately 3,000 feet

of Shasta Series lies below the Horsetown group of reflectors. 1/ The seismic-geologic data interpreted in this section shows the Paskenta nose rooted in a detachment plane lying above basement. The detachment rises in section with the thrust refolding the recumbent upper plate and dragging the lower plate. Originally, the upper plate fold was probably a nontectonic collapse structure. The nose subsequently became part of the synclinal folding of the valley edge, thrusted in continued crustal shortening and sheared along the Paskenta wrench fault trend. Strong reflectors within the Julian Rocks Formation correlate with higher acoustic impedance material found in wells and outcrop. The contact between the Venado Formation and Horsetown group reflectors is shown. It occurs in records as a dominant, coherent reflector. Very clearly multiple reflectors in section and principal reflectors noted are not offset or disturbed under the inferred position of the Black Butte Fault. The Burris Creek Fault is present in section.

Stubin Bridge to Walker Creek Seismic Section (Plate 16). This section lies along an east-west oriented line running along the southern sections of T. 22 N. from R. 6 W. to R. 2 W. The line crosses the Sites anticline/Paskenta nose, and the inferred positions of the Black Butte and the Malton Faults.

Basement (not shown) is deeper than 3 seconds in this area (greater than 15,000 feet). The contact between the Upper Cretaceous and the Lower Cretaceous Shasta Series lies in the lower portions of the record and is distinguished in well and reflector data by the distinctive conglomerate unit of Julian Rocks. Again, the geologic-seismic data interpreted in this section shows the Paskenta nose rooted in a detachment plane lying above basement. The detachment illustrated as a thrust fault is folded, as is the recumbent fold in the Aspilche shale (member of the Lodoga Formation). The syncline in the west part of the record is the beginning of the Fruto Syncline. An unnamed fault confined to bedrock exists east of Julian Rocks. Strong, coherent reflectors in the Upper Cretaceous strata are not disrupted at the inferred position of the Black Butte or Malton Faults.

Hayes-Burnell Seismic Section (Plate 17). This east-west line shows the seismic section from Fruto to the Glenn-Colusa canal north of Willows. The section crosses the Fruto Syncline, Sites Anticline, and the inferred position of the Willows and Malton Fault at bifurcation. The section continues on figure 2-16 and shows the onset of the Greenwood Anticline.

Basement varies from 8,500-foot depth in the eastern portion of the record to about 15,000 feet at midsection (basement not shown). Basement appears to dip off to the west. This line is of interest because it crossed the mapped position of the Sites Anticline. The section is considerably south of the dam near the limits of this study area. The Paskenta nose has apparently turned into and merged with the Sites tectonic structure as shown in this seismic-

1/Horsetown group includes Clark Valley mudstone and Julian Rocks congomerate.

geologic section. The Sites tectonic structure is interpreted as a slightly recumbent, anticlinal fold cut by a thrust fault rooted at depth. The thrust has a lystric surface originating from a basement detachment. Upper Cretaceous Chico Series have no faults disturbing reflectors corresponding with the inferred map position at depth of the Willows, Malton, or Corning Faults.

Sections were oriented approximately east-west and nearly perpendicular to the structural grain. Several additional northeast oriented seismic reflection profiles were also examined. Profiles were generally 4 to 15 miles long and had good resolution to acoustic basement, generally at 3.0 seconds. They were shot by oil exploration or geophysical service companies. Four major tectonic themes are indicated by the profiles and sections.

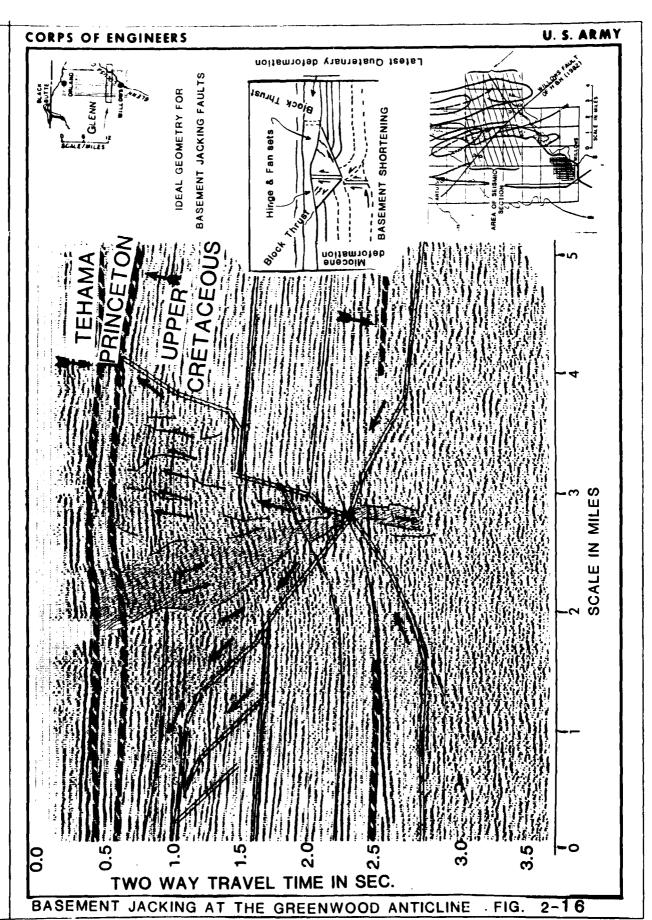
### 2.4.6 Analysis of Geophysical Data.

- a. The western edge of the Great Valley Sequence near Black Butte Dam displays several folded overthrusts which are rooted in the Upper Cretaceous bedrock just above the acoustical basement (plates 14, 15, 16, and 17). The largest overthrust in the study area appears to be associated with the Sites anticline as an attendant listric fault (plate 17). This relationship apparently persists south of the Black Butte study area also (Richard Willingham, Ogle Oil Company and Carl Wentworth, USGS; personal communication January 1985). In the study area, the overthrust block has traveled over the Lower Cretaceous strata in excess of 1 mile. The primary zone of detachment is not observed on the records. The Sites listric thrust does not continue north to the reservoir but merges with the Paskenta structure. The Sites thrust is absent from the Long Hollow section (plates 14 and 15).
- b. Northeast and northwest-trending, high angle, normal and reverse faults cut the folded Cretaceous bedrock near the valley edge but do not cut the Lovejoy Basalt. The overlying Tehama Formation copies bedrock structure as a draping pattern where the Tehama has a thin cover over the Cretaceous bedrock. Northeast-trending faults are visible on the Ham Bridge and Hayes-Burnell Geophysical Sections (plates 15 and 17).
- c. A line of subsurface doubly plunging anticlines appears to mark the western half of the north-central valley. The domes are gas traps, have maximum closure in the Upper Cretaceous strata, and sometimes, but not always, work upward into the base of the Tehama Formation (figure 2-10). The positioning of the domes has been discussed in the general literature and speculated to be linked to the deep seated structure identified by Oliver and Griscom (1980) as the Great Valley Gravity High. Oliver and Griscom explain the gravity high as possibly a tectonically emplaced fragment of oceanic crust. However, they do not discuss the domes. Others have expanded the concept using the proximity of domes near the gravity high to yield evidence for a high angle reverse fault that steps out of the tectonic basement. Using this fault mechanism, the folds result from drag on the reverse slip of the east block. Two other possible causes for the domes are explained below.

- (1) Primary Folding with Secondary Reverse Faulting Model. This model of deformation is developed as the primary response of thin-skin layered competent rock to general compression. As shown on figure 2-16, the folding may have been facilitated by localized or shallow detachment within the suprabasement. In this model, dipping reverse faults that flank the fold were formed during the late stages of fold development and are secondary to the fold structure. Sections shown in the domes on the Long hollow and Hayes Burnell lines indicate this mechanism to be the most probable cause of the domes and faults (plate 14 and figure 2-16).
- (2) Primary Reverse Faulting and Secondary Folding Model. The spatial distribution and geometry of the same folds can also be used to infer the presence of major reverse faults that are the controlling structures in fold development. Bentley (1977) proposed that the Yakima folds in the State of Washington are faulted monoclines that are localized above zones of highangle faults that extend from the basement. He terms this phenomenon "basement jacking." Laubscher (1977) suggests that such folds are narrow structures developed on broader warps. From this, he infers the presence of multiple decollements in the deeper basement and suggests that the broad warps were localized by thrust ramps rising from a deep (approximately 20 km) regional decollement and that the narrow folds formed above a shallow (1- to 3-km deep) decollement. A geometric analysis of Greenwood and Corning anticlines, based on a fault ramp-flexure model (Bruhn, 1979), suggests that the presence of fault ramps below the study area folds would extend to decollements at preferred depths of 10 to 12 km. The presence of crest, fan, and hinge set faults indicate the possibility of basement jacking as the underlying cause of these folds. However, deep decollement planes cannot be seen on the records.

No major concealed north-trending, high angle reverse faults can be interpreted from our analysis of the deep seismic records. In the records examined, the pattern of doming does not suggest drag folding. As shown on figure 2-16, an adequate explanation is found in horizontal shortening already visible in the area. Instead of an overthrust block rooted in the lower infrastructure, the domes are a result of detachment jacking of the shallow basement. An explanation of folding due to transpressive fold-thrust welt development need not be considered due to lack of a major northeast wrench zone (see transpression in glossary).

d. The Paskenta Fault trend is interpreted through profiles and section to continue into the Black Butte area and merge with the Sites anticline. In Cretaceous strata, fold deformation becomes more dramatic in the area of Kirby's Paskenta nose. The causitive mechanism suggested for this is a tear in the overthrust block connecting the Sites decollement with the Paskenta translational fault zone. Profiles on plates 15 and 16 show the squeezed nature of the Paskenta nose. This is a three-dimensional fold with left-lateral slip accounting for horizontal drag to folding as demonstrated on map on plate 15.



- e. The suspected Black Butte Fault cannot be found in the two reflection sections that cross its suspected position at Black Butte Dam. Little evidence for the Malton or Corning Fault exists that supports them as major concealed features. No northwest continuation of the Willows Fault was found and no major tectonic structure was found to cut the Sites decollement. Topographic profiles shown on plate 18 lie across the suspected location of the Malton and Corning Faults. As indicated by broken elevation profile on this plate, the Corning Fault area has indications of possibly reversed or oversteepened drainage.
- 2.5 Summary of Structural Relationships. Analysis and study of Black Butte regional geology and structure indicates the upturned western edge of the northern Sacramento Valley resulted from past crustal shortening. Within the study area listric faulting is responsible for thickening the Upper Cretaceous (Chico) section lying east of Julian Rocks between Nye Canyon and the North Fork of Stony Creek. The high-angle thrust daylighting along the axis of the Sites Anticline is the expression of one such listric surface. Thickening of the Chico seems to be limited along the north by the Paskenta nose structure of Kirby. As such, the Paskenta nose marks a continuation of the upper plate tear identified as the Paskenta fault zone at the Stony Creek fault junction. We believe that this zone dies as it merges into the Sites anticlinal structure much as a bedding fault dies.

All Cretaceous strata are strongly folded; however, there is much less folding in Tertiary strata. The Lovejoy Basalt has only a slight eastern dip and the Tehama Formation is essentially flat-lying. Northwest and northeast-trending, high angle normal faulting began along the western edge of the valley sometime during the time of the Lower Princeton Valley. This faulting was short lived because it does not seem to invade Lovejoy Basalt or Tehama Formation strata. No evidence to support the existence of a north or northwest-branching fault system was found. Faulting is present at Willows along and throughout the Willows Anticline.

The existence of the Black Butte Fault cannot be supported. No surface or subsurface geologic relationships support the proposed fault. A Quaternary stream terrace system dissecting the surface of the Tehama Formation serves to define morphostratigraphic units that lie across all existing and suspected faults. No evidence that these units were disrupted or disturbed could be found. With the exception of the Corning Anticline that indicates some recent doming, the present land surface is not deformed by any suspected or known faults. The existence of the Malton and Corning Faults could not be confirmed by surface or subsurface exploration. Rather, in the case of the Corning Fault, this study concludes the subsurface and near surface disruption of the strata is probably better explained by anastomising fan sets; minor faults resulted from primary folding of the Greenwood and Corning domes with development of secondary reverse faulting. No through-going regional structure cuts across the Sites-Paskenta trend.

Post-Tehama time has been marked by relative geologic stability. An excellent set of Quaternary terraces, 0.5 million years old or younger, lie across all of the above principal tectonic expressions. No evidence for continuing fault activity can be found in the Black Butte area more recent than 0.5 million years ago.

#### SECTION 3. TECTONICS

3.1 Tectonic Setting and History. A summary of the regional setting through Mesozoic and Cenozoic evolution follows. Each of the five informal provinces that are discussed in the regional setting of the project area has a different tectonic history. The major structural features in three of the provinces, the Klamath Mountains, Great Valley, and Sierra Nevada, were largely developed by the end of Mesozoic time some 70 million years ago, and have changed relatively little since then. Many of the Cenozoic features of the Great Valley and Sierra Nevada provinces appear to represent relatively small scale adjustments to changing stress fields that have been localized along major Mesozoic structures and have propagated upward into the overlying late Cenozoic section in increasingly smaller increments. The Cascade province has been largely developed in late Cenozoic time through a combination of basinand range-type extensional tectonism and Cascade volcanism. The northern Coast Ranges province has developed through nearly continuous periods of activity since mid-Mesozoic time. This activity is associated with the evolution of the continental margin.

In terms of the contemporary plate tectonics setting, the northern Sacramento Valley lies considerably inland of, but opposite, the northermost part of the San Andreas transform boundary between the Pacific, Gorda, and North American crustal plates. This is the Mendocino triple junction. The region immediately inland from the sharp westward bend in the San Andreas Fault is an area of concentrated stress marked by frequent earthquakes. Despite the westward turn of the San Andreas Fault some of the regime of right-lateral shear that exists in the northern Coast Ranges continues along trend as the Lake Mountain-Mad River Fault zone. This active zone intersects the coastline in the vicinity of Trinidad Head, north of Eureka, and continues for some distance offshore. As suggested by Herd (1978), it continues toward the alined and similarly oriented Blanco fracture zone.

North of the Mendocino triple junction the offshore Gorda Plate underthrusts the North American Plate beneath the present continental margin. The underthrust plate dips shallowly eastward beneath the present continental shelf and the coastal region north of Cape Mendocino. A highly seismic zone that exists between about 15 and 30 km in depth, beneath the lower Eel River Valley, is thought to be localized within the Gorda Plate. The south edge of the plate is interpreted from gravity evidence as extending beneath the northern Coast Ranges along a southeasterly trend from the vicinity of Cape Mendocino, approximately toward Lake Pillsbury. Two zones of active tectonism exist within the provinces bordering the northern Great Valley. One of these is the zone of right-lateral shear within the northern Coast Ranges that lies adjacent to the San Andreas Fault. This zone represents a wide mobile belt of continuing deformation along the transform boundary between the North American and Pacific crustal plates. The Maacama Fault is one such member. The second zone is the region of generally east-west crustal extension corresponding to the Basin and Range province and overlapping the southern Cascade Range province.

The present structural form of the northern Sacramento Valley is that of an asymmetric synclinorium, with a deep keel of Lower Cretaceous and Jurassic rocks beneath its western margin. Under the central part of the valley an overlapping section of Upper Cretaceous strata extends eastward from the core of the trough of older rocks directly across a platform of crystalline basement, probably oceanic crust. The overlap seems to wedge out northward over the Klamath Mountains and eastward over Sierran rocks. The strike of bedding in the unwarped Great Valley Sequence rocks generally parallels the valley margin. Thus, the Upper Cretaceous rocks define a shallow south-plunging basin underlying the northern valley area. The configuration of the subsurface of the deep keel of Lower Cretaceous strata beneath the west side of the valley was ascertained by geophysical analysis. The east flank of the west side trough has geophysical expression as the steep, linear magnetic and gravity gradients that extend up the axis of the Sacramento Valley, forming the west flank of the Great Valley magnetic and gravity highs. The gradient reflects a basement dicontinuity resulting from juxtapositioning of crust of contrasting magnetic character and density. There are two contrasting interpretations.

One view is that the east flank of the west side trough is faulted against the basement providing the existence of a series of east-up faults extending upward through the Upper Cretaceous section in a zone that closely matches the center of the magnetic and gravity gradient (Harwood and helley, 1982).

Another view sees a continuous basement surface dipping southwestward about 40 degrees to a depth of 15 km under the Coast Ranges. An eastward thinning wedge of Franciscan rock is defined between the basement and the overlying Great Valley sequence. Farther east the Great Valley Sequence is placed directly on the same mafic basement. Thus the Franciscan was emplaced in wedgelike fashion by assemblage thrusting northeastward beneath the Great Valley (Wentworth, Walter, Bartow, and Zoback, 1984). This view has gained widespread interest since 1983 and the Coalinga earthquakes. Postulated tectonic mechanism for Coalinga indicates wedge thrusting is active in the southern Great Valley.

Finally, several centers of persistent seismic activity that occur at unusually great depths for central California (15 to 40 km) exist in locations that would be down dip along the inferred zone of basement underthrusting (subduction), thrusting (wedge theory), or faulting (reverse faulting theory).

The Great Valley magnetic and gravity high terminates northward rather abruptly at about the latitude of the town of Red Bluff, and the matching vest side magnetic and gravity low turns sharply east across the valley trend. Griscom (1973) interprets the alinement of the magnetic gradient that terminates the valley highs as representing a basement fault (possibly corresponding with the Red Bluff fault shown on proprietary seismic reflection data (HMT, 1982, 1983). Concepts such as the Sierran-Klamath block of Page (1966), connecting lower Paleozoic oceanic assemblages between the Klamath-Trinity central metamorphic belt and ultramafic sheet and the Sierran equivalents (Shoofly Formation) or the trench-rift-transform triple junction of Blake and Jones (1981) would

define the magnetic trend as well. (See discussions of Blake and Jones or Schweickert and Snyder in Ruby Volume, 1981.)

In the Great Valley, from the vicinity of Corning southward, the topography and stratigraphy from the Pliocene nonmarine to the oldest Jurassic marine sediments, show the valley as being in a condition of long-term structural depression. It has been only since the Tuscan-Tehama deposition (3.5 million years ago) that the valley has changed posture. Small compressional features still exist from late-Cenozoic time, such as the doubly-plunging, dome-like folds and the Chico Monocline. However, these features can be traced from active tectonics diminishing between 15.6 (Lovejoy) to 3.2 million years ago (Tehama). Valley-down tectonics (east-west extension?) is displayed on ENEalined faults on the northern margin, along the west valley margin at Stony Creek, and some small breaks along the Chico Monocline. Structural deformation continued in this mode until 2.3 million years ago. Since that time, only minor tectonics are documented by structures such as the Red Bluff arch or Corning Domes, which are the youngest structural features in the valley, indicating upwarping at 250,000 years ago or later (Harlan Miller Tate, 1983). Some indications of valley margin deformation at Stony Creek between 125,000 to about 30,000 years ago (ESA, 1980) exist also, and tectonic movement is indicated by the 1975 Oroville earthquake along the Cleveland Hill fault. This geology and seismology on the whole seems to indicate north-south compression, east-west extension.

# 3.2 Contemporary Tectonic Model.

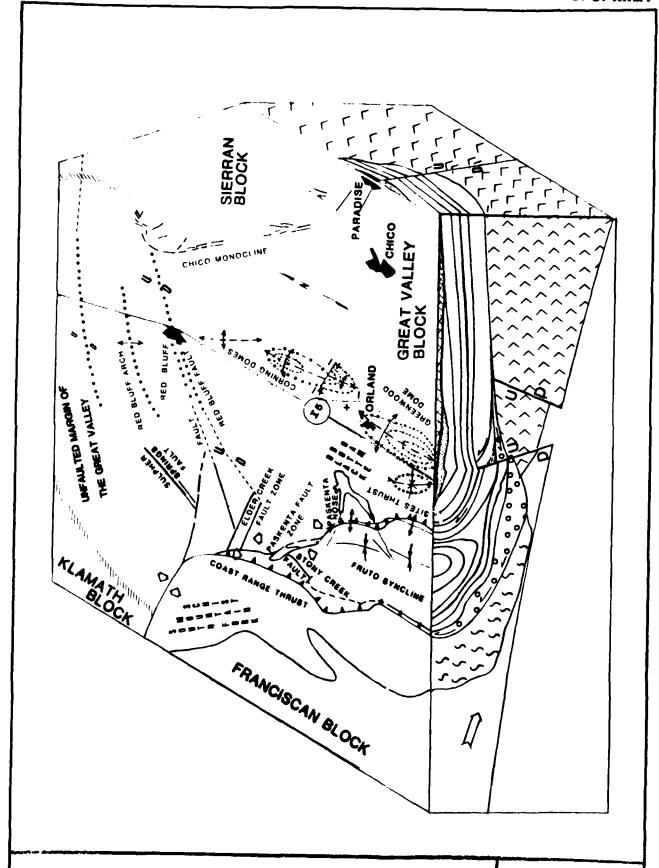
- 3.2.1 General. No consensus exists on the specific geotectonic forces affecting the Great Valley. Prior sections describe geologic evidence regarding state of stress and plate-tectonic setting. Figure 3-1 is a tectonic model which best fits the observations of this study in the Black Butte area.
- 3.2.2 Regional Stress Patterns. Earthquake focal mechanism and in-situ stress measurements provide a useful check on geologic observations. Well constrained fault-plane solutions and stress measurements are available in a publication by Zoback and Zoback (1980) for the western United States. Less well constrained data is sparsely available for the Great Valley and northern California in the fragmented references. Table 3-1 below summarizes contemporary stress.

Focal mechanisms of earthquakes associated with both the 1975 Oroville events east of the valley and 1978 Alder Springs events west of the valley show normal displacements along north-northwest oriented faults. Therefore, general stress in the valley is interpreted to be general north-south compression and east-west extension.

TABLE 3-1
SUMMARY OF CONTEMPORARY STRESS

Province	Maximum Stress Axis	Intermediate Stress Axis	Minimum	Data	Source
San Andreas	n/nne	Vertical	w/wnw	Strong Seismologic	Z&Z* (1980)
Coast Ranges	N/NNE	Vertical	w/wnw	Seismologic	Bolt (1967) ESA (1980)
Sierra Nevada	NNE	Vertical	WNW	Geologic	Z&Z* (1980)
Basin and Range	Vertical	NNE	WNW	Strong Geologic Seismologic	Z&Z* (1980)
Cascade Miller Tate (1982	Vertical	NNE	WNW	Geologic	Harlan
Great Valley	NNE	NNW	Vertical	Weak Seismologic	Lomnitz & Bolt (1967)
Margin	N7 5E	N15W	Vertical	Geologic	Bolt et al (1968)
Interior	NNE	NNW	Vertical	Geologic Seismologic	ESA (1980)

<sup>\*</sup>Zoback and Zoback



TECTONIC MODEL

FIG. 3-1

#### SECTION 4. SEISMOLOGY

4.1 Seismologic Setting. The northern Coast Ranges have a history of active faults and earthquakes. The Great Valley and Sierra Nevada is marked by relative quiescence. The distribution of epicenters is shown on plate 4. Macroseismicity (events greater than magnitude 3.5) is presented on figure 4-1. A lack of seismicity near the project is apparent; the nearest event is approximately 21 km distant. Plate 4 shows the distribution of all earthquakes listed in appendix A and the distribution of only those earthquakes which are believed to have occurred at depths greater than 12 km. A very low level of microseismicity is noted near the project.

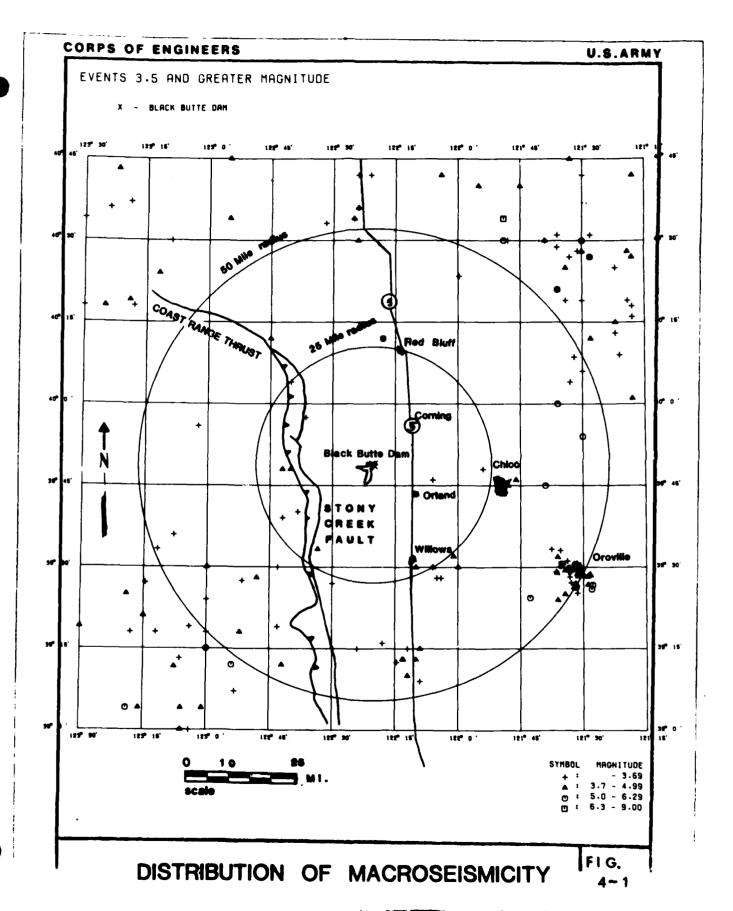
A catalog of earthquakes for the region bounded by north latitudes 39000' and 40045' and by west longitudes 121015' and 123030' is presented as appendix A. The catalog, compiled from three sources, has had obvious duplication removed and has been scrutinized for accuracy only for events greater than magnitude 3.5. The following sources were used:

- o Corps of Engineers Pacific Northwest Catalog. This is a multisource based catalog for the period 1853 through 1980, north of latitude 37000'.
- o California Department of Water Resources. This is a catalog of events recorded by that agency since 1975 from an extensive instrument array. This includes an eight-instrument array installed around Glenn County during the period June 1981 through February 1985. Also contained in this catalog are a limited number of relocated (University of California, Berkeley) events covering the period 1903 through 1973.
- o National Oceanic and Atmospheric Administration. A catalog of macroseismicity covering the period 1877 through early 1983.

Large events have occurred in the regional history. The largest events to affect the site of the dam are the 1903 Willows event at intensity (MM) VI and the 1906 San Andreas event at felt intensity (MM) V. The accuracy and locations of the reported earthquake effects vary according to the population density in the felt area and the historic period pre-1900, preinstrumental period (1901 to 1949), and instrumental data base. Structural features, as understood, presently mirror patterns of microseismicity and some isolated macroseismic trends. However, close examination of the uncertainty in seismic location and depth, net bias, and lack of tectonic understanding weaken correspondence between seismicity, structure, and location as constituting cause. The one exception is the San Andreas fault.

### 4.2 Microzonation.

4.2.1 Northern Coast Ranges. Seismic activity in the northeastern part of this province defines two northwest-trending alinements, which correspond generally to the Maacama and Mad River-Lake Mountain fault trends, and a north-northwest-trending alinement that includes the Alder Springs earthquake



sequence of 1978. A fourth, north-northeast trend, parallels and is nearly coincident with the Coast Range Thrust-Stony Creek Fault boundary.

Both the surface expression of faulting along the Maacama and the Mad River Lake Mountain trends and the few focal mechanisms that have been determined for earthquakes along them are suggestive of right-lateral, strike-slip faulting. Nearly all events along these trends are shallow, with focal depths typically being 6 km or less. The largest earthquakes recorded from this region are the 5.6- and 5.7-magnitude main shocks of the 1969 Santa Rosa earthquake sequence and, probably, the 1898 Mare Island earthquake, estimated to have been of about 6 magnitude (Toppozada and others, 1981). ESA (1982) has previously estimated a maximum credible earthquake (MCE) value of 6-3/4 for the Maacama fault zone. Instrumental earthquake lists show that the adjacent region along the eastern foothills of the Coast Ranges is one of low seismicity but earthquakes occur on a regular basis (Bolt, 1982).

The north-northwest trend of epicenters that is centered at about 39040'N., 122045'W., includes the Alder Springs sequence studied by ESA in 1980. This trend, although also involving shallow events, differs from the two described above in not being associated with any extensive fault trend thus far recognized. First motion solutions determined by S.W. Smith for the 1980 ESA (stated in Harlan Miller Tate, 1983) study indicated normal movement along northwest-alined faults for the source mechanism. The largest events recorded from this area are in the 3.5- to 4.0-magnitude range.

West of the region, the northern Coast Ranges are bordered on the west by the San Andreas fault, having recognized capability for generating an earthquake of 8.3 magnitude.

- 4.2.2 Klamath Mountains. The level of seismicity in the southern Klamath Mountains is generally low. Magnitudes seldom exceed 5.
- 4.2.3 Northern Sierra Nevada. The seismicity of the interior Sierra Nevada province, occurring mainly in proximity to the northerly part of the Melones fault and the Magalia-Cohossett Ridge Fault, might be due to instrumental and recent historic bias. A diffuse seismicity probably occurs across the Sierras. Magnitudes of 3.0 to 4.0 (M1) are scattered through the Sierran foothills. Larger magnitudes (6-1/2) occur to the north; however, they are far from the study area and only as close as the northern Sierran boundary.

## 4.2.4 Sacramento Valley.

a. Western Margin of the Northern Sacramento Valley. This section is the area near the Coast Range Thrust and the area approximately parallel to the Stony Creek Fault.

Two earthquakes of about 4.5 magnitude have occurred near the Coast Range Thrust-Stony Creek Fault during the past 80 years, and a seemingly well defined alinement of epicenters of small, shallow earthquakes occurs paralleling and lying a few kilometers west of these fault trends. Since the Coast Range

Thrust has a moderate eastward dip and is probably cut off at depth by the younger Stony Creek Fault, it is unlikely that these earthquakes originate on the thrust.

One sequence of shallow earthquakes, ranging up to nearly magnitude 5, was recorded by the USGS at a location somewhat east of the Stony Creek Fault and directly on the trace of the Elder Creek Fault. The Elder Creek Fault has not otherwise been identified as active. Alternatively, the sequence may have been centered on the Stony Creek Fault near its intersection with the zones of fracturing along the Elder Creek and Paskenta Faults. The Paskenta Fault zone appears free of seismicity.

b. Eastern Margin of the Northern Sacramento Valley. Four centers of seismic activity exist between the vicinities of Oroville and Red Bluff, along the eastern margin of the northern Sacramento Valley. The Oroville earthquake sequence began on August 1, 1975, with shocks up to 5.7 magnitude. Since then activity has diminished with decreasing frequency. It is the southernmost such center. Earthquakes of this sequence were accompanied by cracking and a few centimeters of differential movement at the surface along the trace of the Cleveland Hill Fault, a seemingly minor structural feature that had hardly been recognized prior to the earthquakes. Marks and Lindh (1978) showed that the foci of shocks of this sequence defined a plane striking north-northwest and dipping 60 degrees west. The shocks originated at depths ranging from the surface down to about 18 kilometers, and exhibited first motions indicating a west-down normal fault mechanism. Activity in this area now extends over a distance of about 25 km, centered approximately at Lake Oroville (Marks and Lindh, 1978; Lester, et al., 1975).

A second center of seismic activity is located near Chico, a few kilometers south of the southernmost surface expression of the Chico Monocline. Two earthquakes of magnitude 4+ were recorded in this area prior to 1981, and the plot for 1981-1982 shows 10 small shocks there. The latter events occurred at depths between 2 and 25 km and were mostly below 17 km. The distribution of hypocenters is suggestive of a moderately west-dipping fault as a source structure with a zone of activity extending about 20 km along a north-north-westerly strike (Bolt, 1968).

The third zone of seismic activity is defined by a scattering of small earth-quakes along the trend of the Chico Monocline. Although there is one shallow-focus shock in this series, most of the events occurred between depths of about 20 and 25 km. The proximity of the epicenters of these events to the surface trace of the Chico Monocline suggests that they originated along the underlying fault in the basement, and further, that the fault is essentially vertical and extends deep into the crust.

A fourth area of seismic activity exists northeast of Red Bluff, and is mostly between the northeast end of the Red Bluff Fault and the Battle Creek Fault. This seismicity is represented by seven small shocks in the 1981-1982 USGS data set and by a 4.3 magnitude event that occurred in 1968. Focal depths range from shallow to more than 20 km.

c. Interior of the Northern Sacramento Valley. Seismicity within the interior of the northern Sacramento Valley is represented by a few scattered small events and by two well-defined, persistent clusters of activity. The epicenter clusters are located near the settlements of Hamilton City and Glenn. The northerly cluster at Hamilton City area is several miles southeast of the south end of the Corning domes as it is defined by the Upper Cretaceous in the subsurface, and the southerly one near Glenn is east of the northern part of the Willows fault. Most of the events in both groups have focal depths between 17 and 35 km, although a few shallow focus events are included with the southerly cluster.

ESA studied the Willows earthquake of 1903 and earthquakes near Glenn during their 1980 study for Glenn Reservoir Complex. For a series of five earthquakes of magnitude 3.0 to 3.9, reported as occurring along a northwest alinement a few miles east of the town of Willows, and two magnitude 4.0 to 4.9 shocks about 10 miles east of Willows, upon resolution and relocation of the events ESA found the northwest trend near Glenn to be unfounded. Although the northwest alinement is not confirmed by ESA relocation of the events, new locations remain very close to the Willows Fault trace as known in the subsurface.

- 4.2.5 Depth Distribution. A few scattered earthquakes in the region, generally between the northwest margin of the Sacramento Valley and Cape Mendocino, have focal depths in the range of 40 to 50 km. These events may be associated with the southerly part of the underthrust Gorda plate. North of Cape Mendocino and beneath the Eel River basin, a well defined, gently eastdipping zone of seismic activity that exists between depths of about 18 and 30 km is interpreted as occurring along conjugate sets of steeply inclined fractures within this lithospheric plate (Smith, 1978). The scattered deep events located farther east and southeast of this region may result from isolated failures within the same plate at a greater depth beneath the overriding North American plate. The shallow events occurring at depth of 17 to 35 km in the interior valley are very often discussed in conjunction with down-dip expresgions of suspected faults (referred to in prior paragraphs). As a whole, this study concludes that the seismic reflection data in the interior valley does not support the existence of east-dipping, reverse high angle faults in the basement nor do they favor reverse faults ramping from deepest basement structures. Scattered seismicity of the interior valley remains unexplained.
- 4.3 Recurrence Intervals. The recurrence level of different magnitudes in a given time span and area is useful to engineering design. The design project life is about 100 years. Bolt (1982) feels, and we concur, that it is an academic exercise of little value to attempt an assessment of earthquake frequency distribution for the 100-year historic record in Glenn County. None-theless an estimate can be made based on the historic record for northern California. One or more San Andreas events should be experienced per 100 years as a source outside the project area. Similarly, based on observations in areas of known faulting away from the San Andreas, "earthquakes with  $M_{\rm L} = 5-1/2$  are perhaps the largest probable in 50- to 100-year periods for northern California." (Bolt, 1982)

If the Willows earthquake of 1903 is representative of such a magnitude and recurrence, then magnitude 5.5 is appropriate for the project life taken at any seismogenic source near the project. So the maximum probable earthquake for the Willows or Stony Creek Faults would be  $M_L = 5-1/2$  or for the Oroville margin about  $M_L = 5-3/4$  (small incremental addition due to increased activity in the foothills).

# SECTION 5. SELECTION OF THE SOURCE AREAS

5.1 Earthquake Potential. Prior to this study, the Sacramento District was provided a report on earthquake potential (Bolt, 1982). The earthquake sources of the region are essentially unchanged from those reported by Bolt (1982) except in minor detail. With exception of that detail, the following summarizes Bolt's report. The San Andreas 1906 event and the Willows 1903 event are the two chief historic earthquakes to affect the site. All faults near the Black Butte study area are pre-Quaternary in age. Various independently authored studies of various exposures along the Coast Range Thrust disclose no evidence of recent tectonic activity since Pliocene time. Bolt (1982) left open the question of activity on faults near the dam for a conservative approach. This study closes that question by showing that the Black Butte Fault and the continuity of the reported Willows Fault system does not exist.

The uncertainties regarding unmapped, north-trending faults branching from the Willows Fault south of the study area have now been sufficiently investigated to clarify and resolve the question. Three major studies, California Water Resources (1982), Harlan Miller Tate (1983, 1984b) and the present study conclude that there is insufficient evidence to support the concept of Harwood and Helley in Open File Report 82-737 proposing the Willows Fault system. The Willows Fault and faulted anticlines do exist in the interior valley; however, they are not connected. The question of age of deformation exists. Mid-Tehama deformation is demonstrable in seismic reflection data (2 million years ago) on the midvalley anticlines, yet the nature and extent of the 0.5-million-year-old deformed Red Bluff material on the Corning dome is questionable. Using Marks and Lindh (1980) or any other plot of earthquakes to associate micro or macroseismicity with these structures would be trifling if not misleading.

In light of the most probable structural connection to the Sites anticline, the Paskenta Fault zone and Paskenta nose are striking structural features. The Paskenta nose can be followed 10 miles from where it emerges from the Sites detachment to Tehama Formation cover. On trend, the Paskenta Fault zone can be followed 10 miles on the ground until it merges into the Stony Creek Fault. The Paskenta Fault is capped in many places by pre-Pliocene sediments, pre-Miocene pyroclastics and the Tehama Formation. Some clustering of microseismicity exists near Paskenta, but Bolt (1982) concludes the evidence is not sufficient to indicate capability. The Paskenta nose and the Sites anticline/fault do not disturb the Tehama Formation or the extensive Quaternary terrace system across it.

For the Stony Creek Fault, ESA (1980) concluded that the lack of Holocene displacement makes that fault an unlikely source of significant earthquakes. Nonetheless, the last major reactivation at 200,000 to 300,000 years ago and a minor episode 30,000 years ago is documented with minor offset in terrace strata. Although unlikely, it is considered a close, capable structure.

Finally, as Bolt states, "two faults to the west of the Coast Range Thrust must be considered since there is no question that either could be the focus of major earthquakes." The first is the Maacama fault located about 95 km to the west of Black Butte Dam. This fault has been identified recently by Bolt as a major strike-slip feature with a length of approximately 120 km. Field study of displacements along the Maacama Fault has established its activity in Holocene time and studies of regional strain and seismicity suggest that a significant earthquake might occur on this fault in the lifetime of a modern engineering project nearby. Further to the west, the San Andreas fault is encountered at a distance of 140 km from Black Butte Dam. Bolt feels it is necessary therefore to consider the possibility of a repetition of the 1906 earthquake on the San Andreas fault so far as shaking at Black Butte Dam is concerned.

Bolt's report, circa 1982, if written in 1985, possibly would be different. One new development that might be addressed is the impact of the Coalinga 1983 and Kettleman Hills 1985 events on how geotectonics are viewed in the Great Valley. The presence of a magnitude 6.7 (Coalinga) event some 35 km away from the San Andreas has caused some to rethink GV tectonics. Appendix D contains in greater detail the position this study has adopted regarding the impact of the Coalinga and Kettleman Hills events on valley edge tectonics. The following summarizes appendix D. Although tectonic wedges might reasonably exist in some fashion along northern California valley edge areas, as they seem to do in southern California near Coalinga, they are not presently proved to exist nor is there reason to believe that crustal strain existing at Coalinga and Kettleman Hills that complement the mechanism for the earthquakes are as follows:

- a. Fold belts in proximity to the San Andreas fault.
- b. Folds that have orientation sympathetic to shear transfer from the SA.
- c. Folds that show en echelon segmentation and have regional extent.
- d. Deformation related to folds on young geomorphic surfaces.
- e. Subsurface northeastward-directed thrust that invades eastward-directed wedges and end in roof structure such as folds.
- f. Reverse faults emanating from basement that are associated with growing folds and localize position of overthrust plate.
  - g. Seismogenic activity in the fold area.

The table below compares this list of what is found at Coalinga with what is present at Black Butte (see appendix D for detail).

TABLE 5-1
FOLD BELT COMPARISON

Structure/Mechanics	Coalinga Area	Black Butte Area
Fold belt proximity to San Andreas	35 km	140 km
Fold orientation compared to San Andreas (N400W)	N600W	North-South
En echelon folds? - regional extent	Yes 100 km	Possibly yes 60 km
Youngest surface deformed related to folds	2,000 y.b.p.	500,000 y.b.p.
Eastward directed thrust invading eastward directed wedges	Present	Absent
Reverse faults associated with folds and emanating from basement	Present	Questionable
Seismogenic activity	Present	Absent

In Coalinga there are 2,000-year-old stream terraces undergoing deformation. At Black Butte, major stream terrace systems extending throughout the 40-km study radius (centered on the dam) were examined in detail and found to show no tectonic deformation for the last 200,000 years. The only deformation found is at least 500,000 years old and lies 20 or more km away from the dam. Very strong confidence can be had in the resolution of the three-dimensional subsurface within a 10-km radius of the dam. No capable structure was found within this zone. Nonetheless, two seismogenic sources, one at Stony Creek Fault and one at Corning Dome, will be considered, and these sources can be visualized as related to wedge tectonics if desired.

Bolt (1982) feels there should be mention of possible earthquakes generated because of crustal loading by water impounded in the reservoir. However, he concludes "the amount of elastic energy added to the system by the reservoir loading is insignificant for earthquakes of engineering consequence." Experience with crustal loading from small lakes (such as Black Butte) vindicate this stand.

Finally, there has been no evidence whatsoever of induced earthquakes associated with Black Butte reservoir during its life since 1963. Especially significant is the fact that the California DWR monitoring for the past 4 years (1978 through 1982) has shown no microearthquakes in the reservoir area. By comparison with various case histories, a dam height of 140 feet is grouped with dams of low significance for induced seismicity.

5.2 Effective Seismic Sources. Table 5-2 below is the result of the study identifying major faults and structure that could possibly be seismogenic.

TABLE 5-2
SEISMIC SOURCES FOR BLACK BUTTE DAM

Name	Distance to Dam	Structural Mechanism	Length	Activity During Project Life?	Recurrence of Faulting
Story Const. Fr. 15	27 l.m	D4===14=	100 km	No	20 000
Stony Creek fault	27 km	Dip-slip		<del>-</del>	30,000 years ago
Willows fault	25 km.	Dip-slip	60 km	Yes	Assume Holocene
Maacama fault	95 km	Strike-slip	120 km	Yes	Holocene
San Andreas fault	140 km	Strike-slip	1000 km	Yes	Holocene
Great Valley	20 km	Thrust	60 km.	No	500,000 years
anticlines					ago
Cleveland Hill					
fault	63 km	Normal-slip	3 km	Yes	Holocene

5.3 <u>Tectonic Blocks</u>. Based on the tectonic model developed in prior sections and figures, we assign the magnitude of events to zones described in table 5-3 below.

TABLE 5-3
MAGNITUDES BY ZONES

Block	Distance	to Dam	Magnitude	
Sierra Nevada	40	km	6.0	
Eastern Coast Ranges	20	km	6.0	
Western Coast Ranges	140	km	8.3	
Northern Sacramento Valley	20	km	5.5	

5.4 Earthquake Source Area. For purposes of design, we follow the technique that the event and source must be localized on a candidate structure. For the western Coast Ranges the overwhelming choice is the San Andreas Fault. For the northern Sacramento Valley and Black Butte Dam, the matter becomes more obscure. The magnitude of the event can be empirically linked to the length of rupture along a candidate fault or the area of fault plane and stress drop. Empirical approaches such as Slemmons' (1977) curves and/or Randall (1973), Kanamori, et al. (1975), and Gibowicz (1973) are useful for estimation. Another obscurity is the matter of tectonic folds existing without a clear, causative fault. Here our magnitude is based on tabulation of limited data and our recent experience. Our premise is that the size of the fold, as indicated by dimensions, wave length, and maximum closure can be useful in

identifying magnitude potential. We believe, in this case, that magnitude potential as indicated by strain/stress accumulation is proportional to elastic energy stored in the deforming plane fold body underlying.

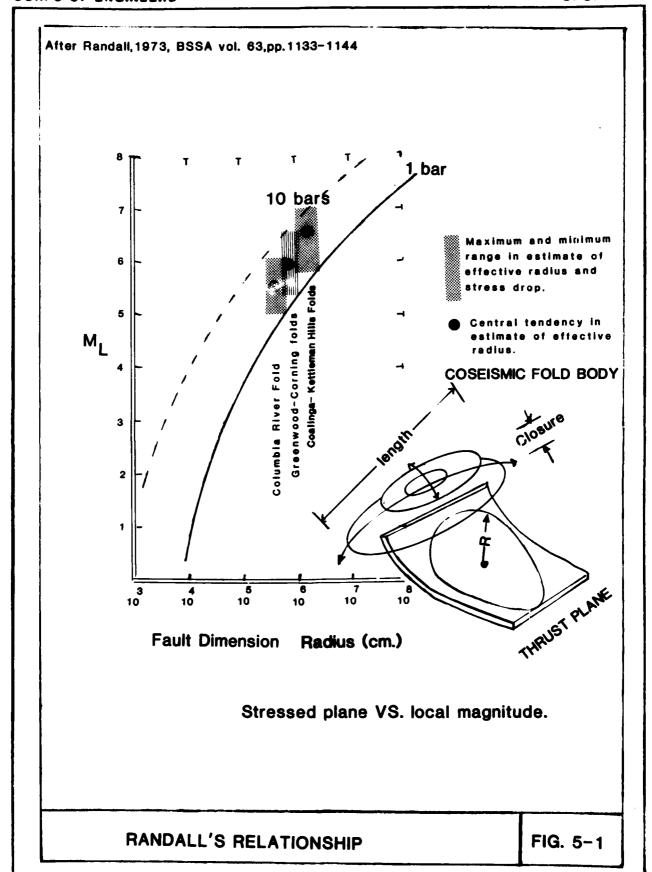
Table 5-4 below, while not exhaustive, serves to suggest credible limits in competent, layered, folded rock by comparing the Corning-Greenwood area with the two other known areas of folded structure that have earthquake history.

Two magnitudes are listed under "Mechanism" on table 5-4. For reverse faulting secondary to folding (2), the maximum stress plane down-dip width is probably no greater than the width of the fold (or wave length). Small surface areas result for the plane of stress accumulation and, if it is seismogenic, only moderate magnitudes are possible. If folding is secondary to faulting (1), then deeper reverse fault planes are involved. The down-dip fault length may vary from 7 km (the immediate basement level) to 14 km (the depth of focus for some deeper interior valley events). Fault length in both cases will be limited to the surface length of fold segments (along axis of complete closure). The deepest events in the interior valley are very few and are most probably related to plate tectonics. In summary, a magnitude 6.3 is the maximum conservative, credible event possible on the Greenwood or Corning anticlines. This is based on the upper uncertainty added to the calculated value for mechanism (1) of table 5-4.

TABLE 5-4
EARTHQUAKE POTENTIAL FROM FOLDS

	Historic Maximum	Segmented Length	Length of Closure		Mechani	sm* (2)
Columbia River Basalts~Umtanum/ Gable Mtn	5.3	20 km	850 m	$M_{L} = 5.5$	5 ( <u>+</u> .5)	$M_{L} = 5.1 \ (\pm .6)$
Kettleman Hills/ Coalinga Anticline	6.7	35 km	1980 m	$M_{\rm L} = 6.7$	( <u>+</u> .3)	$M_{L} = 5.4 (+.4)$
Corning-Greenwood Dome	None	21 km	300 m	$M_{\tilde{L}} = 5.8$	( <u>+</u> .5)	$M_{L} = 5.4 (+.6)$

<sup>\*</sup>Mechanism is either (1) primary reverse faulting with secondary folding or (2) primary folding with secondary reverse faulting. ML is obtained by Randall's relationship for 1-10 bar stress drop on a plane area representing the zone of accumulation of stress underlying the folding strata. Randall's (1973) relationship converts effective radius of the stressed plate to ML at a given stress drop of 10 bars. Figure 5-1 shows this relationship.



# SECTION 6. MAXIMUM EARTHQUAKE

Finally we are in a position to recommend the selection of a maximum earthquake and a source for it. Table 6-1 below shows the candidate structures and gives estimates of their maximum magnitude and distance to the dam from the source.

TABLE 6-1

MAXIMUM EARTHQUAKE MAGNITUDES
FOR BLACK BUTTE DAM

Structure	Distance Source Site	Maximum Earthquake (ML)	Province	Source of Estimate
Stony Creek Fault $\underline{1}/$	27 km	6.5	Border Coast Ranges/ Interior Valley	Bolt, 1982
Willows Fault	25 km	6.25	lnterior Valley	Bolt, 1982
Great Valley Anticlines	20 km	6.3	Interior Valley	This report
Maacama Fault	95 km	7.0	Coast Ranges	Bolt, 1982
San Andreas Fault	140 km	8.3	San Andreas	Bolt, 1982
Cleveland Hill Fault	63 km	6.5	Sierra Nevada	USBR Auburn Dam Studies

1/As shown on the table, Stony Creek Fault poses the most significant hazard to the dam.

The data presented in table 6-1 results from a very conservative approach during the preceding steps. For example, the offset morphostratigraphic surface at Stony Creek Fault has questionable elements, and has 30,000 to 35,000 years age. It is at the limit of guidance governing capability (ETL 1110-2-301 dated 26 August 1983). The distance of the active fault segment from the dam is 27 to 30 km, a distance that is at the border of the near/far field limits for strong ground motion. Other sources are similar in situation. Source to site distance and hypocentral depth of 14 km leads to assigning the maximum event to far field conditions. The recommended MM intensity is VIII for such a maximum event. Using an approach such as referenced in Krinitzsky and Marcuson (1983) gives site acceleration at or below 0.35g as appropriate.

## SECTION 7. CONCLUSIONS AND RECOMMENDATIONS

7.1 Conclusions. Through the Dam Safety Assurance Program, Black Butte Dam was subjected to dynamic analysis studies in a prior investigation. This analysis indicates distress would result in the upstream shell of the embankment. For the dynamic analysis, seismic loading was obtained by contracting the focus of geological and seimological analysis from regional view to near-site view and existing conditions. By applying a conservative approach to resulting open questions regarding activity on nearby inferred faults, an estimate of design acceleration was obtained.

This study answers many, if not all, the open questions left by a prior study by focusing on inferred nearby faults and establishing their existence, nonexistence, and capability. The present study consists of extensive fieldwork and analysis of surface geology, subsurface geology, and geophysical subsurface data. Our evaluation of all data involved has subsequently led to the conclusion that the Black Butte Fault and Willows Fault system as proposed do not exist within 10 miles around the dam. With the near-field elimination of the Black Butte Fault and Willows Fault system as being a concern, other sources were investigated. The tectonic capability of doubly plunging anticlinal folds in the subsurface near the midvalley axis was assessed. There is no visual evidence that these folds are the result of tectonic wedging and thrust faulting at depth. No subsurface evidence of faulting was found to account for a drag folding origin of the anticlines and they have no seismogenic history. Based on minor offsets along a few young Holocene terraces on a Stony Creek Fault segment near Thomes Creek, 27 km from the dam, the maximum (credible) event was chosen as a magnitude 6.5 event. The source is on the youngest segment of the fault as this is judged to be the only capable segment.

Because of the few number of earthquakes experienced and the limited historic earthquake record in Glenn and Tehama Counties, calculating the exceedance chance for the maximum event during the project life span becomes an academic exercise as does existing earthquake recurrence intervals. Instead, historic experience in northern California, as a whole, indicates a local magnitude of 5-3/4 as the maximum event in areas away from the San Andreas Fault zone during any 100-year period.

Based on present results, we find that prior estimates of seismic loading derived by establishing the maximum event on close-in faults like the inferred black Butte Fault have been overly conservative. A reduction in design acceleration level is justified as no viable seismic source is closer than 27 km to black Butte Dam.

- 7.2 Recommendations. The following recommendations are made as a result of this study:
- a. The previously inferred Black Butte Fault does not exist, and should be deleted from geologic maps and given no further consideration.

- b. The MCE for seismic safety analysis of Black Butte Dam is a magnitude 6.5 crustal event on the Stony Creek Fault with a source distance of 27 km.
- c. The inferred Willows Fault system of Harwood and Helley (1982) need not be considered in evaluating seismic sources in the northern Sacramento Valley.
- d. A local magnitude of 5-3/4 as the maximum event, based on historic experience, is recommended to be a substitute for a statistical event and to serve as the 100-year probable earthquake. The source for this event is either near the Coast Ranges-Great Valley contact or east of the Willows Fault that lies southeast of the dam.



\$65 Miles (4)

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#### **GLOSSARY**

- O Soil stratigraphy is used as a technique to define stratigraphic markers in Quaternary deposits and land forms.
- O Geosoil is a weathering profile that formed at or immediately beneath and essentially parallel with the land surface. It has physical characteristics, and stratigraphic relationships that permit its consistent recognition and mapping and whose stratigraphic interval or relationship to immediately older and younger deposits is known quite definitely (Morrison, 1967).
- Morphostratigraphic unit is a distinct stratigraphic unit as a body that is identified primarily from the surface form it displays; it may or may not have distinct lithologic character, and it may or may not be time transgressive in extent.
- Soil stratigraphy is used as stated in Begg (1968) for Glenn County and in Gowan (1967) for Tehama County. Alluvial soils (Xerothents and Xerofluvents) are used as young stratigraphic markers. (They have undeveloped to slightly developed soil profile.) Zonal soils (Haploxeralfs, Durixeralfs, and Palexeralfs) are used as local geosoils and are correlated to stage chronology division (relative time) of the Quaternary by Shlemon (1980), and others in the western Sacramento Valley.
- o <u>Transpression</u> is crustal deformation as an intermediate stage between compression and strike-slip motion. It occurs in zones with oblique compression and wrench faulting.

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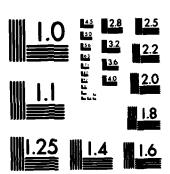
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AREA OF CONCERN:

39.000 TO 40.750 NORTH LATITUDE 121.250 TO 123.500 WEST LONGITUDE

MINIMUM RICHTER MAGNITUDE = -1.0 HYPOCENTER DEPTH LIMITS = 0 TO 999 KM

LOCATION OF BLACK BUTTE DAM

39.813 NORTH LATITUDE, 122.336 WEST LONGITUDE

# EXPLANATION OF EARTHQUAKE CATALOG HEADINGS:

1. OBSERVED,

(SEE "CD" BELOW)

EO # UNIQUE EARTHQUAKE NUMBER ASSIGNED TO EACH EARTHQUAKE IN THE U.S. ARMY CORP OF ENGINEERS MASTER EARTHQUAKE FILE. DATE DATE IS IN GREENWICH MEAN TIME UNLESS OTHERWISE NOTED IN THE YEAR MO-DY COMMENT COLUMN. TIME TIME IS GREENWICH MEAN TIME UNLESS OTHERWISE NOTED IN THE HRMN:SECS CONNENT COLUMN. LAT. LATITUDE IN DECIMAL DEGREES NORTH OF THE EQUATOR. LON. LONGITUDE IN DECIMAL DEGREES WEST OF THE PRIME MERIDIAN. DISTANCE FROM EPICENTER TO SITE IN MILES (MI) AND KILOMETERS (KM) DIST DPTH HYPOCENTER DEPTH IN KILOMETERS. A "\*" AFTER THE DEPTH INDICATES THAT IT WAS DETERMINED BY THE USER RATHER THAN OBSERVED. MMI MAXIMUM MODIFIED MERCALLI INTENSITY. CRM CALCULATED RICHTER MAGNITUDE. MAGNITUDE OF EVENT AS:

CD MAGNITUDE DETERMINATION CODE. (SEE "CRH"ABOVE)

1. BLANK IMPLIES MAGNITUDE WAS OBSERVED,

2. "I" IMPLIES MAGNITUDE WAS CALCULATED FROM INTENSITY,

2. CALCULATED FROM INTENSITY (USUALLY CRM=2/3(MMI)+1), OR 3. GEMERATED BY THE USER (USING A PROJABILITY TECHNIQUE).

3. "G" IMPLIES MAGNITUDE WAS GENERATED BY THE USER.

REF SEISHIC DATA REFERENCE (SEE DOCUMENTATION FOR U.S. ARMY CORPS OF ENGINEERS MASTER EARTHQUAKE FILE).

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1899   12-13   0: 0.0   39.800   121.800   29   46   0   0.00   MOAA   1899   12-20   0: 0.0   39.800   121.800   29   46   0   0.00   MOAA   1891   12-20   0: 0.0   39.800   121.800   29   46   0   0.00   MOAA   1901   11-14   300: 0.0   39.800   121.800   29   46   0   11   2.33   1   MOAA   1901   11-14   300: 0.0   39.500   122.000   28   45   0   VII   4.50   CDMG   CDMG   1904   4-16   920: 0.0   40.500   122.000   47   76   0   VI   4.50   CDMG   REDDING AREA   1904   10-29   0: 0.0   39.500   122.500   50   80   0   0.00   MOAA   1906   5-7   410: 0.0   39.200   122.900   52   84   0   VII   5.67   1   CDMG   CDMG   MOAA   1906   5-7   410: 0.0   39.500   122.300   6   10   0   1   1.67   1   MOAA   1906   5-16   552: 0.0   39.900   122.300   6   10   0   1   1.67   1   MOAA   1909   7-7   1240: 0.0   39.500   122.200   23   37   0   0.00   MOAA   1909   1-7   1240: 0.0   39.500   122.200   23   37   0   0.00   MOAA   1909   1-7   1240: 0.0   39.500   122.200   23   37   0   0.00   MOAA   1909   1-7   1240: 0.0   39.500   122.200   23   37   0   0.00   MOAA   1909   1-7   1240: 0.0   39.500   122.200   23   37   0   0.00   MOAA   1909   1-39   130: 0.0   39.167   122.917   54   87   0   0.00   BRK   1910   3-19   30: 0.0   40.500   121.500   65   104   0   0.00   BRK   1914   10-23   1430: 0.0   40.500   121.500   65   104   0   0.00   BRK   1914   10-23   1430: 0.0   40.500   121.500   65   104   0   0.00   BRK   1915   7-22   0: 0.0   40.500   121.500   65   104   0   0.00   BRK   1915   7-22   0: 0.0   40.500   121.500   65   104   0   0.00   BRK   1915   7-22   0: 0.0   40.417   121.567   53   86   0   0.00   BRK   1915   7-22   0: 0.0   40.417   121.567   53   86   0   0.00   BRK   1915   7-22   0: 0.0   40.417   121.567   53   86   0   0.00   BRK   1915   7-22   0: 0.0   40.507   121.500   65   104   0   0.00   BRK   1915   7-22   0: 0.0   40.507   121.500   65   104   0   0.00   BRK   1915   7-22   0: 0.0   40.507   121.500   65   104   0   0.00   BRK   1915   7-22   0: 0.0   40.567   121.507												
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1920 7-23 355: 0.0 40.500 121.817 55 88 0 VII 5.67 I CDHG SHASTA COUNTY 1920 7-23 1400: 0.0 40.500 121.817 55 88 0 0.00 BRK 1920 7-23 1600: 0.0 40.500 121.817 55 88 0 0.00 BRK 1920 7-23 2000: 0.0 40.500 121.817 55 88 0 0.00 BRK 1920 9-17 620: 0.0 39.000 122.917 64 103 0 0.00 BRK 1920 12-29 959: 0.0 39.500 122.167 24 38 0 V 4.33 I CDHG GLENN COUNTY					60	97						
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1920 7-23 1600: 0.0 40.500 121.817 55 88 0 0.00 BRK 1920 7-23 2000: 0.0 40.500 121.817 55 88 0 0.00 BRK 1920 9-17 620: 0.0 39.000 122.917 64 103 0 0.00 BRK 1920 12-29 959: 0.0 39.500 122.167 24 38 0 V 4.33 I CDMG GLENN COUNTY								VII		I		SHASTA COUNTY
1920 7-23 2000: 0.0 40.500 121.817 55 88 0 0.00 BRK 1920 9-17 620: 0.0 39.000 122.917 64 103 0 0.00 BRK 1920 12-29 959: 0.0 39.500 122.167 24 38 0 V 4.33 I CDMG GLENN COUNTY												
1920 9-17 620: 0.0 39.000 122.917 64 103 0 0.00 BRK 1920 12-29 959: 0.0 39.500 122.167 24 38 0 V 4.33 I CDNG GLENN COUNTY							-					
1920 12-29 959: 0.0 39.500 122.167 24 38 0 V 4.33 I CDMG GLENN COUNTY												
								v		Ţ		CLENN COUNTY

DATE VEAD-MO-DV	TINE	ECC	LAT.	LON.	DIST NI			MMI	CRM C	D:	REF.	COMMENTS
YEAR-NO-DY										-		
1921 12- 4	1657:	0.0	39.000	122.917	64	103	٥		0.00		BRK	
1922 12- 4	1657:	0.0	39.000	122.917	64				0.00		BRK	
1923 10-24						103	0		0.00		BRK	
1923 10-27	1847:	0.0	39.000	122.917	64	103	0		0.00		BRK	
1925 12-18						97	0		0.00		BRK	
1926 10-13						115	0		0.00		BRK	
1927 3- 1						104	0		0.00		BRK	
1927 3- 1						104	0		0.00		BRK	
1928 4-15						28		VIII			CDNG	GLENN COUNTY
1928 4-16						104	0		0.00		BRK	
1928 6- 4						115		VII			CDMG	WEAVERVILLE AREA
1929 6-14						79	0		0.00		BRK	
1929 11- 4						95	0		0.00		BRK	
1929 12- 5						95	0		0.00		BRK	
1930 1-19 1930 3-10						88	0		0.00		BRK	
1930 4-20						54 101	Ö		0.00		NOAA BRK	
1930 4-29						97	Ö		0.00		CDMG	REDDING AREA
1930 10-29						84	Ö		0.00		BRK	REDVING AREA
1930 10-29						84	ŏ		0.00		BRK	
1930 10-29							ŏ	VI	4.50		CDMG	WHITENORE AREA
1930 10-29							ŏ		0.00		BRK	
1930 10-29							Ó	v	4.00		CDMG	REDDING AREA
1930 10-30						86	0		0.00		BRK	
1930 11- 1						84	0				BRK	
1930 11-11						101	0	v	4.33	1	CDMG	FOREST GLEN
1930 11-17						101	0		0.00		BRK	
1930 12- 8						96	0		0.00		BRK	
1930 12-12	1000:	0.0	39.317	123.250	60	96	0		0.00		BRK	
1930 12-12	2200:	0.0	39.317	123.250	60	96	0		0.00		BRK	
1930 12-30	1337:	0.0	40.567	122.417	52	84	0		0.00		BRK	
1931 1-20	1030:	0.0	39.417	123.317	59	95	٥	V	4.00		CDMG	WILLITS AREA
1931 1-24	721:	0.0	39.750	121.800	29	46	0	IV	4.00		CDMG	ORLAND AREA
1931 <b>9</b> -10						45	0		0.00		BRK	
1931 10- 9							0		0.00		BRK	
1931 11-11							0		0.00		BRK	
1931 11-22												
1931 11-22						87	0		0.00	_	BRK	
				123.067		84	0	IV	3.67	I	CDMG	
				122.000		45	0		0.00	_	NOAA	
				123.150		103	0	IV	3.67	I	CDMG	
				123.250		96	0		0.00	_	CDMG	
1932 10-31						110	0		3.67		CDMG	
				122.417		84	0		3.67		CDNG	
1933 6-12 1933 7- 4				121.567 121.500		86	0	14	3.67	Ţ	CDMG	
				121.500		104 81	0	ŤU	0.00	•	CDMG	
				121.500			0	14	3.67 0.00		CDMG CDMG	
				121.500			Ö		0.00		CDMG	
1936 5- 9				121.550		102	ŏ		0.00		CDNG	
				121.550					0.00		CDNG	
				121.600		87			0.00		CDNG	
				121.600		87			0.00		CDMG	
				121.600		87			0.00		CDMG	
			<del>-</del>		_		-		•			

YEAR-MO-DY						DPTH KN	HMI	CRM C	D REF.	CONNENTS
+			~							
1936 5- 9	1841: 0.0	40.350	121.600	54	87			0.00		
	2021: 0.0				96	0			I CDMG	
	218: 0.0				96	0		4.33		
	1115: 0.0				96	0		0.00		
1936 5-15 1936 5-27	704: 0.0 718: 0.0				96 87	0			I CDNG I CDNG	
	1430: 0.0				96	0		4.33		
	2226: 0.0				102	Ö		0.00		
	1245: 0.0				96	Ö			I CDNG	
1936 6-29					96	0		4.33		
1936 6-29					96	0		4.33		
	630: 0.0				87	0		4.33		
	1027: 0.0 1247: 0.0					0		4.33 3 3.67		
	2036: 0.0				96	Ö		4.33		
1936 6-30						ŏ		3.67		
1936 7- 1					113	0		4.33		
	1243: 0.0	40.450	121.300	70	113	0	٧	4.33	I CDMG	
	1550: 0.0				96	0		4.33		
	720: 0.0					0		4.33		
1936 7- 2 1936 7- 6	1738: 0.0					0		5.00		
1936 7- 6					113 113	0		4.33 : 4.33 :		
1936 7- 6						ŏ		4.33		
1936 7- 6					113	Ō		4.33		
1936 7-13	1043: 0.0	40.450	121.300		113	0	V	4.33	I CDMG	
	1714: 0.0					0		0.00	CDMG	
1936 12-25					104	0	V	4.33	I CDMG	
1938 1-11 1938 1-11					112 112	0		4.33		
1938 9- 2					85	Ŏ		0.00		
1938 11-15					85	ŏ		5.00		
1938 11-15					98			0.00		
193 <del>9</del> 6-23					104	0	V	4.33	I CDMG	
1939 8-25						0		0.00	CDMG CDMG	
1939 11-19					74					
1940 2- 8 1941 7- 6				41	66	0	ATI	5.70	CDMG	
1941 7- 6 1941 10-25				60 24	96 38	0		0.00	CDMG CDMG	
	1100: 0.0			54	87	ŏ		0.00	CDMG	
	1223: 0.0	39.150	122.150	47	75	0		0.00	CDNG	
	512: 0.0			47	75	0		0.00	CDMG	
	1642: 0.0			47	76	0		0.00	CDMG	
1942 8-11	530: 0.0			60	96	0		4.33		
1942 9-28 1942 9-29	837: 0.0 829: 0.0			47 40	75 65	0		3.67 1 4.33 1		
1942 3-29				45	72	0		5.00		
1942 11-18				45	72	ŏ		5.00		
1943 10-14	550:15.0	39.250	122.400	39	63	0+	_	3.60	BERK	
1943 10-14				38	61	0		3.70	CDNG	
1943 10-26				47	75	0	IV	3.80	CDMG	
1943 11-14 1943 11-14				40	64	0		0.00	CDMG	
1943 11-14 1943 11-15				24 42	38 67	0	v	0.00 4.00	CDNG CDNG	
1973 11 13		~~/	/	72	•	•	•	7.50	CDNG	

YEAR-MO-DY	TIME HRMN:SECS			MI	KM	KM	HHI	CRM	CD	REF.	COMMENTS
+											
1943 11-15	659: 0.0	39.250	122.200	40	64	0		0.00	)	CDMG	
1943 11-16					68	0	V	3.80	)	CDMG	
1943 11-16					64	0		0.0		CDMG	
1943 11-16					64	0	IA	3.6		CDMG	
1943 11-17					64	0	742	0.0		CDMG	
1943 12- 4					73 87	0	14	4.0		CDMG CDMG	
1944 1-10 1944 1-11					87			0.0		CDMG	
1944 1-11					87			0.0		CDMG	
1944 1-17					87	ŏ		0.0		CDMG	
1944 6- 2					88	ŏ		0.0		CDMG	
1944 7-30					115	ŏ		0.00		CDMG	
1944 8-29					104	0		3.30		CDMG	
1944 8-30					104	0		3.60		CDMG	
1944 8-30					88	0		0.00		CDMG	
1944 8-30					88	0		0.00	)	CDMG	
1944 9- 1					88	0		0.00	)	CDMG	
1944 9- 9					75	0		0.00	)	CDMG	
1944 10- 1					83	0		0.0		CDMG	
1944 12-30					87			0.00	-	CDMG	
	536:10.0				59	0	ΛI	5.00		CDMG	
1945 8-29					87			0.0		CDMG	
1945 8-29					87			0.0		CDMG	
1945 9-8 1945 9-8					110 110			0.0		CDNG CDNG	
1945 9-30					90			0.0		CDMG	
1945 10-16					87			0.0		CDMG	
1945 10-19					87	ŏ		0.0		CDMG	
1945 10-25					87	ō		0.0		CDMG	
1945 12- 4					87	ŏ		0.0		CDMG	
1946 4-19					101	Ó		0.0		CDMG	
1946 4 .3					90	0				CDMG	
1946 6-14	656: 0.0	40.350	121.600	54	87			0.0	)	CDMG	
1946 6-24					87	( 0 0		0.0	)	CDMG	
1 <del>94</del> 6 7- 7					104	0	VI	5.00		CDMG	
1946 7- 8					87			0.0		CDMG	
1947 10- 1						0			•		
1947 11- 7				37	59	0	• • • • • • • • • • • • • • • • • • • •	0.00		CDMG	
1947 11-10				19	31	0		3.90		CDMG	
	2342:12.0 2342:27.0			61 55	98	0	IA	3.6		CDMG	
	1035: 0.0			71	88 115	0	UT	0.0		CDMG CDMG	
	540: 0.0			64	103	ŏ	V.	0.0		CDMG	
	1014: 0.0			60	96	ŏ	v	4.3		CDMG	
1948 7- 8				64	103	ŏ	•	0.0		CDMG	
	1102:48.0			55	88	ŏ	IV	3.6		CDNG	
1948 11-14				26	42	ŏ		0.0		CDMG	
	1330:42.0			65	104	Ŏ		0.0		CGS	
1949 2- 1				65	104	0		0.0		CGS	
1949 3- 3				59	95	0		0.0		CDMG	
	436: 0.0			55	88	0	٧	4.3		CDMG	
1949 11- 1				34	54	0		2.0		CDNG	
1949 11-16					102	0		2.6		CDMG	
1949 11-16	804:20.0	40.517	121.567	63	102	0		3.2	0	CDMG	

DA' YEAR-1	TE NO-DY	TIMI HRMN:	E SECS	LAT.	LON.			DPTH KN	MMI	CRM	CD	REF.	CONHENTS
+													
1950	1- 1	2339:	9.0	40.500	121.500	65	104	0	IV	3.1	0	CDNG	
1950	1-16	1921:3	25.0	40.250	121.367	60	96	0	V	3.8	0	CDMG	
1950	3-13	1717:3	39.0	39.500	122.100	25	40	0	III	3.0	O I	NOAA	
1950					121.500			0		2.5		CDMG	
1950					121.467		102	0	V	5.5	0	CDMG	
1950	3-20	1716:	0.0	40.467	121.500	63	102	0		2.5	0	CDMG	
1950	3-20	1718:	0.0	40.467	121.500	63	102	0		2.6	0	CDMG	
1950	3-20	1800:	0.0	40.467	121.500	63	102	0		2.5	0	CDMG	
1 <b>95</b> 0	3-20	1819:	0.0	40.467	121.500	63	102	0		2.5	0	CDMG	
1950	3-20	1903:	0.0	40.450	121.467	63	102	0		3.4	0	CDMG	
1950	3-20	1935:	0.0	40.467	121.500	63	102	0		2.8		CDMG	
1950					121.500		102	0		2.3		CDMG	
1950					121.500			0		2.8		CDMG	
1950					121.467		102	0		3.6		CDMG	
1950	3-21				121.500		102	0		2.5		CDMG	
1950					121.467		102	0		3.1		CDMG	
1950					121.500		102	0		2.8		CDMG	
1950					121.500			0		2.5		CDMG	
1950					121.500 121.500		102	0		2.6		CDMG	
1950	-						102	0		2.5		CDMG	
1950 1 <b>95</b> 0					121.500 121.500		102 102	0		2.7		CDMG CDMG	
1950					121.500		102	0		2.6		CDMG	
1950					121.500		102	0		2.7		CDMG	
1950					121.467		102	Ŏ		3.0		CDMG	
1950					121.500		102	ŏ		2.8		CDMG	
1950					121.500		102	ŏ		2.7		CDMG	
1950	3-23				121.500		102	ŏ		2.9		CDMG	
1950					121.500	63	102	Ö		2.8		CDMG	
1950					121.467		102	ŏ		3.0		CDMG	
1950					121.467		102	0		3.8		CDMG	
1950	3~23	537:	0.0	40.450	121.467	63	102	0		3.5	)	CDNG	
1950	3-23	547:	0.0	40.467	121.500	63	102	0		2.5	)	CDMG	
1950	3-23	603:	0.0	40.467	121.500	63	102	0		2.8	)	CDMG	
1950	3-23	801:	0.0	40.450	121.467	63	102	0		3.4	)	CDMG	
1950	3-23				121.500	63	102	0		2.7	)	CDMG	
1950	3-23	955:	٥.٥	40.467	121.500	63	102	٥		2.7	0	CDMG	
1950					121.500	63	102	0		2.6	0	CDMG	
1950					121.500	63	102	0		2.6		CDMG	
1950					121.500	63	102	0		2.7		CDMG	
1950					121.500	63	102	0		2.8		CDMG	
1950					121.500	63	102	0		2.7		CDMG	
1950					121.500	63	102	0		2.5		CDMG	
1950	3-24				121.500 121.467	63	102	0		2.80		CDMG	
1950	3-24 3-24				121.500	63 63	102 102	0		3.5		CDMG CDMG	
1950 1950	3-24				121.500	63	102	Ö		2.5		CDMG	
1950	3-24				121.500	63	102	Ö		2.6		CDMG	
1950	3-24				121.500	63	102	ŏ		2.7		CDMG	
1950					121.500	63	102	ŏ		2.7		CDNG	
1950					121.467	63	102	ŏ		3.00		CDMG	
1950					121.500	63	102	ŏ		2.8		CDMG	
1950					121.467	63	102	ŏ		3.50		CDMG	
1950					121.467	63	102	ŏ		3.10		CDMG	
								-					

DATE	<b>MTW</b> 5		1.00				<b></b>	<b></b>	-	DEC	COMMENTS
YEAR-MO-DY	TIME HRMN:SECS	LAT.		DIST			HMI	CKH	CD	REF.	CONNENTS
+											
1950 3-24								2.80			
1950 3-25 1950 3-25	49: 0.0	40.450	121.467	63	102			3.00			
1950 3-25	213: 0.0	40.450	121.46/	63	102 102			3.40	<b>,</b>	CDMG CDMG	
1950 3-26								2.Q	`	CGS	
1950 4-13						Ô		3.50	Ś	CDMC	
1950 7-22					42	ŏ	IV	3.67	, 1	CDMG	
1950 10- 7					99	ō	IV	3.67	, Ī	CDMG CDMG	
1950 10- 7	1911:47.0	39.517	123.067	44		0		3.30	)	CDMG	
1950 11-14					102						
1950 11-14						0		2.60	)	CDNG	
1950 11-14	204:40.℃	40.467	121.500	63	102	0	V	4.10	)	CDMG	
1950 11-14 1950 11-14 1950 11-14	235:50.0	40.467	121.500	63	102	0	V	4.60	)	CDMG	
1950 11-14	246: 0.0	40.467	121.500	63	102	0		3.80	)	CDMG	
1950 11-14	250: 0.0	40.467	121.500	63	102	0		2.80	)	CDMG	
1950 11-14						0		2.60	,	CDMG	
1950 11-14										CDMG	
1950 11-14 1950 11-14						0		2.80	<b>,</b>	CDNG CDNG	
1950 11-14	415: 0.0	40.467	121.500	63	102	Ô		2.00	`	CDMG	
1950 11-14						ŏ		2.60	Ś	CDMG	
1950 11-14						ŏ		3.60	,	CDMG	
1950 11-14											
1950 11-14											
1950 11-14	826: 0.0	40.467	121.500	63	102			3.50	)	CDMG	
1950 11-14						0		3.50 2.90	)	CDMG	
1950 11-14											
1950 11-14											
1950 11-14					102			2.60	)	CDMG	
1950 11-14					102	0		2.60	•	CDMG	
1950 11-14					102	0		3.80	•	CDMG	
1950 11-14					102	0		<b>4.10</b>	,	CDMG	
1950 11-15 1950 11-15											
1950 11-16								3.30			
1950 11-20					94			3.00		CDMG	
	1252:17.0			57	91	ŏ		2.90		CDMG	
	1521:26.0			48	78	Ö		3,60		CDMG	
	854:43.0			68	110	0		3.40		CDMG	
1952 3-7	0: 0.0	40.600	122.400	55	88	0	IV	3.67	ľ	CDMG	
1952 3-23	1336:13.0	40.500	121.650	60	96	0	V	4.33	I	CDMG	
	1301:49.0			47	75	0		3.10		CDMG	
	952:52.0			73	117	0		2.80		CDMG	
1952 8-14				45	73	0		4.00		CDMG	
1952 10-17				66	107	0		4.00		CDMG	
1952 11-15				70	113	0		3.80		CDNG	
1952 11-16 1953 1- 2				70 <b>68</b>	113 110	0		3.80		CDMG	
	1905: 2.0			62	99	0		3.20		CDMG CDMG	
1953 1-24				42	67	Ö		3.90		BERK	
	2212: 0.0			62	99	Ö		2.80		CDMG	
	844:43.0			70	113	ŏ		3.40		CDMG	
	1655: 1.0			71	115	ō		3.30		CDMG	
	2100:22.0			73	117	Ó		2.90		CDMG	

DATE YEAR-NO-DY	TIME HRMN:SECS	LAT.	LON.	DIST NI	DIST KM	DPTH KN	NNI	CRM	CD	REF.	CONNENTS
+											*****
1953 3- 9								2.40	)	CDMG	
1953 4-2	2256:12.0	40.700	121.600					3.20		CDMG	
	1400:48.0	40.600	121.600	67	108	0		2.90		CDNG	
	335:38.0	40.600	121.600	67	108	0		2.10	)	CDNG	
	400124.0	40.467	121.550	62	99	0		2.50		CDNG	
1953 5-25	407:59.0	39.300	123.300			0	VI	3.20	)	CDMG	
1953 6-29	1108:24.0	39.717	122.367	7	11	0	V	3.10	)	CDMC	
1953 6-29	1324: 3.0	40.667	121.917	63	101	0		3.00	)	CDMG	
1953 8- 1	1719:13.0	40.600	121.417	73	117	0		3.00	)	CDMG	
1953 8-23	1221:58.0	40.467	121.717	56	90	0	V	3.30	)	CDMG	
1953 10-18	1019:58.0	40.500	121.600	62	99	0		2.80	)	CDMG	
1953 12-21					80		IV	3.10	)	CDMG	
1953 12-24	717:39.0	40.750	121.550	77	124	0		4.00	)	CDMG	
1953 12-28	644: 8.0	39.400	122.700	34	55	0		3.10	)	CDMG	
	1330: 0.0						III	3.40		CDMG	
1954 2- 4								3.20		CDMG	
1954 2- 4					_			3.40		CDMG	
1954 6-16					97			3.60		BERK	
1954 8-19								3.20		CDMG	
1954 11-10								4.40		CDMG	
1954 11-25								3.80		CDMG	
1954 12-10					116			3.30		CDMG	
1 <b>9</b> 55 8-24								3.60		CDMG	
1955 11-11					101			3.20		CDMG	
1955 11-14					98	0	IV			CDMG	
1955 11-14					96	-		3.10		CDMG	
1955 12- 1	601:38.0	40.467	121.550			-		2.90		CDMG	
1955 12- 1							IV	2.90		CDMG	
1955 12- 1								2.20		CDMG	
1955 12- 5								3.67			
1955 12- 8								3.00		CDMG	
	523: 0.0				94	-		4.33			
1956 2-10								3.00		CDMG	
	2102:11.0							3.30		CDMG	
1956 3-11					99		IV				
1956 3-18						-		3.50			
	1022: 3.0			62	100	0	VI	3.70		CDMG	
	1049:45.0			74	119	0		2.50		CDMG	
	1453:47.0			57	91	0		2.90		CDMG	
1956 4- 5	115:14.0			65	104	0	IV	3.00		CDMG	
	545: 5.0			65	104	0		2.90		CDMG	
	2026:51.0	-		53	86	0		3.20		CDMG	
	907:56.0			63	102	0		3.30		CDMG	
	1751: 5.0			55	89	0	IV	3.00		CDMG	
	1740:28.0			67	108	0		2.60		CDNG	
	1836:48.0			67	108	0	14	3.60		CDMG	
	1850:12.0			67	108	0	•	2.40		CDNG	
	1937:53.0			67	108	0		2.80		CDMG	
	1942:44.0			67	108	0	TA	3.40		CDMG	
	2116:46.0			67 67	108	0		3.00		CDMG	
1956 5-15 1956 5-16	2139:51.0 354:38.0			67 67	108	0		2.90		CDMG	
	534: 2.0			67	106	0		2.70		CDMG CDMG	
	553:45.0			67	108	Ö		2.80		CDNG	
1330 3-10	9991701V	74.54/	154,79/	9/	700	•		2.00	•	CURIC	

DATE YEAR-NO-DY	TIME HRMN:SECS	LAT.	LON.	DIST NI	DIST KN	DPTH KM	HMI	CRM	CD	REF.	COMMENTS
+											
1956 E-16	925.25 A	40 549	101 467	63	400	^		2 7	^	CDMG	
1956 5-16 1956 5-17	2329:23.0				108 108	0		2.70 3.20		CDMG	
	1030: 0.0				108	ŏ	IV	3.6		CDMG	
	230: 3.0				108	Ŏ	•	2.6		CDMC	
	626:56.0				108	Ö		2.40		CDMG	
	510:56.0				104	ō		3.0		CDMG	
1956 6- 3					78	0		2.9	)	CDMG	
1956 6-28				62	99	0		3.10	)	CDMG	
1956 6-29				60	96	0		2.90	)	CDMG	
	150:39.0			73	117	0		2.40	)	CDMG	
	605:32.0				117	0		2.40		CDMG	
	1958: 4.0				98	0		2.60		CDMG	
	1616:46.0				108	0		2.10		CDMG	
	1747:16.0			65	104	0		2.20		CDMG	
	1904:55.0			73	118	0		2.90		CDMG	
	2302: 1.0			79	127	0		3.30		CDMG	
	444:23.0			73	117	0		2.20		CDMG	
1956 8- 3	339:41.0				118 104	0		2.90 3.20		CDMG CDMG	
1956 8- 7					104	Ö		3.2		CDMG	
1956 9-18					106	Ö	•	2.9		CDNG	
1956 10- 4					104	ŏ		2.7		CDMG	
1956 10-23					97	ŏ		2.9		CDMG	
1956 11-25					101	ŏ		3.5		CDMG	
1956 11-28					104	Ö		2.8		CDMG	
1956 12-19					101	0	IV	3.8		CDMG	
1957 1- 7	1239: 0.0	39.250	122.617		67	0		3.3		CDMG	
	415:46.0				93	0	IV	2.9	0	CDMG	
1957 1-22					93	0		2.8	0	CDMG	
1957 3- 2					21	0		3.0		CDMG	
1957 3-28					32	0		3.1		CDMG	
1957 9-25					91	0		2.7		CDMG	
1957 11-22					91	0		2.7		CDMG	
	1148: 2.0				114	0		3.3		CDMG	
	2005:56.0				119	0		2.9		CDMG	
1958 1-27 1958 1-28	1522: 2.0					0		2.9		CDMG	
	258:23.0				/5 35	0	TU	3.4		CDMG	
	1009: 1.0				35 55	ŏ	14	2.1		CDMG CDMG	
	1101:52.0				55	ŏ		3.8		CDMG	
	1945:48.0				91	ŏ		2.7		CDMG	
	2006:49.0				53	ŏ		2.1		CDMG	
1958 9-23	529:30.0	40.100	121.500	48	78	0		2.6		CDMG	
1958 10-24	259:24.0	40.400	121.600	56	90	0		3.0	0	CDMG	
1958 11- 3	32:14.0	40.600	122.000	57	92	0		2.5	0	CDMG	
	1607:35.0				101	0	III	2.8	0	CDMG	
	1319:16.0				124	0		2.9		CDMG	
	1140:54.0				92	0		3.1		CDMG	
	1030:47.0				81	0		2.2		CDMG	
	506:15.0					0		3.1		CDMG	
	1301:46.0					0		2.6		CDNG	
	1944:37.0					0		2.3		CDMG	
_	1709:29.0							2.4		CDMG	
1959 2-17	1330:56.0	40.067	125.000	39	63	0		3.2	J	CDMG	

DATE YEAR-NO-DY	TIME HRMN:SECS	LAT.	LON.	DIST MI	DIST KN	DPTH KM	MMI	CRM	CD	REF.	CONNENTS
+											*****
	1313:50.0				31	0	1	1.6	7 I	NOAA	
	249:49.0				21	0		3.5		CDMG	
	608:22.0				94	0	VI	3.6		CDMG	
	1835:21.0				20	0		3.20		CDMG	
1959 5- 2 1959 5- 9			122.200		47	0		3.20		CDMG	
	1246: 4.0 404:13.0				58 51	0		2.9		CDNG CDNG	
	426:11.0				23	Ö		3.0		CDMG	
	1419:59.0				88	ŏ		2.8		CDMG	
	24:21.0				63	Ō		3.0		CDMG	
1959 9-16					81	0		3.5	0	CDNG	
1959 9-24					73	0	VI	3.6	0	CDNG	
1959 9-27					128			3.0		CDMG	
1959 9-30					49	_		2.5		CDMG	
1959 10- 5					85	0		2.8		CDMG	
1959 11- 1 1959 11-11					39			2.8		CDMG	
1959 11-11			123.300		101 20	0		3.20		CDMG CDMG	
	1903:56.0				90	Ö		2.9		CDMG	
	1706:11.0				103	ŏ		3.0		CDMG	
1960 4-8					110	Ŏ		2.9		CDMG	
1960 4-23					37	0		3.5		CDMG	
1960 4-27	153:53.0	40.600	123.400	78	126	0		2.6	)	CDMG	
1960 7-24	1312: 3.0	39.867	121.867	25	41	0		2.4	)	CDMG	
1960 7-24	1953:47.0	39.200	123.500	75	121	0		2.9		CDMG	
	2135:27.0				104	0		3.0		CDMG	
1960 11- 2					67	0		3.10		CDMG	
1960 11-22					100	0		2.8		CDMG	
1960 12- 8					117	0	IV	3.10		CDMG	
	842:45.0				46	0		3.4		CDMG	
	907:28.0 513:46.0				98	0		3.6		CDNG CDNG	
	902:28.0				126 107	0		4.0		CDMG	
1961 10-14					96	ŏ		2.7		CDMG	
1961 10-14					48	ŏ		3.5		CDMG	
1961 10-15					90	Ō		2.7		CDMG	
1961 12-21				47	76	0		2.7		CDMG	
1962 4- 1			122.067	54	87	0		3.0		CDMG	
	1750: 6.2			73	118	5	VII	5.2		CDMG	
	818: 5.2			63	101	2		2.8		CDMG	
1962 12- 6				55	89	7		3.10		CDMG	
	1637:15.0 1400:32.0			16 55	25 88	0		3.20		CDNG CDNG	
1963 6- 8 1963 6-20				78	125	0		3.0		CDMG	
	535:35.5			51	82	Ö	tv	3.20		CDMG	
	516:46.8			37	60	ŏ		3.3		CDNG	
	2243:16.3			41	66	Ö		3.4		CDMG	
1963 9-25	416:41.3	39.417	123.467	66	107	ŏ	•	3.1		CDMG	
1963 12- 7				6	10	0		3.10		CDMG	
	1044:10.0			50	81	0		3.0		CDMG	
1964 10-27			121.550	45	73	0		3.5		CDMG	
	1825:17.0			25	40	16		3.60		CDNG	
	2128:50.0			25 52	40	0	***	3.20		CDMG	
1965 11- 8	330: 1.1	70.000	121.400	52	83	0	111	3.0	<i>.</i>	NOAA	

DATE			LON.		DIST KN	DPTH KN	MNI CRM CD	REF.	CONNENTS
YEAR-NO-DY	ARMRISECS			MI					
1965 12-12	1257:49.8	40.200	121.467	53	86	٥	4.00	CDMG	
	1831: 0.0				121	33	3.60	CDMG	
	329:36.0				98	16	2.50	CDMG	
	349:55.1				49	20	VI 4.60	CDMG	
	226:20.0			51	82	0	3.40	CDMG	
	1305:39.9				29	0	3.00	CDMG	
	1214:56.3	39.600	122.900	34	54	0	2.90	CDNG	
	1535: 8.1				37	0	3.40	CDMG	
	2030:23.0				78	0	3.40	CDMG	
	1552: 7.5				98	0	IV 4.10	CDMG	
	1351:57.0				56	0	3.10	CDMG	
1967 6-26	1515:35.0	39.300	123.300	63	101	2	VI 3.50	CDMG	
1967 9-11	1857:28.5	40.617	123.317	76	122	0	3.50	CDMG	
1967 11-24	1512:13.6	39.450	122.500	27	43	0	3.60	CDMG	
1967 11-29	255: 2.9	40.400	121.950	45	73	0	3.10	CDMG	
	240:40.0				89	0	3.50	CDMG	
1968 4-29	21:38.6	39.533	122.017	25	41	0	VI 4.70	CDMG	
	1145:57.8					0	3.10	CDMG	
	313: 9.9					0	IV 3.20	CDMG	
	105: 2.5						v 3.30	CDMG	
	107:12.0						3.30	CDMG	
	7 2357:17.2						3.30	CDMG	
	1606:16.0						3.40	CDMG	
	538:55.2						V 3.10	CDMG	
	3 1933:2 <b>8.</b> 0				_		3.30	CDMG	
1970 4-23	526:14.0						IV 3.70	CDMG	
1970 7- 1						_	V 3.60	CDMG	
	1856:59.4						IV 3.50	CDMG	
	1152:29.2						3.00	CDMG	
	5 <b>410:18.</b> 0						3.30	CDMG	
	654:25.0						IV 4.10	CDMG	
	941:41.0						3.30	CDMG	
	943:39.0						3.40	CDMG	
	837:21.4						3.50	CDMG	
	2 1053: 8.0						3.10	CDMG	
	3 1458:32.5						2.80	CDMG	
	935:48.0							CDMG	
1970 12-20					_	_	3.00	CDMG	
	4 1443: 1.1	39.800	122.800	25			3.00	BERK	
	9 1751:24.0						3.00	BERK	
1971 3-2					_		3.00	CDMG CDMG	
1971 7-			123.500				3.50		
	1 1236:34.0						3.60	CDMG CDMG	
1971 7-2							3.20 3.20	CDMG	
	7 546:52.0 6 2229:48.3						IV 3.80	CDMG	
							3.00	CDMG	
1971 8-3	8 1124: 6.3						3.10	CDMG	
1971 10-2			7 123.35				3.10	CDMG	
	3 2137:10.0						3.50	CDMG	
	3 1307:38.						IV 3.60	CDMG	
							3.10	CDNG	
	2 220: 0.0				_	-	3.00	CDNG	
	5 1922:36.0						3.50 III 3.50	CDMG	
1972 3-	5 1510: 2.0	U 37.31	144.71	7 4	v 9.	• 0	111 3.50	<b>√Dild</b>	

								-		
DATE	TIME	LAT.	LON.	DIST		DPTH	MMI	CRM CD	REF.	CONHENTS
YEAR-NO-DY				HI	KM	KM				
+										
1972 3-21								3.00	CDMG	
	1609:22.0				50	0		3.10	CDMG	
	2227:59.0				92	0	IV	3.40	CDMG	
1972 4-11					83	0		3.10	CDMG	
1972 5-25	300:52.0	40.250	123.000	47	75	0		3.10	CDMG	
1972 5-27	2204:16.0	39.400	122.600	32	51	0		2.70	CDMG	
	1838:17.0	40.120	121.430	53	85	0		2.80	BERK	
	618:49.0			55	88	0		3.40	CDMG	
	1553:31.0			45	72	0		3.20	CDMG	
	1442: 9.2	40.017	121.300	57	91	0	V	4.10	CDMG	
	1559:59.4			46	74	0		3.10	CDMG	
1973 2-15				12	19	0		3.00	CDMG	
1973 5-13				52	83	0	V	3.10	CDNG	
	1348:51.2			66	106	0	V	3.80	CDMG	
1973 10-13				54	87	0		3.10	CDMG	
1973 11-12				71	114	0		4.40	CDMG	
1973 11-25				58	94	0		3.50	CDMG	
1974 12-21				58	94	0		3.20	CDMG	
1974 12-27				83	133	0		3.90	CDMG	
1975 1- 5				52	83	17*		~.50	CDWR	OROVILLE DAM
1975 1-13				65	104	5*		0.80	CDWR	MANZANITA LAKE SE
	1735:52.6			30	49	5#		10	CDWR	PRINCETON
	456:54.5			28	45	7*		0.10	CDWR	PRINCETON
	118:48.0			34	54	5#		50	CDWR	HAMLIN CANYON
	1353:24.8			23	37	12#		0.40	CDWR	NORD
	215:41.3			16	25	4*		0.50	CDWR	HAMILTON CITY
	1621:39.3			21	33	4#		0.30	CDWR	GLENN
	1845:33.2			27	44	8*		0.50	CDWR	RICHARDSON SPRINGS
1975 3-19	2059:37.6	39.222	122.478	42	67	6		3.30	USGS	
	2059:38.0	39.233	122.592	42	68	5*		3.40	CDWR	CLEARLAKE OAKS NE
1975 3-20	1506:36.9	39.371	123.059	49	79	1*		3.20	CDWR	POTTER VALLEY
1975 4-10	936:11.4	39.638	121.879	27	44	27*		0.40	CDWR	ORD FERRY
1975 4-18	1632:58.2	39.237	121.826	48	78	5#		2.90	CDWR	SUTTER BUTTES
1975 5-13	1409:19.4	39.788	122.322	2	3	0#		2.90	CDWR	BLACK BUTTE DAM
1975 5-20	1254:27.9	39.658	122.101	16	26	5*		0.50	CDWR	HAMILTON CITY
1975 5-21	1547:14.7	40.008	121.515	45	73	20#		0.30	CDWR	BUTTE MEADOWS SE
1975 5-25	1634:12.4	39.383	123.210	55	89	4#		3.40	CDWR	POTTER VALLEY NW
1975 5-27				17	27	5*		0.70	CDWR	VINA
	1401:42.0			56	90	13*		3.40	CDWR	POTTER VALLEY NW
	851:58.9			42	68	5#		40	CDWR	PARADISE SE
	810: 9.2			31	50	24*		0.10	CDWR	NELSON
	2308: 5.9			18	29	0#		2.30	CDWR	VINA
	816: 2.8			27	44	3*		0.30	CDWR	LOGANDALE
1975 6- 7				48	78	9*		30	CDWR	PALERNO
	1102:10.8			46	74	10*		0.40	CDWR	BUTTE HEADOWS SE
	1639:10.5			50	80	1*		3.20	CDWR	POTTER VALLEY
	843:15.6			39	62	10*		50	CDWR	CHEROKEE
_	231:54.7		_	48	78	8*		40	CDWR	PALERHO
	419:53.1			45	72	6		3.60	USGS	
	443:29.6			45	73	0*		2.60	CDWR	PALERNO
	504:10.6			50	80	6*		20	CDWR	PALERHO
	639:41.1			48	77	7*		10	CDWR	OROVILLE
	1131:24.0			49	79	6*		70	CDWR	PALERNO
1975 6-28	1416:19.1	39.486	121.605	45	72	3*		2.80	CDWR	PALERNO

DATE TIME LAT. LON. DIST DIST DPTH MMI CRM CD REF. COMMENTS YEAR-NO-DY HRMN:SECS MI KM KM CDWR 1975 6-28 1821:59.4 39.486 121.531 48 78 8+ -.10 PALERNO 1975 6-29 2020:39.9 39.478 121.545 48 77 8\* -.20 CDWR PALERNO 0.40 1975 7- 5 857: 9.9 40.138 122.018 28 45 3\* CDWR TUSCAN SPRINGS 1975 7- 5 1819:15.8 39.456 121.535 49 79 7\* CDWR -.30 **PALERNO** 1975 7-6 611: 3.2 39.516 121.504 49 79 6\* CDWR **OROVILLE** -.40 1975 7- 6 1321:60.0 39.513 121.486 50 80 5\* 2.80 CDWR OROVILLE DAM 78 9\* -.20 1975 7- 6 1435:45.9 39.487 121.528 48 CDWR **PALERNO** 1975 7- 7 14:35.5 39.478 121.529 49 79 8\* -.20 CDWR PALERMO 1975 7- 7 1054: 4.7 39.481 121.523 49 79 9\* -.30 CDWR **PALERMO** 1975 7- 7 2104:33.6 40.365 121.680 52 83 4\* 2.30 CDWR LASSEN PEAK SW 1975 7- 8 620:20.7 39.398 122.148 30 49 0# 1975 7- 8 2119:58.8 39.474 121.534 48 78 8# CDWR 0.90 LOGANDALE CDWR -.30 PALERNO 1975 7- 9 556:36.4 39.753 121.656 37 59 5\* -.10 CDWR PARADISE SW 1975 7-10 1258:11.1 40.011 121.490 47 76 8# 0.10 CDWR JONESVILLE SW 1975 7-14 1949: 0.4 39.474 121.522 49 79 9# **CDWR** -.50 **PALERNO** 1975 7-16 1551:35.9 39.971 121.460 48 77 1# CDWR 2.40 PULGA NW 1975 7-20 508:43.1 39.275 123.171 58 93 12\* REDWOOD VALLEY CDWR 3.40 1975 7-20 928: 9.1 39.448 121.589 47 76 1\* CDWR 3.20 PALERNO 1975 7-20 1038:52.2 39.460 121.550 48 78 9\* CDWR PALERNO -.10 1975 7-21 417:37.0 39.464 121.595 46 74 0\* 3.10 CDWR **PALERNO** 1975 7-21 514:49.1 39.499 121.507 49 79 7\* 2.50 CDWR PALERMO 1975 7-22 1232:55.4 39.603 121.815 31 50 24\* CDWR NELSON -.40 1975 7-25 2253:16.4 39.699 122.176 12 19 0+ 1975 8- 1 1545:37.8 39.450 121.531 50 80 7 1975 8- 1 1627:17.8 39.438 121.537 50 80 5 2.90 ORLAND CDWR 3.80 USGS V 4.70 USGS 1975 8- 1 1726:50.1 39.462 121.538 49 79 9 USGS 3.00 1975 8- 1 2020: 4.8 39.439 121.528 50 81 8 4.50 USGS 1975 8- 1 2020:12.9 39.439 121.528 50 81 15 VII 5.70 USGS 1975 8- 1 2025: 0.0 39.438 121.528 50 81 0 4.70 **OROVILLE** 1975 8- 1 2029; 0.0 39.438 121.528 50 81 0 4.60 OROVILLE 1975 8- 1 2032:39.8 39.445 121.507 51 82 5 USGS 3.00 1975 8- 1 2046:18.4 39.473 121.500 50 81 6 3.80 USGS 1975 8- 1 2105:39.8 39.433 121.487 52 84 7 3.00 USGS 50 81 8 1975 8- 1 2121:50.7 39.442 121.528 IV 4.10 USGS 1975 8- 1 2125:59.0 39.474 121.518 50 80 7 3.30 USGS 1975 8- 1 2129:24.1 39.452 121.549 79 49 7 3.60 USGS 49 79 1975 8- 1 2344:41.0 39.486 121.522 7 3.20 USGS 1975 8- 2 52:48.5 39.484 121.509 50 80 7 3.80 USGS 1975 8- 2 631:57.2 39.447 121.484 52 84 6 3.20 USGS 1975 8- 2 1011:53.7 39.490 121.512 49 **79** 7 3.10 USGS 1975 8- 2 1151:50.7 39.473 121.488 51 82 2 3.40 USGS 1975 8- 2 1724:29.2 39.474 121.471 52 83 6 4.30 USGS 1975 8- 2 1743:24.1 39.478 121.474 52 83 6 4.00 USGS 1975 8- 2 1958:36.9 39.448 121.537 50 80 7 3.10 USGS 1975 8- 2 2022:16.3 39.445 121.463 53 85 0 5.10 OROVILLE 1975 8- 2 2035:48.6 39.471 121.482 52 83 6 3.90 USGS 
 1975
 8- 2 2059: 0.0 39.432 121.466
 53
 86
 0

 1975
 8- 2 2059: 2.7 39.406 121.711
 43
 70
 5

 1975
 8- 3 103: 5.8 39.488 121.518
 49
 79
 8

 1975
 8- 3 247: 8.8 39.478 121.501
 50
 81
 7

 1975
 8- 4 947:45.0 39.421 121.523
 51
 82
 8
 5.20 OROVILLE 5.10 USGS 4.60 USGS 4.10 USGS 3.50 3.20 USGS 1975 8- 5 2044:24.5 39.415 121.515 52 83 7 USGS 1975 8-6 350:29.9 39.479 121.524 49 79 7 IV 4.70 USGS 1975 8- 6 1641:52.1 39.497 121.529 48 78 8 3.60 USGS

DATE TIME YEAR-NO-DY HRMN:SECS	LAT.	LON.	DIST NI	DIST KN	DPTH KN	HMI	CRM CD	REF.	CONHENTS
+									
1975 8- 6 2100:33.5	39.440 1	21.485	52	84	9		3.00	USGS	
1975 8- 7 2031:20.4	39.517 1	21.533	47	76	9		3.10	USGS	
1975 8-8 700:50.1	39.502 1	21.512	49	79	8		4.90	USGS	
1975 8- 8 1337:53.9				81	6		3.20	USGS	
	39.446 1		52	84	4	٧	4.30	USGS	
1975 8-11 1559: 5.3			48	77	6		3.60	USGS	
1975 8-16 548: 9.4	39.472 1	21.521	50	80	9		4.00	USGS	
1975 8-21 1407:30.6	39.445 1	22.043	30	48	0#		0.60	CDWR	PRINCETON
1975 9- 1 2141:46.0	40.121 1	21.685	40	65	0#		3.20	CDWR	BUTTE HEADOWS SW
1975 9- 3 2120:53.4	39.494 1	21.542	48	77	3#		2.80	CDWR	PALERNO
1975 9- 4 117: 1.6	39.438 1	21.635	45	73	7*		3.30	CDWR	BIGGS
1975 9- 4 1539:24.9			46	74	0#		2.70	CDWR	PALERNO
1975 9- 4 1603:43.6			54	87	1*		2.70	CDWR	HONCUT
1975 9- 5 2101:38.9			45	73	7*		3.40	CDWR	PALERNO
1975 9- 6 137: 5.2			47	76	7*		3.10	CDWR	PALERNO
1975 9- 6 942:38.1			51	82	3*		2.80	CDWR	PALERMO
1975 9- 6 2110:53.8			42	67	11*		2.80	CDWR	SHIPPEE
	39.652 1		40	64	5*		2.70	CDWR	HANLIN CANYON
	39.456 1		47	75	1+		2.70	CDWR	PALERNO
1975 9- 7 1207:10.5			52	84	0#		2.70	CDWR	BANGOR
1975 9- 7 1235: 6.7			56	90	5*		1.20	CDWR	LONA RICA
1975 9- 9 329:28.4			50	81	0#		2.60	CDWR	PALERHO
1975 9-10 1238:30.3			43	69	5*		3.00	CDWR	OROVILLE
1975 9-10 1723:46.0			43	69	0*		2.90	CDWR	OROVILLE
1975 9-10 1739: 4.5			42	67	4*		3.50	CDWR	SHIPPEE
1975 9-10 1850:39.4			42	68	0*		3.10	CDWR	OROVILLE
1975 9-12 200:47.7			46	74	0#		3.50	CDWR	OROVILLE
1975 9-13 1827:50.3			50	81	0*		2.60	CDWR	PALERNO
1975 9-15 1447:11.1			47	75	1*		3.00	CDWR	BIGGS
1975 9-17 1904:30.4			26	42	0#		0.30	CDWR	RICHARDSON SPRINGS
	39.793 1		26	42	0*		0.40	CDWR	RICHARDSON SPRINGS
1975 9-20 653:14.7			44	71	2* 6*		2.80	CDWR	OROVILLE PALERNO
1975 9-20 836:41.3			50 50	80 81	2*		2.50	CDWR CDWR	
1975 9-20 1222:34.4				75	8*		2.50 2. <b>8</b> 0	CDUR	PALERNO
1975 9-21 1749: <b>58.4</b> 1975 9-22 2125: <b>9.</b> 2			47 50	80	2*		2.70	CDWR	PALERMO PALERMO
	39.840 1		22	36	0#		0.60	CDWR	NORD
	39.485 1		54	87	0#		1.80	CDWR	BANGOR
1975 9-25 1346:23.0			49	79	6*		2.70	CDWR	PALERNO
	39.490 1		47	75	9	IV	4.00	USGS	I ROLANO
	39.511 1		45	72	6*		2.00	CDWR	OROVILLE
1975 9-26 435:52.8			47	75	8*		2.90	CDWR	PALERNO
1975 9-26 536:45.0			50	80	4*		1.90	CDWR	PALERNO
1975 9-26 815: 8.3			46	74	4*		2.30	CDWR	BIGGS
1975 9-26 957:15.2			47	75	10	IV	2.90	USGS	
1975 9-26 957:16.1			46	74	10*		3.20	CDWR	PALERNO
1975 9-26 1249:31.9			48	77	7*		2.30	CDWR	PALERNO
1975 9-26 1255:56.0			50	80	5*		1.70	CDWR	PALERNO
1975 9-27 247:10.2			47	76	0#		2.50	CDWR	PALERNO
1975 9-27 1305:14.8			45	72	1*		2.60	CDWR	OROVILLE
1975 9-27 2234:37.7			43	70	6*		3.90	CDWR	OROVILLE
1975 9-27 2234:38.0			47	76	Ō		4.60		OROVILLE
1975 9-27 2240:14.4			43	69	4+		3.00	CDWR	OROVILLE
1975 9-27 2248: 7.1	39.404 13	21.404	57	92	3*		2.40	CDWR	BANGOR

DATE	TIME		LON.				NNI CRN CD	REF.	CONNENTS
YEAR-NO-DY				MI	KM	KN			
<b>†</b>									
1975 9-27	2251:12.3	39.600	121.496	47	76	0#	2.00	CDWR	OROVILLE DAM
1975 9-27	2253:46.1	39.595	121.486	48			2.10	CDWR	OROVILLE DAM
1975 <b>9</b> -27						_	2.50	CDWR	OROVILLE DAM
	2304:30.4						3.10	USGS	
	2307:43.3				70	0#	2.60	CDWR	OROVILLE
	2328: 4.7					10	3.10	USGS	0000000
1975 9-27 1975 9-28						2*	3.00	CDWR	OROVILLE
1975 9-28	3:49.3		121.607		71 77	0#	2.50	CDWR	OROVILLE
	146:57.0					0# 8#	2.40 2.40	CDWR CDWR	OROVILLE DAM OROVILLE
1975 9-28	422:14.9					_	2.50	CDWR	SHIPPEE
1975 9-28	905: 2.1				72	7*	1.80	CDWR	OROVILLE
	925:37.3			50	80	0#	2.60	CDWR	PALERNO
	930:15.0			47	76	9#	1.80	CDWR	OROVILLE
	1959:50.9			47	76	7#	2.20	CDWR	OROVILLE
1975 9-28	2107:14.3	39.552	121.591	43	70	6#	3.50	CDWR	OROVILLE
1975 9-29	203:36.3	39.478	121.565	47	76	4#	2.70	CDWR	PALERHO
	1009:36.5				73	0#	2.10	CDWR	PALERMO
1975 9-29							2.00	CDWR	OROVILLE
1975 9-30							2.50	CDWR	PALERNO
	2236: 4.3						2.90	CDWR	SHIPPEE
1975 10- 1							2.60	CDWR CDWR	PALERNO BIGHARDSON SDRINGS
1975 10- 1 1975 10- 2							0.10 <b>4</b> 0	CDWR	RICHARDSON SPRINGS OROVILLE DAM
1975 10- 2					_	-	2.80	CDWR	PALERNO
1975 10- 2							3.30	CDWR	PALERNO
1975 10- 2							2.60	CDWR	BANGOR
1975 10- 3							2.70	CDWR	BIGGS
1975 10- 3					_	-	2.90	CDWR	BIGGS
1975 10- 3					-	-	2.50	CDWR	OROVILLE DAM
1975 10- 3					_		2.80	CDWR	BIGGS
1975 10- 3					75	0#	2.50	CDWR	PALERHO
1975 10- 4					80	7*	3.10	CDWR	PALERHO
1975 10- 4					64	4#	2.60	CDWR	SHIPPEE
1975 10- 5						_	0.20	CDWR	HAMILTON CITY
1975 10- 8							2.10	CDWR	BIGGS
1975 10- 9								CDWR	BANGOR
1975 10-10 1975 10-10							IV 3.50 2.50	USGS CDWR	HAMI TH CANVON
1975 10-10							2.20	CDWR	HAMLIN CANYON OROVILLE DAM
1975 10-12							2.40	CDWR	BANGOR
1975 10-12							2.10	CDWR	OROVILLE
1975 10-12						_	20	CDWR	PALERMO
1975 10-13							3.00	CDWR	OROVILLE
1975 10-13							3.00	USGS	
1975 10-13	1606:51.5	39.518	121.544	47	75		3.20	CDWR	OROVILLE
1975 10-13	2130: 8.4	39.537	121.441	52		-	2.50	CDWR	OROVILLE DAM
1975 10-14							1.90	CDWR	OROVILLE DAN
1975 10-14							50	CDWR	BANGOR
1975 10-14							2.00	CDWR	OROVILLE
1975 10-16							2.50	CDWR	PALERNO
1975 10-16							2.20	CDWR	OROVILLE DAM
1975 10-18						_	0.50	CDWR	HONCUT
1975 10-18	1357:42.4	JY.440	121.513	51	82	10*	10	CDWR	PALERNO

	CH I REGG							-		
YEAR-NO-DY		LAT.	LON.	DIST MI	DIST KM	DPTH KM	MMI	CRM CD	REF.	CONNENTS
+										
1975 10-19	1000.12 0	20 605	122 012	10	31	7*		0.20	CDWR	HAMILTON CITY
1975 10-19						7*		0.50	CDWR	PANTHER SPRING SW
1975 10-20						0*		2.90	CDWR	OROVILLE
1975 10-21						0#		30	CDWR	PALERNO
1975 10-23						0#		3.10	CDWR	PRINCETON
1975 10-23						0#		2.50	CDWR	PALERNO
1975 10-27						6*		3.00	CDWR	OROVILLE
1975 10-28					74	10	TTT	3.40	USGS	
1975 10-28						6*		2.50	CDWR	OROVILLE
1975 10-30						6*		50	CDWR	OROVILLE DAM
1975 10-30						2*		50	CDWR	OROVILLE DAM
1975 11- 1						3*		2.80	CDWR	BANGOR
1975 11- 1						1#		50	CDWR	PALERNO
1975 11- 3					77	0#		2.60	CDWR	OROVILLE DAM
1975 11- 3	1603:37.2	39.424	121.618	47	75	0#		2.90	CDWR	PALERNO
1975 11- 3	2023:36.4	39.682	122.041	18	29	1*		2.90	CDWR	HAMILTON CITY
1975 11- 3						1#		2.90	CDWR	PALERNO
1975 11- 4						0#		2.40	CDWR	PALERHO
1975 11- 4						0#		2.40	CDWR	BIGGS
1975 11- 5						9	II	3.30	USGS	
1975 11- 5				48	78	4#		2.40	CDWR	OROVILLE DAM
1975 11- 5						3#		30	CDWR	BANGOR
1975 11- 6					47	0#		0.00	CDWR	RICHARDSON SPRINGS
1975 11- 7					91	3*		60	CDWR	BANGOR
1975 11- 7					87	6#		50	CDWR	BANGOR
1975 11- 7						1#		2.70	CDWR	PALERNO
1975 11- 8						26*		0.50	CDWR	RICHARDSON SPRINGS
1975 11- 8					82 87	1# 0#		2.60 40	CDWR CDWR	BANGOR BANGOR
1975 11- 9 1975 11- 9					79	6*		70	CDWR	PALERNO
1975 11- 9						10*		30	CDWR	OROVILLE
1975 11-10						2*		30	CDWR	BANGOR
1975 11-10						0#		50	CDWR	BANGOR
1975 11-10						8#		60	CDWR	OROVILLE
1975 11-10						2*		70	CDWR	BANGOR
1975 11-10				51	82	4#		30	CDWR	PALERNO
1975 11-12				48	78	7#		40	CDWR	OROVILLE
1975 11-13				25	41	0+		0.50	CDWR	RICHARDSON SPRINGS
1975 11-13				52		3*		1.00	CDWR	PALERHO
1975 11-13				50	81	3+		1.00	CDWR	PALERNO
1975 11-13	2335:53.7	39.415	121.474	53	86	1=		20	CDWR	BANGOR
1975 11-14	1129: 9.4	39.399	121.469	55	88	4+		50	CDWR	BANGOR
1975 11-14				52	83	1*		80	CDWR	PALERNO
1975 11-14				52	84	2*		70	CDWR	BANGOR
1975 11-14				51	82	0#		80	CDWR	PALERNO
1975 11-14				53	86	3*		50	CDWR	BANGOR
1975 11-14				50	80	6#		60	CDWR	PALERNO
1975 11-14				48	78	2#		40	CDWR	PALERNO
1975 11-15				49	79	5	111	4.00	USGS	BANCOS
1975 11-15				53 55	85	0*		2.70	CDWR	BANGOR
1975 11-15 1975 11-15				55 53	88 85	0# 0#		20 30	CDWR	BANGOR
1975 11-15				53 53	86	0#		30	CDWR CDWR	BANGOR
1975 11-15				53	86	1#		70	CDWR	BANGOR BANGOR
12/2 11-12	130113212	JJ. 703	-44.703	93	90	7*		.,,	CDAK	DARUK

DATE YEAR-NO-DY	TIME HRMN:SECS		LON.	DIST	DIST KM	DPTH KN	HMI	CRM CD	REF.	COMMENTS
+										
1975 11-15				52	84	0#		0.00	CDWR	PALERNO
1975 11-15				53		0#		10	CDUR	BANGOR
1975 11-15				53	85	2*		0.00 90	CDWR CDWR	BANGOR PALERNO
1975 11-15 1975 11-16				52 50	83 80	3# 6#		70	CDWR	OROVILLE DAM
1975 11-16				50	_	10*		50	CDWR	PALERNO
1975 11-16						1*		50	CDWR	BANGOR
1975 11-16				52	84	0#		2.90	CDWR	PALERNO
1975 11-16				52	84	1*		30	CDWR	PALERNO
1975 11-17				54	87	2*		40	CDWR	BANGOR
1975 11-17				51	82	0#		30	CDWR	BANGOR
1975 11-18				53	86	0#		40	CDWR	BANGOR
1975 11-18 1975 11-19			121.4/5	52	83	0#		2.90	CDWR	BANGOR
1975 11-19				26 50	42 80	0# 6#		0.20 40	CDWR CDWR	RICHARDSON SPRINGS PALERNO
1975 11-20				53	85	1#		20	CDWR	BANGOR
1975 11-20				53	86	0#		30	CDWR	BANGOR
1975 11-20				53		0#		20	CDWR	BANGOR
1975 11-21				54		4#		70	CDWR	BANGOR
1975 11-22	445:19.2	39.416	121.453	55	88	0#		10	CDWR	BANGOR
1975 11-22					79	0#		3.60	CDWR	BRUSHY MTN.
1975 11-22						0#		30	CDWR	BANGOR
1975 11-22				48		7#		20	CDWR	PALERNO
1975 11-23 1975 11-23				53 53		0# 0#		20	CDWR	BANGOR
1975 11-23				33 19	86 30	5#		2.70 0.40	CDWR CDWR	BANGOR GERBER
1975 11-23				52	83	3*		20	CDWR	PALERMO
1975 11-24				53	85	0#		90	CDWR	BANGOR
1975 11-24				51	82	5#		40	CDWR	PALERNO
1975 11-24				53		4#		60	CDWR	PALERNO
1975 11-24	1410:22.1	39.476	121.516	50	80	10*		40	CDWR	PALERMO
1975 11-25				50	81	4#		30	CDWR	BANGOR
1975 11-26					43	1*		0.10	CDWR	RICHARDSON SPRINGS
1975 11-27						0#		20	CDWR	BANGOR
1975 11-27		-				5*		20	CDWR	PALERNO
1975 11-30 1975 12- 1				48 51		8#		40	CDWR	OROVILLE
1975 12- 1				51 52		5# 0#		60 60	CDWR CDWR	BANGOR
1975 12- 1				53		0#		60	CDWR	BANGOR Palerno
1975 12- 2				52		4#		~.60	CDWR	PALERNO
1975 12- 2				52		1#		30	CDWR	PALERNO
1975 12- 3	133:11.5	39.400	121.515	52	84	6#		1.10	CDWR	PALERNO
1975 12- 3	450:23.6	39.436	121.479	53	85	0*		30	CDWR	BANGOR
1975 12- 3				56	90	5#		80	CDWR	BANGOR
1975 12- 3				53		0#		3.00	CDWR	BANGOR
1975 12- 3				55	88	0#		30	CDWR	BANGOR
1975 12- 4				54		0#		50 50	CDWR	BANGOR
1975 12- 4 1975 12- 5				26 54	42 87	0# 0#		0.50 0.00	CDWR	RICHARDSON SPRINGS
1975 12- 5				53		0#		2.70	CDWR CDWR	BANGOR BANGOR
1975 12- 5	_			49		4#		<b>6</b> 0	CDWR	PALERMO
1975 12- 5				53		1#		50	CDWR	BANGOR
1975 12- 5				53		0#		30	CDWR	BANGOR
1975 12- 6	1404:53.8	39.506	121.474	50	81	0#		2.20	CDWR	OROVILLE DAM

YEAR-NO-DY			LON.	DIST MI		DPTH KN	MMI	CRM CD	REF.	COMMENTS
+										
1975 12- 7	606:15.4	39.414	121.507	52	84	4#		50	CDWR	PALERNO
1975 12- 7	1342: 8.0	39.509	121.509	49	79	7#		60	CDWR	OROVILLE
1975 12- 9	2120: 4.9	39.489	121.489	50	81	0#		1.90	CDWR	BANGOR
1975 12-10	519:45.9	39.489	121.494	50	81	6#		50	CDWR	BANGOR
1975 12-11	852:32.5	39.459	121.492	51	82	0#		2.50	CDWR	BANGOR
1975 12-11					42	0#		0.00	CDWR	RICHARDSON SPRINGS
1975 12-12					78	4#		50	CDWR	OROVILLE DAM
1975 12-13					85	0#		2.80	CDWR	BANGOR
1975 12-13					84	6*		20	CDWR	PALERMO
1975 12-13					87	3#		70	CDWR	BANGOR
1975 12-13					81	6*		30		BANGOR
1975 12-14					86	1#		30	CDWR	BANGOR
1975 12-14					89	5#		0.60	CDWR	SUTTER
1975 12-15	13:53.6	39.925	121.562	42	67			0.60	CDWR	PARADISE NE
1975 12-15 1975 12-15	722:27.7	39.401	121.446	55 47	89 76	5# 5#		90 30	CDWR CDWR	BANGOR PENNINGTON
1975 12-16					80	5*		50	CDWR	BANGOR
1975 12-16					78	8*		50	CDWR	OROVILLE
1975 12-16					81	10*		60	CDWR	PALERNO
1975 12-18					80	7#		70		OROVILLE DAM
1975 12-18					87			70	CDWR	BANGOR
1975 12-18					83	5#		50	CDWR	PALERHO
1975 12-19					86	0#		20	CDWR	BANGOR
1975 12-19					81	0#		2.50	CDWR	PALERNO
1975 12-19					91	14*		20	CDWR	LOMA RICA
1975 12-19					82	5#		50	CDWR	PALERNO
1975 12-20					86	0#		20	CDWR	BANGOR
1975 12-20					73	0#		2.40	CDWR	PALERNO
1975 12-20	1356:46.7	39.493	121.496	50	80	5#		30	CDWR	BANGOR
1975 12-20	1754:44.6	39.285	121.470	59	95	17*		20	CDWR	LOMA RICA
1975 12-21					88	0#		20	CDWR	BANGOR
1975 12-22					80	9*		20	CDWR	PALERNO
1975 12-22					87	2*		60	CDWR	BANGOR
1975 12-22					84			60	CDWR	BANGOR
1975 12-23					75			40	CDWR	OROVILLE
1975 12-23					85	_		30	CDWR	BANGOR
1975 12-24										BANGOR
1975 12-24				53	86	2#		80	CDWR	BANGOR
1975 12-24				49	79	6*		60	CDWR	OROVILLE DAM
1975 12-25				53	85	4*		40	CDWR	BANGOR
1975 12-26 1975 12-26				53 53	86 86	1* 0*		30 40	CDWR CDWR	BANGOR BANGOR
1975 12-26				34	55	11*		0.10	CDWR	HAMLIN CANYON
1975 12-26				54	87	1#		20	CDWR	BANGOR
1975 12-28				53	86	0#		10	CDWR	BANGOR
1975 12-28				53	85	0#		70	CDWR	BANGOR
1975 12-29				54	87	3#		50	CDWR	BANGOR
1975 12-29				53	86	3*		80	CDWR	BANGOR
1975 12-29				52	84	3#		80	CDWR	PALERNO
	1322:13.2			50	80	7*		10	CDWR	PALERMO
<del>-</del> - ·	1022:41.5			58	94	18*		50	CDWR	LONA RICA
	1858:39.4			53	85	0#		2.70	CDWR	BANGOR
	522:59.7			53	86	1*		30	CDWR	BANGOR
1976 1- 2	1022: 9.7	39.453	121.473	52	84	0#		40	CDWR	BANGOR

DAT YEAR-N		TINE HRMN:SECS	LAT.	LON.	DIST NI	DIST KM	DPTH KN	NNI	CRM CD	REF.	CONNENTS
+											
1076	1 - 2	1030:14.6	20 201	101 510	53	85	0*		30	CDWR	PALERNO
	1- 2	917:24.0			52	83	1*		40	CDWR	PALERNO
		1435:38.6			52	83	0#		2.40	CDWR	PALERNO
	1- 4	309:14.9			52	84	0#		80	CDWR	PALERNO
		1048:32.6			54	87	0#		0.00	CDWR	BANGOR
		1313:42.9			47	76	5#		60	CDWR	OROVILLE
		1541:38.1			48	78	6*		0.70	CDWR	OROVILLE
		1550: 4.3			50	80	12*		0.50	CDWR	PALERNO
	1- 5	825:16.7			54	87	1*		80	CDWR	BANGOR
1976	1~ 5	1509:26.6	39.425	121.509	52	83	4#		50	CDWR	PALERNO
		1840:46.4			52	84	0#		30	CDWR	BANGOR
		1926:21.9			58	94	5*		20	CDWR	LOMA RICA
		1308:32.4			57	91	0#		60	CDUR	BANGOR
		1627:21.8			44	71	1#		3.00	CDWR	PARADISE SE
		2158:38.1			52	84	_		30	CDWR	BANGOR
	1- 9	57:26.3			53	86	7*		50	CDWR	HONCUT PALERNO
1976	1- 9	921:58.8			53				50 2.70	CDWR CDWR	BANGOR
1976		1754:26.0					_		0.90	CDWR	LOGANDALE
1976		1449: 3.5 2126:24.6							80	CDWR	BANGOR
1976		2126:24.6							80	CDWR	BANGOR
1976 1 <b>976</b>	_	1125:47.9							90	CDWR	BANGOR
1976	1-11								0.90	CDWR	BANGOR
1976	-	1307:47.0							50	CDWR	BANGOR
1976	1-13	317:35.1							10	CDWR	LOMA RICA
1976		2054:43.3				85	0#		60	CDWR	BANGOR
1976		2159:35.2				69	18*		70	CDWR	BIGGS
1976	1-13	2216:51.0	39.390	121.572	50	81	0#		50	CDWR	PALERNO
1976	1-14	1412:47.7	39.407	121.455	55	88	0#		50	CDWR	BANGOR
1976	1-14	2140:31.7	39.657	122.090	17	27	6*		0.70	CDWR	HAMILTON CITY
1976	1-15	450: 9.7	39.404	121.490	53	86	0#		70	CDWR	BANGOR
1976	1-15	2041:18.1	39.414	121.518	52	83	1*		40	CDWR	PALERNO
1976	1-16		39.412	121.487			0#		40	CDWR	BANGOR
1976	1-16						0#		30	CDWR	BANGOR
1976		2105:23.6							2.10	CDWR	BANGOR
1976		2331:45.2							2.20	CDWR	OROVILLE DAN
1976	1-17								70	CDWR	PALERNO
1976	1-17								60	CDWR	PALERNO
1976	1-17								3.10 <b>4</b> 0	CDWR CDWR	PALERNO Bangor
1976	1-17	1232:48.4							50	CDWR	BANGOR
1976 1976	1-17			121.494					3.30	CDWR	BANGOR
1976		1446:47.8							30	CDWR	BANGOR
1976		1516:12.2							80	CDWR	BANGOR
1976		1534: 5.3							1.00	CDWR	PALERNO
1976	1-20								20	CDWR	BANGOR
1976	1-21								70	CDWR	PALERMO
1976	1-22								30	CDWR	BANGOR
1976	1-22								0.80	CDWR	BANGOR
1976	1-22								30	CDWR	BANGOR
1976		1157:37.8				2 8	4 0#		1.90	CDWR	BANGOR
1976		1212:33.9				3 8	5 0#	,	40	CDWR	BANGOR
1976		1313:18.9						•	2.30	CDWR	BANGOR
1976	1-25					4 8	7 0*	,	20	CDWR	BANGOR

1976 1-25 801:21.9 39.451 121.491 52 83 0*40 CDWR BANGOR 1976 1-26 56: 9.1 39.498 121.515 49 79 8*30 CDWR PALERNO 1976 1-26 755: 3.7 93.405 121.503 53 85 2*60 CDWR PALERNO 1976 1-26 849: 7.9 39.405 121.503 53 85 2*60 CDWR PALERNO 1976 1-26 849: 7.9 39.414 121.509 52 84 1* 1.00 CDWR PALERNO 1976 1-26 1940: 0.0 39.585 122.091 22 35 8* 3.00 CDWR GLENN 1976 1-26 1940: 0.0 39.434 121.559 48 78 10 3.00 USGS 1976 1-26 1940: 0.0 39.434 121.559 48 78 10 3.00 USGS 1976 1-27 1850: 1.3 39.435 121.557 53 86 0* 1.80 CDWR BANGOR 1976 1-27 1850: 1.3 39.405 121.457 53 86 0* 1.80 CDWR BANGOR 1976 1-28 352:19.5 39.409 121.545 50 81 4* 2.60 CDWR DALERNO 1976 1-28 352:19.5 39.419 121.545 50 81 4* 2.60 CDWR BANGOR 1976 1-28 352:19.5 39.419 121.545 50 81 4* 2.60 CDWR BANGOR 1976 1-28 465: 2.6 39.538 121.705 37 60 0*40 CDWR BANGOR 1976 1-28 465: 2.6 39.583 121.705 37 60 0*40 CDWR BANGOR 1976 1-28 352:19.5 39.499 121.550 54 87 24*20 CDWR BANGOR 1976 1-28 352:19.5 39.493 121.550 54 87 24*20 CDWR BANGOR 1976 1-28 409:12.37 39.404 121.550 54 87 24*20 CDWR BANGOR 1976 1-28 409:12.37 39.404 121.550 54 87 24*20 CDWR BANGOR 1976 1-28 352:19.5 39.439 121.457 56 90 5*40 CDWR BANGOR 1976 1-28 352:19.5 39.539 121.455 56 90 5*40 CDWR BANGOR 1976 1-29 315:25.3 39.434 121.579 51 00 0*40 CDWR BANGOR 1976 1-29 315:25.3 39.434 121.579 51 00 0*40 CDWR BANGOR 1976 1-29 315:25.3 39.434 121.579 51 00 0*40 CDWR BANGOR 1976 1-29 315:25.3 39.434 121.579 51 00 0*40 CDWR BANGOR 1976 1-29 315:25.3 39.434 121.457 53 86 1*50 CDWR BANGOR 1976 1-29 315:25.3 39.434 121.457 54 80 5*70 CDWR BANGOR 1976 1-29 123:10.0 39.431 121.470 53 86 1*50 CDWR BANGOR 1976 1-30 36:10.75 39.491 121.450 53 86 0*40 CDWR BANGOR 1976 1-30 36:10.75 39.491 121.450 53 86 0*70 CDWR BANGOR 1976 2-1 15:13.8 39.438 121.450 53 86 0*70 CDWR BANGOR 1976 2-2 1 15:13.8 39.438 121.450 53 86 0*70 CDWR BANGOR 1976 2-2 1 15:13.8 39.438 121.450 53 86 0*70 CDWR BANGOR 1976 2-1 180:13.8 39.438 121.450 53 86 0*00 CDWR BANGOR 1976 2-2	DATE YEAR-MO-D	TIME Y HRMN:SECS	LAT.	LON.	DIST	DIST KM	DPTH KM	HMI	CRM CD	REF.	COMMENTS
1976   1-26   56; 9.1   39,498   121,515   49   79   8*  30   CDWR   PALERNO   1976   1-26   755; 3.7   39,405   121,503   53   85   2*  60   CDWR   PALERNO   1976   1-26   1940; 0.7   39,414   121,509   52   84   1*   1.00   CDWR   PALERNO   1976   1-26   1940; 0.7   39,426   121,501   52   35   8*   3.00   CDWR   PALERNO   1976   1-26   1940; 0.7   39,426   121,501   52   63   5*   3.30   CDWR   PALERNO   1976   1-26   1940; 0.7   39,426   121,501   52   63   5*   3.30   CDWR   PALERNO   1976   1-26   1940; 0.7   39,426   121,501   52   63   5*   3.30   CDWR   PALERNO   1976   1-27   1850; 1.23   39,365   121,487   53   86   0*   1.80   CDWR   BANGOR   1976   1-27   1350; 1.23   39,365   121,387   59   95   4*  80   CDWR   BANGOR   1976   1-28   742;57.8   39,367   121,487   55   89   1*  10   CDWR   BANGOR   1976   1-28   742;57.8   39,367   121,487   56   87   24*  20   CDWR   HONCUT   1976   1-28   742;57.8   39,367   121,487   56   90   5*  40   CDWR   LOWA RICA   1976   1-28   202;54.9   39,493   121,493   50   81   4*  60   CDWR   BANGOR   1976   1-28   202;54.9   39,493   121,493   50   81   4*  60   CDWR   BANGOR   1976   1-29   315;25.3   39,412   12,537   50   80   1*  60   CDWR   BANGOR   1976   1-29   315;25.3   39,431   121,537   50   80   1*  60   CDWR   BANGOR   1976   1-29   493;23.7   39,391   121,490   53   86   1*  10   CDWR   BANGOR   1976   1-29   315;25.3   39,438   121,547   53   86   1*  10   CDWR   BANGOR   1976   1-29   315;25.3   39,438   121,493   50   81   4*  60   CDWR   BANGOR   1976   1-29   315;25.3   39,438   121,493   50   81   4*  60   CDWR   BANGOR   1976   1-29   315;25.3   39,438   121,493   50   81   4*  60   CDWR   BANGOR   1976   1-30   600; 6.7   39,458   121,495   51   82   0*  50   CDWR   BANGOR   1976   2-1   15:31.8   39,438   121,465   53   86   1*  10   CDWR   BANGOR   1976   2-1   15:31.8   39,438   121,475   54   67   0*  50   CDWR   BANGOR   1976   2-1   1934;13.9   39,404   121,475   54   67   0*  5	+										
1976   1-26   56; 9.1   39,498   121,515   49   79   8*  30   CDWR   PALERNO   1976   1-26   201142,5   39,401   121,503   53   85   2*  60   CDWR   PALERNO   1976   1-26   848; 7.9   39,416   121,503   53   85   2*  60   CDWR   PALERNO   1976   1-26   1959; 34,0   39,565   122,091   22   55   89   3.00   CDWR   PALERNO   1976   1-26   1940; 0.7   39,426   121,501   52   83   5*   3.30   CDWR   PALERNO   1976   1-26   1940; 0.7   39,426   121,501   52   83   5*   3.30   CDWR   PALERNO   1976   1-26   1940; 0.7   39,426   121,507   53   86   0*   1.80   CDWR   BANGOR   1976   1-27   1850; 1.23   39,365   121,87   59   95   4*  80   CDWR   BANGOR   1976   1-27   1350; 1.23   39,365   121,87   59   95   4*  80   CDWR   BANGOR   1976   1-28   742;57.8   39,367   121,457   50   81   4*   2.80   CDWR   PALERNO   1976   1-28   742;57.8   39,367   121,457   56   90   5*  40   CDWR   BANGOR   1976   1-28   742;57.8   39,367   121,457   56   90   5*  40   CDWR   BANGOR   1976   1-28   262;54,9   39,493   121,493   50   81   4*  60   CDWR   BANGOR   1976   1-28   200;54,9   39,493   121,493   50   81   4*  60   CDWR   BANGOR   1976   1-29   315;25,3   39,412   12,537   50   80   1*  60   CDWR   BANGOR   1976   1-29   315;25,3   39,412   12,537   50   80   1*  60   CDWR   BANGOR   1976   1-29   315;25,3   39,434   121,537   50   80   1*  60   CDWR   BANGOR   1976   1-29   315;25,3   39,438   121,460   55   88   5*  70   CDWR   BANGOR   1976   1-29   315;25,3   39,438   121,460   55   88   5*  70   CDWR   BANGOR   1976   2-1   15:31.8   39,438   121,465   53   86   1*  10   CDWR   BANGOR   1976   2-1   15:31.8   39,438   121,465   53   86   1*  10   CDWR   BANGOR   1976   2-1   15:31.8   39,438   121,465   53   86   0*  50   CDWR   BANGOR   1976   2-7   1246;27.0   39,528   121,548   52   38   8*  50   CDWR   BANGOR   1976   2-9   139;13.1   39,489   121,475   54   67   67   67   67   67   67   67   6	1976 1-2	5 901.21 0	20 451	121 491	<b>5</b> 2	83	0#		- 40	CDMB	RANGOR
1976   1-26   201   42   5   39   431   12   497   52   84   4*   2.90   CDWR   BANGOR   1976   1-26   346: 7.9   39   415   12   509   52   84   1.00   CDWR   PALERHO   1976   1-26   1940: 0.0   39   435   12   1.559   48   78   10   3.30   CDWR   CLENN   1976   1-26   1940: 0.0   39   434   12   1.559   48   78   10   3.30   CDWR   PALERHO   1976   1-26   1940: 0.0   39   435   12   1.559   48   78   10   3.30   CDWR   PALERHO   1976   1-26   1940: 0.7   39   428   12   1.559   48   78   10   3.30   CDWR   PALERHO   1976   1-26   1940: 0.7   39   428   12   1.559   48   78   10   3.30   CDWR   PALERHO   1976   1-27   1850: 2.3   39   405   12   1.867   53   66   0*   1.80   CDWR   PALERHO   1976   1-28   352: 19.5   39   419   12   1.545   50   81   4*   2.80   CDWR   LONA RICA   1976   1-28   352: 19.5   39   419   12   1.545   50   81   4*   2.80   CDWR   PALERHO   1976   1-28   352: 19.5   39   419   12   1.545   50   81   4*   2.80   CDWR   PALERHO   1976   1-28   352: 19.5   39   439   12   1.545   50   81   4*   2.80   CDWR   PALERHO   1976   1-28   2002: 54   39   439   12   1.557   50   00   4*   4*   4*   4*   4*   4*   4											
1976 1-26 755; 3,7 99,405 121,509 52 84 1* 1.00 CDWR PALERNO 1976 1-26 1159:34.0 39,565 122,091 22 35 8* 3.00 CDWR GLENN 1976 1-26 1940; 0.0 39,434 121,569 48 78 10 3.30 USGS 1976 1-26 21091; 0.7 39,426 121,501 52 83 5* 3.30 CDWR PALERNO 1976 1-26 210914,0 39,435 121,457 53 86 0* 1.80 CDWR BANGOR 1976 1-27 1850; 123 99,365 121,87 59 95 4*80 CDWR BANGOR 1976 1-28 352:19,5 39,365 121,545 50 81 4*80 CDWR BANGOR 1976 1-28 742:55,8 39,325 121,545 50 81 4* 2.80 CDWR BANGOR 1976 1-28 742:55,8 39,325 121,545 50 81 4* 2.80 CDWR PALERNO 1976 1-28 742:57,8 39,367 121,457 56 90 5*40 CDWR BANGOR 1976 1-28 2002:54,9 39,499 121,493 50 81 4*60 CDWR BANGOR 1976 1-28 2002:54,9 39,499 121,493 50 81 4*60 CDWR BANGOR 1976 1-29 315:25,3 39,434 121,537 50 80 1*60 CDWR BANGOR 1976 1-29 315:25,3 39,434 121,537 50 80 1*60 CDWR BANGOR 1976 1-29 315:25,3 39,434 121,537 50 80 1*60 CDWR BANGOR 1976 1-29 315:25,3 39,434 121,537 50 80 1*60 CDWR BANGOR 1976 1-29 315:25,3 39,434 121,537 50 80 1*60 CDWR BANGOR 1976 1-30 499:23,7 39,399 121,460 55 88 5*70 CDWR BANGOR 1976 1-30 305:27,1 39,508 121,460 55 88 5*70 CDWR BANGOR 1976 1-30 305:27,1 39,508 121,460 55 88 5*70 CDWR BANGOR 1976 1-30 600: 6.7 39,455 121,495 51 82 0*50 CDWR BANGOR 1976 2-1 18:31,8 39,438 121,463 53 86 1*10 CDWR BANGOR 1976 2-2 1 18:31,8 39,438 121,463 53 86 0*50 CDWR BANGOR 1976 2-1 836:20,7 39,401 121,475 54 87 0*50 CDWR BANGOR 1976 2-2 1 18:01,89 39,492 121,546 55 28 40 0*50 CDWR BANGOR 1976 2-2 1 18:01,89 39,492 121,546 55 28 80 0*50 CDWR BANGOR 1976 2-2 1 18:01,89 39,492 121,547 54 87 0*50 CDWR BANGOR 1976 2-3 18:01,89 39,492 121,548 52 88 0*50 CDWR BANGOR 1976 2-4 624;24,5 39,455 121,465 53 86 0*50 CDWR BANGOR 1976 2-9 139;13,1 39,459 121,460 53 86 0*50 CDWR BANGOR 1976 2-9 139;13,1 39,459 121,460 53 86 0*50 CDWR BANGOR 1976 2-9 139;13,1 39,459 121,460 53 86 0*50 CDWR BANGOR 1976 2-9 139;13,1 39,459 121,465 55 88 0*70 CDWR BANGOR 1976 2-1 18:44,54 59,455 121,546 55 89 40 0*50 CDWR BANGOR 197											
1976   1-26   2481   7.9   39.414   121.509   52   84   18   1.00   CDMR   PALERNO   1976   1-26   19401 0.0   39.436   121.569   46   78   10   3.30   USGS   1976   1-26   19401 0.7   39.428   121.501   52   83   58   3.30   CDMR   PALERNO   1976   1-26   2108114.0   39.435   121.457   53   86   09   1.80   CDMR   BANGOR   1976   1-27   18501   2.3   39.436   121.357   59   54  80   CDMR   BANGOR   1976   1-28   352:19.5   39.401   21.449   55   89   18  10   CDMR   BANGOR   1976   1-28   352:19.5   39.419   121.545   50   81   49   2.80   CDMR   PALERNO   1976   1-28   352:19.5   39.437   121.540   54   87   24*  20   CDMR   HONCUT   1976   1-28   241:35.1   39.431   121.595   50   81   49  20   CDMR   HONCUT   1976   1-28   241:35.1   39.412   121.529   51   82   2*   2.80   CDMR   BANGOR   1976   1-28   2002:54.9   39.493   121.493   50   81   49  60   CDMR   BANGOR   1976   1-28   2002:54.3   39.434   121.537   50   80   1*  60   CDMR   BANGOR   1976   1-29   315:25.3   39.434   121.537   50   80   1*  60   CDMR   BANGOR   1976   1-29   494:23.7   39.495   121.495   51   82   2*   2.80   CDMR   BANGOR   1976   1-30   325:27.1   39.508   121.595   51   82   2*   2.80   CDMR   BANGOR   1976   1-30   325:27.1   39.508   121.595   51   82   2*   2.80   CDMR   BANGOR   1976   1-30   325:27.1   39.508   121.595   51   82   2*   2.80   CDMR   BANGOR   1976   2-1   15:13.8   39.438   121.463   53   86   0*  20   CDMR   BANGOR   1976   2-1   15:13.8   39.438   121.463   53   86   0*  20   CDMR   BANGOR   1976   2-1   15:13.8   39.438   121.463   53   86   0*  20   CDMR   BANGOR   1976   2-1   16:13.8   39.438   121.463   53   86   0*  20   CDMR   BANGOR   1976   2-1   15:13.8   39.438   121.463   53   86   0*  20   CDMR   BANGOR   1976   2-1   16:13.8   39.438   121.463   53   86   0*  50   CDMR   BANGOR   1976   2-1   16:13.8   39.438   121.463   53   86   0*  50   CDMR   BANGOR   1976   2-1   16:13.8   39.438   121.465   53   86   0*  50   CDMR   BANGOR   1976   2-1											
1976   1-26   1159;34							_				
1976 1-26 1940: 0.0 39.434 121.569 48 78 10 3.30 USDS 1976 1-26 2108:114.0 39.435 121.457 53 86 0* 1.80 CDWR PALERNO 1976 1-27 1850: 2.3 39.365 121.387 59 95 4*80 CDWR PALERNO 1976 1-27 1850: 2.3 39.365 121.387 59 95 4*80 CDWR PALERNO 1976 1-28 352:19.5 39.419 121.545 50 81 4* 2.80 CDWR PALERNO 1976 1-28 352:19.5 39.419 121.545 50 81 4* 2.80 CDWR PALERNO 1976 1-28 742:157.8 39.367 121.457 56 90 5*40 CDWR HONCUT 1976 1-28 656: 2.6 39.583 121.705 37 60 0*40 CDWR HONCUT 1976 1-28 2002:154.9 39.439 121.493 50 81 4*60 CDWR PALERNO 1976 1-28 2041:35.1 39.412 121.557 50 80 1*60 CDWR PALERNO 1976 1-28 2341:35.1 39.412 121.557 50 80 1*60 CDWR PALERNO 1976 1-29 315:25.3 39.434 121.597 50 80 1*60 CDWR PALERNO 1976 1-29 315:25.3 39.434 121.537 50 80 1*60 CDWR PALERNO 1976 1-29 315:25.3 39.434 121.537 50 80 1*60 CDWR PALERNO 1976 1-29 315:25.3 39.434 121.537 50 80 1*60 CDWR PALERNO 1976 1-30 325:27.1 39.508 121.604 59 88 5*70 CDWR BANGOR 1976 1-30 325:27.1 39.508 121.604 49 79 9*50 CDWR PALERNO 1976 1-30 325:27.1 39.508 121.604 49 79 9*50 CDWR BANGOR 1976 2-1 15:31.8 39.438 121.463 53 86 0*40 CDWR BANGOR 1976 2-1 15:31.8 39.438 121.463 53 86 0*50 CDWR BANGOR 1976 2-1 18:36:20.7 39.401 121.475 54 87 0*50 CDWR BANGOR 1976 2-1 18:36:20.7 39.401 121.475 54 87 0*50 CDWR BANGOR 1976 2-1 18:36:20.7 39.508 121.464 52 83 8*50 CDWR BANGOR 1976 2-7 504:19.8 39.922 121.544 43 69 6* 2.60 CDWR BANGOR 1976 2-7 504:19.8 39.922 121.546 52 84 0*50 CDWR BANGOR 1976 2-9 139:13.1 39.459 121.467 51 82 3950 CDWR BANGOR 1976 2-9 139:13.1 39.459 121.467 51 87 4*80 CDWR BANGOR 1976 2-9 139:13.1 39.459 121.467 51 80 0*50 CDWR BANGOR 1976 2-9 139:13.1 39.459 121.467 51 80 0*50 CDWR BANGOR 1976 2-9 139:13.1 39.459 121.467 51 80 0*50 CDWR BANGOR 1976 2-9 139:13.1 39.459 121.467 51 80 0*50 CDWR BANGOR 1976 2-9 139:13.1 39.459 121.467 51 80 0*50 CDWR BANGOR 1976 2-9 139:13.1 39.459 121.467 51 80 0*50 CDWR BANGOR 1976 2-9 139:13.1 39.459 121.467 51 80 0*50 CDWR BANGOR 19											
1976   1-26   1940; 0, 7   39, 428   121,501   52   83   5*   3,30   CDWR   PALERNO   1976   1-27   1850; 2,3   39,365   121,387   59   95   4*  80   CDWR   BANGOR   1976   1-27   1850; 2,3   39,365   121,387   59   95   4*  80   CDWR   BANGOR   1976   1-28   352; 19,5   39,419   121,545   50   81   4*   2.80   CDWR   BANGOR   1976   1-28   352; 19,5   39,419   121,545   50   81   4*   2.80   CDWR   BANGOR   1976   1-28   742;55,8   39,325   121,540   54   67   24*  20   CDWR   HONCUT   1976   1-28   856; 2,6   39,583   121,705   37   60   0*  40   CDWR   SHIPPEE   1976   1-28   2002;54,9   39,493   121,593   50   81   4*  60   CDWR   BANGOR   1976   1-28   2002;54,9   39,493   121,593   51   82   2*   2.80   CDWR   BANGOR   1976   1-29   315;25,3   39,441   21,537   50   80   5*  60   CDWR   BANGOR   1976   1-29   315;25,3   39,434   121,537   50   80   5*  60   CDWR   BANGOR   1976   1-29   349;23,7   39,399   121,460   53   86   5*  70   CDWR   BANGOR   1976   1-29   325;27,1   39,508   212,504   49   79   9*  50   CDWR   BANGOR   1976   2-30   325;27,1   39,508   121,504   53   86   1*  10   CDWR   BANGOR   1976   2-1   186;207   39,401   21,475   54   87   0*  20   CDWR   BANGOR   1976   2-1   186;207   39,401   21,475   54   87   0*  50   CDWR   BANGOR   1976   2-1   186;207   39,401   21,475   54   87   0*  50   CDWR   BANGOR   1976   2-1   186;50,3   39,421   21,475   54   87   0*  50   CDWR   BANGOR   1976   2-2   2144;0   0.39,528   21,546   48   0*  50   CDWR   BANGOR   1976   2-7   1246;6,3   39,581   21,465   52   84   0*  50   CDWR   BANGOR   1976   2-7   504;19,8   39,522   21,548   42   68   14*   0.10   CDWR   BANGOR   1976   2-7   504;19,8   39,522   21,548   42   68   14*   0.10   CDWR   BANGOR   1976   2-9   139;13,1   39,498   121,475   54   87   4*  50   CDWR   BANGOR   1976   2-9   139;13,1   39,498   121,495   50   88   6*  50   CDWR   BANGOR   1976   2-9   139;13,1   39,491   21,465   53   86   6*  50   CDWR   BANGOR   1976   2-											
1976   1-26   2108:114.0   39.435   121.457   53   86   0*   1.80   CDWR   LOMA RICA   1976   1-27   2137:52.1   39.404   121.435   55   89   1*  10   CDWR   BANGOR   1976   1-28   352:19.5   39.419   121.545   50   81   4*   2.80   CDWR   PALERNO   1976   1-28   742:55.8   39.325   121.540   54   87   24*  20   CDWR   CDWR   1976   1-28   742:55.8   39.325   121.540   54   87   24*  20   CDWR   CDWR   1976   1-28   856: 2.6   39.583   121.705   37   60   0*  40   CDWR   CDWR   SANGOR   1976   1-28   2002:54.9   39.493   121.457   56   90   5*  40   CDWR   SANGOR   1976   1-28   2341:35.1   39.412   121.529   51   82   2*   2.80   CDWR   PALERNO   1976   1-29   315:25.3   39.434   121.537   50   80   1*  60   CDWR   PALERNO   1976   1-29   449:23.7   39.399   121.460   55   88   5*  70   CDWR   BANGOR   1976   1-30   3025:27.1   39.508   121.504   49   79   9*  10   CDWR   BANGOR   1976   1-30   600: 6.7   39.455   121.495   51   82   0*  20   CDWR   BANGOR   1976   2-1   15:31.8   39.438   121.463   53   86   0*  40   CDWR   BANGOR   1976   2-1   15:31.8   39.438   121.465   53   86   0*  40   CDWR   BANGOR   1976   2-1   18:31.8   39.438   121.465   53   86   0*  40   CDWR   BANGOR   1976   2-1   18:31.8   39.438   121.465   53   86   0*  50   CDWR   BANGOR   1976   2-1   18:31.8   39.438   121.465   52   84   0*  50   CDWR   BANGOR   1976   2-1   18:31.8   39.438   121.465   52   84   0*  50   CDWR   BANGOR   1976   2-1   18:31.3   39.493   121.475   54   87   0*  50   CDWR   BANGOR   1976   2-5   11:53.9   39.90   121.475   54   87   0*  50   CDWR   BANGOR   1976   2-7   1246:27.0   39.465   121.475   55   82   4*  50   CDWR   BANGOR   1976   2-7   1246:27.0   39.482   121.524   48   77   10*   1.10   CDWR   BANGOR   1976   2-7   1246:27.0   39.482   121.475   55   86   14*  50   CDWR   BANGOR   1976   2-9   1935:48.5   39.493   121.475   55   86   4*  50   CDWR   BANGOR   1976   2-9   1935:48.5   39.493   121.475   55   86   0*  30   CDWR   B						83	5#				PALERNO
1976   1-27   1850; 2.3   39.365   121.387   59   95   4*  80   CDWR   LONA RICA   1976   1-28   352:19.5   39.419   121.545   50   81   4*   2.80   CDWR   PALERNO   1976   1-28   352:19.5   39.419   121.545   50   81   4*   2.80   CDWR   PALERNO   1976   1-28   265: 2.6   39.583   121.540   54   87   24*  20   CDWR   HONCUT   1976   1-28   265: 2.6   39.583   121.705   37   60   0*  40   CDWR   SHIPPEE   1976   1-28   202: 54.9   39.493   121.493   50   81   4*  60   CDWR   SHIPPEE   1976   1-28   2341:351   39.412   121.529   51   82   2*   2.80   CDWR   PALERNO   1976   1-29   315:25.3   39.434   121.537   50   80   1*  60   CDWR   PALERNO   1976   1-29   325:27.1   39.508   121.504   49   79   9*  50   CDWR   BANGOR   1976   1-30   305:27.1   39.508   121.504   49   79   9*  50   CDWR   BANGOR   1976   1-30   600: 6.7   39.455   121.453   53   86   0*  40   CDWR   BANGOR   1976   2-1   18:31.8   39.438   121.463   53   86   0*  40   CDWR   BANGOR   1976   2-1   18:07:56.2   39.533   121.614   43   69   6*   2.60   CDWR   BANGOR   1976   2-1   18:07:56.2   39.533   121.614   43   69   6*   2.60   CDWR   BANGOR   1976   2-2   14:0.0   39.528   121.455   54   87   0*  50   CDWR   BANGOR   1976   2-5   811:53.9   39.404   121.472   54   87   4*  80   CDWR   BANGOR   1976   2-7   504:19.8   39.922   121.548   42   68   14*  80   CDWR   BANGOR   1976   2-7   19:13.1   39.459   121.450   53   85   0*  50   CDWR   BANGOR   1976   2-9   139:13.1   39.459   121.450   53   85   0*  50   CDWR   BANGOR   1976   2-9   139:13.1   39.459   121.450   53   85   0*  50   CDWR   BANGOR   1976   2-9   139:13.1   39.459   121.450   53   85   0*  50   CDWR   BANGOR   1976   2-9   139:13.1   39.459   121.450   53   85   0*  50   CDWR   BANGOR   1976   2-9   139:13.1   39.459   121.450   53   85   0*  50   CDWR   BANGOR   1976   2-9   139:13.1   39.459   121.450   53   85   0*  50   CDWR   BANGOR   1976   2-10   234:16.8   39.431   121.479   53   86   0*  30   CDWR   BAN						86	0#				BANGOR
1976   1-27 2137:52.1 39.404   121.449   55 89						95	4*		80	CDWR	LOMA RICA
1976 1-28 742:55.8 39.325 121.540 54 87 24*20 CDWR HONCUT 1976 1-28 456: 2.6 39.538 121.705 37 60 0*40 CDWR SHIPPEE 1976 1-28 2341:351. 39.493 121.493 50 81 4*60 CDWR SHIPPEE 1976 1-28 2341:351. 39.412 121.529 51 82 2* 2.80 CDWR PALERHO 1976 1-29 315:25.3 39.434 121.537 50 80 1*60 CDWR PALERHO 1976 1-29 315:25.3 39.434 121.537 50 80 1*60 CDWR PALERHO 1976 1-29 349:23.7 39.399 121.460 55 88 5*70 CDWR BANGOR 1976 1-30 325:27.1 39.508 121.504 49 79 9*50 CDWR BANGOR 1976 1-30 325:27.1 39.508 121.504 49 79 9*50 CDWR BANGOR 1976 2-1 15:31.8 39.498 121.495 51 82 0*20 CDWR BANGOR 1976 2-1 1807:56.2 39.533 121.614 43 69 6* 2.60 CDWR BANGOR 1976 2-1 1807:56.2 39.533 121.614 43 69 6* 2.60 CDWR OROVILLE 1976 2-2 2144: 0.0 39.528 121.444 52 83 8*50 CDWR OROVILLE 1976 2-5 811:53.9 39.404 121.472 54 87 4*80 CDWR BANGOR 1976 2-5 811:53.9 39.404 121.472 54 87 4*80 CDWR BANGOR 1976 2-5 81:53.9 39.404 121.472 54 87 7* 0*50 CDWR BANGOR 1976 2-5 81:53.9 39.404 121.472 54 87 7* 0*50 CDWR BANGOR 1976 2-7 504:19.8 39.522 121.544 51 82 3*50 CDWR BANGOR 1976 2-7 504:19.8 39.522 121.544 51 82 3*50 CDWR BANGOR 1976 2-9 133:13.1 39.498 121.524 48 77 10* 1.10 CDWR BANGOR 1976 2-9 139:13.1 39.498 121.524 48 77 10* 1.10 CDWR BANGOR 1976 2-9 139:13.1 39.459 121.460 53 85 0*50 CDWR BANGOR 1976 2-9 133:13.1 39.459 121.551 49 79 0* 2.40 CDWR BANGOR 1976 2-9 133:13.1 39.459 121.460 53 85 0*50 CDWR BANGOR 1976 2-9 133:13.1 39.459 121.460 53 85 0*50 CDWR BANGOR 1976 2-9 133:13.1 39.459 121.460 53 85 0*50 CDWR BANGOR 1976 2-9 133:13.1 39.459 121.460 53 85 0*50 CDWR BANGOR 1976 2-9 133:13.1 39.459 121.450 50 80 5* 2.00 CDWR BANGOR 1976 2-9 133:13.1 39.459 121.450 50 80 5* 2.00 CDWR BANGOR 1976 2-9 133:13.1 39.459 121.450 50 80 5* 2.00 CDWR BANGOR 1976 2-10 426:46.3 39.431 121.479 53 86 0*30 CDWR BANGOR 1976 2-10 426:46.3 39.431 121.479 53 86 0*30 CDWR BANGOR 1976 2-10 426:46.3 39.431 121.479 53 86 0*30 CDWR BANGOR 1976 2-10 1216:44.7 39.458 121.555 58 80 0*30 CDWR BANGOR 1976 2					55	89	1*		10	CDWR	BANGOR
1976   1-28   742:157.8   39.367   121.457   56   90   5*  40   CDWR   LOMA RICA   1976   1-28   2002:54.9   39.493   121.493   50   81   4*  60   CDWR   BANGOR   1976   1-29   315:25.3   39.434   121.529   51   82   2*   2.80   CDWR   PALERHO   1976   1-29   315:25.3   39.434   121.527   50   80   1*  60   CDWR   BANGOR   1976   1-29   122:310.0   39.431   121.470   53   86   1*  10   CDWR   BANGOR   1976   1-30   325:27.1   39.508   121.504   49   79   9*  50   CDWR   BANGOR   1976   1-30   325:27.1   39.508   121.504   49   79   9*  50   CDWR   BANGOR   1976   1-30   360:67.7   39.451   121.475   54   87   0*  50   CDWR   BANGOR   1976   2-1   15:31.8   39.438   121.455   53   86   0*  40   CDWR   BANGOR   1976   2-1   18:36:20.7   39.401   121.475   54   87   0*  50   CDWR   BANGOR   1976   2-2   21:44: 0.3   9.528   121.464   52   83   8*  50   CDWR   BANGOR   1976   2-2   21:44: 0.3   9.528   121.465   52   84   0*  50   CDWR   BANGOR   1976   2-5   81:153.9   39.404   121.472   54   87   4*  80   CDWR   BANGOR   1976   2-5   81:153.9   39.404   121.472   54   87   4*  80   CDWR   BANGOR   1976   2-7   1246:27.0   39.492   121.548   42   68   14*   0.10   CDWR   BANGOR   1976   2-7   1246:27.0   39.492   121.501   50   81   0*  70   CDWR   BANGOR   1976   2-9   139:13.1   39.495   121.515   49   79   0*   2.40   CDWR   BANGOR   1976   2-9   136:47.7   39.486   121.501   50   81   0*  70   CDWR   BANGOR   1976   2-9   134:139.0   39.490   121.459   52   83   0*  50   CDWR   BANGOR   1976   2-9   134:139.0   39.490   121.515   49   79   0*   2.40   CDWR   BANGOR   1976   2-9   134:139.0   39.490   121.459   52   83   0*  50   CDWR   BANGOR   1976   2-9   134:139.0   39.490   121.459   52   83   0*  50   CDWR   BANGOR   1976   2-10   126:46.3   39.413   121.479   53   86   0*  30   CDWR   BANGOR   1976   2-10   426:46.3   39.413   121.479   53   86   0*  30   CDWR   BANGOR   1976   2-10   426:46.3   39.413   121.479   53   86   0*  30   CDWR	1976 1-2	8 352:19.5	39.419	121.545	50	81	4#		2.80	CDWR	PALERNO
1976   1-28   256:   2.6   39.583   121.705   37   60   0*   -40   CDWR   SHIPPEE   1976   1-28   204:   39.493   121.493   50   81   4*  60   CDWR   BANGOR   1976   1-29   315:   25.3   39.494   121.529   51   82   2*   2.80   CDWR   PALERNO   1976   1-29   1223:   10.0   39.491   121.527   50   80   1*  60   CDWR   CDWR   PALERNO   1976   1-29   1223:   10.0   39.491   121.470   53   86   1*  10   CDWR   BANGOR   1976   1-30   325:   27.1   39.508   121.450   49   79   9*  50   CDWR   CROVILLE   1976   1-30   600:   67.   39.455   121.495   51   82   0*  20   CDWR   BANGOR   1976   2-1   15:31.8   39.438   121.463   53   86   0*  40   CDWR   BANGOR   1976   2-1   1807:   56.2   39.533   121.614   43   69   6*   2.60   CDWR   OROVILLE   1976   2-2   2144:   0.0   39.528   121.445   52   83   8*  50   CDWR   CROVILLE   1976   2-5   1257:   46.6   39.528   121.455   54   87   4*  50   CDWR   BANGOR   1976   2-5   1257:   46.6   39.528   121.455   54   87   4*  80   CDWR   BANGOR   1976   2-5   1257:   46.6   39.528   121.455   54   87   4*  80   CDWR   BANGOR   1976   2-7   504:   19.498   121.478   51   82   3*  50   CDWR   BANGOR   1976   2-7   1246:   27.0   39.482   121.501   50   81   6*  50   CDWR   BANGOR   1976   2-9   139:   13.1   39.498   121.450   53   85   0*  50   CDWR   BANGOR   1976   2-9   139:   13.1   39.498   121.450   53   85   0*  50   CDWR   BANGOR   1976   2-9   139:   13.1   39.459   121.450   53   85   0*  50   CDWR   BANGOR   1976   2-9   139:   13.1   39.459   121.450   53   85   0*  50   CDWR   BANGOR   1976   2-9   139:   13.1   39.459   121.450   53   85   0*  50   CDWR   BANGOR   1976   2-9   139:   13.1   39.459   121.450   53   85   0*  50   CDWR   BANGOR   1976   2-9   134:   13.0	1976 1-2	8 742:55.8	39.325	121.540	54	87	24*		20	CDWR	HONCUT
1976 1-28 2002:54.9 39.493 121.493 50 81 4*60 CDWR BANGOR 1976 1-28 2341:35.1 39.412 121.529 51 82 2* 2.80 CDWR PALERNO 1976 1-29 315:25.3 39.434 121.537 50 80 1*60 CDWR PALERNO 1976 1-29 1223:10.0 39.431 121.470 53 86 1*10 CDWR BANGOR 1976 1-30 325:27.1 39.508 121.504 49 79 9*50 CDWR OROVILLE 1976 1-30 600: 6.7 39.455 121.455 51 82 0*20 CDWR BANGOR 1976 2-1 15:31.8 39.438 121.463 53 86 0*40 CDWR BANGOR 1976 2-1 15:31.8 39.438 121.463 53 86 0*40 CDWR BANGOR 1976 2-1 1836:20.7 39.401 121.475 54 87 0*50 CDWR BANGOR 1976 2-1 1807:56.2 39.533 121.614 43 69 6* 2.60 CDWR OROVILLE 1976 2-4 624:24.5 39.465 121.465 52 84 0*50 CDWR BANGOR 1976 2-5 1257:46.6 39.528 121.444 52 83 8*50 CDWR BANGOR 1976 2-5 1257:46.6 39.528 121.444 52 83 8*50 CDWR BANGOR 1976 2-5 1257:46.6 39.528 121.472 54 87 4*80 CDWR BANGOR 1976 2-5 1257:46.6 39.528 121.472 54 87 4*80 CDWR BANGOR 1976 2-5 1257:46.6 39.528 121.524 48 77 10* 1.10 CDWR BANGOR 1976 2-7 504:19.8 39.922 121.548 42 68 14* 0.10 CDWR BANGOR 1976 2-7 1246:27.0 39.482 121.501 50 81 0*70 CDWR BANGOR 1976 2-9 139:13.1 39.469 121.450 53 85 0*50 CDWR BANGOR 1976 2-9 139:13.1 39.469 121.450 53 85 0*50 CDWR BANGOR 1976 2-9 139:13.1 39.469 121.501 50 81 0*70 CDWR BANGOR 1976 2-9 139:13.1 39.469 121.501 50 80 5* 2.00 CDWR BANGOR 1976 2-9 139:13.1 39.469 121.515 49 79 0* 2.40 CDWR BANGOR 1976 2-9 139:13.1 39.469 121.501 50 80 5* 2.00 CDWR BANGOR 1976 2-9 133:13.1 39.459 121.460 53 85 0*50 CDWR BANGOR 1976 2-9 132:148.5 39.501 121.460 53 85 0*50 CDWR BANGOR 1976 2-9 132:148.5 39.501 121.460 53 85 0*50 CDWR BANGOR 1976 2-1 124:41.8 5 89.41 121.472 53 86 0*30 CDWR BANGOR 1976 2-1 124:41.8 5 89.41 121.472 53 86 0*30 CDWR BANGOR 1976 2-10 24:14.6 13.9 4.11 121.472 53 86 0*30 CDWR BANGOR 1976 2-10 24:14.6 39.413 121.460 53 85 0*10 CDWR BANGOR 1976 2-10 12:14:40.7 39.431 121.479 53 86 0*30 CDWR BANGOR 1976 2-10 12:14:41.7 39.468 121.505 50 80 6* 0.70 CDWR BANGOR 1976 2-12 12:26:16.7 39.431 121.467 54 87 8* 8* 8* 8* 8* 8* 8*	1976 1-2	8 742:57.8	39.367	121.457	56	90	5*		40	CDWR	LOMA RICA
1976   1-28   2341:35.1   39.412   121.529   51   82   2*   2.80   CDWR   PALERNO   1976   1-29   449:23.7   39.399   121.460   55   88   5*  70   CDWR   BANGOR   1976   1-29   1223:10.0   39.431   121.470   53   86   1*  10   CDWR   BANGOR   1976   1-30   325:27.1   39.508   121.504   49   79   9*  50   CDWR   CDWR   CDWR   1976   1-30   600: 6.7   39.455   121.495   51   82   0*  20   CDWR   BANGOR   1976   2-1   15:31.8   39.438   121.463   53   86   0*  40   CDWR   BANGOR   1976   2-1   836:20.7   39.401   121.475   54   87   0*  50   CDWR   BANGOR   1976   2-1   1807:56.2   39.533   121.614   43   69   6*   2.60   CDWR   CDWR   CDWR   1976   2-2   2144: 0.0   39.528   121.444   52   83   8*  50   CDWR   BANGOR   1976   2-5   811:53.9   39.404   121.475   54   87   4*  80   CDWR   BANGOR   1976   2-5   811:53.9   39.404   121.475   54   87   4*  80   CDWR   BANGOR   1976   2-6   945:18.1   39.498   121.478   51   82   3*  50   CDWR   BANGOR   1976   2-7   504:19.8   39.922   121.548   42   68   14*   0.10   CDWR   BANGOR   1976   2-9   139:13.1   39.495   121.460   53   85   0*  50   CDWR   BANGOR   1976   2-9   139:13.1   39.496   121.515   49   79   0*   2.40   CDWR   BANGOR   1976   2-9   133:15.2   39.513   121.621   43   70   1*   3.00   CDWR   BANGOR   1976   2-9   133:15.2   39.438   121.501   50   81   0*  70   CDWR   BANGOR   1976   2-9   134:19.0   39.496   121.515   49   79   0*   2.40   CDWR   BANGOR   1976   2-9   134:19.0   39.496   121.479   53   86   0*  50   CDWR   BANGOR   1976   2-9   134:19.0   39.496   121.515   49   79   0*   2.40   CDWR   BANGOR   1976   2-9   134:19.0   39.491   121.479   53   86   0*  30   CDWR   BANGOR   1976   2-10   234:16.8   39.413   121.479   53   86   0*  30   CDWR   BANGOR   1976   2-10   234:16.8   39.413   121.479   53   86   0*  30   CDWR   BANGOR   1976   2-10   234:16.8   39.413   121.479   53   86   0*  80   CDWR   BANGOR   1976   2-10   234:16.3   39.413   121.479   53   86   0*  80   CDWR   BANG		8 856: 2.6	39.583	121.705	37	60	0#				
1976   1-29   315;25,3   39,434   121,537   50   80   1*  60   CDWR   PALERNO   1976   1-29   1223;10.0   39,431   121,470   53   86   1*  10   CDWR   BANGOR   1976   1-30   325;27.1   39,508   121,504   49   79   9*  50   CDWR   CDWR   BANGOR   1976   1-30   600; 6.7   39,455   121,465   51   82   0*  20   CDWR   BANGOR   1976   2-1   15;31.8   39,438   121,463   53   86   0*  40   CDWR   BANGOR   1976   2-1   1807;56.2   39,533   121,614   43   69   6*   2.60   CDWR   CDWR   CDWR   1976   2-2   2144; 0.0   39,528   121,444   52   83   8*  50   CDWR   CROVILLE   DAN   1976   2-2   2144; 0.0   39,528   121,445   52   84   0*  50   CDWR   BANGOR   1976   2-5   811;53.9   39,404   121,472   54   87   4*  80   CDWR   BANGOR   1976   2-5   1257;46.6   39,528   121,455   52   84   0*  50   CDWR   BANGOR   1976   2-5   1257;46.6   39,528   121,472   54   87   4*  80   CDWR   BANGOR   1976   2-7   504;19.8   39,528   121,478   51   82   3*  50   CDWR   BANGOR   1976   2-7   1246;27.0   39,482   121,501   50   81   0*  70   CDWR   BANGOR   1976   2-9   139;13.1   39,459   121,460   53   85   0*  50   CDWR   BANGOR   1976   2-9   139;13.1   39,459   121,460   53   85   0*  50   CDWR   BANGOR   1976   2-9   139;13.1   39,459   121,460   53   85   0*  50   CDWR   BANGOR   1976   2-9   139;13.1   39,459   121,460   53   85   0*  50   CDWR   BANGOR   1976   2-9   1357;48.5   39,496   121,515   49   79   0*   2.40   CDWR   BANGOR   1976   2-9   1357;48.5   39,496   121,501   50   81   0*  50   CDWR   BANGOR   1976   2-9   1357;48.5   39,496   121,498   52   83   0*  50   CDWR   BANGOR   1976   2-10   2062; 6.1   39,491   121,498   52   83   0*  30   CDWR   BANGOR   1976   2-10   2062; 6.1   39,491   121,498   53   86   0*  30   CDWR   BANGOR   1976   2-10   2062; 6.1   39,491   121,498   53   86   0*  30   CDWR   BANGOR   1976   2-14   438;40.0   39,491   121,495   53   86   0*  30   CDWR   BANGOR   1976   2-14   438;40.0   39,491   121,495   53   86   0*											BANGOR
1976   1-29   449:23.7   39.399   121.460   55   88   5*  70   CDWR   BANGOR   1976   1-29   1223:10.0   39.431   121.470   53   86   1*  10   CDWR   BANGOR   1976   1-30   325:27.1   39.508   121.504   49   79   9*  50   CDWR   BANGOR   1976   1-30   600: 6.7   39.455   121.495   51   82   0*  20   CDWR   BANGOR   1976   2-1   15:31.8   39.438   121.463   53   86   0*  40   CDWR   BANGOR   1976   2-1   1807:56.2   39.533   121.614   43   69   6*  50   CDWR   CROVILLE   1976   2-2   2144: 0.0   39.528   121.444   52   83   8*  50   CDWR   CROVILLE   DAN   1976   2-4   624:24.5   39.455   121.455   52   84   0*  50   CDWR   CROVILLE   DAN   1976   2-5   1257:46.6   39.528   121.524   48   77   10*   1.10   CDWR   CROVILLE   1976   2-5   1257:46.6   39.528   121.524   48   77   10*   1.10   CDWR   CROVILLE   1976   2-7   504:19.8   39.922   121.548   42   68   14*   0.10   CDWR   CROVILLE   1976   2-7   504:19.8   39.922   121.548   42   68   14*   0.10   CDWR   CROVILLE   1976   2-7   139:13.1   39.459   121.460   53   85   0*  50   CDWR   BANGOR   1976   2-9   139:13.1   39.459   121.460   53   85   0*  50   CDWR   BANGOR   1976   2-9   133:5.2   39.533   121.621   43   70   1*   3.00   CDWR   BANGOR   1976   2-9   1342:39.0   39.496   121.515   49   79   0*   2.40   CDWR   BANGOR   1976   2-9   1342:39.0   39.490   121.469   52   83   0*  50   CDWR   BANGOR   1976   2-9   1342:39.0   39.490   121.469   52   83   0*  30   CDWR   BANGOR   1976   2-10   2052: 6.1   39.431   121.479   53   86   0*  30   CDWR   BANGOR   1976   2-10   2052: 6.1   39.431   121.479   53   86   0*  30   CDWR   BANGOR   1976   2-10   2052: 6.1   39.431   121.467   54   87   88   88  50   CDWR   BANGOR   1976   2-14   818:40.0   39.521   121.515   58   89   4*  40   CDWR   BANGOR   1976   2-16   1225:36.7   39.408   121.457   55   88   0*  30   CDWR   BANGOR   1976   2-16   1225:36.7   39.408   121.457   55   86   0*  50   CDWR   BANGOR   1976   2-16   1225:36.7   39.408   121.457							_				
1976 1-29 1223:10.0 39.431 121.470 53 86 1* -10 CDWR BANGOR 1976 1-30 325:27.1 39.508 121.504 49 79 9*50 CDWR RORUYLLLE 1976 1-30 600: 6.7 39.455 121.495 51 82 0*20 CDWR BANGOR 1976 2-1 15:31.8 39.438 121.463 53 86 0*40 CDWR BANGOR 1976 2-1 836:20.7 39.401 121.475 54 87 0*50 CDWR BANGOR 1976 2-1 1807:56.2 39.533 121.614 43 69 6* 2.60 CDWR OROYLLLE 1976 2- 2 2144: 0.0 39.528 121.444 52 83 8*50 CDWR OROYLLLE DAN 1976 2- 4 624:24.5 39.465 121.465 52 84 0*50 CDWR BANGOR 1976 2- 5 811:53.9 39.404 121.472 54 87 4*80 CDWR BANGOR 1976 2- 5 1257:46.6 39.528 121.455 52 84 0*50 CDWR BANGOR 1976 2- 5 1257:46.6 39.528 121.524 48 77 10* 1.10 CDWR BANGOR 1976 2- 7 504:19.8 39.922 121.548 42 68 14* 0.10 CDWR PARDISE NE 1976 2- 7 1246:27.0 39.482 121.501 50 81 0*50 CDWR BANGOR 1976 2- 9 139:13.1 39.459 121.460 53 85 0*50 CDWR BANGOR 1976 2- 9 133:3 5.2 39.513 121.621 43 70 1* 3.00 CDWR BANGOR 1976 2- 9 133:3 5.2 39.513 121.621 43 70 1* 3.00 CDWR BANGOR 1976 2- 9 133:3 5.2 39.513 121.621 43 70 1* 3.00 CDWR BANGOR 1976 2- 9 1357:48.5 39.506 121.515 49 79 0* 2.40 CDWR BANGOR 1976 2- 9 1357:48.5 39.506 121.614 44 71 1* 2.30 CDWR BANGOR 1976 2- 9 1357:48.5 39.506 121.614 44 71 1* 2.30 CDWR BANGOR 1976 2-10 2426:46.3 39.413 121.479 53 86 0*30 CDWR BANGOR 1976 2-10 426:46.3 39.413 121.479 53 86 0*30 CDWR BANGOR 1976 2-10 426:46.3 39.413 121.479 53 86 0*30 CDWR BANGOR 1976 2-10 426:46.3 39.413 121.479 53 86 0*30 CDWR BANGOR 1976 2-10 426:46.3 39.413 121.479 53 86 0*30 CDWR BANGOR 1976 2-10 426:46.3 39.413 121.479 53 86 0*30 CDWR BANGOR 1976 2-11 4818:40.0 39.521 121.515 48 7880 CDWR BANGOR 1976 2-12 1226:16.7 39.431 121.485 53 86 0*30 CDWR BANGOR 1976 2-14 818:40.0 39.521 121.515 54 87 89 88 9820 CDWR BANGOR 1976 2-16 115:26.5 39.413 121.479 53 86 0*30 CDWR BANGOR 1976 2-16 115:26.5 39.413 121.479 53 86 0*50 CDWR BANGOR 1976 2-16 115:26.5 39.413 121.479 50 86 0*50 CDWR BANGOR 1976 2-16 1226:36.7 39.431 121.457 54 87 89 8920 CDWR BANGOR 1976 2-16 115:26.5 39.413 121.45											
1976   1-30   325:27.1   39.508   121.504   49   79   9*  50   CDWR   OROVILLE   1976   1-30   600: 6.7   39.455   121.495   51   82   0*  20   CDWR   BANGOR   1976   2-1   15:31.8   39.438   121.463   53   86   0*  40   CDWR   BANGOR   1976   2-1   1807:56.2   39.533   121.614   43   69   6*   2.60   CDWR   OROVILLE   1976   2-2   2144: 0.0   39.528   121.444   52   83   8*  50   CDWR   OROVILLE   DAN   1976   2-4   624:24.5   39.465   121.465   52   84   0*  50   CDWR   OROVILLE   DAN   1976   2-5   811:53.9   39.404   121.472   54   87   4*  80   CDWR   BANGOR   1976   2-5   811:53.9   39.404   121.472   54   87   4*  80   CDWR   OROVILLE   1976   2-6   945:18.1   39.498   121.478   51   82   3*  50   CDWR   BANGOR   1976   2-7   504:19.8   39.922   121.548   42   68   14*   0.10   CDWR   DAR   DAR											
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DAT		TIME		LON.				MMI	CRM CD	REF.	COMMENTS
		HRMN: SECS			MI	KM	KN				
+											
1976	2-25	647:56.7	39.465	121.457	53	85	5#		70	CDWR	BANGOR
1976	2-25	1803:38.5	39.485	121.453	52	84	0#		80	CDWR	BANGOR
1976	2-25	2340: 1.8	39.415	121.478			1*		1.00	CDWR	BANGOR
1976	2-26	741: 5.0	39.445	121.525			7*		50	CDWR	PALERHO
		1645: 0.7					8*		90	CDWR	OROVILLE
		901: 9.2					4*		50	CDWR	BANGOR
		18: 4.6					10*			CDWR	GLENN
		202:48.5					5*		90	CDWR	BANGOR
		533:33.2					4#		70	CDWR	PALERMO
		1053:14.8							80 80	CDWR CDWR	OROVILLE PALERNO
		159:27.6				_	8# 3#		50 50	CDWR	BANGOR
		615:38.6 810: 7.6					5#		20	CDWR	PALERNO
		952:23.9					0#		50	CDWR	BANGOR
		1535:51.2					1#		1.00	CDWR	PALERNO
		850:33.0					3*		50	CDWR	BANGOR
		2116:32.7					_		10	CDWR	PARADISE SW
1976		8: 2.5							0.30	CDWR	PULGA NW
		1207:47.7							70	CDWR	BANGOR
		538:26.7							80	CDWR	BANGOR
		549:21.5							60	CDWR	PALERNO
		7:44.8				82	6*		80	CDWR	PALERNO
		8:33.4				78	7*		80	CDWR	OROVILLE
		908: 5.6				79	8*		60	CDWR	OROVILLE
		2338: 9.8				79	5*		1.00	CDWR	PALERNO
		605:19.5				87	3*		0.60	CDWR	BANGOR
		652:31.9				_			70	CDWR	BANGOR
		714:19.6							2.70	CDWR	BANGOR
		716:32.2							0.40	CDWR	BANGOR
		743:22.4							0.20	CDWR	BANGOR
	3-15					_			0.40	CDWR	BANGOR
		1045:37.0				-			0.60	CDWR	PALERMO
		243:34.8							30	CDWR	BANGOR
		323:59.9							0.80	CDWR	RED BLUFF WEST LOS NOLINOS
		350:34.1							1.20	CDWR CDWR	BANGOR
		414:16.2 1303:25.1							0.00		
									80	CDWR	OROVILLE
1976 1976	3-17	1625: 8.5				_			70	CDWR	OROVILLE
1976	3-19								2.60	CDWR	OROVILLE
1976		1209:37.9					_		60	CDWR	BANGOR
1976		1432:26.2							0.00	CDWR	JONESVILLE SW
1976		843: 0.5							70	CDWR	BANGOR
1976		1157:13.7							80	CDWR	BANGOR
1976		1209:42.2							70	CDWR	BANGOR
1976	3-20	2134:57.4	39.400	121.489	<b>9</b> 53		0#		30	CDWR	BANGOR
1976		2314:39.4	39.402	121.412	2 57		5#		0.10	CDWR	BANGOR
1976	3-21						5#		1.00	CDWR	OROVILLE DAM
1976		1403:26.1							1.00	CDWR	PALERNO
1976	3-22								50	CDWR	PALERNO
1976		2 1213:43.9							10	CDWR	PRINCETON
1976		2048:28.							80	CDWR	PALERNO
1976	3-25			122.03					0.60	CDWR	PRINCETON
1976	3-26	5 230:29.	2 39.396	5 121.47	0 5	5 8	3 1*		60	CDWR	BANGOR

		TIME HRMN:SECS	LAT.	LON.	DIST NI	DIST KM	DPTH KM	HHI CRH CD	REF.	CONHENTS
+										
1976	2 27	1551.40.7	20 500	101 476	E 1	42	0*	2.70	CDWR	OPOUTLIE DAN
1976		1551:42.7 815:31.6			51 49	82 79	12*	30	CDWR	OROVILLE DAM OROVILLE
1976		842: 5.4			48	75 78	8*	0.80	CDWR	OROVILLE
1976		1315:50.6			53	86	2*	70	CDWR	BANGOR
1976		1419:30.2			52	84	1*	60	CDWR	PALERNO
1976		1744:29.4			52	84	2*	50	CDWR	BANGOR
1976		1903: 9.8			53	86	5#	0.80	CDWR	BANGOR
1976		2058:32.3			48	78	8#	90	CDWR	OROVILLE
1976		2303: 9.3			52	84	6#	40	CDWR	PALERNO
1976	4- 1	2307:56.2	39.599	121.439	50	81	0#	50	CDWR	OROVILLE DAM
1976		1804:31.8			53	86	0#	60	CDWR	BANGOR
1976		1329:58.0			39	63	4#	2.80	CDWR	PARADISE SE
1976		944:23.0			54	87	8*	40	CDWR	OROVILLE DAM
1976		1215:49.9				34	19*	0.80	CDWR	GLENN
1976 1976		743: 5.3 1035:52.2			56 48	90 77	0# 10#	80 70	CDWR CDWR	BANGOR OROVILLE
1976		1045:18.1			54	87	8*	50	CDWR	OROVILLE DAM
1976		1757:20.4			50	81	3*	50	CDWR	OROVILLE DAM
1976		1900:21.4			52	83	5*	40	CDWR	PALERNO
1976		2115:21.2			52	84	1*	20	CDWR	PALERNO
1976	4-14				50	81	1*	80	CDWR	PALERNO
1976	4-14	833:39.5			27	43	8*	40	CDWR	GLENN
1976	4-16	625:29.7	39.419	121.476	53	86	0*	70	CDWR	BANGOR
1976	4-16	1305:59.5	39.428	121.492	52	84	4#	50	CDWR	BANGOR
1976	4-16	1711:47.9	39.632	121.711	35	57	0#	3.00	CDWR	HANLIN CANYON
1976		1721:36.9			51	82	0#	2.50	CDWR	BANGOR
1976		825:48.1			55	88	1#	1.00	CDWR	BANGOR
1976		903:38.2			48	77	0#	60	CDWR	OROVILLE
1976		1349:51.0			53	86	2*	0.20	CDWR	JONESVILLE NW
1976		1837:43.7			52	83	0#	60	CDWR	PALERNO
1976		2240:45.1			40	65	5#	1.10	CDWR	GILMORE PEAK
1976		2131:28.6			47	76	5*	0.80	CDWR	LASSEN PEAK SW
1976	4-22	706: 7.7			<b>48</b> 53	78 <b>8</b> 5	6* 5*	60 2.40	CDWR CDWR	OROVILLE BANGOR
1976 1976		1335:38.7 251:46.3			48	77	10*	40	CDWR	OROVILLE
1976		1135: 4.6			47	75	7#	0.30	CDWR	BUTTE MEADOWS SE
1976	4-25	204:43.1			53	86	4#	0.10	CDWR	BANGOR
1976	4-26	349:25.3			53	86	5*	0.30	CDWR	BANGOR
1976		2007:48.9			49	79	3*	10	CDWR	PALERNO
1976	4-27	229:12.0	39.402	121.460	55	88	3#	20	CDWR	BANGOR
1976	4-27	925:15.9			53	86	0#	30	CDWR	BANGOR
1976		2133:54.3			49	79	5*	2.10	CDWR	OROVILLE DAM
1976	4-30			121.479	50	80	5*	30	CDWR	OROVILLE DAM
1976	4-30	705:50.8			53	86	3*	40	CDWR	BANGOR
1976		2051:39.2			53	85	0*	80	CDWR	BANGOR
1976	5- 1	808:53.0			53	85	4#	60	CDWR	BANGOR
1976	5- 2			121.505	50	80	1# 0#	50 80	CDWR	PALERNO
1976 1976		1902:45.9 1216:15.4			52 67	83 108	11*	90 0.30	CDWR CDWR	PALERMO Browns valley
1976	5- 4	623:16.4			49	79	0#	40	CDWR	PALERNO
1976	-	1153: 2.0			52	84	1*	50	CDWR	BANGOR
1976	5- 7	741:43.8			53	86	0#	2.40	CDWR	BANGOR
1976		1305:49.6			51	82	0#	60	CDWR	BANGOR
1976		1348: 3.2			53	85	2*	80	CDWR	BANGOR

DATE	TIME	LAT.	LON.	DIST	DIST	DPTH	NNI CRN CD	REF.	CONNENTS
YEAR-NO-DY	HRMN:SECS			MI	KM	KM			
+									***
1976 5-11	205:32.2	39.507	121.480	50	81	0#	90	CDWR	OROVILLE DAM
	1120:43.6			52	84	7 <b>*</b>	30	CDWR	BANGOR
	2138:21.5			54	87	2*	1.00	CDWR	BANGOR
	1429:15.3			55	88	1#	30	CDWR	BANGOR
	2333:42.2 1536: 7.5			50 53	80 86	0# 3#	80 80	CDWR CDWR	OROVILLE DAM BANGOR
	1713: 8.0			52	83	2*	40	CDWR	BANGOR
	946:46.8			52	84	0#	1.60	CDWR	BANGOR
	2333:35.0			57	91	5#	0.30	CDWR	SUTTER
	110: 3.1			50	81	3*	40	CDWR	OROVILLE DAM
	1115:43.6			52	83	1*	90	CDWR	PALERNO
1976 5-19	1127:23.9	39.487	121.488	50	81	2*	1.00	CDWR	BANGOR
	1320: 2.7	39.482	121.502	50	80	0*	80	CDWR	PALERNO
	727:34.5			50	81	0#	1.30	CDWR	BANGOR
	2321:24.0			32	52	1#	10	CDWR	HAMLIN CANYON
	2116:32.2			53	85	3*	50	CDWR	BANGOR
1976 5-23			121.510	50	81	5*	50	CDWR	PALERHO
	355:25.4			48	78	5*	2.40	CDWR	PALERMO
	1124:29.7 1355:38.8			53 53	86 86	5* 6*	90 60	CDWR CDWR	BANGOR BANGOR
	514:29.9				35	5*	1.60	CDWR	HALL RIDGE
1976 5-30			121.485		86	2*	30	CDWR	BANGOR
	1826:18.0				87	7#	30	CDWR	BANGOR
	2111:31.3				88	0#	60	CDWR	BANGOR
1976 5-31					78	0#	2.50	CDWR	OROVILLE
1976 5-31				54	87	4+	70	CDWR	BANGOR
1976 5-31				53	85	0#	70	CDWR	BANGOR
1976 5-31	831:43.3	39.512	121.522	48	77	4*	2.20	CDWR	OROVILLE
1976 5-31	842:18.5	39.519	121.540	47	76	7#	90	CDWR	OROVILLE
1976 5-31	905:38.1	39.609	121.682	37	60	1#	90	CDWR	SHIPPEE
	1648:58.3				81	8#	1.00	CDWR	PALERNO
1976 6- 4					87	0#	50	CDWR	BANGOR
	956:23.4				80	_	50	CDWR	PALERNO
	1152:34.8				69		70	CDWR	SHIPPEE
	2232:59.7				-	_	0.30	CDWR	NORD
	628:41.7 1239: 7.0				26 79	6 <b>*</b> 7 <b>*</b>	1.30 60	CDWR CDWR	HAMILTON CITY OROVILLE
	1214:18.7				85		0.40	CDWR	BANGOR
1976 6- 8					82	_	80	CDWR	PALERMO
	1111:30.6				49		1.70	CDWR	BUTTE CITY
1976 6- 9					74	2*	0.10	CDWR	OROVILLE
1976 6- 9					64	19#	90	CDWR	PARADISE SE
1976 6- 9	720:29.1	39.512	121.527	48	77		0.10	CDWR	OROVILLE
1976 6- 9	853:58.0	39.413	121.483	53	86	0#	20	CDWR	BANGOR
	1748: 9.7				31	3*	0.30	CDWR	ORD FERRY
	1338:21.5					_	0.60	CDWR	BANGOR
_	637:28.2					_	2.50	CDWR	PALERNO
1976 6-14							0.60	CDWR	PARADISE SE
	1836: 6.5			48			1.90	CDWR	OROVILLE
	1914:36.8						40	CDWR	OROVILLE
	2117:35.3						0.20	CDWR	BANGOR
	2330:25.5 2351:24.7						V 3.80 0.60	USGS	DANCOR
_	2351:24.7						0.50	CDWR CDWR	BANGOR PALERHO
13/0 0-14	#301:3J.U	33.7/3	121.940	77	, 3	J₩	V.3V	CDAR	FALERIO

DATE YEAR-NO-DY			LON.	DIST NI	DIST KM	DPTH KN	HMI	CRM CD	REF.	CONHENTS
+										
										DANGOD
1976 6-15 1976 6-15						0# 0#		0.40 40	CDWR CDWR	BANGOR PALERNO
	357: 3.8 536:15.0			48 19		6*		1.20	CDWR	WILLOWS
	1234:46.1			47		5#		10	CDWR	PALERHO
1976 6-16			121.463	52		1*		40	CDWR	BANGOR
	2234:41.7			26	42	6*		0.00	CDWR	RICHARDSON SPRINGS
	425:26.7			52	83	2*		70	CDWR	BANGOR
1976 6-18	556:47.4	39.447	121.463	53	85	0#		40	CDWR	BANGOR
	955: 4.3			52		3*		80	CDWR	BANGOR
	1141:57.7			34		17*		50	CDWR	HAMLIN CANYON
	1233:19.9			47		5*		50	CDWR	PALERNO
	2301:11.7			51	82	0# 4#		0.50 10	CDWR CDWR	BANGOR BANGOR
	525:30.6			55 50		0*		50	CDWR	BANGOR
	917:13.5 2008:26.1			53		4#		50	CDWR	BANGOR
	1201:19.5			53		2*		40	CDWR	BANGOR
	1556:57.9			39		4*		1.10	CDWR	ONO SW
	2043:19.2			54		4#		90	CDWR	BANGOR
	111:53.2			54	87	1*		60	CDWR	BANGOR
1976 6-25	152:28.7	39.754	121.976	20	32	5#		0.40	CDWR	NORD
1976 6-25	1912:22.4	39.418	121.466	54	87	0#		1.00	CDWR	BANGOR
1976 6-26	401:53.5	39.707	121.967	21	34	11*		0.20	CDWR	ORD FERRY
	603: 0.7			53		1*		0.80	CDWR	BANGÓR
	200:29.1			53		4*		80	CDWR	BANGOR
	1858:13.1			52		8#		10	CDWR	PALERHO
	1949:59.4			50		0#		20	CDWR	PALERNO
	102: 2.9			53		0#		20	CDWR	BANGOR
	703:58.6 838:59.3			47 53	76 86	5 <b>*</b> 4 <b>*</b>		1.90	CDWR CDWR	OROVILLE BANGOR
1976 6-30			121.497	53 53		0#		40	CDWR	BANGOR
	2014:28.7			52		1*		60	CDWR	BANGOR
	1047:18.4			49		7*		30	CDWR	PALERNO
	355:16.2			48		5	v	4.10	USGS	
	759:34.2			52		3*		0.60	CDWR	PALERNO
	17: 1.5			51		5*		60	CDWR	PALERNO
1976 7- 7	5?:21.3	39.859	121.643	37	60	0#		0.30	CDWR	PARADISE SW
	343:48.1	39.542	121.498	48	78	7*		1.60	CDWR	OROVILLE DAM
1976 7- 7	550: 1.5	39.611	122.133	17	28	12*		0.70	CDWR	WILLOWS
	1024:18.1			57	92	0#		0.80	CDWR	BANGOR
1976 7- 9				53	85	4*		30	CDWR	BANGOR
	2227:41.0			50	81	6#		80	CDWR	PALERHO
	2319:20.3			54	87	0#		1.50	CDWR	BANGOR
	2320:24.1 1923:12.7			53 50	86 80	1* 2*		0.80	CDWR CDWR	BANGOR
1976 7-12 1976 7-14	301:57.7			53	85	2*		30	CDWR	PULGA SW BANGOR
1976 7-14				53	86	5*		50	CDWR	BANGOR
	2118:44.1			49	79	5#		70	CDWR	PALERMO
1976 7-15				54	87	1*		80	CDWR	BANGOR
	440:41.0			25	41	0#		50	CDWR	RICHARDSON SPRINGS
	1337:12.6			52	83	9*		90	CDWR	PALERNO
1976 7-17				42	68	34+		60	CDWR	SHIPPEE
1976 7-17	1959:57.3			55	88	1*		40	CDWR	BANGOR
	454:33.1			52	84	5*		0.00	CDWR	PALERNO
1976 7-18	1341: 6.6	39.402	121.477	54	87	3*		0.10	CDWR	BANGOR

DATE	i	TIME	LAT.	LON.	DIST	DIST	DPTH	NNI CRM CD	REF.	COMMENTS
YEAR-MC	YQ-C	HRMN:SECS			MI	KM	KM			
+										
1976 7	7-19	1134:42.9	39.508	121.472	51	82	2#	90	CDWR	OROVILLE DAN
		747:24.0			54		2#	20	CDWR	BANGOR OROVILLE
		1621:15.7			48		10# 9#	80 10	CDWR CDWR	GLENN
		1924: 9.9			27		11*	0.60	CDWR	SUTTER BUTTES
		215:53.6 1226:43.6			55 53	_	4.	70	CDWR	BANGOR
		1257: 1.9			49		_	60	CDWR	PALERNO
		1906:29.2			48		10*	70	CDWR	OROVILLE
		412:25.0						0.10	CDWR	HAMILTON CITY
		953:49.7			57			10	CDWR	JONESVILLE NE
		1230: 0.2			40		0#	1.10	CDWR	CHEROKEE
		1436:57.2			51	82	0#	80	CDWR	BANGOR
1976	7-28	1437: 4.6	39.490	121.471	51	82	1#	70	CDWR	BANGOR
1976	7-28	2003:47.8	39.483	121.493	50		3#	60	CDWR	BANGOR
-		533:43.6					0#	50	CDWR	GLENN
		1130: 1.2						1.40	CDWR	PENNINGTON
		1200: 1.5					_	50	CDWR	BANGOR
		1727:41.3						1.00	CDWR CDWR	PALERNO PULGA SW
		1442:13.1						40 0.40	CDWR	HAMILTON CITY
	8- 1						_	70	CDWR	BANGOR
		1240:46.1 339:42.0				-		20	CDWR	BANGOR
	8- 3 8- 3	913:23.4						1.00	CDWR	OROVILLE DAM
		2028:25.6					_	90	CDWR	PALERNO
	8- 4							70	CDWR	BANGOR
		1405:44.8						1.00	CDWR	OROVILLE DAM
		2021:50.8				_		0.20	CDWR	BANGOR
	8- 6					86	12*	90	CDWR	BANGOR
	8- 7					89	0#	20	CDWR	BANGOR
	8-8	407:42.5				7 44	5*	10	CDWR	CHICO
1976	8-8	1952:45.4	39.487	121.503	50	80	5#	70	CDWR	PALERNO
1976	8- 8	1952:57.9	39.492	121.495	5 50	80	4#	60	CDWR	BANGOR
		1955:22.4				_		60	CDWR	BANGOR
		1504:21.1						80	CDWR	BANGOR
		1515: 1.1					_	80	CDWR	BANGOR OROVILLE
		1538:56.7					_	50 10	CDWR CDWR	BANGOR
		1746:47.3 439:15.3						60	CDWR	CHEROKEE
		828:29.							CDWR	BANGOR
		1113:36.9							CDWR	BANGOR
		1314:35.							CDWR	OROVILLE
1976		758:36.0							CDWR	BANGOR
1976		708:49.							CDWR	BANGOR
1976		1529:50.							CDWR	BANGOR
1976		833: 7.							CDWR	BANGOR
1976		1440:15.				0 6	4 10*	0.40	CDWR	WEST OF BIGGS
1976	8-16	619:22.	4 39.450	121.46	5 5	3 8			COWR	BANGOR
1976		5 1511:58.							CDWR	PALERNO
1976		5 2313:42.							CDWR	PALERNO
1976		1035:23.				-		_	CDWR	PALERNO
1976	8-19								CDWR	OROVILLE DAN
1976	8-19		7 39.46	5 121.47	5 5				CDWR	BANGOR
1976		815: 3.				9 7 2 8		2.90 3.30	USGS CDWR	BANGOR
1976	8~1S	9 815: 4.	/ 37.46	0 121.4/	<b>.</b> 3	2 8	a 1#	3.30	ODWK	PRROOM

LON. DIST DIST DPTH MMI CRM CD REF. TIME LAT. COMMENTS YEAR-MO-DY HRMN:SECS MI KN KN 1976 8-19 824:17.5 39.459 121.474 52 84 4# -.60 CDWR BANGOR 1976 8-19 1516:21.8 39.422 121.475 2\* -.50 BANGOR 53 86 CDWR 1976 8-20 126:57.8 39.463 121.475 52 84 8\* -.50 CDWR BANGOR 1976 8-21 2237:13.2 39.482 121.621 45 72 5# -.70 CDWR **PALERNO** 1976 8-24 710:11.2 39.368 121.386 59 95 5# -.10 COWR LOMA RICA 1976 8-24 715:48.5 39.424 121.475 53 86 1\* 0.20 CDWR BANGOR 1976 8-26 952:53.5 39.211 122.483 42 68 0# 2.00 CDWR WILBUR SPRINGS NW 1976 8-27 1217:13.7 39.427 121.527 82 -.30 CDWR 5\* 51 PALERNO 1976 8-30 1455: 2.9 39.402 121.532 83 7+ -.30 CDWR 52 **PALERNO** 76 1976 8-31 1929:35.6 39.518 121.536 47 6# 2.90 CDWR OROVILLE 86 79 1976 9- 1 927:40.5 39.398 121.492 53 6# -.70 CDWR BANGOR 1976 9-17 1421:20.1 39.501 121.510 49 8# -.60 CDWR **OROVILLE** 51 86 78 1976 9-18 951:36.4 40.267 122.261 32 5# 0.40 CDWR HOOKER 1976 9-19 2230: 1.0 39.414 121.475 53 4\* -.80 CDWR BANGOR 1976 9-27 1012:21.9 39.473 121.538 48 78 9\* -.80 CDWR **PALERNO** 1976 9-28 810:11.2 39.403 121.490 53 86 6\* -.30 CDWR BANGOR 1976 9-28 1639:43.1 39.416 121.482 86 9# -.50 BANGOR 53 CDWR 1976 10- 1 336:23.5 39.998 121.500 46 74 1# 0.40 CDWR PULGA NW 1976 10- 4 640:13.6 39.415 121.452 55 88 2\* 1.00 CDWR BANGOR 1976 10- 5 916: 0.1 39.410 121.472 54 87 0# -.40 CDWR BANGOR 85 1976 10- 6 1820: 1.8 39.396 121.502 53 4+ -.40 CDWR **PALERMO** 1976 10- 6 1905: 7.0 39.412 121.498 53 85 0# -.40 CDWR BANGOR 1976 10- 7 1741:57.0 39.351 121.451 57 92 0# -.90 CDWR LOMA RICA 1976 10- 8 314:35.8 39.410 121.437 89 0# 55 -.30 CDWR BANGOR 1976 10- 8 1506:31.1 40.040 121.360 54 87 1# 0.00 CDWR JONESVILLE SE 1976 10- 8 2331;25.1 39.394 121.502 85 -.60 CDWR 53 5\* **PALERNO** 1976 10- 9 219:22.2 39.531 121.477 50 -.70 80 5# CDWR OROVILLE DAM 1976 10-10 816:15.9 39.387 121.681 73 45 5# -.40 CDWR BIGGS -.90 1976 10-11 108: 3.4 39.399 121.493 86 1# CDWR 53 BANGOR 87 1\* 1976 10-12 1542:32.2 39.426 121.461 -.60 54 CDWR BANGOR 65 6\* 1976 10-15 530:25.3 39.835 121.582 40 -.50 CDWR PARADISE SE -.70 1976 10-16 121:14.8 39.407 121.483 53 86 0# CDWR BANGOR 1976 10-21 636:12.1 39.403 121.476 54 87 3# 0.80 CDWR BANGOR 1976 10-22 1254:24.0 39.403 121.470 54 5# 87 -.90 CDWR BANGOR 1976 10-24 7:31.3 39.404 121.483 53 86 2\* -.30 CDWR BANGOR 1976 10-24 25: 8.3 39.410 121.477 53 86 1\* -.90 CDWR BANGOR 1976 10-24 348:21.7 39.405 121.474 54 87 1\* -.40 CDWR BANGOR 1976 10-24 1014:24.6 39.409 121.458 55 8.8 0# -.40 CDWR BANGOR 1976 10-26 324:22.6 39.404 121.479 53 86 1# -.60 CDWR BANGOR 1976 10-27 918:55.8 39.388 121.541 52 83 5# - .20 CDWR **PALERNO** 1976 10-31 1707:12.1 39.412 121.848 38 61 17\* 0.30 CDWR WEST OF BIGGS 1976 11- 4 238:52.0 39.473 121.500 50 81 3# -.10 CDWR PALERMO 1976 11- 8 606: 8.6 39.669 121.926 24 39 14# 0.30 CDWR ORD FERRY 1976 11- 8 2003:50.7 40.030 121.630 40 65 5# 0.50 CDWR BUTTE MEADOWS SW 1976 11- 8 2315:34.4 39.513 121.497 49 79 7\* -.80 CDWR OROVILLE DAM 1976 11-10 814:57.2 39.411 121.484 53 86 0# -.80 CDWR BANGOR 1976 11-14 907:47.1 39.401 121.535 52 83 5# 0.00 CDWR **PALERMO** 1976 11-15 1717:30.6 39.394 121.667 46 74 5# 0.30 CDWR BIGGS 1976 11-19 38:41.4 39.510 121.493 50 80 5# -.70 CDWR OROVILLE DAM 1976 11-22 1852: 9.9 39.454 121.488 52 83 4. -.60 CDWR BANGOR 1976 11-23 657: 6.6 39.397 121.476 54 87 0# -.60 CDWR BANGOR 1976 11-23 1258: 6.3 39.468 121.479 52 83 1\* 0.60 CDWR **BANGOR** 1976 12- 8 1246:22.0 39.812 122.613 15 24 19 3.00 USGS

V 3.30

USGS

1977 1- 9 2324:39.5 39.500 121.643 43 69 2

DATE			LON.				MMI CRM CD	RLF.	CONNENTS
YEAR-NO-DY	DENETION			NI	KN	KN			
	2327:43.9				80	7*	2.30		PALERNO
	628: 1.5				86	0*	0.30		BANGOR
	330:13.7 1925:19.5				85 78	0 <b>*</b> 8 <b>*</b>	2.50 0.70		BANGOR OROVILLE
	2009:58.3				84	8*	3.70		FRENCH GULCH SE
	2042:40.0				81	4*	0.10		BANGOR
	732:54.8				82	0#	0.90		BANGOR
	418: 2.8				81	4#	0.00		BANGOR
	726:45.7				85	5#	0.20		BANGOR
	645:41.6				85	2#	0.50		BANGOR
	1248:33.1 1455:57.2				81	4*	0.30		BANGOR
	421:29.5				86 83	1* 4*	0.30 0.50		BANGOR BANGOR
	1323:26.0				82	1*	0.80		BANGOR
	1037:31.7			53	85	2*	0.20		BANGOR
	1102:14.9				87	1*	1.00		BANGOR
	826:14.0				85	5#	0.40		BANGOR
1977 1-25	157:52.7	39.405	121.482	53	86	0#	0.50		BANGOR
	1642:17.5				35	5*	1.00		GLENN
	337: 6.8				78	7*	0.10		OROVILLE
	1516:32.8				85	6*	0.00		PALERNO
	439:56.5				86 76	0# 8#	0.30		BANGOR
	1736:18.2 635:25.4				83	_	10 2.80		OROVILLE BANGOR
	2259:44.5				85	2*	10		BANGOR
	631:32.8				85	0#	0.50		BANGOR
	2217: 2.3				79	7#	20		OROVILLE DAM
1977 2- 3	20:13.4	39.438	121.460	53	86	2*	0.40		BANGOR
	1919:58.4	39.483	121.495	50	81	0*	0.00		BANGOR
	1449: 2.2			60	96	6#	10		LONA RICA
	1632:14.5			23	37	6#	1.00		GLENN
	454:54.9				83	1#	0.70		BANGOR
	644:30.9				87	3*	10		BANGOR
	645: 9.9 750:58.9				86	2*	0.80		BANGOR
	1235:13.6				<b>8</b> 6 <b>8</b> 9	2* 0*	0.50 0.00		BANGOR
1977 2-11						-	0.00		BANGOR BANGOR
1977 2-11				50	80	6*	0.10		OROVILLE DAM
1977 2-11	206:12.2	39.412	121.472	54	87	2*	0.40		BANGOR
1977 2-13				55	88	0#	10		BANGOR
	152:15.6			19	30	5*	0.80		ORD FERRY
	346: 7.5			45	73	10#	0.30		OROVILLE
	1435:25.8 551:14.7			53	86	4*	0.30		BANGOR
	1109:15.0			54 55	87 88	3# 10#	0.50 3.50		BANGOR
	1940: 0.4			67	108	5*	2.70		POTTER VALLEY NW MANZANITA LAKE SE
1977 2-26		-		50	80	5*	0.70		OROVILLE DAM
	553: 2.2			53	86	0*	0.60		BANGOR
	1424: 9.8			53	85	1+	0.00		PALERNO
1977 2-26	1934: 3.2	39.412	121.465	54	87	1#	0.20		BANGOR
	1313:54.8	39.445	121.481	52	84	0#	0.80		BANGOR
	541: 9.3			52	83	4*	0.40		BANGOR
	1755:32.6				86	0#	2.80		BANGOR
1977 3- 3	219: 5.7	<b>39.</b> 510	121.521	48	78	7*	0.20		OROVILLE

DATE Year-Mo-Dy	TIME HRMN:SECS	LON.	DIST NI	DIST KM	DPTH KN	NMI	CRM CD	REF.	CONNENTS
+		 							
	630:31.6			87	0#		0.00		BANGOR
	2229:37.4		53	86	4*		0.60		BANGOR
	2155:27.3		£ .	80	6*		0.80		PALERNO
	2354:11.6		54	87	1*		0.10		BANGOR
1977 3-12			53	85	3*		1.20		BANGOR
	1020:50.3		34	55	2*		1.00		HAMLIN CANYON
	413:26.7		53	85	7*		0.20		BANGOR
	1037:15.3		53	86	0#		0.40		BANGOR
	2338:39.4		49	79	7*		0.10		PALERNO
	626:21.0		50	81	2#		10		BANGOR
	2237:21.5		55	89	0#		2.60		BANGOR
	455: 5.8		50	80	8*		0.80		OROVILLE DAM
	1936: 2.6		55	88	0#		0.80		BANGOR
	1555: 4.1		50	80	_		0.70		OROVILLE DAN
	1608:54.8	 	49	79	9#		30		OROVILLE
	1941:36.8		50	80	7*		0.60		OROVILLE DAM
	350:46.2 16:12.2		32 49	51	9*		1.70		BUTTE CITY
	1202:50.4		48	79	7*		0.30		OROVILLE
			_	78	9*				OROVILLE
	1836:34.2 1249:41.8		49	79	8*		60		OROVILLE DAM
	1728:25.1		53	86	1*		0.40		BANGOR
	1020:44.5		55	89	0#		20	nece	BANGOR
			29	47	5			USGS	DAL COMO
	1750:42.9		50	81	10#		20		PALERMO
			52	84	5#		0.00		OROVILLE DAM
	1042:31.9		27	44	10#		1.90		LLANO SECO
	700:25.9		53	85	3#		0.20		BANGOR
	1937:37.1		51	82	10*		0.30		OROVILLE DAM
	1329:14.0		55	88	0#		20		BANGOR
	1819:30.6		52 53	83	0*		0.40		BANGOR
	1139:40.1			86	3#		1.10		BANGOR
	2257:11.4 517: 3.6		50 51	81	5*		10 0.30		OROVILLE DAM
	1057:58.8		49	82 79	4* 7*				BANGOR
	2310:14.8		51	82	4#		0.80		PALERHO BANGOR
	1231:23.1		53	85	8*		0.90		BANGOR
			_		_		0.80		
1977 4-22 1977 4-22	15:52.6 527:16.0		53 45	86 72	0 <del>*</del> 7*		10		BANGOR OROVILLE
	1324:56.4		<b>5</b> 3	86	1*		0.80		
	2314:17.5		55 55	88	3#		0.60		BANGOR BANGOR
	2225:49.2		53	86	1#		0.90		BANGOR
1977 4-27	426:29.2		53	86	1#		1.90		BANGOR
	1745:52.2		53	86	0#		1.20		BANGOR
1977 5- 4	611:11.4		53	85	3#		3.10		BANGOR
1977 5- 4	615:36.4		54	87	0#		2.90		BANGOR
1977 5- 4	659:10.3		53	86	0#		3.40		BANGOR
1977 5- 4	826:12.3		54	87	0#		0.20		BANGOR
	1342:24.2		53	86	0#		10		BANGOR
	1019:40.7		54	87	0#		0.10		BANGOR
	1641:44.3		53	85	0#		0.30		PALERNO
1977 5- 6	701:53.1		54	87	0#		1.50		BANGOR
1977 5- 6	822:52.2		54	87	0#		0.30		BANGOR
	1343:59.6		54	87	0#		0.50		BANGOR
1977 5-8	30:57.9		53	86	4*		0.70		BANGOR

DATE YEAR-NO-DY	HRNN:SECS			DIST NI	DIST KN	DPTH KN	NNI	CRN CD	REF.	COMMENTS	
*											
1977 5-10						-		0.10		BANGOR	
	1104:40.2					2*		0.40		BANGOR	
1977 5-11			121.481			0#		10		BANGOR	DAM
1977 5-11 1977 5-13	1644: 7.2				81 29	0# 9#		2.40 0.80		OROVILLE HAMILTON	
1977 5-13						> <del>*</del> 7*		0.30		OROVILLE	
1977 5-16					81	2*		10		BANGOR	Dan
1977 5-17						4#		1.00		BANGOR	
1977 5-18					86	0#		1.70		BANGOR	
1977 5-18						4#		0.30		BANGOR	
	2157:41.4				36	3#		2.30		GLENN	
	2349: 7.1 53:55.5				55	18*		0.50		PARADISE	SW
	657:51.4				81 87	5# 0#		30 0.50		BANGOR	
	518: 4.4				87	0#		0.70		BANGOR BANGOR	
	804:15.9				30	6*		1.60		GLENN	
	422:24.1					9*		10		OROVILLE	
	1056:43.9				80	5#		0.20		OROVILLE	
1977 5-25	1018:58.4	39.414	121.467	54	87	1#		1.60		BANGOR	
1977 5-26					82	7#		10		PALERNO	
1977 6- 2								0.00		OROVILLE	
1977 6- 2						-		0.50		BANGOR	
1977 6- 2					89	5*		0.40		BANGOR	
1977 6- 6 1977 6- 8	2019: 8.8				86 87	1 <b>*</b> 0 <b>*</b>		0.00		BANGOR BANGOR	
1977 6-8					90	1*		0.80		BANGOR	
1977 6-8						1*		1.80		PULGA NW	
	1736:26.6				87	0#		0.90		BANGOR	
1977 6-10	1121: 8.8	39.465	121.482	52	83	2*		30		BANGOR	
1977 6-12					29	6#		0.60		HAMILTON	CITY
1977 6-13					88	0#		0.50		BANGOR	
1977 6-17					76	8#		50		OROVILLE	
1977 6-21					89	0#		70		BANGOR	
1977 6-22 1977 7- 4	1128:37.6					0*		0.80		BANGOR	
1977 7- 5			_	_		6 <b>*</b> 1*		0.50		PALERMO Bangor	
	430:22.2				70	_		1.00		BIGGS	
	1554:30.7			50	80	6*		0.40		PALERNO	
1977 7-11	837:15.8	39.797	121.691	34	55	9#		2.70		PARADISE	SW
1977 7-14				48	78	8#		2.20		OROVILLE	
	423: 8.5			54	87	2*		20		BANGOR	
1977 7-16				25	40	6*		0.60		GLENN	
	1953:48.6			51	82	6#		0.40		PALERNO	
1977 7-18 1977 7-19				_	86 87	0*		1.30		BANGOR	
1977 7-19				29	46	2* 0*		1.40 2.70		BANGOR	co
1977 7-20				53	86	1*		1.40		LLANO SEG	<i>,</i> ,,
	1914:49.2				81	8*		0.40		PALERMO	
	1926: 9.8			50	81	4#		0.00		BANGOR	
1977 7-22	918:40.4	39.374	121.478	55	88	13*		0.80		LONA RICA	A
	1927:16.7			54	87	0#		0.80		BANGOR	
1977 7-29			122.004	19		_		0.80		HAMILTON	CITY
	1421:20.1					_		0.20		PALERMO	
1977 8- 1	744:41.2	40.156	121.420	54	87	0#		2.40		JONESVIL	LE NW

DATE YEAR-MO-DY	TIME HRMN:SECS		LON.	DIST NI	DIST KM	DPTH KN	NNI CRH CD	REF.	COMMENTS
+									
1977 8~ 1	2257:22.4	29 401	121 452	55	89	1*	0.50		BANGOR
- <del>-</del>	608:31.3				88	0*	0.10		BANGOR
	831:23.7				87	3*	10		BANGOR
	406:27.6				81	5#	0.50		PALERMO
	1035:28.8				82	5*	3.00		PALERNO
1977 8~ 6	1059:13.2	39.415	121.466	54	87	3#	30		BANGOR
	237:31.6				89	0#	20		BANGOR
	1931:55.5				82	1#	0.80		BANGOR
	2002:24.7				38	6 <b>*</b>	1.90		LOGAN RIDGE
	734:28.3			32 54	52 87	7 <b>*</b> 0 <b>*</b>	2.30 2.90		PANTHER SPRING SE WHITHORE SE
	1748:12.8 144: 3.9				84	3*	0.50		PALERNO
	1531:51.6				84	0#	0.90		PALERMO
	1346:59.9				86	0#	10		BANGOR
	351:13.3				42	8#	2.70		LOG SPRING
	1629:22.1				72	0*	1.60		POTTER VALLEY NE
1977 8-28	1839:26.6	39.428	121.494	52	84	5*	0.00		BANGOR
	1233:12.2				67	0#	2.80		LAKE PILLSBURY
	2257:44.3				82	8*	1.10		PALERNO
	1017:49.8				85	5#	0.00		BANGOR
	1615:15.7				43	12*	0.90		GLENN
	2044:10.4				88	3*	0.50		BANGOR
	2219:50.6				84	2*	0.40		BANGOR
	1145:23.8			42	68	5 <b>*</b>	3.00		FOUTS SPRINGS
	1153:22.6 1824:14.2				69 32	5* 4*	2.80 0.00		CLEARLAKE OAKS NW HAMILTON CITY
1977 9- 4					87	0*	0.70		BANGOR
1977 9- 8				49	79	7*	0.00		PALERNO
	1325:56.9				76	0*	0.40		BERRY CREEK
	428:28.5				89	0*	2.50		BANGOR
1977 9-13					89	0#	2.30		BANGOR
	1034:48.8				37	9*	1.60		ORD FERRY
	558:17.0		-		96	7*	2.50		NT. HARKNESS SW
1977 9-15	853:13.8	40.277	121.372	60	97	3*	1.10		NT. HARKNESS SE
	1241:47.8				87	0#	10		BANGOR
	701:16.4				87	0*	0.60		BANGOR
1977 9-20				52	84	9#	20		PALERNO
	1712:52.7			53	85	4#	0.30		BANGOR
	2210:14.2			54 55	87	0# 3#	0.30		BANGOR
	1233: 2.9 1200:15.0			55 54	88 87	2*	0.10 0.50		BANGOR BANGOR
	435:34.1			50	80	5*	40		PALERNO
	2112:19.5			55	88	3#	20		BANGOR
1977 10- 3				51	82	6*	2.50		PALERNO
1977 10- 8	18:15.7			51	82	3#	70		BANGOR
1977 10- 8	32:59.1			37	59	11	3.10	USGS	
1977 10-11				50	80	5#	0.40	-	BANGOR
1977 10-16	909:27.6	39.554	121.583	44	71	5#	0.90		OROVILLE
1977 10-16				43	69	5*	1.10		OROVILLE
1977 10-16				42	68	4+	0.60		OROVILLE
1977 10-18				50	81	2*	0.40		OROVILLE DAM
1977 10-18				53	85	0#	20		BANGOR
1977 10-18				54	87	3*	0.10		BANGOR
1977 10-19	1244:41.0	37.900	121.434	53	85	3*	0.90		BANGOR

DATE YEAR-NO-DY	TIME HRMN:SECS	LAT.	LON.	DIST NI	DIST KM	DPTH KN	MMI	CRN CD	REF.	CONNENTS
+										
1977 10-20	502:59.7	39.508	121.482	50	81	1*		0.00		OROVILLE DAM
1977 10-24						0#		0.30		BANGOR
1977 10-25	2000: 5.5	39.392	121.501	53	86	6*		0.10		PALERHO
1977 10-30					78	10#		0.70		PALERNO
1977 10-31					90	0#		0.30		BANGOR
1977 10-31					77	10*		0.30		OROVILLE
1977 11- 1					81	5*		50		BANGOR
1977 11- 3				53	85	0#		0.30		BANGOR
1977 11- 5				48	77	9*		20		OROVILLE
1977 11- 5				53	86	0#		0.20		BANGOR
1977 11- 7					86	6#		50		BANGOR
1977 11-10					64	2*		3.20		FOUTS SPRINGS
1977 11-10				53	86	0#		3.00		BANGOR BANGOR
1977 11-11					87 81	0# 3#		0.80		PALERNO
1977 11-12 1977 11-14					86	0#		0.30		BANGOR
1977 11-14					31	5*		0.60		GLENN
1977 11-14								0.60		BANGOR
1977 11-18						_		1.20		BANGOR
	1527:19.3							3.70		BRUSHY MTM.
	1058:21.5							0.70		ORD FERRY
	1228:43.1					_		0.50		BANGOR
	932:49.0							0.30		BANGOR
	1250:12.1							40		BANGOR
	1414:57.5							2.70		BANGOR
	953:15.3							0.80		BANGOR
	403:11.4				_			1.90		PALERNO
	939: 7.8							1.90		BANGOR
	1146:23.8							0.10		BANGOR
	1043:46.4							50		BANGOR
	842:41.8							1.00		BANGOR
	953:45.7							0.00		BANGOR
	1437: 8.2					-		0.70		OROVILLE
	1818: 0.7				_			60		OROVILLE
	458:10.4				-			60		BANGOR
	433:49.7							2.60		BUTTE HEADOWS NE
1977 12-13	935: 8.7	40.238	121.559	50	81	0#		0.70		BUTTE HEADOWS NE
	1408:55.2							0.70		PALERNO
1977 12-19	1207:10.6	39.521	121.896	31	50	5#		0.00		LLANO SECO
	2257:54.6				89	2*		0.40		BANGOR
	527:56.1				87	7 *		0.60		BANGOR
1977 12-20	849:43.0	39.390	121.436	5€	90	0#		0.40		BANGOR
1977 12-21	424:42.3	39.441	123.095	48	77	7 0*		3.30		POTTER VALLEY NE
1977 12-22	2 607: 3.8	39.450	121.942	33	<b>5</b> 3	12*		2.20		BUTTE CITY
1977 12-22	2 1718:10.9	39.341	121.593	52				30		HONCUT
1977 12-22	2 1729:43.9					8*		10		BANGOR
1977 12-23			121.515			-		30		PALERNO
	3 1941:15.9							10		BANGOR
	5 1106:36.7							1.10		WILLOWS
1977 12-27	7 1233: 8.7	39.452	121.387	57				1.10		BANGOR
1977 12-20	620:51.9	39.440	121.496	52				30		BANGOR
1977 12-20	8 1412:20.6	39.423	122.020	32				0.80		PRINCETON
1977 12-20	8 2042:52.8	39.461	121.485	5 52				0.10		BANGOR
1977 12-29	9 150:24.4	39.575	122.300	17	7 27	7 7*		1.80		STONE VALLEY

DATE Year-ho-dy			LON.	DIST MI	DIST KM	DPTH KN	NNI CRN CD	REF.	COMMENTS
+									*
1977 12-29	344: 7.5	39.445	121.542	49	79	1*	0.00		PALERNO
1977 12-30	114: 2.2	40.276	121.672	48	77	5*	1.60		LASSEN PEAK SW
1977 12-30				54	87	2*	0.30		BANGOR
1977 12-31					84	8*	0.10		PALERNO
1978 1- 1				38	61	5*	0.90		PANTHER SPRING SE
1978 1- 2 1978 1- 4				53 54	86 87	4# 0#	0.20 1.00		BANGOR BANGOR
	1042:47.2			42	67	9*	2.50		SANHEDRIN MTM.
	342:10.3			54	87	1*	10		BANGOR
	1247:46.7			54	87	6*	10		BANGOR
	1158:42.2				86	0#	20		BANGOR
	1451:48.6			42	68	23*	0.30		PENNINGTON
1978 1-11	808:37.0	39.383	121.361	60	96	3#	1.20		RACKERBY
	1911:26.8			53	86	3*	40		BANGOR
	2233:56.7			51	82	9*	0.00		BANGOR
	1432: 4.0			53	86	2*	0.80		OROVILLE DAN
	1442:18.8			23	37	10#	0.90		GLENN
	650:15.7			53	86	6*	10		BANGOR
	650:55.5 2059:18.1			53 48	86 77	5* 18*	0.30 1.90		BANGOR POTATO HILL
	2337:57.7			11	17	9#	1.10		ORLAND
	408:37.3			34	55	2*	1.30		HULL MOUNTAIN
	1126: 9.7			54	87	5	2.80	GS	36
	1539:23.8			56	90	5*	10		BANGOR
	2336:55.5			44	71	8#	60		OROVILLE
1978 1-27	155:54.2	39.445	121.447	54	87	0#	0.20		BANGOR
1978 1-28	1907:48.9	39.628	122.066	19	31	5*	1.10		HAMILTON CITY
1978 1-30	813: 1.1	39.405	122.308	28	45	16*	1.70		LOGAN RIDGE
	1928:29.5			55	88	0#	60		BANGOR
	2209:57.9			50	81	2#	1.00		OROVILLE DAN
	24:13.2			54	87	5*	0.00		OROVILLE DAN
	1055:34.1			54	87	6*	10		BANGOR
	2140:58.2			27	43	4*	1.70		RED BLUFF EAST
	2328:35.4				39	8#	1.00		RED BLUFF EAST
1978 2- 2 1978 2- 2					80 80	0# 0#	2.90 1.90		JONESVILLE SW JONESVILLE SW
1978 2- 5	947:55.0			53	86	1*	50		BANGOR
	1731:53.4			25	40	1#	2.30		RED BLUFF EAST
	2311:14.3			26	42	0#	2.20		RED BLUFF EAST
	2328:26.4			28	45	5*	1.20		RED BLUFF EAST
	2335:20.0	40.172	122.171	26	42	7#	0.90		RED BLUFF EAST
1978 2- 6	319:24.8			27	43	10*	1.10		RED BLUFF EAST
	1259:58.0			50	80	2#	0.10		PALERNO
1978 2- 7	220:48.0			53	86	5*	20		BANGOR
	1339:25.9			37	60	9#	3.20		HULL MOUNTAIN
1978 2-8	56: 0.1			53	86	5#	20		BANGOR
• • • • • • • • • • • • • • • • • • • •	1135:17.5			55 56	89 90	3 <b>+</b>	0.80		BANGOR
1978 2- 8 1978 2-11	1206: 8.5 738:18.8			56 54	87	5# 4#	0.10 30		BANGOR
	1556:13.3			53	85	1*	1.40		BANGOR Jonesville se
1978 2-11	608:45.7			25	40	2	3.30	USGS	vonestible se
	2354: 7.7			24	38	5*	1.10	1030	RED BLUFF EAST
	1450:27.1			53	86	8*	2.00		JONESVILLE NW
	1523:36.4	40.330	122.073	39	62	0#	1.50		DALES

DA1 YEAR-1		TIME HRMN:SECS		LON.	DIST		DPTH KM	INN	CRM CD	REF.	COMMENTS
1978	2-17	735:26.9	40.082	121.457	50	81	0#		3.10		JONESVILLE SW
		1339:44.7			50	81	8*		2.50		JONESVILLE SW
		222:20.4			52				0.70		PALERNO
1978	2-22	859:37.5	39.380	121.497	54	87			20		BANGOR
		810: 2.8			32	52			0.80		BUTTE CITY
		2354:44.5			27	43			1.20		RED BLUFF EAST
		551:28.7			26	42			1.20		RED BLUFF EAST
1978		603: 0.0			43	69			1.40		LAKE PILLSBURY
		1402:24.4			25	40			0.90		LLANO SECO
		645:55.6			49	79			0.60		PALERNO
1978		1007:26.1			22	35			1.30	60	ALDER SPRINGS
1978		1620:52.8			45	73			3.00	GS	36
		2123:49.6			53	86			0.50		BANGOR
1978		1355:14.7 232:18.6			50 54	81 87	7* 9*		0.00		PALERHO OROVILLE DAM
1978		114:56.3				85			20		PALERMO
1978		1659:42.8							0.90		ORD FERRY
		203:39.5							1.90		PARADISE NW
		354:35.1							1.50		JONESVILLE SW
		29:57.3							0.50		WEST OF BIGGS
		456:18.1							1.60		GLENN
		547:49.4							1.20		BUTTE CITY
		7:30.0							0.20		OROVILLE DAM
		1552:47.1							20		PALERNO
1978		2012:42.0				83	0#		0.70		BANGOR
		413:32.0				84	8*		0.50		BANGOR
		1143:54.2					1#		0.00		BANGOR
1978	3-13	1806: 6.2	39.405	121.472	54	87	2#		50		BANGOR
1978	3-15	935:45.2	39.234	121.756	50	81	27*		0.60		SUTTER BUTTES
1978	3-15	1612:31.9	39.508	121.990	28	45	9#		1.10		LLANO SECO
		38:56.4				87	4*		10		OROVILLE DAM
		946:47.0					_		0.70		BANGOR
		1112:27.6				_			40		BANGOR
		1547:59.3							1.70		MANZANITA LAKE SE
		1208: 7.9							0.10		LOMA RICA
1978		1813:12.9							1.70		PARADISE NE
1978		331:51.9					_		0.50		CHICO
1978	3-22								3.20	GS	36
1978		840:33.4							3.10		YOLLA BOLLY SW Yolla Bolly SW
1978		1049:46.2							3.10		
1978	3-23	17: 5.3		122.058					1.40		GLENN
1978		2037:42.6							1.40		JONESVILLE NW PALERMO
1978 1978	3-26			123.128					4.30		UKIAH
1978	3-26			123.126					3.10		COW HOU TAIN
1978		119: 9.6							3.60		COW HOU TAIN
1978	3-26								3.40		UKIAH
1978	3-26								3.40		UKIAH
1978		1620:30.5							0.60		BANGOR
1978	3-27								0.10		BANGOR
1978	3-27			121.484					1.30		BANGOR
1978	3-27			121.486					0.50		BANGOR
1978		1615:33.4							3.30		COW HOUNTAIN
1978		1343:35.7							1.10		BANGOR

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DA' YEAR-		TIME HRMN:SECS	LAT.	LON.	DIST NI	DIST KN	DPTH KM	HHI	CRM CD	REF.	COMMENTS
+											
							_				
1978		614:53.6			21	34	7*		1.60		WILLOWS
1978		950:43.8			52	83	0#		0.50		BANGOR
1978		1446:53.8			56	90	0*		1.50		BANGOR
1978		2342:49.1			32	52	21*		0.40		CHICO
1978	4- 3	402:37.6			55	88	0*		0.70		BANGOR
1978 1978	4- 4	2334:54.8 251:11.3			50 15	81 24	5* 5*		0.50 1.60		OROVILLE DAM HAMILTON CITY
1978		1438: 1.9		-	52	84	5#		1.40		UPPER LAKE
1978		1555:38.4			53	85	5*		1.70		UPPER LAKE
1978	4- 8	922:18.1			62	100	2*		3.30		UKIAH
1978		912:40.1			53	86	5*		0.00		BANGOR
1978	4-11				33	53	1*		2.00		MENDOCINO PASS
1978	4-11	214: 6.5	39.665	122.671	21	33	0#		3.10		ALDER SPRINGS
1978	4-12	236:24.1	39.416	121.475	53	86	0#		0.60		BANGOR
1978		2334:11.3			50	80	8#		0.70		OROVILLE DAM
1978	4-15	155:14.7			53	86	0#		0.70		BANGOR
1978		2150:19.0			57	91	2#		0.90		BANGOR
1978	4-18	341:19.7			22	36	8*		3.40		ALDER SPRINGS
1978		854:17.3			27	44	5		3.00	USGS	
1978		854:20.1			22	36	12*		3.50		ALDER SPRINGS
1978		1711:40.4			20	32	2*		3.20		ALDER SPRINGS
1978 1978	4-18	1723:52.7			17	27	0#		3.10		CHRONE
1978	4-19	138:18.1 419:50.8			19 24	30 38	0 <b>*</b> 5		3.50 3.20	USGS	ALDER SPRINGS
1978		419:51.9			21	33	13*		3.40	0363	ALDER SPRINGS
1978	4-20	42:51.1			22	35	1*		1.40		RICHARDSON SPGS NW
1978	4-20	650:44.3			53	86	3*		0.50		OROVILLE DAM
1978		1154:26.6			53	86	1*		0.20		BANGOR
1978		1435:19.0			21	33	0#		1.70		ALDER SPRINGS
1978	4-25	132:59.7			17	28	5		2.90	USGS	
1978	4-25	133: 0.3	39.696	122.615	17	27	0#		3.30		CHRONE
1978	4-25	1334:44.0	39.413	121.474	53	86	1*		0.50		BANGOR
1978	4-27	2017:57.1	39.497	121.398	55	88	0#		1.40		BANGOR
1978	4-28	912:59.2			16	26	4#		1.90		CHRONE
1978		1215: 5.6			53	86	3#		0.60		BANGOR
1978		1200:34.0			29	47	5#		2.00		PLASKETT MEADOWS
1978		1454:59.3			34	54	1*		1.70		HULL MOUNTAIN
1978	5- 1	37:43.8			53	86	0#		1.00		BANGOR
1978	5- 1	558:16.9 400:49.8			22 54	35	9# 0#		1.00		GLENN
1978 1978	5~ 2 5~ 3	437: 5.1			50	87 81	4*		1.20		BANGOR BANGOR
1978		1723:23.3			27	44	8#		3.20		RED BLUFF EAST
1978		2132:51.4			39	62	10#		2.10		LODOGA SW
1978		1438:13.9			52	84	9#		0.70		BANGOR
1978		1153:35.4			22	35	12*		1.20		GERBER
1978		1231:33.4			24	39	7*		1.70		RED BLUFF EAST
1978	5- 7	1314:26.2	40.164	121.477	52	83	10#		1.00		JONESVILLE NW
1978	5- 8	129:59.5	39.708	121.798	30	48	11#		1.70		CHICO
1978	5- 8	945:10.4	39.482	121.995	29	47	13#		1.90		BUTTE CITY
1978	5-11	703:50.8			17	27	5#		2.80		CHRONE
1978	5-11	707:14.4			17	27	12*		2.40		CHRONE
1978		1601:10.8			17	28	2*		1.70		ALDER SPRINGS
1978	5-13	433:17.6			45	72	10#		1.10		OROVILLE
1978	5-13	707:31.6	39.682	122.652	19	31	0#		2.40		ALDER SPRINGS

DATE		TIME	LAT.	LON.				MMI	CRM CD	REF.	COMMENTS
		HRMN:SECS			MI	KM	KH				
+								***			
1978 9		516:14.3			53		0#		0.80		BANGOR
		654:19.9			19		0#		1.80		ALDER SPRINGS PALERNO
		955:13.4			52 31	<b>84</b> 50	0# 4#		2.40		SAINT JOHN MTM.
		2211:53.1 2238:45.6			31 34	-	4#		3.20		CROCKETT PEAK
		1221:40.0			64				1.70		MT. HARKNESS NW
		240:53.1			53		0#		0.70		BANGOR
		2232: 2.6			55		5#		1.00		BANGOR
		1426:18.4			53	85	7*		0.10		BANGOR
	5-24	1950:22.6	40.026	122.128	19	30	5#		1.80		GERBER
1978	5-25	354:23.2	39.402	121.490	53	86	0#		1.10		BANGOR
	5-25	2258:51.5	39.469	121.968	31	50			1.10		BUTTE CITY
		1938:59.0			50		2*		0.30		PALERNO
		1613: 7.9		-	35		0#		1.90		DALES
_		1658:48.0			52				0.20		BANGOR
		2034:59.6			31		_		1.30		PRINCETON
		403:22.1			50		9*		0.40 0.50		OROVILLE DAM OROVILLE
		1103:37.1 2027:31.1			48 53		-		2.10		BANGOR
		1205:43.1							0.90		BANGOR
		537:18.0							0.60		BANGOR
		2328:42.0			-				0.40		PALERNO
	6- 4			122.245		47	12*		1.70		LOGANDALE
		2234:14.9				76	5#		0.60		PALERNO
	6- 6	1815:29.1	39.185	121.441	65	104	8#		0.90		BROWNS VALLEY
		2307:37.1				38	7#		2.30		PANTHER SPRING SW
	6-15								1.10		CHICO
1978	6-17	209:18.8	39.444	121.486					0.80		BANGOR
	6-17								0.70		BANGOR
_		1534:27.6							0.60		LONA RICA
		958: 5.5							0.60		BANGOR
	6-21						_		0.30		PALERNO
		1119:57.6							3.70 30		LEECH LAKE MTN. PALERNO
		1125: 8.3 213: 8.8							1.00		GLENN
	6-26 6-27								0.80		BANGOR
		746:29.3					14*		1.30		FOSTER ISLAND
1978	7- 2	1355:30.1	39.400	121.539	51				0.10		PALERMO
		1926:56.9							30		BANGOR
1978		207:33.4				3 86	3#		0.10		BANGOR
1978		1010:53.9				1 39			0.80		GLENN
		2054: 2.2							0.60		PALERNO
1978		1056:13.7							0.60		BANGOR
1978		704:22.1							0.30		BANGOR
1978		718:32.4							2.10		BANGOR
1978	7-11			121.496					0.00		BANGOR
1978		1242:43.9							0.40 1.40		BANGOR OREGON HOUSE
		1407:57.6 2 1007:46.6							1.20		GLENN
1978 1978	7-14	832: 4.	, 33.300 1 39 410	121.464	5 5				1.00		BANGOR
1978	7-14			2 121.54					0.50		PALERNO
1978		1449:28.							0.20		BANGOR
1978		5 554: 4.0							2.80		BANGOR
1978		5 838:32.0							1.90		DALES

DATE YEAR-NO-DY		LAT.	LON.	DIST NI	DIST KM	DPTH KN	NNI CRH CD	REF.	CONHENTS
+									
1978 7-15	1918:28 0	29 555	122.038	24	38	11*	1.00		GLENN
1978 7-16			122.035		41	9*	0.60		GLENN
	1207: 6.1			55	88	0#	40		HONCUT
	1342:57.3			55	88	9#	0.90		BANGOR
	1447:59.4			50	81	6#	1.20		PALERMO
	1846:43.9	40.272	121.419	58	93	6*	1.10		NT. HARKNESS SW
	506:51.0			53	86	6#	0.10		PALERNO
	733:36.0			54	87	4#	3.30		BANGOR
	539:42.5			56	90	0#	0.50		BANGOR
	552:44.6			21	33	0#	3.00		RAGLIN RIDGE
	807:24.5			21	33	12 <b>*</b> 10 <b>*</b>	2.10 1.60		LOWREY LOGAN RIDGE
	833: 8.2 534:33.9			22 55	35 89	0*	0.50		BANGOR
	1657:41.0			50		5*	0.60		PALERNO
1978 8- 3				46	74	2*	0.20		PENNINGTON
1978 8- 3				50		5*	0.10		PALERNO
	1232: 8.9			52	84	0#	0.80		BANGOR
1978 8-15	2009: 7.0	39.774	121.700	34	55	18*	0.50		PARADISE SW
	1939:27.1			56	90	0#	0.50		BANGOR
	2317:33.1			23	37	5*	1.50		RICHARDSON SPGS NW
	1945:25.8			52	84	0#	1.00		BANGOR
	1515:15.4			27	44	0#	1.70		FELKNER HILL
	1814:48.7			52	83	10*	0.70		OROVILLE DAM
	1842:52.1			22	36	1*	1.60		ALDER SPRINGS
	1844:26.5			30	48	1*	1.60		KNEECAP RIDGE
	407:55.6 2141:14.3			53 53	86 86	4# 0#	0.00 2.60		BANGOR BANGOR
	1612:24.3			56	90	1*	0.50		RACKERBY
	2008:37.2			55	89	0*	0.70		BANGOR
	2103:21.5			52	84	4*	0.70		BANGOR
	521:44.1			50	80	6*	0.20		PALERMO
	1042:40.5			52	84	7*	0.00		PALERNO
1978 9- 2	825:55.8	39.447	121.487	52	83	1*	0.10		BANGOR
	1009:30.4			55	88	0#	0.40		BANGOR
	1327:50.6			28		5*	1.40		KNEECAP RIDGE
	1006:52.9			9	14	6*	1.50		KIRKWOOD
	136:45.4			50	81	0#	0.60		PALERNO
	1820:18.7			53	85	0#	0.90		BANGOR
	514:36.5			45	73	3*	0.70		OROVILLE
	1211:39.7			53	86	0#	1.00		BANGOR
	1247:12.1 1748:19.2			53 54	86 87	0# 0#	40 1.30		BANGOR BANGOR
1978 9-13 1978 9-15	329:31.8			53	85	0*	2.20		BANGOR
	2113: 6.9			57	91	5#	1.80		JONESVILLE NW
1978 9-17	746:42.7			53	86	0#	1.80		BANGOR
1978 9-17				53	86	1*	0.80		BANGOR
	259:13.1			59	95	15#	1.10		JONESVILLE NE
	1124: 9.7			55	88	0#	0.50		BANGOR
	2208:28.4			52	84	4*	0.00		BANGOR
	2236:31.8			51	82	5#	0.80		PALERNO
1978 9-24	359: 7.2			57	91	0#	0.20		BANGOR
	1606:24.8			53	85	6#	0.50		BANGOR
1978 9-29 1978 10- 2			121.854	25 55	41	0*	2.20		RICHARDSON SPRINGS
19/8 10- 2	2331:32.0	37. <b>9</b> 32	141.426	55	89	12*	0.20		BANGOR

DATE YEAR-NO-DY	TIME HRMM:SECS	LAT.	LON.	DIST	DIST	DPTH KM	MMI	CRM CD	REF.	COMMENTS
+										**
1978 10- 3	522:41.0	39,404	121.456	55	88	3#		0.20		BANGOR
1978 10- 4				52		5#		0.30	CDWR	BANGOR
1978 10-10				52		16*		1.00		PALERNO
1978 10-10	2212:45.3	40.005	121.982	23	37	8#		1.60		PANTHER SPRING SW
1978 10-13	929:17.9	39.415	121.484	53		4#		1.70		BANGOR
1978 10-13						0*		0.60		BANGOR
1978 10-17				53		4*		1.40		BANGOR
1978 10-18						5*		0.50		CHICO
1978 10-21						1#		0.90		BANGOR
1978 10-21						5*		0.30		PALERMO OROVILLE
1978 10-21				48		10*		0.20		BANGOR
1978 10-22					_	5*		0.30 0.50		BANGOR
1978 10-22						10*		0.20		BANGOR
1978 10-23 1978 10-23					88 50	5*		0.80		CHICO
1978 10-25						0#		0.40		BANGOR
1978 10-28		_			-	4*		1.40		BANGOR
1978 10-28								1.60		BUTTE MEADOWS SW
1978 11- 4						12*		1.30		GLENN
1978 11- 7								0.00		BANGOR
1978 11- 7								0.60		BANGOR
1978 11-12				-				3.60		CROCKETT PEAK
1978 11-15					59	20*		0.20		PARADISE SW
1978 11-16						4#		0.70		BANGOR
1978 11-16					85	5*		0.30		PALERNO
1978 11-18	2136:42.8	39.399	121.524	52	84	7#		1.10		PALERNO
1978 11-22	1040:19.1	39.378	121.535	52	84	6#		0.10		PALERNO
1978 11-24	648:40.5	39.711	121.701	35	56	7*		0.40		HAMLIN CANYON
1978 11-24	712:17.1	39.732	121.652					0.90		HAMLIN CANYON
1978 11-24						1*		0.50		BANGOR
1978 11-28						_		0.20		BANGOR
1978 11-28						_		1.10		HAMLIN CANYON
1978 11-29								1.30		PALERHO
1978 12- 1								2.50		HAMILTON CITY
1978 12- 2								1.70		CHICO
1978 12- 3	19:52.9	39./15	121.683	35				0.10 1.80		HAMLIN CANYON FOSTER ISLAND
1978 12- 7	1834:22.3 1957:26.1	39.819	122.083	14		9# 3#		0.60		PALERMO
	505:46.1							0.60		BANGOR
	702:29.8							2.10		RICHARDSON SPRINGS
	422:54.3							0.40		BANGOR
	652:25.5							0.40		PALERNO
	437:24.0							0.90		NELSON
	904:40.2							0.50		OROVILLE
1978 12-15	2024:34.1	39.446	121.461	53				0.20		BANGOR
1978 12-19			121.446					0.30		LOMA RICA
	141: 6.1							0.70		BANGOR
1978 12-21	917:35.0	39.721	122.082	2 19		_		1.30		HAMILTON CITY
1978 12-22	708:32.5	40.264	121.430	5				1.50		NT. HARKNESS SW
1978 12-22	733:15.6	39.456	121.539	49		8*		0.10		PALERNO
1978 12-22	1524:52.0	39.776	121.889	9 24	4 36	27*		1.30		NORD
1978 12-27	7 522:12.9	39.71	121.669	3 3				0.60		HAMLIN CANYON
1978 12-21	7 1323:13.7	7 39.474	121.484	4 5:	_			0.50		BANGOR
1978 12-20	452:58.	3 <b>39.4</b> 19	5 121.50	5 5	2 84	74	<b>)</b>	0.60		PALERNO

DATE YEAR-NO-DY	TIME HRMN:SECS	LAT.	LON.	DIST NI	DIST KN	DPTH KN	HHI	CRN CD	REF.	CONHENTS
+										
1979 1- 2	349:37.6	39.517	122,061	25	40	12*		0.70		GLENN
· -	1448: 0.7			24		10#		0.70		LLANO SECO
	725:53.6			51	82	5#		0.40		PALERMO
	1400:54.5			29	46	5#		0.60		RICHARDSON SPRINGS
_	510:51.1	39.694	121.661	37	59	19#		0.40		HAMLIM CANYON
	1811:20.1			55	88	0#		0.10		BANGOR
	327:52.9			54	87	3#		40		BANGOR
	1507:53.4			15	24	4#		1.50		HAMILTON CITY BANGOR
	2014:29.0 238: 6.3			53 55	86 89	0#		40 0.80		BANGOR
	652:11.9			55 55		3#		0.10		BANGOR
	1905:46.1			52		12*		1.20		PALERNO
	626: 5.1			52				0.30		BANGOR
	513:20.9			50		7#		0.90		OROVILLE DAM
	1322:21.9			48	77	9#		0.90		OROVILLE
	1502:35.4	39.553	121.508	48		_		2.60		OROVILLE
	121: 6.9			55	88	_		1.00		BANGOR
	912:38.3			14	23	3#		1.80		CHRONE
	1024:27.9			25	41	11*		0.80		GLENN
	2028:19.4		-	28	45	12*		2.50		LOGANDALE
	2111: 3.6			30 25	49 40	10# 9#		0.60 0.90		LLANO SECO GLENN
	2315:21.9 554:31.5			53	86	2*		0.40		BANGOR
	1009:29.5			2 <del>9</del>		12*		1.50		LOGANDALE
	2140:12.7			45	72	9*		2.40		MANOR SLOUGH
	631: 1.1			45	73	7*		1.20		NO QUAD FOR THIS EF
	1015:29.0			68	109	5*		0.80		WHEATLAND
1979 2- 6	828: 3.9			31	50	14*		2.10		PRINCETON
1979 2- 8	406: 4.2			55	88	1*		0.00		BANGOR
1979 2-8	1717:17.4			25	41	4#		1.00		NO QUAD FOR THIS EF
1979 2- 9	2247:22.7	39.423	121.496	52	84	7#		1.10		BANGOR
	1854: 1.7			44		11*		0.80		BIGGS
	830:40.8			27		8*		1.00		NO QUAD FOR THIS EF
	556:12.8			50		3*		0.10		BANGOR
1979 2-22	359:36.5			54				0.70		BANGOR
				54		1*		0.30		BANGOR
1979 2-24 1979 2-24	347:58.3 841: 0.8			55 53	88 85	0# 5#		1.00		BANGOR BANGOR
	1340:22.2			39	62	25*		0.80		NO QUAD FOR THIS EF
	2015:20.1			24	39	21*		0.90		NORD
1979 3- 2	828:20.2			51	82	5*		0.20		PALERNO
1979 3- 3				54	87	3*		0.50		BANGOR
	1944:46.7			22	36	5#		1.00		NORD
1979 3-4	212:40.5	39.393	121.480	54	87	9*		0.40		BANGOR
	1052:12.4			27	43	16*		1.40		LLANO SECO
	1924:54.0			22	36	2*		0.70		GLENN
	1330:22.9			43	70	1*		3.10		FOUTS SPRINGS
	1914:18.3			47	75	8#		0.10		PALERMO
	2047:16.1			55	88	5*		40		BANGOR
	1649:38.2			54	87 50	1* 7*		0.10		BANGOR HEST OF RICCS
	501:30.1 2339: 4.9			37 <b>5</b> 0	60 81	/# 4#		0.40		WEST OF BIGGS BANGOR
	1009:14.2			51	82	3*		0.50		BANGOR
	1929:41.6			54	87	1*		1.20		BANGOR
· ·					<b>J.</b>	- "				

DATE VEAR-MO-DV	TIME HRMN:SECS	LAT.	LON.	DIST NI	DIST KN	DPTH KN	HMI	CRM CI	REF.	COMMENTS
+										
1979 4- 6	2025:51.9	39.398	121.473	54	87	0#		0.70		BANGOR
	2027:41.5			53	86	3#		0.50		BANGOR
	2339: 1.4			27	43	5#		1.70		CHICO
	2344:22.7			51	82	4*		1.10		BANGOR
	1449:52.7			54	87	0#		1.40		BANGOR
	938:54.5			24	39	8#		1.30		GLENN
	1028:10.5			32	51	24*		0.90		NELSON
	1030:14.8			41	66	7#		0.40		CHEROKEE
	1101:51.2			23	37	9*		0.60		GLENN WILLOWS
	758:17.9			16	25	9*		2.80		BANGOR
	1840: 0.7 1035:27.9			53 <b>49</b>	86 79	4 <b>*</b> 1*		0.20		OROVILLE
	1035:27.3			49		1*		0.10		OROVILLE
	1224:50.0			53	86	1*		0.90		BANGOR
	1318:47.6			52	83	3*		0.90		BANGOR
	515:20.5			40		5#		1.40		BUTTE MEADOWS SW
	1352:36.7			29		11*		0.60		LLANO SECO
	1439:18.7			29				0.70		CHICO
	505:31.7			53		0#		0.60		BANGOR
	621:30.2			54		0#		0.90		BANGOR
	1019:50.4			51	82	8*		1.00		BANGOR
1979 5- 4			121.976		33	7#		1.30		ORD FERRY
	1743:13.5				38	7#		1.50		GLENN
1979 5- 4	2214:54.3	39.403	121.508	52	84	2#		0.50		PALERNO
1979 5-13	220:24.1	39.610	121.386	53	85	17*		0.50		OROVILLE DAM
1979 5-13					87	15*		20		FORBESTOWN
1979 5-19	5 1926:24.5	39.613	121.392					10		OROVILLE DAM
1979 5-10			121.506					0.50		PALERNO
	5 949:46.4							0.80		BANGOR
	7 1148:56.2					_		1.10		LLANO SECO
	l 1439: 4.5							1.30		PALERNO
	2 1252:25.9					-		0.90		BANGOR
	4 1706:19.0							0.90		LLANO SECO
	5 1338:28.5							0.40		BANGOR
	5 1858:32.3							1.70		HAMILTON CITY
	8 1014:19.6					_		0.00		BANGOR
	2 1358:50.5				87			0.30		BANGOR
	3 455:11.4 9 650:27.4					_		1.00		PALERNO PRINCETON
	3 2017:33.8							2.60		MANZANITA LAKE SW
	3 2017:53.6 3 2025:53.2							1.90		MANZANITA LAKE SW
	5 1111: <b>3</b> 7.9							0.50		BANGOR
	9 727:31.0							1.90		HOOKER
	6 903:10.3							2.90		JONESVILLE SW
	9 1529: 3.0							0.50		BANGOR
	3 1742:21.6							0.60		OROVILLE
	6 1457:49.4							2.60		PALERMO
	7 408:38.3							1.50		BANGOR
	7 416:48.6				80	7*		0.50		OROVILLE DAM
1979 7-1			121.505		9 79	9#		0.70		OROVILLE
	5 1446: 0.5				3 80	5 5*		2.60		BANGOR
	5 1517:18.				3 80	8*	•	2.80		BANGOR
	5 1645:42.				2 84	8 8	•	0.30		BANGOR
	5 1702:28.	8 39.42	5 121 <b>.48</b> 6	5 5	3 8	5 74	<b>,</b>	0.50		BANGOR

DATE YEAR-MO-DY				MI	KM	KM		REF.	CONNENTS
1979 7-15					85		 0.50		BANGOR
	550: 7.1						0.20		BANGOR
	740:53.6						0.70		BANGOR
	2053:16.7				86		0.30		BANGOR
	835:10.8						3.10		HALL RIDGE
	2220:11.0					0#	3.10 0.50		OROVILLE
and the second s	2225:30.0				81	1*	0.60		OROVILLE DAM
1979 7-22	102:48.7	39.495	121.489	50	81	6*	0.20		BANGOR
1979 7-23	858:57.6	39.736	122.001	19	30	6#	1.50		HAMILTON CITY
1979 7-26	1016: 7.7	39.583	122.053	22	35	7*	1.20		GLENN
1979 7-29	2314:12.6	39.746	121.848	26	42	8#	0.90		CHICO
1979 8- 2	1526: 7.7	39.615	122.368	14	22	9#	2.70		STONE VALLEY
1979 8- 5	712:48.9	39.612	122.206	16	25	8*	1.30		WILLOWS
1979 8- 6	707:39.1	39.429	121.459	54	87	5*	20		BANGOR
1979 8-8					84	4*	0.50		BANGOR
1979 8- 9					80	7*	0.50 3.00		BANGOR
1979 8-11					84	5*	3.00		PALERNO
1979 8-11					86	0#	1.10		BANGOR
1979 8-11							0.70		PALERNO
1979 8-11							0.30		BANGOR
1979 8-12			121.408				0.10		BANGOR
1979 8-13					82	8*	0.80		BANGOR
	423:21.4				88	3*	0.20		BANGOR
1979 8-25					87	3*	0.60		BANGOR
1979 9-4					89	0#	0.70		BANGOR
1979 9- 7	1149:54.8	39.426	121.471	53			0.90		BANGOR
1979 9-21	1828: 0.1	39.425	121.542	50	81	0#	0.50		PALERHO
1979 10- 7	617:37.8	39.416	121.540	50	81	1#	0.10		PALERNO
1979 10-13	503:46.0	39.412	121.482	53	86	4+	0.50		BANGOR
1979 10-17	208:39.0	39.390	121.455	55	89	4*	10		BANGOR
1979 10-27	2259:36.7	40.385	122.105	41	66	5*	2.50		NO QUAD FOR THIS EF
1979 10-28	622:59.8	39.754	121.881	25	40	17#	1.50		NORD
1979 10-28	1748:54.8	39.414	121.495	53	85	3#	2.90		BANGOR
1979 10-28	1801:37.2	39.403	121.469	54	87	1*	0.80		BANGOR
1979 10-31	17: 4.8	39.636	121.572	42	68	17#	0.50		CHEROKEE
1979 10-31	2006:25.3	39.426	121.516	51	82	6*	3.10		PALERNO
1979 11- 1	1230: 4.1	39.407	121.484	53	86	3*	0.60		BANGOR
1979 11- 2	345:48.8	39.403	121.474	54	87	0#	0.90		BANGOR
1979 11- 2	844:36.9	39.405	121.472	54	87	1*	0.40		BANGOR
1979 11- 3	1317:23.0	39.394	121.446	55	89	0*	10		BANGOR
1979 11- 3	1749:53.2	39.416	121.475	53	86	2*	0.90		BANGOR
1979 11- 6	2017:47.0	39.440	121.616	46	74	0#	J.60		PALERNO
1979 11-14	622:16.6	39.391	121.471	55	88	6*	2.70		BANGOR
1979 11-14	751:21.4	39.3 <del>9</del> 2	121.475	54	87	1*	0.20		BANGOR
1979 11-21	2335:25.5	40.243	121.377	59	95	8=	2.70		JONESVILLE NW
1979 11-22				37	59	19+	2.80		PLASKETT RIDGE
1979 11-24	844:32.1	39.732	121.984	19	31	7+	1.70		ORD FERRY
1979 11-25				25	41	9*	1.10		GLENN
1979 11-27				53	86	0*	0.30		BANGOR
1979 11-28				28	45	5*	2.50		LOGANDALE
1979 11-28				54	87	0*	0.80		BANGOR
1979 11-28				54	87	6.	0.40		BANGOR
1979 12- 1				49	79	8=	0.50		OROVILLE DAM
1979 12- 1	402:44.6	39.797	122.709	20	32	5*	2.10		HALL RIDGE

YEAR-NO-DY	HRMN:SECS				DIST KN		NNI CRN CD	REF.	CONNENTS
+									
1979 12- 1						11*	0.90		PARADISE SW
1979 12-12						0#	1.00		BANGOR
1979 12-15							1.80		WILLOWS HAMILTON CITY
	252:42.6 255: 9.9						2.30 1.50		HAMILTON CITY
1980 1-11						5*	0.50		BANGOR
1980 1-14							2.80		LAKE PILLSBURY
	1531: 2.5					8*	2.70		JONESVILLE NW
	1541:27.8					0#	2.70		HONCUT
1980 1-30					86	5#	0.60		BANGOR
1980 2-4						8#	2.30		JONESVILLE SW
	2157:38.1				35	7#	1.50		WILLOWS
	1613:19.7				_	_	2.60		OROVILLE
	707:22.5 1239:48.9						2.30		LLANO SECO
	2335:30.8					2* 7*	1.90 2.60		NAXWELL Dales
	2243:10.6						2.40		JONESVILLE NW
1980 4-24							2.40		HAMILTON CITY
	1118:15.7						2.00		LAKE PILLSBURY
1980 4-28	620: 4.4	39.645	121.907	25	41	9#	1.60		ORD FERRY
1980 4-29	1350:22.6	40.437	123.480	75	120	5	3.00	GS	36
	416: 1.2						1.70		BANGOR
1980 5-12							3.10	CDWR	HALL RIDGE
	1146:49.8						1.00		BANGOR
	1910:54.1						1.10		RACKERBY
	229:27.5 236:40.0						3.40 2.40		THATCHER RIDGE
	2157:17.5					-	2,90		BANGOR KNEECAP RIDGE
	758:12.1								GILSIZER SLOUGH
	1612:31.6						2.60		CROCKETT PEAK
	2132:18.0				46	3#	1.00		NELSON
	1017:38.7				86	5#	0.30		BANGOR
	209:19.5						1.60		BANGOR
	1008:11.2						0.30		BANGOR
	932:59.6						0.40		BANGOR
	633:32.9 1710:43.1						2.40		PALERNO
	2011:56.0			30 54	49 87	8 <b>*</b> 0*	1.50		KNEECAP RIDGE
	2131: 5.5					1#	2.20 0.50		BANGOR BANGOR
	1030:33.5			52		0#	0.50		BANGOR
	1257:47.5			53		3#	1.40		BANGOR
1980 8-30	701:10.9	39.417	121.483	53	85	2*	1.30		BANGOR
	1019:41.5					5#	2.10		JONESVILLE NW
	524: 2.0			53		8*	2.20		PALERNO
	938:52.1					4*	2.10		OREGON HOUSE
	40:13.9					1#	2.60		SHIPPEE
	943:10.5 1609:27.7					204	1.10		SUTTER
	1355: 7.3					11 <b>*</b> 5 <b>*</b>	1.80 3.20	GS	LLANO SECO 36
	2337:34.0			48		9#	0.50	43	OROVILLE
	347:26.8				_	11#	0.80		LLANO SECO
	2038: 7.7					12*	2.60		CORNING
	710: 5.1					12*	3.10		GERBER
1980 11-15			121.984		39		2.30		LLANO SECO

YEAR-NO-DY	TIME HRMN:SECS	LAT.	LON.	DIST NI	DIST KM	DPTH KM	NMI CRM CD	REF.	COMMENTS
+									
1980 11-18					87	1*	0.40		BANGOR
1980 11-19					86	8#	0.60		BANGOR
1980 11-19					90	0#	0.30		BANGOR
1980 11-19					75	6#	2.90		OROVILLE
1980 11-19					86	1#	0.00		BANGOR
1980 11-21				40	64	5#	1.20		SHIPPEE
1980 11-24					66	5	3.30	USGS	
1980 11-24					67		3.80		WILLIAMS
1980 11-26					66	31*	3.10		WILLIAMS
1980 11-28					67	9*	3.00		WILLIAMS
1980 11-29					68	27*	2.40		WILLIAMS
1980 11-29					87	1*	1.50		BANGOR PARADISE SW
1980 11-30					60	7# 9#	2.30		BUTTE MEADOWS SW
1980 11-30 1980 12- 1					64 56		2.10 2.80		SITES
1980 12- 1					82	5	2.80	HECE	31123
1980 12- 9					88	5 <b>*</b>	1.70	0505	UPPER LAKE
1980 12-10					21	_	2.60		FOSTER ISLAND
1980 12-12				35	56	20*	2.30		MAXWELL
1980 12-12				36	58	11*	2.50		MAXWELL
1980 12-12					42	10*	2.20		LLANO SECO
1980 12-12					62	5	3.40	USGS	DUNNO DEVO
1980 12-12					68	15*	3.60	1000	WILLIAMS
1980 12-14					66	15*	3.20		WILLIAMS
1980 12-15				30	49	12*	2.90		LOGANDALE
1980 12-20				21	34	11*	2.00		GLENN
1980 12-23				39	62	0#	1.30		PARADISE SE
1980 12-25				32	51	6#	0.40		NELSON
1980 12-27				17	28	10*	1.60		HAMILTON CITY
1980 12-31	1:36.8	39.906	123.075	40	64	14*	3.00		LEECH LAKE MTN.
	1605:33.4			54	87	1#	0.30		BANGOR
1981 1- 6	803:16.6	39.554	121.975	26	42	17*	2.60		LLANO SECO
1981 1-6	813:24.7	39.559	122.011	25	40	11*	2.30		GLENN
1981 1-6	1305:18.2	39.459	121.505	50	81	7*	0.10		PALERNO
1981 1- 7	2005:28.7	39.573	121.958	26	42	0#	2.20		LLANO SECO
	2134:46.0			39	63	5#	1.10		MANOR SLOUGH
	2315:39.6			43	70	5#	10		CHEROKEE
	1406:29.3			60	97	0#	0.20		OREGON HOUSE
	1313:28.6			48	77	7*	3.40		OROVILLE
	1046:10.7			21	33	6*	2.70		RED BANK
	1147:50.5			25	41	7*	2.00		PANTHER SPRING SW
	410:33.7 1840:54.8			54 21	87 33	0# 24#	1.70 1.50		BANGOR
	841:33.1			48	78	24× 5	3 <b>.5</b> 0	USGS	HAMILTON CITY
	1424:31.2			37	60	4*	1.00	0363	MOULTON WEIR
	1847:44.8			27	43	8*	2.30		CHICO WEIK
	2305: 2.4			39	63	5#	3.00		LAKE PILLSBURY
	2305: 2.4			30	49	10#	3.00		SITES
	1457:39.5			53	86	0#	0.80		BANGOR
1981 2-26	241:23.9			29	47	9*	2.00		CHICO
	1742: 0.7			25	40	10*	0.60		LLANO SECO
	457:10.9			53	85	8*	1.90		BANGOR
	855:33.8			55	88	1*	0.10		BANGOR
	1011:50.4			55	89	5*	0.00		BANGOR
							• •		

DATE TIME YEAR-NO-DY HRMN:	E LAT.	LON. DIST	DIST	DPTH KM	NNI CR	M CD	REF.	COMMENTS
+								
1981 3-14 1850::	26.5 40.188 12	1.362 58	93	8*	2.	70		JONESVILLE NE
	35.7 39.411 12			4#		60		BANGOR
	44.2 39.689 12		30	10*		80		HAMILTON CITY
	58.8 39.474 12			0*		80		BANGOR
	35.9 39.432 12		85	0#		30		BANGOR
	58.1 39.521 12		41	21*		80		GLENN
	47.6 39.364 12			3*		30		LOMA RICA
	22.8 39.711 12		29	28*		70		ALDER SPRINGS
	47.0 39.376 12 9.9 40.437 12		87 70	7 <b>*</b> 9*		60 20		BANGOR
	26.3 39.397 12		87	3× 1∗		20 80		NO QUAD FOR THIS EP BANGOR
	4.7 39.226 12		68	7*		<b>7</b> 0		CLEARLAKE OAKS NE
	8.4 40.278 12		84	10*		80		LASSEN PEAK SE
	6.6 39.412 12		85	0#		90		BANGOR
	37.8 39.397 12		87	5*		40		BANGOR
	11.4 40.079 12		79	10*		10		JONESVILLE SW
1981 4-28 2053:2	26.0 39.401 12	1.453 55	89	5#	٥.	80		BANGOR
	45.9 39.763 12		50	26*		40		RICHARDSON SPRINGS
	7.7 39.910 12		37	0#		00		BALL MOUNTAIN
	10.5 39.957 12			1*		00		BLUENOSE RIDGE
	47.8 39.604 12			4#		00		HULL NOUNTAIN
	4.1 39.633 12		51	1#		00		PLASKETT RIDGE
	11.7 39.670 12		43 33	0#		00 50		PLASKETT HEADOWS STONE VALLEY
	36.7 39.523 12 55.5 39.365 12		33 94	21 <b>*</b> 5*		<del>4</del> 0		LOMA RICA
	26.5 40.465 12		98	4#		90		LASSEN PEAK NE
	57.0 39.999 12		41	11*		30		RICHARDSON SPGS NW
	32.0 40.165 12		_	19*		20		YOLLA BOLLY NW
	39.7 40.077 12			0#		00		BLACK ROCK MTN NE
	51.3 39.475 12			6#		00		CROCKETT PEAK
	0.8 39.441 12			2*		10		BANGOR
	44.6 39.859 12			36*		00		FOSTER ISLAND
1981 5-24 918:	47.1 39.376 12	2.836 40	65	5*	0.	00		CROCKETT PEAK
1981 5-24 1051:	6.0 39.787 12	2.608 14	23	18#	٥.	00		NEWVILLE
	7.7 39.593 12	<del>-</del>		17#	0.	00		KNEECAP RIDGE
	5.8 39.622 12			7*		00		KNEECAP RIDGE
	26.6 40.003 12							PANTHER SPRING SW
	53.0 <b>39.899</b> 12		44	0#		00		BALL NOUNTAIN
	51.0 39.929 12		61	0#		00		LEECH LAKE HTN.
	19.0 39.351 12 24.0 <b>39.858</b> 12		90	0#		00		REDWOOD VALLEY
	24.0 39.856 12 30.2 39.373 12		50	3* 9*		00		MENDOCINO PASS
	30.2 39.373 12 11.9 39.508 12		97 72	5#		00 00		OREGON HOUSE SANHEDRIN MTN.
	3.1 39.933 12		60	0*		00		LEECH LAKE MTM.
	40.4 39.906 12		64	1*		00		LEECH LAKE MIN.
	15.7 39.533 12		72	5*		00		SANHEDRIN MTN.
	15.1 39.383 12		89	19*		40		POTTER VALLEY NW
	33.2 39.811 12			3#		00		HALL RIDGE
1981 6- 4 1854:	<b>20.2 39.811 12</b>			0#		00		HALL RIDGE
	53.9 39.593 12		69	0*	1.	40		OROVILLE
	25.6 <b>39.775</b> 12		46	13*	0.	00		LOG SPRING
	26.9 39.347 12			7#		70		OREGON HOUSE
	47.5 39.971 12			0#		00		BUCK ROCK
1981 6-8 135:	5.9 40.090 12	23.065 43	69	1#	0.	.00		BLACK ROCK HTN SE

		TIME HRMN:SECS	LAT.	LON.	DIST MI	DIST KM	DPTH KM	NHI CRN CD	REF.	COMMENTS
+										
1981	6- <b>9</b>	1558: 4.3	39 410	122.649	32	52	0*	0.00		SAINT JOHN MTN.
1981		2344:21.9				41	8*	1.60		PANTHER SPRING SW
1981		1234:51.0				42	9*	0.70		LLANO SECO
1981		1702:26.5			48	77	0#	0.00		COVELO EAST
1981		2204:12.0			46	74	0#	0.00		BRUSHY MTM.
1981		2230: 5.6			34	55	0#	0.00		PLASKETT RIDGE
1981	6-15	750:12.0			23	37	1*	0.00		ALDER SPRINGS
1981	6-15	1130:50.5	39.677	122.756	24	39	4#	0.00		PLASKETT NEADOWS
1981		701: 9.8			43	69	5*	1.90		COVELO EAST
1981		839:29.0			19	31	5#	0.00		HALL RIDGE
1981		339:54.7			31	50	5*	0.00		MENDOCINO PASS
1981		1804:29.1			48	78	0#	0.00		BLUENOSE RIDGE
1981		2337:57.1			53	86	0* 3*	2.00		BANGOR Lona Rica
1981 1981		1454:12.2 244: 6.7				92 38	10#	0.50 1.60		GERBER
1981		1133:19.2				76	1#	1.60		CLEARLAKE OAKS NW
1981		2204: 6.5			9	15	5*	0.00		FLOURNOY
1981		107:34.8			27	43	15*	0.00		PLASKETT MEADOWS
1981		1155:22.4			36	58	0#	0.00		HULL HOUNTAIN
1981		1456:26.1			53	86	0#	2.90		BANGOR
1981		1850:14.1			54	87	2*	1.00		BANGOR
1981		1926:40.8			17	28	10*	0.00		HALL RIDGE
1981	7- 1	446:54.3	39.458	121.539	49	79	4*	0.10		PALERMO
1981	7- 2	128:21.9	39.835	122.659	17	28	10	3.20	USGS	
1981	7- 2	131:35.7	39.854	122.743	22	35	0#	0.00		HALL RIDGE
1981	7- 2	144: 8.5				87	0#	0.50		BANGOR
1981	7- 2	152:58.9			21	33	1#	0.00		HALL RIDGE
1981	7- 2	337:47.2				34	0*	0.00		HALL RIDGE
1981	7- 2	541:39.9			21	34	10*	1.60		HALL RIDGE
1981	7- 2	548:58.0				33	0#	0.00		HALL RIDGE
1981		1120:44.1			21	34	0*	0.00		HALL RIDGE
1981		1711:28.3			22	35	22 <b>*</b> 1 <b>*</b>	0.00		RILEY RIDGE HALL RIDGE
1981 1981		2350:22.7			21 21	34 34	0#	0.00 0.00		HALL RIDGE
1981	7-3	127:10.6			56	90	0*	0.10		BANGOR
1981	7- 4				40	65	4#	0.00		NEWHOUSE RIDGE
1981	7- 4	320:45.8			33	53	12*	1.40		STONYFORD
1981		2257:45.5			33	53	14*	1.30		HAMLIN CANYON
1981	7- 8	817:30.9			53	85	5*	1.90		LASSEN PEAK NW
1981	7- 9	1636:27.9			19	30	2*	0.00		HALL RIDGE
1981	7- 9	1707:49.3	39.846	122.732	21	34	3*	0.00		HALL RIDGE
1981	7-10	408:56.3			32	52	0#	0.00		PLASKETT RIDGE
1981	7-11	330:47.4			13	21	33*	0.00		NEWVILLE
1981	7-11	359:40.2			34	54	9*	0.00		KNEECAP RIDGE
1981	7-11	411:52.9			43	69	0#	0.00		POTTER VALLEY NE
1981	7-12	200:56.9			17	28	2*	0.00		RILEY RIDGE
1981	7-12				17	28	6#	0.00		HALL RIDGE
1981		1108:42.6 1115:20.5			54 55	87 89	0#	0.40		BANGOR DOTTER WALLEY MA
1981 1981		1527:38.6			53	85	1*	0.00 0.90	CDUP	POTTER VALLEY NW
1981		1856:30.5			33	53	0*	0.90	CDWR	BANGOR DI ASVETT BIDGE
1981	7-14	229:21.6			26	42	1*	0.00		PLASKETT RIDGE FELKNER HILL
1981	7-14				45	73	2*	0.00		LAKE PILLSBURY
1981		2130:57.2			53	86	0#	1.90		BANGOR
-								· -		

DATE		LAT.	LON.				MMI	CRM CD	REF.	COMMENTS
+	HRMN:SECS			MI	KM	KM				
								0.00		DI ACKETT DIDER
	2300:20.7					0* 14*		0.00		PLASKETT RIDGE LOMA RICA
1981 7-15 1981 7-16	2308:43.1		121.430	58 53	94 86	0*		1.30		BANGOR
	1056:13.5			25	41	3*		1.90		RICHARDSON SPRINGS
	1112: 8.7			30	48	0#		0.00		PLASKETT NEADOWS
	233:43.1			30	49	0*		0.00		PLASKETT RIDGE
	234:13.1			34	54	1*		0.00		PLASKETT RIDGE
1981 7-17	423:43.4	39.476	121.441	53	85	5#		0.00		BANGOR
	1343:11.8			18	29	2*		0.00		HALL RIDGE
	1307:11.5			43	70	0#		1.30		OROVILLE
	2303:12.1			43	70	3#		1.90		OROVILLE
	440: 2.6			36	58	0#		0.20		SHIPPEE
	909:54.9			37	60	12*		2.40		DALES
	1410:10.8			22	36	13#		1.20		NORD
	1821:12.5			42	68	1* 1*		0.00		SANHEDRIN HTN. LOG SPRING
	123:19.5			24	39 52	20*		0.00		HAMLIN CANYON
	1208:37.1					20 <del>*</del>		0.70		LOMA RICA
	723:27.8					-		0.00		FELKNER HILL
	724:50.8							0.20		BANGOR
	1008:30.1							0.00		HULL HOUNTAIN
	1105: 2.5				_			0.00		ELK HOUNTAIN
	1105:15.4				77	0#		0.00		ELK HOUNTAIN
	540:50.7				59	0#		0.00		HULL HOUNTAIN
	810:15.2				53	2*		0.00		HULL HOUNTAIN
1981 7-29	928:11.9	39.623	123.023	39				0.00		SANHEDRIN MTM.
1981 7-30	1151:28.4	39.304	123.101	54	-			2.20		POTTER VALLEY
1981 7-30	1733:38.0	40.191	122.551	29	46	5#		0.00		OXBOW BRIDGE
	111:50.7							0.00		POTTER VALLEY NE
	1634:27.2							0.00	•	MENDOCINO PASS
	2206: 4.4							0.00		RILEY RIDGE
	2212:56.8							0.00		MENDOCINO PASS
	1626:44.6							0.40		NELSON
	1623:21.1							0.00		YOLLA BOLLY NE
	925: 1.1 1048:15.0							0.00		POTTER VALLEY NE LOG SPRING
	1522:55.2							0.00		CROCKETT PEAK
1981 8- 4								0.00		HALL RIDGE
	515: 7.0							0.00		NEWHOUSE RIDGE
	1035:13.2							0.00		JAMISON RIDGE
	1836:58.3				38	4*		0.00		ALDER SPRINGS
1981 8- 6	359:42.8	39.702	122.940	33	53	10*		0.00		PLASKETT RIDGE
1981 8- 6	412:59.2	39.883	122.608	16	25	4#		0.00		PASKENTA
1981 8- 7	7 1802:42.7	40.014	121.931	25	41	8#		1.30		PANTHER SPRING SW
1981 8-10	423:23.6	39.934	123.123	43	69	9#		0.00		LEECH LAKE MTN.
1981 8-13					49	0#		0.00		BUCK ROCK
	2 1615: 3.7							0.00		YOLLA BOLLY SW
	1447:10.1							0.00		LEECH LAKE MTM.
	1705: 7.1							0.00		ALDER SPRINGS
	2130:12.5							0.00		RILEY RIDGE
	5 550: 1.7					-		0.00		POTTER VALLEY NE
	7 1909:22.2							0.40 0.00		BANGOR LOG SPRING
	8 1518: <i>3</i> 0.5 0 1650:57.8							0.00		BLACK ROCK MTN NW
1981 8-20	0 1030:3/.0	70.100		, 52	. 01	. •		····		PERON NOON DIN NA

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DATE YEAR-NO-DY	TIME HRMN:SECS		LON.	DIST MI		DPTH KN	MMI	CRM CD	REF.	COMMENTS
+										
1981 8-21	321:58.0	39.979	123.132	44	71	0#		0.00		BLUENOSE RIDGE
1981 8-22						0#		0.00		LOG SPRING
	1418:48.6			14		0#		0.00		CHRONE
1981 8-24	134:33.5	39.602	123.211	49	79	1#		0.00		BRUSHY MTN.
	445: 3.4			52	83	1*		0.00		POTTER VALLEY NW
	2140:53.2			27		0#		0.00		PLASKETT NEADOWS
	2259:33.7			23	37	5#		0.00		LOG SPRING
	1517:27.3			57	91	1#		0.00		REDWOOD VALLEY
	1610:27.7			12 43	19 70	5# 5#		0.00		ORLAND SANHEDRIN MTN.
	2118:37.0 2349:59.7							0.00		JAMISON RIDGE
	2345: 9.9							0.10	CDWR	
	1036:19.1			4		_		0.00	<b>VD</b> w.i.	BLACK BUTTE DAN
•	1135:21.8			53				1.50		BANGOR
	1749:33.4			43	70			0.00		YOLLA BOLLY NW
	1240:23.2			22	36	0#		0.00		RILEY RIDGE
	1544:45.5			28	45	10#		0.00		FELKNER HILL
	1646:40.6			28	45	0#		0.00		BALL HOUNTAIN
	333:21.2			24	38			0.00		BALL HOUNTAIN
	341:26.7			25	40			0.00		BALL HOUNTAIN
	1329:35.8			21	34	5#		1.70		RED BANK
	2101:17.9			48	77	5#		0.00		CHANCHELULLA PEAK S
	1728:52.3			62	99	7*		2.90		LONA RICA
	158: 3.1			30	49	30 <b>*</b>		1.40		PRINCETON
	817:47.7 2113:32.0			23 62	37 99	0*		0.00		LOG SPRING UKIAH
	2239:15.8			29	46	8*		0.00		BALL MOUNTAIN
	6:31.6			47	75	4#		0.00		POTTER VALLEY NE
1981 9-26	A			24	38	0*		0.00		LOG SPRING
	2036: 9.0			31	50	-		0.00		MENDOCINO PASS
1981 9-27				48	78	2*		0.00		POTTER VALLEY NE
	2249:10.4			36	58	22#		0.90		HAMLIN CANYON
1981 9-28	1946:15.5	39.325	123.029	50	81	4*		0.00		POTTER VALLEY
1981 10- 3	1256:56.6	39.870	123.108	41	66	0#		0.00		NEWHOUSE RIDGE
1981 10- 6								0.00		NEWVILLE
1981 10- 8				56				1.10		JONESVILLE NW
1981 10- 9					53			0.00		BUCK ROCK
1981 10-10				15	24	30#		0.00		CHRONE
1981 10-10 1981 10-11				50	81	0#		3.10		PALO CEDRO
1981 10-11				33 33	53 53	5# 0#		0.00		STONYFORD MENDOCINO PASS
1981 10-11			122.937	49	79	0#		0.00		ELK HOUNTAIN
1981 10-12				42	67	23*		2.40		OROVILLE
1981 10-12				42	67	5*		0.00		LEECH LAKE MTN.
1981 10-12				36	58	8*		1.60		HAMLIN CANYON
1981 10-14	50:35.9	40.377	122.190	40	64	5#		1.60		NO QUAD FOR THIS EF
1981 10-14				34	54	0#		0.00		MENDOCINO PASS
1981 10-15				18	29	0#		0.00		HALL RIDGE
1981 10-15				4	6	5#		0.00		SEHORN CREEK
1981 10-18				23	37	4*		0.00		ALDER SPRINGS
1981 10-18				23	37	5*		0.00		LOG SPRING
1981 10-19				25	41	5*		0.00		BALL MOUNTAIN
1981 10-23 1981 10-23				26 26	42 42	3* 2*		0.00		BALL HOUNTAIN
1201 10-73	1/07120.7	33.63/	141.000	<b>∠</b> 0	42	4		1.40		ORD FERRY

1981 10-24	DATE YEAR-NO-DY	TIME HRMN:SECS			MI	KN	KM	CRM CD	REF.	COMMENTS
1981 10-24 420: 7.4 39.420 122.346 27 44 10* 0.00 LOGAR RIDGE 1981 10-25 458: 4.6 39.836 121.825 27 44 24* 2.50 RICHARDSON SPRINGS 1981 10-25 602:24:1 39.958 122.540 15 24 14* 0.00 COVELD EAST 1981 10-25 502:1:1.8 39.958 122.540 15 24 14* 0.00 COVELD EAST 1981 10-26 532:1:1.8 39.958 122.540 15 24 14* 0.00 COVELD EAST 1981 10-26 2047:40.5 39.886 123.142 43 59 0* 0.60 RICHARDSON SPRINGS 1981 10-26 2308:28.8 39.846 122.300 3 5 5* 0.60 RICHARDSON SPRINGS 1981 10-26 2308:28.8 39.646 122.300 3 5 5* 0.00 BLUENDES RIDGE 1981 10-31 1806:30.7 39.658 122.872 83 45 1* 0.00 PLASKETT MEADOWS 1981 10-31 1007:554 40.025 122.80 32 52 0* 0.00 PLASKETT MEADOWS 1981 11- 1 1007:554 40.025 122.901 34 54 6* 0.00 PLASKETT MEADOWS 1981 11- 2 1916:23.0 39.591 123.050 39 62 1* 0.00 LEECH LAKE HTM. 1981 11- 3 107:554 6.0161 122.096 27 44 9* 2.50 THIOCHARDSON SPRINGS 1981 11- 5 1072:28.2 40.161 122.096 27 44 9* 2.50 THIOCHARDSON SPRINGS 1981 11- 5 10527:28.2 40.161 122.096 27 44 9* 2.50 THIOCHARDSON SPRINGS 1981 11- 7 1700: 2.4 39.652 122.644 75 59 0* 0.00 PLASKETT RIDGE 1981 11- 7 1700: 2.4 39.652 122.964 37 59 0* 0.00 PLASKETT RIDGE 1981 11- 7 1700: 2.4 39.652 122.964 37 59 0* 0.00 PLASKETT RIDGE 1981 11- 7 1700: 4.8 39.630 122.964 37 59 0* 0.00 PLASKETT RIDGE 1981 11- 7 1905: 6.8 39.630 122.742 23 37 0* 0.00 PLASKETT RIDGE 1981 11- 9 1458:11.5 39.915 122.742 23 37 0* 0.00 PLASKETT RIDGE 1981 11- 9 1503:40.7 39.918 122.742 23 37 0* 0.00 PLASKETT RIDGE 1981 11- 9 1503:40.7 39.918 122.742 23 37 0* 0.00 RILEY RIDGE 1981 11- 10 247:11.6 39.751 122.009 18 29 8* 1.10 FOSTER ISLAND 1981 11-10 1407:42.2 39.752 122.009 18 29 8* 1.10 FOSTER ISLAND 1981 11-10 1407:42.2 39.752 122.009 18 29 8* 1.10 FOSTER ISLAND 1981 11-10 1407:42.2 39.752 122.009 18 29 8* 1.10 FOSTER ISLAND 1981 11-10 1407:42.2 39.752 122.009 18 29 8* 1.10 FOSTER ISLAND 1981 11-10 1407:42.2 39.752 122.009 18 29 8* 1.10 FOSTER ISLAND 1981 11-10 1407:42.2 39.752 122.009 18 29 8* 1.10 FOSTER ISLAND 1981 11-10 1407:42.2 39.752 122.009 18 29 8* 1.10 FOSTER ISLAND 1981 11-10 1407:42.2 39										HULL HOUNTAIN
1981 10-24 1241;33.0 39,869 123.053 39 62 5* 0.00 MEMMOUSE RIDGE 1981 10-25 458: 4.6 39,835 121.825 27 44 24* 2.50 RICHARDSON SPRINGS 1981 10-25 502;131.8 39,816 123.227 47 76 1* 0.00 PASKENTA 1981 10-26 521:11.8 39,816 123.227 47 76 1* 0.00 RICHARDSON SPRINGS 1981 10-25 1318:31.2 39,777 121.793 29 47 5* 0.60 RICHARDSON SPRINGS 1981 10-26 2047;40.5 39,886 123.142 43 69 0* 0.00 BLUEMEGSE RIDGE 1981 10-26 2047;40.5 39,886 123.142 43 69 0* 0.00 BLACK BUTTE DAM 1981 10-30 2216;24.2 39.558 122.827 28 45 1* 0.00 PLASKETT RIDGE 1981 11- 1 1007;56.4 40.025 122.880 32 52 0* 0.00 PLASKETT RIDGE 1981 11- 1 1916;123.0 39,919 123.050 39 52 1* 0.00 PLASKETT RIDGE 1981 11- 2 1916;23.0 39,919 123.052 39 52 0* 0.00 FLACK MITL NOUNTAIN 1981 11- 2 1222;54.6 39,647 123.092 42 67 0* 0.00 HULL MOUNTAIN 1981 11- 5 1027;28.2 40.161 122.096 27 44 9* 2.50 TUSCAN SPRINGS 1981 11- 5 1163:135.6 39,668 122.1910 32 52 0* 0.00 FULGE ALSKETT RIDGE 1981 11- 7 1700; 6.8 39,668 122.1912 32 52 0* 0.00 FULGE ALSKETT RIDGE 1981 11- 7 1700; 6.8 39,630 122.1912 33 53 10* 0.00 FULGE ALSKETT RIDGE 1981 11- 7 1700; 6.8 39,630 122.147 23 37 0* 0.00 FULGE ALSKETT RIDGE 1981 11- 7 1936; 9.1 39,915 122.747 23 37 0* 0.00 FULGE ALSKETT RIDGE 1981 11- 9 1365; 9.1 39,915 122.747 23 37 0* 0.00 RILEY RIDGE 1981 11- 9 1503:40.7 39,918 122.742 23 37 0* 0.00 RILEY RIDGE 1981 11- 9 1503:40.7 39,918 122.742 23 37 0* 0.00 RILEY RIDGE 1981 11- 10 407;42.2 39,752 122.099 14 22 35* 1.10 FUNCAN FILL RIDGE 1981 11-10 1407;42.2 39,752 122.099 14 22 35* 1.10 FUNCAN FILL RIDGE 1981 11-10 1407;42.2 39,752 122.099 14 22 35* 1.10 FUNCAN FILL RIDGE 1981 11-10 1407;42.2 39,752 122.099 14 25 35* 1* 0.00 FULG MILL RIDGE 1981 11-10 1407;42.2 39,752 122.099 14 25 35* 1* 0.00 FULG RIDGE 1981 11-10 1407;42.2 39,752 122.099 14 25 35* 0* 0.00 FULG RIDGE 1981 11-10 1407;42.2 39,752 122.099 14 25 35* 0* 0.00 FULG RIDGE 1981 11-10 1407;42.2 39,752 122.099 14 25 35* 0* 0.00 FULG RIDGE 1981 11-10 1407;42.2 39,752 122.099 14 25 35* 0* 0.00 FULG RIDGE 1981 11-10 1407;42.3 39,563 122.740 39 55* 122						44	10#			
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1981   10-26   2047;40.5   39.886   123.142   43   59   0*   0.00   0.										
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1981   11- 1   1007-156.4   40.025   122.880   32   52   0+ 0.00   0.00   1981   11- 2   1916:23.0   39,919   123.050   39   62   1+ 0.00   0.00   1981   11- 2   2012: 1.0   39.607   122.901   34   54   6* 0.00   0.00   1981   11- 4   2226:54.6   39.647   123.092   42   67   0* 0.00   0.00   1981   11- 5   1027:28.2   40.161   122.096   27   44   9* 2.50   1981   11- 5   1051:35.6   39.668   122.910   32   52   0* 0.00   0.00   1981   11- 5   1451:35.6   39.668   122.910   32   52   0* 0.00   0.00   1981   11- 7   1700: 6.8   39.630   122.964   37   59   0* 0.00   0.00   1981   11- 7   1700: 6.8   39.630   122.912   33   33   10* 0.00   1081   10.00   1981   11- 7   1700: 6.8   39.630   122.912   33   33   10* 0.00   1081   11- 9   1306: 9.1   39.915   122.747   23   37   0* 0.00   1881   11- 9   1306: 9.1   39.915   122.747   23   37   0* 0.00   1881   11- 9   1306: 9.1   39.915   122.742   23   37   0* 0.00   1881   11- 9   1306: 9.1   39.918   122.742   23   37   0* 0.00   1882   1881   11- 9   1393: 13- 39.918   122.742   23   37   0* 0.00   1882   1881   11- 9   1396: 9.1   39.918   122.742   23   37   0* 0.00   1882   1881   11- 9   1396: 9.1   39.918   122.742   23   37   0* 0.00   1882   11- 9   1395: 37   39.315   121.411   60   97   2* 0.30   1004   1881   11- 9   1306: 9.1   39.751   122.009   18   29   8* 1.10   FOSTER ISLAND   1981   11-10   1407: 42.2   39.752   122.009   18   29   8* 1.10   FOSTER ISLAND   1981   11-10   1407: 42.2   39.752   122.009   18   29   8* 1.10   FOSTER ISLAND   1981   11-10   1407: 42.2   39.752   122.009   18   29   37   0* 0.00   1888   188										
1981   11- 2   1916:23.0   39.919   123.050   39   62   1*   0.00   LEECH LAKE MTN.     1981   11- 2   2012: 1.0   39.607   122.901   34   54   6*   0.00   HULL MOUNTAIN     1981   11- 5   1027:28.2   40.161   122.096   27   44   9*   2.50   TUSCAN SPRINGS     1981   11- 5   1451:35.6   39.668   122.910   32   52   0*   0.00   HULL MOUNTAIN     1981   11- 6   1806: 7.0   39.591   122.964   37   59   0*   0.00   HULL MOUNTAIN     1981   11- 7   1700: 6.8   39.630   122.912   33   53   10*   0.00   FRUTO     1981   11- 7   1700: 6.8   39.630   122.912   33   53   10*   0.00   PLASKETT RIDGE     1981   11- 7   1306: 9.1   39.915   122.742   23   37   0*   0.00   JULIAN ROCKS     1981   11- 9   1306: 9.1   39.915   122.742   23   37   0*   0.00   RILEY RIDGE     1981   11- 9   1503:40.7   39.918   122.742   23   37   0*   0.00   RILEY RIDGE     1981   11- 9   1503:37.3   39.315   121.411   60   97   2*   0.30   LOMA RICA     1981   11- 10   247:16.2   39.751   122.009   18   29   8*   1.10   FOSTER ISLAND     1981   11-10   1407:42.2   39.752   122.089   14   22   35*   1.10   FOSTER ISLAND     1981   11-10   1407:42.2   39.752   122.089   14   22   35*   1.10   FOSTER ISLAND     1981   11-10   1947: 9.7   39.917   122.743   23   37   0*   0.00   RILEY RIDGE     1981   11-10   1947: 9.7   39.917   122.750   23   37   0*   0.00   RILEY RIDGE     1981   11-12   234:45.4   39.648   122.868   32   51   5*   0.00   MELL RIDGE     1981   11-25   504:37.4   39.554   122.903   34   55   0*   0.00   MELL RIDGE     1981   11-25   504:37.4   39.554   122.904   38   61   5*   0.00   MULL MOUNTAIN     1981   11-25   1027:25.4   39.584   122.966   32   51   5*   0.00   MULL MOUNTAIN     1981   11-25   1027:25.4   39.584   122.966   32   51   5*   0.00   MELL RIDGE     1981   11-26   103:344.5   39.891   122.994   34   55   0*   0.00   MELL RIDGE     1981   11-27   103:36:30.4   39.891   122.994   34   55   0*   0.00   MELL RIDGE     1981   11-30   36:299.3   39.975   122.994   34   55   0*   0.00   MELL RIDGE     1981   12- 1										
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DATE YEAR-MO-DY		LAT.	LON.	DIST MI	DIST KM	DPTH KM	HHI	CRM CD	REF.	COMMENTS
+										
1981 12- 9	558:59.1	39.434	122.696	32	52	9*		2.50		SAINT JOHN HTN.
1981 12- 9					47	13*		1.50		STONYFORD
1981 12- 9	1032:26.3	39.887	122.696	20	32	6*		0.00		RILEY RIDGE
1981 12-10				35	56	1*		1.50		PARADISE SW
1981 12-12				15	24	26*		0.00		NEWVILLE
1981 12-12				52	83	7*		2.10		BANGOR
1981 12-13				22	35	24*		0.00		RILEY RIDGE
1981 12-14				36	58 64	0* 5*		0.00 2.00		PLASKETT RIDGE POTATO HILL
1981 12-15 1981 12-15				40 28	45	5*		1.90		STONYFORD
1981 12-16				33	53	0#		0.00		SAINT JOHN MTN.
1981 12-20				28	45	9*		0.00		RED BLUFF EAST
1981 12-21				16	26	10*		0.00		HALL RIDGE
1981 12-22				53	85	0#		0.00		POTTER VALLEY
1981 12-23				23	37	3*		0.00		BALL MOUNTAIN
1981 12-24				31	50	9*		0.00		YOLLA BOLLY SE
	1831:46.3	39.956	123.179	46	74	5		0.80		BLUENOSE RIDGE
	1536:51.4			17	27	8		30		CHRONE
	523:24.2			34	55	0		1.00		HULL MOUNTAIN
	1448:24.7			35	57	0		1.40		COLD FORK
	1045:52.4			52	83	0		1.30		ELK HOUNTAIN
	2118:33.2			34	54	1		0.80		PLASKETT RIDGE
	249:44.1			27	43	0		0.60		PLASKETT MEADOWS
	1902:48.2			23	37	2		0.70		RILEY RIDGE
	1905:52.3			28	45	10		0.30		YOLLA BOLLY SE
	1510:11.9			40	65 22	0 20		0.50 0.10		LEECH LAKE MTN. FRUTO
	1938: 1.9 644:11.8			14 32	22 52	16		1.10		PRINCETON
1982 2- 6 1982 2- 6	923:28.1			32 25	41	5		0.80		GLENN
	2030:42.9			33	53	<b>25</b>		2.70		PRINCETON
1982 2- 7				53	85	1		1.20		BANGOR
1982 2- 9			121.644	39	63	10		1.30		BUTTE MEADOWS SW
	1524:20.1			50	81	18		3.30		DUBAKELLA MTN SE
	2352:49.8			55	88	0		0.90		BANGOR
1982 2-26	5:55.0	39.452	122.033	30	48	5		1.10		PRINCETON
1982 3- 7	952:13.5	39.412	121.470	54	87	1		1.70		BANGOR
1982 3-7	953:49.9	39.395	121.480	54	87	1		1.50		BANGOR
1982 3- 7				53	86	7		0.50		BANGOR
1982 3-16				41	66	7		1.90		SANHEDRIN MTN.
	1152:46.1			37	59	5	IV	3.20	USGS	
	1152:47.3			37	59	8		2.90		PLASKETT RIDGE
	1206:10.1			38	61	5		2.80	USGS	
	1206:11.6			35	57	9		3.20		PLASKETT RIDGE
	1451:10.5 1210: 2.9			29 12	47 19	8 1		1.70 3.30	USGS	BUCK ROCK
	1210: 2.5			15	24	8		3.20	0363	PASKENTA
1982 3-23				22	35	Ö		1.40		HALL RIDGE
1982 3-23				20	32	ŏ		2.00		HALL RIDGE
	1558: 1.2			19	31	11		1.20		HALL RIDGE
1982 4-10				30	49	16		2.30		PRINCETON
	1646:56.1			42	67	5		0.90		THATCHER RIDGE
	2350:33.8			48	77	12		1.00		OROVILLE
	947:47.8			25	41	20		1.20		GLENN
1982 5- 6	2223:47.3	40.345	122.087	39	63	4		1.50		DALES

DATE YEAR-NO-DY	TIME HRMN:SECS	LAT.	LON.	DIST	DIST KN	DPTH KM	HNI	CRM CD	REF.	COMMENTS
+										
1982 5-10	1015:54.7	39.398	122.880	41	66	٥		1.10		LAKE PILLSBURY
	1058:49.3			40				0.90		CHEROKEE
	1714:33.9			52				1.10		POTTER VALLEY
	1250:56.4			50	80	8		0.60		OROVILLE DAM
	1318: 1.6			53	86	4		1.40		BANGOR
1982 5-15	2226: 0.3	39.862	122.330	3	5	5		1.70		BLACK BUTTE DAM
1982 5-15	2237:38.9	39.755	122.564	13	21	20		10		NEWVILLE
1982 5-17	834:30.8	39.429	121.485	53	85	6		1.00		BANGOR
1982 5-17	837:30.6	39.428	121.490	52	84	13		0.40		BANGOR
1982 5-19	15:20.9	39.622	123.011	39		5		0.50		SANHEDRIN MTN.
1982 5-19	2016:11.0	40.008	122.617	20		8		3.00		LOWREY
1982 5-19	2055:12.2	40.082	122.875	34				1.00		YOLLA BOLLY SE
	403:39.5			20				0.20		RILEY RIDGE
	405: 1.4			20				0.20		RILEY RIDGE
	1214:25.9			24				0.10		RILEY RIDGE
	1616:58.4			32				0.70		KNEECAP RIDGE
	2112:25.4			46				2.10		NO QUADRANGLE FOUND
	800:45.0							1.50		LOWREY
	3 1117:11.8							0.70		BUTTE MEADOWS NE
	1134:42.0							1.70		LOWREY
	700: 8.5							2.90	USGS	
	700: 9.6							3.00		OROVILLE
	702:40.2							3.00		OROVILLE
	926:39.6							1.30		OROVILLE
	7 317: 7.7							0.40		BANGOR
	7 1442:27.8							0.60		OROVILLE DAM
	1604:27.3							2.10		GLENN
	3 1046: 2.4							3.00		LAKE PILLSBURY
	3 1048: 5.7							1.70		POTTER VALLEY NE
	3 1345: 5.4							1.30		POTTER VALLEY NE
	2248:10.6							0.60		CHEROKEE
	5 2343: 7.7							0.20		CHEROKEE
1982 7- 9			121.995					1.50		BUTTE CITY
	1211:25.1							1.30		GLENN
	3 1745:18.8							1.70		POTTER VALLEY NE
	827:53.0							1.40		DALES
	5 211:22.0					1		0.50		YUBA CITY RILEY RIDGE
	8 128:19.1 8 606:55.4							3.10 1.40		RAGLIN RIDGE
1982 7-18	8 2232:11.3							2.30		RILEY RIDGE
	8 2245:22.8							3.60	USGS	NIDE! NIDEL
	8 2245:22.0 8 2245:24.1							3.40	0505	RILEY RIDGE
1982 7-2								1.90		BALL MOUNTAIN
1982 8-								1.10		HULL MOUNTAIN
1982 8~								1.60		HULL MOUNTAIN
1982 8-								0.50		KNEECAP RIDGE
1982 8-								0.60		KNEECAP RIDGE
1982 8-								0.70		KNEECAP RIDGE
1982 8-								3.20		ALDER SPRINGS
1982 8~								0.80		HULL MOUNTAIN
1982 8-					-			0.30		KNEECAP RIDGE
	1 1010:52.3							0.90		KNEECAP RIDGE
1982 8- 1982 8-	1 1010:52.3	, 33.300 ; 39 551	122.894	, 3:				1.10		HULL HOUNTAIN
1982 8-	1 1025: 3.3	39.526	122.95	39		_		1.80		HULL HOUNTAIN

	ONTALOG							_		_
DATE	TIME	LAT.	LON.	DIST	DIST	DPTH	MMI	CRM CD	REF.	CONNENTS
YEAR-MO-DY	HRMN:SECS			MI		KN				
+										
1982 8- 2	8:21.0	39.272	122.983	51	82	0		1.30		ELK MOUNTAIN
1982 8- 3	1424:46.1	39.846	122.693	19	31	0		0.60		HALL RIDGE
	1644:26.4							0.40		BUCK ROCK
	2349:42.7				79	5		1.50		POTTER VALLEY
	511:56.8				25			1.30		FOSTER ISLAND
1982 8- 9	1148:35.1	39.922	122.891	30	49	11		1.00		BUCK ROCK
	102:37.0				99	6		1.10		LOMA RICA
1982 8-17	1036: 4.9	39.503	123.084	45	73	2		0.90		SANHEDRIN NTN.
1982 8-17	2059: 0.1	39.566	122.885	34	54	7		2.30		HULL HOUNTAIN
1982 8-18	1039: 9.9	39.930	122.757	24	38	8		0.50		BALL NOUNTAIN
1982 8-19	125:22.4	39.318	121.398	61	98	12		1.40		LOMA RICA
1982 8-22	134:50.6	39.797	122.969	34	54	1		0.20		MENDOCINO PASS
1982 8-22	525:48.9	39.785	122.893	30	48	0		0.40		NENDOCINO PASS
	720:59.4				49			0.90		MENDOCINO PASS
	1819:48.5				24			0.10		NEWVILLE
1982 8-22					74			1.10		POTTER VALLEY NE
1982 8-23					60			1.10		LEECH LAKE MTN.
	1826:57.8				32			1.50		GLENN
	1054:26.9				82	1		1.10		BRUSHY MTN.
	1101:59.5				73	4		2.70		BRUSHY HTN.
	148:56.7				41			0.50		FELKNER HILL
	433:18.9				58	1		0.90		YOLLA BOLLY SW
	1025:48.0				68	0		0.90		THATCHER RIDGE
	913:23.6				53			0.90		PLASKETT RIDGE
	1153:22.9				54	0		0.20		YOLLA BOLLY SE
1982 8-27	1520:57.9	39.299	121.414		98	5		1.20		LOMA RICA
	1230: 2.2				53	2		0.30		MENDOCINO PASS
1982 8-31	827: 7.9	39.933	122.662	19	31	16		0.70		RILEY RIDGE
1982 8-31	1835:49.3	39.780	122.618	15	24	16		0.40		NEWVILLE
1982 8-31	2101: 5.3	40.122	122.980	40	65	0		1.50		YOLLA BOLLY SW
	122:39.7				30			40	CDWR	RILEY RIDGE
1982 9- 1					56	5*		0.30		NITCHELL GULCH
1982 9- 1	140: 1.2	39.946	122.644	19	30			0.40	CDWR	RILEY RIDGE
	218: 0.9							0.10	CDWR	RILEY RIDGE
	221: 5.6				34	14*		40	CDWR	LOWREY
	436:55.1							0.40		NEWVILLE
	1858:23.8			21	34	5	III	4.00	USGS	
	1905:42.7				38	11*		0.50	CDWR	FELKNER HILL
	2029:42.7				36	11*		0.20	CDWR	FELKNER HILL
	2333:46.4			24	39	11#		10	CDWR	FELKNER HILL
	2355:23.9				37	10#		0.30	CDWR	ELK CREEK
1982 9-4	100: 6.4			20	32	0#		0.20	CDWR	ELK CREEK
1982 9-4	218:19.9				80	0#		1.20	CDWR	ELK MOUNTAIN
1982 9- 4					45	7*		0.50	CDWR	YOLLA BOLLY SE
	1948:25.9			20	32	3*		0.60	CDWR	ELK CREEK
1982 9-5					60	3#		1.20	CDWR	NEWHOUSE RIDGE
	2054:48.6			51	82	1*		0.90	CDWR	ELK HOUNTAIN
1982 9- 7	653:28.1			19	31	7*		1.20	CDWR	ORD FERRY
1982 9- 7				19	31	7*		0.90	CDWR	ORD FERRY
	1251:53.1			41	66	0#		1.10	CDWR	BLACK ROCK NTN SE
	1425:42.5			32	51	6#		0.90	CDWR	HULL HOUNTAIN
	1746:29.6			33	53	0#		1.30	CDWR	HULL MOUNTAIN
	907:56.6			48	78	10#		1.90	CDWR	OROVILLE
1982 9-8	1532: 7.7	39.507	121.511	49	79	10#		1.80	CDWR	OROVILLE

YEAR-NO-DY	TIME HRMN:SECS	LAT.		DIST	DIST KM	DPTH KN	HMI	CRM CD	REF.	COMMENTS
+										
1982 9-10	844: 4.4	39.573	122.578	21	34	13*		0.30	CDWR	ELK CREEK
	651:36.7					_		3.10	CDWR	ONO SE
	739:40.1				_			0.60	CDWR	BUCK ROCK
	744: 2.6			45				1.10	CDWR	BRUSHY MTN.
	746:39.1					_		0.90	CDWR	BRUSHY MTN.
1982 9-16				56		-		1.10	CDWR	UPPER LAKE
	1500:51.6 1533:55.0			19				0.00	CDWR	CHRONE
	1958:10.3			48 29	78 46	5 <b>*</b> 0 <b>*</b>		3.30 0.50	CDWR CDWR	BRUSHY MTM. PLASKETT MEADOWS
	1510: 5.9			25	41	2*		30	CDWR	FELKNER HILL
	6:10.9			9				1.00	CDWR	KIRKWOOD
	131:14.9							0.90	CDWR	PANTHER SPRING SW
	135:18.7			14	22			1.00	CDWR	FOSTER ISLAND
	530:23.8			27	43	10*		40	CDWR	STONYFORD
	1128:20.9			24	39			1.00	CDWR	GLENN
	1815:26.7			30	_			0.60	CDWR	YOLLA BOLLY SE
	204:41.1		_	29		6#		0.10	CDWR	LOG SPRING
	342: 7.5			45				0.20	CDWR	JAMISON RIDGE
	607:15.8							0.60	CDWR	POTTER VALLEY
	801:32.0 1835:31.3							2.70 1.10	CDWR CDWR	FOSTER ISLAND LONA RICA
	1946:33.3							0.50	CDWR	HALL RIDGE
1982 9-21			121.415					1.00	CDWR	LONA RICA
	1744:40.4			24		_		2.50	CDWR	RED BLUFF EAST
•	1219:35.8							1.50	CDWR	FOSTER ISLAND
	2308: 9.1			39				1.00	CDWR	THATCHER RIDGE
1982 9-24	2314:17.7	39.616	123.188	47	76	0#		1.40	CDWR	BRUSHY MTN.
1982 9-25	35:27.9	39.381	123.232	57	91	4#		1.70	CDWR	POTTER VALLEY NW
	719:59.3			16		36#		0.90	CDWR	FOSTER ISLAND
	745:32.6			19		0#		0.00	CDWR	ELK CREEK
	711:13.0			22				0.10	CDWR	ELK CREEK
	126:56.0			23				1.10	CDWR	LOS NOLINOS
	737:18.8			29				0.40	CDWR	RAGLIN RIDGE
	738:44.9							1.50	CDWR	BANGOR
	819: 3.7 1842:32.6					_		1.50	CDWR	CROCKETT PEAK
	2052:17.3			21		25* 5*		80 0.40	CDWR CDWR	HALL RIDGE
	303:25.6			17		28*		0.50	CDWR	ELK CREEK FRUTO
	851:24.8			21	33			10	CDWR	HALL RIDGE
	1410:30.1			41	66	5#		0.50	CDWR	OROVILLE
	1436: 1.9			30	48	17*		0.70	CDWR	SAINT JOHN MTM.
1982 10- 7	216: 1.2	39.549	122.659	25	40	12*		0.00	CDWR	FELKNER HILL
	1712:14.4			16		28#		1.70	CDWR	FOSTER ISLAND
	2231: 8.1			16				1.70	CDWR	FOSTER ISLAND
	332:15.6			27		_		20	CDWR	FELKNER HILL
	307:49.5			24		3*		0.20	CDWR	LOG SPRING
	1345:55.2			29		11*		50	CDWR	SAINT JOHN HTM.
	944:58.6 2131: 9.5			33 13				0.70	CDWR	BUCK ROCK
	1545:59.6			23		35 <b>*</b> 16 <b>*</b>		0.00 1.90	CDWR	CHRONE
	2000:58.9			44				1.90	CDWR CDWR	ELK CREEK SANHEDRIN MTN.
	559:47.0			43				0.90	CDWR	CHANCHELULLA PEAK S
	1202:42.5			46				2.10	CDWR	BUTTE HEADOWS NW
	917:39.8			27		-		20	CDWR	FELKNER HILL

1

	ON 1 N200							_		-
DATE	TIME	LAT.	LON.	DIST	DIST	DPTH	MMI	CRM CD	REF.	COMMENTS
YEAR-MO-DY				MI	KM	KM				
+										
1982 10-28	2342:51.1	39.782	123.036	37	60	0#		1.00	CDWR	NEWHOUSE RIDGE
1982 10-29	1431: 1.0	39.766	121.742	32	51	25*		1.40	CDWR	PARADISE SW
1982 11- 8	356:34.5	40.437	122.861	51	82	5#		1.70	CDWR	CHANCHELULLA PEAK N
1982 11- 8	752:19.9	39.407	121.475	54	87	5#		1.70	CDWR	BANGOR
1982 11- 8					87	5*		1.30	CDWR	BANGOR
1982 11- 8					35	0#		40	CDWR	ELK CREEK
1982 11- 9	10:52.6	39.881	122.737	22	35	0#		0.00	CDWR	RILEY RIDGE
1982 11-10	1041:39.0	39.786	122.962	34	54	0#		0.60	CDWR	MENDOCINO PASS
1982 11-11					28	8*		1.40	CDWR	VINA
1982 11-14	638:36.6	39.885	122.899	30	49	0#		0.20	CDWR	BUCK ROCK
1982 11-14	1324:59.3	39.674	122.651	19	31	5*		0.00	CDWR	ALDER SPRINGS
1982 11-15					21	10		3.10	USGS	
1982 11-15	348:42.2	39.785	122.715	21				3.30	CDWR	HALL RIDGE
1982 11-15								1.00	CDWR	BUTTE CITY
1982 11-15								1.10	CDWR	POTATO HILL
1982 11-19								0.30	CDWR	PLASKETT HEADOWS
1982 11-21								0.50	CDWR	BALL MOUNTAIN
1982 11-21								1.80	CDWR	BUTTE CITY
1982 12- 2						5		3.10	USGS	
1982 12- 2								3.40	CDWR	WILLIAMS
1982 12- 2						9*		2.50	CDWR	WILLIAMS
1982 12- 2						9*		2.60	CDWR	CORTINA CREEK
1982 12- 2								1.10	CDWR	THATCHER RIDGE
1982 12- 2						0#		1.20	CDWR	THATCHER RIDGE
1982 12- 2								2.60	CDWR	WILLIAMS
1982 12- 2								2.30	CDWR	WILLIAMS
1982 12- 4								2.80	CDWR	WILLIAMS
1982 12- 5								3.20	CDWR	WILLIAMS
1982 12- 5						-		3.20	CDWR	CORTINA CREEK
1982 12- 7								1.90	CDWR	BALL HOUNTAIN
1982 12-11						23*		0.50	CDWR	FELKNER HILL
1982 12-15						5#		0.40	CDWR	BALL MOUNTAIN
1982 12-19						0#		0.20	CDWR	ELK CREEK
1982 12-19						1*		1.80	CDWR	JULIAN ROCKS
1982 12-21								0.70	CDWR	NEWHOUSE RIDGE
1982 12-23	1633:26.8	39.565	122.673	25	40			1.10		
1982 12-23				32	51	1#		0.60	CDWR	PLASKETT RIDGE
1982 12-23				19		25*		0.00	CDWR	ELK CREEK
1982 12-28				21		0#		0.20	CDWR	HALL RIDGE
1982 12-28				37	59	18*		1.40	CDWR	HANLIN CANYON
1983 1- 2	723:18.1	39.554	123.064	43	69	0#		1.20	CDWR	SANHEDRIN HTN.
	1827:46.0			43	69	0#		1.20	CDWR	LEECH LAKE HTM.
	1054: 8.9			38	61	0#		1.00	CDWR	THATCHER RIDGE
	1411: 0.8			34	54	3#		0.40	CDWR	BUCK ROCK
	851:27.9			21	34	1*		0.20	CDWR	HALL RIDGE
	1123:57.0			27	44	10#		1.50	CDWR	RICHARDSON SPRINGS
	2339: 5.4			33	53	1*		0.30	CDWR	BUCK ROCK
	725: 3.6			50	80	4#		1.70	CDWR	POTTER VALLEY NE
	850:14.0			38	61	0#		0.80	CDWR	HULL MOUNTAIN
_	1513: 4.2			20	32	0#		0.00	CDWR	ELK CREEK
	1410:59.6			48	77	1*		1.30	CDWR	BRUSHY MTM.
	1436:34.1			30	48	5#		40	CDWR	BUCK ROCK
	1601: 0.3			18	29	35#		1.80	CDWR	HAMILTON CITY
1983 1-16	1927:28.7	39.847	122.722	21	33	5*		0.00	CDWR	HALL RIDGE

Teach   Pop   Harmises   MI	DA'		TIME	LAT.	LON.	DIST	DIST KM	DPTH KN	NNI	CRM CD	REF.	COMMENTS
1983 1-20 1832: 7.3 39.705 122.878 30 48 0° 0.40 CDWR PLASKETT RIDGE 1983 1-31 526:57.6 39.430 121.991 32 52 26° 1.60 CDWR BUTTE CITY 1983 1-31 526:57.6 39.430 121.991 32 52 26° 0.70 CDWR BUTTE CITY 1983 2-16 656:56.4 39.416 121.459 54 87 5° 0.70 CDWR BUTTE CITY 1983 2-17 634:41.1 39.835 122.756 21 33 0° 0.50 CDWR MILL RIDGE 1983 2-17 929:112.5 39.727 122.673 29 47 0° 0.40 CDWR PLASKETT NEADOWS 1983 2-19 1983 1-19 39.727 122.673 29 47 0° 0.40 CDWR PLASKETT NEADOWS 1983 2-19 1983 1-19 39.340 121.463 55 99 5° 1.30 CDWR COVELO EAST 1983 2-19 1213:41.6 39.826 122.500 9 14 30° 3.00 CDWR SHORK CREEK 1983 2-19 213:41.6 39.826 122.500 9 14 30° 3.00 CDWR SHORK CREEK 1983 2-22 145:21.9 39.240 123.049 57 11 4° 2.30 CDWR COW MOUNTAIN 1983 3-5 1641:29.1 39.829 122.812 29 47 5° 0.90 CDWR RIBECAP RIDGE 1983 3-7 8.1311:41.7 39.20 122.812 29 47 5° 0.90 CDWR WWILLE 1983 3-18 550:18.8 40.212 122.155 30 48 9° 2.70 CDWR HULL MOUNTAIN 1983 3-18 550:18.8 40.212 122.155 30 48 9° 2.70 CDWR TUSCAN PRINGS 1983 3-19 602:42.2 39.784 122.059 13 21 16° 1.10 CDWR FUNCAN PRINGS 1983 3-20 242:30.6 39.990 123.044 26 7° 2° 1.40 CDWR LEECH LAKEN TIN. 1983 3-23 27:35.2 39.402 122.155 30 48 9° 2.70 CDWR WILL MOUNTAIN 1983 3-23 27:35.2 39.402 122.155 30 48 9° 2.70 CDWR WILL MOUNTAIN 1983 3-23 27:35.2 39.402 122.156 30 40 6° 2° 1.40 CDWR LEECH LAKEN TIN. 1983 3-23 249:151.8 39.500 121.486 50 80 7° 1.40 CDWR LEECH LAKEN TIN. 1983 3-23 249:151.8 39.500 121.486 50 80 7° 1.40 CDWR LEECH LAKEN TIN. 1983 3-23 249:151.8 39.500 121.486 50 80 7° 1.40 CDWR DWR LEECH LAKEN TIN. 1983 3-24 46:151:3.2 39.935 122.682 37 47 5° 8.00 CDWR DWR LEECH LAKEN TIN. 1983 3-24 61551:3.2 39.935 122.682 39 60 60 60 CDWR DWR LEECH LAKEN TIN. 1983 3-23 47:151.8 39.500 121.486 50 80 7° 1.40 CDWR LEECH LAKEN TIN. 1983 3-22 505:53.9 39.500 121.486 50 80 7° 1.40 CDWR PRINCSTOW LAKEN TIN. 1983 3-24 61551:3.3 39.500 122.800 47 75 5° 80 0.60 CDWR PRINCETOW LAKEN TIN. 1983 3-29 30.200 121.486 50 80 7° 1.40 CDWR PRINCETOW LAKEN TIN. 1983 4-2 1225:542.6 40.390 122.800 47 75 5° 80 0.60 CDWR												
1943   1-24   2000;38.9   39,623   122.600   19   31   10   10   10   10   10   10   11   193   1-3   15   56:655.6   39,430   121.991   32   52   26   1.60   CDWR   BANGOR   1943   2-16   656:56.4   39,416   121.459   54   87   50   0.70   CDWR   BANGOR   1943   2-17   284:137.4   40,028   122.754   27   43   50   0.80   CDWR   COWR   LEKERT   1943   2-17   292:112.5   39,727   122.873   29   47   00   0.40   CDWR   HALL RIDGE   1983   2-17   292:112.5   39,224   121.463   55   89   55   1.00   CDWR   COWELD EAST   1983   2-18   1406:551.2   39,324   121.500   9   14   300   3.00   CDWR   COWELD EAST   1983   2-18   146:12.5   39,224   121.363   57   91   40   2.30   CDWR   COWELD EAST   1983   2-26   452:26.8   39.599   122.812   29   47   50   0.90   CDWR   COWERD REWNITCH   1983   3-5   1641:29.1   39,825   122.812   29   47   50   0.90   CDWR   REWVILLE   1983   3-18   1313:14.7   39,920   122.812   29   47   50   0.90   CDWR   REWVILLE   1983   3-18   550:18.8   40.212   122.099   13   21   160   1.10   CDWR   FOSTER ISLAND   1983   3-19   602:42.2   39,764   122.099   13   21   160   1.10   CDWR   FOSTER ISLAND   1983   3-20   242:30.6   39,990   123.094   42   67   20   1.40   CDWR   FOSTER ISLAND   1983   3-20   242:30.6   39,990   123.094   42   67   20   1.40   CDWR   REWORD REWNITCH   1983   3-23   247:51.6   39,500   21.496   50   60   70   1.40   CDWR   BANGOR   CDWR   HALL RIDGE   1983   3-23   247:51.6   39,500   21.496   50   60   70   70   70   70   70   70   7	<b></b>		~~~~~~									
1983   1-31   526:57.6   39.430   121.991   32   52   26								_				
1983   2-16   656:56.4   39.416   121.459   54   87   5*   0.70   CDWR   BARCOR     1983   2-17   634:41.1   39.435   122.726   21   33   0*   0.50   CDWR     1983   2-17   205:43.4   39.805   122.208   47   75   0*   1.30   CDWR     1983   2-18   1406:51.2   39.346   121.463   55   89   5*   1.00   CDWR     1983   2-19   1213:41.6   39.826   122.500   9   14   30*   3.00   CDWR     1983   2-22   146:21.9   39.210   123.049   57   91   4*   2.30   CDWR     1983   2-22   146:21.9   39.210   123.049   57   91   4*   2.30   CDWR     1983   3-26   452:26.8   39.599   122.812   29   47   5*   0.90   CDWR     1983   3-5   1641:29.1   39.826   122.812   29   47   5*   0.90   CDWR     1983   3-18   550:18.8   40.212   122.105   30   48   9*   2.70   CDWR     1983   3-19   602:42.2   39.784   122.099   13   21   16*   1.10   CDWR     1983   3-20   242:30.6   39.595   122.842   35   56   1*   1.00   CDWR     1983   3-20   250:50:59   39.535   122.842   35   56   1*   1.00   CDWR     1983   3-20   250:50:59   39.535   122.842   35   56   1*   1.00   CDWR     1983   3-20   250:50:59   39.535   122.842   35   56   1*   1.00   CDWR     1983   3-21   327:30.2   39.402   122.105   30   48   9*   2.70   CDWR     1983   3-21   327:30.2   39.402   122.955   47   71   5*   1.30   CDWR     1983   3-23   249:41.7   39.736   122.412   29   47   12*   2.20   CDWR     1983   3-23   249:41.7   39.736   122.412   29   47   12*   2.20   CDWR     1983   3-23   249:41.7   39.736   122.412   29   47   12*   2.20   CDWR     1983   3-23   249:41.7   39.736   122.412   29   47   12*   2.20   CDWR     1983   3-23   249:41.7   39.736   123.042   47   71   4*   1.80   CDWR     1983   3-24   45:38:1   39.506   121.496   50   80   5*   -60   CDWR     1983   3-25   45:39.39   39.555   22.864   39   39   35:35   30.99   30.906												
1983   2-12   824;37.4   40.022   122.754   27   43   58   0.80   CDWR   YOLLA BOLLY SE   1983   2-17   634;41.1   39.835   122.266   21   33   0*   0.50   CDWR   MALL RIDGE   1983   2-18   20:17   20:13:3.727   122.873   29   47   0*   0.40   CDWR   CDWR   CDWELD EAST   1983   2-18   1406;51.2   39.344   121.463   55   89   5*   1.00   CDWR   SERORN CREEK   1983   2-19   1213;41.6   39.826   122.500   9   14   30*   3.00   CDWR   SERORN CREEK   1983   2-22   146;21.9   39.210   123.049   57   91   4*   2.30   CDWR   COWELD EAST   1983   2-22   146;21.9   39.210   123.049   57   91   4*   2.30   CDWR   SERORN CREEK   1983   2-26   452;26.8   39.599   122.812   29   47   5*   0.90   CDWR   KRECAP RIDGE   1983   3-8   1131;14.7   39.920   122.813   27   43   12*   0.20   CDWR   KRECAP RIDGE   1983   3-8   153;14.7   39.920   122.813   27   43   12*   0.20   CDWR   MEWYILLE   1983   3-17   26:20.5   39.595   122.862   35   56   1*   1.00   CDWR   MEWYILLE   1983   3-19   602;42.2   39.784   122.099   13   21   16*   1.10   CDWR   FOSTER ISLAND   1983   3-19   602;42.2   39.784   122.099   13   21   16*   1.10   CDWR   FOSTER ISLAND   1983   3-20   242;30.6   39.990   123.094   42   67   2*   1.40   CDWR   HULL HOUNTAIN   1983   3-23   243;41.7   39.273   122.975   47   11*   1.30   CDWR   HULL HOUNTAIN   1983   3-23   243;41.7   39.273   122.975   47   12*   2.20   CDWR   BLOSSOM   BL												
1983   2-17   634:41.1   39.835   122.726   21   33   00   0.50   CDWR   MALL RIDGE   1983   2-17   292:12.5   39.727   122.873   29   47   00   0.40   CDWR   COVELO EAST   1983   2-18   1406:51.2   39.384   121.463   55   89   58   1.00   CDWR   COVELO EAST   1983   2-19   1213:41.6   39.826   122.500   9   14   30.3   3.00   CDWR   COVELO EAST   1983   2-22   146:21.9   39.210   123.049   57   91   48   2.30   CDWR   COVELO EAST   1983   2-26   452:12.6   39.599   122.812   29   47   58   0.90   CDWR   KENDRY CREEK   1983   3-5   1641:29.1   39.829   122.567   12   20   198   0.40   CDWR   MEWILLE   1983   3-6   1131:147   39.920   122.819   27   43   128   0.20   CDWR   MEWILLE   1983   3-18   550:18.8   40.212   122.105   30   48   98   2.70   CDWR   HULL MOUNTAIN   1983   3-19   602:42.2   39.784   122.099   13   21   168   1.10   CDWR   CDWR   HULL MOUNTAIN   1983   3-20   205:15.39   39.535   122.864   39   62   68   1.30   CDWR   LEECH LAKE MTN.   1983   3-21   327:30.2   39.402   122.975   47   71   58   1.30   CDWR   LAKE PILLSBURY   1983   3-23   297:51.8   39.500   121.496   50   80   78   1.40   CDWR   BANGOR   1983   3-23   249:141.7   39.731   122.772   44   71   48   1.80   CDWR   BANGOR   1983   3-23   249:141.7   39.731   122.772   44   71   48   1.80   CDWR   BANGOR   1983   3-23   249:141.7   39.731   122.772   44   71   48   1.80   CDWR   LEECH LAKE MTN.   1983   3-28   46:18.1   39.776   122.698   19   31   00   0.60   CDWR   LEECH LAKE MTN.   1983   3-28   46:18.1   39.776   122.698   19   31   00   0.60   CDWR   LEECH LAKE MTN.   1983   4-2   35:51.3   39.961   122.608   27   44   11   11   11   11   11   11   1												
1983   2-17   292:12.5   39.727   122.873   29   47   0+   0.40   CDWR   PLASKETT MEADOWS   1983   2-18   1406:51.2   39.384   121.463   55   89   5+   1.00   CDWR   CDWR   CDWR   CDWR   1983   2-19   1213:41.6   39.826   122.500   9   14   30+   3.00   CDWR   SENDRIN CREEK   CDWR   1983   2-22   146:21.9   39.210   123.049   57   91   4+   2.30   CDWR								_				
1983   2-18   1406:51.2   29:384   121:463   55   89   5+   1.00   CDWR   SAMGOR   CREEK   1983   2-19   1213:41.6   39:826   122:500   9   14   30+   3.00   CDWR   SEHORN CREEK   1983   2-22   146:21.9   39:210   123:049   57   91   4+   2.30   CDWR								0#				
1983 2-19 1213:41.6 39.826 122.500 9 14 30* 3.00 CDWR SEHORN CREEK 1983 2-22 146:21.9 39.210 123.049 57 91 4* 2.30 CDWR COW NOUNTAIN 1983 2-26 452:26.8 39.599 122.812 29 47 5* 0.90 CDWR KNEECAP RIDGE 1983 3-5 1641:29.1 39.829 122.567 12 20 19* 0.40 CDWR KNEECAP RIDGE 1983 3-7 8 1131:14.7 39.920 122.819 27 43 12* 0.20 CDWR BALL HOUNTAIN 1983 3-17 28:20.5 39.535 122.882 35 56 1* 1.00 CDWR BALL HOUNTAIN 1983 3-18 550:18.8 40.212 122.105 30 48 9* 2.70 CDWR HULL HOUNTAIN 1983 3-19 602:42.2 39.784 122.099 13 21 16* 1.10 CDWR FOSTER ISLAND 1983 3-20 242:30.6 39.990 123.084 42 67 2* 1.40 CDWR LEECH LAKE HTM. 1983 3-21 327:30.2 39.402 122.975 44 71 5* 1.30 CDWR LAKE PILLSBURY 1983 3-23 129: 31 40.230 122.412 29 47 12* 2.20 CDWR BANGGR 1983 3-23 247:51.8 39.500 121.496 50 80 7* 1.40 CDWR BANGGR 1983 3-23 249:41.7 39.273 122.772 44 71 4* 1.80 CDWR BANGGR 1983 3-26 1551: 3.2 39.935 123.090 41 66 2* 0.60 CDWR BANGGR 1983 3-26 46:83.1 39.776 122.698 19 31 0* 0.60 CDWR HALL RIDGE 1983 3-0 227:12.3 39.624 122.603 19 31 0* 0.60 CDWR HALL RIDGE 1983 4-3 355:33.0 39.975 122.842 29 47 5 3.10 USCS 1993 4-7 355:33.0 39.975 122.842 29 47 5 3.10 USCS 1993 4-7 355:33.0 39.975 122.842 29 47 5 3.10 USCS 1993 4-7 355:33.0 39.975 122.842 29 47 5 3.10 USCS 1993 4-7 355:33.0 39.975 122.842 29 47 5 3.10 USCS 1993 4-7 355:33.0 39.975 122.842 29 47 5 3.10 USCS 1993 4-7 355:33.0 39.975 122.842 29 47 5 3.10 USCS 1993 4-7 355:33.0 39.975 122.842 29 47 5 3.10 USCS 1993 4-7 355:33.0 39.975 122.842 29 47 5 3.10 USCS 1993 4-7 355:33.0 39.975 122.842 29 47 5 3.10 USCS 1993 4-7 30:13.10 122.806 27 44 11* 3.10 CDWR PRINCETON 1993 4-7 30:13.0 39.690 123.207 47 76 5* 0.80 CDWR PRINCETON 1993 4-7 10:15.50.3 39.119 122.807 56 90 15* 3.70 CDWR PRINCETON 1993 4-21 90:158.1 39.934 122.774 40 65 5* 2.70 CDWR PRINCETON 1993 4-21 90:158.1 39.934 122.776 40 65 5* 2.70 CDWR PRINCETON 1993 4-22 2055:42.6 40.390 121.996 50 75 5 5 0.00 CDWR PRINCETON 1993 4-22 2055:42.6 40.390 121.996 50 75 5 5 0.00 CDWR PRINCETON 1993 5-29 121:51.7.3 39.694 122.907 30 49 5* 0.40 CDWR	1983	2-17	2205:43.4	39.809	123.208	47	75	0#		1.30	CDWR	COVELO EAST
1983   2-22   146:21.9   39.210   123.049   57   91   4+   2.30   CDWR   COW MOUNTAIN   1983   2-26   452:26.8   39.599   122.812   29   47   5*   0.90   CDWR   KRECGAP RIDGE   1983   3-5   1641:29.1   39.829   122.567   12   20   19*   0.40   CDWR   MEWVILLE   1983   3-17   28:20.5   39.595   122.862   35   56   1*   1.00   CDWR   MEWVILLE   1983   3-17   28:20.5   39.595   122.862   35   56   1*   1.00   CDWR   MULL MOUNTAIN   1983   3-18   550:18.8   40.212   122.105   30   48   9*   2.70   CDWR   TUSCAN SPRINGS   1983   3-20   202:42:20.6   39.990   123.084   42   67   2*   1.40   CDWR   LEECH LAKE NTN.   1983   3-20   202:30.6   39.990   123.084   42   67   2*   1.40   CDWR   LEECH LAKE NTN.   1983   3-20   30:553.0   39.535   122.964   39   62   6*   1.30   CDWR   LEECH LAKE NTN.   1983   3-23   129: 3.1   40.230   122.475   47   75   1.30   CDWR   BLOSSON   1983   3-23   249:41.7   39.273   122.772   44   71   4*   1.80   CDWR   BLOSSON   1983   3-23   249:41.7   39.273   122.772   44   71   4*   1.80   CDWR   DOTATO HILL   1983   3-27   1651:30.9   39.550   121.496   50   80   5*   -60   CDWR   CDWR   CDWR   LEECH LAKE NTN.   1983   3-23   249:41.7   39.273   122.772   44   71   4*   1.80   CDWR   DOTATO HILL   1983   3-23   249:41.7   39.273   122.772   44   71   4*   1.80   CDWR   DOTATO HILL   1983   3-23   249:41.7   39.273   122.698   19   31   0*   -60   CDWR   CDWR   CDWR   LEECH LAKE NTN.   1983   3-30   2227:12.3   39.624   122.608   19   31   0*   -60   CDWR   EECH LAKE NTN.   1983   4-3   355:33.0   39.975   122.842   29   47   5   5   0.80   CDWR   EECH LAKE NTN.   1983   4-7   304:13.0   39.650   122.806   29   46   4*   2.30   CDWR   DOTATO HILL   1983   4-7   304:13.0   39.650   122.807   47   75   5*   1.60   CDWR   LAKEPORT   1983   4-7   304:13.0   39.650   122.807   5*   5*   1.60   CDWR   LAKEPORT   1983   4-21   901:58.1   39.396   122.806   39.50   123.100   47   75   5*   1.60   CDWR   EECH LAKE NTN.   1983   4-22   905:542.6   40.390   122.807   5*   48   49   40.60   CDWR   EE		2-18	1406:51.2	39.384	121.463	55	89	5*		1.00	CDWR	BANGOR
1983 2-26 452:26.8 39.599 122.812 29 47 5* 0.90 CDWR MRECAP RIDGE 1983 3-5 1611:29.1 39.829 122.857 12 20 19* 0.40 CDWR MEWILLE 1983 3-8 1131:14.7 39.920 122.819 7 43 12* 0.20 CDWR BALL HOWITAIN 1983 3-17 28:20.5 39.535 122.862 35 56 1* 1.00 CDWR HULL HOWITAIN 1983 3-18 550:18.8 40.212 122.105 30 48 9* 2.70 CDWR TUSCAM SPRINGS 1983 3-19 602:42.2 39.784 122.099 13 21 16* 1.10 CDWR FOSTER ISLAND 1983 3-20 505:539.3 93.535 122.964 39 62 6* 1.30 CDWR HULL HOWITAIN 1983 3-20 505:539.3 93.535 122.964 39 62 6* 1.30 CDWR LEECH LAKE HTM. 1983 3-21 327:30.2 39.402 122.975 44 71 5* 1.30 CDWR LAKE PILLSBURY 1983 3-23 129: 3.1 40.230 122.412 29 47 12* 2.20 CDWR BANGOR 1983 3-23 247:51.8 39.500 121.496 50 80 7* 1.40 CDWR BANGOR POTATO HILL 1983 3-26 1551: 3.2 39.935 123.090 41 66 2* 0.60 CDWR DATO HILL 1983 3-26 1551: 3.2 39.935 123.090 41 66 2* 0.60 CDWR DATO HILL 1983 3-26 1551: 3.2 39.935 122.968 19 31 0* 0.60 CDWR HALK RIDGE ELECK LAKE HTM. 1983 3-20 227:12.3 39.624 122.603 19 31 0* 0.60 CDWR CROVILLE DAK HILL 1983 3-30 227:12.3 39.624 122.603 19 31 0* 0.60 CDWR CROVILLE DAK HILL 1983 4-5 1252: 5.4 39.463 122.056 29 46 4* 2.30 CDWR POTATO HILL 1983 4-7 735:50.3 39.119 122.808 27 44 11* 3.10 CDWR POTATO HILL 1983 4-7 735:50.3 39.119 122.808 27 44 11* 3.10 CDWR POTATO HILL 1983 4-15 408:56.2 39.647 122.508 27 44 11* 3.10 CDWR POTATO HILL 1983 4-2 50.55142.7 40.491 121.942 51 82 23* 3.10 CDWR POTATO HILL 1983 4-22 2055:42.7 40.491 121.942 51 82 23* 3.10 CDWR FOSTER ISLAND 1983 4-2 1015:50.3 39.761 122.2056 29 46 4* 2.30 CDWR POTATO HILL 1983 4-22 1015:50.3 39.761 122.2056 29 46 4* 2.30 CDWR POTATO HILL 1983 4-22 1015:50.3 39.761 122.808 27 44 11* 1.30 CDWR POTATO HILL 1983 4-22 1015:50.3 39.761 122.808 27 44 11* 1.30 CDWR POTATO HILL 1983 4-22 1015:50.3 39.761 122.807 30 8.30 8.30 8.30 8.30 8.30 8.30 8.30 8						9	_	30*			CDWR	SEHORN CREEK
1983 3-5 1641:29.1 39.829 122.567 12 20 19* 0.40 CDWR MEWVILLE 1983 3-8 1131:14.7 39.920 122.819 27 43 12* 0.20 CDWR MILL MOUNTAIN 1983 3-17 28:20.5 39.595 122.822 35 56 1* 1.00 CDWR MULL MOUNTAIN 1983 3-18 550:18.8 40.212 122.105 30 48 9* 2.70 CDWR TUSCAN SPRINGS 1983 3-19 602:42.2 39.784 122.099 13 21 16* 1.10 CDWR FOSTER ISLAND 1983 3-20 242:30.6 39.990 123.084 42 67 2* 1.40 CDWR FOSTER ISLAND 1983 3-20 505:53.9 39.535 122.964 39 62 6* 1.30 CDWR HULL MOUNTAIN 1983 3-21 327:30.2 39.402 122.975 44 71 5* 1.30 CDWR HULL MOUNTAIN 1983 3-23 129: 3.1 40.230 122.975 44 71 5* 1.30 CDWR HULL MOUNTAIN 1983 3-23 249:41.7 39.273 122.772 44 71 2* 2.20 CDWR BANGGR 1993 3-23 249:41.7 39.273 122.772 44 71 4* 1.80 CDWR BANGGR 1983 3-26 46:38.1 39.76 122.698 19 31 0*60 CDWR BANGGR 1983 3-28 46:38.1 39.76 122.698 19 31 0*60 CDWR HULL MOUNTAIN 1983 3-28 46:38.1 39.76 122.698 19 31 0*60 CDWR HULL MOUNTAIN 1983 4- 1 537: 7.4 39.410 123.042 47 75 5* 0.80 CDWR ELK CREEK 1983 4- 3 355:33.3 39.975 122.842 29 47 5 5* 3.10 USCS 1983 4- 7 735:50.3 39.119 122.807 75 5* 0.80 CDWR POTTER VALLEY NE 1983 4- 7 304:13.0 39.650 123.207 47 76 5* 1.60 CDWR BALL MOUNTAIN 1983 4- 7 304:13.0 39.650 123.207 47 76 5* 1.60 CDWR BALL MOUNTAIN 1983 4- 7 304:13.0 39.650 123.207 47 76 5* 1.60 CDWR PRINCETON 1983 4- 7 304:30.3 39.951 122.802 74 47 11 1* 1.30 CDWR LAKEPORT 1983 4- 1 5408:56.2 39.647 122.583 17 28 0* 0.00 CDWR PRINCETON 1983 4- 2 2055:42.6 40.390 121.996 43 70 5 3.50 USGS 1983 4- 22 2055:42.6 40.390 121.996 43 70 5 3.50 USGS 1983 4- 22 1015:50.3 39.761 122.802 30 48 0* 0.60 CDWR PRINCETON NW 1983 4- 22 1015:50.3 39.761 122.802 30 48 0* 0.60 CDWR PRINCETON NW 1983 4- 22 1015:50.3 39.761 122.903 30 49 5* 0.40 CDWR RUSECAP RIDGE 1983 4- 22 1015:50.3 39.761 122.905 30 48 0* 0.60 CDWR BALL MOUNTAIN 1983 4- 22 1015:50.3 39.584 122.776 40 65 5* 0.40 CDWR RUSECAP RIDGE 1983 5- 9 115:17.3 39.564 122.976 40 65 5* 0.40 CDWR RUSECAP RIDGE 1983 5- 9 127:9.5 39.404 121.459 54 87 4* 1.40 CDWR BANGOR 1983 5- 9 127:9.5 39.404 121.459 54 87 4* 1.40 CDW												
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	1983	6-11	239: 3.6	39.563	121.984	25	41	9*		2.20	CDWR	LLANO SECO

DAT	Έ	TIME	LAT.	LON.	DIST		DPTH	MMI CRM CD	REF.	COMMENTS
YEAR-M	O-DY	HRMN: SECS			MI	KM	KM			
+										
1983	6-11	1645:40.8	39.744	121.958	21	33	0#	2.20	CDWR	ORD FERRY
1933	6-15	1138:59.6	40.092	122.783	30	49	5#	0.80	CDWR	YOLLA BOLLY SE
1983	6-21	1612:39.5	39.692	122.871	30	48	0#	0.50	CDWR	PLASKETT HEADOWS
1983	6-23	900:58.6	39.282	122.076	39	63	24#	3.10	CDWR	NOULTON WEIR
1983	6-28	227:45.8	39.918	121.406	50	80	5*	1.30	CDWR	PULGA NW
1983	6-29	126: 0.7	39.939	122.872	30	48	5#	1.70	CDWR	BALL MOUNTAIN
1983	6-30	812:25.7	39.943	122.885	30	49	5#	0.50	CDWR	BUCK ROCK
		1452:14.0			27	44	11*	1.80	CDWR	LLANO SECO
		1731:54.7			22	36	1#	0.30	CDWR	RILEY RIDGE
		46:25.7			34	55	0#	1.20	CDWR	BUCK ROCK
		445: 0.7			48	78	5#	1.40	CDWR	BLACK ROCK MTN NE
		858:36.6			31	50	9#	1.00	CDWR	SAINT JOHN HTM.
		947: 2.9			30	48	8#	0.40	CDWR	FELKNER HILL
		1607:53.2			25	41	30*	1.10	CDWR	LLANO SECO
		2234:51.4			42	67	8*	1.60	CDWR	LAKE PILLSBURY
		110:22.6			37	59	2*	0.90	CDWR	BUCK ROCK
	7-11	209:18.1			19	31	5# 1#	0.90	CDWR	ELK CREEK
		1610:38.5			25	41 41	5*	0.10 0.50	CDWR	PLASKETT MEADOWS PLASKETT MEADOWS
		1449:56.2			25	41		0.30	CDWR	
		1450:32.4			25		5#		CDWR	PLASKETT MEADOWS CROCKETT PEAK
		1500:47.5			36	58 27	1#	1.10	CDWR	FOSTER ISLAND
		1110:36.0			17	27 47	26# 0#	1.50	CDWR	
		1722:13.5 933:57.9			29 35	57	1*	0.30 0.80	CDWR CDWR	PLASKETT HEADOWS PLASKETT RIDGE
		2316:51.7			48	77	0#	1.10	CDWR	COVELO EAST
		1540:28.0			17	27	23*	0.10	CDWR	PASKENTA
		1545:50.9			13	21	27*	2.00	CDWR	PASKENTA
		1745:54.1			28	45	0#	0.70	CDWR	KNEECAP RIDGE
		2215:57.0			21	33	31*	0.00	CDWR	RILEY RIDGE
		2311:48.8			36	58	5#	1.10	CDWR	LODOGA SW
		122: 4.6			34	54	0#	0.20	CDWR	HULL MOUNTAIN
		536:27.0			16	25	34*	1.80	CDWR	PASKENTA
		1156:14.5			27	44	6#	0.40	CDWR	PLASKETT NEADOWS
		1546:19.8			32	52	22*	1.40	CDWR	NELSON
		1416:23.7			16	25	28*	1.90	CDWR	HAMILTON CITY
		2021:47.6			32	51	0#	0.70	CDWR	PLASKETT RIDGE
		1818:49.9			19	31	6#	0.50	CDWR	CHRONE
		1129:55.9			19	30	1#	60	CDWR	CHRONE
		1314:17.3			39	63	0#	1.10	CDWR	HULL MOUNTAIN
		1441:37.0			50	80	0#	1.10	CDWR	PULGA SW
	8-17	452:38.9	39.689	122.766	24	39	0#	0.30	CDWR	PLASKETT NEADOWS
1983	8-17	624:30.3	40.381	122.114	41	66	9#	1.50	CDWR	NO QUADRANGLE FOUNT
	8-18	722:28.3			50	80	0#	1.20	CDWR	POTTER VALLEY NE
1983	8-21	1950:26.5	40.278	122.902	44	71	0#	1.40	CDWR	CHANCHELULLA PEAK S
1983	8-22	1402:34.0	39.970	122.764	25	41	0*	0.90	CDWR	BALL HOUNTAIN
		1607: 3.4			42	68	12*	2.10	CDWR	BLACK ROCK NTN SE
1983	8-26	1815:17.0			19	30	19*	0.60	CDWR	HAMILTON CITY
	8-28	503:51.3			33	53	3*	1.30	CDWR	HAMLIN CANYON
	9- 1	339:27.3			26	42	1*	0.10	CDWR	PLASKETT HEADOWS
1983	9- 3	1126:45.1			21	34	1#	0.20	CDWR	HALL RIDGE
	9- 4	155: 4.2			22	36	1#	0.20	CDWR	ALDER SPRINGS
		1920:56.8			24	39	6*	1.40	CDWR	NORD
		2042: 3.9			21	34	0#	0.50	CDWR	FELKNER HILL
1983	9-14	822:30.5	39.695	122.816	27	43	0*	0.70	CDWR	PLASKETT NEADOWS

YEAR-MO-DY				MI	KN	KN	HMI	CRM CD	REF.	COMMENTS
+										
1983 9-14	1013: 0.8	39.751	121.949	21	34	7*		1.30	CDWR	NORD
1983 9-18					77	2*		1.90	CDWR	ELK MOUNTAIN
1983 9-19								0.80	CDWR	FRUTO NE
1983 9-21								0.40	CDWR	PLASKETT HEADOWS
1983 9-23								30	CDWR	HALL RIDGE
1983 9-24								1.30	CDWR	BUCK ROCK
1983 10- 2								0.60	CDWR	PLASKETT NEADOWS
1983 10- 2								0.30	CDWR	PLASKETT HEADOWS
1983 10- 2								0.10	CDWR	ALDER SPRINGS
1983 10- 5								1.40	CDWR	POTTER VALLEY NE
1983 10- 5								1.70	CDWR	PALERMO
1983 10- 7								0.90	CDWR	PLASKETT RIDGE LOG SPRING
1983 10- 8						0# 10#		0.30	CDWR CDWR	SEHORN CREEK
1983 10- 8					_			1.50	CDWR	POTTER VALLEY
1983 10-12 1983 10-17						-		2.30	CDWR	POTTER VALLEY NW
1983 10-17								1.10	CDWR	LAKE PILLSBURY
1983 10-19								10	CDWR	HULL MOUNTAIN
1983 10-22								1.20	CDWR	POTTER VALLEY
1983 10-26						_		2.10	CDWR	YOLLA BOLLY SE
1983 10-26								0.30	CDWR	RILEY RIDGE
1983 10-26								0.80	CDWR	PALERNO
1983 10-28						0#		1.00	CDWR	YOLLA BOLLY SW
1983 10-31						4#		0.70	CDWR	PARADISE SE
1983 11- 1					82	7#		1.50	CDWR	BANGOR
1983 11- 7	1758:29.8	39.480	121.505	50	80	6#		1.30	CDWR	PALERNO
1983 11- 8	144:49.6	39.274	121.464	60	96	15#		1.60	CDWR	LOMA RICA
1983 11- 8	2236:48.3	40.390	123.170	60	96	5*		2.20	CDWR	DUBAKELLA MTN NW
1983 11-13	2210:25.3	39.960	122.835	29	46	3*		1.50	CDWR	BALL MOUNTAIN
1983 11-17					67	4#		2.00	CDWR	NEWHOUSE RIDGE
1983 11-17	2310:49.5	39.929	122.918	32	51	5*		0.90	CDWR	BUCK ROCK
1983 11-18								1.60	CDWR	POTTER VALLEY NE
1983 12- &								0.50	CDWR	ALDER SPRINGS
1983 12-13								1.90	CDWR	BLUENOSE RIDGE
1983 12-15								1.50	CDWR	RICHARDSON SPGS NW
	1543:45.8							0.10	CDWR	MENDOCINO PASS
	101: 7.7							1.00	CDWR	ONO NE
	2205:20.1							10 2.30	CDWR CDWR	LOG SPRING WEST OF GERBER
								20	CDWR	CHRONE
						_		0.10	CDWR	PASKENTA
1984 1-16 1984 1-18								0.50	CDWR	LOG SPRING
	1227:26.5							1.60	CDWR	HAMILTON CITY
1984 1-29								0.40	CDWR	BERRY CREEK
	2015:58.7					-		0.90	CDWR	THATCHER RIDGE
	2136:54.7							1.50	CDWR	JAMISON RIDGE
	2144: 8.4							0.00	CDWR	OROVILLE DAM
	843:24.3							0.00	CDWR	LOG SPRING
	1225:44.4							0.30	CDWR	BALL HOUNTAIN
	1440:29.6				32	12*		1.70	CDWR	HALL RIDGE
	1606:58.5				3 (	) 3*		0.50	CDWR	ALDER SPRINGS
	2124:55.3				5 50	0+		1.50	CDWR	CROCKETT PEAK
1984 2-11					8:	1#		1.70	CDWR	POTTER VALLEY NW
1984 2-14	1359:13.6	40.149	122.90	7 36	6	l 1*		1.40	CDWR	YOLLA BOLLY NW

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DA' VFAR-		TIME HRMN:SECS	LAT.	LON.	DIST MI	DIST KM	DPTH KM	MMI	CRM CD	REF.	COMMENTS
+											
1984		1639:55.7	39.519	122.757	30	49	5#		10	CDWR	KNEECAP RIDGE
1984		1748:59.6			32	52	3*		0.60	CDWR	KNEECAP RIDGE
1984		2249:17.0			26	42	1*		0.30	CDWR	FELKNER HILL
1984		2004: 4.8			37	60	4*		1.00	CDWR	YOLLA BOLLY SW
1984	5-16	136:20.1			32	52	4*		0.40	CDWR	KNEECAP RIDGE
1984		626:34.7			29	46	0#		20	CDWR	PLASKETT MEADOWS
1984	5-16	1552:58.8	40.117	122.749	30	49	27*		1.80	CDWR	RAGLIN RIDGE
1984	5-16	2303:19.2	39.972	121.879	27	43	1*		1.40	CDWR	RICHARDSON SPGS NW
1984	5-17	2100:57.3	40.019	121.928	26	42	8#		1.80	CDWR	PANTHER SPRING SW
1984		900:53.7			35	56	4*		0.00	CDWR	HULL MOUNTAIN
1984	5-19	2139: 4.4	39.607	123.102	43	70	5*		1.00	CDWR	SANHEDRIN MTM.
1984	5-20	255:18.3	39.895	123.238	48	78	1*		1.60	CDWR	BLUENOSE RIDGE
1984	5-22	1618:20.2			29	46	5*		0.10	CDWR	PLASKETT MEADOWS
1984	5-26	500:55.5			35	57	9*		1.60	CDWR	CROCKETT PEAK
1984		2304:19.0			17	27	26*		1.60	CDWR	FOSTER ISLAND
1984	5-28	2318:38.3	39.731	122.024	17	28	12*		1.60	CDWR	HAMILTON CITY
1984	6- 2	2222: 8.4	39.294	122.982	50		3*		2.20	CDWR	ELK HOUNTAIN
1984	6- 3	235:21.8	39.668	122.764			_		0.30	CDWR	PLASKETT MEADOWS
1984		2355:36.7							1.60	CDWR	BANGOR
1984		1409:38.0							1.20	CDWR	PULGA NW
1984		630:50.0					_		1.70	CDWR	OXBOW BRIDGE
1984		1845:55.9							1.90	CDWR	POTTER VALLEY NE
1984		1915:57.1					_		1.70	CDWR	POTTER VALLEY NE
1984		1056: 3.1					_		1.00	CDWR	BUCK ROCK
1984		658:30.5							0.70	CDWR	OROVILLE
1984		1809:38.9							1.80	CDWR	BROWNS VALLEY
1984		1154:45.5							0.10	CDWR	LOG SPRING
1984		1938: 2.1							20	CDWR	NEWVILLE
1984		1943:42.6							20	CDWR	HALL RIDGE
1984		2341:39.0							0.90	CDWR	FRUTO
1984		125:18.6					_		1.00	CDWR	BUCK ROCK
1984		1331:14.0					_		0.50	CDWR	LOG SPRING
1984		1348:24.2					-		1.90	CDWR	TUSCAN SPRINGS
1984		634:58.5							2.00	CDWR	FOUTS SPRINGS
1984		1743:13.8							0.40	CDWR	PLASKETT HEADOWS
1984		1231:59.4							0.90 1.60	CDWR	ALDER SPRINGS RICHARDSON SPGS NW
		953:25.0					17*			CDWR CDWR	HULL MOUNTAIN
1984	7-25	620:30.2	40 023	121 925	34 26		_		0.50 1.60	CDWR	PANTHER SPRING SW
1984 1984		1728:21.2							1.80	CDWR	PANTHER SPRING SW
1984		2127:25.2					_		0.50	CDWR	PLASKETT MEADOWS
1984	7-31						-		1.50	CDWR	POTATO HILL
1984		1705:58.5							0.90	CDWR	COLD FORK
1984	8- 6								1.80	CDWR	PALERNO
1984	8-6								1.10	CDWR	LAKE PILLSBURY
	8- 9						-		0.10	CDWR	ELK CREEK
1984	8-10								1.00	CDWR	YOLLA BOLLY NE
1984		1123:50.0							0.90	CDWR	LAKE PILLSBURY
1984 1 <b>98</b> 4		1123:50.6					_		1.50	CDWR	REDWOOD VALLEY
1984	8-18								0.00	CDWR	LOG SPRING
1984		1659:32.1							1.10	CDWR	HAMILTON CITY
1984		1659:52.1 1141:58.6							30	CDWR	ALDER SPRINGS
1984		5 1957: 9.0							0.80	CDWR	YOLLA BOLLY SE
1984		1239:37.7							0.60	CDWR	FELKNER HILL
. 704	~						•			-	

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		TIME HRMN:SECS		LON.		DIST KM	DPTH KM	MMI	CRM CD	REF.	COMMENTS
+											
1984	8-28	1253: 0.7	39.476	122.792	34	54	2*		2.30	CDWR	CROCKETT PEAK
1984	8-29	618: 2.5	40.016	122.567	19	30	4#		0.30	CDWR	LOWREY
1984	8-30	58: 6.2	39.690	122.938	33	53	0#		0.70	CDWR	PLASKETT RIDGE
1984	8-30	1054:11.6	39.519	122.745	30	48	0#		0.50	CDWR	FELKNER HILL
1984	8-30	1116:57.1	39.994	122.592	19	30	7#		0.90	CDWR	PASKENTA
1984	9- 4	1835:53.8	39.418	121.386	58	93	5#		0.90	CDWR	BANGOR
1984	9- 5	1114: 1.1	40.032	122.631	22	35	5#		1.70	CDWR	RAGLIN RIDGE
1984	9- 5	1131:59.5	40.013	122.798	28	45	1#		0.60	CDWR	YOLLA BOLLY SE
1984	9- 5	1301:10.2	39.978	122.731	24	38	5#		0.00	CDWR	RILEY RIDGE
1984	9- 5	1511:46.2	39.445	123.119	49	79	2*		1.00	CDWR	POTTER VALLEY NE
1984	9- 7	753:17.8	40.209	122.117	30	48	8*		2.10	CDWR	TUSCAN SPRINGS
1984	9- 8	9:15.8	40.208	122.123	29	47	11*		1.90	CDWR	TUSCAN SPRINGS
1984		417: 9.7				44	8*		1.60	CDWR	RED BLUFF EAST
1984	9- 9	603:18.2	39.761	122.702	20	32	0#		0.00	CDWR	HALL RIDGE
1984	9-13	1539:56.0	39.954	122.867	30	48	5#		0.90	CDWR	BALL HOUNTAIN
1984	9-13	2135:35.1	39 <b>.9</b> 97	121.897	27	43	11*		2.60	CDWR	RICHARDSON SPGS NW
1984	9-15	2257:52.7	40.020	121.920	26	42	9*		1.80	CDWR	PANTHER SPRING SW
1984	9-18	305:37.8	39.366	123.084	50	81	1*		1.20	CDWR	POTTER VALLEY
1984	9-19	236:50.3	39.644	123.044	40	64	0#		0.90	CDWR	THATCHER RIDGE
1984	9-21	1217:43.5	39.775	122.903	30	49	5*		0.90	CDWR	MENDOCINO PASS
1984	9-24	125:43.3	40.036	122.078	21	33	36*		2.50	CDWR	LOS NOLINOS
1984	9-25	1713:13.5	40.132	122.308	22	35	25*		2.70	CDWR	RED BLUFF WEST
1984	9-25	2151:46.5	39.665	123.013	37	60	0#		0.80	CDWR	THATCHER RIDGE
1984	9-29	318:29.3	39.505	122.057	26	42	15*		0.80	CDWR	GLENN

APPENDIX B

**STRATIGRAPHY** 

## APPENDIX B

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#### STRATIGRAPHY

General Introduction. Stratigraphy in the study area represents deposition on an emerging continental margin. Succession began with deep marine turbidites (Great Valley (GV) Sequence) deposited in an elongated forearc basin lying between a subduction zone and parallel volcanic chain. After synclinal collapse and the accretion of the basin onto the continental margin, marine sedimentation dwindled and subaerial deposition began (Dickinson, 1981; Hackel, 1966). The sequence then is oceanic and continental basement overlain by assemblages of lithologically heterogenous metasedimentary, metavolcanic, and ophiolitic rock (Franciscan Complex) juxtaposed against marine sedimentary rock of simple stratigraphic order (Great Valley Sequence). These suprabasement stratigraphies are in turn overlain by marine shelf and lower submarine canyon fill deposits (Tertiary marine sequence) and upper canyon deposits. These consist of pelites, sandstone, and conglomerates of upper valley mixed load rivers laying down meander belt and flood plain deposits (Tuscam-Tehama and Upper Princeton fill).

Basement Complex. From west to east across northern California the basement rocks consist of the principally metamorphic Franciscan Complex, an ophiolite sequence, the sedimentary marine (turbidites) Great Valley Sequence, and the granitic Sierra Nevada Batholith. For the most part these groups are coeval, representing different terrains collectively formed during Late Mesozoic subduction.

Franciscan Complex. The Franciscan Complex forms the core of the California Coast Ranges west of the project area. These deep marine sediments, mafic marine volcanic, and metamorphic rocks have been tectonically scraped from a subducting plate and dragged downward. This has resulted in a melange of chaotic units represented by graywacke, siltstone, shale, volcanic rocks, radiolarian chert, and serpentinite, much of which has been metamorphosed to blueschist facies by high pressure-low temperature (Bailey and others, 1964). The Franciscan Complex is principally Late Jurassic and Cretaceous in age making it coeval with the Great Valley Sequence. However the Franciscan lies structurally below the Ophiolite and Great Valley Sequence. The Franciscan Complex has been tectronically thrust eastward beneath the Ophiolite Complex and Great Valley Sequence by imbricate thrust faulting (Maxwell, 1974; Raney, 1976; Suppe, 1979).

Ophiolite Complex. The Ophiolite Complex is believed to be Jurassic oceanic crust upon which the Great Valley Sequence was deposited. The rock comprising the Ophiolite Complex is best described as a serpentinite melange. It is a chaotic complex of volcanic breccia, bedded radiolarian chert, pillow basalt, variable sized blocks of basalt commonly altered to greenstone, and highly weathered and sheared serpentinite. The serpentinite is the most abundant rock type, is highly weathered and sheared, and contains mudstone-graywacke blocks (Fritz, 1975). Although the Great Valley Sequence is believed to be deposited directly on the Ophiolite, the contact is nearly everywhere faulted along the Stony Creek fault.

Sierra Nevada Batholith. Most of the Sierra Nevada batholith is composed of numerous plutons of granodioritic composition that were emplaced between Middle Triassic and Middle Cretaceous times into older Paleozoic marine sedimentary strata. These rocks are bordered along the western side of the Sierra by a belt of deformed metamorphic rocks. The Sierra Nevada was the principal source of detritus for the Great Valley Sequence. Batholithic or related rocks of the Sierra Nevada may form much of the crystalline basement at great depth beneath the Sacramento Valley.

Great Valley Sequence. The Mesozoic rocks of the Great Valley Sequence are generally exposed in a long band along the western margin of the Great Valley. The sequence is composed of interbedded and intertonguing mudstone, sandstone, and conglomerate. The sequence is up to 24,000 feet thick in the area of the project. Rocks of the Great Valley Sequence are coeval with, but structurally overlie, those of the Franciscan Complex. In contrast to the Franciscan Complex, the Great Valley Sequence consists of clastic sedimentary rocks that occur in simple stratigraphic order, are folded and faulted locally but not disrupted into a melange, and are affected only by mild metamorphism (Dickinson and others, 1969). The great bulk of material comprising the sequence is believed to be derived from the igneous and metamorphic terrains of the Klamath Mountains and Sierra Nevada to the north and east of the Great Valley, respectively.

The Great Valley Sequence was primarily deposited as deep marine turbidites in a fore-arc basin associated with Late Mesozoic subduction (Dickinson, 1971; Ingersol, 1976). This depositional environment resulted in coarse-grained (sandstone and conglomerate) tongue-shaped lenticular beds set in a continuum of fine-grained material (Ingersol, and others 1977). This lateral variability of the continuity of the coarse-grained units has given rise to many local lithologic names which lack regional continuity and are generally uncorrelatable. The Great Valley Sequence has long been divided into three principal units based on faunal criteria (Anderson, 1933): the Upper Jurassic-Knoxville series, the Lower Cretaceous Shasta series, and the Upper Cretaceous Chico series. More recent work has relied on petrologic criteria of the coarse-grained portions to subdivide the Great Valley Sequence into formations (e.g., Dickinson & Rich, 1972) and to subdivide on texturally mappable criteria. Ingersol and others (1977) combined these methods by broadly subdividing the mapped coarser-grained units, based on internally consistent petrology, to establish five formational units. The lower three of these units were adapted for this study: Stony Creek, Lodoga, and Boxer Formations. The Stony Creek Formation includes all of the Knoxville Series and the lower Shasta Series while the Lodoga Formation comprises the upper Shasta Series. The Boxer Formation forms the lowermost unit of the Chico Series. Nomenclature used here for the remaining exposed portion of the Chico Series is that originally used by Kirby (1943), in ascending order: Venado, Yolo, Sites, Funks, Guinda, and Forbes formations. The uppermost unit of the Chico Series, the Kione Formation, is not exposed in the study area but is identified in the subsurface east of the project (see plate 5).

Throughout most of the immediate project area the Upper Cretaceous Chico series is heavily blanketed by the late Pliocene Tehama Formation and younger alluvial terrace materials. Hence, heavy emphasis was placed on correlating and extrapolating the limited number of outcrops to more extensive and better known exposures adjacent to the project area.

Stony Creek Formation. The Stony Creek Formation consists of dark greenish-gray to black, thinly bedded to laminated siltstone and mudstone containing minor thin to thick beds and lenses of sand and conglomerate (Ingersol, 1977). The lower half of the unit contains a basaltic sandstone member and a variably thick, pebble-cobble conglomerate, while the upper half of the unit is predominately siltstone. The Stony Creek Fault forms the lower contact of the unit. Along Stony Creek the formation is 15,000 to 18,000 feet thick.

Lodoga Formation The Lodoga Formation consists of numerous thin to thick, fine- to medium-grained lenticular sandstone beds, or sets of sandstone-silt-stone beds, at closely spaced stratigraphic horizons (Ingersol and others, 1977). In contrast the underlying Stony Creek Formation is predominately mudstone with scattered sandstone beds. Additionally the sandstones of the Lodoga Formation are gray and brown in color. The contact between the two formations is generally sharp or gradational over a vertical distance of 18 to 90 feet. The Lodoga Formation is about 8,000 to 10,000 feet thick.

Boxer Formation. The lower portion of the Boxer Formation consists predominately of medium to dark gray, thin bedded to laminated siltstone and mudstone with a conglomerate unit at the base, while the upper portion has randomly spaced beds, rarely up to 3 feet thick, of fine-grained sandstone (Ingersol and others, 1977). The contact between the conglomerate unit and the Lodoga Formation lies in back of the prominent north-south-trending ridges at Julian Rock where the conglomerate is approximately 280 feet thick. The Boxer Formation has previously been referred to as the Julian Rocks Formation by Chuber (1901). The formation is approximately 5,000 feet thick in the Julian Rocks area. This is the basal unit of the Chico Series. (The Boxer is composed of Julian Rocks and Clark Valley mudstone on some sections presented in this report because of acoustical character and other distinctions.)

Venado Formation. The remaining portion of the Chico Series consists of alternating formations dominated by sandstone (Venado, Sites, and Guinda Formations) or siltstone and shale (Yolo, Funks, and Forbes Formations). Ingersol and others (1977) note that the sandstone units are lenticular along strike and are therefore not everywhere present, leaving no means for differentiating the similar appearing siltstone/shale formations. However, in the project area, each of the individual formations appear represented; hence the terminology of Kirby (1943) was used in this study rather than the more general terminology of Ingersol and others (1977). The Venado Formation consists predominately of sandstone with varying amounts of siltstone and mudstone interbeds. The sandstone is light olive-gray to pale olive, thick bedded to massive, and fine- to medium-grained. The formation tends to form prominent strike ridges in the topography. Faulted thickness in the Stony Creek area is approximately 600 to 900 feet; beneath the dam the strata is up to 2,000 feet thick.

Yolo Formation. The Yolo Formation consists primarily of siltstone and mudstone with numerous thin layers of sandstone. The siltstone and mudstone is medium to dark gray and thin bedded to laminated. The sandstone is fine-grained and tends to be flaggy. In the lowermost part of the section the sandstone and siltstone is rhythmically bedded (Kirby, 1943; Ingersol and others, 1977). The Yolo Formation tends to form low rounded hills with few outcrops. In the Stony Creek area the faulted thickness of the Yolo appears to be 900 feet. Beneath the dam its sectional thickness is about 600 feet.

Sites Formation. The Sites Formation is mostly concretionary sandstone. The sandstone is light olive-gray to pale olive, thin to thick bedded, and fineto medium-grained (Kirby and Ingersol). Interbeds of greenish gray carbonaceous siltstone are common. The Sites Formation is thick to the south (Nye Canyon area) but thins dramatically immediately north of the project area. The Sites formation is about 2,500 feet thick in unfaulted section beneath the dam and thins gradually to the east.

Funks Formation. The Funks Formation consists primarily of greenish gray siltstone and mudstone similar to that of the Yolo Formation. These rocks are interbedded with fine-grained sandstone. The Funks Formation also tends to form low rounded hills with few outcrops. The Funks Formation appears to be 400 feet thick beneath the dam.

Guinda Formation. The Guinda Formation consists mostly of massive to well bedded sandstone alternating with thin bedded siltstones and shales. The sandstone beds weather from gray to buff, are fineto medium-grained, and are laterally persistent over distances of hundreds to thousands of feet. Calcareous concretions are abundant in the upper portion of the formation, commonly reaching 3 feet in diameter and rarely reaching 15 feet in diameter. The siltstone and shales, interbedded with the sandstones, are gray, generally less than 5 cm thick, and strongly laminated. They occasionally contain small concretions and rarely display ripple marks and convoluted bedding (Hoggart and Ward, 1984). The Guinda sandstone is well exposed approximately 1 mile upstream from the dam where wave erosion has exposed it along the shoreline between Orland and Black Buttes. A few small outcrops are also found on the west flank of the buttes where the Guinda Formation is unconformably overlain by the Lovejoy Basalt which caps the buttes. Shadows reveal that the sandstone beds strike parallel to the trend of the buttes (N. 100 W.), indicating the existence of the sandstone along the entire west face of the buttes. At exposed section on the buttes, the Guinda is about 500 feet thick.

Dobbins Shale. The Dobbins Shale is a distinctive mudstone formation lying conformably upon the Guinda Formation. Kirby (1943) identified this lithologic unit as constituting the base of the Forbes Formation. Pessagno (1976) elevated the Dobbins Shale to member status as the basal member of the Forbes Formation. However, others have considered the Dobbins Shale as a unique unit separated from the Forbes Formation by an unconformity. These relationships were not determinable in the area of Black Butte Dam where a limited amount of outcrop is found between the dam and the buttes. The Dobbins Shale is well recognized in the subsurface east of the dam where at 3 miles distance its

upper surface is found at the 4,300-foot depth in gas wells. The Dobbins Shale is also readily recognized by its foraminiferal assemblage, and it marks the top of the "G" foraminiferal zone of Goudkoff (1945). This important marker identifies the bottom of the Forbes Formation which is the major gas producing formation. East of the dam the Dobbins shale appears to be 300 feet thick in the subsurface and thickens slightly at the damsite.

Forbes Formation. The Forbes Formation consists mainly of carbonaceous siltstone and shales with occasional thin beds of fairly soft sandstone. The siltstone and shale is generally well bedded to laminated and has a greenish gray to light gray color. The sandstone beds are lenticular and for the most part are fine grained.

The sandstone locally contains pebbles and fragments of rock derived from older Cretaceous beds to the west (Kirby, 1943). The Forbes Formation is approximately 2,700 feet thick in the subsurface east of Black Butte Dam where it is the major gas producing formation in the Black Butte/Malton Gas Field.

In the subsurface, beneath R. 3 W. and eastward, the Forbes Formation is conformably overlain by the Kione Formation while in the area of Black Butte Dam (R. 4 W.) the top of the Forbes Formation is marked by an erosional unconformity and is overlain by Tertiary and Quaternary deposits. At the dam site, the Forbes Formation constitutes the foundation of the control tower, transition section, and a portion of the main dam embankment and tunnel, where it is characteristically massive to thin bedded, gray, soft shale with thin sandstone and limestone interbeds.

Kione Formation. The Kione Formation is not exposed in the project area but is known to exist in the subsurface some 4 miles east of the dam. Early Tertiary erosion has removed it near the dam (see plate 5). The Kione consists of fine- to coarse-grained sandstone interbedded with gray siltstones and shales. Some sand lenses have good lateral development which serve as major gas reservoirs (Barger and Sullivan, 1966). The Kione Formation is the uppermost unit in the Chico Series and thus marks the top of the Cretaceous section. This surface is marked by a major erosional unconformity. The Kione formation thickness is extremely variable in the subsurface sections showing the stratigraphy of the northern Valley (see, for example, California Division of Mines and Geology Bulletin 181).

Paleogene Units. A major erosional unconformity marks the boundary between the Cretaceous formations and the overlying Tertiary deposits. This unconformity marks a period of time (Paleocene) in which a huge submarine valley (lower Princeton Valley) was eroded into the Cretaceous units in the area of the present Sacramento Valley (see figure 2-4). This submarine valley became infilled in the early Eocene and is now occupied in the subsurface by sandstone and siltstone of the Capay Formation. Additional middle to late Eocene formations conformably overlie the Capay Formation and unconformably underlie the Lovejoy Basalt. These formations, Capay, Ione, Domingine, and Nortonville, are all found at depth in the gas wells in the central Sacramento Valley. They are generally marine deposits but locally, and particularly around their

periphery, may include nonmarine facies. None of these Eocene formations are exposed in outcrop along the western margin of the valley in the area of Black Butte Dam. While not exposed, these units are discussed here to complete the regional geologic section and history and to present the possible relationship with the Black Butte Formation.

The Black Butte Formation is one additional Paleogene Formation which lies directly beneath the Lovejoy Basalt at the dam site and was exposed extensively during foundation preparation for the dam. The following description of these units relies in part on descriptions presented by Redwine (1979) and Boyd (1956).

Capay Formation. The Capay Formation consists principally of thick to very thick bedded, internally massive claystone, siltstone, and sandstone beds with a chaotic basal conglomerate unit. The pelitic rocks are greenish, the sandstone beds are buff to brown and generally massive, and the conglomerate is dark to very dark gray with a sandstone matrix.

This conglomerate is similar to conglomerates found in the Upper Cretaceous section suggesting that the conglomerate unit was a very large submarine slide or submarine mass flow (Redwine 1972). Overall, the Capay Formation was deposited in the submarine lower Princeton Valley during early Eocene where it reaches a thickness of over 2,000 feet. Typical thickness of the formation beyond the confines of the principal submarine valley is typically 300 feet. (See figure 2-4 for the extent of the Princeton Valley and its relationship to Black Butte Dam.)

Ione Formation. The Ione Formation is apparently absent in the subsurface east of the project area. It consists for the most part of white to yellowish white, highly quartzose kaolinitic sandstone with some claystone interbeds that lie conformably on the Capay Formation. The Ione Formation is mostly early Eocene and partly middle Eocene in age.

Domengine Formation. The Domengine Formation is principally sandstone but occasionally contains interbeds of shale, siltstone, and lignite. The Domengine Formation lies conformably upon the Capay Formation and upon the Ione Formation where that formation exists. The formation is approximately 170 feet thick in the Johnson Unit No. 6 well, 5 miles east of Black Butte Dam. The Domengine Formation is considered to be middle Eocene age.

Nortonville Formation. The Nortonville Formation consists principally of siltstone with minor shale and some carbonaceous material. The siltstone is brown to tan, firm, massive, fine to very fine grained, highly quartzose, and micromicaceous. The Nortonville Formation lies conformably upon the Domengine Formation and is late Eccene in age. The formation represents relatively deep marine sediments and is the youngest marine formation of the last major marine transgression of the Sacramento Valley. It is approximately 100 feet thick in the Johnson Unit No. 6 well, east of Black Butte Dam.

Markley Sandstone Formation. The Markley Sandstone is well documented far south of the study area, in the Mount Diablo area. The Markley Sandstone is not recognized in the subsurface immediately to the east of Black Butte Dam, however similar sandstone is widely recognized in the Sacramento Valley subsurface. This seemingly widespread, previously unnamed, sandstone lying between the Nortonville Formation and the Lovejoy Basalt (Harding and others, 1900) to the east and southeast of the study area was considered by Van den Berge (1960) to be part of his nonmarine Nord Formation. Van den Berge considered the deposits (Black Butte Formation) beneath the basalt at Black Butte Dam to be an outcrop of his Nord Formation. Redwine (1972) correlated and considered most of Van den Berge's Nord Formation to be the Markley Formation; however, in the "Hamilton Nord" No. 1 well, located due east of the study area and near the Sacramento River (sec. 9, T. 22 N., R. 1 W.), Redwine considered it to be part of the Domengine Formation based on correlation to adjacent wells.

The Markley sandstone consists principally of grayish brown to yellowish brown, arkosic sandstone with abundant large muscovite flakes. It rests conformably upon the Nortonville Formation, is thought to be late Eocene in age, and is probably a marine to nonmarine deposit.

Black Butte Formation. The name Black Butte Formation was coined by the Corps of Engineers for a distinctive sequence of rocks found at the damsite underlying the Lovejoy Basalt and overlying the Upper Cretaceous formations. The sequence consists of a basal conglomerate overlain by mudstone and sandstone. The formation was exposed extensively during foundation preparation for the dam and pertinent structures. A somewhat degraded exposure of the sandstone can be seen in the spillway and a good exposure of the mudstone can be seen in the right bank below the intake tower bridge. There are no natural exposures of the conglomerate and those excavated during construction were covered by the structures. The foundation report (Corps of Engineers, 1963) gives a detailed account of the rock types as follows:

"The deposit is well consolidated and poorly cemented to uncemented although some zones of the sandstone and conglomerate are moderately cemented. The conglomerate consists of rounded particles from fine sand to 6-inch cobbles. In most excavations the matrix consists of fine to coarse sand and the deposit is slightly to moderately pervious. The grading of the constituents and presence of some cementing agent, together with silty zones make the deposit less pervious than would be supposed. In general, the material is fairly well graded from fine sand through 2-1/2-inch gravel with a lesser amount of cobbles up to 6-inch size. The particles are hard and fresh and are composed of quartzite, chert, volcanics, jasper, vein quartz and metamorphics. Despite the multicolored particles, the conglomerate is typically either blue or tan. The deposit varies in thickness from a minimum of 10 feet in the stream channel to a maximum of 75 feet in the left abutment where it interfingers with sandstone... Sandstone and mudstone are grayish green to blue-

gray or tan. The mudstone is composed of clays, silts and fine sand and frequently has a pseudo-granular texture; that is, it has the appearance of fine-grained sandstone but breaks down to a slick, shiny clay when moistened and rubbed. The material is well consolidated and uncemented except in the upper few feet where heat from the basalt flow has baked it into a moderately hard, well indurated rock having much the same appearance as the basalt. In fact, in some localities, particularly the right abutment core trench, the mudstone has a definite columnar structure in the upper 3 or 4 feet of the baked zone. In the outlet tunnel random zones occur where it is a light chocolate brown color, is moderately hard and is assumed to have been thermally altered even though the zones are often completely surrounded by softer mudstone. The unaltered mudstone is typically massive, soft and can usually be crumbled under hand pressure when dry. It tends to air slake badly after exposure for a few weeks. The sandstone is poorly cemented, soft and crumbly to moderately cemented and moderately hard. It is friable even in the moderately cemented zones. Although megascopically blue or bluish gray, examination under a hand lens indicates it is composed of vari-colored metamorphic and volcanic rocks and contains a large percentage of white quartz. The rounded particles are fine to medium grained with occasional zones of coarse particles up to one-quarter inch. Frequently there are silty or clayey zones, some in which the clay has a pseudo-granular texture. It is usually the silty and clayey zones which are least cemented. Where sandstone is in contact with the basalt, it, too, has been well indurated by heat and is black, dark gray or brown. Downward from the basalt contact, it gradually becomes less indurated and grades to the typical bluish color. The sandstone is massive with no indication of bedding."

The Black Butte Formation probably represents a part of one of the Eocene formations or possibly a facies change of one of the Eocene formations with the depositional environment being high on the west margin of the lower Princeton submarine valley. There are several possible correlations.

The only known outcropping of the Black Butte Formation in the project area is at the damsite where it attains a maximum thickness of 100 feet. The position of the Cretaceous rocks and Lovejoy Basalt higher on the buttes limits the thickness of the Black Butte sediments (beneath colluvium) in some areas, to less than 40 feet. Hence, the Black Butte Formation at the dam site may occupy an erosional low in the topography existing at the time of its deposition.

In the existing literature, Van den Berge (1968) has made the only known correlation of deposits at the damsite to another known Eocene Formation, specifically his Nord Formation which he describes as variegated bentonitic clays. While not directly addressing the damsite outcrops, Redwine (1972), suggests that most of the "Nord Formation" belongs to the Markley Formation and locally to the Domengine Formation (see paragraph on Markley Formation). Redwine

acknowledged the possible existence of an unnamed post-Markley, pre-Lovejoy Basalt volcaniclastic formation, at least locally, directly beneath the basalt. He represents these deposits as a volcanic episode culminating with the Lovejoy Basalt. Such relationships would correlate the Black Butte Formation with the Domengine, Markley, or a younger volcaniclastic (possibly redefined "Nord") formation.

Descriptions in the literature of the Capay Formation (Boyd, 1956, see Capay section; Redwine, 1972) bear a remarkable resemblance to the Black Butte Formation. Additionally, the Capay Formation, as far as is known, abuts the sides of lower Princeton Valley rather than drapes the sides. No megafossil or microfossil studies were performed in connection with any study and no other field evidence has been found to further establish the age of the Black Butte Formation or its correlation to other formations. This study concludes that the Black Butte Formation is a younger volcaniclastic formation, of Oligocene or early Miocene age, and that both its upper and lower contacts are marked by unconformities. At the dam, the Black Butte Formation unconformably lies upon the Upper Cretaceous rocks and is overlain unconformably by the Lovejoy Basalt.

Neogene Units. After deposition of the Eocene formations a relatively smooth, erosional, subaerial surface developed in the present area of the Sacramento Valley upon which the Lovejoy Basalt spread. The basalt apparently issued in a south and southwesterly direction from the Northern Sierra Nevada. The upper Princeton Valley was then subaerially eroded and then backfilled with the Upper Princeton Valley fill. Major subaerial erosion occurred once again after which the fluvial Neroly Formation was deposited. After another similar cycle the Tehama Formation was deposited.

Lovejoy Basalt. The Lovejoy Basalt is a discontinuous unit found in the subsurface of the Sacramento Valley and in outcrops around its margins. In the subsurface the basalt is found as the rimrock of the upper Princeton Valley (Redwine, 1972). This relationship may also hold for the limited surface exposures of the basalt along the western margin of the Sacramento Valley. The Lovejoy Basalt is the dominant geologic feature at the dam site, capping and leaving near vertical scarps on the flat topped Orland and Black Buttes (which rise to over 1000 feet in elevation), and is conspicuous in the topography of smooth and rounded hills. On the whole the buttes are formed of Cretaceous sandstone beds of the Guinda Formation, capped by 40 to 80 feet of basalt, and dissected by Hambright and Stony Creeks. From the buttes the basalt dips gently eastward at nearly 4 degrees and appears to extend into the subsurface. Magnetic modeling performed as part of this study (see geophysical studies section) suggests the basalt is a continuous flow dipping to the east in the subsurface. Two miles east of the dam, the Lovejoy Basalt is found in gas wells at approximately 1,200 feet below sea level with a thickness of up to 120 feet ("Black Butte" 27-3, Grace Petroleum Co., well). Paleomagnetic analysis performed in connection with magnetic modeling on Lovejoy Basalt samples taken from Orland Buttes concluded: (a) magnetization on the basalt was very strong, (b) the basalt was formed during a reversed magnetic epoch, and (c) the basalt contains slight variations in declination (average 155 degrees) and inclinations (average -52 degrees).

The Black Butte Dam Foundation Report (Corps of Engineers, 1963) offers a good description of the basalt as found in the study area:

"Two separate flows have occurred, and underlying the basalt, in localized areas at least, is a volcanic breccia which reaches a maximum thickness of 16 feet. The lower boundary of the basalt is marked by a thin vesicular zone. The top flow, which averages 45 feet in thickness is characterized by its crude columnar structure and darker color. The columns are mostly 3 to 12 inches in diameter. They are very irregular with nonplanar sides that curve and twist. Other low angle joints divide the columns into blocks that generally range from 3 to 24 inches. The fractures have been tightly healed with zeolites, quartz and lime, and in unweathered rock the columnar structure is not readily apparent. However, with only slight weathering the fracture fillings are removed or their bonding strength is lost. The structure then becomes quite definite, at least in its breaking characteristics. When unweathered, the rock is as apt to break across the healed fractures as along them. When slightly weathered, breakage of the rock is controlled by the cooling fractures with resultant small, crude, roughly equidimensional columnar fragments, the most common size being 3 to o inch. The top flow is black, dense, aphanitic, hard, and extremely brittle. The lower flow, averaging about 20 feet in thickness, differs from the upper flow in its lighter (medium to dark gray) color, lack of any columnar structure, blocky fracturing and in being slightly softer and less brittle. Fractures of the lower flow are generally filled with soft, brown to yellow-brown zeolites. Some fracture fillings are as much as 2 or 3 inches wide. The material is usually flaky and has little strength. Fractures are spaced from a few inches to several feet and average about 1.5 to 2 feet. They occur in a blocky or rectangular pattern. The contact between the two flows is not a sharp break. On close inspection it appears to be gradational but from a distance there is a definite change in structure which can be located within a space of one or two feet. Petrographic analysis indicates that 27 percent of the basalt is andesine, 14 percent is augite and the ground mass, which accounts for 59 percent of the volume, consists of a dark opaque volcanic glass. The high percentage of glass accounts for the extreme brittleness and conchoidal fracturing featured by the basalt. The base of the basalt is marked by a layer of breccia which ranges in thickness from 1 to 16 feet where observed in excavations. In the thinner zones it apears only to be a vesicular layer but frequently contains pockets of loose pyroclastics and baked mudstone. In the thicker sections, particularly in the outlet tunnel and left abutment core trench, the flow breccia consists of vesicular basalt boulders of various sizes up to 25 feet surrounded by a mixture of smaller fragments and volcanic ash together with inclusions of mudstone. It is sometimes completely loose but frequently welded together in a fairly stable

mass. Weathering has turned the breccia rust-brown but has not been intense. Vesicles and surfaces of the larger fragments are coated with zeolites and calcareous material. The only evidence of flow direction is in the streaking out of some inclusions of mudstone, in an easterly direction."

This study also recognizes sparse megascopic olivine crystals.

Durrell (1959) states the origin of Lovejoy Basalts to be related to the Cohassett Ridge and Big Chico Creek flows. As part of this study, and to aid in correlating the basalt at Orland/Black Buttes, two samples of basalt were radiometrically dated (K-Ar). One sample, taken near the abutment of the intake control tower bridge, yielded an age of 15.2.+ 0.2 million years. The other sample, collected from Big Chico Creek (sec. 35, T. 23 N., R. 2 E.) on the eastern margin of the Sacramento Valley, yielded an age of 13.6+ 0.2 m.y. Both samples were formed during a single reverse-magnetic polarity episode and have slight diversity in paleomagnetic inclination but similar declamation. Thus, the Orland Buttes basalt indicates coeval age but with tectonic between site deformation. Hence, based on these two dates, the Lovejoy Basalt is believed to be mid-Miocene spanning a 1.5 million-year-old flow episode. This differs from most previously published ages: Dalrymple (1964), early Miocene; Van den Berge (1968), late Oligocene or early Miocene; Durrell (1959), upper Eocene and lower Oligocene. Throughout its occurrence, the Lovejoy Basalt and related underlying volcaniclastic interval rests unconformably upon all older formations, the ancestral Sacramento Valley having been eroded to a fairly level subaerial basin. Along the northeastern margin of the valley as many as 15 separate flows have been identified. Apparently only the largest flows crossed the valley as generally fewer numbers of flows are recognized to the west and south; two at Orland and Black Buttes.

Upper Princeton Valley Fill. The upper Princeton Valley fill typically comsists of variegated sandstone with locally varying amounts of interbedded shale and siltstone and occasional conglomerate. While commonly greenish brown to gray, the inclusion of volcaniclastic minerals causes the varying degree of color. The shale and siltstone are also somewhat variable in color, moderately soft, and have a waxy texture. These materials were deposited in the upper Princeton Valley in a subaerial environment principally as channel deposits of an ancestral Sacramento Valley. The age of the fill is post-upper Princeton Valley erosion, which in turn is post-Lovejoy Basalt, and is pre-Neroly Formation. Hence, the fill is considered upper Miocene and possibly lower Pliocene in age. The upper Princeton Valley fill is not identified in the subsurface stratigraphy east of the dam site. However, Redwine (1972) notes the existence of a thin section of the basal part found in the "Altofer" No. 1 well (sec. 14, T. 22 N., R. 4 W.) four miles southeast of Black Butte Dam. A thin section of the fill may be included, possibly locally, with the basal section of the Tehama Formation, as shown in the section on plate 5. As in the lower Princeton Valley, the fill abuts rather than drapes the upper Princeton Valley walls.

Neroly Formation. The Neroly Formation consists of locally variable amounts of interbedded tuffaceous sandstone and tuffaceous shales with occasional conglomerate. These sediments commonly have a bluish color due to montmorillonite coating, but where not tainted are typically greenish gray and dark gray. These predominately volcaniclastic sediments unconformably overlie the upper Princeton Valley fill. The Neroly was deposited on an erosional surface characterized as a nearly level fluvial flood plain. As with the upper Princeton Valley fill, the Neroly Formation is not specifically identified in the subsurface east of the damsite. The deposits belonging to the Neroly Formation may be included, possibly locally, in the basal section of the Tehama Formation. Due to its stratigraphic position between the upper Princeton Valley fill and the Tehama Formation, the Neroly Formation is here considered to be possibly early Pliocene and probably middle Pliocene in age.

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# APPENDIX C

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#### WILLOWS FAULT SYSTEM

Introduction. Willows Fault system was identified by Harwood and Helley (1982) in United States Geological Survey (USGS) Open File Report 82-737. Bolt (1982) relied upon the faults as identified by Harwood and Helley to postulate a "close in" fault source for design earthquake of his report. A large part of our study was devoted to examining the field evidence for the existence of this fault system.

Willows Fault System. The Willows Fault system was defined by Harwood and Helley based on contouring of stratigraphic units obtained from analysis of well log data. Some surface mapping and seismic reflection profiles run by Seisdata Services Inc. added to their confidence. Their concept of the Willows System is shown on plate 6. We conclude that the Willows Fault system, as defined, does not exist. Points we take exception to are noted following each quote from Harwood and Helley's open file text (excerpted below).

"The main stem of the Willows fault was discovered in the subsurface rocks of the northern Valley when it was penetrated by the Marathon Oil Company (formerly Ohio Oil Company "Capital Company" No. 1 well during development of the Willows-Beehive Bend gas field in the late 1950's (Calif. Div. Oil and Gas, 1960; Alkire, 1962; 1968). From the discovery well, Redwine (1972) traced the Willows fault southeast to Sutter Buttes and he suggested that it extended northwest of the discovery well possibly connecting with the surface fault mapped west of the Orland Buttes (Anderson and Russell, 1939; Jennings and Strand, 1960). Redwine (1972) documented displacement on this 40 km (25 mi) segment of the main stem of the fault and concluded that it dipped 740 or steeper to the east and showed reverse, east-side-up movement that decreased upward in the geologic section. He found that the Princeton submarine channel was localized, in part, by "movement on the fault and that vertical separation in the discovery well varied from about 488 m (1610 ft) on top of the Cretaceous rocks to about 477 m (1575 ft) on the top of the Eocene Capay Formation. At Orland Buttes, the Willows fault offsets the Miocene Lovejoy Basalt and the Pliocene Tehama Formation.

Fault at Orland Buttes. The existence of a fault at Orland Buttes is discussed in the section on Black Butte Fault. A conclusive line of evidence has been developed to refute the existence of a high-angle fault here. Neither the Lovejoy Basalt nor the Tehama Formation are offset. Thus the Willows trend, as shown by Open File Report 82-737, is broken, and this study concludes the Willows Fault extends no farther than Artois (section 20-3) as indicated by Alkire's (1902) analysis.

"Data from a number of sources indicate that the Willows fault is far more extensive and complex than previously thought. The first clue that the Willows fault branched into a multi-strand fault system was provided by an analysis of seismicity of the northern Valley and Sierran foothills after the Oroville earthquake. Marks and Lindh (1978) located a number of small-magnitude earthquakes that originated near the discovery well in the Willows-Beehive Bend gas field and extended north for a distance of about 30 km (20 mi) rather than following the known, northwest trend of the Willows fault. The trend of seismic events suggested that a north-trending fault splayed off from the main stem of the Willows fault and passed west of the Corning domes.

Trend of Seismic Events. In our judgment, Marks and Lindh (1978) do not indicate the conclusion stated by Harwood and Helley (1982). Figure 2-14 from their article shows the relationship of Corning domes to the seismicity. Marks and Lindh (1978) do not discuss seismicity in the western valley.

"Analysis of well records in the Corning gas fields by the Sacramento Petroleum Association (1962), however, did not identify a fault west of the Corning domes, but that study did show an anticlinal fold in the area with a 121 m (400 ft) of maximum closure on the base of the Tehama Formation in the north dome and a steeply dipping southeast-trending fault located at the north end of the south Corning dome.

"Recent seismic reflection profiles in the area have identified a major north-trending, steeply east-dipping reverse fault that passes west of the Corning domes and the Greenwood anticline. Displacement of reflecting horizons increases with depth on this north-trending fault indicating progressive deformation through time similar to the pattern of deformation on the main stem of the Willows fault in the Willows-Beehive Bend gas field. From well-log data, the pattern of earthquakes shown by Marks and Lindh (1978), surficial mapping, and the seismic reflection profiles, we conclude that the Corning domes and the Greenwood anticline formed by east-side-up drag on the north-trending Corning fault which splays off from the main stem of the Willows fault."

Corning and Greenwood Folds. We do not interpret the Corning domes and Greenwood anticline as drag folds. There is no data present east of and at the 1-5 corridor to justify the map position shown by Harwood and Helley (1982) for the fault (see the Hayes-Burnell and Long Hollow Seismic Section in the Geophysical Study section of this report, paragraph 2.4).

"The location of this north-trending splay fault north of Red Bluff is unknown, if, in fact, it extends north of the Red Bluff fault.

"The youngest deposits deformed by the Corning fault are gravels of the Pleistocene Red Bluff Formation, the age of which has been bracketed by overlying and underlying volcanic rocks radiometrically dated at 0.45 and 1.09 m.y., respectively (Harwood and others, 1981).

"The location of the Willows fault system north and northwest of Orland Buttes is not closely controlled by direct evidence in either the surface or subsurface rocks. Wells are sparse in this area of the Valley, particularly wells located west of the probable trace of the main stem of the Willows fault, and, for that reason, the structure contours are very generalized and of little value in locating even major structures. Several pieces of evidence, however, point to a pattern of deformation associated with the Corning fault and the main stem of the Willows fault to the south and east.

"First, the north northwest-trending fault, shown as the Malton fault east of Orland Buttes and extending at least 32 km (20 mi) northward toward Red Bank, was crossed by two seismic lines run by Seisdata Services, Inc. Their profiles show a steeply east-dipping fault with east-side-up displacement that increases with depth. The broad anticlinal dome identified by Redwine (1972, section AA') southeast of Red Bank lies east of the Malton fault and appears to bear the same genetic relationship to that fault as the Corning domes bear to the Corning fault."

Seisdata Lines. At Orland Buttes two proprietary seismic lines (not Seisdata) fail to confirm the Malton or Willows faults as presented. The fault found on the Seisdata line north of Black Butte, identified as the Malton, is in fact a northeast-trending high-angle fault (east side up) identified as Burris Creek fault by this study.

"The base of the Tehama Formation occurs at an elevation of -17 m (55 ft) in the McCulloch Oil Corporation 'McCulloch Sunray Anchordogry' no. 1 well (sec. 7 T25N R4) east of the Malton Fault and at -297 m (980 ft) in the Occident Petroleum Corporation 'Harris' no. 1 well (sec. 23 T26N R4W) west of the fault. This difference in elevation on the base of the Tehama is due to the combined effects of doming east of and offset on the Malton fault."

Citing these wells for existence of the Malton fault is ambiguous. Occidental Petroleum Harris No. 1 lies north of the map limits shown by the authors on the map plate of Open File Report 82-737. This well is possibly more meaningful to the development of south-side-down movement on the Red Bluff Fault, a position sponsored by both the authors and other investigators (Jennings (1976) and Peppard and Associates (1981)). Redwine (1972) indicates a broad fold, unfaulted in section, with a wave length of 8 miles, not the 2- to 4-mile wave lengths of Greenwood or Corning folds.

"Relatively extensive deposits of Red Bluff gravels cap the interfluves east of the Malton fault, but only a few scattered patches of Red Bluff gravel occur west of the fault. The western limit of the extensive Red Bluff deposits is generally aligned along the Malton fault trace suggesting that the Red Bluff was stripped from the area west of the Malton fault either by deformation on that fault or by uplift on the Willows fault to the west. The northern extent of the Halton fault is not known." Red Buff Deposits. It would possibly be more fitting to explain the western limit of the Red Bluff Formation as controlled by source and waning of southward tilting of the Great Valley. East-side-up faulting does not need to be implied. Steele's (1979) mapping of the Red Bluff Formation indicates extensive deposits west of the hypothetical Willows and Malton Faults. The Black Butte and Malton gas fields are two such areas blanketed by Red Bluff gravels.

"Our projection of the Willows fault into the Cold Fork and Elder creek faults, and possibly into the Paskenta fault mapped by Jones and others (1969) is based primarily on the outcrop pattern of the Tehama Formation. North of Elder Creek, the Tehama Formation dips gently east and the Nomlaki Tuff Member is at the base of the Tehama or is, at most, a few tens of feet above the base of the Pliocene section. South of Elder Creek, however, the Tehama dips more steeply into the valley and the Nomlaki Tuff is a few hundred feet above the base of the Tehama. This outcrop pattern of the Tehama Formation suggests that the Great Valley Sequence was topographically higher and projected farther east into the Valley north of the Willows fault prior to deposition of the Pliocene rocks. The position of the Nomlaki Tuff relative to the base of the Tehama on opposite sides of the Willows fault indicates that the Tehama filled a topographic low southwest of the fault prior to eruption of the Nomlaki Tuff about 3.4 m.y. ago. We interpret the topographic low, reflected by the thicker basal part of the Tehama, to be the result of east-side-up movement on the Willows fault prior to and possibly during deposition of the early phases of the Tehama Formation.

We believe the position of the basal Tehama deposits was controlled by sedimentary bench-land topography near topographic gaps in the strike valley walls of the Pliocene foothills locating fanheads. Along the flanks of the fan, the Nomlaki Tuff is thin and occurs near Cretaceous bedrock. In the core of the fan it thickens and is up further in the Tehama section. We suggest that there are two different fans north and south of Elder Creek.

"If this interpretation is correct, Late Cenozoic movement on the Willows-Elder Creek fault system has been significantly different, in style and movement of displacement, from the Cretaceous movement on the Elder Creek-Cold Fork-Paskenta fault system outlined by Jones and Irwin (1971)."

"Jones and Irwin (1971) inferred at least 96 km (60 miles) of left-lateral displacement of the Early Cretaceous (Valanginian) shoreline along the combined Cold Fork-Elder Creek-Paskenta faults. They concluded that this deformation commenced shortly after deposition of the Valanginian rocks and continued concurrently with deposition until at least mid Late Cretaceous time. Well documented vertical displacement of Late Cretaceous and younger rocks on the Willows fault in the Beehive-Bend gas field is not incompatible with left-lateral displacement on the Elder Creek fault system to the northwest, but it does indicate that the inferred lateral displacement was accompanied

by a major component of east-side-up vertical movement. This vertical movement is consistent with the interpretation that the Elder Creek fault system represents tear faults in the upper plate of the Coast Range Thrust (Jones and Irwin, 1971) along which the Klamath Mountain terrane moved upward and westward over the Coast Range province."

In conclusion of the review of evidence for the Willows Fault system, we wish to address the Seisdata lines referred to in Open File 82-7370. Three out of five lines crossing the map positions of the Willows Fault system have been examined by this study. (See time sections in Geophysical Study section of main report). The existence of faults in two Seisdata time sections and other oil company time sections examined is not denied. However, neither the orientation and position of faults nor the density of line coverage allow us confidently to define through-going faults, or to indicate the presence of a multiple branching fault system.

APPENDIX D

TECTORIC WEDGE THROWY

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#### TECTONIC WEDGE THEORY

Introduction. This appendix discusses what effect the tectonic wedge theory has on seismic hazard analysis at Black Butte Dam. It should be emphasized that tectonic wedging for the eastern margin of the Coast Ranges is an emerging theory and has not yet been widely accepted as the tectonic mechanism responsible for the structure of the western edge of the Great Valley and eastern Coast Ranges.

Coherent Structural Interpretation. Chief spokesmen of this theory are pursuing a coherent structural interpretation for the entire west side of the Great Valley of California. This interpretation suggests that the Franciscan complex and steeply dipping Great Valley Sequence is underlain by a subhorizontal to west-dipping thrust fault that nowhere reaches the ground surface. Eastward thrusting of the Great Valley Sequence and Franciscan Complex across this surface has produced an east-dipping thrust in the overlying rocks one of which is the Coast Range Thrust (Wentworth letter of 11 October 1985; Wentworth, et. al, 1984; and Wentworth, personal communication). Much of the movement on these thrusts should have occurred in the Cretaceous and early Tertiary time. Within this thrust environment, the Coalinga earthquake of 1983 (magnitude 6.7) occurred. The tectonic wedge theory has two elements of concern to seismic hazard at Black Butte Dam. One element consists of the validity of such a tectonic model in northern California, a distance of 275 miles from where the model seeks substantiation. The second element is the validity of carrying crustal strain, cause of the Coalinga earthquake, into the northern California environment. An examination of each element leads to the conclusion that although tectonic wedges might exist along the northern California western valley edge, there is no reason to believe that the crustal strain existing at Coalinga exists at the wedge nearest Black Butte Dam.

Pervasive Structure. Examination of the first element starts with the pervasive structure along the western edge of the Great Valley and within the Coast Ranges. As shown on the sketch map in figure D-1, this structure consists of the Coast Range Thrust at the eastern range front and the San Andreas zone of transform faulting. The San Andreas is a recent tectonic entity, being created when subduction along the coast was extinguished (arch switch off). The San Andreas lengthened along the California coast margin north and south of the point of contact between the Pacific and North American plate. In the southern ranges this first occurred 25 million years ago (Dickinson and Snyder, 1978). In the northern ranges the event occurred less than 5 million years ago.

The Coast Range Thrust separates two coeval terrains, the Franciscan Complex (F, see sketch) and Great Valley Sequence (K) along the eastern range front. Ophiolite (Jb) often overlies the thrust. It is generally held that Great Valley Sequence was deposited on ophiolite basement along the Mesozoic continental margin. As indicated in sections illustrated in the figure, the Coast Range Thrust (symbol CRT) can be visualized as defining the upper plate on top of an eastward underthrusting wedge of Franciscan material. The driving mechanism for this wedge, or any wedges in the Great Valley Sequence located

further eastward, must come from the west. Eastward thrusting of the Franciscan Complex is responsible for the primary Franciscan wedge (F) and is located in the sketch sections as possibly occurring all along the valley. Faults typically associated with the top of the Coast Range ophiolite (SCF) appear to mark another wedge; thrust faults in the Cretaceous section (ST) mark other wedges. The compressional episode that caused the thrusting was definitely shut off at the onset of transform boundary tectonics. Franciscan thrusting derives from collisional tectonics prior to transform faulting when the subduction zone existed further west (Wentworth, et. al., 1984). Thus, in regional aspect, the state of stress in the crust for thrusting tectonics preceded the state of stress for transform tectonics and they are different. Similarities in geometry of the Coast Range Thrust and the Franciscan Complex-Great Valley Sequence can be followed in sections from the south to the north along the western valley edge.

Stress and Mechanism. We are now led to examine the second element. Can the conditions viewed as explaining the 1983 Coalinga event or the Kettleman Hills earthquake of 1985 be found at/near the tectonics and structure around Black Butte Dam? To answer this it is first necessary to look at the cause and mechanism of the 1983 and 1985 events. Figure D-2 shows sections normal to the Coalinga and Kettleman Hills anticlines. These sections are taken from various publication releases of USGS authors and wedging theory spokesmen. The Coalinga section is taken normal to the anticlinal axis and lies 7 km southeast of the 1983 event. The structure defined is based on interpretation obtained during the USGS Deep Crustal Study, No. 9540-02191; principal investigator, Dr. C. Wentworth. Investigators feel the 1983 Coalinga earthquake occurred at the east margin of the Coast Range beneath the Coalinga anticline and was located some 35 km northeast of the San Andreas Fault. This anticline is the northwest segment of a  $100 \ \mathrm{km}$  long zone of young anticlines associated with the San Andreas zone. Folding northeast of the San Andreas in the Diablo Range and Tremblor Range is indeed young. Harding (1976) suggests a regular progression of late Cenozoic folding away from the San Andreas, the youngest folds outboard of the older folds which lie nearer to the San Andreas. The Kettleman Hills anticline began growing in late Pliocene time and has folded a thick section of Pleistocene alluvial sediments. At the Coalinga nose, the northwest fold trend evident in the Kettleman Hills is superimposed against the existing westerly trend of older San Andreas folding.1/ Study of stream terrace deformation south of the Coalinga nose (King and Stein, 1984) suggests deformation and recent movement beneath the Coalinga anticline. The Coalinga, Kettleman Hills, and Lost Hills anticlinal deformation trend is accepted as being controlled by the San Andreas in movement much like Riedel shear is associated with master fault movement (see figure D-4 - Riedel shear).

As shown on the sketch figure D-2, northeast directed thrusts (named Coalinga Thrusts), terminate beneath the anticline, at a depth of 10 km (point C on figure), in a series of upward-splayed reverse faults above which the anticline

1/See Woodring, Stewart, Richards, 1940, and Dibblee, 1973 and 1971; references in Wentworth, et. al, 1984.

Schematic cross sections across Coast Ranges- Great Valley boundary in the north adapted from Wentworth et.al., 1984. In the south from Page, 1981.

1

# Explanation

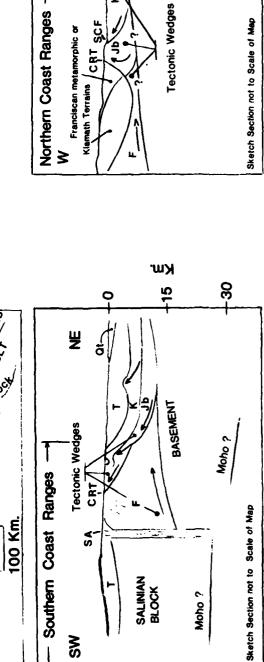
K- Great Valley sequence

T- Miocene marine sediments and volcanics

Qt- Pleist, and Plio. clastic sediments Jb- Great Valley Ophiolite

F- Franciscan Complex

CRT- Coast Range Thrust SCF- Stony Creek Thrust ST- Sites Thrust SA- San Andreas Fault



Kw

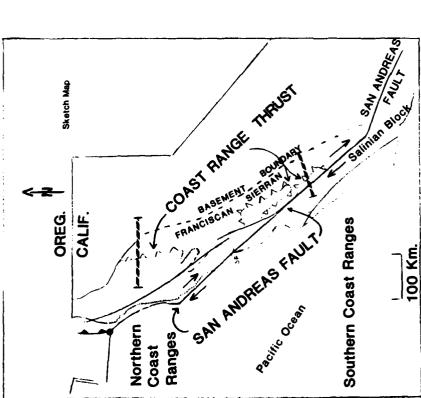
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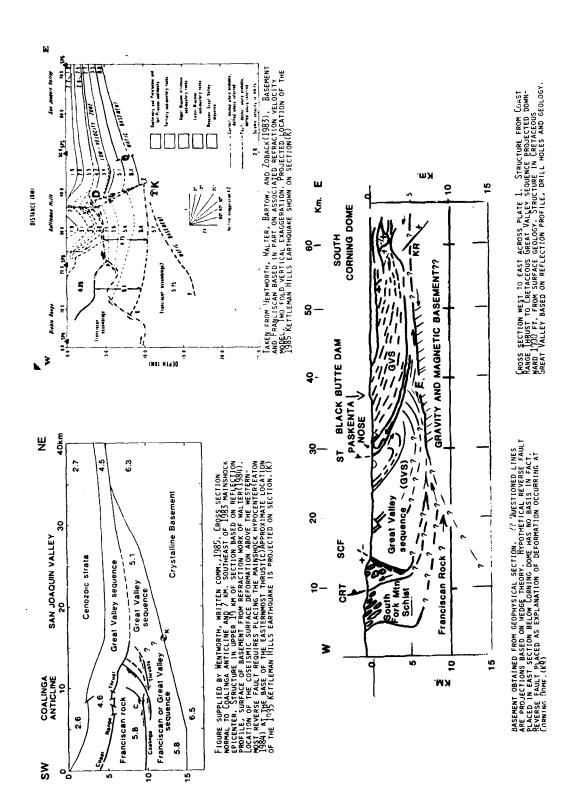
CRT SCF

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STRUCTURE ALONG COAST RANGE -- GREAT VALLEY CONTACT



TECTONIC WEDGING IN CALIFORNIA FIG. D-2

has grown. The Coalinga main shock occurred at the base of the reverse fault splay with focal mechanism that closely resembles the strike of the fold and dip of the fault (Eaton, et. al. 1984).

Thrusting of the type responsible for the growth of the Coalinga Anticline probably extends the length of the Coalinga-Kettleman Hills trend. Wentworth sees tear faults in the subsurface as en echelon steps connecting the thrust. At Kettleman Hills the thrust is not rooted and the fold is underlain by a subhorizontal thrust that terminates beneath the east limb of the anticline (point D in figure). The position of the easternmost ramp on the thrust surface was probably controlled by reverse fault deformation of the lower plate (Wentworth, Walter, Bartow, and Zoback, 1984). The driving force for thrusting is believed by Wentworth not to be superficial but deep crusted. His interpretation is a reverse fault penetrating from basement and controlling the location of the thrust ramp and fold. Mechanisms for the Kettleman Hills event suggest high angle reverse faulting (Eaton, personal communication, 1985) at the basement, consistent with shallow east thrusting. Reverse faults are localized where the basement bends to plunge beneath the Diablo Range.

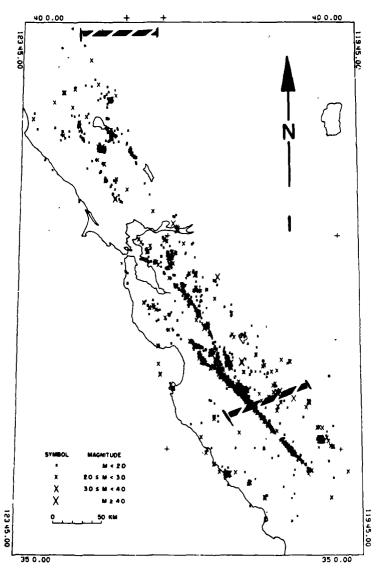
At this point the major structures present in the sections that complement the mechanism reported at the Coalinga and Kettleman Hills events can be listed:

- a. Fold belts in proximity to the San Andreas fault.
- b. Fold belts that have orientation sympathetic to shear transfer from the San Andreas.
- c. Fold belts that have en echelon nature and presumably are of some extent regionally.
  - d. Young surfaces of deformation that demonstrate recent growth in folds.
- e. Northeast directed thrusts invading eastward directing wedges in the subsurface; thrusts that do not penetrate surface but end in roof structure of fold.
- f. Reverse faults emanating from basement that are associated with growing folds and localize position of overthrust plate and fold.
  - g. Seismogenic activity in fold area (see figure D-3).

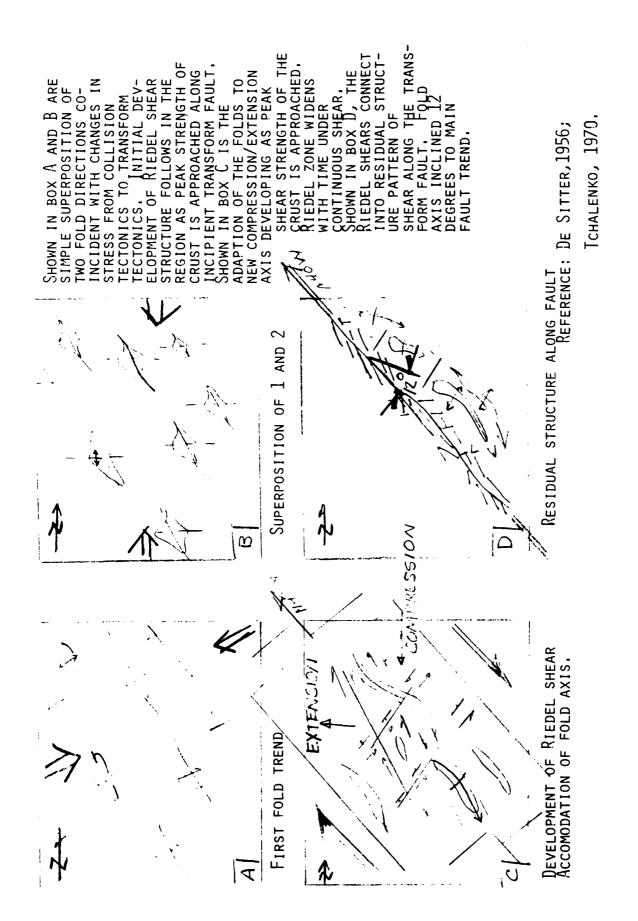
Comparison of Structure. The lower section presented in figure D-2 is a west to east section across Black Butte Dam study area shown on map plate 1. The section runs along the Glenn-Tehama County line. The section is supplemented with depth information obtained through the geophysical studies. The quality of subsurface data in the lower section is less than that obtained from the deep crustal studies of the USGS from which the upper sections are drawn. Surface geology is projected to 1,000-foot depth and indicates the Coast Range Thrust (CRT) and Stony Creek fault (SF) are steeply dipping. They are projected in the crust as flattening eastward to conform with wedge theory. The

Great Valley Sequence (Jurassic and Lower Cretaceous) is steeply east dipping and compressed against the ophiolite belt that separates the Coast Range Thrust from the Great Valley Sequence. There is no evidence of shallow east directed thrusts or folds in the lower Great Valley Sequence. A west-directed thrust exists near the Sites anticline and Paskenta nose (labelled Sites Thrust-ST). A possible thrust of Franciscan ophiolite or schist, of Great Valley Sequence exists above the basement flexure at point E. Gravity indicates this to be lighter than the overlying material. If this material is thrusted to the position shown, the wedge has not penetrated the overlying Cretaceous Great Valley Sequence. A fold exists to the east that can be included as part of the line of anticlines and domes trending north-south down the medial northern Sacramento Valley. At KR on the section, a high angle reverse fault, positive on the east, is drawn as penetrating from the basement into the upper Cretaceous. This fault has not been seen in any records of seismic reflection data but is shown as an explanation for the fold and is consistent with the wedge theory. The projection of the fault to the surface is the point of flexure on the 0.5-million-year-old Redding topography near Corning.

In summation, the Black Butte section, shown in figure D-2, has general wedge-type elements similar to Coalinga-Kettleman Hills structure but lacks the east-directed subhorizontal thrust with overlaying coseismic folds. The general axis of principal compression in the northern Sacramento Valley is north-south and not east-west. The San Andreas fault is very distant to the existing valley folds and, we believe, too distant to activate the folds/faults through stress transfer on conjugate shear lines.



Epicenters of earthquakes in central California for the year 1976, based on data f.om USGS Central California Network (McHugh and Lister, 1978). From Eaton (in press).



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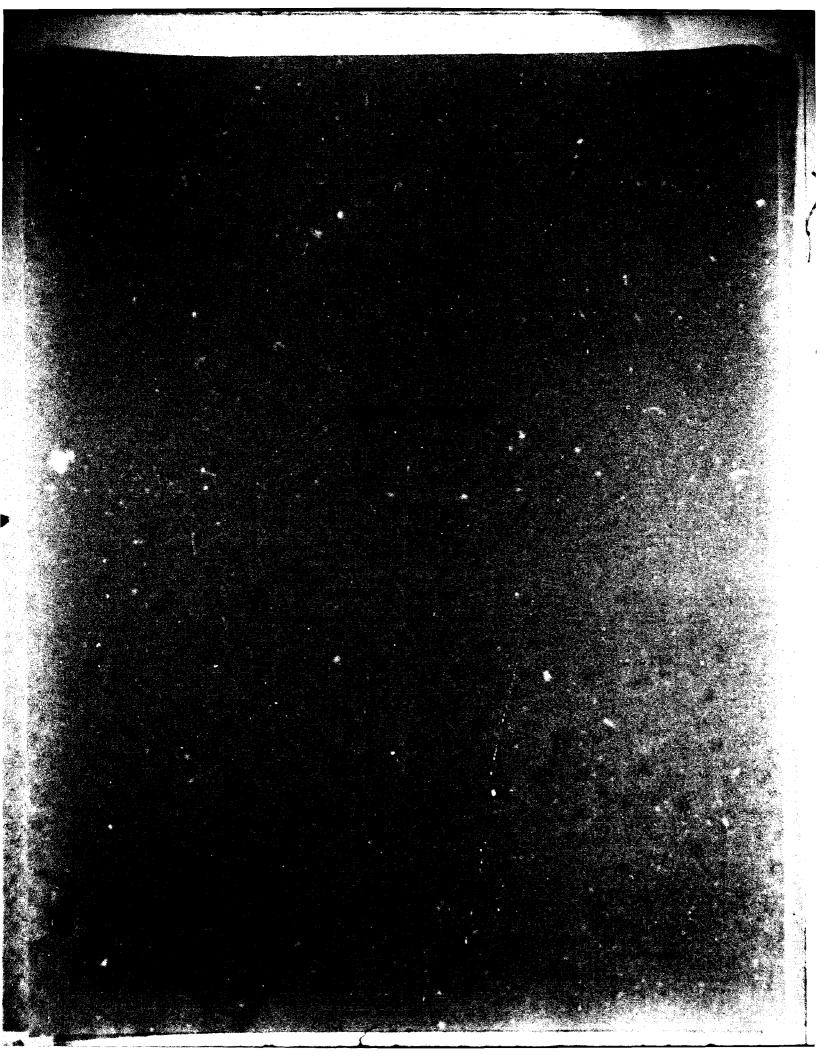
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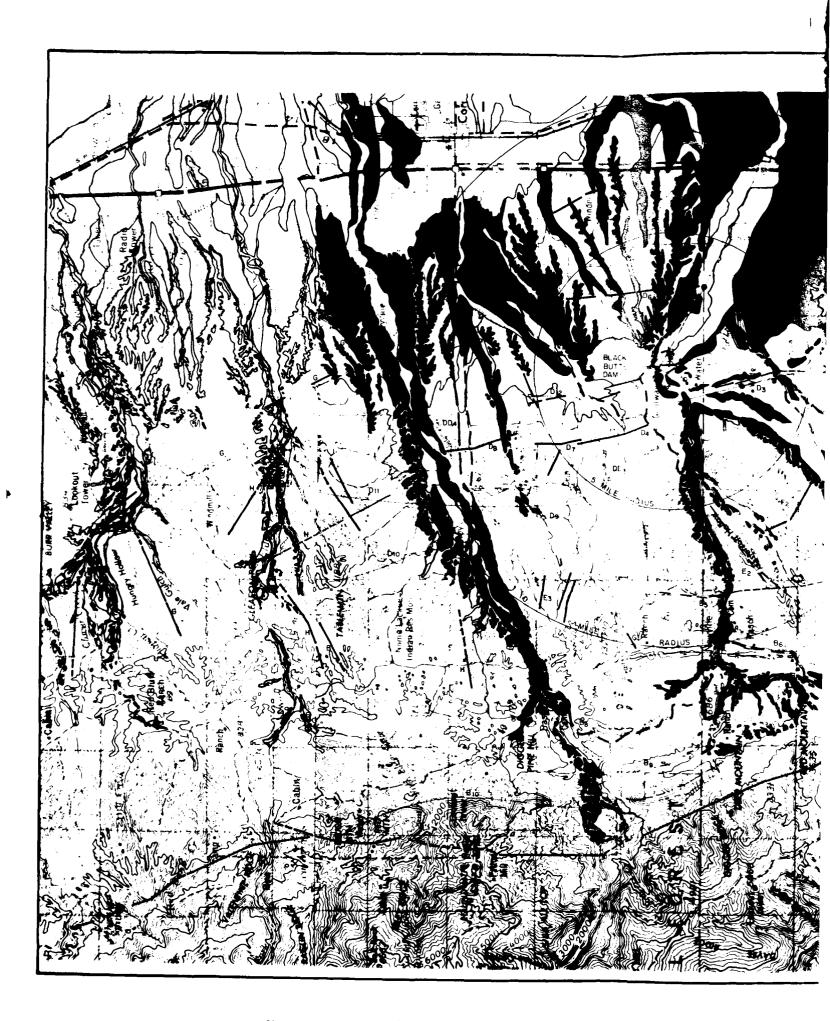
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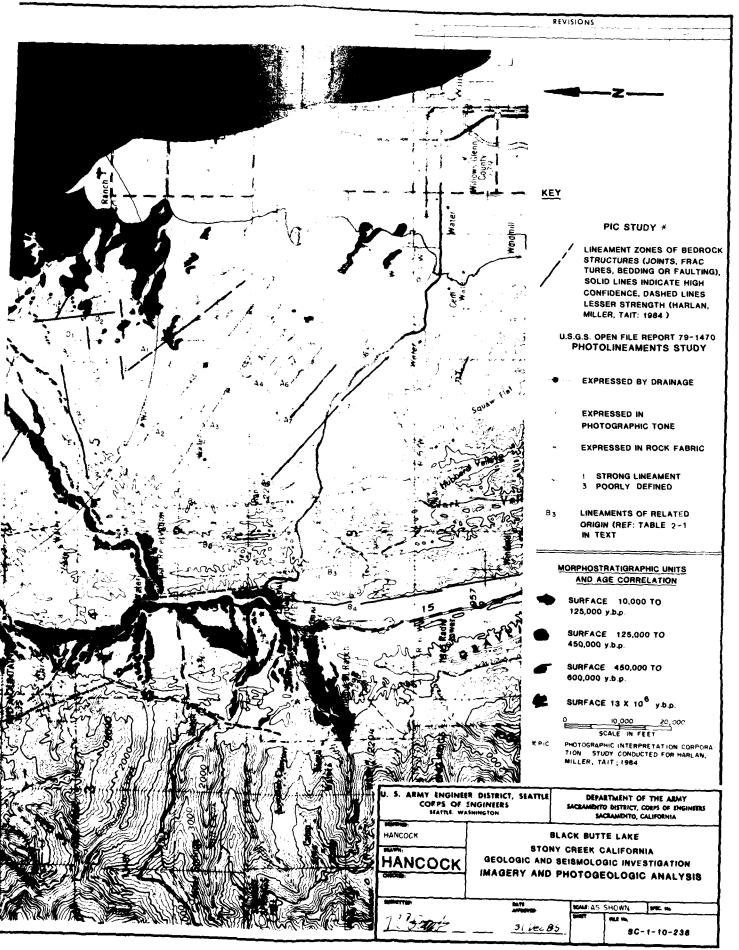
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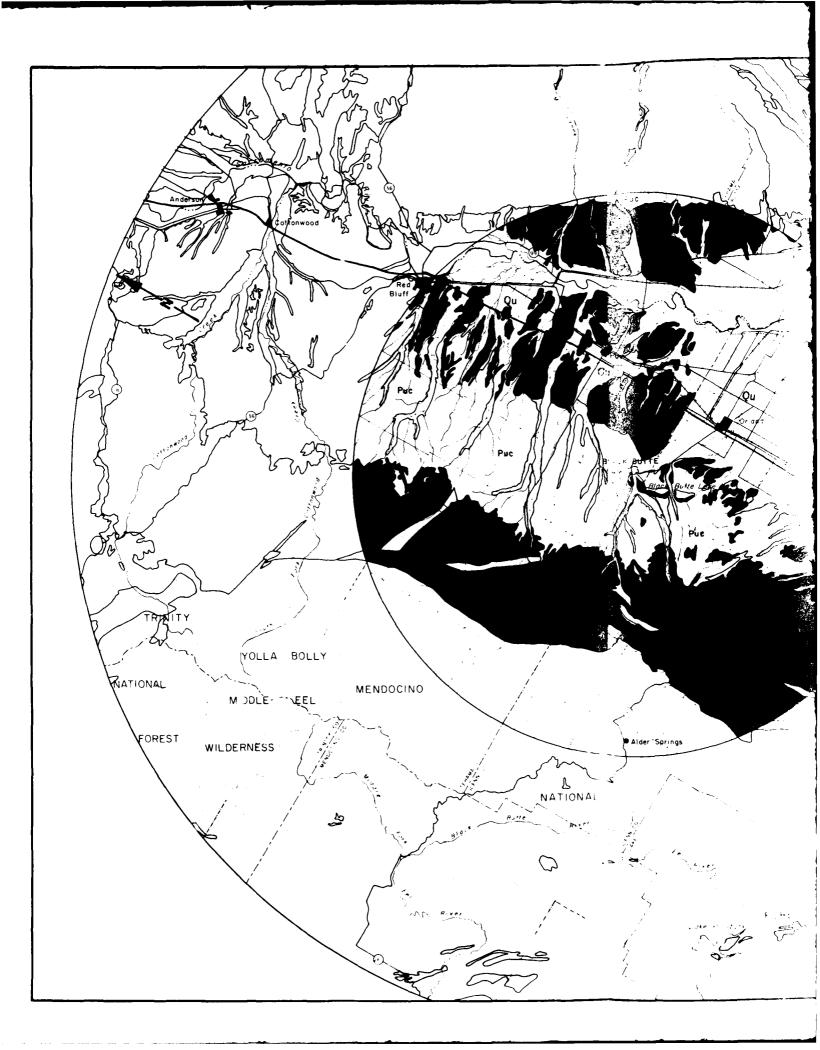
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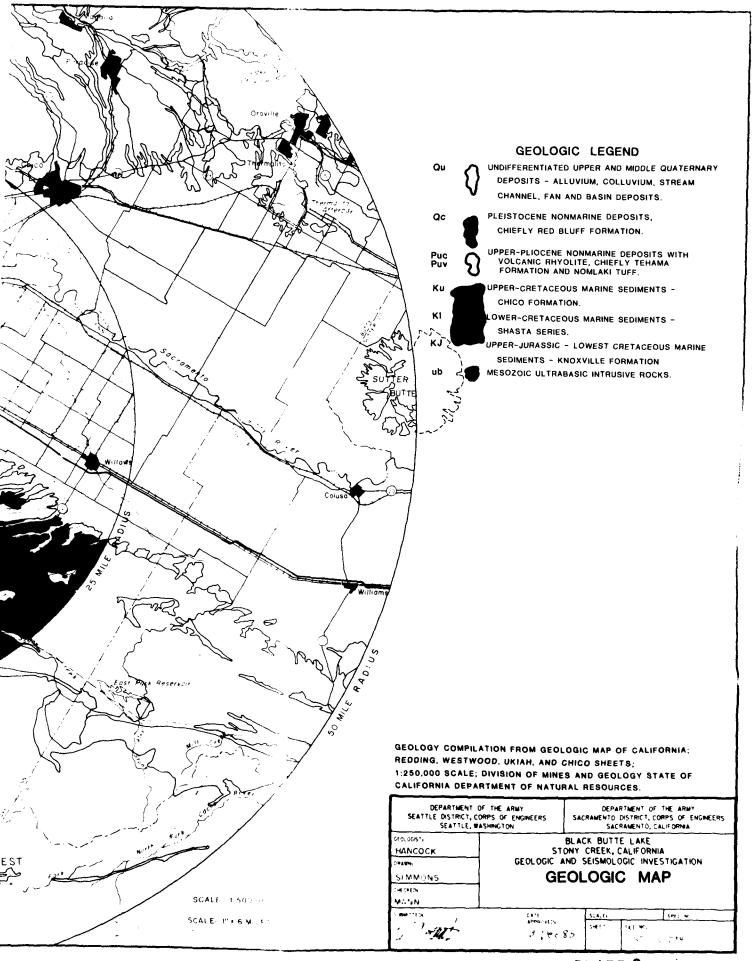
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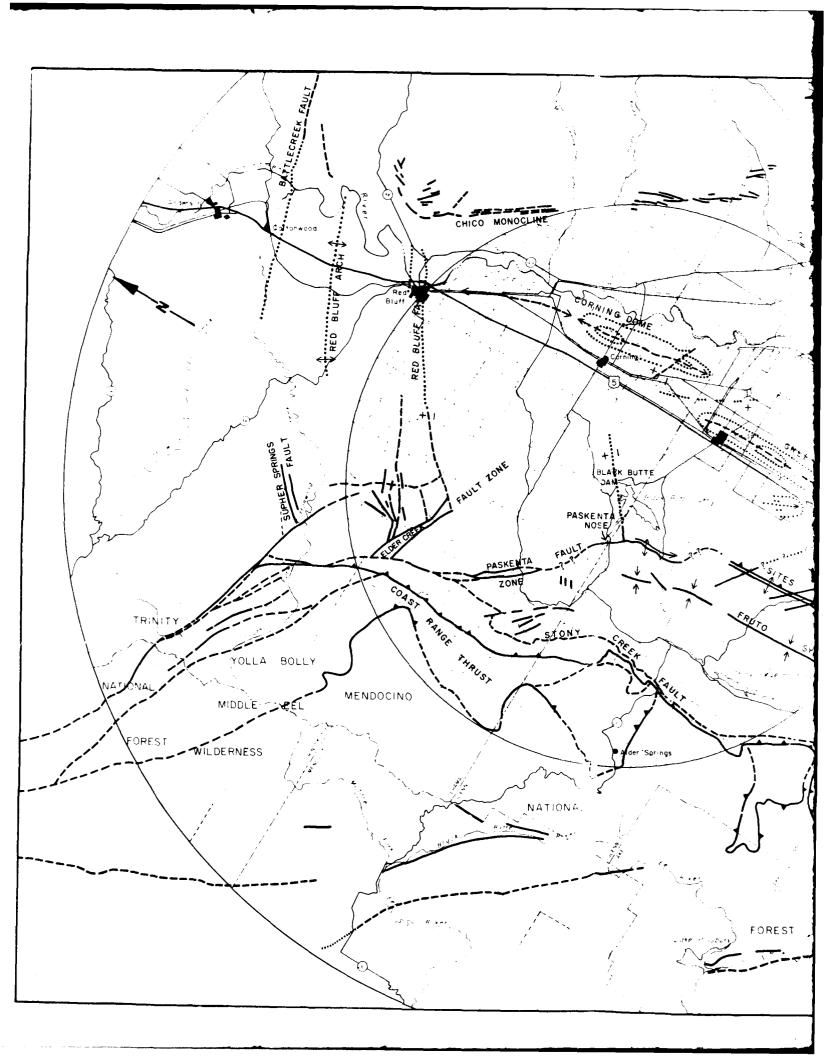


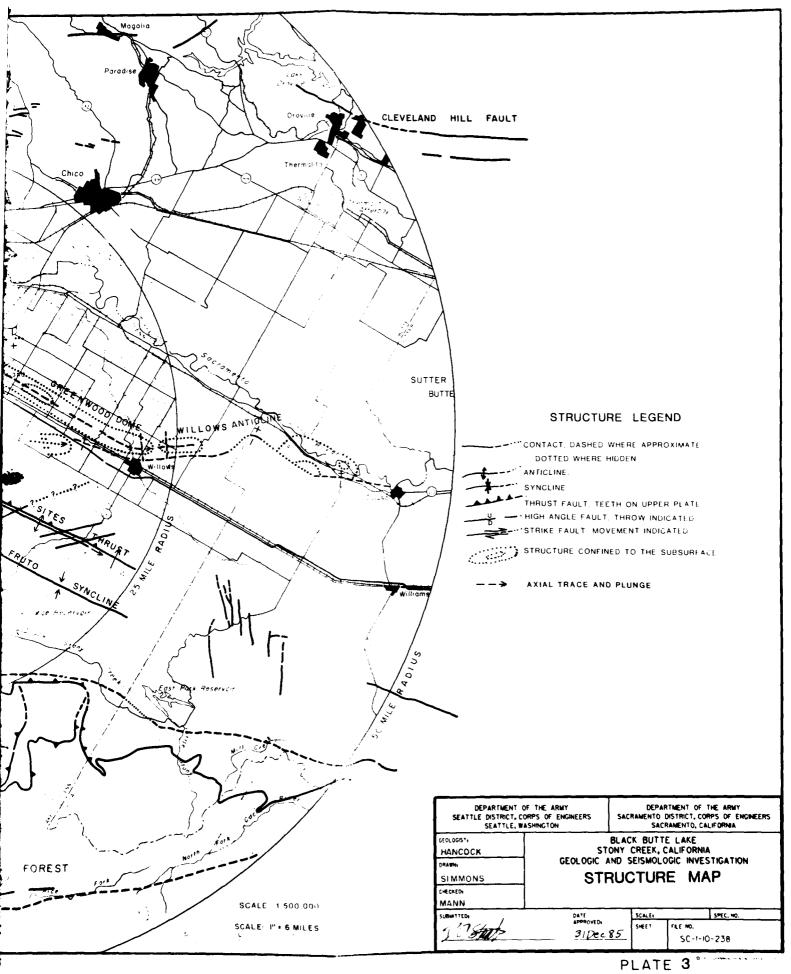


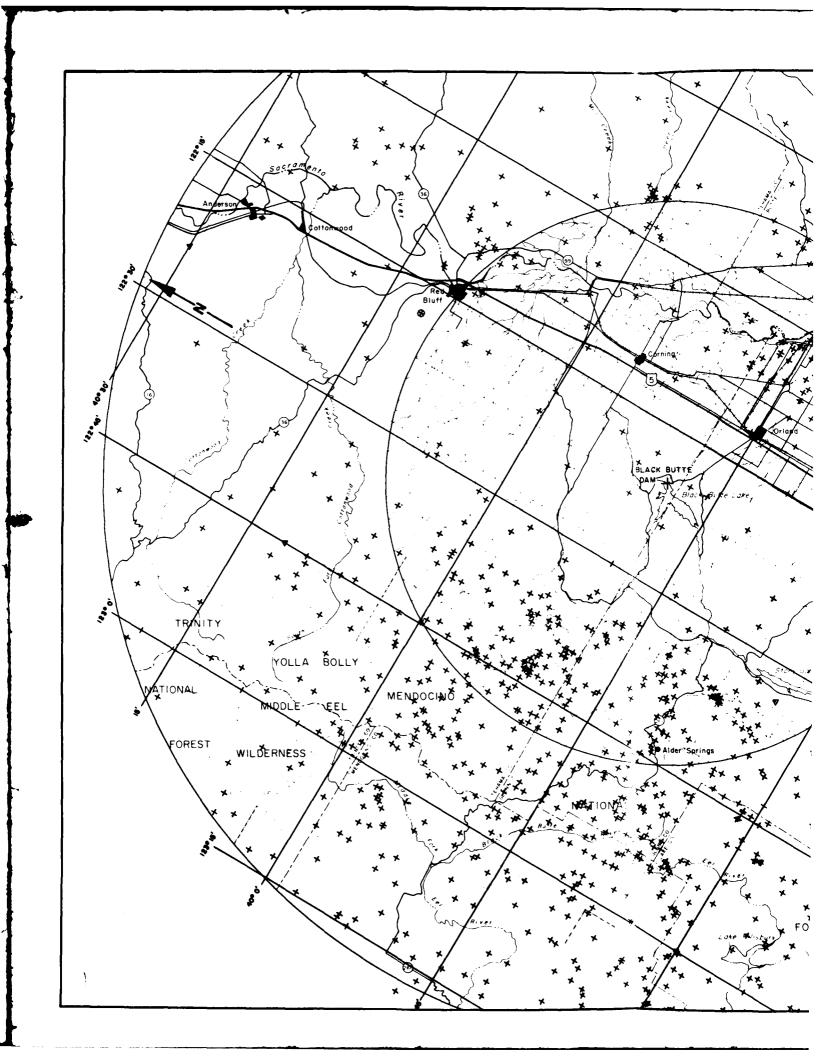


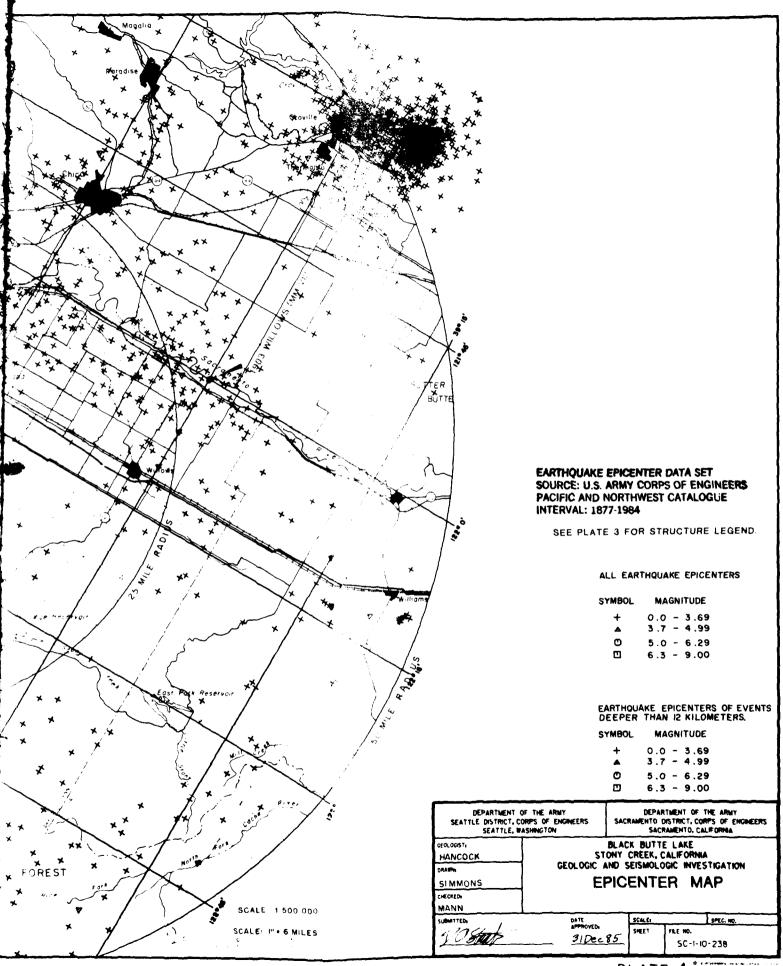


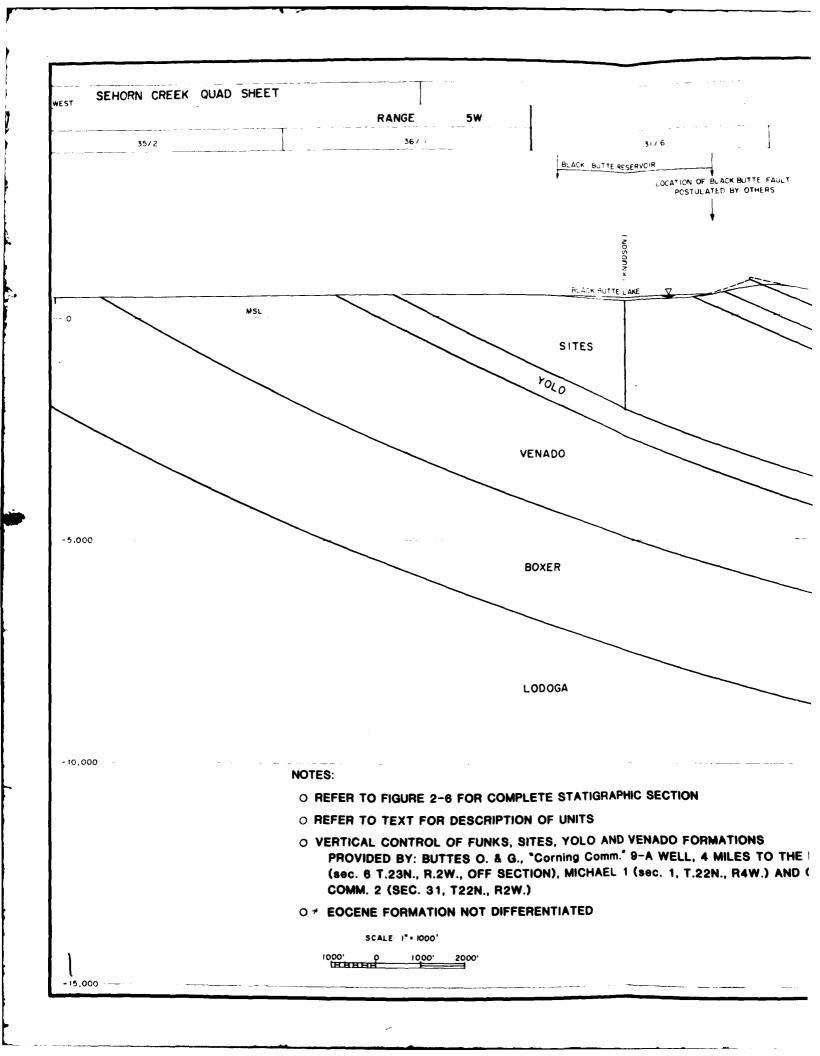


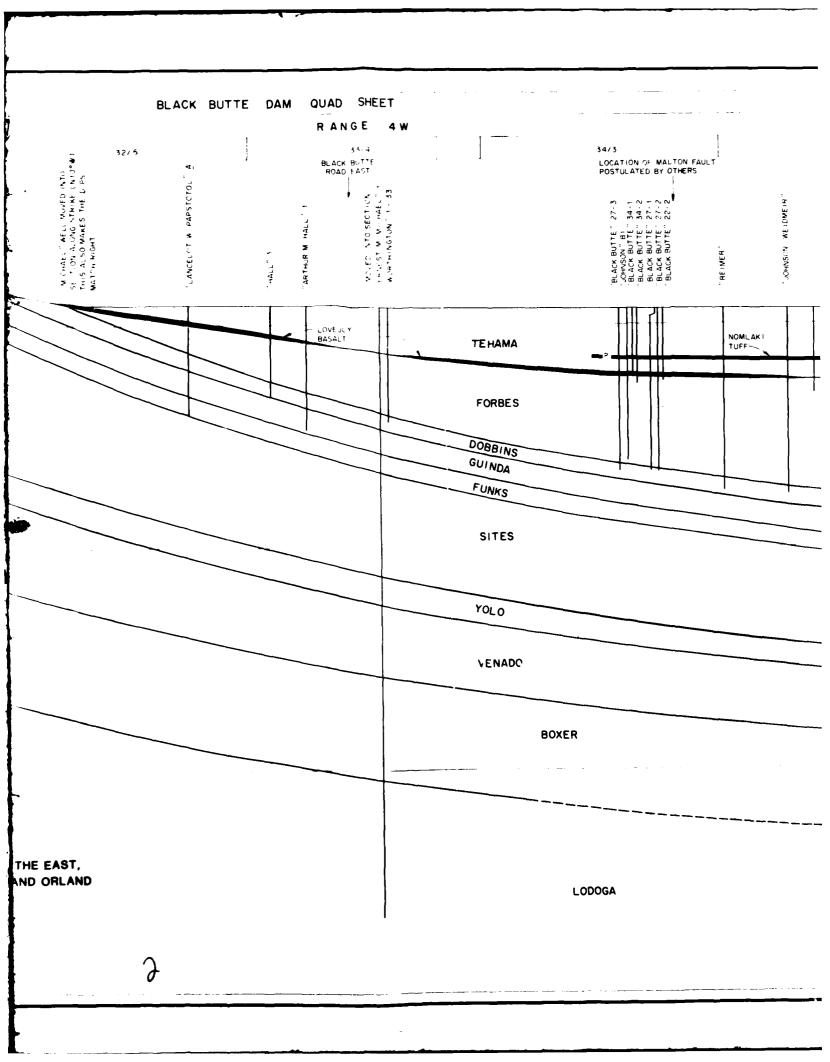


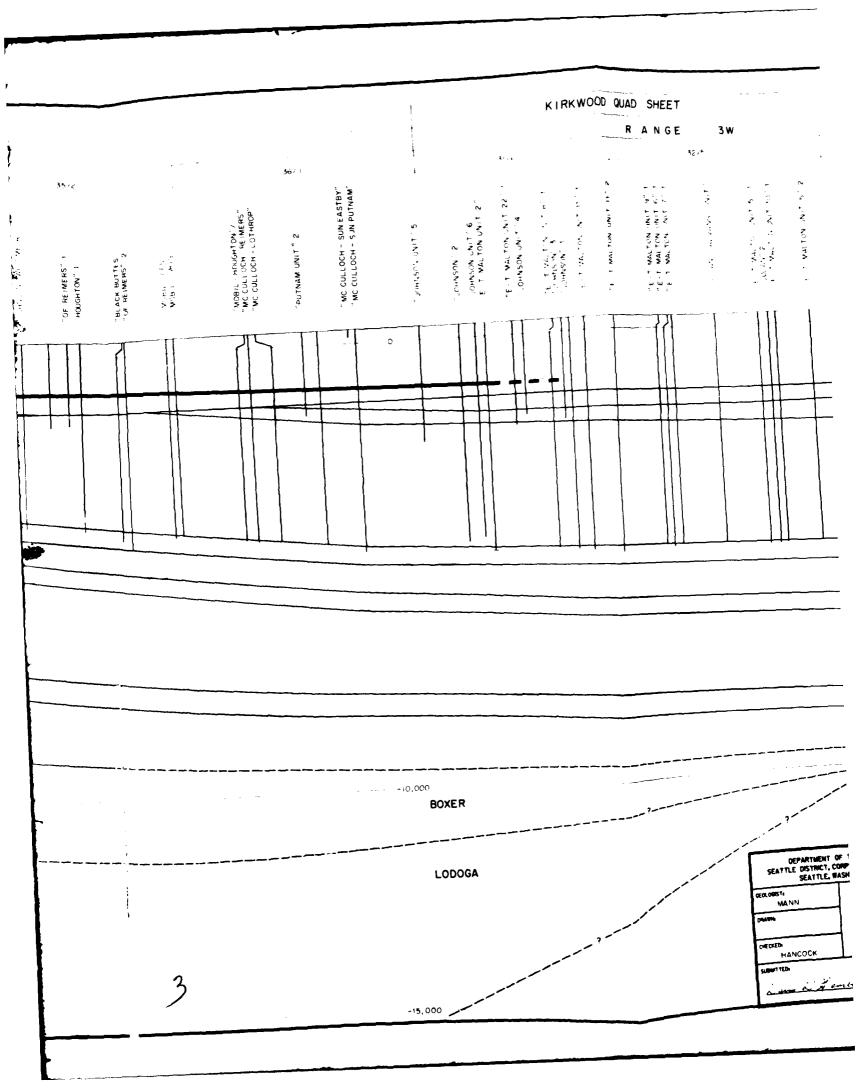


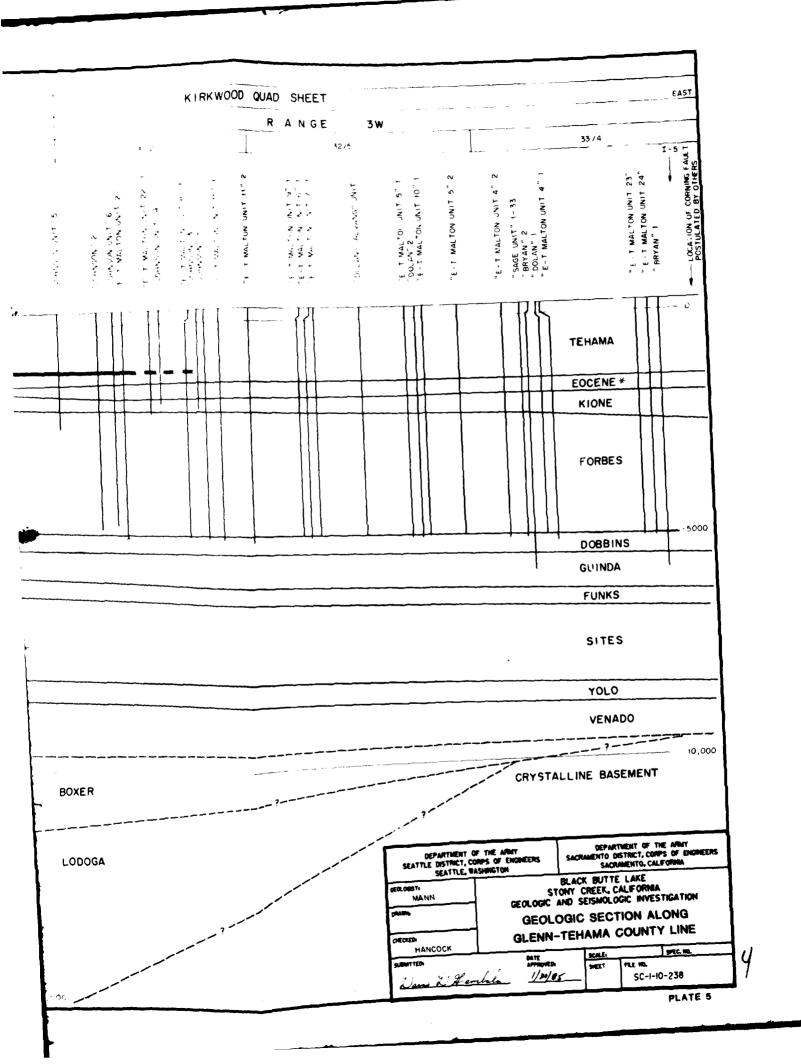


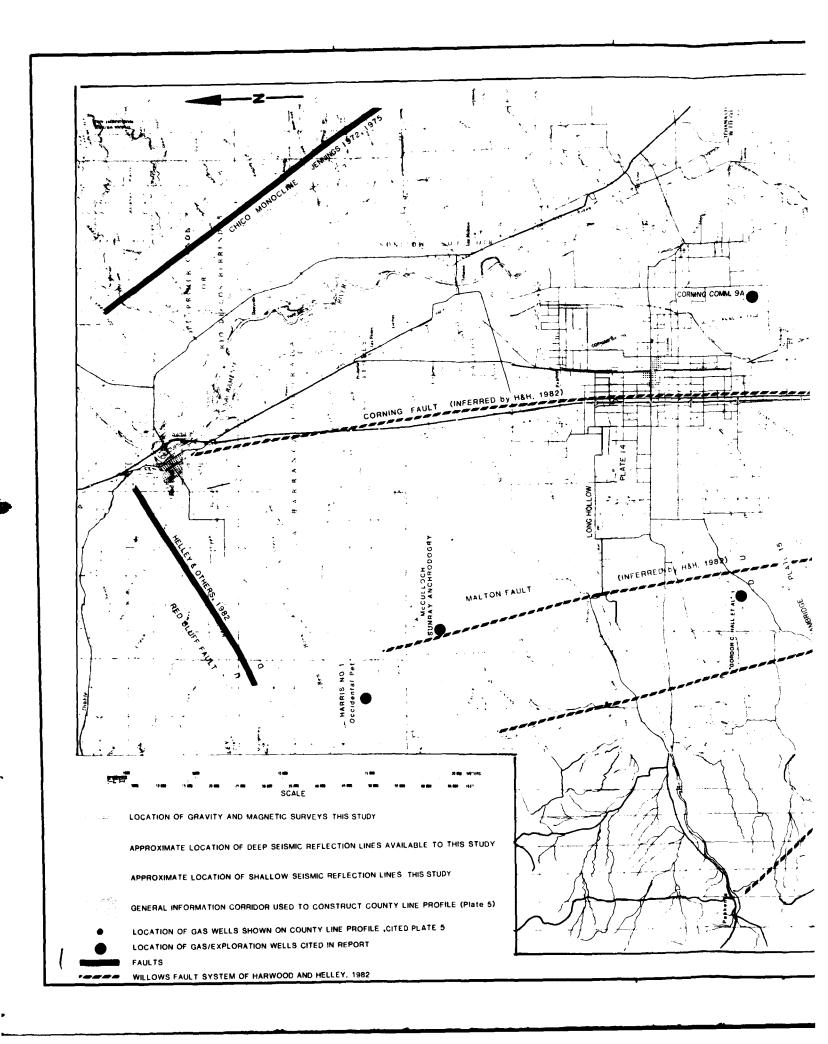


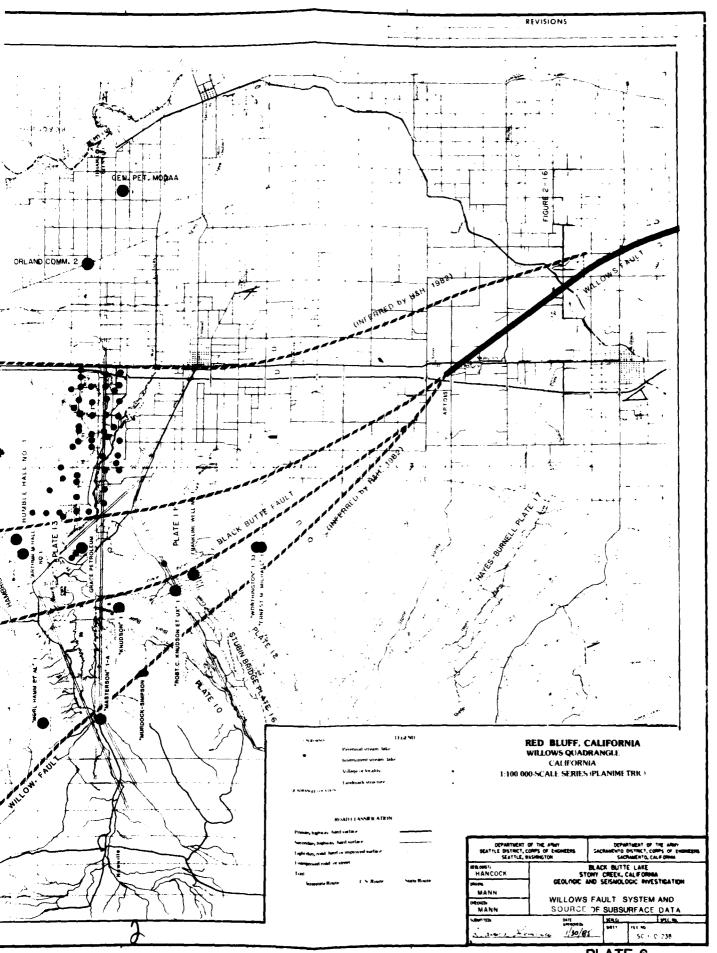


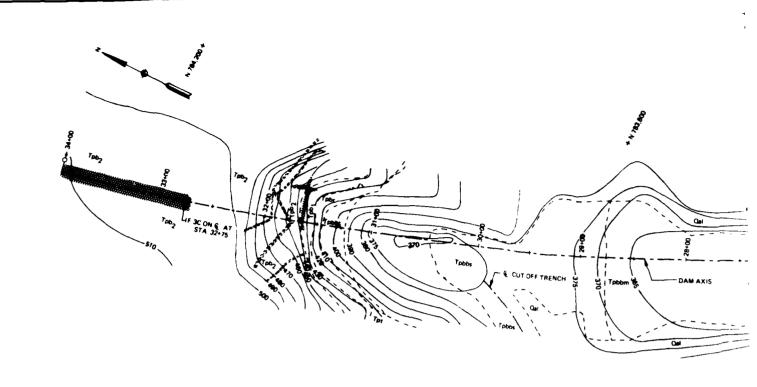


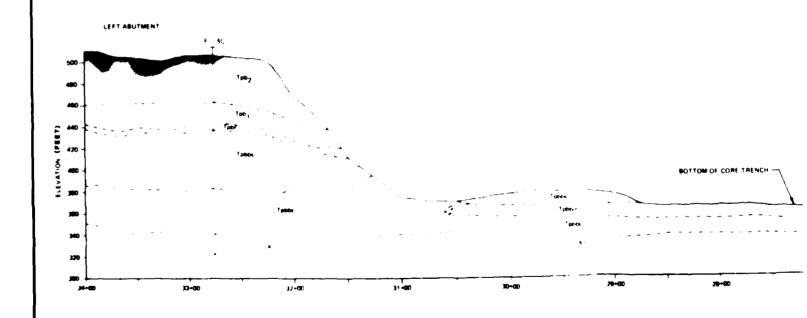












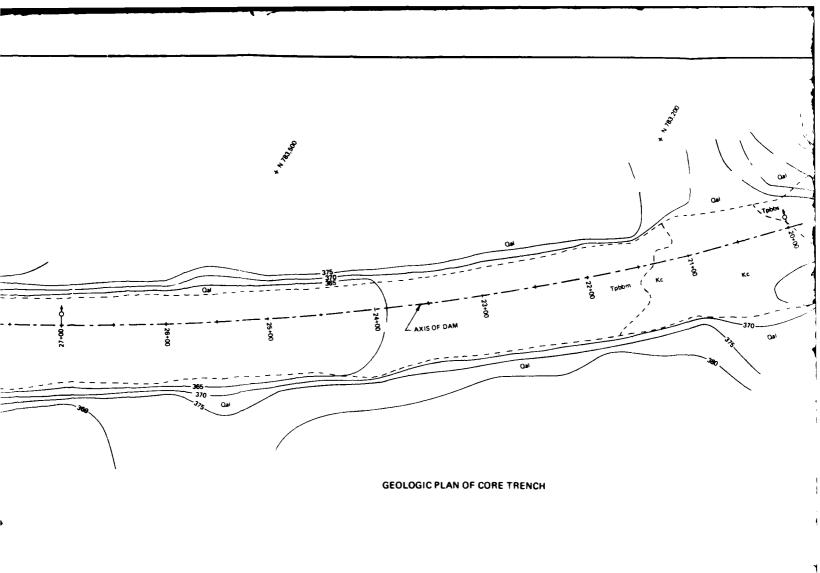
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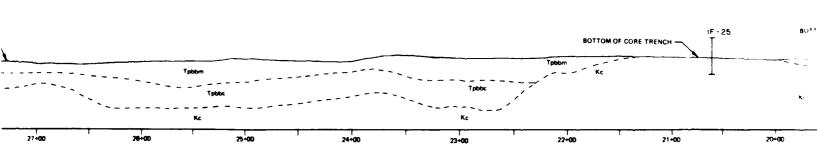
TRACE OF JOINT PLANE

BADLY FRACTURED ROCK NO PREFERRED ORIENTATION OF RACTURES

CLAY SEAM WEATHERED FRACTURES OF FRACTURE IN FILLING

- JO OVERBUNDEN IME ENT ALLIE ME HEAM CHANNEL AND FLOOD PLAN HEAM CHANNEL AND A FEW COBBLES, FREGIENT HEAT AT HIS ATHE AT THE STLTY SAND DEPOSITS FAIRLY MELL HADEN IP TO KIND SEE MARD, STREAM ROUNDED, COMMONEL WOLLDNING AND SANTY TE WERE KEIN QUARTZ AND GRANITIC HOLKS PYRE ONSER LATE, NEMENTED, PERVIOUS
- THE AN A. . . A. DEPOS TOMPOSED OF LEAN TO FAT CLAY WITH VARYING PROFIDE ONS OF LINE TO TOARSE SAND, STREAM ROUNDED GRAVELS, AMELIANGULAR FRAGMENTS AND BOULDERS OF BASALT, TAN, FIRM, MODERATELY CONSCLUDATED NOT INDURATED





### **GEOLOGIC SECTION ALONG CENTERLINE OF CORE TRENCH**

BASALT - DARK GRAY TO BLACK, APHANITIC TO FINE GRAINED, DENSE, HARD WHEN UNWEATHERED, MASSIVE, BRITTLE AND GLASSY. TWO FLOWS PRESENT - TOBI, LOWER FLOW CHARACTERIZED BY LIGHTER COLOR, COARSER TEXTURE, AND BLOCKY JOINTING; Tpb2, UPPER FLOW DISTINGUISHED BY DARKER COLOR, COLUMNAR STRUCTURE, FINER TEXTURE AND EXTREME BRITTLENESS

VOLCANIC BRECCIA - BROWN TO GRAY, COMPOSED OF BAGALT BOULDERS JP TO 10 OR MORE IN DIAMETER, SOME VESICULAR, IN A MATRIX OF VOLCANIC EJECTA WITH INCLUSIONS OF MUDSTONE FRAGMENTS ARE GENERALLY WELDED TOGETHER, DEPOSIT IS PERVIOUS AND STAINED THROUGHOUT WITH IRON OXIDE MATRIX IS SOFT TO MODERATELY HARD SOME ZONES ARE LOOSE AND CRUMBLY.

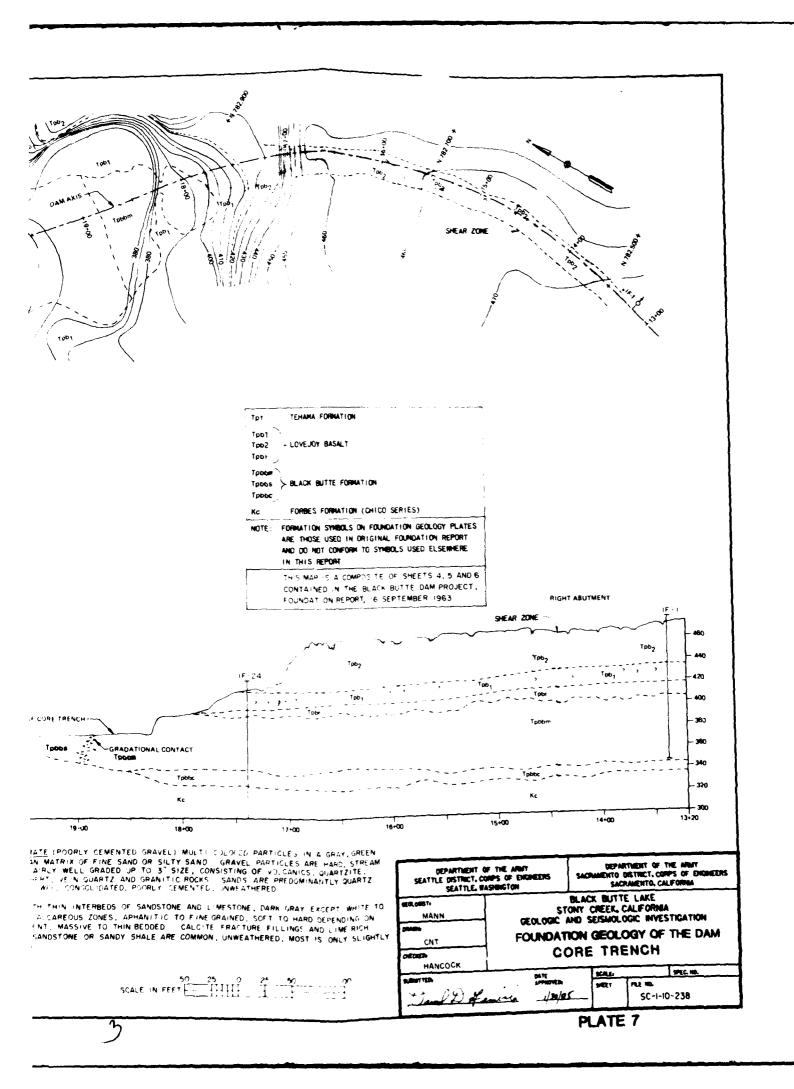
MUDSTONE - COMPOSED OF LEAN CLAY, SILT AND FINE SAND, BLUE, GRAY, GRAY-GREEN OR TAN, MASSIVE OR IRREGULARLY BEDDED, SOFT TO MODERATELY HARD - MOST CAN BE SCRATCHEL WITH A FINGERNAIL OR CARVED EASILY WITH A KNIFE - MOIST TO DRY, WELL CONSOLIDATED WITH LITTLE OR NO INDURATION

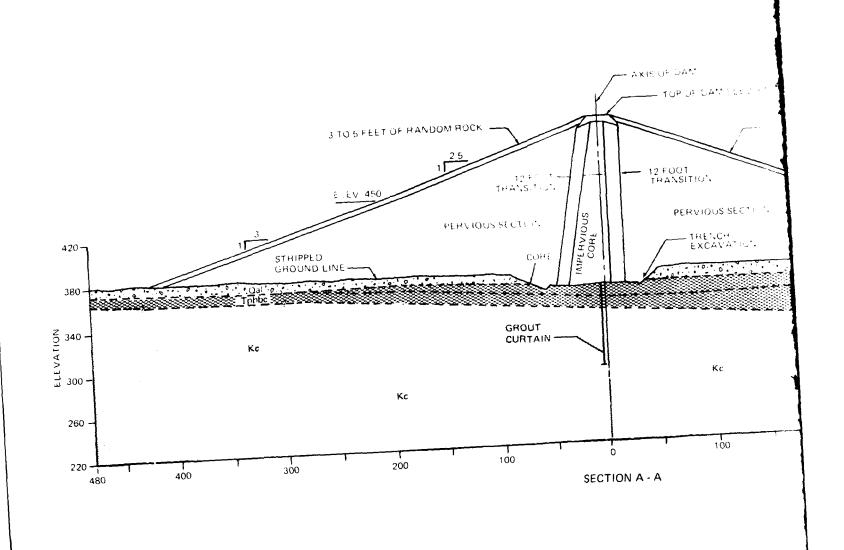
Tpbbs SANDSTONE - BLUE, GRAY, GRAY-GREEN, FINE GRAINED, MASSIVE. SOFT, FRIARLE, EASILY CRUSHED, COMPOSED MOSTLY OF ROUNDED QUARTZ PARTICLES, FREQUENTLY CONTAINS A SMALL AMOUNT OF LEAN CLAY OR SILT, WELL CONSOLIDATED, POORLY TO MODERA-TELY CEMENTED, UNWEATHERED.

CONGLO Tpbbc BLUE D

DE PER

SHALE BROWN LIME ZONES FRACTI

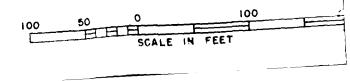




RECENT ALLUVIUM

BLACK BUTTE FORMATION

FORBES FORMATION



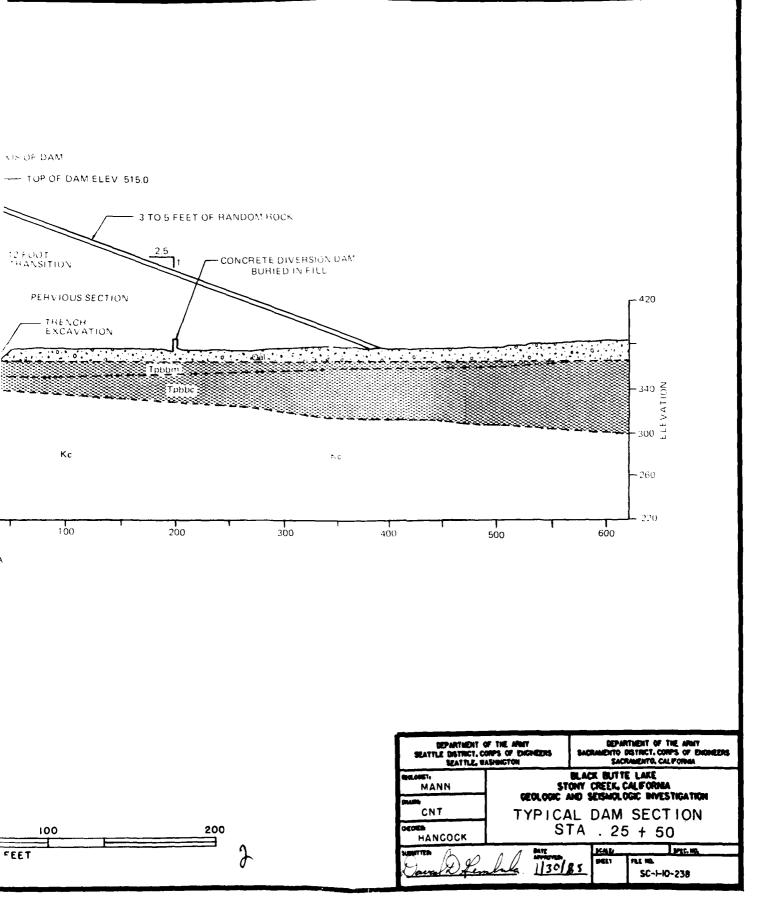
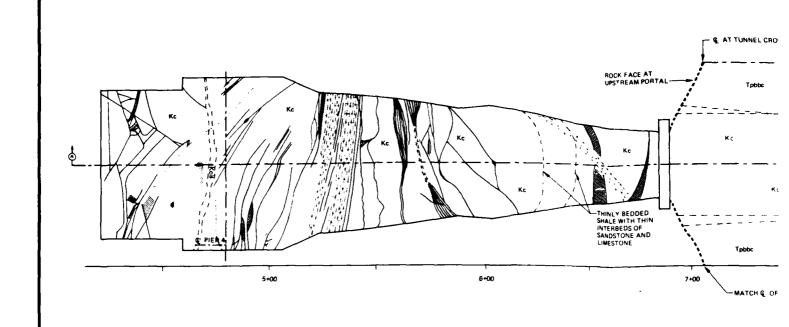
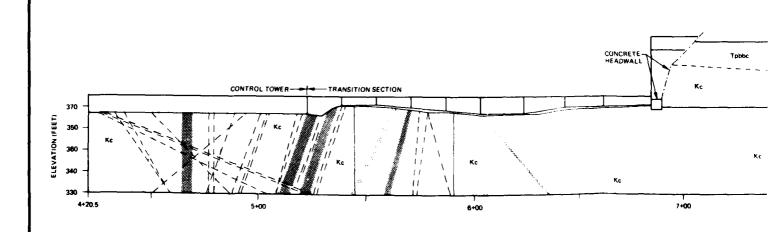
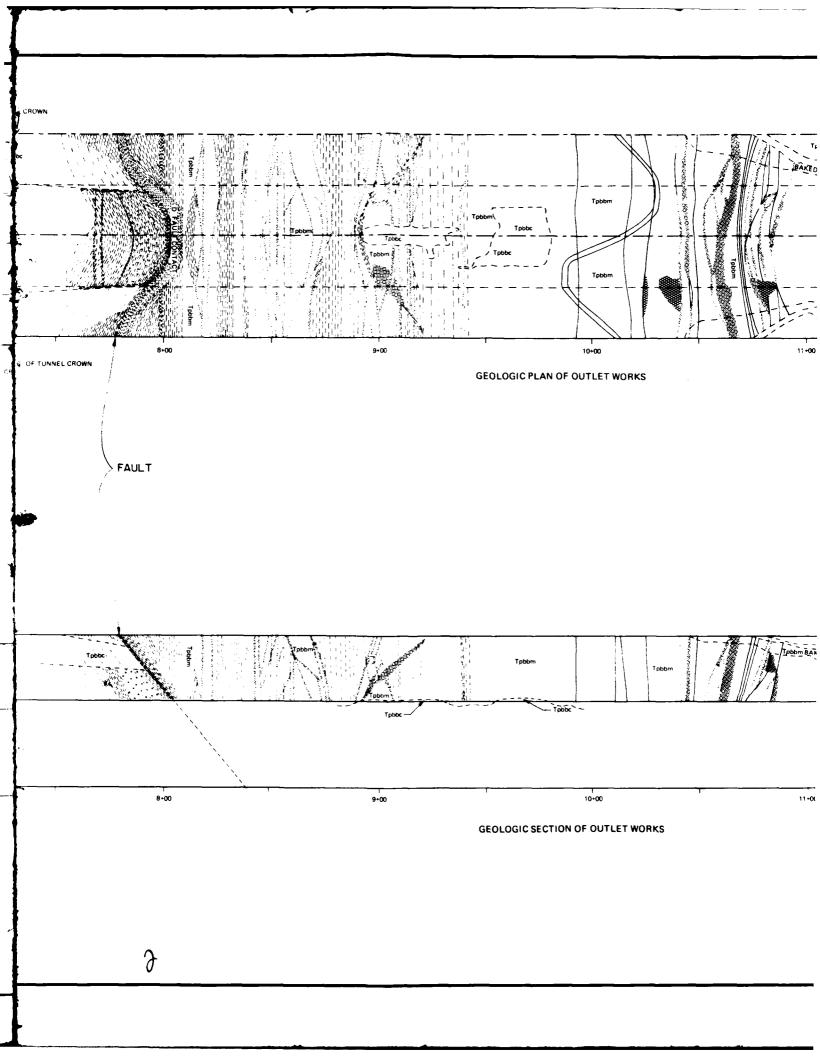
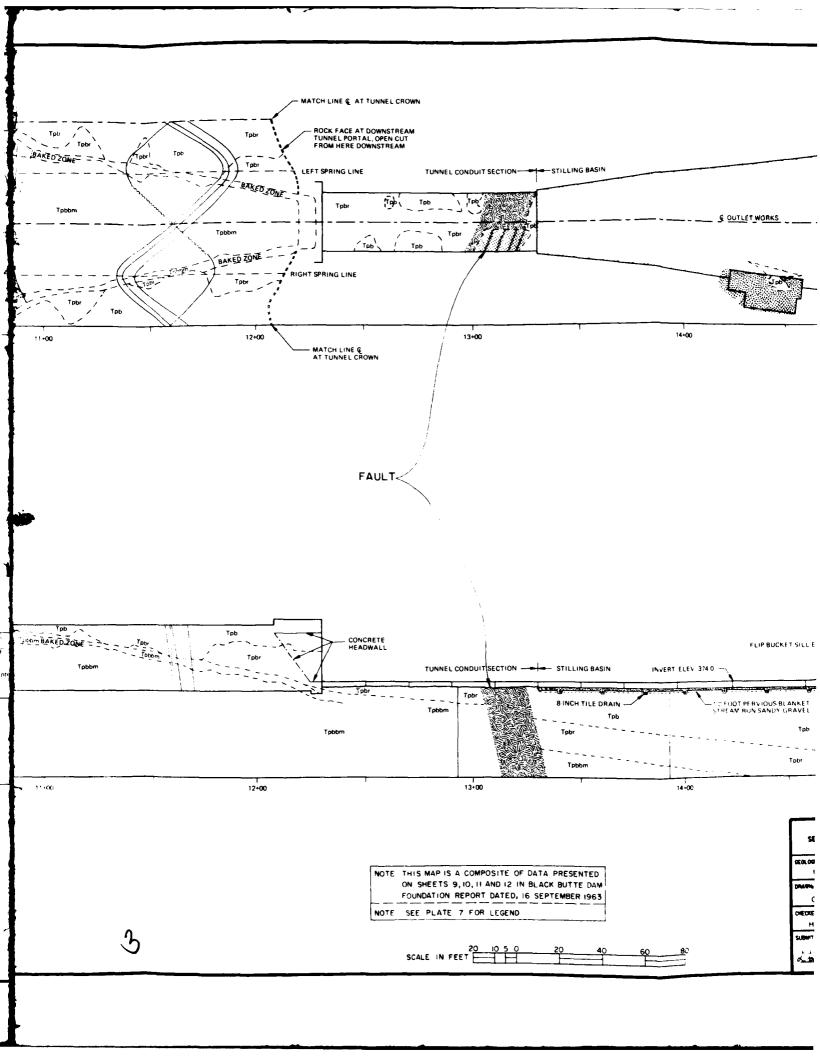


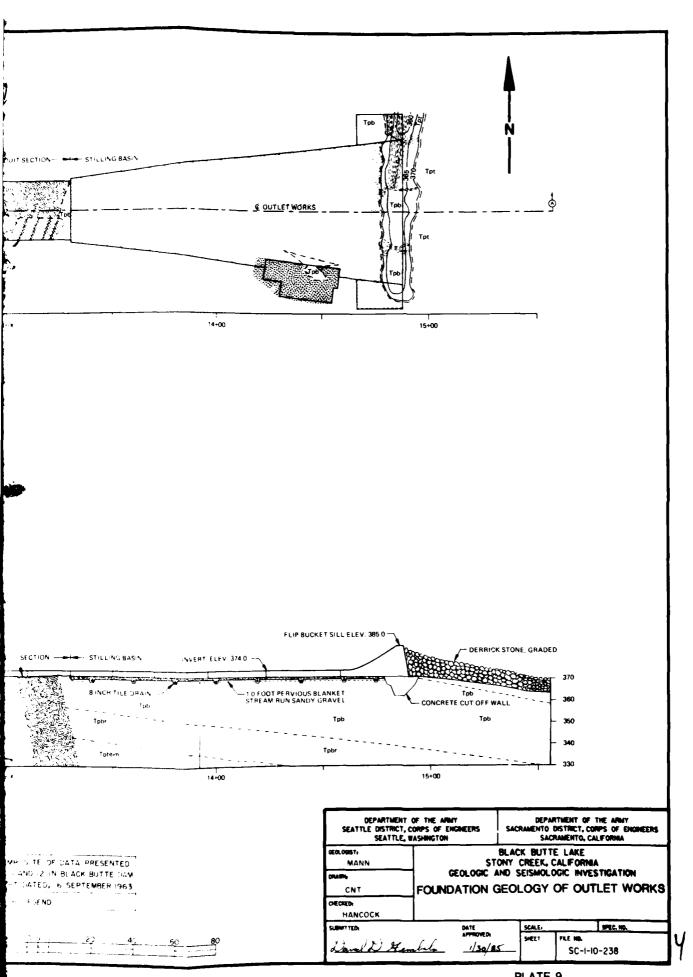
PLATE 8

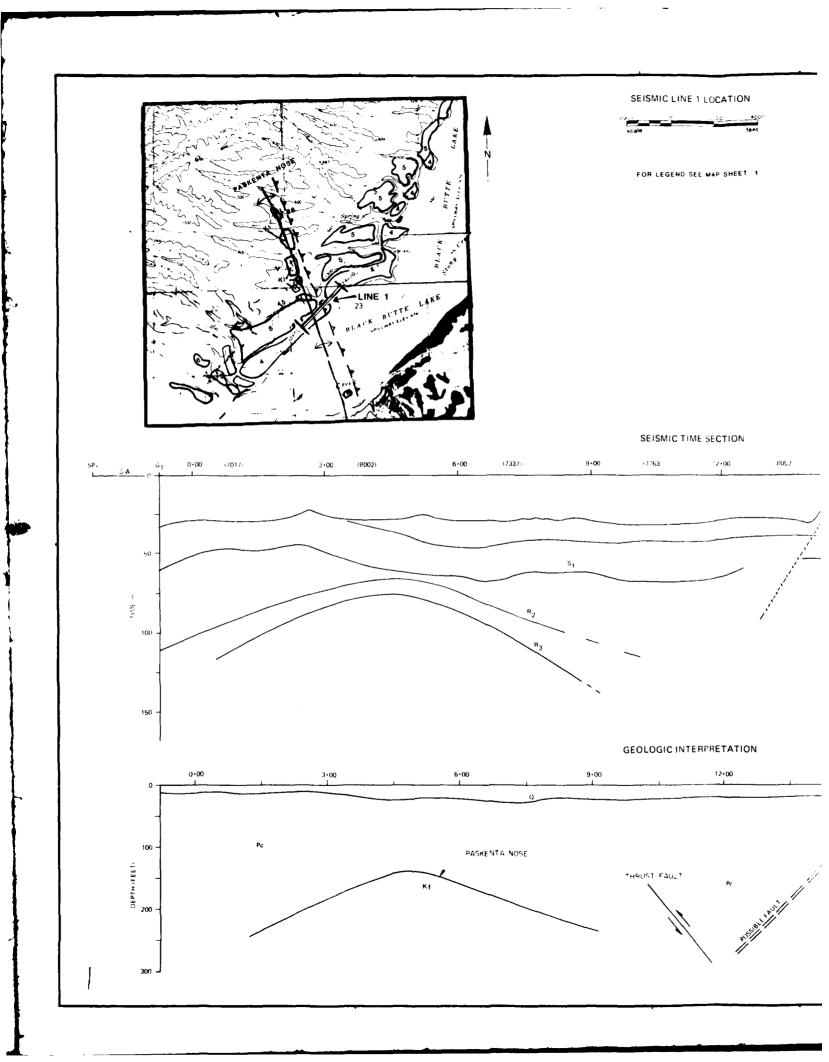


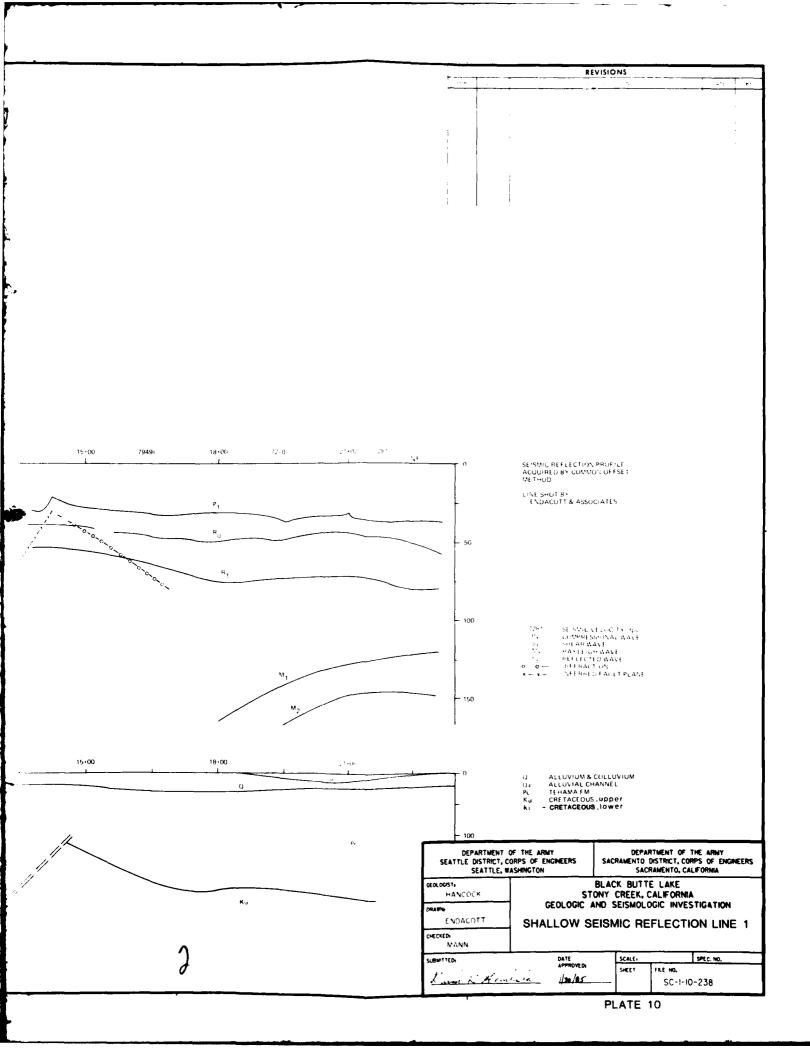


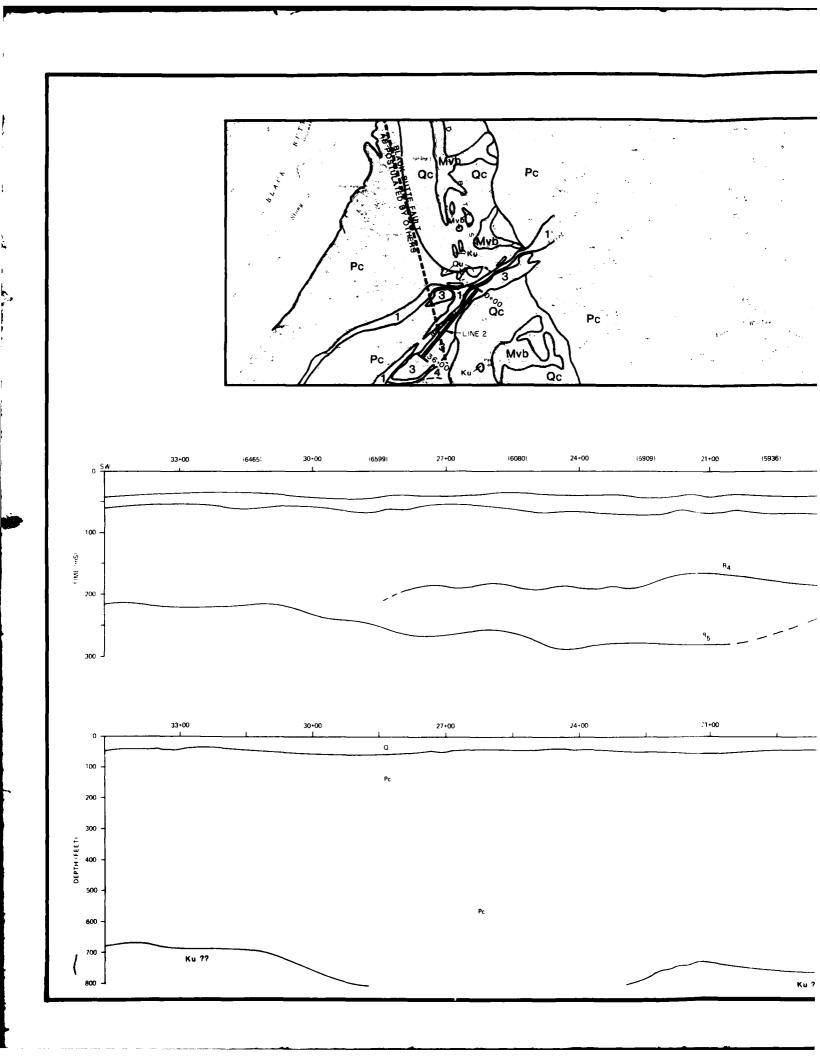


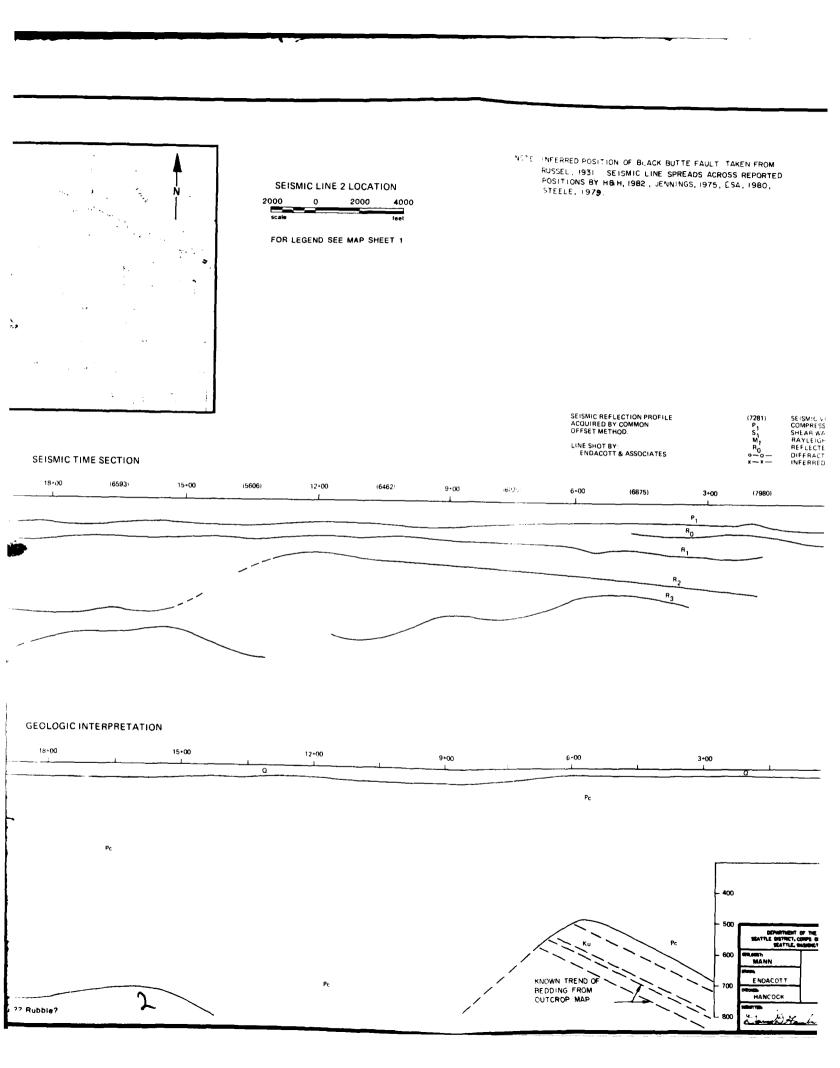


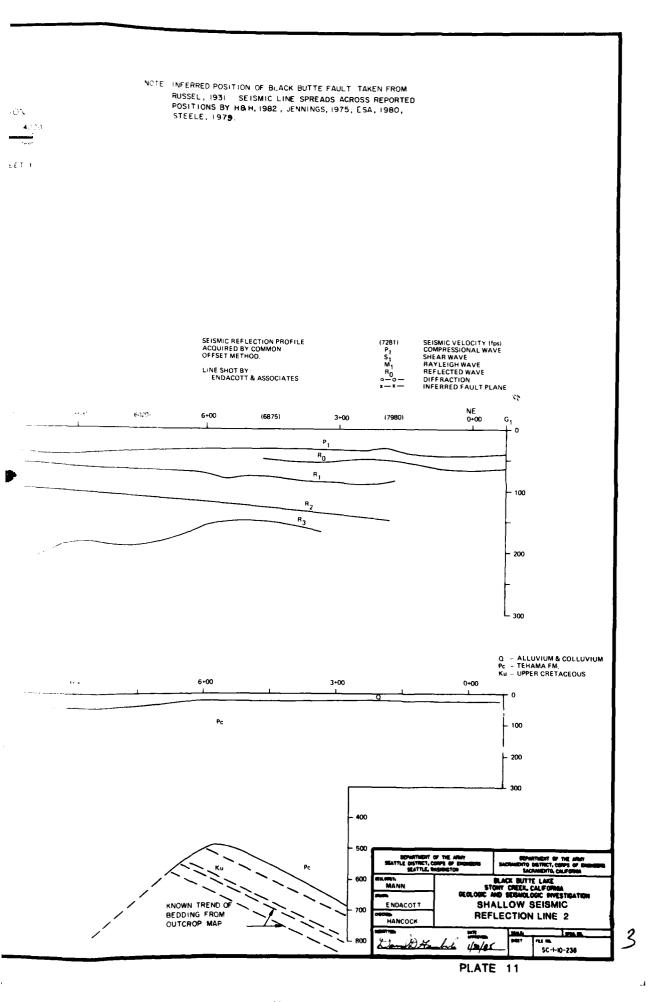


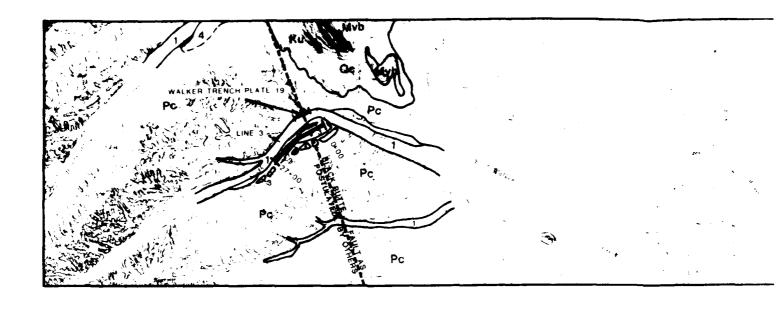


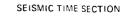


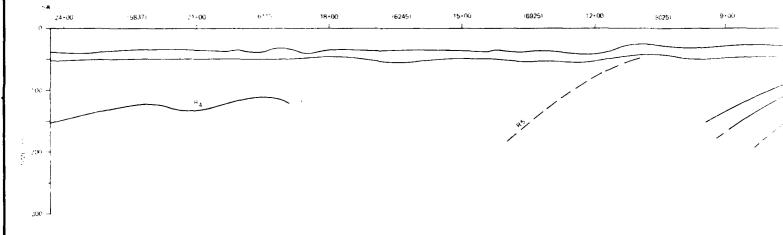




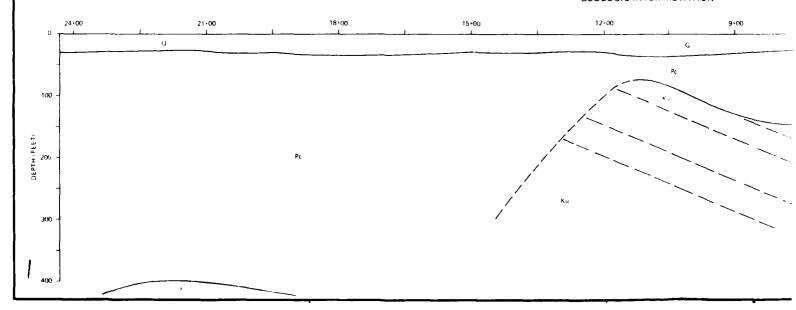


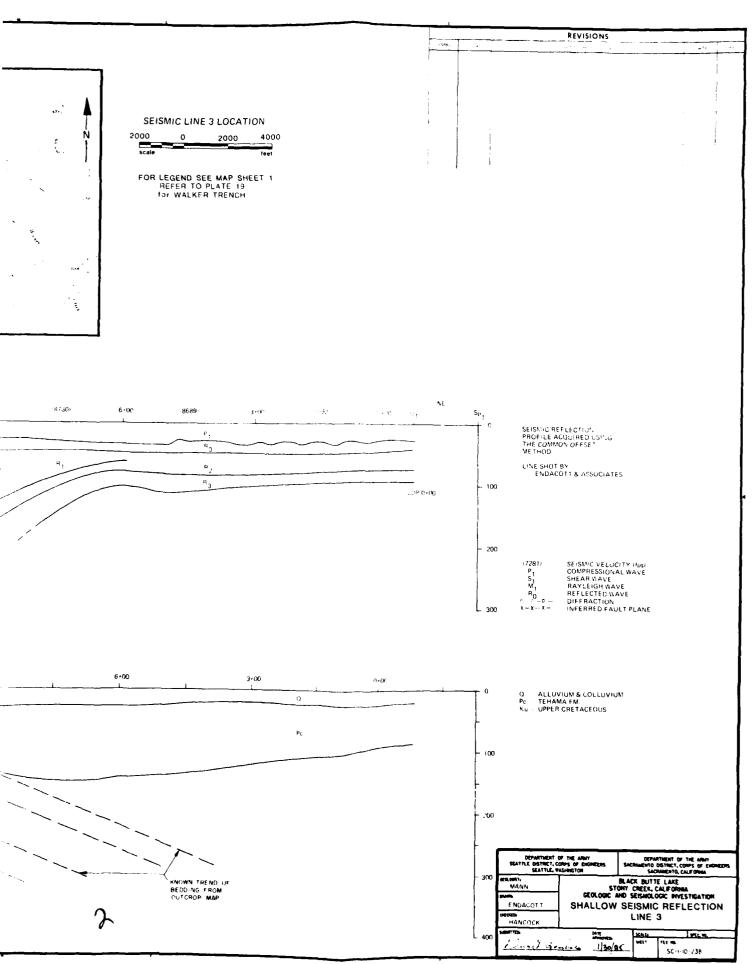


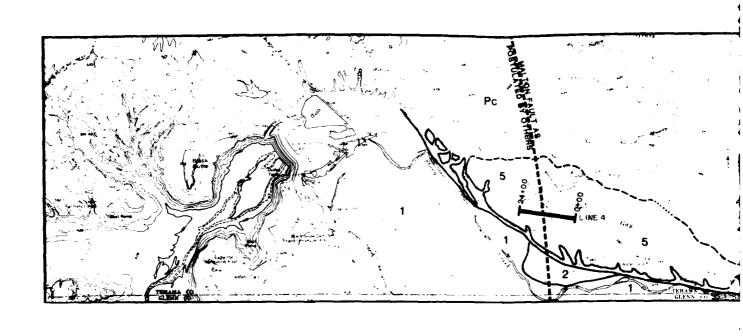




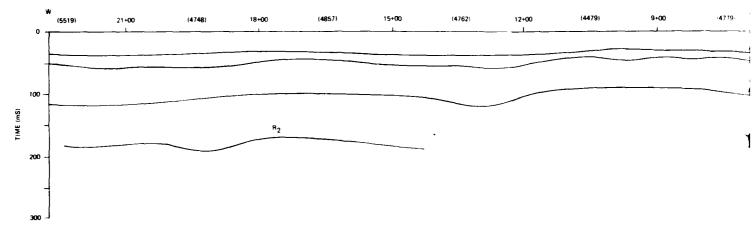
#### GEOLOGIC INTERPRETATION



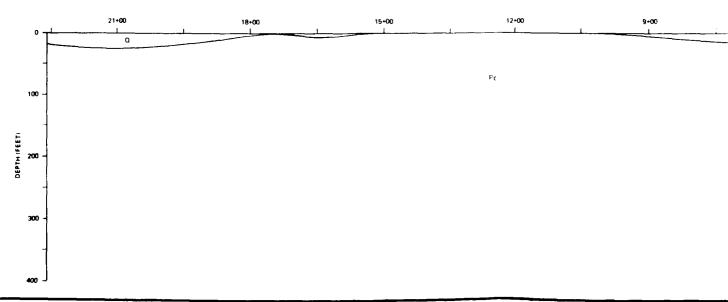


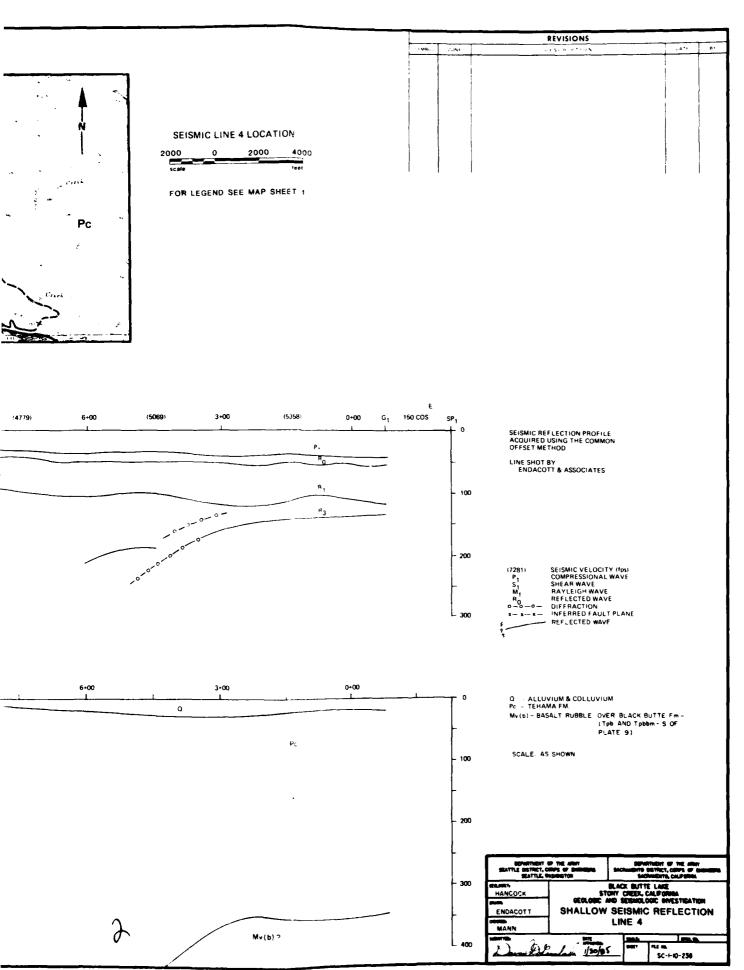




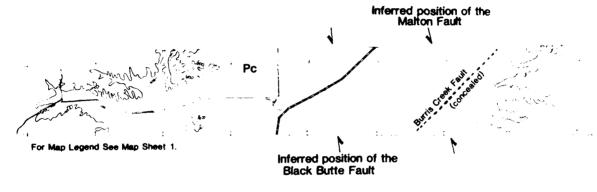


#### GEOLOGIC INTERPRETATION

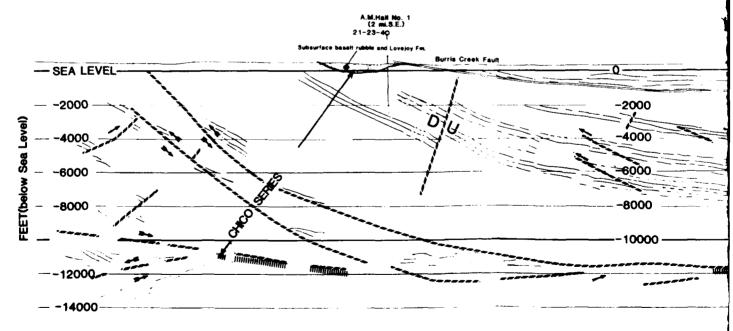




0 2000
Approximate scale in feet.
Contour interval in Feet.
Map strip location shown on Plate 6



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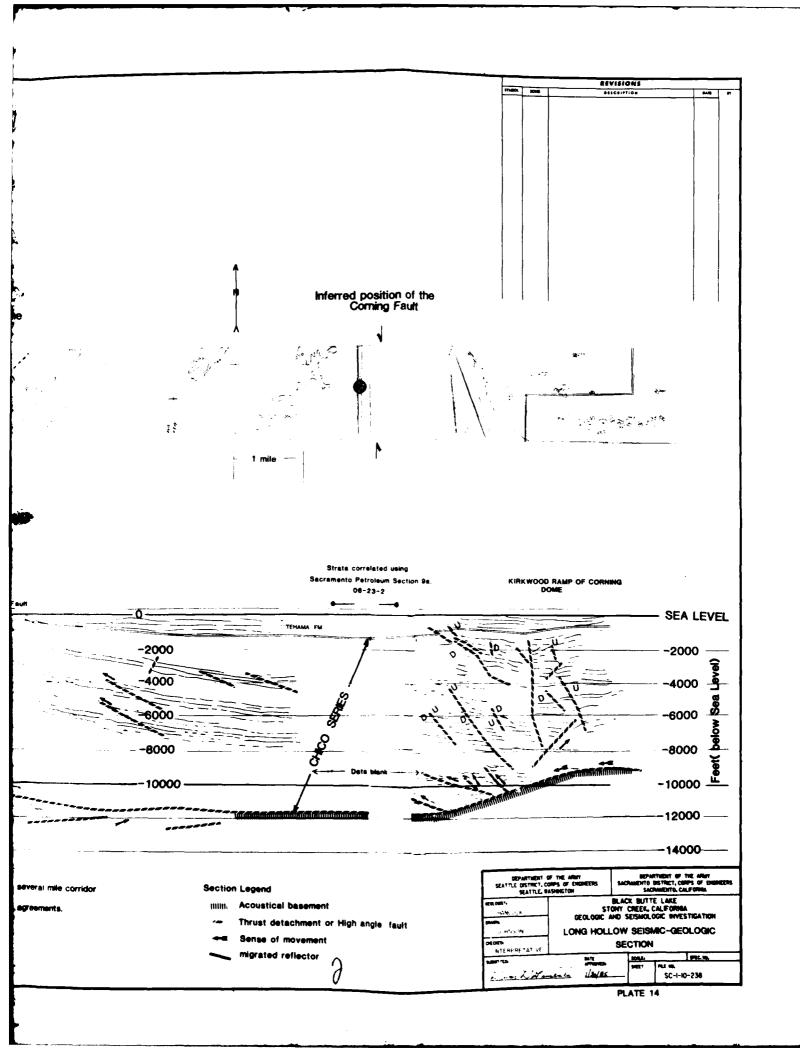
Seismic section constructed from geology, well and reflection data located along several mile corridor either side of the center of map shown above.

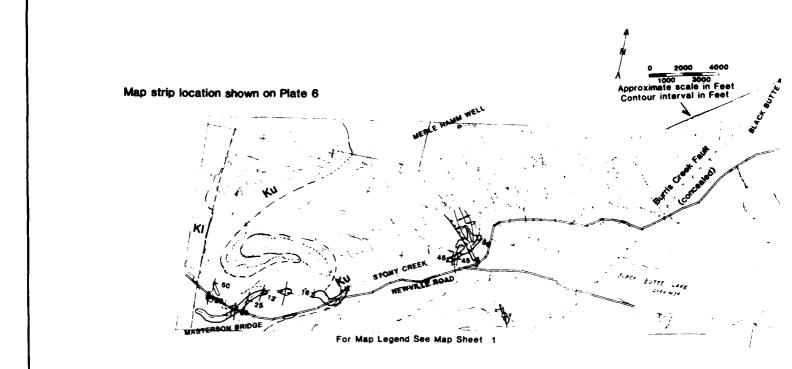
Exact location of reflection data not shown due to proprietary nature and lending agreements, Horizontal to vertical scale has 1.2 V. to 1 H. exaggeration.

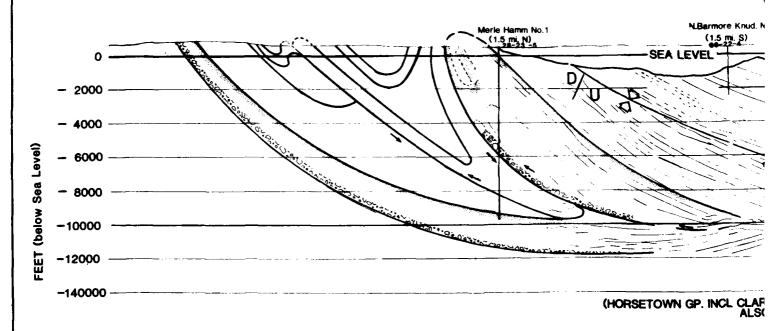
Section Le

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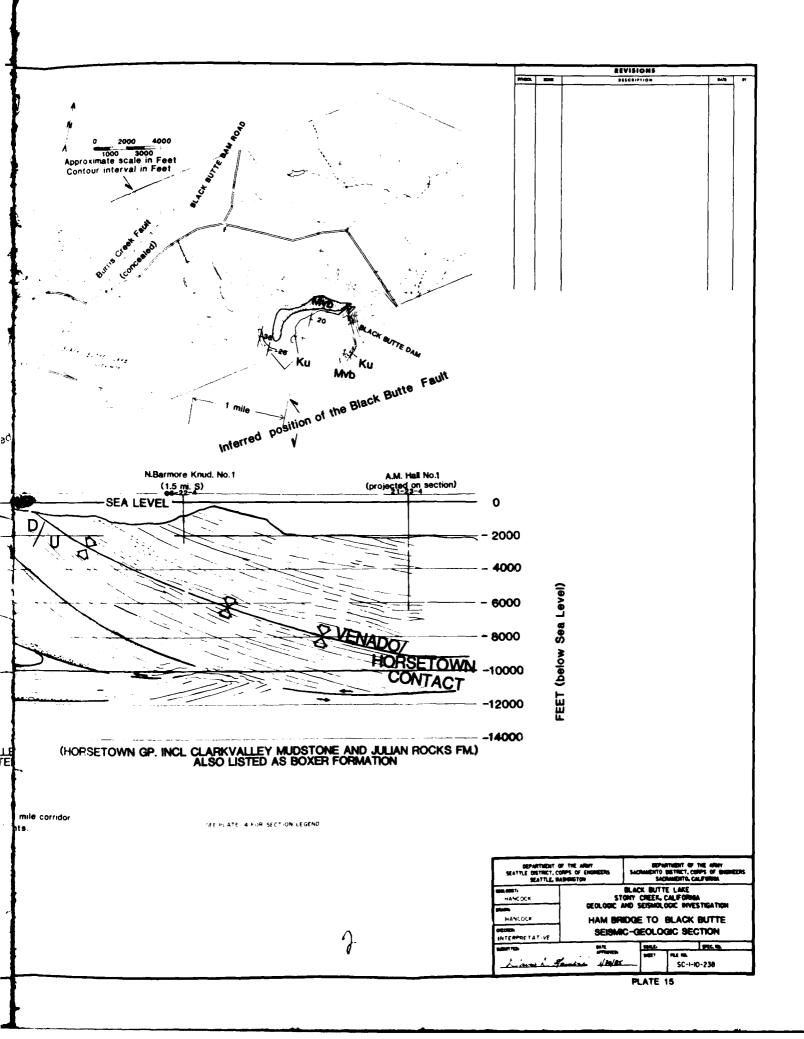
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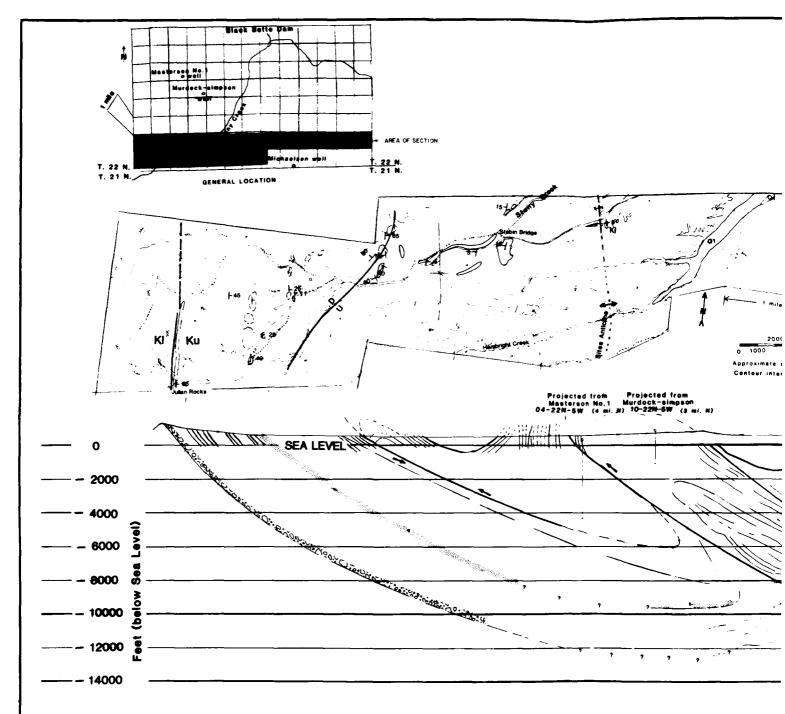






Seismic section constructed from geology, well and reflection data located along several mile corridor either side of the center of map shown above. Exact location of reflection data not shown due to proprietary nature and lending agreements. Horizontal to vertical ratio approx. 1:1



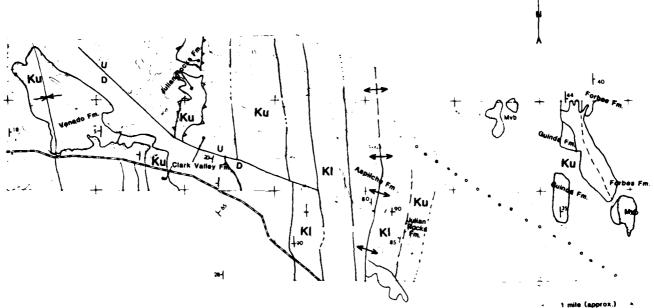


Selamic section constructed from geology, well and reflection data located along a several mile corridor near the range line dividing T, 22 M, and T. 21 M. Exact location of reflection data not shown due to proprietary nature and lending agreements.

Horizontal to vertical ratio approx. 1:1

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e <del>s</del> f	2000 0 1000 3000 4000		For Map Leger	d See Map Sheet 1	P (m)		
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0 2000 4000
1000 3000
Approximate scale in feet varies thru record length.
Contour interval in Feet.



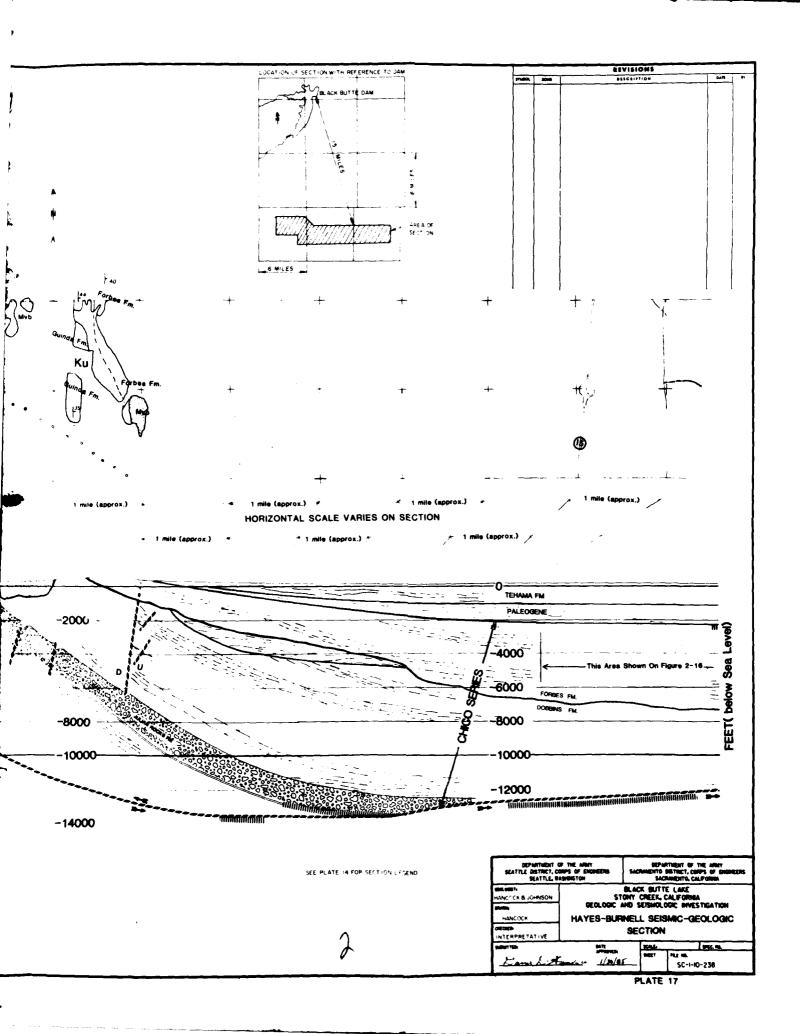
For Map Legend See Map Sheet 1.

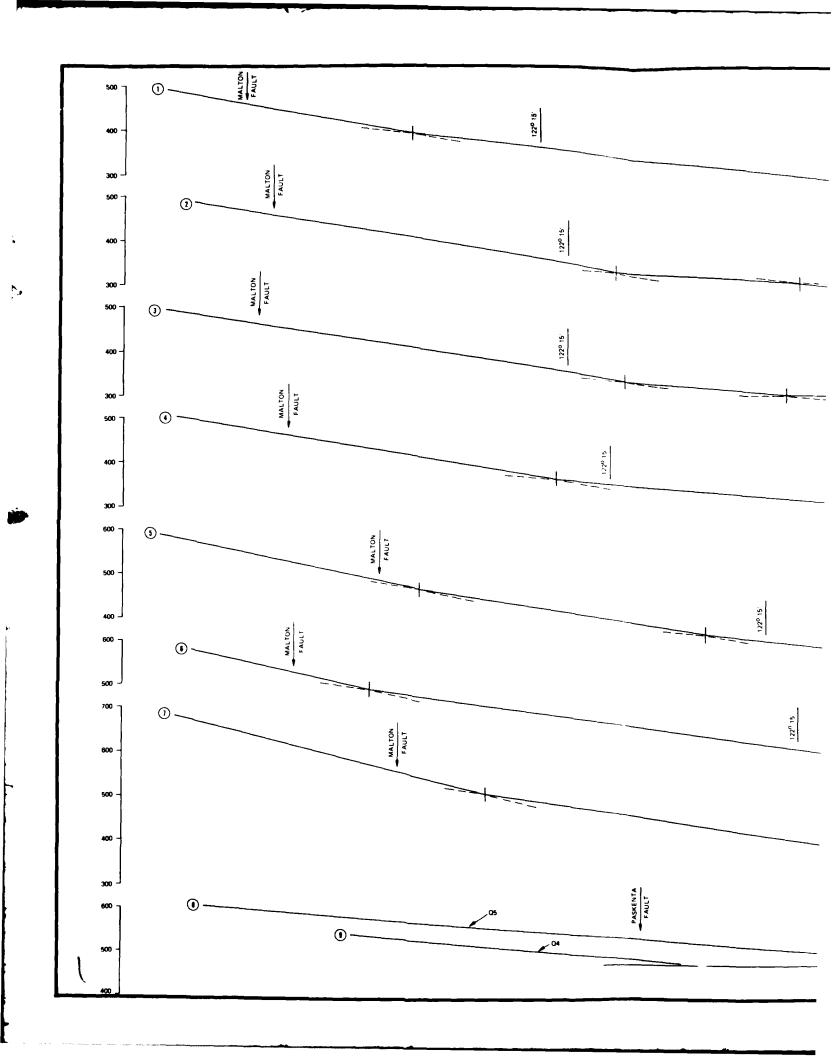
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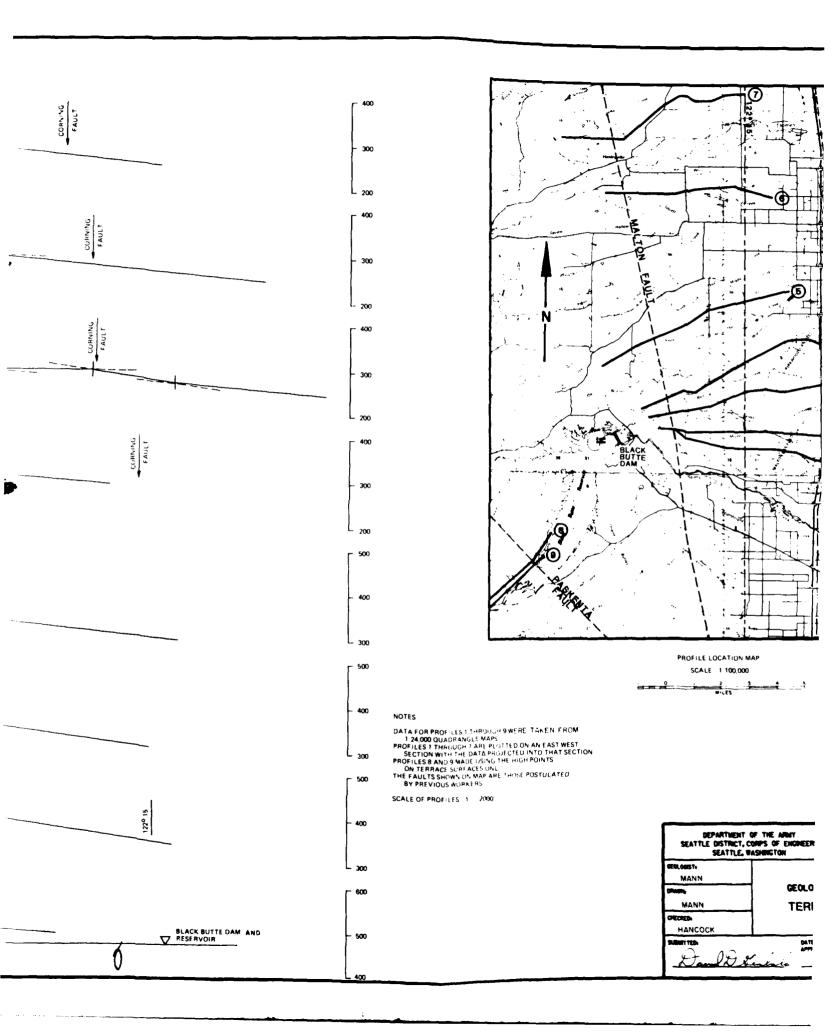
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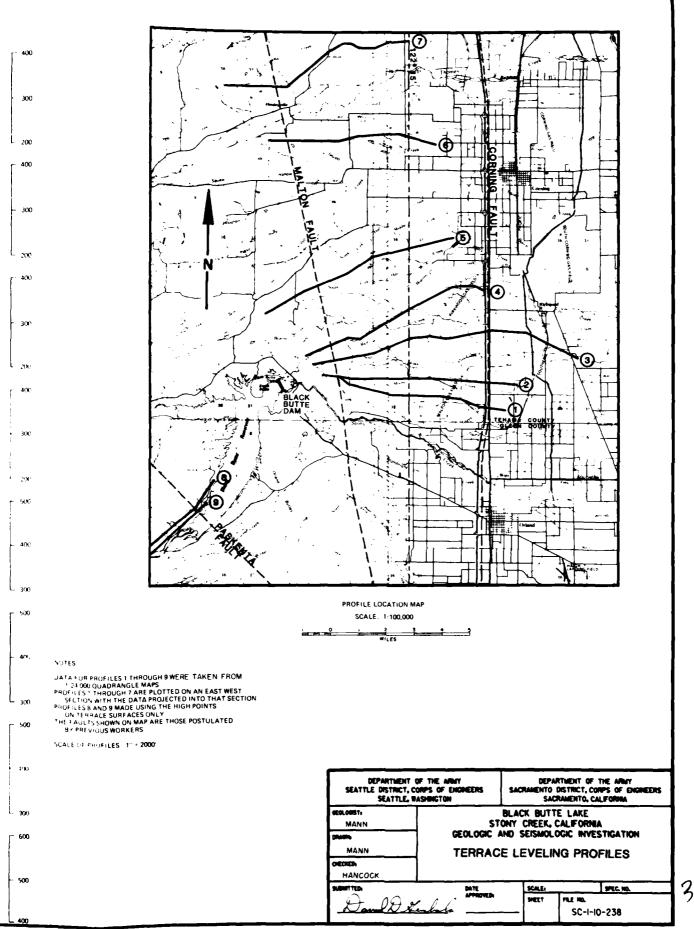
Seismic section constructed from geology, well and reflection data located along several mile corridor either side of the center of map shown above.

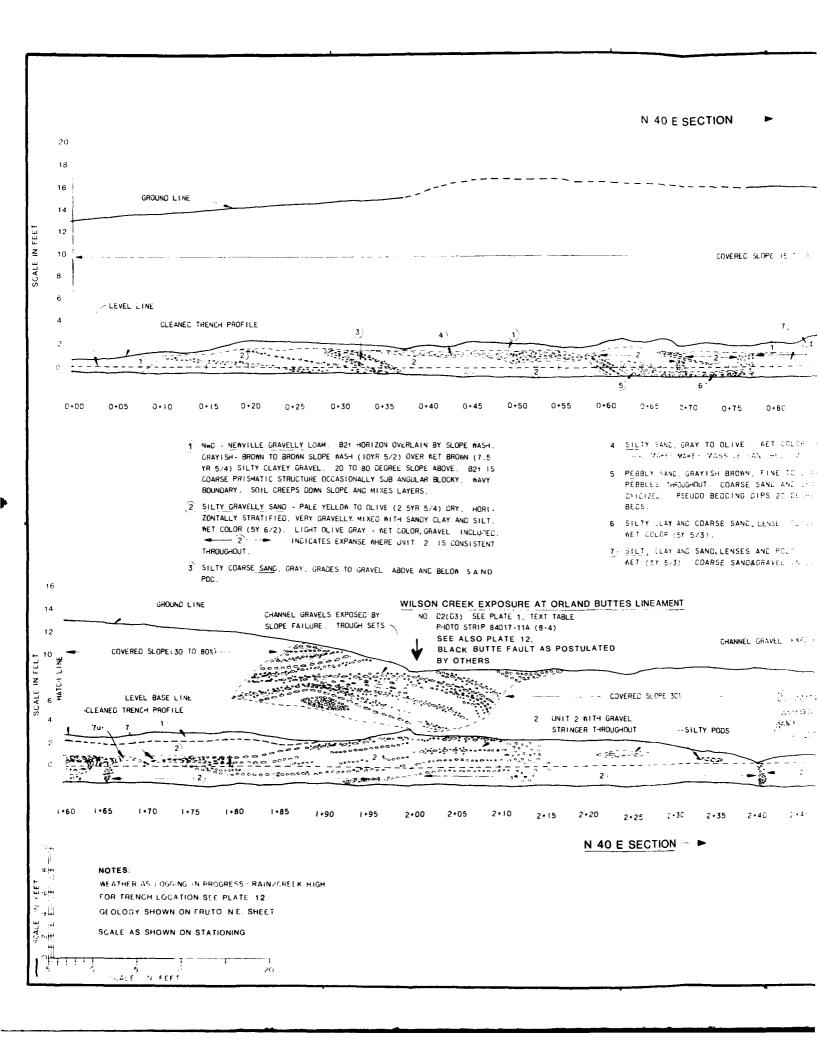
Exact location of reflection data not shown due to proprietary nature and lending agreements. Horizontal to vertical scale approx. 1:1

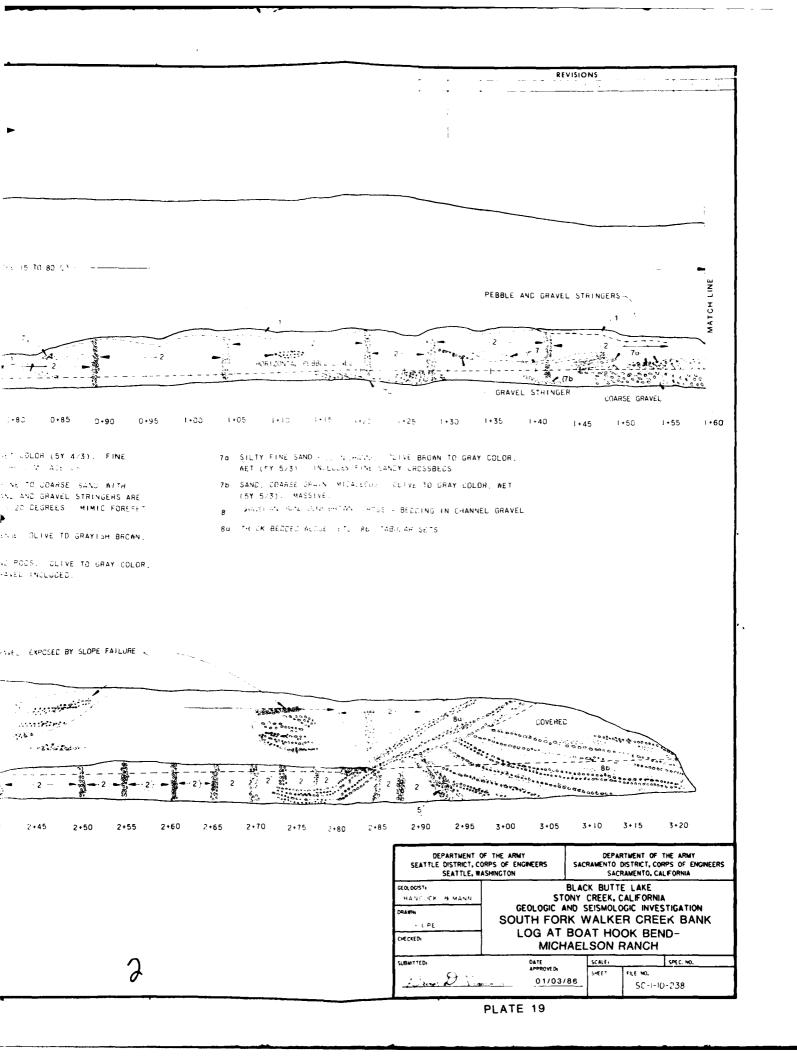












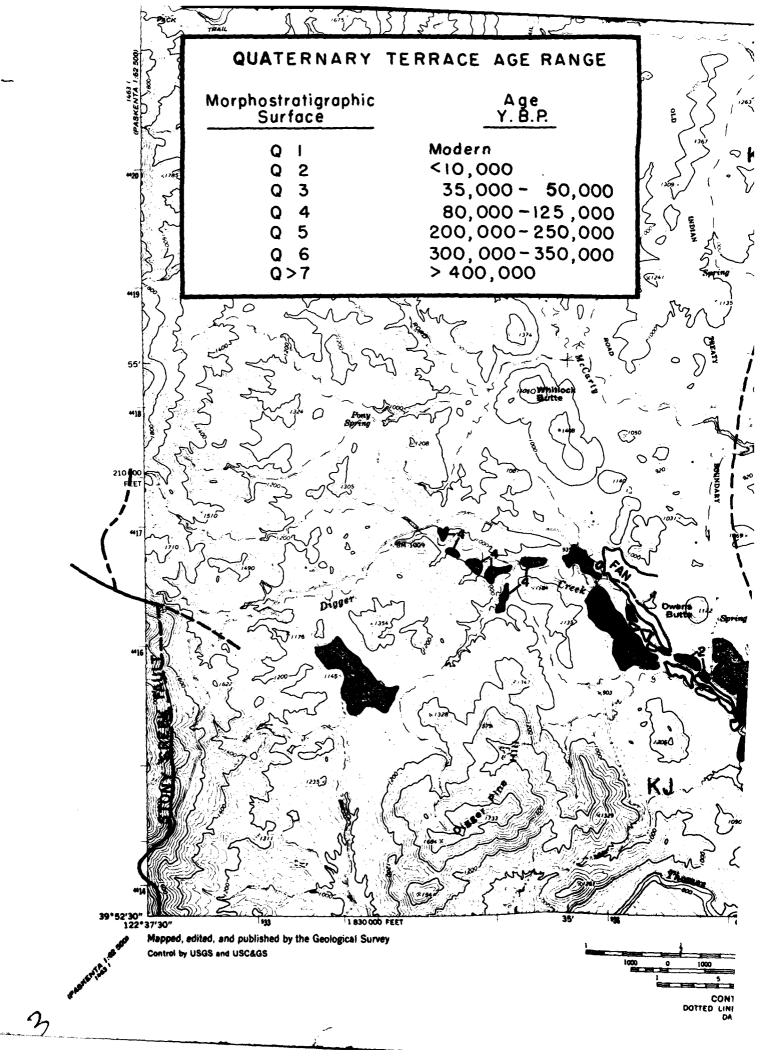
# LEGEND

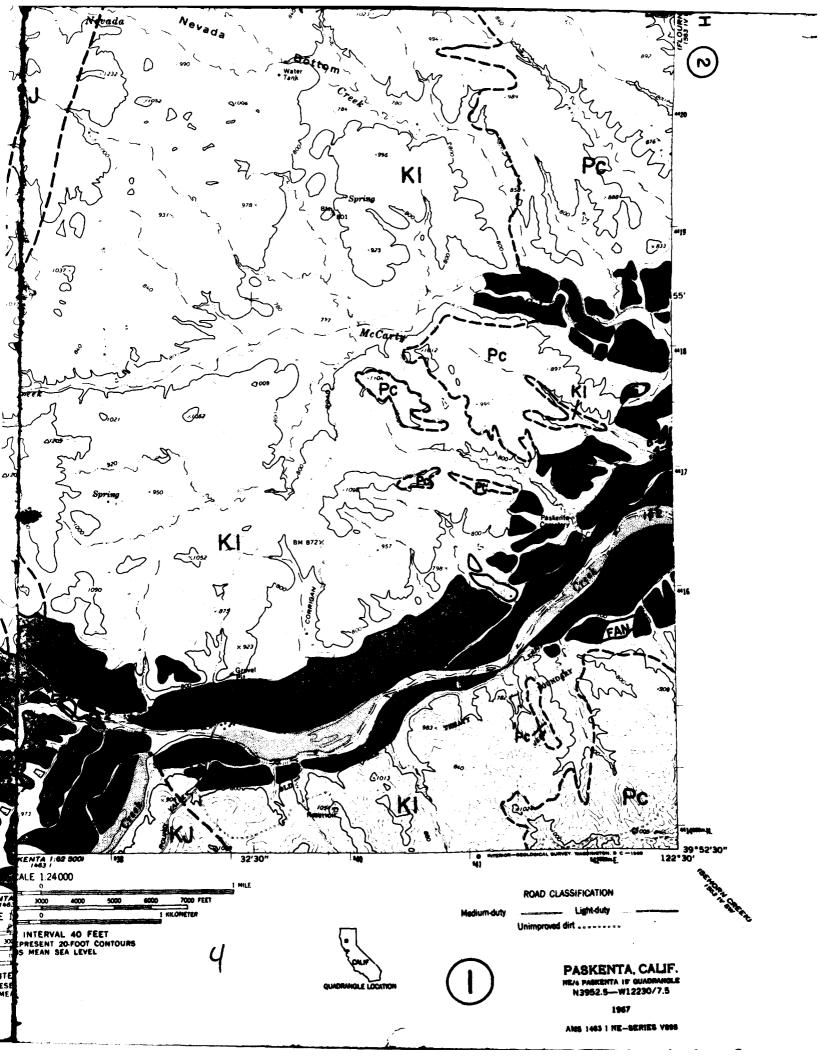
Qc——Colluvium; shown on the slopes of Orland Buttes only.	
Q———Quaternary Terrace Sequence; chronologically numbered with I being the youngest. 'Q' omitted for clarity.	RAGLIN RIDGE
Qu——Undifferentiated Quaternary Deposits; principally multiple terrace levels with local side stream fan deposits and fan deposits in the southwest area.	1111502
Pc——Tehama Formation; Pliocene alluvial fan deposits.	
Pvp——Nomlaki Tuff; found in lower part of Tehama Fan.	RILEY
Mvb——Basalt; found capping Orland Buttes. Correlates to Lovejoy Formation on east side of Sacramento Valley.	RIDGE
Ku——Upper Cretaceous sedimentary basement; includes Kione, Forbes, Dobbins Shale, Guinda, Funks, Sites, Yolo, Venado, and Boxer (Julian Rocks) Formations. (Chico Series)	
KI——Lower Cretaceous sedimentary basement; includes Lodoga (Horsetown) and the upper part of Stony Creek Formations. (Shasta Series)	HALL RIDGE
KJ——Upper Jurassic lowest Cretaceous sedimentary basement; includes lower part of Stony Creek Formation. (Knoxville Series)	
Refer to text for detailed description of units.	ALDER SPRING
Contact; dashed where approximately located.	
Anticline Anticline	FELKN HILL
Syncline	
D Thrust Fault; teeth on upper plate.	
	7

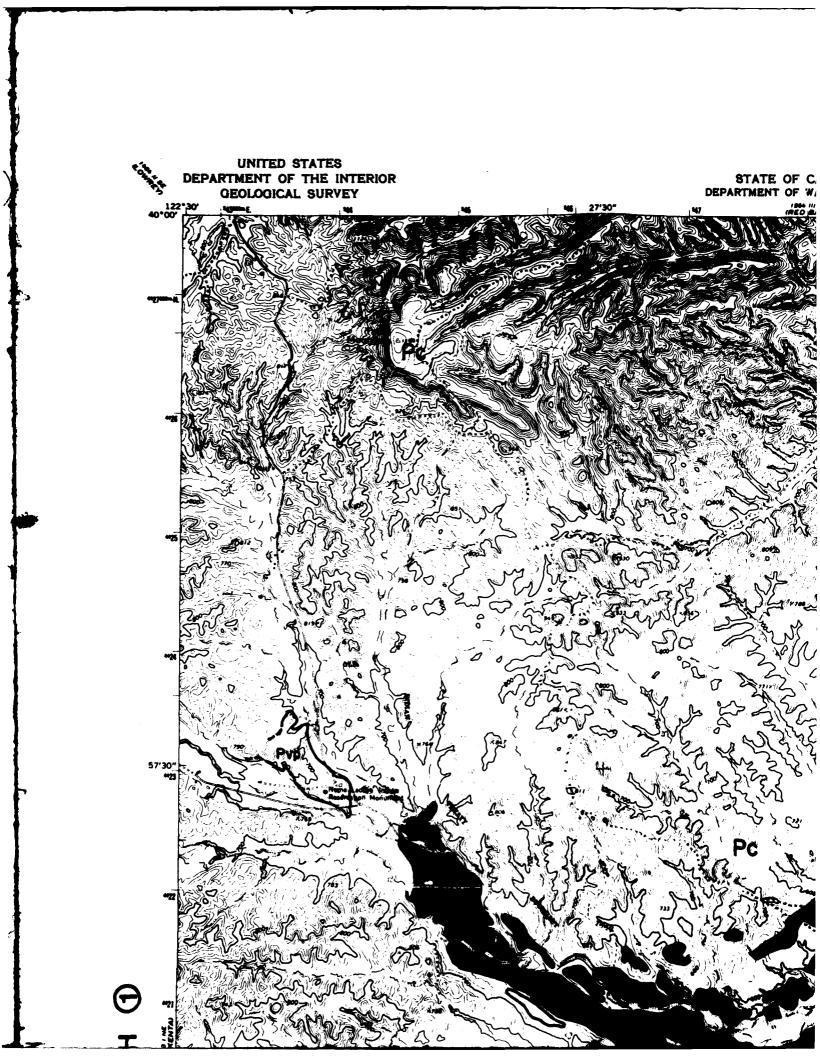
## MAP INDEX

7.5 QUADRANGLES

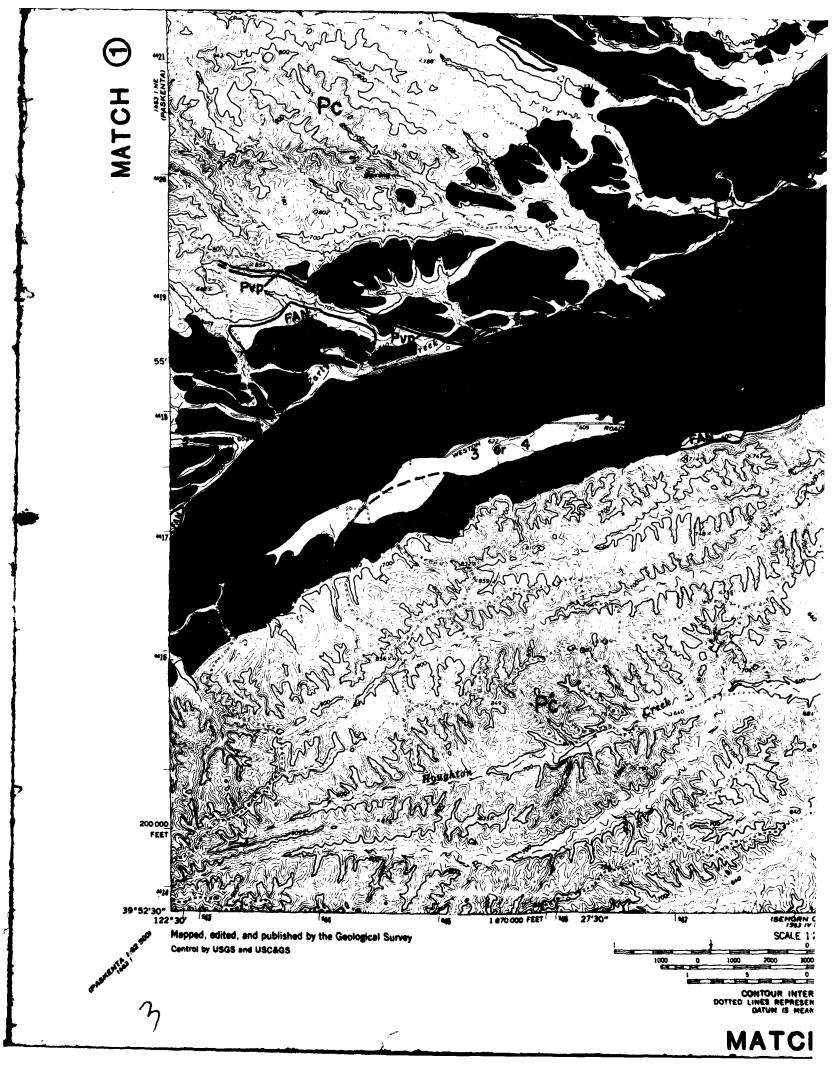
	7.5 QUADRANGLES						
only. imbered y.	RAGLIN RIDGE	LOWREY	RED BANK	WEST OF GERBER	GERBER	LOS MOLINOS	
to ley. Kione, Venado,	RILEY RIDGE	SHEET  ()  PASKENTA	SHEET ② FLOURNOY	SHEET  3  HENLEYVILLE	CORNING	VINA	
Lodoga ormations.	HALL RIDGE	NEWVILLE	SHEET    SEHORN  CREEK	SHEET  S  BLACK BUTTE DAM	SHEET  6  KIRKWOOD	FOSTER ISLAND	
· located.	ALDER SPRINGS	CHROME	SHEET  T  JULIAN ROCKS	SHEET  8  FRUTO NE	SHEET	HAMILTON CITY	
	FELKNER HILL	ELK CREEK	FRUTO	STONE Valley	WILLOWS	GLENN	



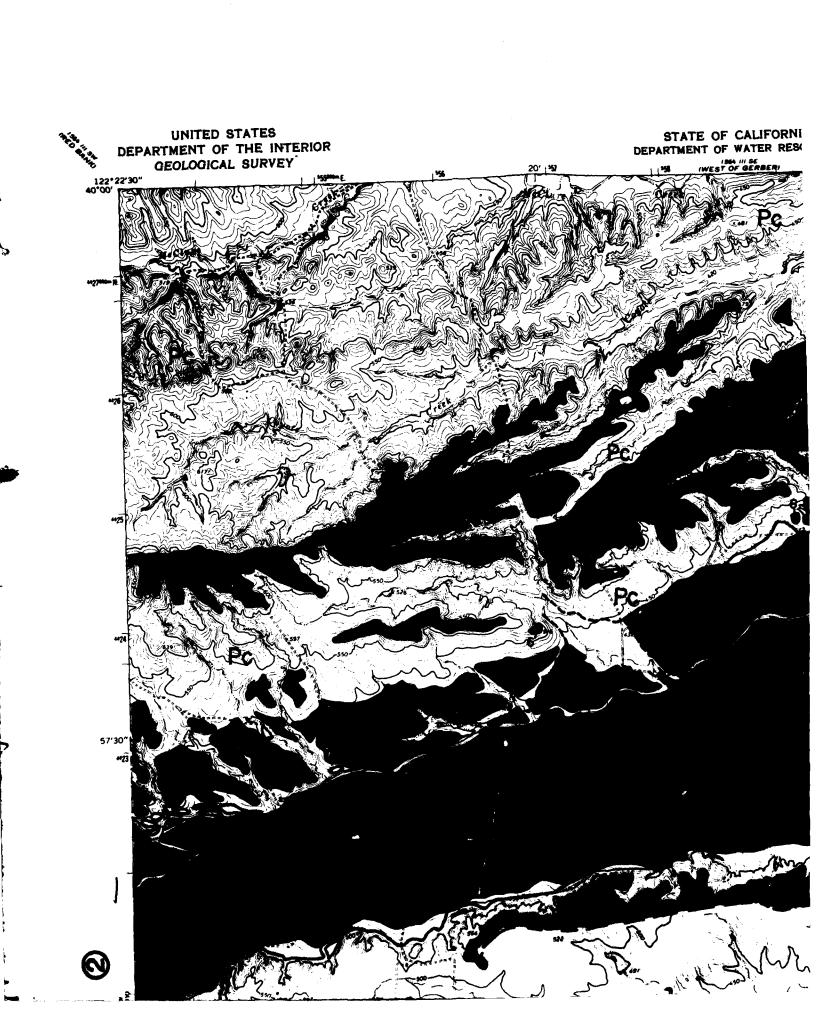


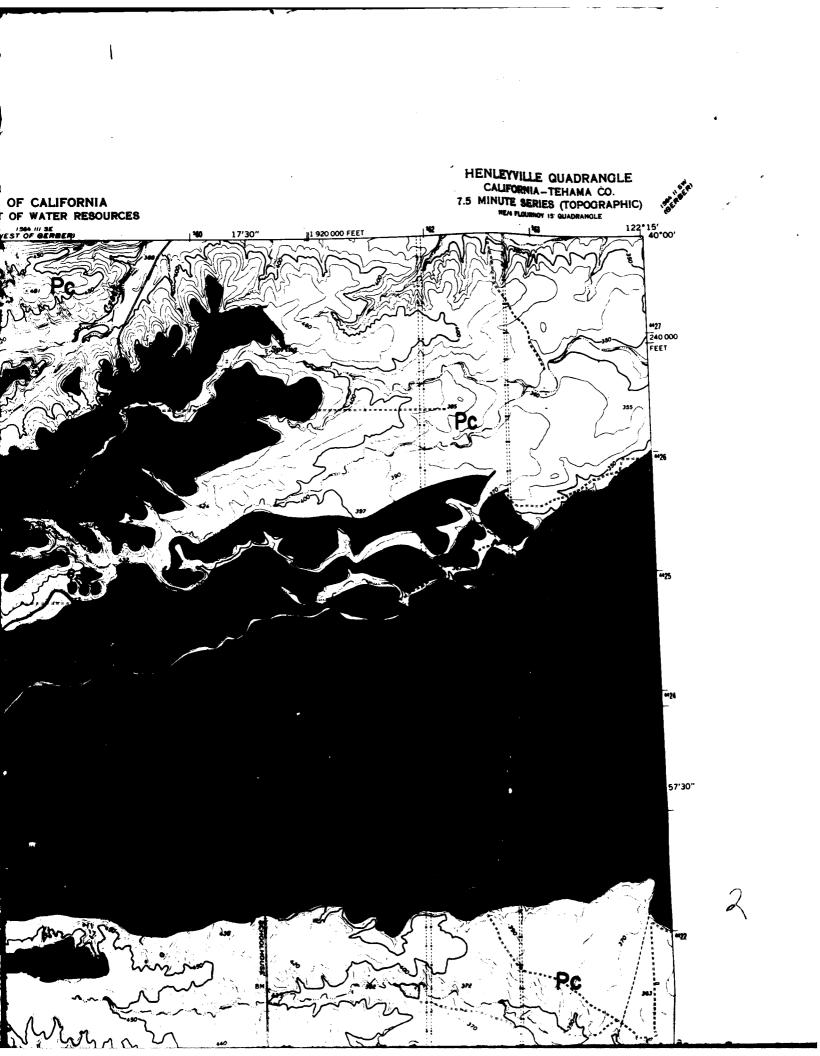


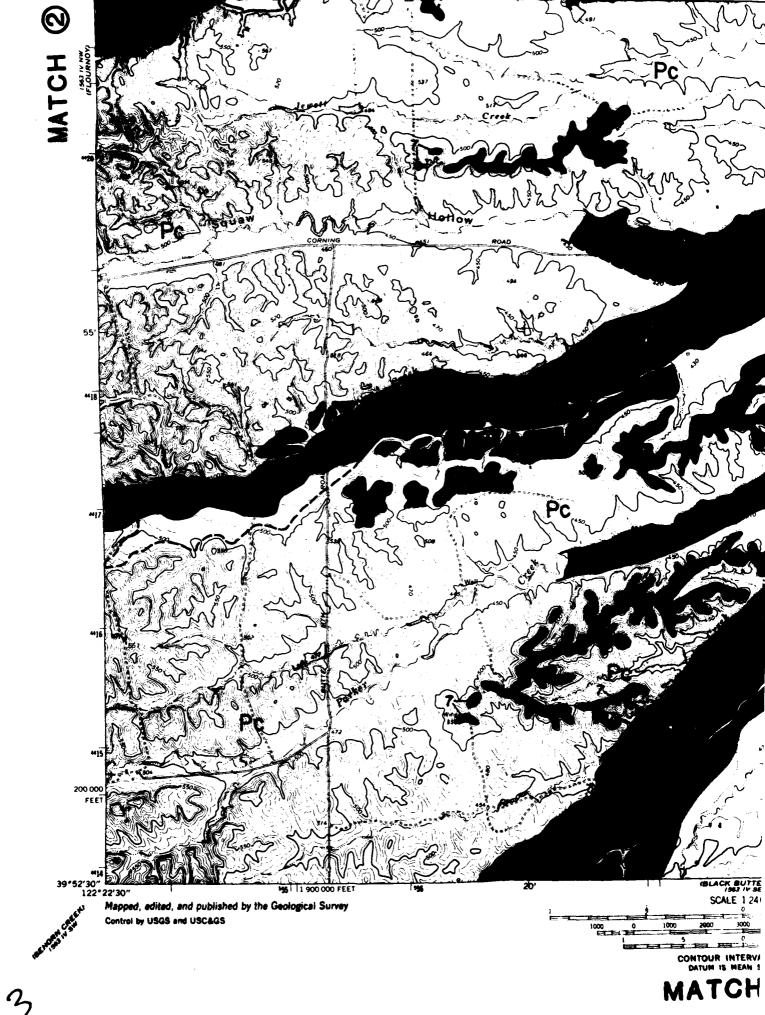
FLOURNOY QUADRANGLE CALIFORNIA - TEHAMA CO.
7.5 MINUTE SERIES (TOPOGRAPHIC)
HWA PLOUBBOY 15' GUADRANGLE OF CALIFORNIA OF WATER RESOURCES



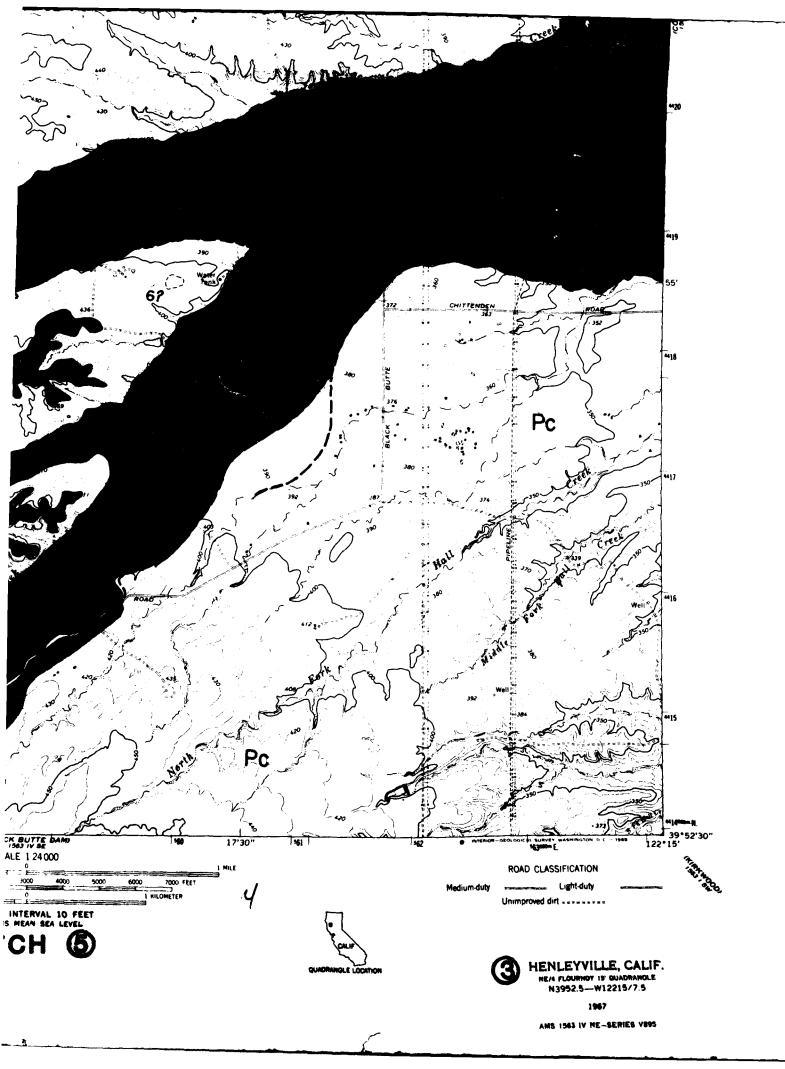




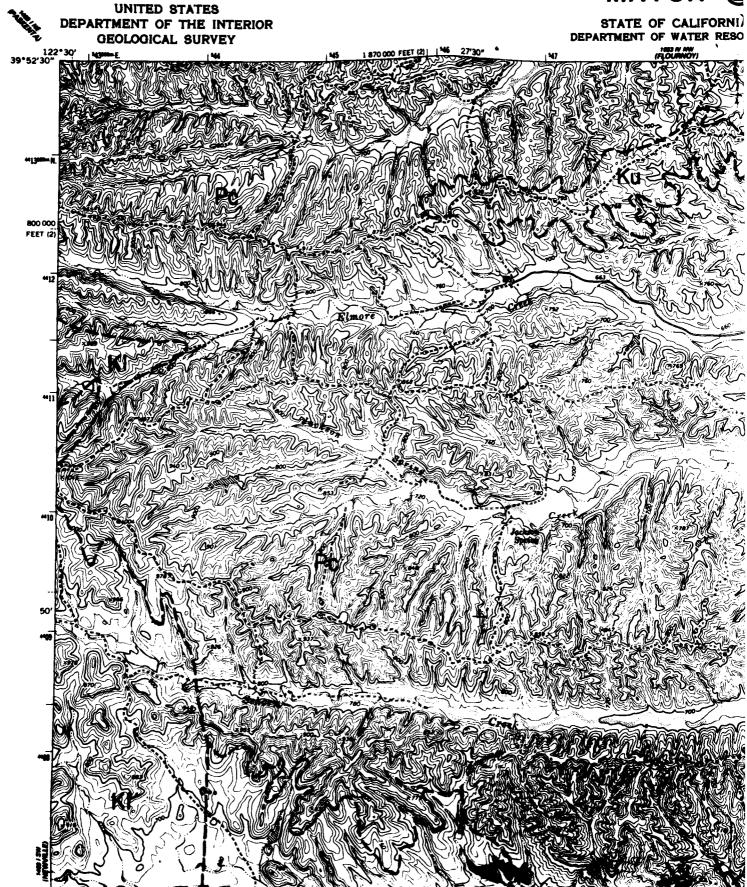




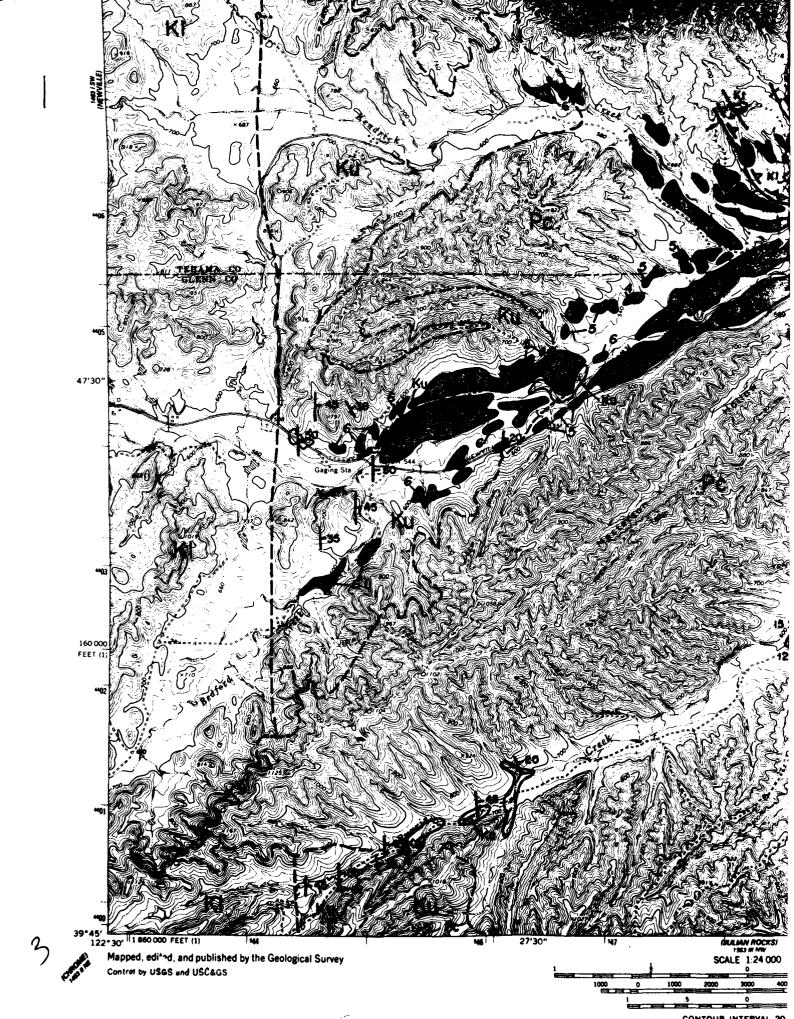
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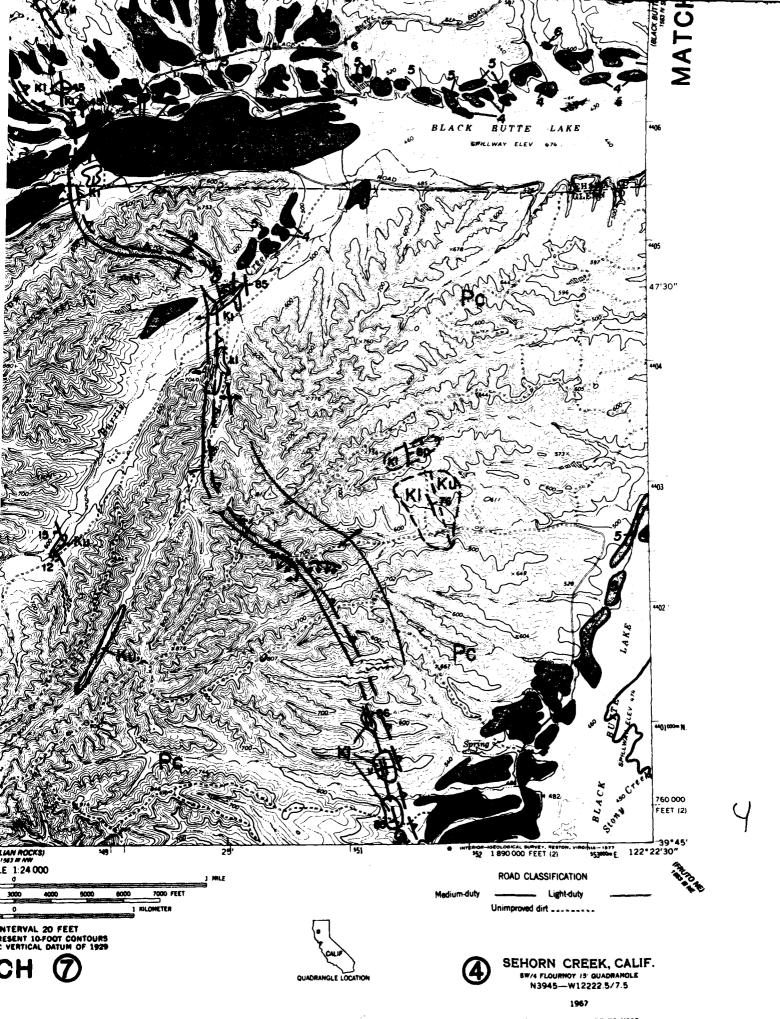
## MATCH @



'CH 2 SEHORN CREEK QUADRANGLE CALIFORNIA
7.5 MINUTE SERIES (TOPOGRAPHIC)
8W/4 FLOURNOY 15' QUADRANQLE CALIFORNIA WATER RESOURCES 122\*22'30" 190 000

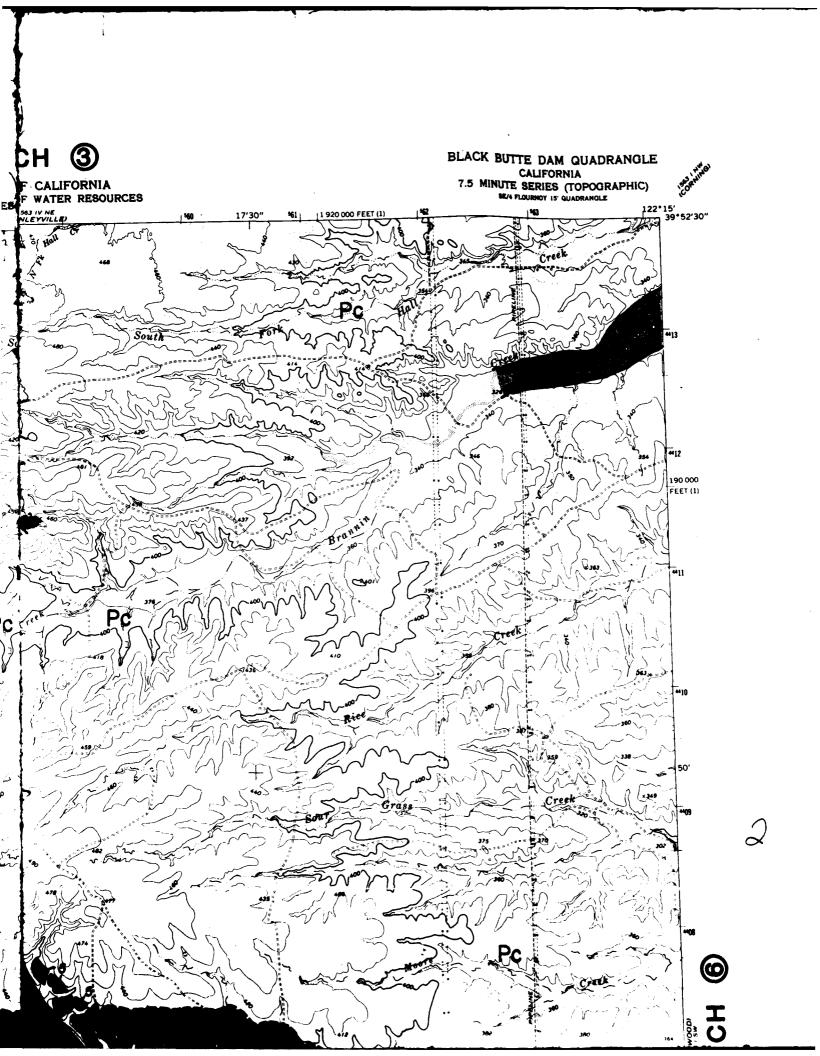


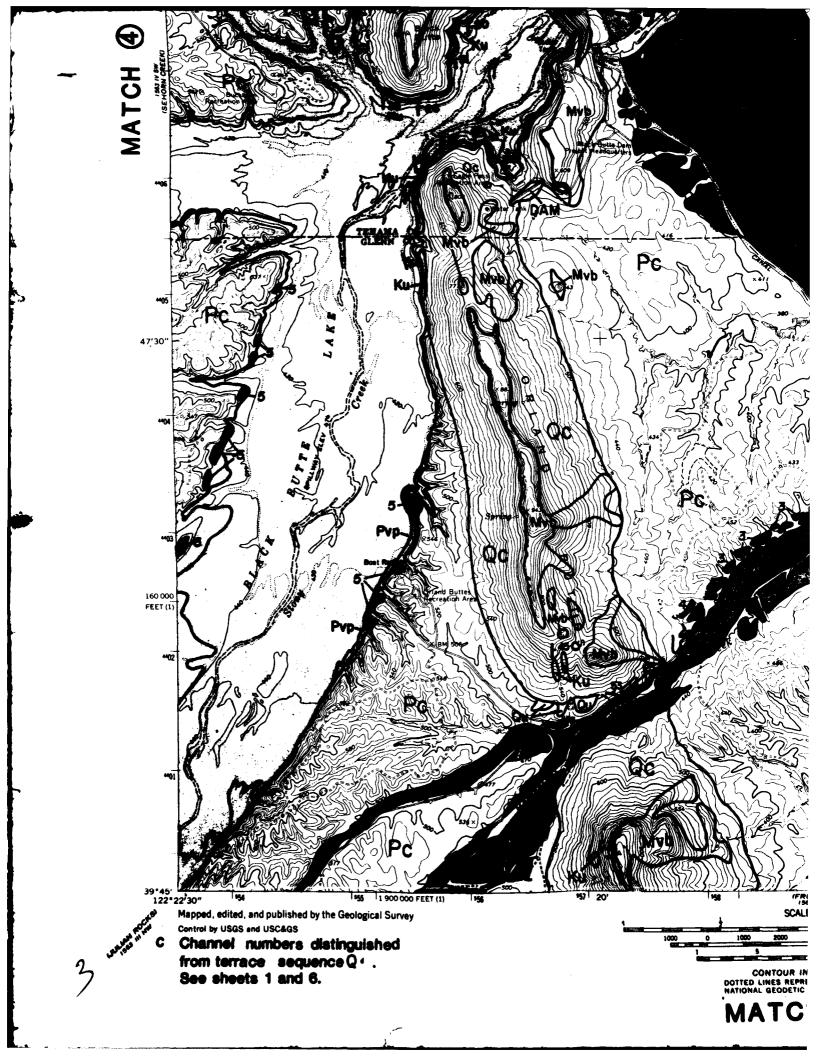
CONTOUR INTERVAL 20
DOTTED LINES REPRESENT 10-FO

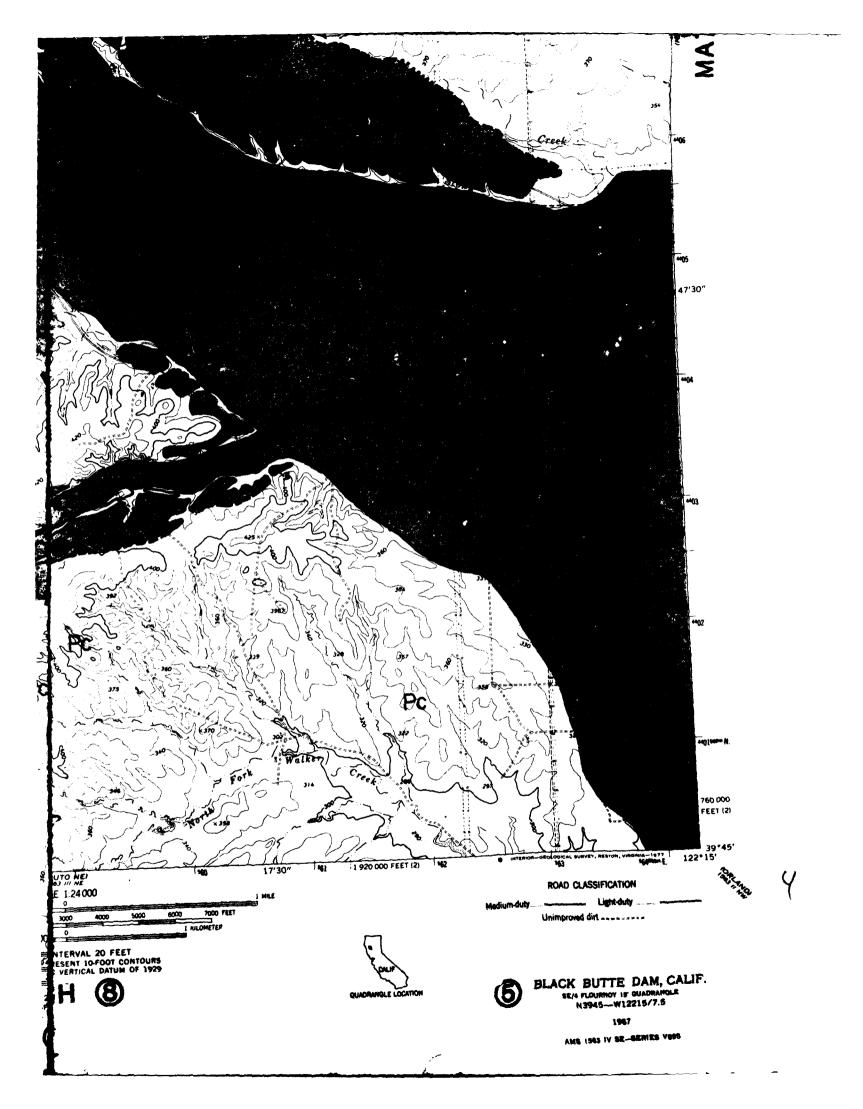


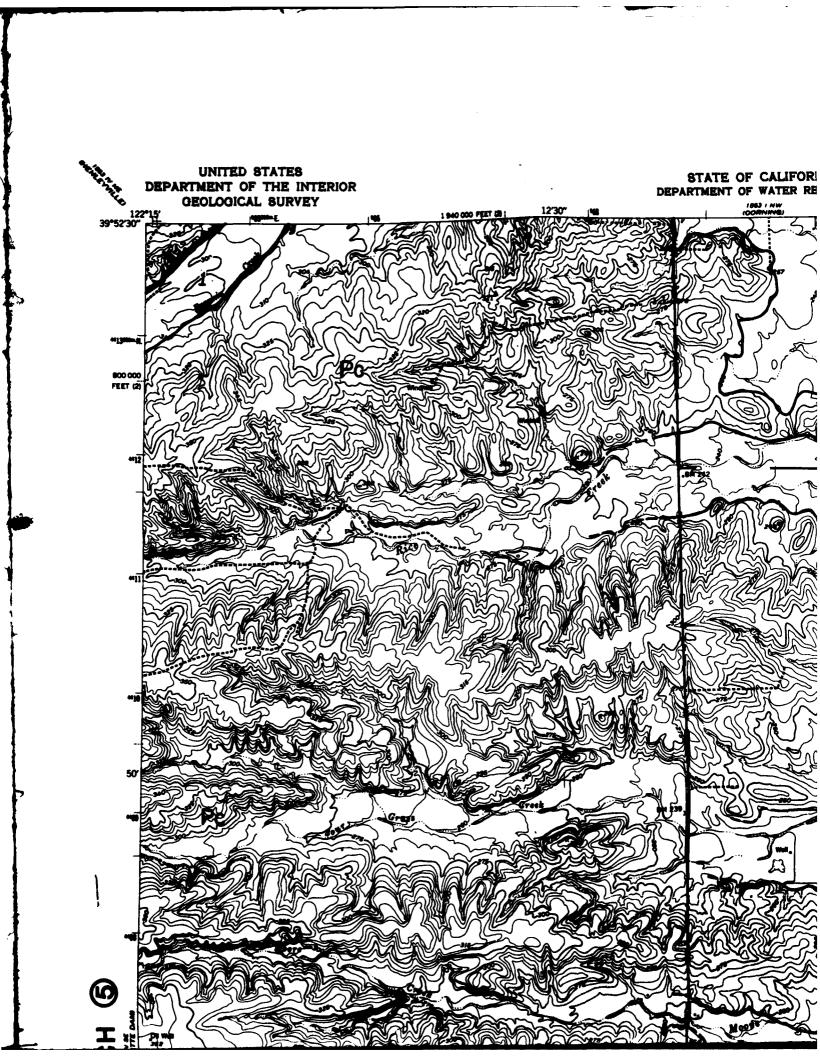
AMS 1563 IV SW-SERIES V895

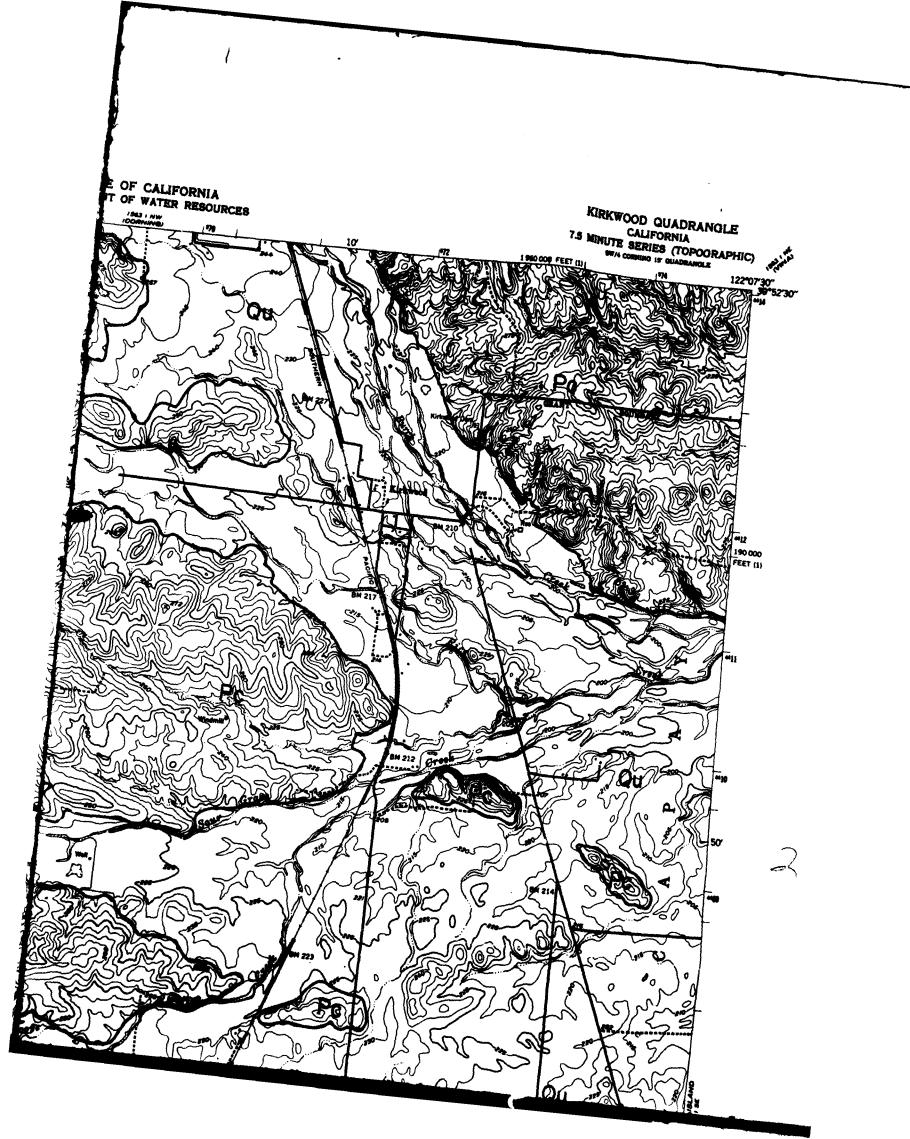
MATCH 3 UNITED STATES STATE OF CALIFORNI.
DEPARTMENT OF WATER RESC **DEPARTMENT OF THE INTERIOR GEOLOGICAL SURVEY** 122\*22'30" 39\*52'30" | 800 000 FEET (2)

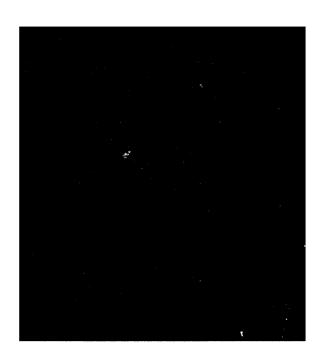


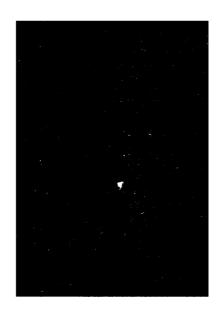




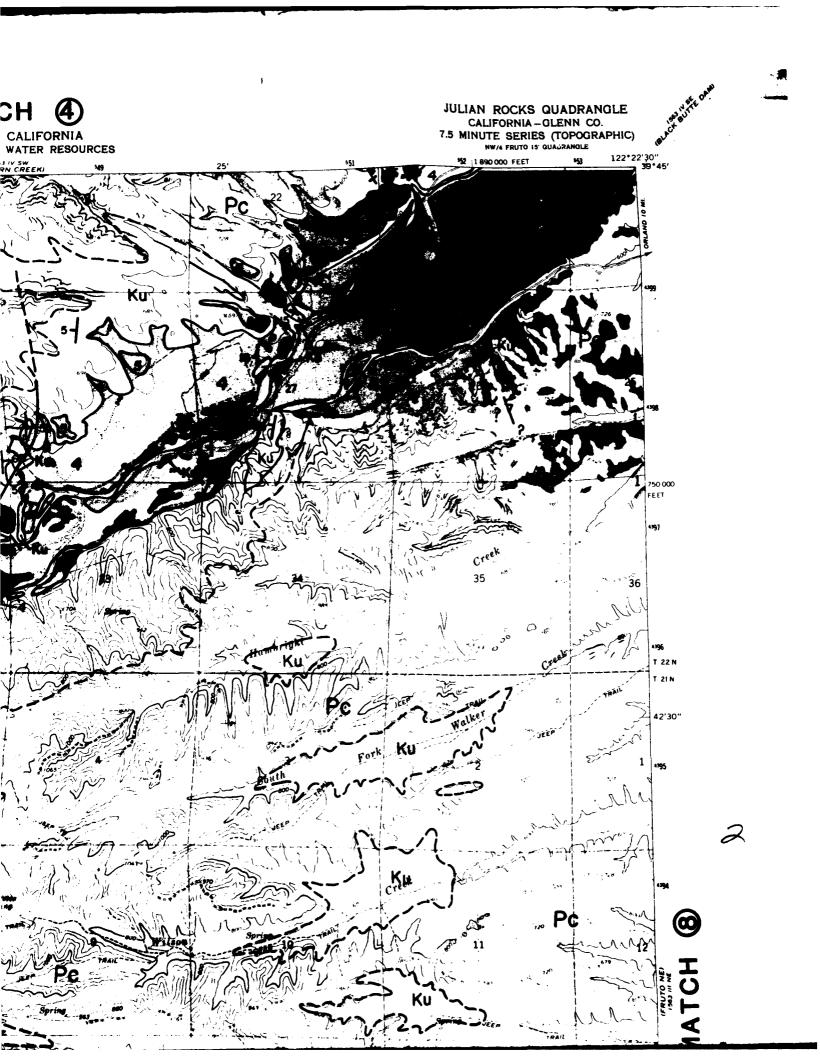


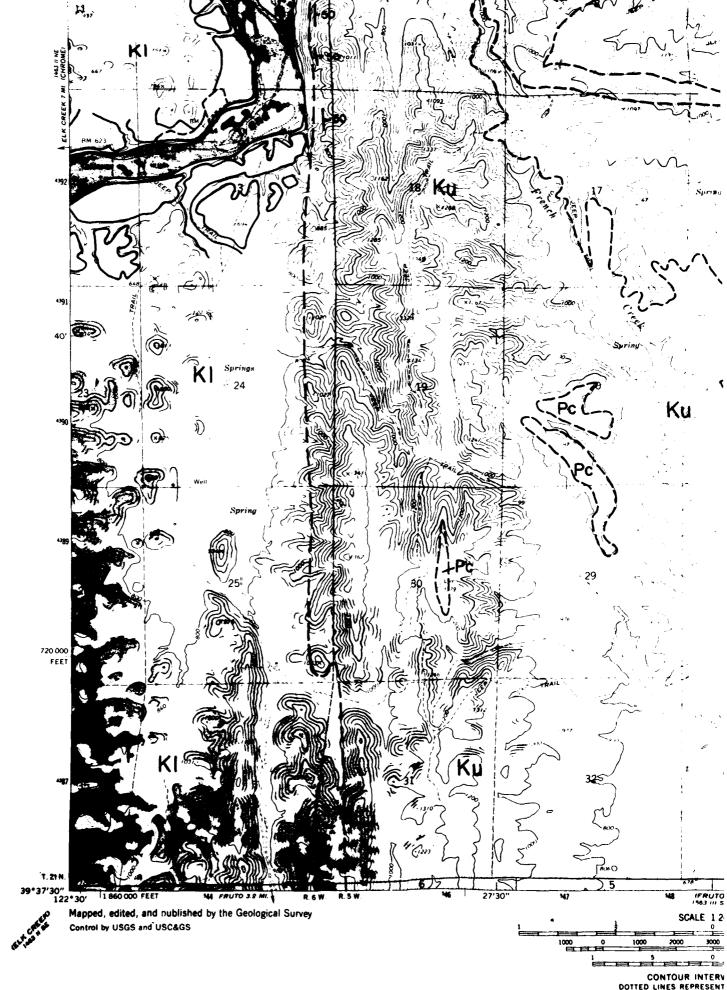




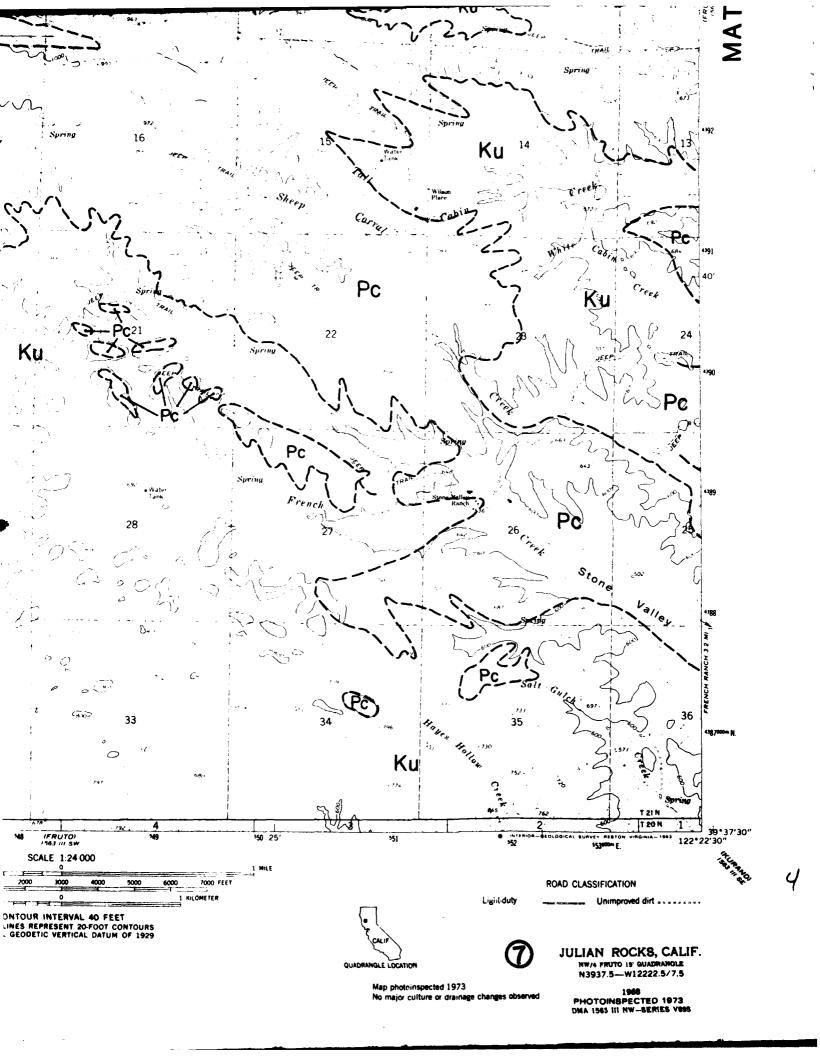


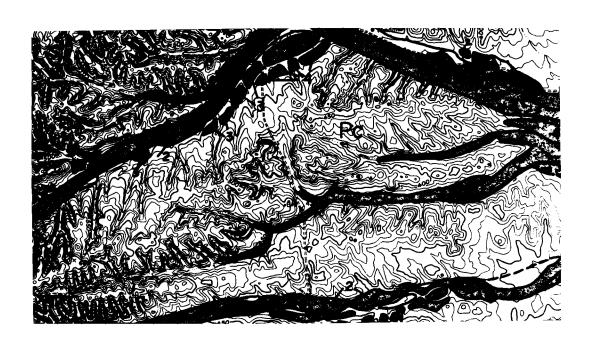
MATCH UNITED STATES STATE OF CALIFORNIA
DEPARTMENT OF WATER RESOURC DEPARTMENT OF THE INTERIOR **GEOLOGICAL SURVEY** 122°30′ 39°45′ T 22 N T 21 N 42'30'

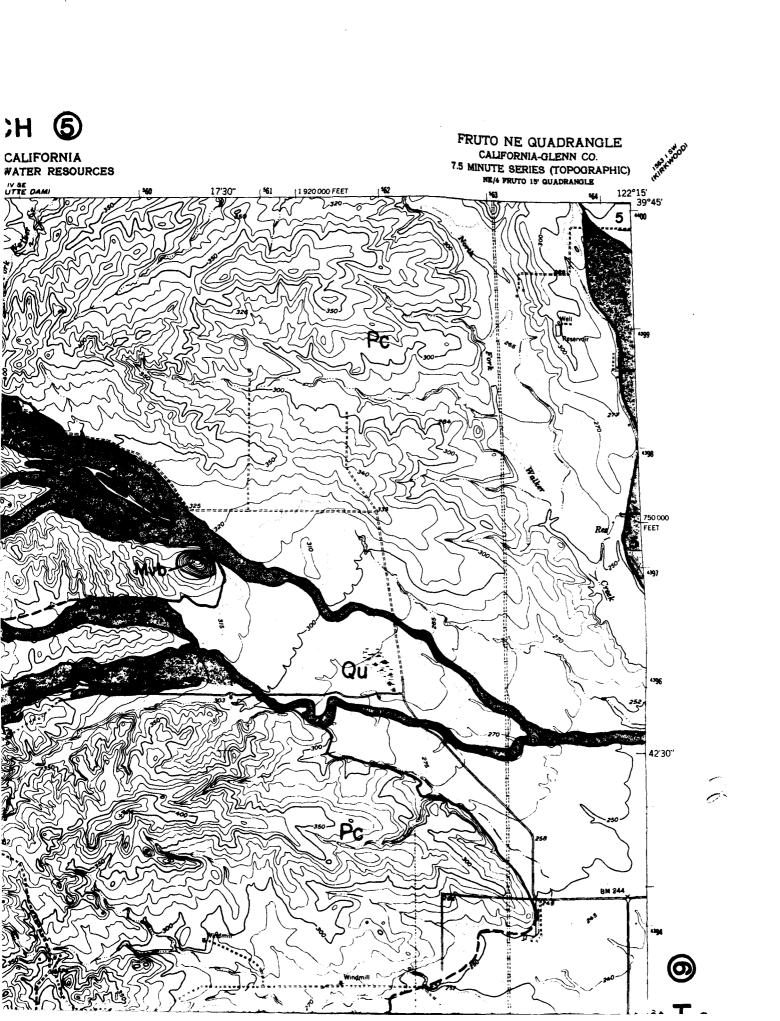


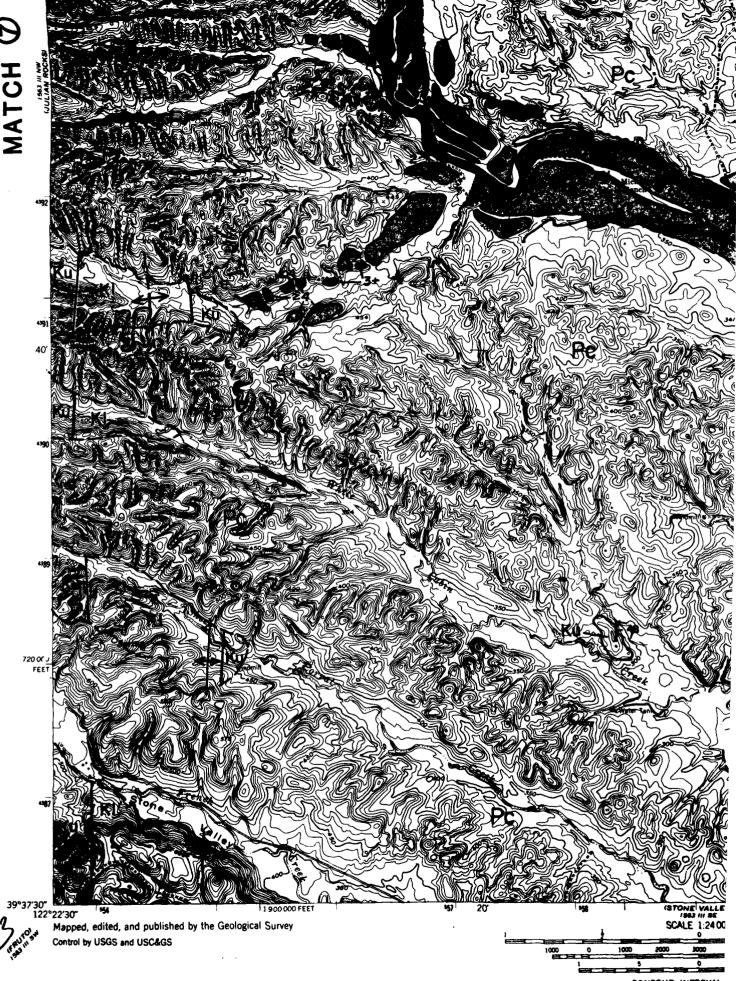


CONTOUR INTERV DOTTED LINES REPRESENT NATIONAL GEODETIC VERTI

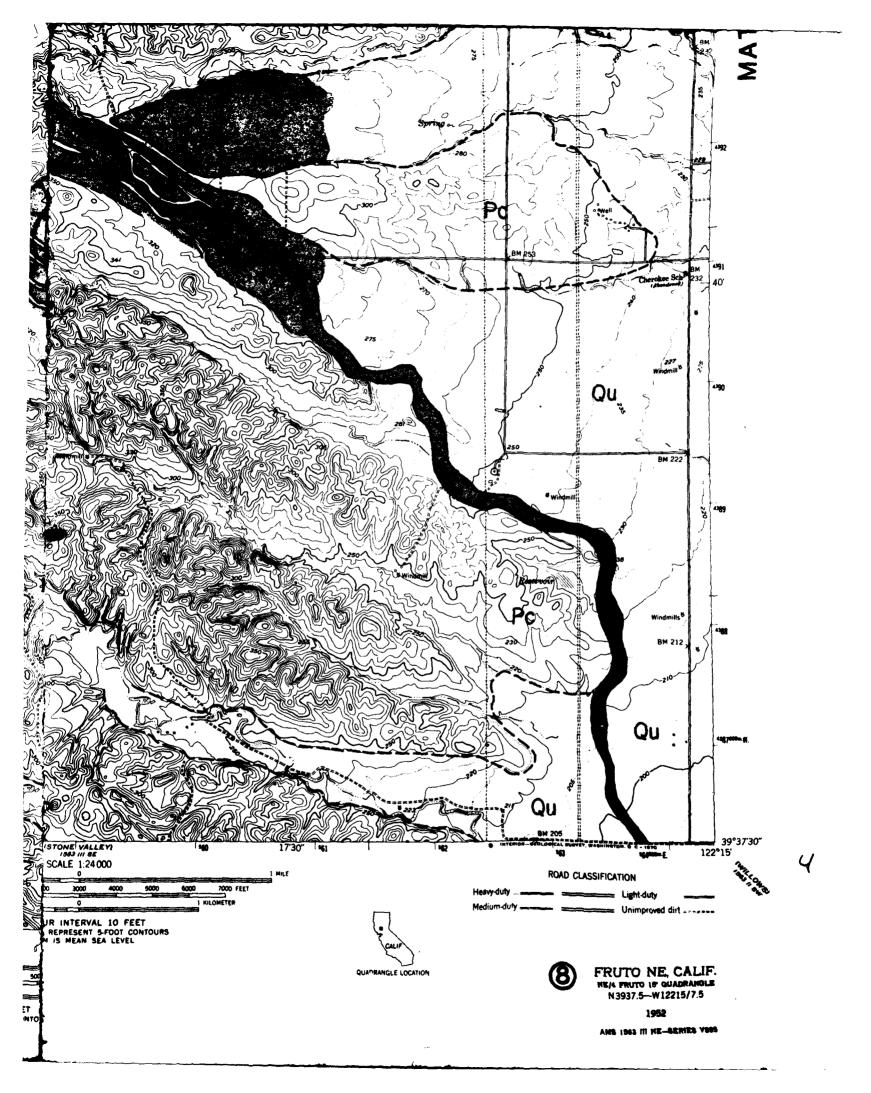


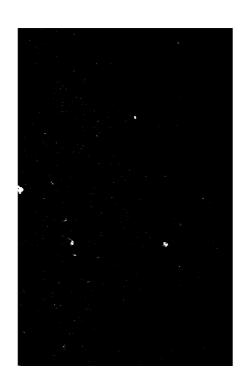


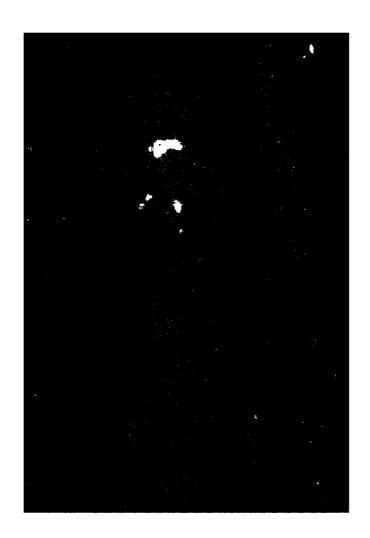


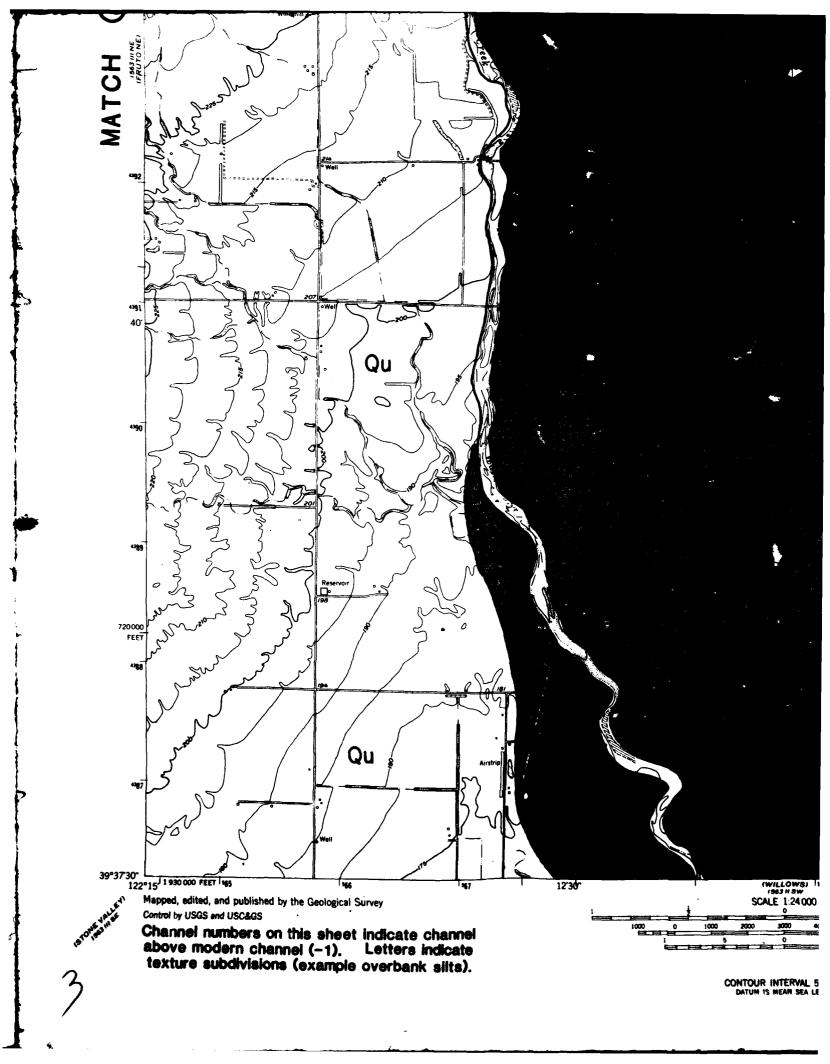


CONTOUR INTERVAL DOTTED LINES REPRESENT 5-DATUM IS MEAN SE

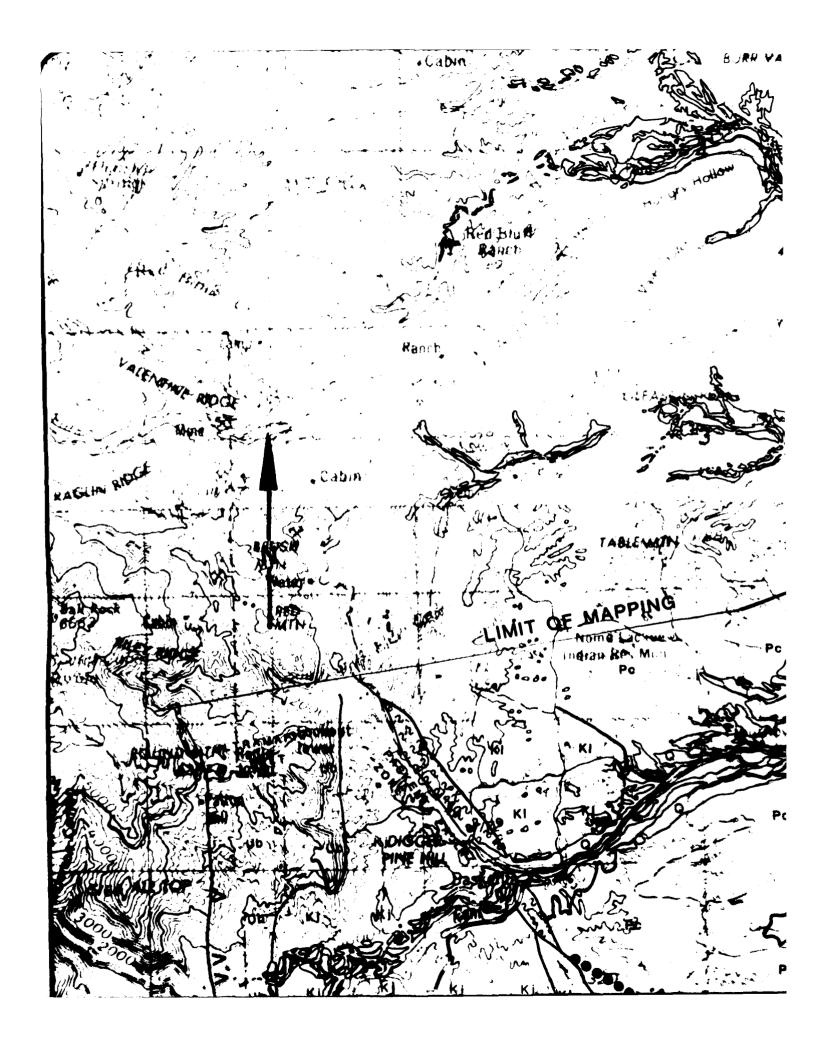


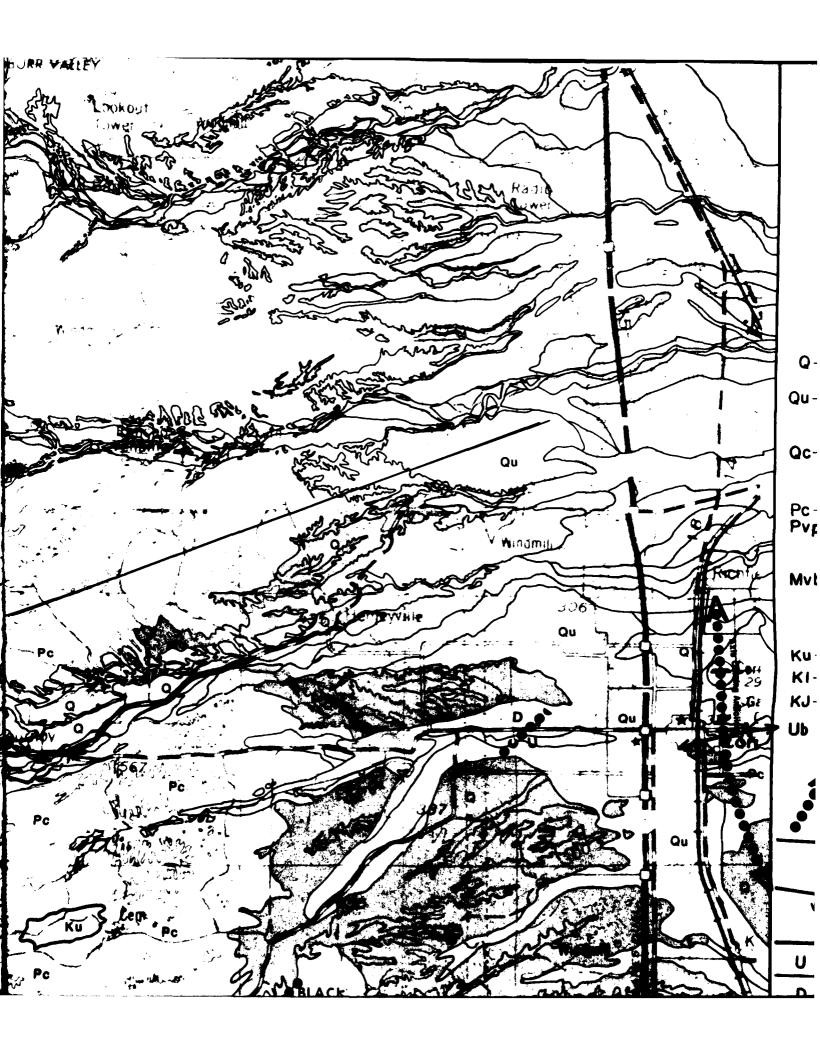


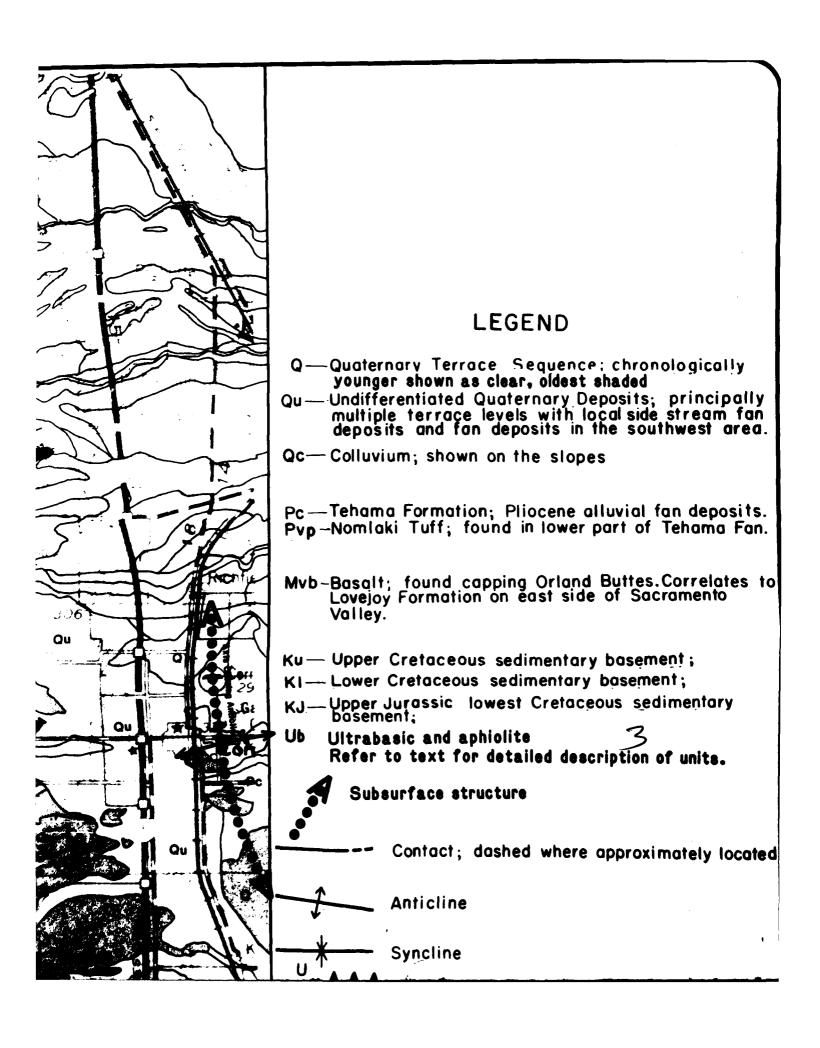


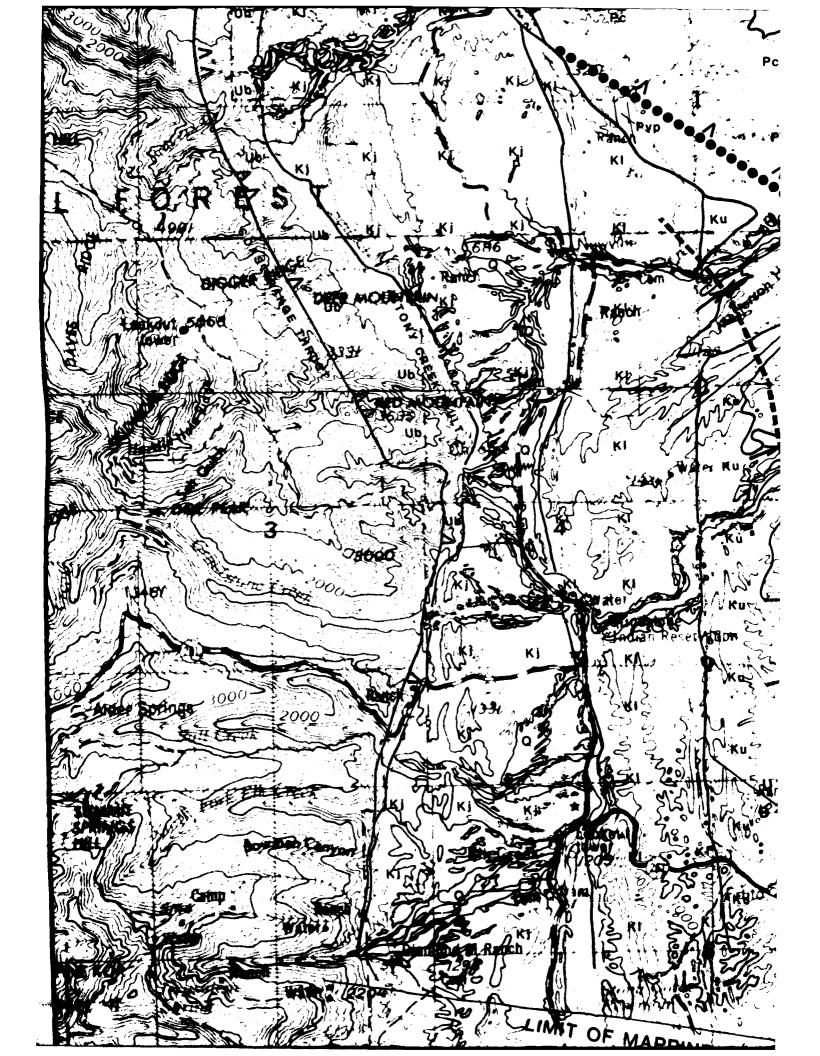


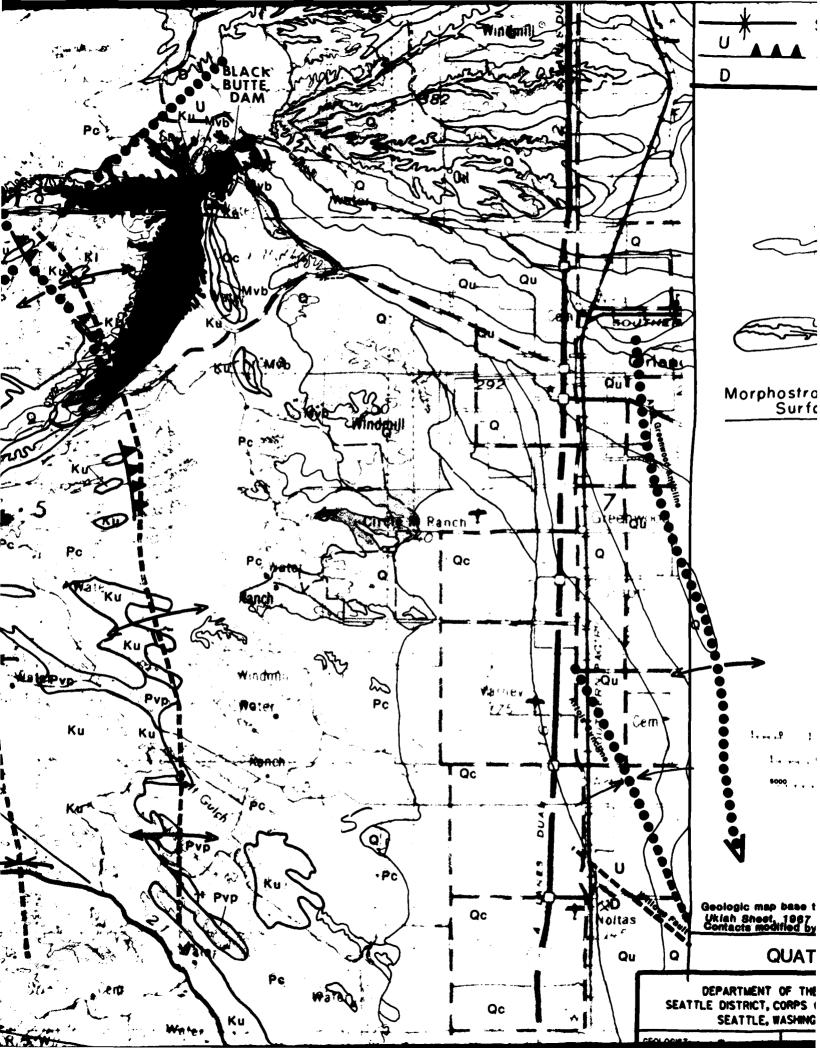


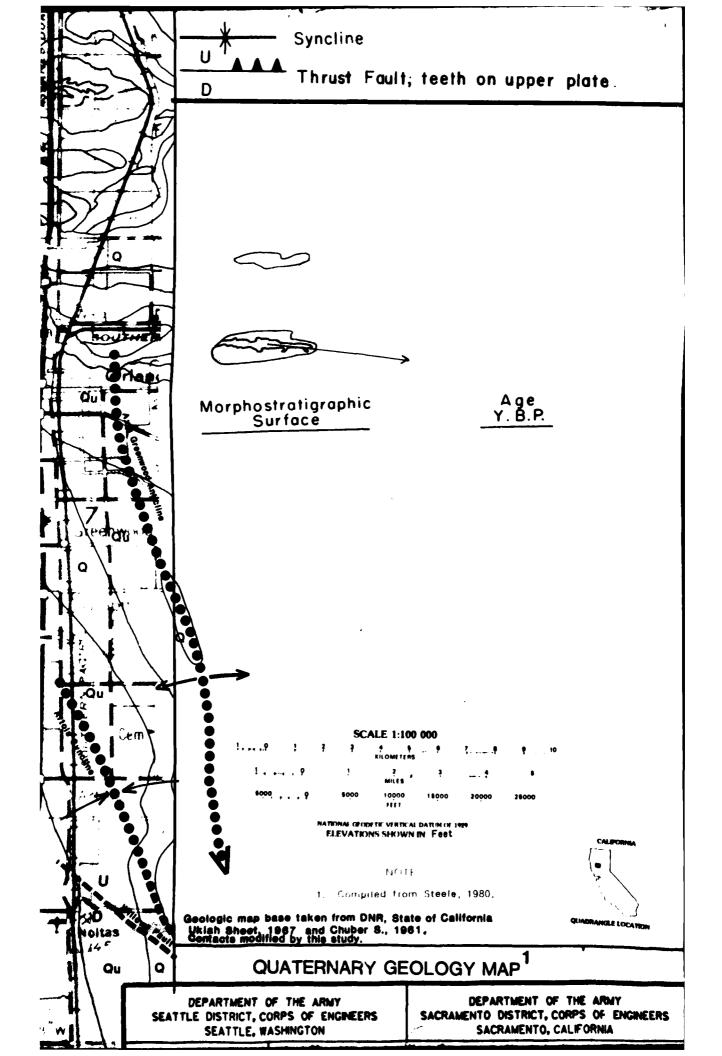


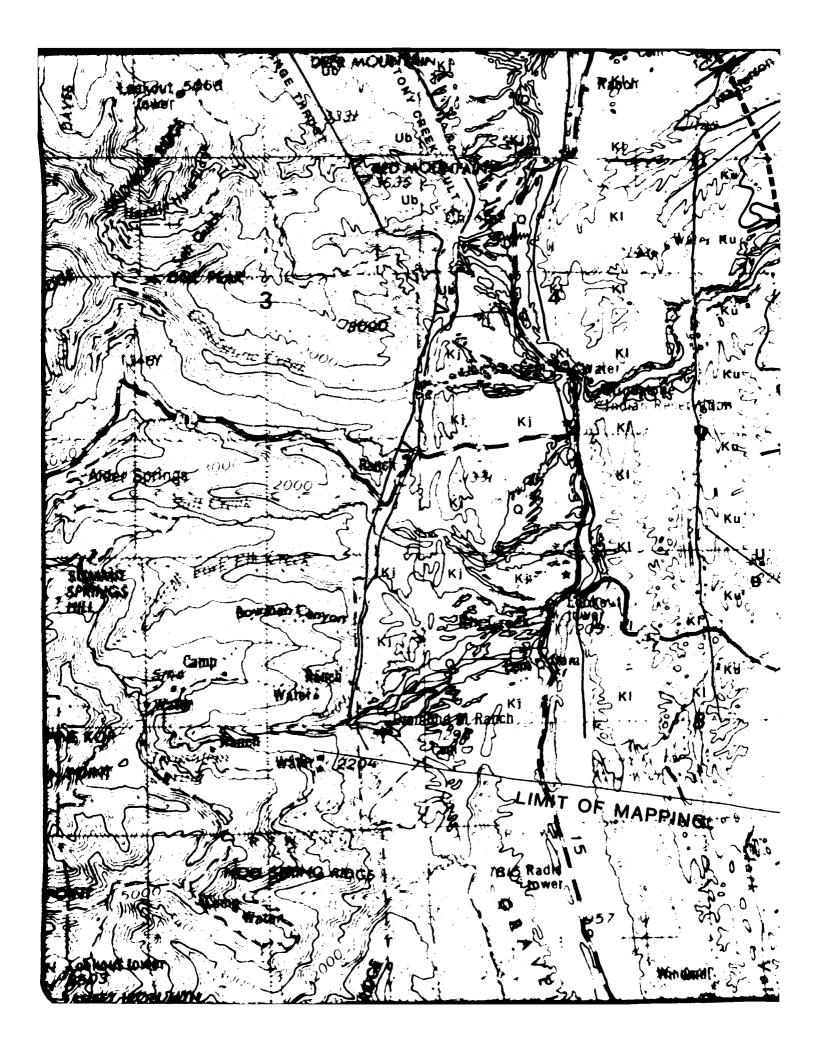


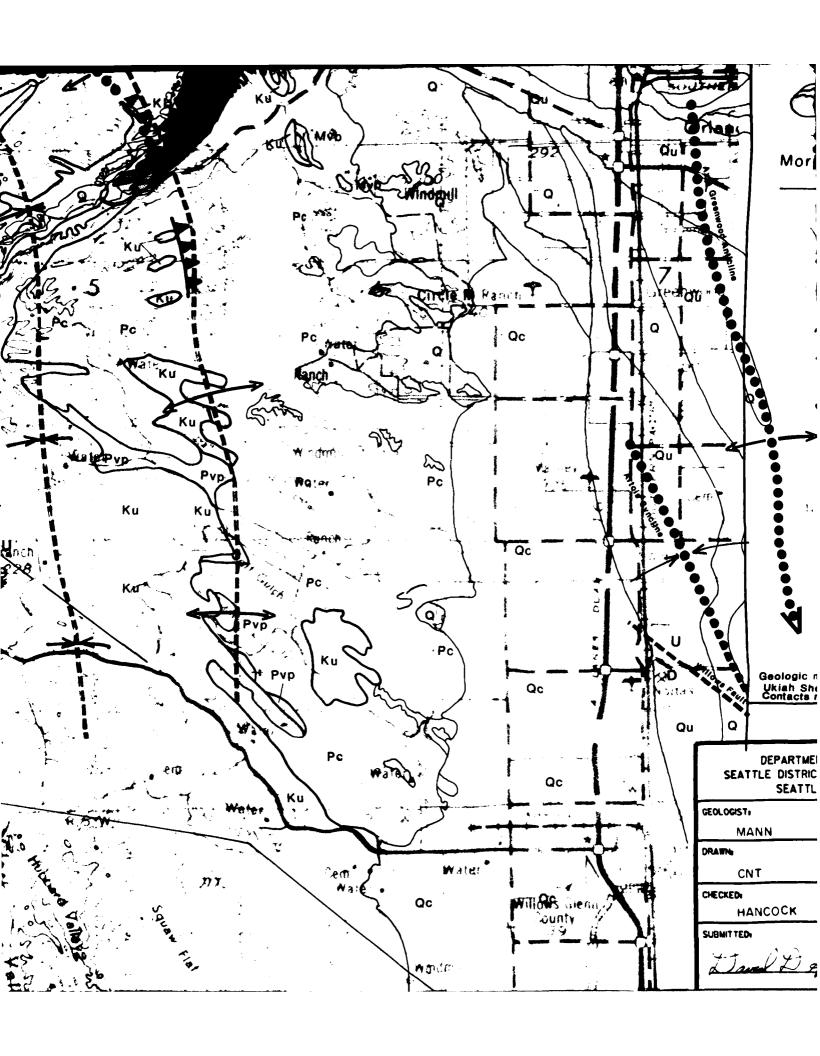












BLACK BUITE LAKE STONY CREEK CALIFORNIA GEOLOGIC AND SETSHOLOGIC HUESTIGATION(U) CORPS OF EMGINEERS SEATTLE HA SEATTLE DISTRICT HE HANCOCK ET AL. JAN 86 F/G 13/2 AD-A172 511 UNCLASSIFIED



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	BEDROCK GEOLOGY MAP
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