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THERMOSPHERE STRUCTURE VARIATIONS DURING HIGH SOLAR AND MAGNETIC ACTIVITY CONDITIONS

Jeffrey M. Forbes

Boston College Chestnut Hill, MA 02167

30 September 1985

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DONALD D.

Chief, Atmospheric Structure Branch

FOR THE COMMANDER

McCLATCHEY ROBERT

Director, Atmospheric Sciences Division

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1. MAGNETIC STORM VARIATIONS IN THERMOSPHERIC DENSITIES, TEMPERATURES, AND WINDS

At the NOAA "Workshop on Satellite Drag," March 18-19, 1982, Boulder, Colorado, a paper entitled "Geomagnetic Storm Variations and Prediction of Low-Perigee Satellite Ephemerides" by J.M. Forbes was presented. A written version has appeared in the proceedings of the workshop. In this spatial/temporal characteristics of density paper, the perturbations in the lower thermosphere (150-250Km) during geomagnetic disturbances are discussed within the context of the low-perigee satellite ephemeris prediction problem. The variability of atmospheric density (satellite drag) is significantly enhanced at high latitudes due to the proximity of magnetosphere/ionosphere source mechanisms for the neutral atmosphere disturbances. In addition, for precise orbital computations winds in excess of about 300msec⁻¹(which commonly occur at high latitudes) must also be considered as part of the drag effect. It is demostrated that precise low-perigee ephemeris predictions require accurate forecasting of the phase as well as spatial structures of atmospheric density and wind disturbances. It was recommended that future advances in predictive capability of atmospheric drag necessitate (1) Comprehensive numerical modelling efforts which tie together all of the relevant

high-latitude processes in a self-consistent manner, and which can be utilized to provide insight into the development of an improved magnetic index of density variations, and guide the development of semi-empirical models for predicting stellite drag variations; and (2) Concerted experimental efforts to obtain simultaneous drag and electrodynamic data (incoherent scatter measurements of temperature and drifts; satellite measurements of particles, fields and currents; and ground magnetometer measurements).

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Dr. Forbes and Mr. J. Slowey attended a Technical Interface Meeting at AFSCF, Sunnyvale, CA on 26-27 April 1982, which was held for the purpose of coordinating AFSCF's Data System Modernization (DSM) efforts with its prime A point of discussion concerned the contractor. IBM. Jacchia-Roberts analytic model, which was originally fitted to the J70 model, whereas a second version of the model was developed by IBM which was fitted to the J71 model. J. Slowey recommended that the version of the Jacchia-Roberts model fitted to the J70 model be retained in the system, but that equations representing low altitude geomagnetic activity in the J71 be substituted for those in the J70 model. Dr. Forbes gave some reasons why it would not be advisable to add the MSIS-77 model to their system at this time. These included:

- (a) The MSIS represents a fit to composition and temperature measurements, not total mass densities;
- (b) Data utilized to construct MSIS covers a limited range of solar activity and altitude, as compared to the Jacchia models;
- (c) The geomagnetic effect in MSIS-77 is parameterized by the daily Ap index, which is incapable of reflecting time variability as well as the 3-hour Kp index utilized in the Jacchia models;
- (d) MSIS-77 requires about 25% more CPU time that J71; and
- (e) The Millstone Hill (MH) temperatures used in MSIS-77 originated from the 1-pulse experiment at MH, which contain a 50K bias (based on more accurate 2-pulse data analyzed since development of MSIS-77.).

Dr. Forbes also presented comparisons of prediction errors for various models and satellite orbits which have been examined over the years within AFGL. In particular, the Harris-Priester Model, which was discussed for possible consideration by SCF and IBM as part of DSM was shown to exhibit considerably larger prediction errors and about twice the CPU time than J71.

Dr. Forbes and A. Kantor of AFGL attended a meeting of the DARPA Committee for Autonomous Positioning of GPS at Hq ONR on May 1982. At this meeting possible satellites, in

addition to GPS, for which the SHAD (Stellar Horizon Atmospheric Dispersion) navigation system could be used were discussed. Some of these satellites are classified. Density variability in the 20-50Km altitude region could provide a limitation to the navigation accuracy. A. Kantor discussed his planned study of day-to-day and seasonal variabilities, including the statistical quantities to be provided, and Dr. Forbes presented his plans to provide information on the tidal variations. In addition, in response to Dr. F. Quelle's (ONR) request for possible science that could be derived from the SHAD experiment, Dr. Forbes offered a number of suggestions, falling within the following scientific areas:

- (a) Global Climatology
- (b) Polar Stratospheric Warmings
- (c) Planetary-Scale Waves
- (d) Equatorial Waves
- (e) Mesospheric Particulate Layers

(f) Minor Constituents

which were subsequently incorporated with other recommendations into a letter from Dr. Champion to Dr. Quelle.

During 1982 the capability was developed to derive neutral thermospheric temperatures from ionospheric

parameters (Te, Ti, Ne) measured by incoherent scatter radars at Millstone Hill (42°) and Arecibo $(18^{\circ}N)$. The purpose of this project was to examine the equatorward propagation and extension of high-latitude geomagnetic disturbance effects. To develop the above capability involved the following tasks:

- Acquisition of the following magnetic tapes of Arecibo data from the World Data Center:
 - (a) 19/129 (August 1974 May 1977)
 - (b) N35 (December 1971 December 1972)
 - (c) N12 (July 1966 December 1970)

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The tapes initially acquired contained formatting errors and additional problems. After extensive attempts at analysis and conversations with the World Data Center, reformatted tapes were acquired and subsequently analyzed.

(2) Development of programs to read and plot Te, Ti, and Ne off acquired tapes. A sample plot is illustrated in Figure 1. For every set of vertical profile measurements (available on the order of every 15 minutes) of Te, Ti and Ne, a determination had to be made of the quality and acceptability of the data for further analysis. Sometimes instrument or power



Figure 1. Sample vertical profiles of ionospheric parameters derived from Arecibo incoherent scatter measurements on day 260 of 1974.

problems suggested the rejection of data. Occasionally, unrealistic or inexplicable observations (for instance Ti > Te) had to be rejected. The editing of Te, Ti and Ne data prior to neutral temperature determinations (see following paragraph) is a tedious step in the analysis. ビビアンビア

- (3) Acquisition of NEUTEMP program from Dr. W. Oliver of Millstone Hill and program redevelopment for use on the AFGL CDC machine for calculation of Tn. The Millstone Hill facility uses a Harris computer. The NEUTEMP program iteratively solves the ion thermal balance equation (given Te, Ti, Ne) to produce neutral thermospheric (exospheric) temperatures.
- (4) Tailoring of the Multiple Linear Regression Analysis (MLRA) Program and development of plotting package for Millstone Hill neutral exospheric temperature analysis on the CDC machine.

The first Arecibo experiment analyzed, and the one used to develop and test the above programs, is day 260 of 1974. This day was characterized by an F10.7 value of 89 and a daily magnetic Ap index of 3 (Kp less that 2- throughout the day). The neutral exospheric temperatures obtained are illustrated in the lefthand portion of Figure 2. The MLRA yielded the following Fourier components: mean = 842K,





diurnal = 158K, semidiurnal = 24K, and terdiurnal = 22K. Local times of maxima were: diurnal = (15.2), semidiurnal = (3.0, 15.0) and terdiurnal = (4.0, 12.0, 20.0). The following day (September 18) was characterized by significant geomagnetic activity, with Kp values of up to 5+. The quiet-day least-squares fit was extrapolated into the second day to delineate the different temperature response during the magnetically disturbed period (See Figure 2). Temperature differences from the quiet-day behavior in the range of 100-200K are clearly evident.

Exospheric temperatures were also derived from incoherent scatter data at Millstone Hill $(42^{O}N)$ during the geomagnetically disturbed (Kp < 50) day of June 3, 1971. A magnetically quiet day (Kp < 2-; F10.7 = 108.3) with similar F10.7 was also analyzed as a reference day. Unfortunately, the quiet-day data contained strong oscillations suggesting the presence of gravity waves, rendering questionable the use of a fit to these data as a quiet-day reference. Therefore, the separation of the magnetic storm contribution from the June 3, 1971 data was not possible.

The next step in the incoherent scatter project was to have been the analysis of several storms simultaneously observed at Millstone Hill and Arecibo during the solar maximum period 1979-1981. However, besides the March 22, 1979, storm to be described below in conjunction with the

analysis of satellite accelerometer-measured densities, this effort was never completed. The reasons for this were twofold. First, initial postponement of the effort occurred due to the higher priority given to the interpretation of satellite winds and densities during geomagnetically active periods. described below. Then, approximately as simultaneous with the principal investigator moving to Boston University, a new and unanticipated effort relating to the analysis of data and development of inputs to the Air Force Reference Atmosphere in the 80-120Km height region was given first priority by the Air Force. The subsequent subcontract to Dr. Forbes at Boston University only provided for (a) the 80-120Km region data analysis and (b) satellite density and wind analysis during magnetically disturbed conditions, and did not fund completion of the incoherent scatter work. Work completed under item (b) is reported below and the results of (a) are included in Section 3.

The work involving analysis and interpretation of satellite accelerometer-measured densities and winds under geomagnetically- disturbed conditions is reported fully in the following publications:

Marcos, F.A., and J.M. Forbes, Thermospheric Winds from the Satellite Electrostatic Triaxial Accelerometer System, J. Geophys. Res., 90, 6543-6552, 1985.

Forbes, J.M., and F.A. Marcos, Thermospheric Winds, Densities, and Temperatures during an Isolated Magnetic Storm. (CDAW-6 Interval), J. Geophys. Res., accepted for publication, manuscript in revision, 1985.

In the first paper a new thermospheric wind measurement technique is reported which is based on a Satellite Electrostatic Triaxial Accelerometer (SETA) System capable of accurately measuring accelerations in the satellite's in-track, cross-track, and radial directions. Winds measured between 170 and 210 Km during a 5-day period of mostly high geomagnetic activity are analyzed to demonstrate the potential contributions of SETA data to studies of thermospheric dynamics. The data are consistent with a by a two-cell polar circulation pattern characterized trans-polar flow parallel to the 1600 h/0400 h geomagnetic local time meridian, and return flows in the late morning and late evening sectors. The flow pattern was found to be asymmetric in that it is displaced about 5-10⁰ latitude towards the noon (MLT) sector, and the evening cell is somewhat more diffuse than the morning cell. The system also covers a greater area of the polar cap and is more intense during active $(Kp > 5^{\circ})$ than guiet 3°) (Kp geomagnetic conditions. Average trans-polar flow velocities are characteristically of order 150 \pm 75 ms⁻¹ for Kp ~~ 2^o and 375 \pm 100 ms⁻¹ for Kp ~~ 5°, the stated variabilities

representing 15 deviations. Details may be found in the above-cited publication.

In the second work, thermospheric winds and densities form SETA (Satellite Electrostatic Triaxial Accelerometer) measurements and exospheric temperatures form Millstone Hill incoherent scatter radar data are examined during an isolated magnetic storm occurring on March 22, 1979, a period which coincides with the CDAW-6 interval. A polar thermospheric wind circulation suggestive of a two-cell horizontal is reflected convection pattern in the cross-track acceleration measurements. Winds are generally trans-polar and appear to flow parallel to the 1600h/0400 MLT meridian with return flows in the morning and evening sectors near 57.5° + 7.5° and 70° + 7.5° geomagnetic latitude, respectively. Winds are of order 100-200 ms⁻¹control during quiet periods, and attain maximum speeds between 300 and 600 ms⁻¹during the storm. The two-cell pattern is distinctively more well-ordered in geomagnetic rather than geographic coordinates. The time delay between maximum magnetic disturbance as reflected by the 3-hour ap index and full set-up of the circulation pattern is somewhere between 0-3 hours, the uncertainty being due to the discrete nature of ap and the 90-minute orbital period. The circulation pattern persists almost unattenuated for about 6 hours after

the magnetic disturbance has returned to quiet levels, reflecting the so-called 'flywheel' effect.

The density response is highly asymmetric with respect to its day/night behavior. At high geomagnetic latitudes (> 60°) the daytime density variation, about a 40% increase form quiet levels, distinctively reflects the response to an increase in magnetic activity with a time delay of less than 3 hours. At lower latitudes the response is smaller (~~20%) and less well-defined, occuring with time delay of about 6 + 2 hours near 10-20° latitude. The nighttime density response, on the other hand, is not so well defined poleward as it is equatorward of 60° latitude. The time delay increases from about 2 \pm 2 hours to 6 \pm 2 hours from high to low latitudes.

An exospheric temperature response occurs at Millstone Hill $(42^{\circ}N)$ with an amplitude of 210 K and time delay of between 0 and 2.5 hours.

Latitude structures of the density response at successive times following the substorm peak suggest the equatorward propagation of a disturbance with a phase speed of between 300 and 600 ms-1. A deep depression in the density at high latitudes (>70°) is evident in conjunction with this phenomenon. The more efficient propagation of the disturbance to lower latitudes during the night may be connected with the 'midnight surge' effect, a reinforcement

of southward winds associated with high latitude heating and solar EUV heating, as opposed to cancellations between these meridional flows during the daytime. Details of this study may be found in the above-cited manuscript.

2. ATMOSPHERIC TIDES STUDIES

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Atmospheric tides constitute much of the meteorology and variability in atmospheric fields above 70 Km altitude, and represent a major uncertainty in specifying such as environmental conditions for a variety of Air Force Experimental data often suffer from inadequate missions. coverage in local time and/or latitude to properly delineate tidal motions, so greater reliance on theory and numerical simulations must be made in this discipline of study. The following work was performed to better understand the nature oscillations and the uncertainty which of tidal thev contribute to density and temperature specifications in the upper atmosphere.

A paper entitled "Neutral Temperatures from Thompson Scatter Measurements: Comparisons with the CIRA (1972)" by J. M. Forbes, M. E. Hagan, and K. S. W. Champion was prepared for presentation at the COSPAR Meeting, Ottowa, Canada, May, 1982, and was published in Space Research, Volume 3, 1983. In this paper, Neutral exospheric and lower thermospheric (100-130 km) temperatures from Thomson scatter measurements

Millstone Hill (42^ON) are compared with CIRA at temperatures with a view towards identifying deficiencies in the CIRA and recommending revisions. CIRA is found to model the observed diurnal mean temperatures (T_0) to within 10% over a wide range of solar conditions (75< F10.7< 250), but consistently underestimates the diurnal temperatures with maximum deviations approaching 50% of observed amplitudes (180-240 K) at solar maximum $(1200 \text{ K} < T_0 < 1400 \text{ K})$. The observed semidiurnal amplitudes, which lie in the range of 20K-80K, are always underestimated and frequently by more than 50%. In the lower thermosphere, tidal oscillations of temperature of order 20K-40K occur which are not modelled by CIRA. In addition, an analysis of exospheric temperatures at Millstone Hill during a magnetic disturbance indicates a response within 1-2 hours from storm onset, whereas CIRA assumes a 6.7 hour delay. As a result, and although some of these deficiences are addressed by the more recent MSIS model, there exists a sufficient data base to recommend several additional revisions to the CIRA temperatures at this time.

A review of thermosphere tides was prepared as AFGL Report TR-82-0264 entitled "Tides in the Thermosphere: A Review" by J.M. Forbes and K.S.W. Champion. In it a comprehensive review of recent theoretical and observational accomplishments relating to diurnal and semidiurnal tidal

oscillations of neutral winds, temperature, density, and composition above 100 Km is presented. Topics emphasized include: (1) Recent theoretical studies; (2) Solar cycle, seasonal, and latitudinal variations in tidal oscillations of temperature and winds as inferred from Thomson scatter measurements: (3) Tidal variations in total mass density and composition as inferred from satellite accelerometer and mass (4) Comparison of spectrometer measurements; recent theoretical models with the above observations; (5) The relative influence of in-situ and propagating tides in determining the total semidiurnal thermospheric tide; and (6) Propagating tides of lower atmosphere origin as a source of mean momentum and heat in the lower thermosphere.

As an input to the SHAD (Stellar Horizon Atmospheric Dispersion) Program, a draft report was written by Dr. Forbes entitled "Density Variability due to Atmospheric Tides below 140 Km". In this report equations and a computer algorithm were derived for the tidal oscillations in density compatible with temperatures from the numerical tidal model of Forbes. Vertical profiles and tables of tidal amplitude and phase corresponding to 0° , 30° , and 60° latitude under equinox conditions were included for the diurnal and semidiurnal components, respectively. Diurnal amplitudes were found to be of order 1-2% in the stratopause region (ca. 50 km) and 5-15% in the mesopause region, attaining

amplitudes of 7-25% near 110 Km where the first diurnal propagating (1,1) mode begins to experience severe molecular dissipation. The phase structure at 0° and 30° latitude in fact revealed a nominal 25-30 Km vertical wavelength characteristic of the (1,1) mode. A constancy of phase with height above 50 Km for latitudes >~ 30° was noted to be indicative of the trapped (1,-2) tidal mode.

semidiurnal oscillations found The were to be significantly smaller than the diurnal component below 100 Km, of order .2-1% around the stratopause and 2-6% in the mesopause region. Amplitudes of 20-25% are attained in the vicinity of 120 Km. Semidiurnal vertical wavelengths below are much larger than for the diurnal 70 Km tide. characteristic of the fundamental (2,2) mode. Above 80 Km the presence of the shorter-wavelength (2,4) mode is evident, due primarily to mode coupling processes involving the mean zonal wind and meridional temperature gradient distributions.

Over most of the 0-140 Km altitude regime the largest density perturbations generally occur at equatorial latitudes for the diurnal and semidiurnal tides. However, considerable latitude structure is exhibited at any given level. A discussion was also included of the possible effect of non-migrating tidal components. (i.e., those which do not migrate with the apparent motion of the sun). It was noted that wind data from the Jicamarca, Peru (12⁰s) incoherent

scatter radar indicate the presence of non-migrating components with amplitudes as much as an order of magnitude larger than the migrating contribution. Thus the potential exists (at least at some longitudes) of significantly larger tidal density variations in the low-latitude stratosphere than given for the migrating components. It was recommended that SHAD density data from the HEAO experiment, and in the possession of the MIT Draper Laboratories, should be able to non-migrating tidal components reveal in stratospheric density if they exist.

The density work was subsequently expanded upon to provide inputs to the point-analysis and Air Force Reference Atmosphere efforts, and was written up in as a proposed AFGL report entitled "Estimates of Point-Density Errors due to Atmospheric Tides between 70 and 120 Km" by J.M. Forbes. This latter report includes various contours of "percent density departures from mean values due to tides" as a function of height, latitude, local time and season.

3. TEMPERATURES, PRESSURES, AND DENSITIES IN THE 80-120 Km REGION

3.1 Introduction

At the present time the standard reference for temperature, density, and wind specification between 80 and

120 Km is the 1972 COSPAR International Reference Atmosphere (CIRA 1972). Temperatures and densities are given as monthly averages at 5 Km height increments every 10° in latitude for the Northern Hemisphere. However, data are extremely sparse in this region; as Groves (see CIRA 1972) points out, CIRA 1972 suffers from the same limitations as CIRA 1965 above 60 Km:

 Data from all longitudes are combined without consideration of longitudinal effects. Most data are from N. American sites, and so any longitudinal bias would be towards the W. Hemisphere.

- (ii) Insufficient S. Hemisphere data were available for developing a separate model. Therefore, S. Hemisphere data were combined with N. Hemisphere data with a six-month change of date.
- (iii) Due to insufficient data, no account was taken of local time in development of the models. Consequently the temperature and density fields may be diurnally-biased as well.

The current Air Force Reference Atmosphere (Cole and Kantor, 1978) only extends to 90Km. Development of the temperature and density inputs between 80 and 120 Km for the new Air Force Reference Atmosphere (AFRA) has involved building upon the existing CIRA 1972 model data base (see section 3.2). Since the greatest abundance of experimental data is in the form of temperatures, and since temperature is one of the meteorologial fields (besides winds) for which we have a firmer theoretical and intuitive base for understanding its behavior and structure, the temperature field is the basis for development of the model. Following the procedure

followed in CIRA, given a specification of the pressure field at some lower boundary, the density field at higher altitudes were derived from the barometric and ideal gas laws. Available experimental density determinations are used to provide consistency checks on these densities. A preliminary analysis of this data was written up in the form of a report "Temperature Structure of the 80 to 120Km Region" by J.M. Forbes, as input to the COSPAR working Group on the new Cospar International Reference Atmosphere. The report has been published in the MAP Handbook Series, Volume XVI.

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3.2. Description of the Temperature Data Base

The following is a description of the data which has formed the basis for the 80-120 Km specification of temperature, density and pressure recommended for the new AFRA:

(i) The CIRA 1972 temperature model between 80 and 120 Km was based heavily on data from the NASA Sounding Meteorological Rocket Program (MSRP) collected prior to 1967. The primary techniques were pitot tube and grenade measurements. The MSRP was phased out in 1973. Between 1967 and 1973 32 pitot tube and 135 rocket grenade experiments were conducted which generally yielded temperature data

above 80Km. The soundings were made at Wallops I. ($38^{\circ}N$, $75^{\circ}W$), Ft. Churchill ($59^{\circ}N$, $94^{\circ}W$), Pt. Barrow ($71^{\circ}N$, $157^{\circ}W$), Natal ($6^{\circ}S$, $35^{\circ}W$), Arecibo ($18^{\circ}N$, $67^{\circ}W$), Arenosillo ($37^{\circ}N$), Eglin (30° , $87^{\circ}W$), and Kourou ($5^{\circ}N$, $53^{\circ}W$) (Smith <u>et al</u>, 1967, 1968, 1969, 1970, 1971).

(ii) Gaigerov et al(1984) present extensive analyses of temperatures from Soviet rocket measurements (mostly grenade method) north and south of the equator in the Eastern Hemisphere. Most of the Southern Hemisphere data were collected after 1970, and hence were not included in the 1972 CIRA. Gaigerov et al also discuss the results of several intercomparisons with other independent measurements of temperature, and report that adjustments have been made for any biases which might have existed in their raw data. Monthly average data between 80 and 90-100Km are available from Heiss I. (81°N, 58°E), Volgograd (48°N, 44^oE), Molodezhnaya (68^oS, 45^oE), and from research vessel soundings in the Pacific near 0° . 20°S, 40°S, and 50°S. Keeping in mind known deficiencies in the Soviet data, this information

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now enables us to examine possible longitudinal and latitudinal asymmetries in the structure of the mesopause region. S. Hemisphere data may in fact introduce an undesired bias in the AFRA if mixed with N. Hemisphere data, and so this aspect of the analysis has been pursued with caution.

(iii) Five years (1970-1975) of temperature profiles from incoherent scatter measurements in the E-region (100-130Km) over Arecibo, Puerto Rico (18⁰N. 67⁰W) and Millstone Hill, Massachusetts (42⁰N, 71^OW) were available for analysis. These data have been analyzed in parts, mostly with regard to the semidiurnal oscillation by Salah (1974), Salah et al (1975), Salah and Wand (1974), and Wand (1976, 1983).This investigator has pooled all the available data, separated mean and semidiurnal tidal components, and constructed monthly A total of over 1,500 profiles (each) averages. are available from Arecibo and Millstone Hill. These data are considered extremely important in terms of "matching" the rocket-based temperature of structure the 80-100Km region with the satellite-based density and temperature fields above 150Km.

A list of stations from which rocket and radar data has originated for the new AFRA is included in Table I. A preliminary analysis was first performed to ascertain whether evidence exists for longitudinal and latitudinal asymmetries in the temperature structure of the 80Km to 100Km region, and whether sufficient data are available to delineate these dependences in a reference atmosphere.

TABLE I

Locations of rocket* measurements and incoherent scatter radar** measurements which form the basis data for the new AFRA between 80Km and 120Km.

"Western Hemisphere"

"Eastern Hemisphere"

	Thule	(76 ⁰ N, 69 ⁰ W)
	Pt. Barrow	$(71^{\circ}N, 157^{\circ}W)$
	Ft. Churchill	$(59^{\circ}N, 94^{\circ}W)$
k 🖈 👘	Millstone Hill	$(42^{\circ}N, 71^{\circ}W)$
	Wallops I.	$(38^{\circ}N, 75^{\circ}W)$
	White Sands	$(32^{\circ}N, 106^{\circ}W)$
	Eglin	$(30^{\circ}N, 87^{\circ}W)$
	Cape Kennedy	$(28^{\circ}N, 80^{\circ}W)$
	Barking Sands	$(22^{\circ}N, 159^{\circ}W)$
t *	Arecibo	$(18^{\circ}N, 67^{\circ}W)$
	Antigua	$(17^{\circ}N, 62^{\circ}W)$
	Ft. Sherman	$(9^{\circ}N, 80^{\circ}W)$
	Kourou	$(5^{\circ}N, 53^{\circ}W)$
	Natal	$(6^{\circ}S, 35^{\circ}W)$
	Ascension I.	$(8^{\circ}S, 14^{\circ}W)$

Heiss I.	$(81^{\circ}N, 58^{\circ}E)$
Volgograd	$(48^{\circ}N, 44^{\circ}E)$
Sardinia	$(40^{\circ}N, 10^{\circ}E)$
Guam	$(13^{\circ}N, 145^{\circ}E)$
Kwajalein	(9 ⁰ N,168 ⁰ E)
Thumba	$(8^{\circ}N, 77^{\circ}E)$
Res. Vessels	(0 ⁰)
Res. Vessels	(20 ⁰)
Carnavon	$(25^{\circ}S, 114^{\circ}E)$
Woomera	$(31^{\circ}S, 136^{\circ}E)$
Res. Vessels	(50 ⁰ S)
Kerguelen I.	(49 ⁰ S)
Res. Vessels	(50 ⁰)
Molodezhnaya	(68 ⁰ S, 45 ⁰ E)

- With some exceptions, data are generally available between * 80-100Km
- Data generally available between 100-130Km

In Figure 3 variations in monthly temperatures at 85Km for individual stations are compared. The comparison between Barrow $(71^{\circ}N. 157^{\circ}W)$ $(81^{\circ}N. 58^{\circ}E)$ and Heiss T. Pt. suggests 5-10⁰K higher temperatures at Heiss I. in the summer and 5-10° cooler temperatures between January and Although these two stations are separated by 10° March. in latitude, the discrepancy is opposite to what one might the positive (negative) pole-to-equator from expect temperature gradient assumed to exist in Northern Hemisphere summer (winter) months. An examination of vertical structures at the two stations indicates that the summer mesopause minimum is near 90Km at Heiss I. as opposed to 85Km at Pt. Barrow, and this is in itself is an important contributor to their differences in monthly behavior at 85Km.

Although the Volgograd ($48o_N$, $44^{O}E$) data during summer exhibit 10-15^oK higher temperatures that Ft. (59[°]N, 94[°]W), their 11[°] separation Churchill in latitude is sufficient to account for more than half of this difference assuming a realistic latitude gradient in temperature (see following figures). Figure 3 also suggests a much larger temperature gradient in the eastern than western hemisphere, but it must be remembered that Ft. 120 Churchill and Pt. Barrow are only separated by latitude, whereas the separation between Heiss I. and Volgorad is 330 in latitude. Obviously, to make any

convincing statements about latitude structure we should examine all possible data at a given height. This will be done below. Before leaving Figure 3, note that Southern Hemisphere data at 40° S and 50° agree quite well with the Northern Hemisphere data at similar latitudes.

In Figures 4 and 5 the latitude structures of temperature during summer (mostly July) and winter (mostly January), respectively, are depicted at 80, 90, and 100Km. An obvious feature of these plots is that the Eastern/Western/Northern/Southern Hemisphere data collectively delineate fairly well-defined patterns.

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Figure 6 depicts a selection of temperature measurements between 100 and 130Km at various latitudes during July and January. Comparisons are made with the MSIS-83 model (Hedin et al, 1983), as this is a likely candidate for the new AFRA above 100Km or so. The data shown in Figure 6 during January are extremely consistent with each other and with the MSIS-83 and do not exhibit any significant latitude model. structure. During July, however, there appears to exist a significant positive equator-to-pole temperature gradient. At 115Km, the temperature varies from about 268^OK at Kwajalein $(9^{\circ}N)$ to $320^{\circ}K$ at Arecibo $(18^{\circ}N)$ to $370^{\circ}K$ at Wallops I. (38^oN) and Millstone Hill (42^oN). A small temperature difference (20°K) of this sense between 18°N and 42°N is specified in the MSIS-83 model. Compared with the earlier AFRA, the present data set provides improved



Figure 3. Temperature vs. month at 85Km for various stations which allow examination of possible longitudinal or hemispherical asymmetries within specific latitude belts. CIRA 1972 values are shown for comparison.

seasonal, latitudinal, and longitudinal coverage in the 80-100Km region, and a combination of incoherent scatter and rocket data in the 100-120Km region which allows a much improved delineation of lower thermosphere temperature At the same time, although some individual structure. station comparisons indicate measurable asymmetries in longitude and latitude, data are still insufficient to separate these effects; that is, to provide a reliable description of latitude structure as а function of longitude, or of longitude structure at any given latitude. It is therefore recommended that in order to obtain the best description of Northern Hemisphere temperatures, pressures, and densities between 80 and 120Km, one must combine together data from all longitudes and from the Southern Hemisphere (with a 6-month change of date) in order to obtain even a minimally acceptable distribution of data. Thus the AFRA model input as it is provided here is not capable of distinguishing longitudinal or hemispheric dependences if they exist. Soviet data (which has been bias-corrected) has also been included in the data set. Inclusion of Southern Hemisphere and/or Soviet data may arouse some concern. As it turns out, both of these data show remarkable consistency with the North American data, and in some cases fill crucial gaps or extend the data set and model credibility to higher latitudes than would otherwise be unattainable. High



Figure 4. Temperature vs. latitude for measurements representative of July north of the equator in the Western (\bullet) and Eastern (χ) Hemispheres, and January south of the equator in the Eastern Hemisphere (+). Where data under these exact conditions were no available in the Western Hemisphere, data points form 8° August (\Box), 6° August (Δ), August (\bullet), and June (\odot) were inserted to allow a more complete delineation of the latitude structure. CIRA 1972 values are shown for comparison.

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Figure 5. Temperature vs. latitude for measurements representative of January north of the equator in the Western (\bigcirc) and Eastern (\checkmark) Hemispheres. Where data under these exact conditions were not available in the Western Hemisphere, data points from 6°S February (\Box), 6° December (Δ), and February (O) were inserted toa llow a more complete delineation of the latitude structure. CIRA 1972 values are shown for comparison.



Figure 6. Vertical structures of temperature from 100Km to 130Km at various latitudes. MSIS-83 values are shown for comparison.

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latitudes show the greatest degree of variability, so it was important that every possible piece of information be considered in development of the model.

3.3 Method of Analysis

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The data set (consisting of monthly averages at the observing sites) is distributed unevenly in latitude, height, and by month. In order to construct monthly tables at 15° latitude increment, it is necessary to fit the data in a least-squares sense using appropriate functions, and then to use these formulas to specify the data at 150 latitude increments. For a given month and altitude the best latitudinal fit with the least number of terms was found to be given by a function with the following form (0 = colatitude):

 $T(z) = T_0(z) + A_1 \cos^2 \theta + A_2 \cos^3 \theta + A_3 \cos^4 \theta$

The fit was made to the data set from 70 to 120Km every 5Km for every month of the year. (A total of 11 X 12 = 121 fits). The polynomials were then used to generate temperature tables every 15° latitude, every 5Km (70-120Km), for every month. These data do not, however, represent a smooth profile at any given latitude. The next step was to fit an analytic function to every 15° latitude data set (for every month) between 70Km or so and 120Km such

that the temperature and temperature derivative match the existing USAF Reference corresponding values from the Atmosphere at some lower boundary. Specifically, the following procedure was found to work well: Only temperature data from 80 to 120Km in 5Km increments was included in the fit. Additional data points at 68, 72, 74, and 78Km were obtained from the old AFRA tables. A fourth-order polynomial in altitude was fit to all these data points. In all cases the fit was extremely close to the data points in the 70-90Km region, and formed a well-defined minimum (the mesopause) Tables were then generated every 2Km from the near 90Km. polynomial fit. A final adjustment at the lower boundary made such that the 68Km point (generated from the was polynomial) was replaced by the actual temperature value from the old AFRA tables, and the value at 70Km was replaced with the value obtained by linearly interpolating between the value at 68Km and the polynomial-generated value at 72Km. This ensured exact matching of temperature with the old AFRA very close approximate) at 68Km, (but matching with temperatures and gradients from the old AFRA between 70 and the newly-defined mesopause 80Km, merging into region (85-95Km) and lower-thermosphere (100-120Km) regions above.

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It is important to note two points relating to the method analysis as adopted above. First, no attempt was made to impose or fit any functional dependence vs. month at a

given height or latitude. Second, since the temperature field at 120Km specified by the upper portion of the AFRA (which will probably be the MSIS-83 model) is dynamic and dependent on solar, geomagnetic, and other variations, a single match could not be performed at these levels to conform with the tabular nature of the model below 120Km.

3.4 Tables of Temperature, Pressure, and Density

Tables of monthly Temperature, Pressure, and Density extending from 70Km to 120Km in 2Km height increments at 0⁰. 15⁰N. 30⁰N. 45⁰N, 60⁰N, and 75⁰N are appended to this report. These tables are in exactly the same format as the existing Air Force Reference Atmosphere Tables (Cole and Kantor, 1978). The pressure tables were calculated using the barometric law and a reference pressure at 68Km equal to that given by the AFRA (Cole and Kantor, 1978). The latitude dependence of the acceleration due to gravity at sea level was taken from Cole and Kantor (1978), and the formula for the height dependence of g was taken from the 1966 U.S. Standard Atmosphere Supplements. The height dependence of the mean molecular weight above 90 Km was adopted from the MSIS model (Hedin et al, 1983) for equinox conditions, an F10.7 cm solar flux of 100 and Ap=15; the values for M from 80 to 120Km in steps of 10Km were as follows: 28.96, 28.90,

28.46, 27.42, 26.23 in a.m.u.. A linear interpolation was utilized at intermediate altitudes.

3.5 Variability

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In addition to the marked seasonal variability depicted in the temperature, pressure, and density tables, within a given month or even a single day there exist significant variations about mean values due to gravity waves, tides, and perhaps other sources. Yet, while the Tables are based on temperature because it is the more globally available measured field, the variability characteristics of total mass density are of much greater Air Force operational concern. Therefore, density information from several typical stations quantitatively characterize used here to the are Specifically, variability of density between 70 and 110Km. the data employed originate from NASA MSRP data from Pt. Barrow (71^oN), Ft. Churchill (59^oN), Wallops I. (38^oN) and the Natal $(6^{\circ}S)$ - Ascension $(8^{\circ}S)$ pair. (Smith et al, 1964; 1966,1967,1968,1969,1970,1971; Theon et al, 1972). The monthly and local time distributions from each of these sites is given in Figure 7. Of the 227 total soundings grenade soundings represented, 207 (which are yield temperature and horizontal winds as the primary variable, from which pressure and density are derived) and 20 pitot



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probe soundings (which actually measure density). While there exist scattered falling-sphere density data which will enable important consistency checks on the current density NASA MSRP tabulations. the data represent the best statistical base to provide information on density variability as a function of season and latitude.

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Percent Standard deviations of temperature and density from seasonal-mean values at Pt. Barrow, Ft. Churchill, and Wallops I. are given in Tables II-IV. Seasonal means are utilized to provide better statistics given the small number of soundings available. Values based on the combined stations of Natal (6° S) and Ascension I (8° S) are based on annual means in Table V due to the small seasonal variations present near the equator (Theon <u>et al</u>, 1972). Note that the standard deviations for temperature are generally smaller than for density, the latter values typically lying between 5-20% between 70 and 110Km.

Theoretical considerations must be invoked to explain the origin of the above variations. Note from Figure 7 that in toto the data are distributed over all parts of the day, with some bias towards the 1200-2000 hour local time period. However, within a given season the local time distribution is much more uneven, essentially precluding determination of tidal components; furthermore, by the same token, the

tabulated standard deviations do not fully account for tidal Since tides represent global oscillations with effects. fairly well-known excitation and propagation characteristics, useful theoretical estimates of this variability source can be obtained. The model of Forbes (1982a,b) has been utilized to produce such estimates. Figure 8 illustrates percent density departures from diurnal mean values for 80 and 100Km between 0 and 60⁰ latitude. An important feature to note tidal is the small-amplitude oscillation predicated polewardof about 54⁰ latitude. This implies that the density variations listed in the TABLES for Pt. Barrow (71°N) and Ft. Churchill (59°N) may be presumed to be predominantly a reflection of gravity wave effects. In fact, the height dependence of density standard deviations at Pt. Barrow is similar to that of radar echoes (and hence turbulent irregularities) observed with the MST Radar at Poker Flat (65⁰N;

THELE II

PERCENT STANDARD DEVIATIONS OF SEASONAL - MEAN TEMERATURES AND DENSITIES AT PT. BARRON (710A)

		MINTER			SUMMER		SPRING/FALL				
ALT (Rm)	TEMP ST.D.	DENSITY ST.D.	NO. OBSV.	TEAP ST.D.	DENSITY ST.D.	NO. OBSV.	TIMP ST.D.	DENSITY ST.D.	NO. OBSV.		
70	6	20	18	3	6	10	9	24	16		
74	7	19	18	2	7	9	5	26	16		
78	9	17	18	2	8	9	6	29	16		
82	10	13	17	2	7	8	9	29	16		
86	9	12	17	4	8	8	13	26	13		
90	9	10	15	7	10	7	14	21	11		
94	5	7	6	24	14	2	2	11	3		

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TABLE III

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PERCENT STANDARD DEVIATIONS OF SEASONAL - MEAN TEMERATURES AND DENSITIES AT FT. CHURCHILL (599N)

		MINISTR			SUMMER		SPRING/PALL				
ALT (Rm)	TEMP ST.D.	DENSITY ST.D.	NO. CBSV.	TEMP ST.D.	DENSITY ST.D.	NO. CB <i>S</i> V.	TEMP ST.D.	DENSITY ST.D.	NO. CB <i>S</i> V.		
70	5	17	29	2	6	12	3	19	13		
74	7	17	29	3	6	12	3	20	13		
78	8	16	28	4	7	12	5	21	13		
82	9	14	27	6	9	12	6	19	13		
86	6	14	26	6	9	12	6	8	12		
90	11	13	19	10	15	7	12	15	9		
94	5	10	11 .	.5	8	2	10	3	4		
98	3	9	4	-	-	1					
102	7	13	4	-	-	1					
106	6	7	4	-	-	1					
110	16	16	4	-	-	1					
114	16	8	4	-	-	1					
118	15	11	4	-	-	I					

TIBLE IV

PERCENT STANDARD DEVIATIONS OF SEASONAL - MEAN TEMPERATURE AND DENSITIES AT WALLOPS I. (38°N)

WINTER					SIMER		SPRING/PALL				
ALT (Rm)	TEMP ST.D.	DENSITY ST.D.	NO. OBSV.	TEMP ST.D.	DENSITY ST.D.	NO. OBSV.	TEMP ST.D.	DENSITY ST.D.	NO. OBSV.		
70	4	9	28	4	9	23	4	13	38		
74	7	10	25	5	8	23	6	13	38		
78	7	8	21	7	9	20	7	14	36		
82	7	ĭ1	16	6	9	18	8	12	32		
86	6	10	10	8	14	16	8	10	31		
90	iı	18	6	8	13	13	14	14	19		
94	-	-	i	6	8	5	3	6	3		
98			_	9	1	4	3	.6	2		
102				14	5	3	7	1	2		
106				8	19	3	2	6	2		
110				11	27	3	.1	10	2		
114				6	25	2	4	11	2		
118				5	27	2	8	11	2		

			1/151	as v				
PERCE	t S	MDARD	DEVIA	TIONS	OP	ANNUAL -	MEAN	
TEMERATURES	AND	DENSIT	ies ai	I NATA	- J	ASCENSION	1 (6 TO	8°5)

ALT (RR)	TEMP ST.D.	DENSITY ST.D.	NO. CBSV.
70	4	6	32
74	5	9	31
78	6	7	30
82	5	10	30
86	8	10	30
90	11	12	29
94	14	13	12
98	22	16	3
102	23	.1	2
106	-	-	1

Balsey et al, 1983); that is, summer is characterized by the predominant occurence of echoes between 80 and 100Km, where as during

winter the echoes appear to originate predominantly below 80Km. The turbulence arises from convective instabilities associated with gravity waves whose level of breakdown is modulated by the seasonally-dependent mean winds of the troposphere and stratosphere.

The latitude dependence of gravity-wave related perturbations in the 70-100Km height region is not known. However, there is no reason to expect that gravity wave activity should increase significantly with decreasing latitude. The data in Figure 8, on the other hand, predict that tidal effects increase markedly as one gets closer to

the equator, suggesting that tides predominate equatorward of some intermediate latitude. The standard deviations of temperature and density in the Tables are in fact consistent with the temperature computations presented by Forbes(1982a,b) and the densities in Figure 8. At middle latitudes, say 30-55° latitude, it is not unreasonable to assume that tides and gravity waves contribute about equally to the observed variability of the region.

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In the 100-120Km regime gravity waves most likely exist with amplitudes no greater than their convectively-limited values at lower levels, whereas tides experience exponential growth with height throughout the region. In addition, other sources of tides such as UV and EUV solar radiation absorption assume importance above 95-100Km. Estimated point density errors at 120Km as presented in Figure 9 commonly occur in the 15-30 % range. This order of magnitude is supported by the semidiurnal temperature perturbations obtained from an analysis of Millstone Hill (42⁰N) and Arecibo (18⁰N) temperatures performed by Dr. Forbes and presented in Figure 10. Figure 10 illustrates seasonal averages of the semidiurnal temperature oscillations as determined from 25 days of simultaneous measurements at Millstone Hill and Arecibo. These data represent a subset of the profiles utilized in constructing the monthly and diurnal-mean profiles discussed in Section 3.2. The



Figure 8. Estimated Percent Density derivations from the mean at equinox due to atmospheric tides.



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Figure 9 Estimated percent density derivations from mean at equinox due to atmospheric tides.

Millstone Hill amplitude profile is characterized by a peak near 110-115Km which varies in amplitude from 45K (summer, 60K (winter). The Arecibo equinox) to temperatures generally do not exhibit a well-defined peak. Both stations exhibit a downward phase progression with characteristic vertical scales of 30-45Km. Further, the Millstone Hill phases lead those at Arecibo by about 4 hours in summer, and lag the Arecibo phases by 4-6 hours during winter. These amplitude and phase characteristics suggest that the observations reflect the strong presence of (2,4) and (2,5) semidurnal tidal modes propagating upwards from below 100Km.

3.6 Discussion of Densities

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Since total mass densities represent derived quantities appended tables, it is necessary to perform in the consistency checks for validation purposes, and to ascertain what new information these results provide. The average density profiles corresponding to the standard deviations listed in Tables III and IV are used to accomplish this. In addition, since MSIS-83 is a likely candidate for the AFRA at satellite altitudes, comparisons with this model are also made here to aid in the formulation of a means to merge the dynamic model above and the tables at lower heights (see also Section 3.4). These intercomparisons are performed at 90, 110, and 120Km in Figures 11, 12, and 13, respectively.



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Figure 10 Seasonal averages of semidiurnal temperatures derived from 25 simultaneous daytime experiments at Arecibo and Millstone Hill.

The present model is seen to agree extremely well with at 90Km (Figure 11). Both models indicate a the AFRA predominantly annual variation, with a greater amplitude at $60^{\circ}N$ than $30^{\circ}N$. The only marked difference is that at 30^ON the AFRA predicts the density maximum to occur between February and April, whereas the present analysis indicates maximum densities during the April-June period; both models indicate a minimum in density to occur during the August-October period. At 60⁰ latitude the density maxima for both models occur during April-May with minima during The annual variation predicted by MSIS-83 December-January. in good agreement at 60°N, 30°N MSIS-83 but at is contains little variation with month of the year. The densities at Pt. Barrow appear to contain a significant semiannual component, and about half the variability indicated by the other sources at 60°N. The Pt. Barrow data, on the other hand, appear to confirm the annual variation with amplitude and mean value about 20% less than the other models examined.

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At 110Km (Figure 12), the current model indicates a more or less annual variation at high latitudes, but with maximum during December-April and minimum during June-July ... an approximately 180° phase difference from the annual cycle at 90Km! The data at Ft. Churchill are in agreement with the phase and amplitude of this variation, but with a slightly



Figure 11. Data comparisons at 90 Km.



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Figure 12. Data comparisons at 110 Km.

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Figure 13. Data comparisons at 120 Km.

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smaller mean value. The MSIS-83 model also agrees with the sense of this variation, except with a significantly smaller amplitude. The density variation at 30° latitude now reflects a significant semiannual component. This result is not inconsistent with data at Wallops I., except that during winter months there does not exist sufficient data at this altitude to confirm the expected decrease in density. The MSIS-83 variation is still annual at 30°N, with only a 10-12% amplitude and 30% smaller mean value.

At 120Km the semiannual variation at 30° latitude is particularly predominant, while at 60° latitude the distortion of the annual variation also indicates presence of a significant semiannual component as well. The Ft. Churchill data and MSIS-83 model yield similar overall magnitudes similar to the current results at $60^{\circ}N$, but are approximately 30% less at 30° latitude.

3.7 Merging with MSIS-83 above 120 Km

The above results suggest that MSIS-83 is in reasonably good agreement with available data at 90Km, but differs measurably at 110Km and 120Km, particularly at 30° latitude where 20-30% underestimates are evident. Practical use of the new AFRA requires a functionally efficient algorithm for

merging the lower-altitude tables (<120Km) with the MSIS-83 model above. In addition, the MSIS-83 contains other geophysical variations such as those connected with solar and geomagnetic activity which are not relevant at lower altitudes. It is recommended that, at least initially, the merging of these two model segments be accomplished as follows:

(1) Store tabular data below 110Km

- (2) Use MSIS-83 computer algorithm down to 130Km
- (3) Between 110Km and 130Km determine density and pressure by linearly interpolating the natural logarithms of these fields, and determine temperature by direct linear interpolation.

The above method will naturally damp out solar - geomagnetic variations from 130Km to 110Km. However, internal consistency pressure, density, between and temperature will not necessarily be maintained between 110Km and 130Km. This should not be cause for concern for operational utilization of the model. In addition, the first derivatives of these fields will not match at 110Km and 130Km. After initial experimentation with the merging method described above, more sophisticated methods of smoothing, or modification of the exact altitudes of matching, may warrant examination.

Unless specific solar, geomagnetic, and geophysical inputs to MSIS-83 are given, it is not possible to provide a usable or useful set of tables covering the 80-200km regime, unless these would be constructed solely for illustrative purposes; this is because the matching conditions would have to be changed for each set of solar geophysical conditions specified. Rather, the model must by given in computer algorithm form for those who would ultimately make operational use of the model. For these reasons, the tables given here only extend to 120Km, and a recommended method of merging has been provided for implementation in the working version of the new AFRA.

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ר	5.4005 3.9583 2.8512 2.8438 1.4585	1. 8378 7. 3486 5. 1945 3. 6631 2. 5793	1.8148 1.2759 8.9882 6.3476 4.5613	3.2007 2.3006 1.6735 1.2318 9.1866	6.9594 5.3649 4.2189 3.3685 2.7453 2.2797 2.2797
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ר	5.6637 4.1696 3.8463 2.2893 1.5981	1. 1300 8. 6572 5. 6777 3. 9778 2. 7738	1.9267 1.3371 9.2925 6.4885 4.5476	3.2186 2.3844 1.6741 1.2357 9.2848	7.1053 5.5429 4.4005 3.5099 2.9441 2.4783
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ר	5.7415 4.1968 3.8491 2.1979 1.5733	1.1181 7.9014 5.5547 3.8919 2.7268	1.9823 1.3327 9.3878 6.6558 4.7056	3.4493 2.5311 1.8856 1.4279 1.6955	8.5119 5.5553 4.5723 3.8218 3.2388 3.2388
7	5.8038 4.3113 3.1325 2.2618 1.6239	1. 1001 8. 2535 5. 8534 4. 1425 2. 9292	2.0722 1.4692 10.4649 7.4906	3.9448 2.9078 2.1714 1.6453 1.2050	9.8942 6.3462 5.28782 4.3399 3.6784
Σ	5.9100 4.3331 3.1534 2.2790 1.6362	1. 1675 8. 2862 5. 8553 4. 1235 2. 8986	2.8367 1.4337 18.1388 7.2143 5.1793	3.7586 2.7629 2.8639 1.5668	9.5141 7.6196 5.2115 5.1525 4.3445 3.7218
Œ	5.8127 4.2668 3.1862 2.2474 1.6157	1. 1354 8. 2286 5. 8283 5. 8283 4. 1281 2. 9894	2.0543 1.4548 10.3378 7.3967 5.3378	3.8084 2.8786 2.1497 1.6342 1.2638	9.9228 7.9218 6.4443 5.3248 4.4608 3.9858
£	5.4482 4.8846 2.9258 2.1257 1.5354	1. 1835 7. 8916 5. 6236 3. 9948 2. 8321	2.0004 1.4235 7.2499 5.2237	3.7979 2.7916 2.9891 1.5728	9.4185 7.4745 6.8296 4.9465 4.1216 3.4884
u.	5. 1267 3. 7778 2. 7648 2. 6698 1. 4492	1.8376 7.3786 5.2138 3.6652 2.5665	1.7936 1.2537 8.7944 6.2028 4.4128	3.1726 2.3122 1.7128 1.2995	7.7237 6.1447 4.9771 4.1918 3.4362 2.9233
ר	4.9737 3.6788 2.6941 1.9642 1.4238	1.0243 7.3324 5.2226 3.7050 2.6220	1.8558 1.3136 9.3482 6.6789 4.8161	3.5005 2.5849 1.9324 1.1309	8.8612 7.6563 5.7692 4.6988 3.9115 3.3872
RLT (KN)	888 888 888 888 888 888 888 888 888 88	88 . 988 94 . 988 988 . 988 988 . 988 988 . 988	98.9 98 98.9 98 98.998 98.998 98.998 98.998	198.900 182.900 184.900 184.900 186.900	118.000 112.000 114.000 116.000 118.000 118.000

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'n ę ę 1 4.4967 3.3476 2.4827 1.8353 1.3526 . 1428 . 1032 . 3713 . 0664 . 0716 6625 5626 6969 6327 9183 3753 2659 4638 8887 4551 3003 Ο 0 ~ 0 0 0 --01 5782 5745 5746 5746 5067 3062 3062 9429 9952 9963 144 7948 6475 9475 9665 7288 7288 7288 72.00 70.000 9453 5887 2686 1746 3891 T 0 ~ n o o 4 0 N - -000---. ო ~ 10 8561 3262 8888 8332 10.0050 7.2281 5.1568 3.6723 2.6110 3.7178 2.7830 2.1135 1.6297 1.2760 0145 1886 7073 5705 5705 6887 00230 0 4 0 N - -. ന ŝ 9859 8531 8531 8531 8531 8531 8238 2845 5382 5382 7582 7582 7582 51-13 6328 5328 5366 2284 2284 9858 6342 6486 5763 7398 0731 ŝ 0 ~ n n n 2.9591 2.1425 1.5085 1.1983 .9169 5.7208 5.7208 4.7046 3.9056 3.4137 5.6265 4.1637 2.9757 2.1464 1.5266 8083 2084 2088 2088 2088 2088 2088 673 œ 10 4 2001 2415 2415 2408 3405 3405 9177 9177 6773 4386 2156 3655 3655 5466 5700 5327 5327 5327 5327 8666 4156 3448 5638 5638 5638 1900 7 -0000 0 * 0 N ю. ii t 9753 9753 9753 9723 9637 9644 5142 5142 5578 9916 2074 2074 2074 2074 6223 6667 6967 6367 2224 8453 7 0400-N 00 00 4 N 04 , . 8493 4218 9918 9697 9593 2849 8227 3798 4478 7558 4700 7896 1474 2745 2745 8615 3.5776 2.6238 1.9666 1.4922 1.1584 9165 3877 3877 8612 8569 2847 6838 Σ 00040 41--00 - - - - - - - n d . ო 9640 6785 2285 1237 2871 11.7616 8.4617 5.9697 4.2236 2.9786 5.7125 4.2235 3.1887 2.2662 1.6368 2.0900 1.4790 1.0465 7.4464 5.3418 3.8786 2.8482 2.1153 1.6665 1.2315 **61**13 æ ത് 2195 8647 8647 8463 8463 8463 8213 4.8068 2.9348 2.1759 1.6358 1.2475 9778 9778 9778 9778 9778 9778 6087 6055 6655 5227 9672 6275 1191 9952 1462 4004 T - ~ D + O ม่อ่งง่-ത് 10.1938 7.3377 5.2842 3.7777 3.7777 2.0911 9121 9657 9657 9659 9659 5836 6235 6235 4682 1259 8789 9892 6686 6687 9173 3417 400--10 T ο N 0040 Ø 9.9262 7.2601 5.2143 3.7592 2.7639 3.9196 2.9149 2.1954 1.6765 5721 5721 5166 5166 5555 95655 9429 9971 9678 9284 3281 8224 1728 6343 4685 5745 **9807** - 0 N io n , <u>.</u> 880 F <u>....</u> 22222 88288 88188 88388 8

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	٥	3.4966 2.6334 1.9736 1.4705 1.8989	8. 1437 6. 0000 4. 4 167 3. 2348 2. 35 12	1.7199 1.2519 9.1227 6.6643 4.8982	3.6186 2.6898 2.6898 1.5434 1.1931	9.3598 7.4579 6.8372 4.9668 4.1498 3.5176
Y PRESSURE (MB) AT 60 N	Z	3.8185 2.8659 2.1427 1.5958 1.1812	8.6992 6.3069 4.6331 3.3498 2.4879	1.7217 1.2265 8.7234 6.2862 4.4208	3. 1894 2. 3857 1. 6942 1. 2048 . 9619	7.4617 5.9189 4.7815 3.9408 3.3238 2.833
	0	4.3757 3.2458 2.3959 1.7581 1.2825	9.3022 6.7117 4.8201 3.4406 2.4628	1.7578 1.2549 8.9982 6.4865 4.7133	3.4365 2.5651 1.9294 1.4726 1.1412	8.9831 7.1827 5.8329 4.8983 4.6213 3.4892
	S	5.1010 3.7491 2.7371 1.9776 1.4152	18.8228 7.8352 4.9081 3.3947 2.3436	1.6187 1.1217 7.8389 5.5342 3.9643	2.8834 2.1367 1.6135 1.2438 .9756	7.8826 6.3475 5.2518 4.4877 3.7328 3.7328 3.2348
	œ	6.5462 4.7613 3.4295 2.4487 1.7332	12.1678 8.4724 5.8572 4.6217 2.7466	1.8088 1.2086 8.6293 5.8972 4.6646	2.8355 2.0110 1.4556 1.8782 1.8188	6.3763 5.9946 4.1718 3.4951 2.9999
	ר	7.2366 5.2641 3.7834 2.6843 1.8799	13.0020 8.8871 6.8128 4.6337 2.6902	1.7889 1.1988 7.9714 5.3847 3.6659	2.5648 1.8188 1.3198 .9787 .7434	5.7777 5.5727 3.7277 3. 9806 2. 6826 2. 5826
	ר	7.3875 5.4275 3.9297 2.7921 1.9461	13.2048 8.9171 5.8851 3.8464 2.4064	1.6166 1.6574 7.8234 4.7332 3.2073	2.3426 1.7885 1.2776 9785 .7655	6. 1876 4. 9543 4. 9982 3. 4816 2. 8072 2. 4378
	E	6.6188 4.8918 3.5751 2.5888 1.8395	12.9410 6.9024 6.1788 4.2848 2.9419	1.9139 1.2895 8.7354 5.9783 4.1353	2.9899 2.6875 1.5295 1.1453 .8705	6.8369 5.4641 5.4641 3.6746 3.6746 3.6894 2.6336
MONTH	Œ	5.7022 4.2342 3.1217 2.2031 1.6563	11.9158 8.5841 6.0245 4.2481 2.9082	2.0701 1.4416 10.0548 7.0378 4.9577	3.5238 2.5332 1.6478 1.3692 1.8321	7.9173 6.1849 4.9168 3.9782 3.2726 2.7352
	E	4.6428 3.4546 2.55933 1.8864 1.3831	18.8828 7.3862 5.2638 3.7665 2.6845	1.9825 1.3431 9.4667 6.6689 4.7889	3.3387 2.3863 1.7247 1.2648 .9416	7.141 5.5284 4.3671 3.5216 2.8979 2.4312 2.4312
	łد.	4. 8276 3. 8892 2. 2484 1. 5687 1. 2252	8.9919 6.5534 3.4388 2.4537 2.4537 2.4538	1. 7623 1. 2538 8. 9048 6. 32 17 4. 4904	3.2157 2.3198 1.6999 1.2516 .925	7.2346 5.6683 4.5355 3.7661 3.6982 2.6267
	7	3.5748 2.6884 1.5826 1.15326	8.2340 6.6475 4.4177 3.2118 2.3235	1.6734 1.2034 8.6746 6.2521 4.5253	3.2949 2.4211 1.8981 1.3568 1.8356	8.8088 6.3063 5.1532 4.2246 3.5225 2.9849
	RLT (KN)	88.58 88.58	888888 88888 888888	88.5.5.88 88	100.000 100.000 101.000 100.000 100.000 100.000	118.000 112.000 114.000 114.000 118.000 118.000

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MONTHLY PRESSURE (MB) AT 75 N

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٥	3.0417 2.3002 1.7435 1.3114 9.8113	7.2946 5.3918 3.9589 2.8999 2.8998	1.5191 1.6961 7.9184 5.7173 4.1515	3. 8335 2. 2383 1. 67 16 1. 266 1 9. 7377	7.6135 6.8534 4.8961 4.8261 3.3647 2.8549
Z	3.2534 2.4517 1.8435 1.3709 18.2788	7.6089 5.5972 4.0081 2.9053 2.1354	1.5282 1.6878 7.7231 5.4766 3.8931	2.7809 2.0050 1.4056 1.6088 8.2434	6.3726 5.8341 5.8341 4.8655 3.3541 2.8243 2.4233
0	3.7798 2.8004 2.8004 1.5288 11.1788	8. 1286 5. 8657 3. 8134 2. 1478	1.5282 1.6874 7.7644 5.5693 4.8261	2.9367 2.1684 1.6233 1.23348 9.5361	7.4826 5.9712 4.8440 3.9009 3.3384
S	5.2218 3.8358 2.7928 2.6074 14.2468	9.9759 6.9007 4.7194 3.2014 2.1593	1.4551 .9837 6.7124 4.6365 3.2582	2.3322 1.7862 1.2761 .9762 7.6269	6.0829 4.0416 4.0859 3.4301 2.9216 2.5188
Œ	7.0421 5.1071 3.6513 2.5612 2.5812 18.6300	12.4518 8.5188 5.7781 3.8744 2.5835	1.7171 1.1395 7.5826 5.6813 3.4424	2.3685 1.6615 1.1938 .8796 6.6662	5.1784 4.1362 3.3858 2.8383 2.4289 2.197
7	8.2652 6.8237 6.8237 4.2081 3.6088 20.6238	13.8498 9.1133 5.8878 3.7446 2.3553	1.4728 .9223 5.8272 3.7376 2.4483	1. 6442 1. 1358 . 8886 . 5929 4. 4746	3.4678 2.7351 2.7351 1.8559 1.3006 1.3448
7	8.2296 6.8488 4.3441 3.8437 20.7268	13.6810 8.7572 5.4465 3.3113 1.9847	1. 1869 . 7168 4. 4285 2. 8185 1. 8593	1.2716 .9023 .6622 .5613 3.8959	3.0000 2.5127 2.6714 1.7300 1.4008
£	6.9657 5.1411 3.7424 2.6816 18.8966	13.8868 8.9894 5.9783 3.9472 2.5833	1.6814 1.6945 7.1731 4.7366 3.2993	2.2000 1.5500 1.1279 .8374 0.3728	4.9655 3.9543 3.2127 2.6577 2.2346
Œ	5.3474 3.9766 2.9367 2.1474 15.5498	11.1388 7.8992 5.5455 3.8598 2.659	1.8319 1.2552 8.6131 5.9314 4.1178	2.8864 2.8524 1.4835 1.8928 8.1928	6.2681 4.8889 3.8866 3.1457 2.5981 2.5981 2.1655
E	4.1544 3.1656 2.3118 1.7675 12.5106	9.8899 6.3444 3.2984 2.3186	1.6070 1.1108 7.6596 5.2768 3.6489	2.5366 1.7831 1.2714 .9225 6.8249	5.1563 3.9642 3.1469 2.5396 2.5396 1.7562
íL.	3.2344 2.4349 1.8282 1.3644 1.3644 19.1210	7.4512 5.4454 3.9461 2.8373 2.8373	1, 4327 1, 6682 7, 6748 4, 9556 3, 4798	2.4547 1.7488 1.2625 .9269 0.9365	5.3028 4.1436 3.3119 2.7049 2.2562
7	2.9689 2.2537 1.7868 1.2842 9.0011	7.1197 5.2398 3.8233 2.7698 1.9899	1.4219 1.8112 7.1854 5.1854 3.0465	2.6288 1.9833 1.4018 1.8488 7.9749	6.1834 4.8838 3.9328 3.2226 2.6885 2.2789
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5	88288	****	88288	<u> </u>	1128 1128
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٥	9.1961 6.9458 5.1199 3.7466 2.7211	1.9639 1.4876 1.6835 7.1123 7.1123 5.6171	3.5205 2.4500 1.7078 1.1851 8.2119 5.6899 3.9359 3.9359 1.9084 1.3421	9.5136 6.8129 6.8129 4.9317 3.6155 2.6833 2.6833 2.6833	
Z	9.2664 5.9331 3.6965 3.6965 2.6786	1.9334 1.3994 1.3994 7.1313 7.1313 5.6853	3.6127 2.5527 1.7976 1.2615 8.8216 6.1505 4.2065 2.9568 2.9568 1.4348	10.0010 7.1131 5.8775 3.0605 2.6857 1.9942	
0	9, 3062 6, 9569 5, 0757 3, 6817 2, 6553		3.5661 2.4798 1.2364 8.7187 6.1463 5.1463 2.1535 1.5298 1.5298	10.9288 7.8653 5.7897 4.1891 3.1867 2.3327	
S	9.1881 5.8338 5.6832 3.6371 2.6278	1.8895 1.3522 .9643 6.8528 4.8527	3.4243 2.4838 1.6838 1.1751 8.1874 5.6947 3.9479 2.7419 1.3373	9.4235 6.7869 4.8136 3.4998 2.5763 1.9231	
œ	9.0200 6.6877 4.8922 3.5528 2.5625	1.8395 1.3146 .9368 6.6594 4.7272	3.3515 2.3786 1.6783 1.6783 1.6783 8.4466 8.117 4.2885 3.8629 3.8629 1.5948	11.6248 8.5373 6.3238 6.3238 4.7272 3.3679 3.3679 2.7287	
ר	8.8666 6.5468 3.5298 2.5298 2.5056	1.8528 1.3308 0492 6.7398	3.3432 2.3327 1.1243 7.7771 7.7771 3.7019 3.7019 1.7764 1.7764	8.7442 6.2187 6.2187 3.2585 2.4865 2.4863	
Ъ	9.0000 6.7321 4.9326 3.5877 2.5927	1.8637 1.3327 9493 6.7361 4.7648	3.3597 2.3583 1.1572 8.6075 3.9613 3.9613 1.9578 1.3881	9.9118 7.1398 5.1924 3.8191 2.8415 2.1465	
٤	9.3743 7.8169 5.1112 3.7819 2.6558	1.9127 1.3678 .9749 6.9338 4.⊴225	3.4844 2.4564 1.7276 1.2126 8.4901 5.9337 2.8766 2.8766 1.4949	9.8837 7.8851 5.8976 3.6283 2.6478 1.9638	
Œ	9.5888 7.2262 5.2874 3.8378 2.7647	1.9868 1.4112 1.0028 7.0882 5.0982	3.5232 2.4728 1.7345 1.2169 6.6025 4.2151 2.1659 1.5614	16.7738 7.7951 5.6692 4.1962 3.1268 2.3574	
Σ	9.5085 7.2097 5.3138 3.8885 2.8247 2.8247	2.0405 1.4645 1.8461 7.4318 5.2308	3.6972 2.5949 1.2560 8.6563 5.9877 4.1297 1.9853 1.9853	9.700 6.9813 5.9318 3.6703 2.7214 2.7214	
۱Ľ.	9.8895 6.9272 5.1754 3.8362 2.8183	2.0519 1.4786 1.8552 7.4515 5.2166	3.6866 2.4698 1.6792 1.1354 7.6498 5.1486 3.4614 1.5949 1.9986	7.6595 5.4162 3.8875 2.8356 2.1012 1.5823	
Ъ	8.9995 5.1435 3.8284 2.8141	2.0561 1.4881 1.6677 7.5880 7.5880 5.3461	3.7319 2.5881 1.7728 1.2118 8.2487 5.6128 3.8128 3.8128 1.7887 1.2414	8.7003 6.1906 4.4008 3.2038 2.4233 1.8274	
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MONTHLY DENSITY (KG/M°3) AT 15 N

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0	9.8429 6.8189 5.8398 3.6999 2.6983	1.9571 1.4111 1.8128 7.2342 5.1474	3.6472 2.5718 1.8682 1.2698 8.9091 8.9091 5.2534 4.3833 3.6853 3.6853 3.6853 1.5541	11.1530 8.0795 5.9124 4.3779 3.2793 3.2793 2.4878
Z	9, 6918 5, 7957 5, 6655 3, 6648 2, 6665	1.9343 1.3956 1.8932 7.1829 5.1239	3.6426 2.5742 1.8141 1.2753 8.9467 8.9467 8.2687 3.6691 2.1582 1.5278	10.8818 7.8278 5.6891 4.1863 3.1198 2.3333
0	9.1751 6.7796 4.9856 3.6388 2.6391	1.9034 1.3058 1.3058 0.9762 4.9306	3.4965 2.4665 1.7392 1.2257 8.6474 6.1874 5.1874 5.1874 1.3141 3.6613 2.1848 1.5786	11.3738 8.3189 6.1356 4.5777 3.4544 2.6369
S	9.8845 6.6434 4.8942 3.5778 2.5971	1.8748 1.3445 .9598 6.8211 4.8267	3.4838 2.3865 1.6789 1.1689 8.1649 8.1649 3.9921 3.9921 2.8833 1.9888 1.4888	10.1120 7.3278 5.3696 3.9813 2.9888 2.2888
œ	9.6005 6.6286 6.6286 3.5681 3.5681 2.5889	1.8598 1.3314 0489 6.7366 4.7366	3.3644 2.3659 1.1715 8.2645 8.2645 2.1208 2.1208 1.5345	11.2060 8.2000 6.1529 4.6322 3.5261 2.7138
ר	8.8654 6.5172 4.8217 3.5391 2.5798	1.8006 1.3414 .9576 5.7955 4.7935	3.3649 2.3466 1.6326 1.1326 7.8662 3.7892 1.3267 3.7892 1.8586 1.3168	9.4214 6.8105 4.9832 3.6912 2.7702 2.1056
J	9.2672 6.8251 5.8163 3.6594 2.6531	1.9127 1.3728 .9888 6.9753 4.9477	3.4998 2.4046 1.7339 1.7339 8.5693 8.5693 8.5693 8.5693 2.1254 2.1254 1.5187	10.9370 7.9445 5.8297 4.3245 3.2452 3.2452 2.4640
Σ	9.4899 7.0008 5.1845 3.7797 2.7385	1.9748 1.4168 1.8119 7.2856 5.1144	3.5186 2.5482 1.7913 1.2573 8.8179 8.8179 6.1827 4.3281 3.6399 2.1443 1.5221	18.8848 7.8522 5.7231 4.2197 3.1488 2.3797
Œ	9.3815 7.6398 5.1798 3.7881 2.7376	1.9702 1.4098 1.8032 7.1103 5.6239	3.5389 2.4836 1.7419 6.6438 6.6438 3.0163 2.1562 1.5458	11. 1938 8. 1871 6. 6486 4. 5194 3. 4148 2. 6696
£	9.2611 6.9791 5.1586 3.7845 2.7364	1.9947 1.4336 1.8246 7.2811 5.1481	3.6215 2.5324 1.7056 8.5468 8.5468 8.5468 2.6286 1.4339	16.2478 7.4625 5.4698 4.6656 3.6636 3.6636
L.	9.0004 6.7976 5.0824 3.7720	2.0202 1.4571 1.6408 7.3596 5.1543	3.5766 2.4573 1.6791 1.6791 7.7617 7.7617 3.5694 3.5694 1.1898 1.1898	8.4285 6.8467 4.4862 3.2685 2.4491 1.8678
ר	8.9383 6.7669 5.6598 3.7588 2.7588	2.0240 1.4636 1.6521 7.4859 5.2826	3.6982 2.5668 1.7732 1.2280 8.3964 8.3964 1.9253 1.9253 1.3586	9.6958 7.0070 5.1319 3.8122 2.8714 2.1935
RLT (KH)	88.5 88.6 88.6 88.6 88.6 8 8 8 8 8 8 8 8 8 8	80.000 80.0000 80.00000 80.00000 80.0000 80.0000 80.0000 80.0000 80.0000 80.0000 80.00000 80.00000 80.00000 80.00000 80.00000000	86.28 86.29 86	118.800 112.800 114.800 116.800 116.800

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٥	8.8449 6.8867 4.5197 3.3393 2.4549	1.7977 1.3187 9.5265 6.9994 4.9851	3.5985 2.5755 1.8447 1.3283 9.4467	6.7635 4.8365 3.4724 2.5837 1.8169	1.3275 9.7816 7.2725 5.4644 4.1483 3.1836
Z	8. 1780 6. 1495 4. 5889 3. 3916 2. 4959	1.8257 1.3278 9.5982 6.8999 4.9258	3.5822 2.4749 1.7442 1.2278 8.6302	6.8762 4.2793 3.8312 2.1624 1.5565	1. 1317 8. 3197 6. 1899 4. 0645 3. 3396
0	8.3683 6.2657 4.6661 3.4472 2.5265	1.8371 1.3254 9.4962 6.7624 4.7998	3.3887 2.3758 1.6686 1.1736 8.2828	5.8788 4.1776 2.9983 2.1727 1.5912	1. 1784 8. 8266 5. 6987 5. 1338 3. 9846 3. 1293
S	8.7369 6.5842 4.8478 3.5825 2.6262	1.9083 1.3747 9.8239 6.9782 4.9144	3.4485 2.4869 1.6795 1.1728 8.2281	5.7882 4.6946 2.9232 2.1603	1. 1373 8.5047 6.4423 4.9417 3.8387 3.8178
œ	8.8642 6.5464 4.8648 3.5637 2.6194	1.8979 1.3638 9.7195 6.8778 4.8332	3.3785 2.3462 1.6272 1.1279 7.8305	5.4629 3.8224 2.6969 1.9226	1.0116 7.4758 5.0004 4.2589 3.2785 3.2785
ר	9.3451 6.9216 5.1533 3.8827 2.7822	2.8157 1.4486 18.3168 7.2982 5.1155	3.5699 2.4758 1.7152 1.1881 8.2334	5.7563 4.6316 2.8478 2.8478 1.4698	1.0732 7.9438 5.9588 4.5281 3.4866 2.7183
ر	9.5100 7.1164 5.2008 3.8861 2.8373	2.8562 1.4795 10.5818 7.5254 5.3282	3.7586 2.6394 1.8527 1.3010 9.1300	6.4653 4.5757 3.2616 2.3439	1.2459 9.2272 5.2379 5.2379 4.0161 3.1156
ε	9.4404 7.1696 5.2944 3.9101 2.8635	2.8982 1.4985 18.7168 7.6899 5.3699	3.7693 2.6384 1.8325 1.2768 8.9888	6.2287 4.3784 3.6911 2.2664 1.5921	1.1621 8.5659 6.4256 6.4256 4.8716 3.7414 2.9103
E	9.2295 6.9987 5.2024 3.8412 2.8138	2.0444 1.4740 10.5580 7.5121 5.3174	3.7468 2.6259 1.8384 1.2875 9.6347	6.3623 4.4916 3.1962 2.2942 1.0648	1.2199 9.6488 5.7912 5.1618 3.9788 3.9788
E	8.5465 6.4988 4.8418 3.5839 2.6339	1.9231 1.3939 10.0428 7.1889 5.1289	3.6283 2.5573 1.7989 1.2647 8.8985	6.2767 4.4328 3.1521 2.2583 1.6334	1 1932 8.8122 5.3825 4.9783 3.8695 3.8695
LL.	8. 1048 5. 0848 4. 5593 3. 3924 2. 5659	1.8348 1.3386 9.5665 6.8174 4.8176	3.3787 2.3563 1.6296 1.1266	5.4001 3.7622 2.6382 1.8682 1.3384	
ר	7.8321 5.8786 4.4812 3.2777 2.4238	1.7785 1.2943 9.3456 6.6979 4.7671	3.3728 2.3784 1.6626 1.1649 8.1761	5. 7556 4. 06 18 2. 8993 2. 8742 1. 5648	1. 1942 8. 1982 6. 1635 4. 6925 3. 6179 2. 8239
ALT (KH)	888 888 888 888 888 888 888 888 888 88	88 . 996 84 . 996 86 . 986 86 . 986 86 . 986	88.88 88.98 88.98 88.98 88.98 88 88 88 88 88 88 88 88 88 88 88 88 8	196 . 996 194 . 996 194 . 996 196 . 996	118.000 112.000 114.000 116.000 118.000 118.000

MONTHLY DENSITY (KG/M°3) AT 30 N

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MONTHLY DENSITY (KG/M"3) AT 45 N

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٥	6.8681 5.1884 3.8931 2.9183 2.1674	1. 6885 1. 1898 8. 7589 6. 4389 4. 7848	3.4313 2.4987 1.8948 1.8948 1.3968 1.3968 9.4476 6.8368 4.9488	3.5842 2.6128 1.9151 1.4148 10.5248 7.9933 5.9947 4.5536 3.5578
Z	7.1452 5.3962 4.8059 3.8499 2.2764	1.6805 1.2458 9.1274 6.6426 4.8014	3.4475 2.4555 1.7414 1.2385 8.6888 8.6888	3.6501 2.1832 1.5722 1.5722 8.4282 6.2757 6.2757 3.6262 3.6262 3.6262
0	7.6526 5.7417 4.3111 3.2131 2.3767	1.7433 1.2681 9.1496 6.5549 4.6658	3.3049 2.32849 1.6593 1.6593 1.1551 8.1668 5.8088 4.1385	2.9705 2.1674 1.5944 8.9419 6.8216 4.1186 3.2582 4.1186
S	8.4824 6.2938 4.7276 3.5198 2.5965	1.6943 1.3673 9.7686 6.9831 4.8391	3.3717 2.3343 1.6161 1.1205 7.8128 3.8778 3.8778	2.7745 2.0121 1.4793 8.3478 6.4077 6.4077 3.9328 3.9328
Œ	9. 1423 6. 7063 5. 8536 3. 7249 2. 7222	1.9715 1.4147 10.0008 7.0008 7.0008	3.4378 2.3629 1.6188 1.1048 7.5442 5.1613 3.5483	2.4485 1.7118 1.2115 6.3316 6.3316 3.5157 2.6883 2.6883 2.6883
ר	9.8592 7.3668 5.4968 4.8519 2.9589	2. 1367 1. 5238 10. 7818 7. 5453 5. 2353	3.6875 2.4683 1.6859 1.1516 7.8925 5.4358 3.7639	2.6337 1.8053 1.3386 7.1918 5.3878 3.1548 3.1548 3.1548
ר	9.8282 7.7757 5.8567 4.3638 3.2854	2.3250 1.6014 11.7280 8.1751 5.6474	3.8696 2.6376 1.7962 1.2277 8.4377 5.854 1.8926	2.9829 2.8873 2.8873 1.5233 6.5886 6.5886 3.9787 3.9787
Σ	9.7826 7.4369 5.6867 4.1856 3.6914	2.2573 1.6282 11.6899 8.1844 5.7122	3.9529 2.7125 1.8561 1.2692 8.7835 5.9992 5.9992 4.1578	2.9138 2.0080 1.4885 1.4885 8.0580 6.0035 4.6312 3.5882 3.5882 3.5882 3.5882
œ	8.8889 6.7285 5.6577 3.7732 3.7732 2.7025	2.0495 1.4908 10.7540 7.6045 5.4652	3.8563 2.7816 1.8875 1.3178 9.1976 6.4483 4.5202	3.1971 2.2821 2.2821 6.646 6.6441 3.8693 3.8693 3.8693
£	8.8918 6.1136 4.5788 3.4128 2.5287	1.8639 1.3635 9.9484 7.2064 5.1928	3.7211 2.6497 1.8818 1.3323 9.4279 6.6733 4.7286	3.3051 2.3085 1.7281 9.2191 5.1369 3.9832 3.9832 3.9832
L.	7.4169 5.6622 4.2839 3.1483 2.3333	1.7238 1.2049 9.2227 6.6789 4.9049	3.4342 2.4358 1.7206 1.2116 8.5212 5.9938 4.2149	2. 1215 2. 1215 1. 5246 8. 1171 8. 1171 3. 4511 2. 5394 3. 4511
ר	7.8288 5.2846 3.9787 2.9819 2.2335	1.0484 1.2143 8.8898 6.4691 4.6799	3.3688 2.4883 1.7181 1.2239 8.7226 8.7226 4.4584	3.2001 2.3184 1.6946 9.3297 9.3297 7.6377 5.3787 4.1478
ALT (KN)	800 800 800 800 800 800 800 800 800 800	88.82.92.88 86.82 86.82 86.82 86.83	8.52 8.62 8.62 8.62 8.62 8.62 8.62 8.62 8.6	164.900 166.906 116.906 114.906 114.906 114.906 118.906 118.906 118.906 118.906 118.906

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٥	5.1192 3.9163 2.9778 2.2579 1.7057	1.2833 9.6651 7.1537 5.2965 3.9836	2.8085 2.8816 1.5894 1.5894 7.8739	5.6814 4.6922 2.9386 2.1495 1.5723	1.1506 8.6287 6.4758 4.9177 3.7755 3.7755 2.9318
Z	5.6696 4.3106 3.2704 2.4755 1.8675	1.4029 10.4790 7.7793 5.7351 4.1958	3.8447 2.1878 1.5597 1.1949 7.7862	5.4003 3.8220 2.6702 1.8862 1.3378	. 9578 6. 9348 5. 8945 3. 7885 2. 8588
0	6.7653 5.6852 3.8275 2.8643 2.1302	1.5731 11.5310 8.3999 6.8622 4.3517	3.1061 2.2025 1.5594 1.1815 7.7955	5.5312 3.9333 2.8174 2.8174 2.6357 1.4835	1.0038 8.1732 6.1725 6.1725 4.7175 3.6497 2.8574
S	8. 1589 6. 8535 4. 5085 3. 48 18 2. 5188	1.8202 13.1618 9.3445 6.5568 4.5486	3.1202 2.1345 1.4541 1.4541 0.0011	4.6864 3.2612 2.2988 1.6452 1.1953	
œ	10.6540 8.8231 5.9426 4.3374 3.1628	2.2738 16.1838 11.4118 7.9781 5.5161	3.7838 2.5702 1.7356 1.1667 7.8206	5.2528 3.5268 2.3892 1.6343 1.1328	
J	11.7020 8.8784 6.6389 4.8951 3.5562	2.5442 17.9158 12.4268 8.4944 5.7317	3.8248 2.5201 1.6614 1.6913 7.1998	4. 7643 3. 1707 2. 1586 1. 4768 1. 8313	. 7328 5.3883 3.9828 2.9238 2.2268 1.7223
J	11.6838 8.8875 6.8967 5.1139 3.7683	2.7186 19.8008 12.9838 8.6585 5.6537	3.6372 2.3158 1.4741 .9443 6.1318	4.0469 2.7204 1.8704 1.3160 .9486	
Σ	18.2678 7.8342 5.9528 4.4784 3.3128	2.4189 17.3800 12.2860 8.5499 5.8622	3.9686 2.6547 1.7665 1.1731 1.1731 7.8115	5.2325 3.5318 2.4151 1.6778 1.1853	.8523 6.2356 4.6435 3.5178 2.7879 2.7879 2.1172
Œ	8.7448 6.6376 5.6375 3.7725 3.7725 2.8101	2.0739 15.1540 10.9640 7.8547 5.5737	3.9281 2.7389 1.8927 1.3069 9.8147	6.2234 4.3007 2.9921 2.6007 1.4901	1.8085 7.7658 5.7194 4.2728 3.2353 2.4842
E	7.1200 5.3500 4.6327 3.0262 2.2006	1.6795 12.4956 9.1845 5.6381 4.8052	3.4531 2.4582 1.7304 1.2231 8.5031	5.9722 4.1488 2.8863 2.8156	1.6031 5.1756 5.1756 3.8163 2.6335 2.6335 2.6335
L.	6.0837 4.5928 3.4691 2.6125 1.9065	1.4658 10.8918 8.6515 5.9154 4.3167	3.1272 2.2443 1.6683 1.1339 7.9933	5.6117 3.9221 2.7446 1.9273 1.3016	
ר	5.2763 4.8127 3.8526 2.3155 1.7497	1.3158 9.8382 7.3111 5.3078 3.9578	2.8822 2.6966 1.4946 1.6065 7.6185	5.4221 3.8512 2.7444 1.9059 1.4182	1.8321 7.5856 5.6373 4.2482 3.2298 2.4899
LT (KM)	888 888 888 888 888 888 888 888 888 88	80. 9 00 80. 9 00 80. 9 00 80. 9 00 80. 9 00	88888	900 900 900 900 900 900 900 900 900 900	119.000 112.000 114.000 116.000 116.000 118.000
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MONTHLY DENSITY (KG/M°3) AT 50 N

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MONTHLY DENSITY (KG/M"3) AT 75 N

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9	4.3783 3.3313 2.5069 1.9632 1.4995	1.1301 8.5982 6.4432 4.7928 3.5344	2.5864 1.8741 1.3586 1.3588 9.6721 5.9889 4.9259 3.5862 2.5847	1.7999 4.3035 9.5287 7.8379 5.2582 3.9796 3.9796 3.9493 3.8493
z	4. 7011 3. 6153 2. 7366 2. 9987 1. 5948	1.2057 9.0717 6.7812 5.0321 3.7021	2.6992 1.9458 1.9458 9.8467 6.9386 6.9386 3.3788 2.3557	1.6493 1.1626 8.2759 5.9579 5.9579 3.2184 2.4185
0	5.8318 4.3593 3.2895 2.4694 1.8431	1.3666 10.8478 7.3311 5.3893 3.8154	2.7239 1.9294 1.3623 9.5967 6.7642 3.3778 3.3778 2.4056	1.7286 1.2544 5.2544 5.1411 3.9148 3.9148 3.9260 2.3584
v	8.3338 6.2287 4.7049 3.5181 2.5970	1.8872 13.4900 9.4759 6.5503 4.4591	2.9998 1.9998 8.7954 5.8775 3.9631 2.7018 1.8714	1.3199 .9481 5.9489 5.1722 3.9255 3.8293 2.3769 1.8927
Œ	11.4158 8.7514 6.4865 4.7446 3.4228	2.4376 17.1200 11.8760 8.1325 5.3005	3.6664 2.4417 1.6667 1.6667 6.9631 6.9631 2.9926 1.9968	1.3590 .9290 6.4886 4.528 3.3543 3.3543 1.8664 1.8664
7	13.3236 16.2856 7.7645 5.7517 4.1689	2.9500 28.3500 13.6928 8.9881 5.7581	3.6223 2.2439 1.3817 8.5103 5.2801 5.2801 3.3179 2.1167 1.3804	. 9216 . 6306 4. 4199 3. 1711 2. 3262 1. 7429 1. 3314
7	12.9150 18.6018 7.7381 5.8504 4.3171	3.6756 21.6776 13.8656 8.7635 8.7635 5.3447	3.1755 1.8566 1.8566 1.8856 3.8811 3.8811 2.4116 1.5429	. 6974 . 4915 3. 5683 2. 6593 1. 5893 1. 5893 1. 2782 1. 6357
Σ	10.7860 8.2971 6.3362 4.7799 3.5331	2.5567 18.2578 12.7869 8.6532 5.7745	3.7878 2.4459 1.5689 18.8448 5.4686 5.4686 5.4586 1.934 1.8411	1.2359 .8751 .8751 6.2298 4.5277 3.3588 3.3588 2.5483 1.9567 1.5338
Œ	8. 1964 6. 1822 4. 7884 3. 5555 2. 6599	1.9675 14.3898 18.3758 7.3028 7.3028 5.1995	3.5153 2.4838 11.4918 7.7868 5.2814 3.5984 2.4688	1.7852 1.1965 5.1565 6.1569 4.5132 3.3633 2.5468 2.5468
Σ	6.2895 4.7195 3.5981 2.7281 2.8568	1.5368 11.3868 8.3497 6.8597 4.3463	3.0827 2.1568 1.4552 16.2748 7.8190 4.7747 3.2389 2.2842	1.5183 1.9444 5.3191 5.1846 3.7344 2.7321 2.8319 1.5351
L.	4.7930 3.6899 2.7548 2.1883 1.5951	1.2043 9.0313 6.7171 4.9534 3.6158	2.6136 1.8646 1.3178 9.2297 9.2297 9.419 4.419 3.8586 2.1888	1.4611 1.0199 5.1364 3.7282 2.7348 2.6425 1.5488
ר	4.2828 3.2312 2.4972 1.9235	1.1243 8.5612 6.3666 4.7265 3.4696	2.5211 1.2874 9.6915 6.3921 4.4789 3.1324 2.1985	1.5537 1.1074 7.9824 5.8218 4.3054 2.4554 2.4554
ALT (KH)	800 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 100 10 1	88 99 99 99 99 99 99 99 99 99 99 99 99 9	96.99 96 96 96 96 96 96 96 96	106.000 108.000 110.000 112.000 114.000 116.000 118.000

