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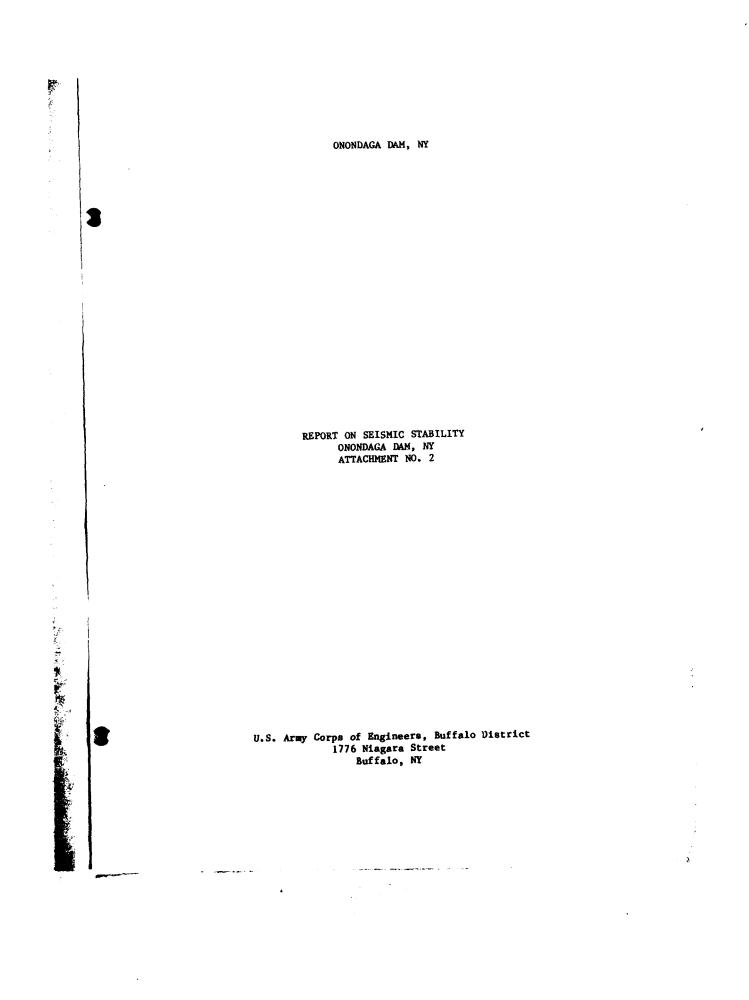
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about Onondaga Dam. Appendix B contains the rationale for the selected soil parameters used in the analysis. Appendix C contains hand calculations of the critical slip surfaces. Appendix D contains the computer input files and samply output for the slip circle analysis. Appendix E contains the wave runup analys for the storage pool. Appendix F contains the slope protection calculations. Attachment No. 1 is the Concrete Spillway Stability analysis and Attachment No. 2 is the seismic evaluation.

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PREFACE

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This study was conducted by H & A of New York and was authorized by the Department of the Army, Buffalo District, Corps of Engineers under Contract No. DACW 49-81-D-0011, Work Orders No. 2 and No. 3, dated October 13, 1981 and June 15, 1982, respectively.

The work was accomplished and the report written by Thomas X. Grasso, Staff Geologist, with assistance on seismologic evaluations provided by Dr. Edward F. Chiburis, Consultant, and Dr. Edward B. Kinner of Haley & Aldrich, Inc. All work was performed under the general technical supervision of Joseph J. Rixner, Manager of H & A of New York.

This work order was performed under the technical administration of the Buffalo District, Corps of Engineers, Geotechnical Section.

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GEOLOGICAL AND SEISMOLOGICAL INVESTIGATIONS AT ONONDAGA DAM, NEW YORK

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PART I: INTRODUCTION

Purpose and Scope

1. The purpose of this study was to investigate and evaluate the potential earthquake hazard at Onondaga Dam located in Central New York. This study was conducted in compliance with ER1110-2-1806 "Earthquake Design and Analysis for Corps of Engineers Dams" (by the Office of the Chief of Engineers, dated April 1977).

2. The study included geological, seismological, and subsurface investigations, and geotechnical engineering analyses. The study consisted of the following elements:

- a. an evaluation of the regional and local geology,
- b. performance of subsurface explorations and laboratory testing to further define the nature and density of soil deposits,
- <u>c</u>. an evaluation of the regional and local tectonic history with respect to structural deformation including faults,
- d. a review of historical regional seismicity,
- e. the determination of the maximum earthquake that will effect the site,
- f. an assessment of earthquake effects on Onondaga Dam, including an evaluation of liquefaction potential of the subsurface soils.

General Description of Onondaga Dam

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3. Onondaga Dam is located on Onondaga Creek, a tributary to the Seneca River in Onondaga County, New York (Figure 1). It is about four miles southwest of the southern limits of the City of Syracuse, New York and about 3,000 feet south of N.Y. Route 8°.

4. Onondaga Dam is a flood control structure consisting of a rolled earth fill embankment with a concrete overflow side channel spillway on the right (east) abutment. The overall length of the rolled fill embankment is 1,780 feet and the maximum height is 67 feet. An 1,100 ft. long spillway channel has been cut into rock along the right abutment. Rock cuts, ranging from about 15 to 40 ft. high, are exposed along the spillway channel.

5. The dam was completed in August 1949. It was designed and constructed by the Syracuse District of the U.S. Army Corps of Engineers.

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PART II: GEOLOGY

Physiography

6. Figure 2 shows the physiographic provinces of New York State. Onondaga Dam is located in the northern part of the Allegheny (Appalachian) Plateau Physiographic Province. The physiographic province immediately north of the site is the Lake Plain Physiographic Province.

Lake Plain Province

7. Bordering Lakes Erie and Ontario is a region of low relief known as the Lake Plain Province. This plain is underlain by sedimentary rocks of Late Ordovician to Devonian age that dip gently southward at 30 to 50 feet per mile. Refer to Table I for a geologic time scale. This province is bordered on the South by the Appalachian Plateau with a more or less abrupt boundary.

8. Three escarpments or "cuestas" strike east-west across the Lake Ontario portion of the plain, especially west of Rochester to the Niagara Peninsula. They are, from north to south:

> <u>a</u>. Medina Escarpment - Formed in the resistant Grimsby and Thorold Sandstones. This ridge is of minor topographic relief being only some 30 to 50 feet high. Wherever streams cross this feature, waterfalls result such as those at Medina, Albion and Holley. The escarpment cannot be recognized east of the Genesee River. The line of the Barge Canal rests atop this escarpment from Spencerport to Lockport.

> b. Lockport or Niagara Escarpment - This ridge is the most prominent bedrock controlled topographic feature on the Lake Ontario Plain. The Lockport Dolostone forms the caprock of this "mountain ridge", the crest of which lies about 300 feet above Lake Ontario. Niagara Falls,

the Upper Falls of the Genesee at Rochester, and many other minor falls in between are also formed on the Lockport Dolostone.

<u>c</u>. Onondaga Escarpment - The southernmost escarpment is about 50 feet in height, trending eastward from Buffalo through Batavia to Auburn. This escarpment is formed of resistant Onondaga Limestone. A line of abandoned and existing quarries marks its trace across the state, along with relatively small waterfalls on the numerous streams that cross the escarpment.

9. The remainder of the Lake Plain is underlain predominantly by shales, with a few relatively thin sandstone and carbonate interbeds forming the prominent east-west escarpments described above.

10. In general, the rocks are structurally simple, dipping gently southward and forming the north-northwest flank of the Appalachian (Allegheny) Synclinorium or Basin. This singly-dipping, homoclinal structure is interrupted locally by a few faults and a few minor low, open folds. The most significant of these is the Clarendon-Linden Fault, which strikes southwestward from Lake Ontario across northwest Monroe County and southeast Orleans County, through Batavia in eastern Genesee County and to Attica in central Wyoming County. It is a nearly vertical fault system, downthrown to the west, cutting the Precambrian basement, and having a displacement ranging from about 30 to 90 meters (100 to 275 ft.), (Hutchinson, et al. 1979).

11. A local fault structure of perhaps significant proportions occurs in northern Cayuga County but mainly in the subsurface, although one or more of these faults are thought to have a surface expression. Not much is known about this structure at present. A new gas field, the Blue Tail Rooster, is producing from traps in this structure. A well-developed, nearly vertical joint system of two sets, one essentially north-south and the other east-west is the only major, regional, geologic structure. Locally, a third set of joints may be present in places.

12. The Paleozoic rocks of the Lake Plain are mantled by a substantial thickness of glacial deposits. The most dramatic of these is the drumlin field of glacial till lying between Rochester and Syracuse. Several morainal deposits and glacial lake shoreline deposits, such as the offshore bar of Lake Iroquois are also present. Ridge Road (Rt. 104) runs along this bar.

Allegheney (Appalachian) Plateau

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13. The Allegheny Plateau covers nearly a third of the entire State. as shown on Figure 2, and is its largest physiographic province. It is a region of moderate to high relief with summits at Elevations 2000 to 4000 feet above sea level. The Catskill Mountains make up the eastern portion of the plateau while the remainder is often referred to as the Southern Tier. The entire region is underlain by Devonian shales, siltstones, and sandstones. However, the tops of the hills, south of Olean, expose sandstones and conglomerates of Mississippian and Pennsylvanian age (Figure 2 and Table 1). The eastern border of this province is the Catskill Front, an escarpment composed of Middle Devonian, massive, red sandstone of the Catskill Delta. The entire Catskill region is held up by these coarse deltaic sandstones. The Helderberg Escarpment of Lower Devonian Limestones stretches from the Albany area westward to Syracuse forming the northern and eastern boundary. West of Syracuse, in the Finger Lakes and Genesee region, the boundary with the Lake Plain province is not abrupt, but more gradual, occurring in the northern portion of the Finger Lakes across the Middle Devonian Hamilton Group. In western New York and southeast of Lake Erie, the Portage Escarpment of Upper Devonian siltstones and sandstones forms the western border.

14. The strata of the Allegheney Plateau dip gently southward at about 50 to 80 feet per mile. In south central New York, gentle warping of the Devonian rocks occurred and produced several gentle open folds along with minor faulting. Deep wells and geophysical seismic surveys, penetrating through the Upper Silurian in this region, do not reveal

the large surface anticlines at depth. This suggests that the Upper Silurian salt beds may have "absorbed" the deformation. Aside from the Clarendon-Linden structure, no major faults are known in the Allegheny Plateau. A few small reverse faults of minor displacement are known in central New York, one in particular being the Portland Point Fault north of Ithaca along Cayuga Lake. Several thin, ultrabasic, peridotite dikes are known in central New York near Ithaca and Syracuse. These dikes have migrated upward along joint planes, are of Mesozoic age, and are of limited extent. The joint system is extremely regular, especially in Central New York, and comprises a north-south set and an east-west set. A third set is well developed only in the southwestern part of the province. Most of the streams in the bedrock gorges are controlled by the north-south and east-west joints.

15. Glacial deposits are relatively thin on the upland portions of this province, but the north-south valleys are normally deep U-shaped glacial troughs that are partially filled by thick moraine, lacustrine, and glacial outwash deposits. Some of the glacial troughs are plugged at both ends, thereby forming Finger Lakes, while other valleys have subsequently drained. A major east-west trending divide in drainage between the north flowing St. Lawrence and the south flowing Susquehanna systems occurs along the Valley Heads Moraine, which loops across the province at the southern extremities of the Finger Lakes and neighboring valleys.

16. In the Onondaga (Tully) Valley, the Tully Moraine, south of the site and west of the Village of Tully, is the local development of the Valley Heads Moraine. Onondaga Creek flows north off the moraine toward Lake Ontario and the Tioughnioga River flows south through Cortland to the Susquehanna River.

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Drainage

17. The major drainage system through the site is Onondaga Creek, consisting of two branches; the west branch rising in the vicinity of Cedarville Village and the main branch rising at the Tully Moraine. Both branches join just south of the Onondaga Dam (Figures 3 and 4) and flow north past the dam into Onondaga Lake, a tributary of the Seneca River and part of the St. Lawrence System.

18. Onondaga Creek occupies a deep U-shaped glaciated valley that once impounded a finger lake in the waning stages of glaciation. Therefore the valley contains glaciofluvial sand and gravel as terraces on both sides probably representing hanging deltas. Lacustrine deposits floor the valley in the vicinity of the site and probably underlie the sands and gravels of the hanging deltas.

19. The highest delta terrace is approximately at El. 800 feet, just southeast of Nichols Corners, while the lowest is approximately at El. 520 feet near Nedrow.

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Central New York Stratigraphy

20. The rock formations of central New York are structurally simple and abundantly fossiliferous, thereby lending themselves to detailed stratigraphic studies over the last 100 years. The stratigraphic sequence dips gently southward at approximately 40 to 50 feet per mile such that the oldest units are exposed in the north and progressively younger formations are exposed southward.

21. Table II presents the generalized stratigraphic column for central New York. Figure 3 is a geologic map of the central New York region.

22. A brief description of the various rock formations in central New York is contained in the following paragraphs. A more detailed discussion of thicknesses, lithologies, and relationships of the various series, groups and formations noted on Table II is presented in Appendix A.

Upper Ordovician System

23. The oldest stratigraphic unit in central New York, exposed along the south shore of Lake Ontario, is the Queenston Formation. It encompasses a maximum of 700 feet of red, crumbly shale and red siltstone and sandstone in this region. The coarser beds increase in thickness eastward and southward.

Silurian System

24. Silurian rocks surround the Appalachian Plateau on the north and are found in the Lake Ontario Plain, and on the east in the westernmost Hudson Valley Province. The older units of the Silurian sequence consist of the Grimsby Sandstone and Clinton and Lockport Groups of the Albion and Niagaran Series. These units are overlain by shales, dolomites and limestones of the Cayugan Series.

Devonian System

25. Devonian rocks are by far the most extensively exposed in New York State. Devonian strata underlie all of the Allegheny Plateau and regions adjacent to it, such as the Lake Erie Plain, southern Lake Ontario Plain, and the westernmost Hudson Valley Province.

26. Carbonates dominate the Lower Devonian and lower part of the Middle Devonian while clastic deposits predominate in the remainder of the section. The clastics are fine-textured marine rocks at the bottom, grading upward to coarse, red continental sandstones. These deposits, are known as the Catskill Delta and are a wedge of clastic sediments that thicken and become coarser to the east. The New York Devonian is subdivided into the Ulsterian Series (Lower Devonian), Erian Series (Middle Devonian) and Senecan and Chautauquan Series (Upper Devonian).

Onondaga Dam Site Stratigraphy

27. The stratigraphic sequence in the vicinity of the dam belongs to the Lower and Middle Devonian System described on Table II. A more detailed stratigraphic column along the Syracuse Meridian in the vicinity of the dam is summarized on Table III, and Figure 4 is a geologic map of the area surrounding the dam.

28. Older Silurian units are exposed to the north of the dam in Syracuse and younger Upper Devonian units cap the hills south of the dam, and form the valley sides and floor south of the Village of Tully.

29. A brief description of the stratigraphy exposed at the location of the dam is contained in the following paragraphs. Appendix B presents an expanded discussion of this stratigraphy, along with a discussion of the underlying or surrounding stratigraphy in the general vicinity of the dam site, as shown on Figure 4.

Lower Devonian - Ulsterian Series

Helderberg Group

30. The Helderberg group is composed of two formations; the Manlius Formation overlying the Rondout Formation (Chrysler Dolostone).

31. The Manlius Formation is subdivided, from older to younger, into the Olney Limestone Member, the Elmwood A, B, and C Dolostone Members, the Clark Reservation Limestone Member, and the Jamesville Limestone Member.

- a. The Olney Member is approximately 30 ft. of fine-grained, even bedded limestones; thin bedded in the lower portion, somewhat more thickly bedded in the upper portion.
- b. The Elmwood Member overlies the Olney and consists of a series of dolostones, waterlimes, and limestones. It is subdivided into three units. Elmwood A is a drab yellowish-brown, thin-bedded and mud cracked waterlime approximately 6 ft. thick. Elmwood B is a fine-grained,

blue limestone, approximately 3 ft. thick. Elmwood C, the uppermost unit, is approximately 4 ft. thick, and is a waterlime similar to Elmwood A.

- c. The Clark Reservation Limestone Member overlies the Elmwood and consists of a fine-grained, white weathering limestone characterized by a diagonal fracture system. It is approximately 5 ft. thick.
- d. The Jamesville Limestone Member overlies the Clark Reservation Member. It is approximately 25 ft. of fine-grained, dark blue, brittle limestone in thin even beds. It is exposed in the lower 12 ft. of the rock exposed in the spillway cut at the east end of the dam.

Tristates Group

32. The Oriskany Sandstone is the only formation of this group exposed in Central New York. It is not well developed in the vicinity of the dam. Only a few inches of reddish-orange, sandy layers of this unit are exposed at the base of the Onondaga Limestone in the spillway cut.

Middle Devonian - Erian Series

Onondaga Limestone

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33. The Onondaga Limestone is the lowest Middle Devonian formation within the Syracuse area. It is a series of light bluish-gray semicrystalline limestones in even continuous layers from 1 inch to 2 feet thick, separated by thin seams of dark calcarerous shale. Flattened nodules of dark blue or black chert, sometimes in continuous sheets or beds, are unevenly distributed throughout the formation, but occur mainly in the upper part of the formation. The total formation thickness is approximately 70 ft.

34. The Onondaga is comprised of five members which, from older to younger, are the Edgecliff Member, the Nedrow Member, the Moorehouse Member, the Tioga Bentonite Member, and the Seneca Member.

- <u>a</u>. The Edgecliff Member is a coarse, crystalline limestone with abundant fossils. The base of the member is a disconformity; i.e., it was deposited over an erosion surface that had removed previously deposited rock strata. The Member is approximately 15 ft. thick and may be seen in the spillway cut.
- b. The Nedrow Member is characterized by thin-bedded, shaley limestone. It is medium gray, very fine-grained, and contains a little chert unevenly distributed at top. It is approximately 10 ft. thick and is gradational with the overlying Moorehouse Member. Both the Nedrow and the Moorehouse units are exposed in the spillway cut.
- <u>c</u>. The Moorehouse Member is a medium gray, very fine-grained limestone, with 2-in. to 5-ft. thick beds separated in many places by thin shaley partings. Total thickness of the deposit is approximately 25 ft. Chert is common in the Member, but varies in amount.
- d. The Tioga Bentonite Member is a volcanic ash layer 6 to
 9 inches thick separating the Moorehouse and Seneca Members.
 Neither the Tioga nor the overlying Seneca Member were observed in the spillway cut.
- e. The Seneca Member is similar in lithology to the upper portion of the Moorehouse Member, but is characterized by a slightly different fossil assemblage. Chert is less common in the upper part of the Member and the limestone beds become darker gray. The Member is approximately 25 ft. thick.

35. The Onondaga Limestone is overlain by the Hamilton Group which includes, from older to younger, the Marcellus Formation, the Skaneateles Formation, the Ludlowville Formation, and the Moscow Formation. These formations are approximately 1,100 ft. thick in total and are comprised

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mostly of silty shales and siltstones. Refer to Appendix B for further details. The formations of the Hamilton Group are generally exposed in ravines and gullys on the sides of the Onondaga Valley at elevations above the spillway.

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Depositional History

36. In general, the sedimentary rocks of central New York, and at the site in particular, were deposited in a shallow sea that flooded most of New York State. The sea was bounded on the east by a landmass that occupied the Hudson Valley region. This landmass was uplifted into a region of high relief during the Late Ordovician and once again in the Middle and Late Devonian Periods. These are discussed further in Part III of this report. Between these periods of uplift the landmass was an area of low relief and may even have been flooded by shallow seas from time to time. The material comprising the shales, siltstones and sandstones, such as the Hamilton Group, was sediment being transported off this landmass by westward flowing streams. When this landmass was high, coarse-grained clastics such as sandstones reached central New York, when it was low, fine-grained clastics such as shale were deposited. When this area was very low or even flooded, the sea cleared allowing for the deposition of carbonates such as the Helderberg Group and the Onondaga Limestone.

37. The abundance and diversity of organic remains attests to normal marine conditions prevailing during deposition. The sea was probably warm, shallow, well lighted, and of normal salinity and oxygen content. Notable exceptions to this would be conditions represented by the Queenston and Grimsby Formations, Salina Group, and the black and dark shales of the Marcellus Formation.

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Glaciation

38. During the Pleistocene Epoch, the study area was covered by glacial ice approximately 2 miles thick. The erosional effect of the ice was to deeply scour the Valley of Onondaga Creek into a deep U-shaped trough.

39. As the ice waned northward, its margin paused momentarily south of the site, depositing the Tully Moraine, near Tully, New York. With continued northward withdrawal of the ice, a thin veneer of ground moraine or glacial till was laid down on top of bedrock. The ice margin then paused along the northern edge of the Allegheny Plateau causing a proglacial lake to form south of the ice front and north of the Tully Moraine. Since the ice acted as a dam, the impounded waters were forced to flow eastward to the adjacent north-south valley (Butternut Creek Valley) across the sharp ridge between the two. Therefore, a dramatic series of west to east meltwater channels were carved across the highland ridge between the Onondaga Valley on the west and the Butternut Creek Valley on the east. As the ice margin slowly withdrew northward, successively lower channels were cut across the northerly sloping face of the Helderberg-Onondaga Escarpment marking the northern rim of the Allegheny Plateau. One of the more spectacular of these is the extinct waterfall of Clark Reservation State Park west of Jamesville. A large horseshoe shaped rim of a once high waterfall and plunge basin below (Green Lake) is all that remains of this meltwater channel.

40. As lower meltwater channels were carved, the lake in the Onondaga Valley was lowered until the ice finally pulled back from the rim of the Allegheny Plateau allowing the pent up waters to escape eastward to the Mohawk River draining the lake from the Onondaga Valley.

41. When the lake was present in the valley south of Syracuse, and as it gradually receded, it received meltwater from waters discharging eastward from Marcellus, through Pumpkin Hollow, to Cedervale (Figure 3). As the lake received meltwaters from the west, a dramatic series of glacial lake deltas were deposited in the vicinity of the site. These are the nearly flat terraces so pronounced in the immediate vicinity

of the dam, as well as at Gwills Corners, South Onondaga, Nichols Corners and Ironsides on the south side of the valley. They are known as hanging deltas and are shown on Figure 4. The west end of the dam is tied into one of the largest hanging deltas in the Onondaga Valley having a surface elevation of approximately 620 feet. The hanging deltas are excellent sources of coarse sand and gravel and are being actively quarried northwest of South Onondaga. Many smaller hanging deltas occur at lower elevations north of the site on both sides of the valley.

42. The surfaces of the hanging deltas correspond favorably with the threshold elevations of the various west to east meltwater channels, described above, leading eastward to Butternut Creek Valley. Therefore, the surface elevation of the lowering proglacial lake can be mapped and recognized.

43. At the dam, the soil profile, based on the geologic history, should reflect till above bedrock with lacustrine deposits above till in the middle of the valley but coarser deltaic deposits on the sides of the valley. Above this should be coarse deltaic sand and gravel as the deltas came closer to the site as a result of the lowering lake. Finally, overbank deposits and organic peat from Onondaga Creek should cap the section. A detailed discussion of the soil profile beneath the dam is presented in a later section of this report.

PART III: FIELD AND LABORATORY INVESTIGATIONS

Subsurface Exploration Program

44. Subsurface conditions in the vicinity of the dam were investigated prior to dam construction. Nine test borings, Nos. DH-1 through DH-4, DH-6, DH-8, DH-10, DH-11, and DH-16, were drilled approximately along the dam centerline as shown on Figure 5, Site and Subsurface Exploration Plan. A subsurface profile along the dam centerline, prepared by the Corps of Engineers, is shown on Figure C-1, Profile Along Dam Centerline, contained in Appendix C. Strata lines and geologic descriptions were added to Figure C-1 by H & A of New York.

45. The logs of the original test borings generally indicate compact foundation soils. However, near-surface soil and one deeper strata in DH-3, are noted to be semi-compact. The criteria used in the design phase explorations for the dam to evaluate the relative density of the soils is not known. Therefore, no specific quantitative data relative to soil density, such as blow counts, were available to make evaluations of the liquefaction susceptibility of the subsurface soils. Such data is required to assess liquefaction susceptibility and foundation stability under design earthquake loadings.

46. For these reasons, H & A of New York recommended that a limited subsurface exploration program be conducted to obtain relative density data in order to conduct liquefaction susceptibility analyses. It was recommended that three test borings be made from the dam crest, through the embankment and into the foundation soils. Standard penetration test results would be obtained and split spoon soil samples recovered for laboratory soil testing. By drilling the borings from the dam crest, data on soil type and density for the embankment materials could be obtained and added to the permanent record for the dam.

47. The three test borings, Nos. D82-101 through D82-103, were conducted during the period 28 July through 27 August 1982 by Parratt-Wolff, Inc. of East Syracuse, New York. These test borings were drilled generally in accordance with Corps of Engineers "Specifications for Subsurface Drilling Sampling and Testing". H & A of New York field

personnel monitored the test borings, described soil and rock samples, and prepared written test boring logs in accordance with standard Corps of Engineers procedures. Completed test boring logs are contained in Appendix D. As-drilled locations and ground surface elevations are shown on Figure 6. They were surveyed by Rowell and Associates, P.C., of North Syracuse, New York, using base information contained in "Plans for Onondaga Dam, Onondaga Creek, Syracuse, N.Y.", dated 1947. Elevations are referenced to Mean Sea Level Datum (MSL).

48. The test borings were drilled with truck-mounted Mobile Drill B-56, CME 55 and/or CME 75 drill rigs. Hollow stem augers with 3-3/4 in. inside diameters were used to drill through the dam embankment materials. Augers were fitted with a properly sized plug to prevent soils from entering the augers during drilling. Split spoon samples were obtained through the augers. After the augers were advanced to a depth of 60 to 78 ft., NW-size (3-3/16-in. I.D.) flush-joint casing was then set within the augers and the boring continued by driving the NW casing with a 350-1b. hammer. The casing was driven to the depth at which a sample was to be obtained and washed out using a fishtail bit. Split spoon samples were obtained below the bottom of the casing. When driving of the NW casing became difficult, BW size (2.75-in. I.D.) flush joint casing was set within the NW casing and the boring continued by driving the BW casing. In D82-103, driving of the BW casing became difficult at a depth of 150 ft. This hole was completed from this depth using an open hole filled with drilling fluid consisting of bentonite powder manufactured by N.L. Baroid. The test boring logs in Appendix A contain notes relative to the specific steps taken to advance each of the test borings.

49. Split spoon samples were taken in all test borings. When soils containing an appreciable gravel content were encountered in the embankment, a 2-7/8-in. I.D. by 3.5-in. O.D. split spoon was used. It was driven into the soil with a 350-lb. hammer free falling 18 inches. This large diameter spcon could only be used through the hollow stem augers. When sampling in all other instances, a standard 1-3/8-in. I.D. by 2.0-in. O.D. split spoon, was used. It was driven by a 140-lb. safety hammer free falling 30 inches. All split spoons were generally

driven 24 inches unless refusal was encountered. Refusal was defined as 50 blows of either hammer with less than 0.5 ft. of spoon penetration. Boulder or cobble obstructions were occasionally encountered and were removed by means of fishtail bits or tricone roller bits. The test boring logs contain notes relative to specific sampling procedures and obstruction removal procedures used for each of the test borings.

50. Sampling was conducted continuously in each of the test borings from the bottom of the embankment soils to a depth of 150 ft. or to the bottom of the hole as in D82-101. Continuous sampling was also conducted through the embankment materials in D82-102. At depths below 150 ft. and through the embankment materials in D82-101 and D82-103, sampling was conducted at 5-ft. intervals. Test boring D82-101 was terminated at a depth of 105.7 ft., D82-102 at 173.0 ft., and D82-103 at 184.8 ft.

51. When sampling below the subsurface water level, efforts were taken to generally maintain a positive head of water within the casing to prevent the soils below the casing from running into the casing. Running conditions were still encountered, however, within D82-102 at a depth of 63.0 ft.

52. Bedrock was encountered in D82-101 at a depth of 103.3 ft. During core drilling operations, it was discovered that the bottom portion of the casing had become bent when the bedrock was encountered, and only 1.9 ft. of rock was cored before a BX-size (1.375-in. I.D.) core barrel was sheared off within the hole. The core barrel was later retrieved and the hole terminated.

53. Upon completion of each test boring, the hole was filled with grout from the bottom of the hole to the top of the embankment. The casing was removed as the grout was placed. The grout consisted of a cement and bentonite mixture containing a small amount of sand.

Laboratory Testing Program

54. Laboratory soil tests were performed on selected samples obtained from the test borings in order to assist in classification of the samples. Natural water contents, Atterberg Limits, and grain size distribution tests were performed. Grain size distributions were obtained by means of mechanical sieve analyses and/or hydrometer analyses. All tests were conducted in accordance with accepted practice, generally as recommended by ASTM.

55. A total of five natural water contents and five Atterberg Limits determinations were made on fine-grained soils exhibiting some degree of plasticity (i.e., containing clay). Four of the samples tested were from the uppermost natural soil strata and one was from a cohesive layer within the lacustrine deposits found in D82-103. In addition, a hydrometer analysis was performed on each of these samples. A total of thirteen mechanical sieve analyses were performed. Four of the samples were from the embankment materials and nine from underlying natural soil strata. Representative samples were selected from most of the major soil strata.

56. Laboratory test results for all of the above tests are contained in Appendix E.

PART IV: SUBSURFACE CONDITIONS ALONG DAM CENTERLINE

57. The recent test borings were designated, from east to west, D82-101, D82-102, and D82-103 as shown on Figure 6. They encountered five significant soil deposits and bedrock. These were as follows:

. Dam Embankment Materials

- . Fluvial Overbank Deposits
- . Deltaic Deposits
- . Lacustrine Deposits
- . Glacial Till
- . Bedrock

58. Bedrock was encountered in the eastern-most boring only, D82-101. Not all soil deposits were encountered in each test boring. A subsurface profile along the dam centerline developed from both the recent test borings and the original design test borings is shown on Figure 6. The following paragraphs describe the major soil deposits encountered in the recent borings, beginning with the embankment materials and progressing downward. Refer to Paragraphs 38 through 43 for a discussion of the geologic origin of the deposits.

Dam Embankment Materials

59. The dam embankment materials may generally be described as brown, silty coarse to fine SAND, little gravel, to coarse to fine sandy GRAVEL, little silt, with an occasional layer of medium to fine SAND, trace gravel, trace silt.

- <u>a</u>. Embankment materials were encountered to depths of 55.0, 65.5, and 65.5 ft. in borings D82-101, D82-102, and D82-103, respectively.
- <u>b</u>. Penetration resistance values were generally obtained with a 2-7/8 in. I.D. by 3.5-in. O.D. split spoon sampler driven by a 350-lb. hammer falling freely

18 inches. Blow counts for one foot of penetration ranged from 19 to refusal, defined as 50 blows with less than 0.5 ft. penetration. Generally, measured penetration resistances exceeded 50 blows/ft. A granular soil of this average penetration resistance can be described as very compact.

- <u>c</u>. Grain size distribution curves of typical embankment samples are contained in Appendix E as Test Nos. 1, 4, 5 and 9. For the samples tested, the embankment materials consisted of 34 to 56% gravel, 31 to 39% sand, and 13 to 28% silt.
- <u>d</u>. As a result of the above visual descriptions and laboratory test results, the dam embankment materials would be classified as GM, SM, or SP soils according to the Unified Soils Classification System (USCS).

Fluvial Overbank Deposits

60. The fluvial overbank deposits consist of gray and brown, fine sandy SILT, little clay, to silty CLAY, with an occasional layer of silty fine SAND. Small amounts of organic material are also present. They occur immediately below the embankment material, and are the original near-surface soils of the valley bottom.

- a. The deposit was 11.0, 6.5, and 5.0 ft. thick in borings D82-101, D82-102, and D82-103, respectively.
- b. Penetration resistance values were generally obtained with a 1-3/8-in. I.D. by 2.0-in. O.D. standard split spoon driven by a 140-lb. safety hammer falling freely 30 inches. Standard penetration resistance values of 1.0 ft. of penetration ranged from 17 to 34, with an average of 24. A cohesive soil of this average penetration resistance can be described as stiff; a granular soil as medium compact.

 d. As a result of the above visual descriptions and laboratory test results, the deltaic deposits would be classified as GM, GW, SM, and SP soils according to the USCS.

Lacustrine Deposits

62. Lacustrine sediments underly the deltaic deposits. These sediments consist of red-brown, silty fine SAND, trace coarse to medium sand and fine gravel, to silty CLAY, trace coarse to fine sand, with occasional layers of coarse to fine SAND, little gravel, little silt.

- a. The lacustrine deposit was encountered in D82-102 from a depth of 120.5 to 154.0 ft. and in D82-103 from 101.5 to 181.0 ft. It was not encountered in D82-101. The deposit was 33.5 and 79.5 ft. thick in D82-102 and D82-103, respectively.
- b. Samples were recovered with a standard split spoon sampler. Standard penetration resistance values ranged from 32 to 128 blows/ft. with an average of 77 blows/ft. A granular soil with this average penetration resistance would be described as hard.
- <u>c</u>. Grain size distributions of representative samples of coarser portions of this deposit (silty fine SAND) are contained in Appendix E as Test Nos. 8, 12, and 13. These soils contained 60 to 72% fine sand and 28 to 40% silt. Atterberg Limits No. A-5 and hydrometer analysis H-5 reported in Appendix E were performed on a sample of silty clay soil recovered in D82-103 from the uppermost portion of the deposit. This sample had a natural water content of 13.6%, a plastic limit of 12.9%, and a liquid limit of 21.7%, and consisted of 10% sand, 30% silt, and 60% clay.
- <u>d</u>. As a result of the above visual descriptions and laboratory test results, the lacustrine deposits would be classified as SM, SP, ML, and CL soils according to the USCS.

Glacial Till

63. Glacial till underlies the lacustrine deposit. Descriptions of recovered samples range from gray, gravelly, coarse to fine SAND, little silt, to brown, fine sandy SILT, little coarse to medium sand and gravel. Numerous cobbles and boulders were indicated during casing advance and sampling.

- <u>a</u>. This deposit was penetrated 19.0 and 3.8 ft. by D82-102 and D82-103, respectively. These borings terminated within this deposit. Glacial till was not encountered in D82-101.
- b. Standard penetration resistance values ranged from 71 blows/ft. to refusal. A granular soil of this penetration resistance can be described as very compact.
- <u>c</u>. No grain size distributions of this deposit were obtained.
- <u>d</u>. As a result of the above visual descriptions, the glacial till would be described as SM, GM and GW soils according to the USCS.

Bedrock

64. Bedrock was encountered in only one test boring; D82-101 at a depth of 103.3 ft. Attempts to core the rock met with limited success due to bent casing at the bottom of the hole. A run of 1.9 ft. was made using a BXM double tube core barrel with 1.6 ft. of rock core recovered. The recovered rock core has been described as a dark gray, slightly weathered, moderately hard, dense to finely crystalline, flatlying LIMESTONE, with a few stylolites. It has been identified as the Manlius Limestone, Olney Member.

Subsurface Water

65. At the time of the dam construction, a series of piezometers and settlement gages were installed. The settlement gages were constructed such that subsurface water level readings could also be obtained. Readings have been obtained at semi-regular intervals since the dam was constructed. These readings indicate a subsurface water level of approximately EL 460± to 465±, closely corresponding to the original ground surface. During the recent subsurface exploration program, regular water level observations were made during drilling operations and the water level in adjacent settlement gages was measured. These levels, noted on the subsurface profile on Figure 6, also correspond closely with the original ground surface.

Subsurface Profile

66. The subsurface conditions along the centerline of the dam are quite variable. The previous subsurface explorations reveal that the soils of the east and west dam abutments are deltaic deposits. These deltaic deposits form the terraces rising above the Onondaga Creek flood plain. However, the recent test borings along with the previous test borings show that, within the limits of the original flood plain beneath the dam, a deposit of fluvial overbank sediments overlie deltatic deposits, which in turn overlie lacustrine deposits. Glacial till and bedrock lie below the lacustrine deposits. The depths and thicknesses of these deposits vary from east to west.

- <u>a</u>. The silt and clay overbank deposits were encountered between test borings DH-6 and DH-2. The overbank sediments are approximately 5 to 13 ft. thick at test boring locations, within the flood plain, and pinch out to the east and west at the valley sides.
- b. Sand and gravel deltaic deposits underlie the obervank deposits in the central portion of the

dam and make up the east and west abutments. At the test boring locations, the deltaic sediments beneath the dam appear to vary in thickness from approximately 30 to 55 ft. This deposit thins to the east where the bedrock surface rises. The deltaic deposits of the west abutment consist of two units; a coarse-grained upper portion and a fine-grained lower portion. The deltaic deposits of the east abutment are predominantly granular and overlie bedrock.

- <u>c</u>. Underlying the deltaic deposit, along approximately the western one-half of the dam centerline, is a lacustrine deposit consisting of silt and fine sand. Based on test boring results, it appears the deposit thickens to 80 ft. and extends beneath the terrace at the west abutment. Towards the east it pinches out due to the rising bedrock surface.
- <u>d</u>. Underlying the lacustrine deposit, is a deposit of glacial till, a homogeneous mixture of silt, sand, and gravel.
 The top of the glacial till deposit slopes to the west as shown by the test borings and is probably closely underlain by bedrock.
- e. Test borings along approximately the eastern one-third of the dam and the east abutment encountered limestone bedrock immediately beneath the deltaic deposit. The bedrock surface appears to drop off sharply to the west, as rock was not encountered west of boring DH-2. It is H & A of New York's opinion that the rock cored in the bottom of DH-3 is a boulder in the till.

PART V: TECTONIC HISTORY

Orogenies

67. Orogenic events, or mountain building episodes, have not occurred per se in central New York. However, orogenies have occurred within and outside the state boundaries that have effected the depositional and structural history of the region.

The orogenic events are as follows:

- a. Taconic Orogeny This is the earliest mountain building episode to strike the state and took place in the Late Ordovician and Early Silurian Periods (Table I). It resulted from compressive forces that became established as ancestral Africa and Europe began to move closer to ancestral North America. These compressive forces folded and thrust faulted the area of eastern New York and western New England forming the Taconic Mountains of New York, Berkshire Hills of Massachusetts, and the Green Mountains of Vermont. Streams flowing westward of this mountainous highland (sometimes called Vermontia) deposited a westward migrating delta or clastic wedge across the state. It is known as the Queenston Delta and the Queenston and Grimsby Formation represent the terrestrial uppermost portion of this delta.
- b. Acadian Orogeny The Acadian Orogeny took place during Middle and Late Devonian times immediately after the deposition of the Onondaga Limestone. It continued into the succeeding Mississippian Period as well. The Acadian Orogeny struck the same general area as the earlier Taconic Orogeny. However, from the amount of clastic sediment shed from the highland area and other data, this orogeny was larger than the Taconic. It came about as a result of the collision of ancestral Africa and Europe with northeastern North America. This orogeny was the

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climax of the Appalachian Mountains uplift in the north. Another delta of clastic wedge was shed westward off the Acadian Landmass known as the Catskill Delta. It begins with the clastic sedimentation of the Hamilton Group and continues into the Upper Devonian (Table II). By the end of the Devonian Period this delta had prograded westward across the entire state or nearly so.

- <u>c</u>. Appalachian Orogeny This orogeny struck mostly the southern Appalachian chain as ancestral Africa rotated to collide with southeastern North America. This took place primarily during the Late Paleozoic Era, the Pennsylvanian and Permian Periods. The effect on central New York from this orogeny was to impart to the strata a dip of 40 to 50 feet per mile to the south. In addition, the joint system was formed at this time as well as the minor faults and folds discussed below.
- <u>d</u>. Mesozoic Events After the Appalachian Orogeny the present Atlantic began to open as Europe and Africa rifted from North America and drifted gradually eastward. Tensional forces resulting from this breakup formed the downfaulted Triassic Lowland Province of southeastern New York (Figure 2). In addition, upward migrating magmas formed the Palisades Sill of the Hudson and also may have been responsible for the many small dikes that occur throughout central New York. These dikes are shown on Figure 3, and discussed below.

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Structural Deformation

68. Superimposed on the regional dip approximately $1/2^{\circ}$ to the south are local folds, consisting of low amplitude anticlines and synclines, and faults of low displacement.

Folds

69. A broad anticline and syncline is exposed in the Chittenango Member of the Marcellus Formation on the east side of N.Y. Route 173 just south of its intersection with Slate Hill Road, south of Marcellus.

70. Along the east side of Interstate 81, south of the junction of Interstate 481, there are a series of gentle anticlines and synclines. The height of these folds is approximately 30 feet. They are a few hundred feet wide and expose the Manlius Formation (Olney, Elmwood, Clark Reservations and Jamesville Members).

71. A gentle anticline is exposed in the spillway cut at the dam. The Manlius-Onondaga contact rises to a maximum of 12 feet above the floor of the valley from the spillway northward, then descends further northward such that it is near creek level at the northern extremity of the exposure.

72. A monocline in the Onondaga Limestone is exposed on the south side of the Seneca Turnpike (NY Route 173) at Fillmore Corner, which is at the intersection of Sweet Road, about 2.5 miles east of Jamesville. The Onondaga is horizontal on each side of the monocline but is bent nearly vertical within the structure. The Onondaga is some 23 feet lower on the downside of the monocline. Outside of the Onondaga Valley region, a broad anticline is exposed along Cayuga Lake north of Ithaca, known as the Portland Point anticline.

73. The folds in the Syracuse and Central New York region are related either to mild compression as a result of the Appalachian Orogeny or to removal of salt and gypsum in the Salina and Bertie Groups due to migrating ground water. Thus some of the smaller folds may be subsidence structures and not uplift structures formed by compression.

Faults

74. A number of small faults occur in the Syracuse - Central New York region as revealed by a search of the literature augmented by field work. All of these are small in lateral extent and displacement or offset. Further, they strike approximately N70°W, dip mainly southward, are mostly thrust faults that form due to compression, and therefore, are probably Late Paleozoic Appalachian Orogeny structures (approximately 220 million years old). Overlying glacial deposits are not offset indicating that no motion has taken place on these faults at least since the last retreat of the ice, approximately 14,000 years ago.

75. A small thrust fault of 3 feet vertical displacement occurs in the Old Malley quarries south of Marcellus and a few hundred yards north of the anticline described above in the Chittenango Shale Member. The fault cuts the Onondaga Limestone.

76. A thrust fault, with a throw of about 5 feet, in the Onondaga Limestone and extending into the overlying Union Springs Member can be seen in the drainage ditch on the east side of U.S. Route 11, south of the Nedrow interchange to Interstate 81 at the intersection of Kennedy-Camping Road. The Union Springs Shale Member is highly contorted and tightly folded, apparently absorbing the fault motion. This fault is the closest thrust fault to the dam, a distance of approximately 2.5 miles, and cannot be traced beyond this one exposure.

77. A thrust fault in the Jamesville Quarry on the east side of the Butternut Creek Valley, north of Jamesville, has a vertical displacement of 16.5 feet in the Onondaga Limestone. This structure probably grades eastward into the monocline described above, exposed at Fillmore Corner.

78. One of the largest faults in the Syracuse area formerly was exposed in the old Russell's Quarry, now buried near the junction of N.Y. 173 and LaFayette Road in the southeast corner of the Syracuse West quadrangle. The fault plane dips south 20° and displaces the Manlius and Onondaga limestones 42 feet. South of Jamesville, on the east side of the Butternut Creek Valley, in the Onondaga County Penitentiary Quarry, a thrust fault of about 31 feet may be the same fault as the one

formerly exposed in Russell's Quarry. If so, the fault is approximately 3.5 miles long parallel to strike.

79. Four small thrust faults with throws of 2 to 4 feet cut the Bertie Group in the Fiddlers Green Gorge at Jamesville. Three thrust faults of comparable strike, dip, and displacement to those described above have been discovered on the Manlius Quadrangle to the east.

80. A series of small normal faults, down thrown to the north, can be seen along the southwest side of the Tully Valley in Lords Hill Ravine and Cass Hill Ravine, south (upstream) of U.S. Route 20, having a total displacement of approximately 45 feet. The fault planes strike $N60^{\circ}W$, and are nearly vertical. The Staghorn Point and Joshua Submembers are cut by the faults.

81. The unusual thickness of the Pompey Member in Bare Mountain Ravine may also be due to normal faulting, but the lack of outstanding key beds makes this hypothesis difficult to test.

Joints

82. The joint system in the Syracuse region contains two main sets, one nearly north-south and the other nearly east-west. Most of the joints in these sets are vertical or nearly so. Two minor sets that strike northeast and northwest are also present.

83. Inclined joints that strike about $N70^{\circ}$ to $75^{\circ}W$ and dip 30° to 80° either north or south are locally developed in the limestones in the Syracuse area. Some of the inclined joints are continuous for many tens of feet while others are limited to a single bed or group of beds and end against a shaley parting. They are well developed in parts of the Manilus Formation such as the Clark Reservation Limestone Member and the Onondaga Limestone, especially the Nedrow Member.

Dikes

84. In central New York, small dikes from a few inches to a few feet wide occupy the North-South joint set. They are composed of kimberlite

and alnoite, a high temperature, ultra-basic, igneous rock, high in iron and magnesium silicates and low in aluminum silicates such as potash feldspars.

85. One dike occurs in North Syracuse near Schiller Park, one at the DeWitt Reservoir 3 miles east of Syracuse, and another near Otisco Lake, 15 miles southwest of Syracuse. A relatively large number are exposed in Ithaca and on both sides of Cayuga Lake north of Ithaca to Taughannock Falls State Park.

PART VI: EARTHQUAKE ACTIVITY

Historic Earthquakes

86. More than 330 earthquakes with a maximum Modified Mercalli Intensity (I_0) greater than II are known to have occurred in New York State between 1720 and 1980 (Mitronovas, 1981). Figure 7 is a seismicity map for the northeast U.S. and adjacent Canada, showing all known historical and recorded events of maximum intensity greater than or equal to III for the time period 1534 to 1977. Figure 7A is a seismic zone map for the contiguous United States and Puerto Rico with the corresponding location of the dam site. An abbreviated form of the Modified Mercalli scale is shown on Figure 8, along with approximate relationships between intensity, Richter magnitude, and peak horizontal ground acceleration.

87. All earthquakes with a maximum intensity greater than or equal to IV that have occurred within 200 kilometers of the Onondaga site are listed in Table IV in order of decreasing epicentral intensity.

Geographic Distribution of Earthquakes

88. As discussed by Mitronovas (1981) and shown on Figure 9, New York State can be subdivided into three areas of relatively high seismic activity, separated by a large area of very low or no activity in the center of the State (which includes the Onondaga Dam site). These regions were designated A, B, C, and D by Mitronovas. Over 90 percent of all seismic events, and all events of maximum intensity greater than V, fall into regions A, B, and C, which have a combined area of less than 60 percent of that of the State. Table V gives a chronological list of earthquakes within New York State of maximum intensity VI or greater. The following discussions of the seismicity of the four regions of New York State are taken from Mitronovas (1981).

89. Region A is an area of moderate seismic activity. It contains the Ramapo fault system, which has been associated with recent seismic activity. One of the earliest documented large earthquakes in New York ($I_0 = VII$) occurred in Region A near New York City on 18 or 19 December 1737.

90. The Adirondack area, region B, is the most seismically active region of the state. It is part of a larger seismic area which trends northwest-southeast and extends well into Canada. There is evidence that the Adirondacks are undergoing doming, or upward movement (Isachsen, 1975). The doming may be linked to the seismic activity of the area, but this is not known for certain. The largest known earth-quake within New York State ($I_o = VIII$) occurred in this region near Massena, New York on 5 September 1944.

91. The number of known earthquakes within region C is smaller than in either of regions A or B. However, one of the two largest earthquakes in the state, near Attica, occurred in region C on 12 August 1928. The general location of this event suggests that it was associated with the southern end of the Clarendon-Linden fault system. Past and present seismicity of region C is not restricted to this fault system alone, however.

92. The central part of the state, region D, has very little seismic activity, both at present and historically. This large aseismic area separates the seismically active areas of the state into three distinct regions.

93. The available data (Mitronovas, 1981) indicates secular variation, or periodicity over time, in seismic activity in New York since 1920. Periods of greater activity, in terms of total number of events per unit of time, between 1760 and 1790, 1830 and 1880, and 1910 through the present, are separated by periods of less activity between 1790 and 1830 and 1880 and 1910. An increase since 1910 is consistent with the improvement in detection capability. Chiburis (1981) has indicated that the Adirondack region has exhibited a decreased seismic activity since 1925. The reasons for time-dependent seismicity in the northeast U.S. are not known.

PART VII: EVALUATION FOR ACTIVE FAULTS

94. Earlier sections of this report have established that mapped faults in the site vicinity are ancient ones which date back to the Appalachian Orogeny of Late Paleozoic time (over 220 million years ago). Glacial overburden and urban construction have made it all but impossible to trace faults for any great distances. The evidence indicates that all the faults are inactive at the present time and are small local occurrences of limited extent.

Assocation of Earthquakes with Tectonism and Faults

95. The earthquake distribution in New York State bears little or no proven relationship to known active or inactive faults. In region A, there is some activity that corresponds to the Ramapo fault system, but not all activity is confined to this system. In region B, the relationship to known faults or other tectonic features is not clear. The correlation between epicenters and known surface faults or other brittle structures still remains to be determined. In region C, the past and present seismicity is not restricted to the Clarendon-Lindon fault system alone. Indeed, although the seismicity extends both east and west of the fault system, it does not extend northward along the fault system, which is known to extend into Lake Ontario. The relationship of most of the epicenters in region C to tectonic structures is, as in region B, not yet known.

96. The Attica seismic region (region C) has produced several events for which some possible fault types and orientations have been obtained, based on seismological data. Hermann (1978) determined similar focal mechanisms for both the 1 January 1966 and 13 June 1967 events. A plane striking NNE parallels the trend of the Clarendon-Linden fault structure. If this is indeed the fault plane, then the motion along this plane has approximately equal components of right-lateral strike slip and reverse

faulting. Composite fault-plane solutions obtained by Fletcher and Sykes (1977) for Attica events between 1972 and 1975 have one plane also oriented NNE, but their solution indicates predominantly reverse faulting.

Lineations

97. Lineations or linear features, are those that are found in tonal changes in satellite imagery, and in the alignment of rivers, terrace boundaries, etc. They may be the result of a multitude of causes and may be unrelated to faults.

98. The Brittle Structures Map of New York State, compiled by Isachsen & McKendree (1977) was reviewed to note locations of any lineations in the site vicinity. Several lineations were noted. An attempt was made to check these in the field but nothing could be discerned. They appear to be unrelated to faults but may be due to jointing or glacial erosion and deposition. None of the faults or linears show any evidence of activity in the area of the project.

PART VIII: INTENSITY OF EARTHQUAKE SHAKING AT ONONDAGA DAM

Earthquake Intensities and Attenuation

99. All earthquakes with a maximum intensity greater than or equal to IV that have occurred within 200 kilometers (km) of the site are shown in Table IV. Because larger earthquakes at distances greater than 200 km may produce significant site intensities, all 2705 events in the Chiburis (1981) earthquake catalog have been attenuated to the Onondaga site. The attenuation relation used was that of Cornell and Merz (1975):

 $I = I_0 + 3.211 - 1.3 \ln (distance in km)$

100. Table VI shows the specific events which produce site intensities greater than or equal to IV when attenuated to the site.

101. As the data clearly show, the maximum shaking intensity experienced at the Onondaga Dam site has been IV to V. (Intensities noted in Table VI are in arabic numerals because of fractional values). The peak horizontal ground acceleration was therefore in the range of 1 to 3 percent of the acceleration due to gravity (0.01g to 0.03g). The event that produced the highest intensity at the site was the 1663 La Malbaie earthquake ($I_0 = X$) at a distance of about 700 km from the site. Because of the early occurrence at this event, the assigned intensity of X, based on historical accounts (written or otherwise) is at best only an estimate. Of more concern would be the Attica and Massena earthquakes, recorded in 1929 and 1944, at distances of 180 km and 250 km from the site, respectively. The magnitude of the Massena event was reported to be 5.9, whereas the magnitude of the Attica event was estimated from historical records to range from 5.5 to 5.9 (Chiburis, 1981). However, the attenuated intensities at the dam are still only on the order of IV. A possible faulting mechanism for the Attica region was discussed earlier in this report (para. 96).

102. Isoseismals from the 1 March 1925 intensity IX event centered in La Malbaie, Quebec are shown in Figure 10. This event resulted in

intensity IV shaking at the site, and would have caused rattling of dishes and windows and disturbance of doors and cars near the site. The reduction in five levels of intensity from La Malbaie to central New York for the 1925 event is consistent with the attenuation of the 1663 La Malbaie earthquake discussed in the previous paragraph.

103. It is concluded that earthquakes with epicentral intensity greater than IV will not have epicenters near the Onondaga site. The nearest regions of concern are the Attica and Massena regions, which may produce intensities of IV to V at the site. The focal mechanism of the Attica region is most probably reverse or thrust faulting with north-easterly trending fault planes.

Frequency of Earthquake Occurrence and Seismic Risk

104. A seismic risk analysis was performed to aid in the selection of a level of earthquake shaking for use in assessing the seismic stability of Onondaga dam. The analysis was used to estimate the probability of exceeding various levels of earthquake shaking within a specified time period. These calculated probabilities allow for selection of an "analysis" earthquake that is consistent in terms of seismic risk with other similar structures.

105. The procedure used to calculate the probability of exceedence of specific levels of earthquake shaking at the site was to first divide the northeastern United States and adjacent areas into more than 50 zones or potential sources of seismic activity. Each zone has its own assigned level of seismic activity based on historical seismicity. The zones are quadrilateral in shape and are based on the regionalization given in Chiburis (1981). It is assumed that the locations of potential earthquake point sources (epicenters) are randomly distributed with each zone. In addition to the specific zones, a low-level "background" seismicity was assumed, randomly distributed over the entire northeast area. This allows for the possibility of earthquakes occurring at any location, not just in zones where historic seismicity is relatively concentrated.

106. The probability of occurrence of intensity i \geq I₁ shaking at the site was calculated based on the "Total Probability Theorem" as

$$P(i \ge I_1) = \sum_{\substack{n=1\\n \text{ source}}}^{n} \sum_{\substack{m \in I \\n \text{ source}}} P(i \ge I(m) P(m)$$

in which:

P (i $\ge I_1$) denotes the probability of the occurrence of intensity i $\ge I_1$ at the site; P (m) denotes the probability of occurrence of an earthquake of magnitude m or greater in a source zone; P (i $\ge I_1 \mid m$) denotes the probability of occurrence of intensity i $\ge I_1$ at site given the occurrence of a

magnitude m event in the source zone.

107. The probability calculations were performed on a computer using a program developed by McGuire (1977). The procedures used to calculate the needed parameters for each zone are described in the following two paragraphs.

108. The probability of occurrence of an earthquake of magnitude m or greater within a source zone, P(m), is determined from the assigned recurrence relationship for each zone. The recurrence relationship is of the form:

 $\log N = A-Bm$

in which N is the average number of events per year of magnitude m or greater for each specific source zone, and A and B are constants for the zone. The probability of occurrence of an earthquake of magnitude m or greater is determined by numerical integration of the above recurrence relationship for each zone. The "B" parameter was assumed to be equal for all zones except La Malbaie, which exhibits much more seismic activity than other areas in the northeast. The "A" parameter was determined from historical seismic activity for each zone.

109. The probability of intensity i shaking at the site given a magnitude m event in a source zone, P ($i \ge I_1 \mid m$), is related to the attenuation relationship for the site area. Attenuation is assumed to follow a normal probability distribution. The mean value of the distribution was based on Cornell and Merz (1975) as presented in paragraph 99. Chiburis (1979) compared the Cornell and Merz relationship to observed attenuation for six northeastern events, as shown on Figure 11. The relationship compares well with the observed data in the near and far fields, although it is not as good in the intermediate range. To account for the variability in the observed attenuation in the analysis, a standard deviation of 0.4 was assumed. This is more conservative than the 0.2 value used by Cornell and Merz.

110. The probability of exceedence results calculated with the computer program are based on a one-year time period. The probability of exceedence of a specified intensity I_1 in T years can be calculated from the following:

 $P(i \ge I_1) = 1 - e^{-\lambda} T$

in which λ is the mean number of events per year of intensity I_1 . For the levels of seismic risk considered herein, λ is also equal to the probability of exceedence of intensity I_1 in one year.

111.. Figure 12 depicts the results of the analysis, that is, the probability that specified earthquake intensities and corresponding peak horizontal ground accelerations will be exceeded at the site in a 50-year period. Intensity was converted to peak horizontal ground acceleration in rock using the relationship shown on Figure 8. The following is a brief summary of selected results:

PEAK HORIZONTAL GROUND ACCELERATION IN ROCK	INTENSITY	PROBABILITY OF EXCEEDENCE IN 50 YEARS
0.02g	v	7.0%
0.05g	VI	1.0%
0.10g	VII	0.8%

On the average, intensity V, VI and VII shaking would be expected to recur at the site every 650, 4,400 and 65,000 years, respectively.

PART IX: ASSESSMENT OF EARTHQUAKE EFFECTS ON ONONDAGA DAM

112. Based on the results of the seismic risk analysis and current engineering practice, the peak horizontal ground accelerations in rock recommended for use in assessing the seismic stability of the Onondaga Dam is 0.05g. This acceleration corresponds to an intensity of shaking of VI, which is one intensity level higher than the site is known to have experienced historically. The probability that this level of shaking would be exceeded in a 50-year period is 1.0 percent based on the analysis performed.

113. The duration of strong motion in an earthquake is often defined as the time interval between the times that a 0.05g acceleration is first and last achieved. This definition would result in a duration of zero for Onondaga Dam. Nuttli (1979) has defined duration as the time interval between the times that one-half the "sustained maximum" (third highest acceleration peak) acceleration is first and last achieved. This definition leads to longer durations for small earthquakes. Using this definition of duration and considering that the epicenter of the causative earthquake is likely to be more than 100 km from the site, the recommended duration of shaking is approximately 30 seconds.

114. The recommended level of earthquake shaking could only cause embankment or spillway instability, if a marginal static stability condition exists for either the embankment dam or concrete spillway. Based on the calculations presented in the U.S. Engineer Office Design Analysis for the dam (Ref. 17), the static stability of the embankment and spillway appears adequate. Significant settlement or liquefaction of foundation soils c uld only occur if the foundation soils are cohesionless and very loose to loose in relative density. In addition, for liquefaction to occur, the foundation soils must be saturated. The logs of the design phase (1945) borings indicate that foundation soils at the dam are predominantly compact to very compact, saturated cohesionless soils, but the criteria used to evaluate the relative density

of the soils is not known. No penetration test results or quantitative data of any kind were available. The recent test borings were undertaken as part of this study in order to provide such data. Engineering studies relative to the seismic stability of the dam were then performed using the data obtained from the borings.

115. The subsurface conditions along the dam centerline are shown on Figure 6 and have been summarized in paragraphs 57 to 66. As noted in paragraph 65, subsurface water levels are generally between El. 460 to 465.

116. A plot of Standard Penetration Resistance (N-value) versus sample depth below the top of the embankment for the recent test borings is provided in Figure 13. N-values obtained by means of standard methods, 1-3/8 in. I.D. 2.0-in. O.D. split spoon driven by a 140-lb. hammer free falling 30 inches, are shown on the plot without correction. N-values obtained by means of the large split spoon, 2.875-in. I.D. 3.5-in. O.D. driven by a 350-lb. hammer free falling 18 inches, have been converted to standard values for use on the plot. The conversion has been made on the basis of sampler-hammer ratios which take into account sampler size, hammer weight, and drop height by the following equations:

> $R_s = (D_0^3 - D_1^3)/WH$ (144) ft.²/lb. for granular soils $R_s = (D_0^2 - D_1^2)/WH$ (12) ft./lb. for cohesive soils

where;

R_s = Sampler - Hammer Ratio D_o = Outside Diameter of Sampler (in.) D_i = Inside Diameter of Sampler Shoe (in.) W = Weight of Hammer (lb.) H = Height of Drop (in.)

Calculations of the sampler-hammer ratios and the conversion charts used are contained in Appendix F.

117. In addition, projected N-values have been used for samples which encountered refusal before reaching the required one foot of

penetration. In these cases the recorded penetration resistance values were projected by direct proportion to the one foot standard for use on the plot.

118. The liquefaction potential of the saturated cohesionless foundation soils was analyzed in accordance with the methodology of Seed and Idriss (1981). The calculations performed for the assessment of liquefaction potential are provided in Appendix F. A summary of the analytical procedure is as follows:

- 1. Calculate the effective ($\overline{\mathcal{O}} v_0$) and total ($\mathcal{O} v_0$) overburden pressures below the crest and toe of the dam.
- 2. Calculate the seismic shear stresses (\mathcal{T}) imposed by the earthquake on the foundation soils and obtain the seismic shear stress ratio, \mathcal{T} / $\tilde{\mathcal{O}}$ v_o, as a function of depth below the toe and crest of dam.
- 3. Correct Standard Penetration Resistance values to N_1 Values, the N-value at an effective overburden pressure of one ton per sq. ft. (tsf.)
- 4. Plot $T / \overline{\sigma} v_o$ versus N₁ and compare to data obtained from previous earthquakes for known cases of liquefaction and nonliquefaction.

119. To determine the seismic shear stress as a function of depth, it is necessary to estimate the amplification of the bedrock earthquake motions as cuased by the presence of the overburden soils. The amplification by the dense soils at the site was estimated by taking the ratio of the response spectra peak for "stiff" soil conditions to the peak for "rock", for the spectra shown on Figure 14. The ratio of the "stiff" soil peak to the "rock" peak was 1.16. An amplification factor of 1.2 was assumed in the analysis, resulting in a peak horizontal ground acceleration of 0.06g at the natural ground surface.

120. The results of the liquefaction analysis are shown on Figures 15 and 15A, which shows the plot of $\overline{\zeta}$ / $\overline{\mathfrak{S}}$ v_o versus N₁ for the dam foundation soils. The figures show that even the lowest N₁ values obygained at the site lie

to the "safe" side of the boundary line separating regions of observed liquefaction from nonliquefaction. Therefore, liquefaction of the foundation soils is not believed to be possible under the recommended levels of earthquake shaking.

121. It should be noted that the N_1 range shown on Figures 15 and 15A is the range of the <u>lowest</u> N_1 values encountered at the site. The average N_1 for the foundation soils would be significantly greater than the highest end of the range shown.

122. The embankment soils were disclosed by the borings to be medium compact to very compact, as noted earlier. In addition, these soils were found to be partially saturated. The results of the foregoing analyses provide a basis for concluding that the embankment soils would not be liquefaction-susceptible for the recommended level of shaking if they were to become saturated due to an impoundment.

PART X: CONCLUSISONS

123. The Onondaga Dam is located in a region of exposed gentle, southerly sloping Lower and Middle Devonian limestones and shales of the Helderberg Group, Onondaga Limestone and Hamilton Group. Superimposed on the dip are some minor, small faults and folds of Late Paleozoic age. Therefore, the dam is situated in an area that is structurally simple and tectonically stable.

124. The region has been glaciated, and the valley in which the dam is located is a U-shaped glacial trough. Upon retreat of the ice, large, hanging delta terraces of sand and gravel were deposited in the valley in a gradually receding proglacial lake that escaped eastward through meltwater channels.

125. The dam is located in an area that historically can be described as nearly aseismic. The intensity of ground shaking recommended for use in assessing the seismic stability of the dam and its foundations is VI, with a corresponding peak horizontal ground acceleration of 0.05g in rock and 0.06g at the surface of natural ground where the dam overlies soil deposits. The average length of time between occurrences of this degree of shaking at the site is 4,400 years, based on probability analyses. The calculated probability that the above intensity will be exceeded at the site in a 50-year period is 1.0 percent. Possible sources of earthquakes which might produce this degree of shaking are the Attica and Massena, New York areas and the La Malbaie, Quebec area.

126. To obtain qualitative and quantitative data for use in the assessment of the stability of the dam under earthquake shaking, three test borings were drilled from the crest of the embankment into the foundation soils as a part of this study.

127. The approximately 65-ft. high embankment consists of medium compact to very compact sands and gravels that were not saturated at the time of this study. Immediately below the embankment is a relatively thin layer of saturated, stiff to hard, clayey silt and silty clay.

Below this layer, foundation soils consist primarily of saturated, medium compact to very compact sand with varying amounts of gravel and silt, to a depth of approximately 35 to 65 ft. below the base of the dam. Below this granular deposit is a medium compact to very compact fine grained deposit of variable thickness overlying glacial till. On the east side, bedrock is relatively close to the bottom of the embankment.

128. The static stability of the embankment and foundation soils is adequate based on the type and relative density of the soils present and the geometry of the dam. Likewise, the stability of the concrete spillway, which bears on rock, also appears to be adequate.

129. The liquefaction susceptibility of the soils at the site was analyzed for an earthquake producing 0.06g peak acceleration at the surface of natural ground. The embankment and foundation soils are not considered to be susceptible to liquefaction for this level of earthquake shaking. Minor seismically-induced settlement of the embankment and subsoils may occur, but the settlement will not be detrimental to the performance of the structure.

130. As a result of the analyses conducted, it is the opinion of H & A of New York that a more in-depth analysis of the stability of the dam due to liquefaction or other earthquake effects is not warranted.

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ONONDAGA DAM, NEW YORK

TABLES

U. S. Army Engineer District, Buffalo 1776 Niagara Street Buffalo, New York 14207

	T	ABLE I		
	GEOLOG	IC TIME SCALE		
			TIME BEFOR (MILLIONS	
ERA	PERIOD	EPOCH	BEGINNING	DURATION
CENOZOIC	Quarternary	Holocene (recent)	0.01	
		Pleistocene	3±	2.99
ſ	Tertiary	Pliocene Miocene	<u>5</u> 25	2 [±] 20
	ł	Oligocene	34	9
	t t	Eocene	49	15
		Paleocene	64	15
	Cretaceous		130	66
MESOZOIC	Jurassic		180	50
	Triassic		220	40
	Permian		270	50
	Pennsylvania n		325	55
PALEOZOIC	Mississippian		355	30
	Devonian		410	55
	Silurian		430	20
	Ordovician		500	70
	Cambrian		600	100
PRECAMBRIAN				
FORM	ATION OF THE EART	H	4,600	
NOTE: Table	derived from var	ious geologic s	ources.	- -

TABLE I

	GENER	ALIZED STRATIGR	APHIC COLUMN		PPROXIMATE
	SYSTEM	SERIES	GROUP		HICKNESS N FEET
	Upper Devonian	Chautauquan Senecan	Conewango Conneaut Canadaway West Falls Sonyea	(Many formations recognized in each group)	3000 2600 670 to 800
IAN			Genesee	Tully Ls.	900 to 1450 15 to 100
DEVONI	Middle Devonian	Erian	Hamilton	Moscow Ludlowville Skaneateles	160 to 240 300 300 to 500 100 to 500
				Onondaga Ls.	70 to 100
	Lower	Ulsterian	Tristates	Oriskany Ss.	0 to 10
	Devonian		Helderberg	Manlius Ls. Chrysler Dol.	0 to 200 15
				Chrysler Dol. Cobleskill Ls.	15 10
IAN	Upper	Cayugan	Bertie	Williamsville Dol. Scajaquada Dol. Fiddlers Green Dol	50
I L U R	Silurian		Salina	Camillus Sh. Syracuse Sh. Vernon Sh.	200 200 600
S	Middle Silurian	Niagaran	Lockport Clinton	Sconondoa Ls. (Many formations)	80 to 150 320
	Lower Silurian	Albion	Medína	Grimsby Ss.	30
	Upper Ordovician			Queenston Sh.	700

TABLE II

System	SERIES	GROUP	IAN - ONONDAGA		APPROXIMATE FHICKNESS IN FEET
Upper Devonian	Senecan		Tully Ls.		30
Jeromium	builden		Moscow	Windom Sh.	165
	1			Portland Point Ls.	12
			Ludlowville	Owasco Ss. Spafford Sh. Ivy Point Ss. Otisco Sh. Centerfield Ls & Ss.	2 27 37 209 25
Middle Devonian	Erian	Hamilton	Skaneateles	Butternut Sh. Pompey Sh.& Siltst. Delphi Station Sh.& Siltst. Mottville Ls.& Sh	135 60 125 6
			Marcellus	Cardiff Sh. Chittenango Bl. Sh. Cherry Valley Ls. Union Springs Bl. Sh.	180 80 3 12
			Onondaga Limestone	Seneca Ls. Tioga Bentonite Moorehouse Ls! Nedrow Ls! Edgecliff Ls!	25 .4 25 10
		Tristates	Oriskany Ss*		0 to .2
Lower Devonian	Ulsterian	Helderberg	Manlius	Jamesville Ls* Clark Reserv. Ls. Elmwood Dol. Olney Ls.	25 5 10 30
			Rondout	Chrysler Dol	30

TABLE III

	TABLE IV				
EARTHQUAKES WITHIN	200 KM OF	ONONDAGA	DAM	SITE	

Year	Date	Loca Lat(N)	tion Long(W)	Maximum* Intensity	Richter Magnitude	Location
1929	Aug. 12	42.9	78.4	VIII	5.8	Attica, NY
1954	Feb. 21	41.2	75.9	VII		Wilkes-Barre, PA
1840	Jan. 16	43.0	75.0	VI		Herkimer, NY
1853	Mar. 12	43.7	75.5	VI		Lowville, NY
1857	Oct. 23	43.2	78.6	VI		Buffalo, NY
1954	Feb. 24	41.2	75.9	VI		Wilkes-Barre, PA
1966	Jan. l	42.8	78.2	VI	4.7	Attica, NY
1966	Jun. 13	42.9	78.2	VI	3.9	Attica, NY
1916	Feb. 2	42.9	74.0	v		Mohawk Valley, NY
1916	Feb. 3	43.0	74.0	v		Mohawk Valley, NY
1922	Dec. 8	44.4	75.1	v		S. of Canton, NY
1929	Dec. 2	42.8	78.3	v		Attica, NY
1952	Aug. 25	43.0	74.5	v		Johnstown, NY
1955	Jan. 21	43.0	73.8	v		Malta, NY
1955	Aug. 16	42.9	78.3	v		Attica, NY
1971	May 23	43.8	74.5	v	3.9	Blue Mt. Lake, NY
1971	May 23	43.9	74.5	v	3.6	Blue Mt. Lake, NY
1971	Jul. 10	43.9	74.5	v	3.4	Blue Mt. Lake, NY
1907	Jan. 24	42,8	74.0	IV		Schenectady, NY
1927	Mar. 14	44.6	75.4	IV		Canton, NY
1929	Dec. 3	42.8	78.3	IV		Attica, NY
1933	Oct. 29	43.0	74.7	IV		St. Johnsville, NY
1946	Oct. 28	41.5	76.6	IV	3.6	LaPorte, PA
1946	Nov. 10	42.9	77.5	IV	3.1	Canandaigua, NY
1950	Mar. 20	41.5	75.8	IV	3.3	NW of Scranton, PA
1954	Jan. 31	42.9	77.3	IV		Canandaigua, NY
1954	Feb. l	43.0	76.7	IV	3.3	Montezuma, NY
1958	May 6	42.7	73.8	IV		Albany, NY
1960	Jan. 22	41.5	75.5	IV	3.4	Carbondale, PA
1963	Mar. 2	41.5	75.7	IV	3.4	Scranton, PA
1963	Jul. 2	42.6	73.8	IV	3.3	Albany, NY
1965	Jul. 16	43.0	78.1	IV	3.2	Attica, NY
1965	Aug. 27	43.0	78.1	IV	3.1	Attica, NY
1969	Aug. 13	43.3	78.2	IV	2.8	Albion, NY
1971	Jun. 21	43.9	74.5	IV	3.3	Blue Mt. Lake, NY
1973	Jul. 15	43.9	74.4	IV	3.4	Blue Mt. Lake, NY
1973	Jul. 15	43.9	74.4	IV	3.2	Blue Mt. Lake, NY
1973	Jul. 16	43.8	74.5	IV	3.3	Blue Mt. Lake, NY
1975	Nov. 3	43.9	74.6	IV	3.9	Raquette Lake, NY

* Modified Mercalli Intensity

Ref.: Chiburis (1981)

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TABLE V

Chronological list of significant (1 $_{\odot}$ $\stackrel{*}{\scriptstyle\sim}$ VI) earthquakes within NeW York

		Loca	tion	(1)	(2)	(3)
Year	Date	Lat (N)	Long (W)	Region ⁽¹⁾	<u>1</u> (2)	<u>м</u> (3)
1737	Dec. 18	40.8	74.0	A	VII	
1775	July 6	43.5	73.5	в	VI	
1853	Mar. 12	43.7	75.5	В	VI	
1857	Oct. 23	43.2	78.6	С	VI	
1867	Dec. 18	44.7	75.2	В	VI	
1877	Nov. 4	45.3	74.0	В	VII	
1884	Aug. 10	40.6	74.0	А	VII	
1897	May 27	44.5	73.5	В	VI	
1929	Aug. 12	42.9	78.3	С	VIII	5.9
1931	Apr. 20	43.4	73.7	в	VII	5.0
1944	Sept. 5	45.0	74.9	В	VIII	5.9
1944	Sept. 5	45.0	74.9	В	-	4.6
1957	Mar. 23	41.6	74.8	A	VI	4.8
1966	Jan. l	42.8	78.2	с	VI	4.7
1967	Jun. 13	42.9	78.2	С	VI	3.9
1974	Jun. 7	41.6	73.9	Α	-	3.3
1975	Jun. 9	44.9	73.6	в	-	4.2

1. See Figure 6.

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I = Maximum Intensity (Modified Mercalli Scale)
 M = Generalized Magnitude

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Ref.: Mitronovas (1981)

TABLE VI

LIST OF EARTHQUAKES HAVING ATTENUATED INTENSITIES GREATER THAN OR EQUAL TO IV AT THE ONONDAGA DAM SITE

Ettewin	REGTON	AT EPICENTER Intensity (I) Magni	CENTER Magnitude	DISTANCE FROM SITE (KM)	ATTENUATED2)
		1	2000 11601		
Feb. 5, 1663	La Malbaie, Quebec	10		705	4.7
Aug. 12, 1929	Attica, N. Y.	8	5.8	182	4.4
Sept. 5, 1944	Massena, N. Y.	8	5.9	252	4.0
Oct. 20, 1870	Baie-StPaul, Quebec	6		668	3.8
Mar. 1, 1925	La Malbaie, Quebec	6	6.6	705	3.7
June 11, 1638	La Malbaie, Quebec	6		707	3.7
1534	Les Eboulements, Quebec	6		713	3.7
Nov. 18, 1929	Grand Banks, NF	10	7.2	1608	3.6
Sept. 16, 1732	Montreal, Quebec	8		352	3.6

(1) Modified Mercalli Intensity(2) At Dam Site

Ref: Chiburis (1981)

ONONDAGA DAM, NEW YORK

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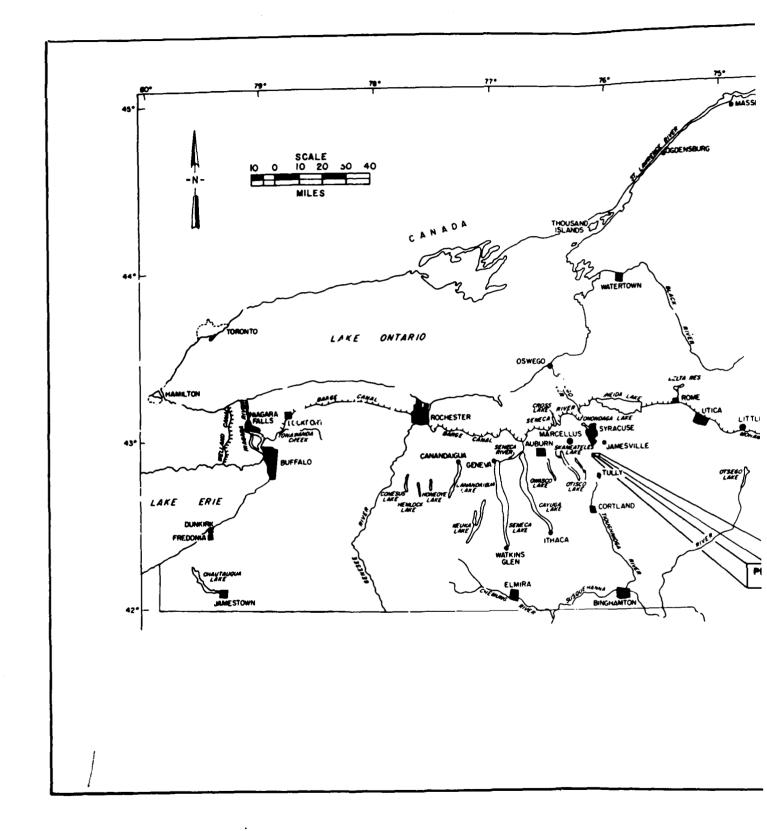
.

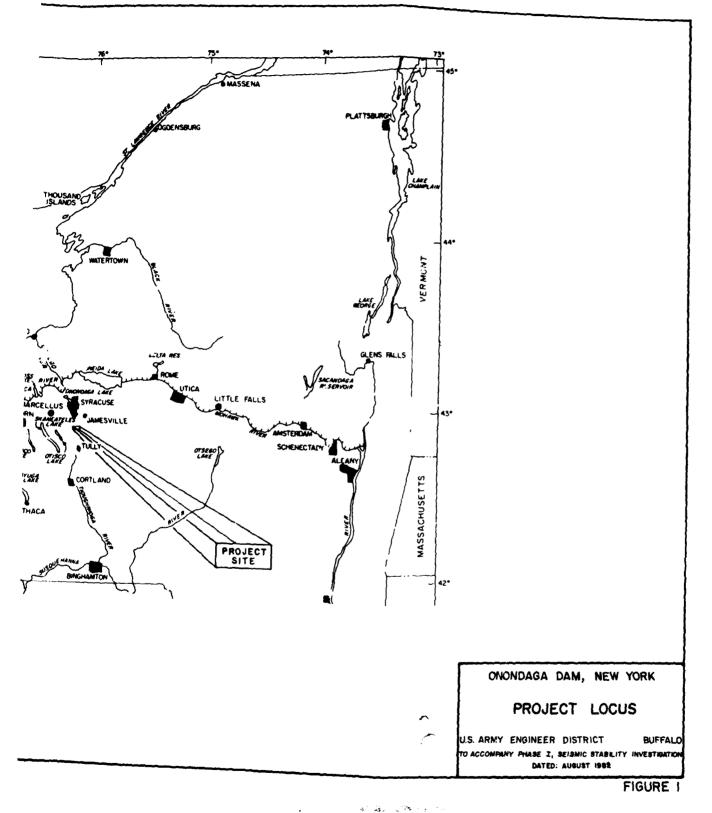
FIGURES

U. S. Army Engineer District, Buffalo 1776 Niagara Street Buffalo, New York 14207

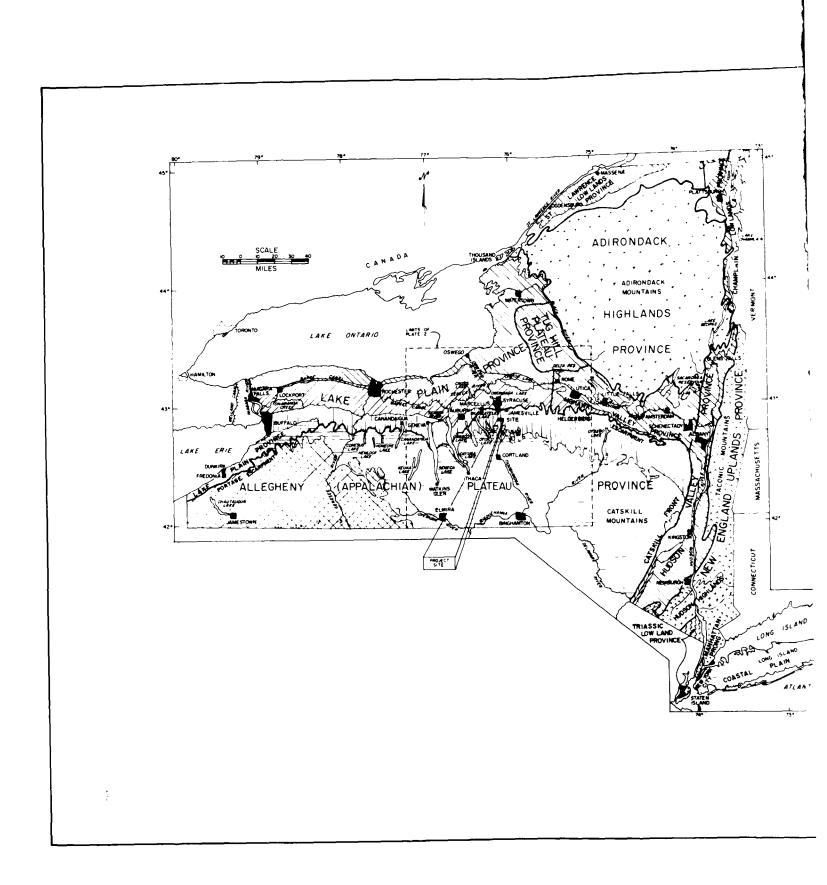
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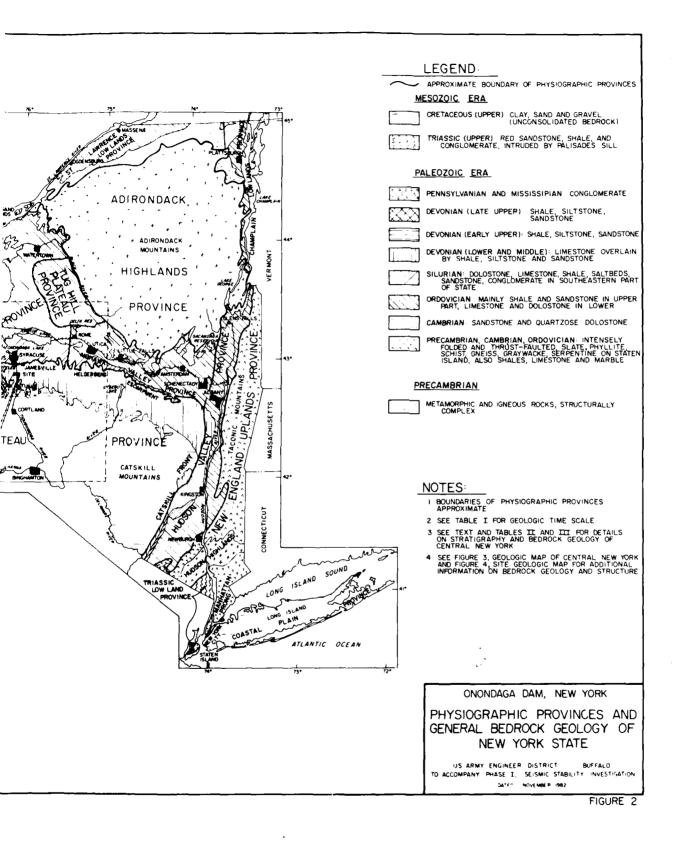
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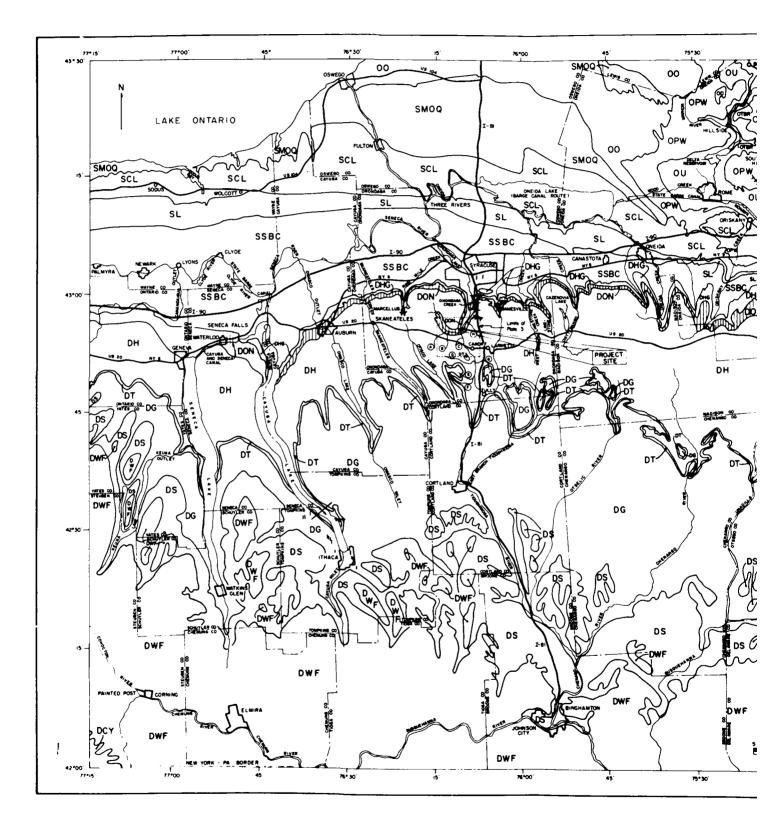


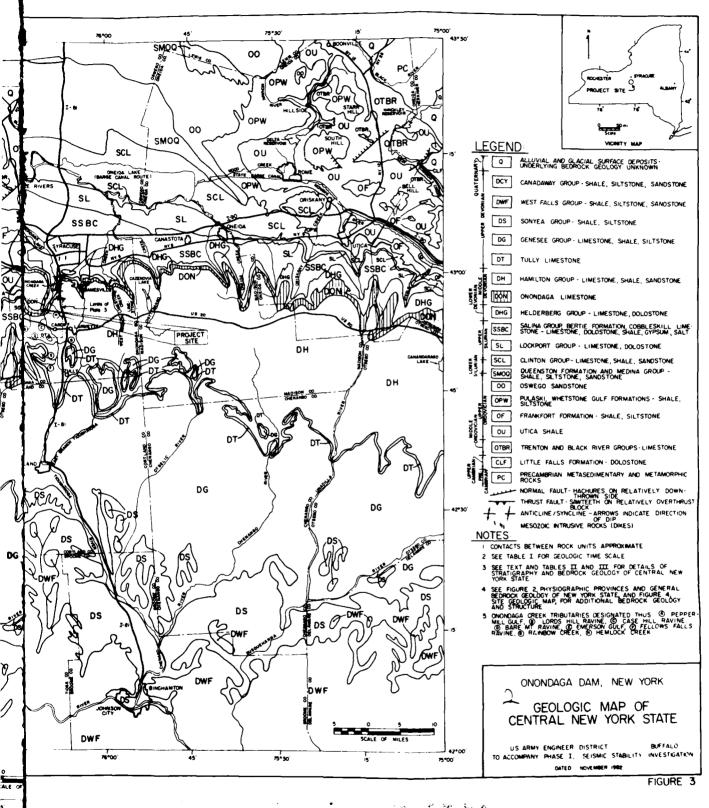


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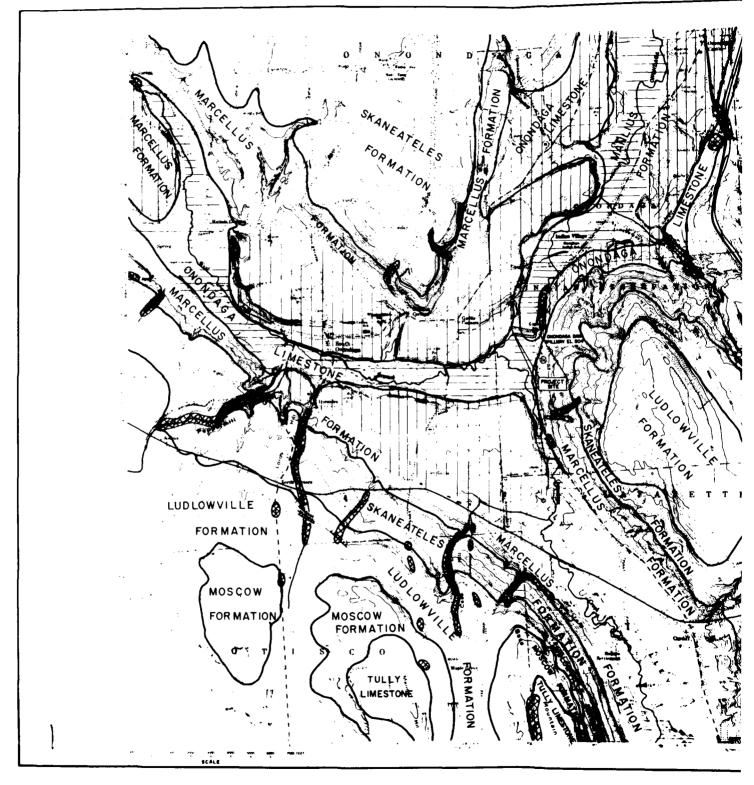




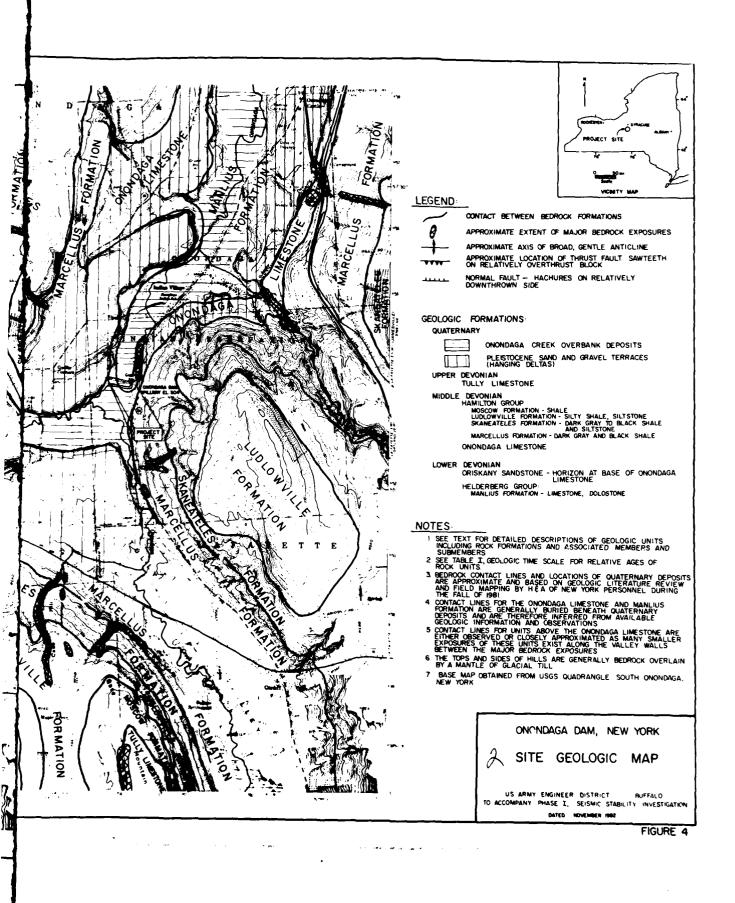


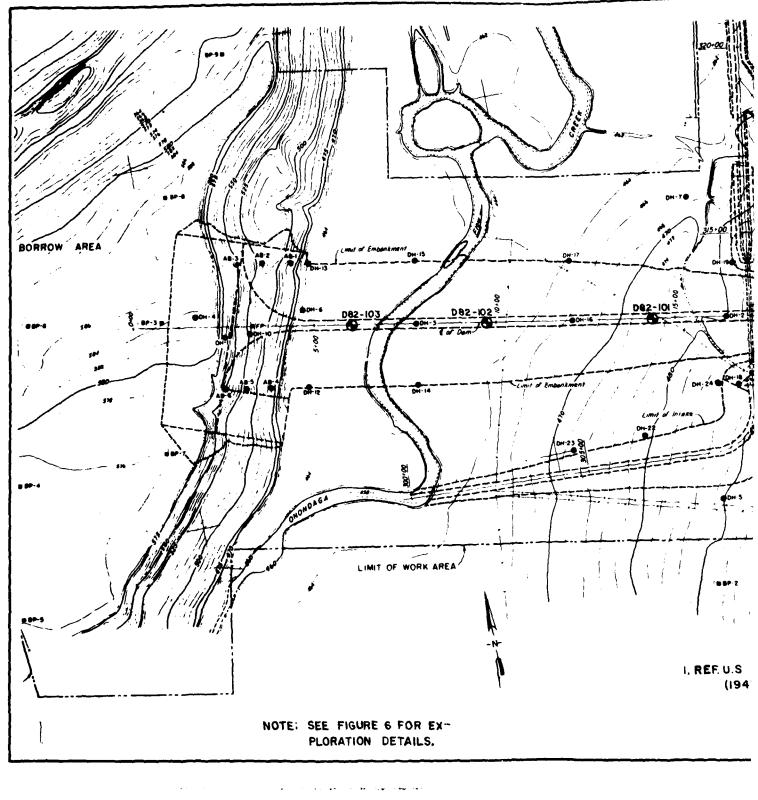


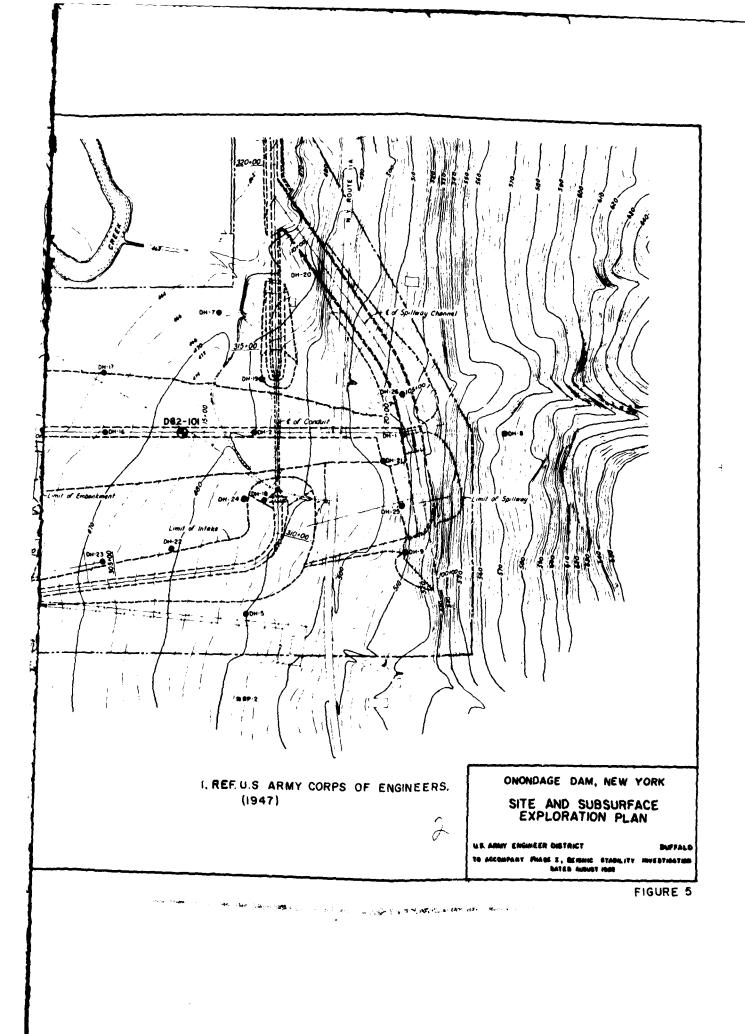
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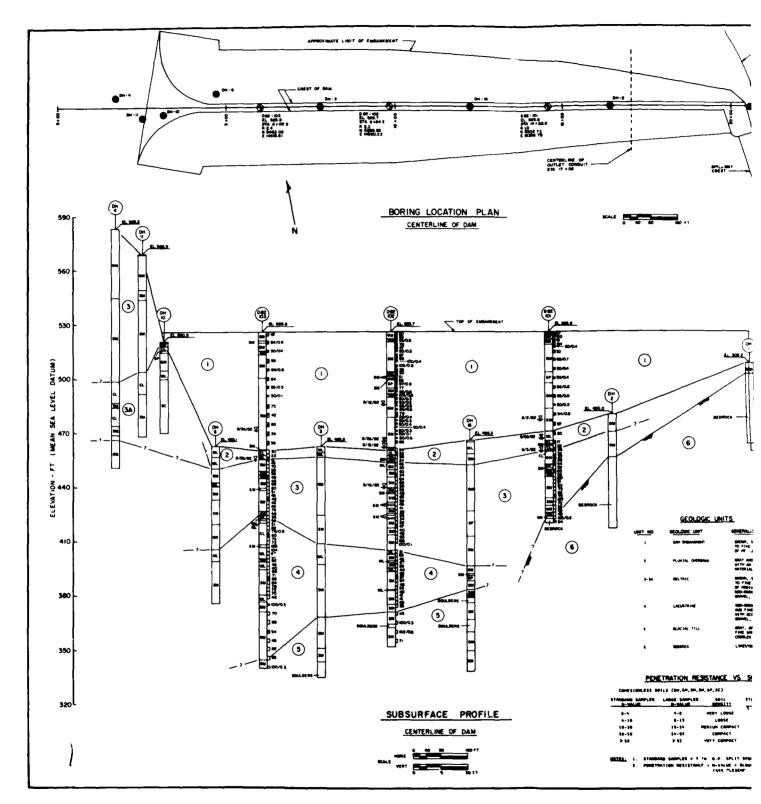


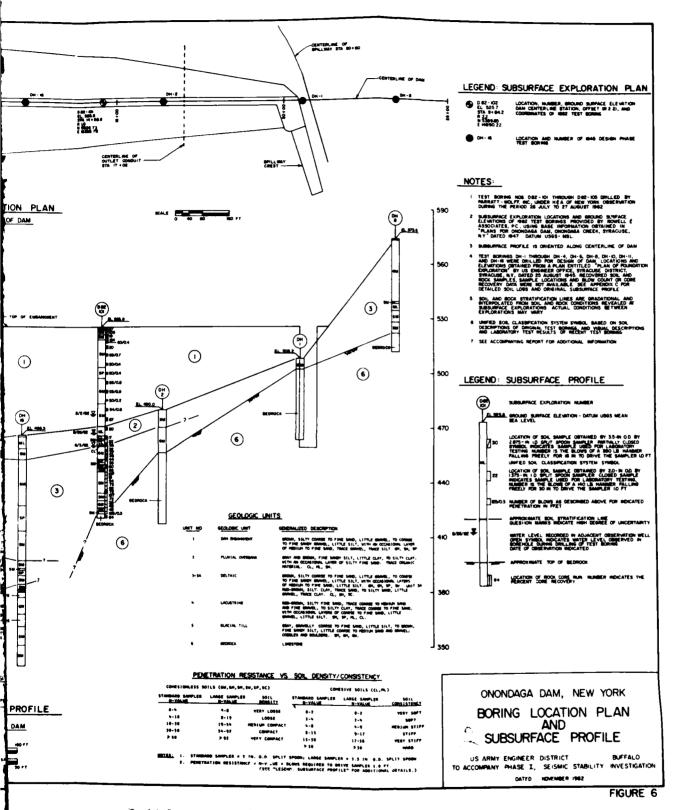
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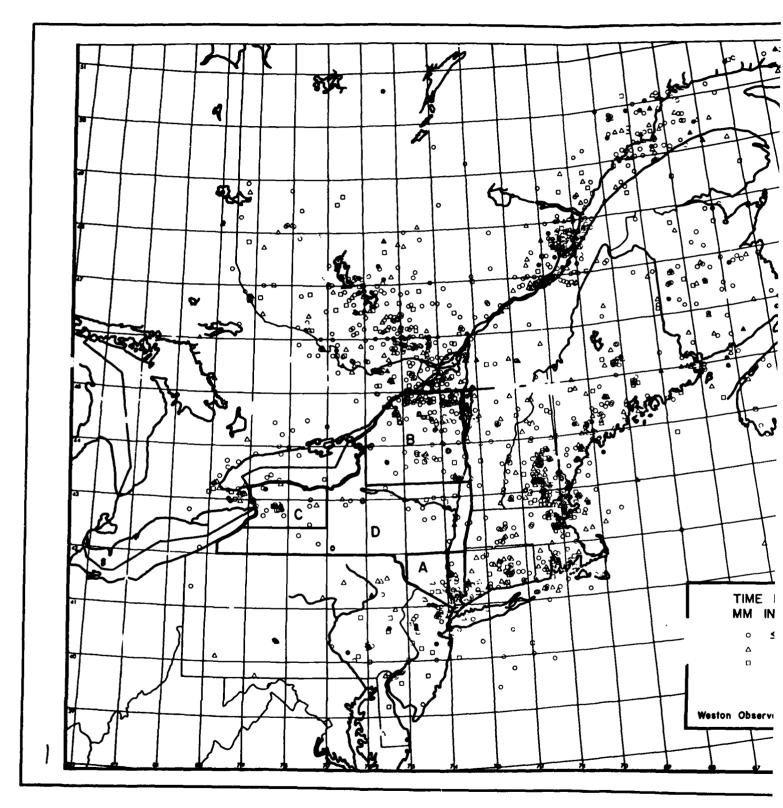




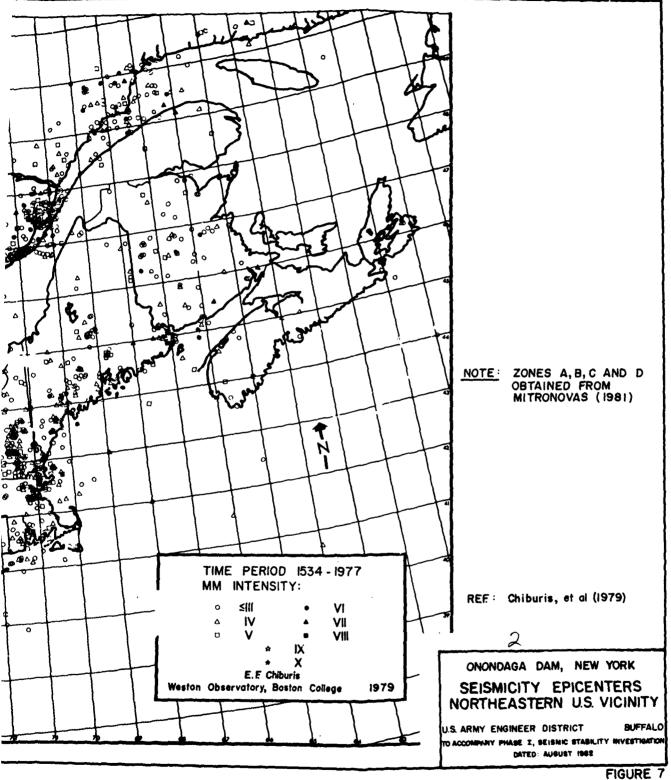


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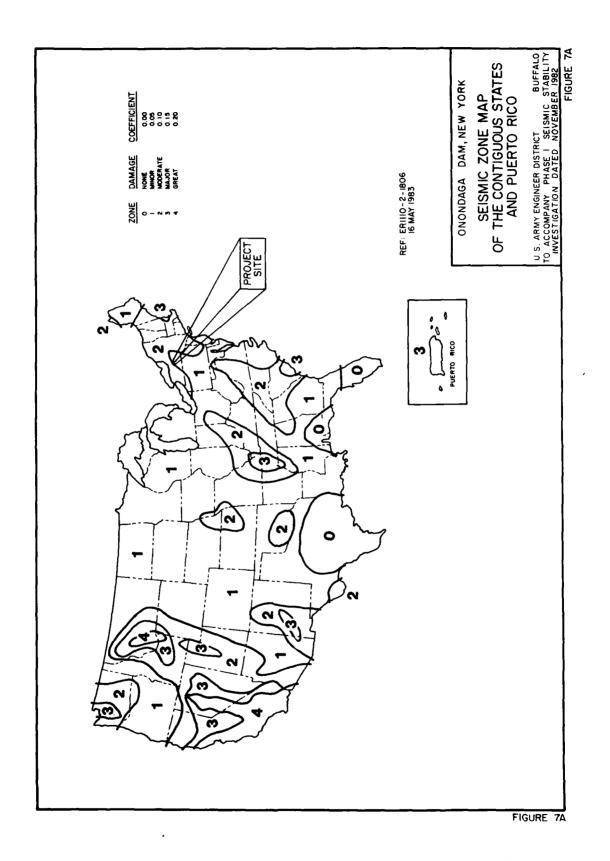
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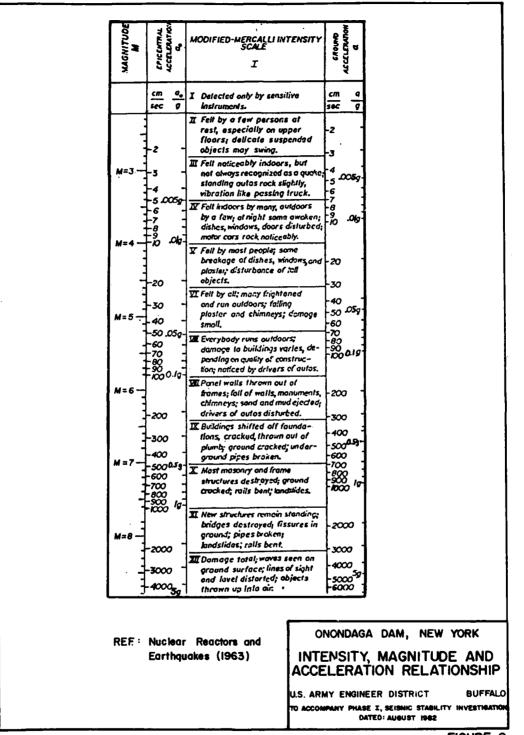


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FIGURE 8

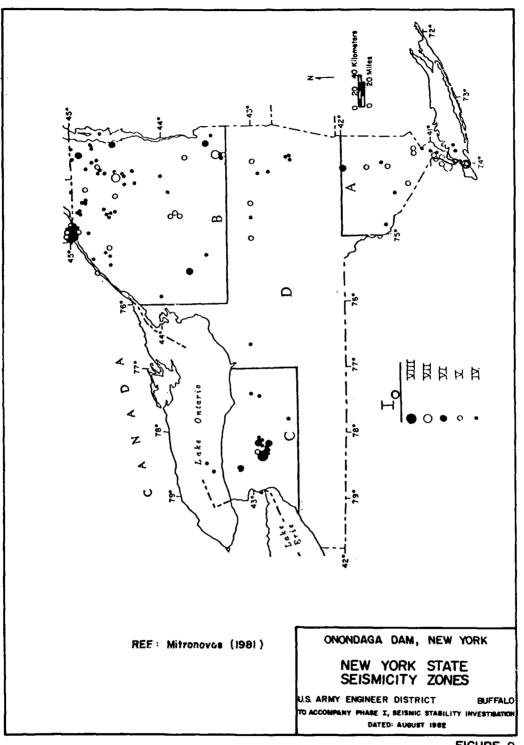


FIGURE 9

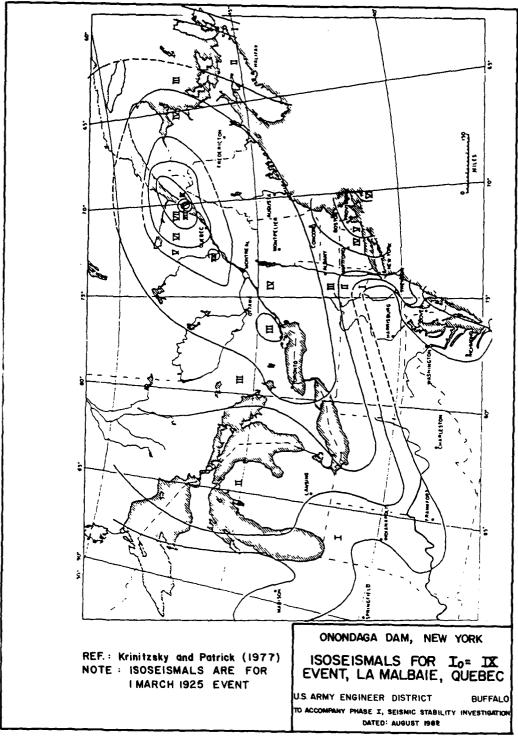
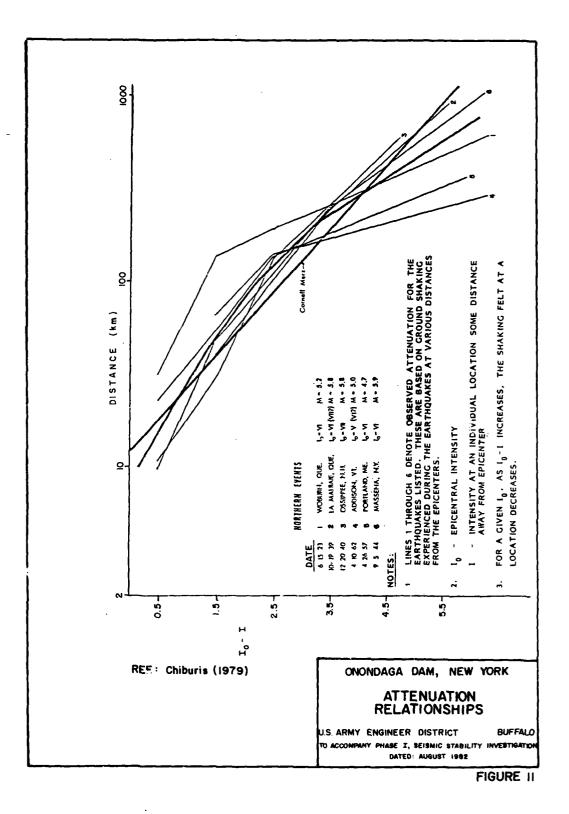


FIGURE 10



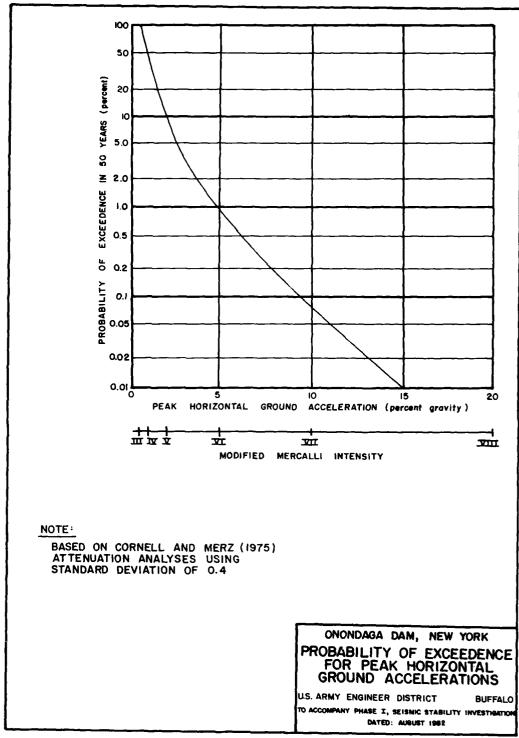
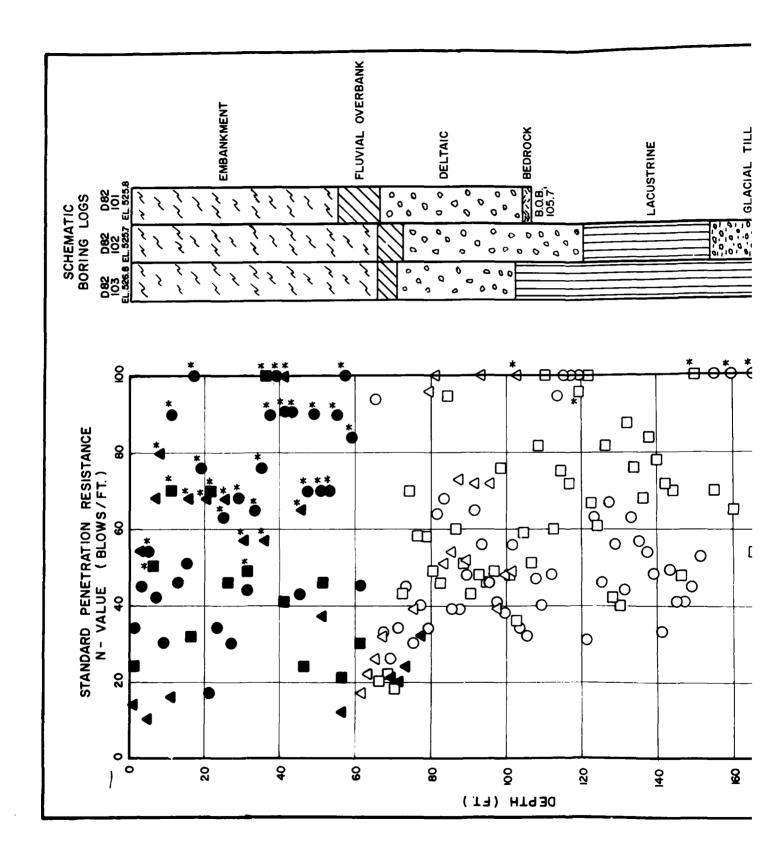
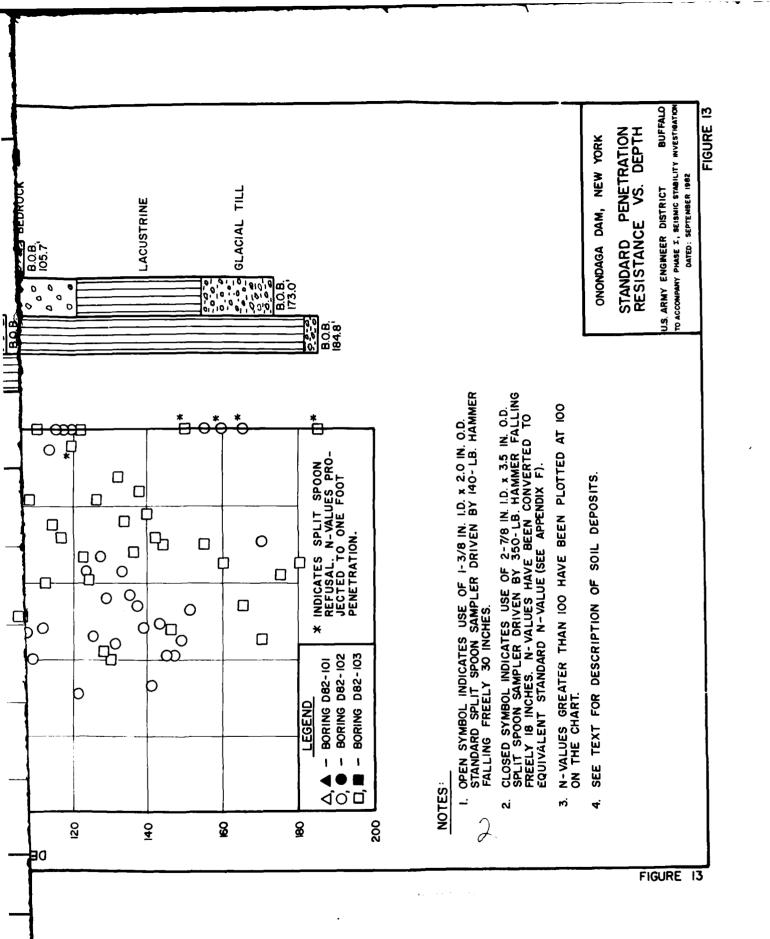
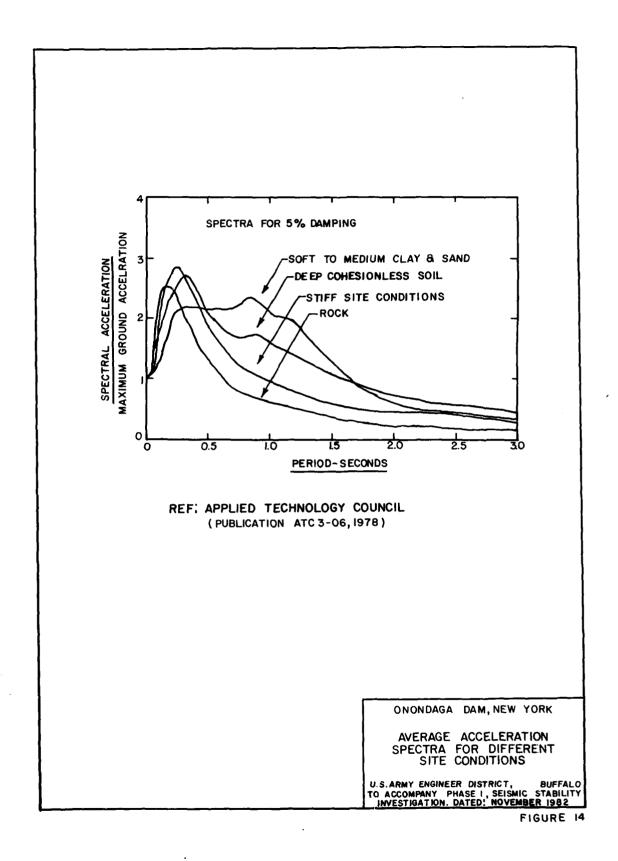


FIGURE 12







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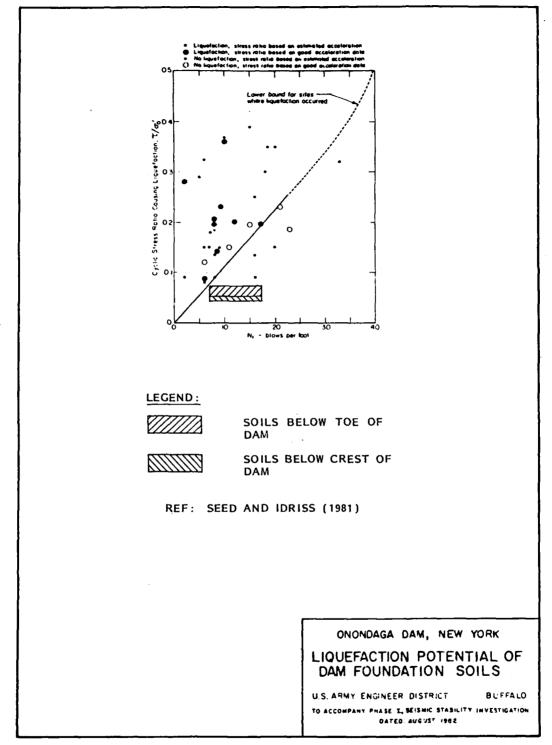


FIGURE 15

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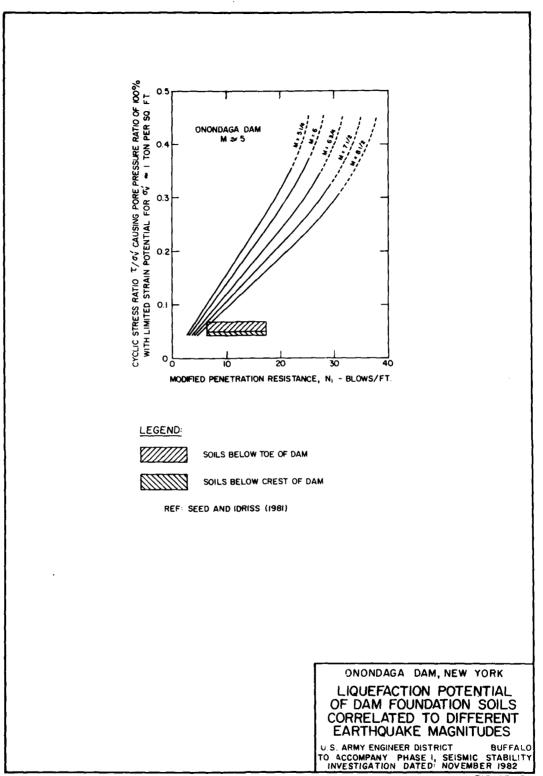


FIGURE 15A



APPENDIX A

CENTRAL NEW YORK STRATIGRAPHY

A1. Table II presents the generalized stratigraphic column for central New York and Figure 3 is a geologic map of the central New York Region.

A2. As shown on Table II, the rocks in central New York range in age from Late Ordovician along the south shore of Lake Ontario to Late Devonian at the New York-Pennsylvania border. Refer to Table I for the geologic time scale.

Upper Ordovician System

A3. The oldest stratigraphic unit in central New York, exposed along the south shore of Lake Ontario, is the Queenston Formation. It encompasses a maximum of 700 feet of red, crumbly shale and red siltstone and sandstone in this region. The coarser beds increase in thickness eastward and southward.

Silurian System

A4. Silurian rocks surround the Appalachian Plateau on the north and are found in the Lake Ontario Plain, and on the east in the westernmost Hudson Valley Province. In the Helderberg region in the northeast corner of the Allegheny Plateau, Lower Devonian rocks rest upon the Middle Ordovician and the normally intervening Silurian strata are entirely missing.

A5. For convenience, the New York Silurian can be subdivided into three parts. A lower portion, mainly of red sandstone beds (Medina Group); a middle part consisting of shales and carbonates (Clinton and Lockport Groups); and an upper division of more unusual strata, mostly hypersaline evaporites and brackish lagoonal deposits (Salina and Bertie Groups).

A6. Complexities in the Silurian rocks include numerous facies changes from west to east, the presence of many depositional breaks (disconformities), and poor exposures especially in the Upper Silurian section. Small fossils called <u>conodonts</u> and <u>ostracods</u> have been used to develop refined correlations in the Silurian Sy em of New York State, and individual strata can now be identified from rock core of exploratory borings. These rocks have a simple geologic structure, dipping gently southward at 40 to 50 feet per mile (65 to 80 feet in the subsurface) south of Lake Ontario.

Albion Series (Lower Silurian)

A7. Lower Silurian rocks are exposed in a narrow east-west belt parallel to the south shore of Lake Ontario and just south of the exposure belt of Queenston rocks, pinching out by unconformity east of the Oswego River. The line of Barge Canal eastward from Lockport to Spencerport follows the outcrops of this unit.

A8. The Medina Group is the sole Albion-aged unit in New York State. It is made up of, in the Niagara region, the Whirlpool Sandstone (25 feet thick) below, Power Glen Formation (30 feet thick) and the Grimsby Formation or Red Medina Sandstone above (50 to 70 feet thick). East of Lockport, the Whirlpool and Power Glen rocks pass by facies change into the lower Grimsby Formation. At Rochester, the Grimsby Formation rests directly on the Queenston Formation.

A9. The Grimsby Formation consists of red, crossbedded, medium bedded to massive, medium to coarse, locally conglomeritic, arkosic sandstone with some red silty shale. East of Holly and Medina, the massive sands become thin and interbedded with finer silty shales. To the west, the silty shales become thicker.

Niagaran Series (Middle Silurian)

A10. Niagaran rocks are exposed in an east-west belt from the Niagara River through Rochester to the Helderberg region anl in a northeast-

southwest belt in the western Hudson Valley Province from the Kingston area to Port Jervis. Niagaran rocks are missing around the Northeast perimeter of the Allegheny Plateau. On the Lake Ontario Plain, the Niagaran Series embraces the Clinton Group below and the Lockport Group above.

All. The Clinton Group is a complex suite of lithologies and facies changes, with numerous disconformities. Therefore, no Clinton formation traverses the state without some change in character. In general, coarse clastic sedimentary rocks dominate this section in east central New York, with fine shales and carbonates being restricted to western New York. The section is about 120 feet thick at Niagara, 180 feet in the Rochester Gorge, 320 feet in the Oneida Lake region, and some 220 feet at its easternmost outcrops in Oneida and Herkimer Counties. It can be divided into a lower, middle and upper part based on fossil zones.

Al2. Lower Clinton rocks represent a transgressive sequence of rock types, including the basal Kodak (Thorold) Sandstone (six feet thick), Oneida conglomerate (12 to 50 feet), the Neahga (six feet), Maplewood (20 feet), Bear Creek (10 feet), and Sodus (60 feet) shales, and finally terminating upward with Reynales (7 to 17 feet) and Wolcott Limestone (20 feet).

Al3. Middle Clinton formations are restricted to central and east central New York, none being found west of the Sodus Bay - Seneca Lake meridian. The Sauquoit Formation is the sole Middle Clinton unit (0 to 80 feet), consisting of shales with interbedded sandstone and siltstones. The Sauquoit strata change eastward into the Otsquago Sandstone (100 to 130 feet). At the top of the Sauquoit, the fourth Clinton iron ore bed is found, the Westmoreland, south of Rome and Utica. The Otsquago and the eastern and coarsest Sauquoit beds have yielded several gas and salt water shows in Madison and Otsego Counties.

Al4. Upper Clinton rocks, unlike the Middle Clinton units, are found across New York, east to west. The Irondequoit Limestone (18 to 30 feet), Williamson Black Shale (0 to 30 feet) and Willowvale Shale

(0 to 80 feet) are all facies of one another comprising the lowermost portion. In eastern New York, the Kirkland Hematite rests on top of the Dawes Dolostone (ten feet), the uppermost Clinton iron ore. The most significant portion of the Upper Clinton is made up of the Rochester Shale and its eastern facies, the Herkimer Sandstone. Rochester rocks are relatively thick, gray to dark gray, calcareous, dolomitic, silty shale containing numerous limestone and dolostone beds and lenses, especially in the upper portion. In the Genesee region, an upper dolostone is so prominent that a separate mappable unit, the Gates Dolostone, is recognized. The Rochester Shale is about 90 feet thick in the west and increases to 140 feet thick in central New York, north of Syracuse.

Al5. The Herkimer Sandstone is 70 feet of gray calcareous, thinbedded sandstone, interbedded with dark gray sandy and silty shale, and is found only in the Oneida-Utica region. To the east, this mandstone becomes a coarse, medium-bedded, gray and red sandstone with minor shale partings.

Al6. The Lockport Group is a thick (75 to 200 feet, east to west) sequence of dolostones and sandy dolostones with lesser amounts of limestone. The Lockport rocks are extremely resistant and form the Niagara or Lockport Escarpment, a major topographic feature extending from Rochester to the Niagara Gorge. East of the Rome meridian, the Lockport passes into the Ilion Formation of shales with interbedded dolostones.

Cayugan Series (Upper Silurian)

Al7. Rocks of Late Silurian age belong essentially to the Salina and Bertie Groups. Exposures of the Salina Group are relatively rare as it is composed primarily of weak shales and was, therefore, usually stripped by glacial erosion. The Bertie rocks are more resistant carbonates, and therefore overlying glacial deposits are thin. The Salina Group is about 300 feet thick near Buffalo, increasing to about 950 to 1000 feet near Syracuse, thereafter thinning rapidly to a featheredge near Richfield

Springs in Otsego County. Three formations are included in the Salina Group: Vernon Shale, salt of the Syracuse Formation, and the Camillus Formation.

Al8. The Vernon Shale is composed of variegated red and green shale and siltstone totalling 300 to 400 feet in thickness. The base of the Vernon is Late Niagaran in age in eastern and central New York, and becomes progressively younger westward.

Al9. The Syracuse Formation consists of three to four hundred feet of shale, dolostone and anhydrite which thicken rapidly southward to about 1,300 feet beneath Schuyler and Chemung Counties. With distance to the south, numerous salt beds occur, some reaching a maximum thickness of 400 feet. These salt beds rarely, if ever, outcrop at the surface, but the updip limit of the major salt beds is nearest the surface at Syracuse. This limit then strikes southwestward across the state, meaning that salt lies considerably deeper below the Genesee Valley than below Syracuse.

A20. Overlying the Syracuse Formation are 100 to 200 feet of shales, dolostones, and shaley dolostones of the Camillus Formation, containing a large quantity of interbedded anhydrite and gypsum.

A21. The remainder of the Cayugan Series above the Salina Group is composed of the Bertie Group dolostones (50 feet), Akron or Cobleskill limestones and dolostone (10 feet), and lower Chrysler Dolostone (30 feet). The Silurian-Devonian boundary occurs in the middle of the Chrysler Formation. Caverns, caves and other solution features are also common in these units.

Devonian System

A22. Devonian rocks are by far the most extensively exposed in New York State. Devonian strata underlie all of the Allegheny Plateau and regions adjacent to it, such as the Lake Erie Plain, southern Lake Ontario

Plain, and westernmost Hudson Valley Province. In addition, the beds are generally structurally simple, and the abundant fossil fauna in such rocks have led to a detailed stratigraphic description over the last fifty years.

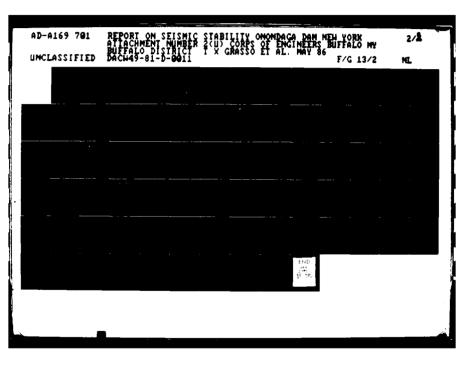
A23. Carbonates dominate the Lower Devonian and lower part of the Middle Devonian while clastic deposits predominate in the remainder of the section. The clastics are fine-textured marine rocks at the bottom, grading upward to coarse, red continental sandstones. These deposits, are known as the Catskill Delta, and are a wedge of clastic sediments that thicken and become coarser to the east. The New York Devonian is subdivided into the Ulsterian Series (Lower Devonian), Erian Series (Middle Devonian) and Senecan and Chautauquan Series (Upper Devonian).

Ulsterian Series (Lower Devonian)

A24. Lower Devonian exposures are confined to the area west of Seneca Lake, along the north edge of the Allegheny Plateau and south along the north edge of the Allegheny Plateau and south along its eastern perimeter to Port Jervis. This largely carbonate sequence is divided into the Helderberg Group below and the Tristates Group above.

A25. The Helderberg Group is a complex of intertonguing carbonate facies. Lagoonal dolostones with minor evaporites pass eastward into an offshore shelf facies dominated by limestones, some with reefs. The Helderberg Group is about 240 feet thick in the Helderberg Mountains southwest of Albany, increasing southward to some 450 feet at Port Jervis near the New Jersey border. Westward the group thins to about 100 feet near Syracuse and is absent west of Geneva. Solution features and caverns (e.g. Howe Caverns) are common in the Helderberg Group.

A26. The Helderberg Group is succeeded in an upward direction by the Tristates Group of which the Oriskany Sandstone is the lowest unit. This formation is, in turn, succeeded by the Esopus Shale, Carlisle Center





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Shale and Schoharie Limestone in eastern New York only. The Tristates Group is about 160 feet thick in the Albany region, thickening southward to about 750 feet at Port Jervis, and except for isolated patches of Oriskany, disappears west of Richfield Springs.

A27. The Oriskany Sandstone is the most important formation in the Tristates Group. It underlies much of the Allegheny Plateau, except in the southwest part of the state, and is 0 to 50 feet of quartzose sandstone, with calcite and silica cement. This sandstone is porous and permeable in siliceous zones and tight where calcareous. In the west, it is a prolific gas producer. Eastward in the Allegheny Plateau, there is no known gas production in the Oriskany Sandstone.

Erian Series (Middle Devonian)

A28. Two stratigraphic units make up the Middle Devonian in New York State, the Onondaga Limestone and the Hamilton Group. They are exposed across the entire state, in a continuous belt, from Lake Erie on the west to the New Jersey border near Port Jervis in the southeast.

A29. The Onondaga Limestone is a regional, blanket carbonate deposit, thinning from about 175 feet at Buffalo to only some 70 feet at Syracuse, but thickens eastward again to about 250 feet at Catskill and 300 feet at Port Jervis. It is a thin to medium bedded, medium to coarse textured, gray limestone, often containing bedded chert and chert nodules. Onondaga rocks are extensively quarried across the state.

A30. The Hamilton Group is a clastic wedge (Catskill Delta) deposited by westward flowing streams from a rising mountainous region to the east and southeast of the present Catskill region. The group is about 265 feet thick in western New York, 1,100 feet in central New York and 2,500 to 3,000 feet thick in the Catskills. At the toe of the delta complex, in an ocean basin in what is now western New York, fine-grained, bituminous, dark to black Hamilton shales were deposited. Uppermost Hamilton beds in western New York are gray calcareous shales. The shales pass eastward

into coarse, silty shales and sandstones. These, in turn, grade eastward and interfinger with mostly terrestrial, crossbedded, red and green sandstones, shales and siltstones.

A31. Three thin but persistent carbonate units (key beds) pierce the clastic wedge and were deposited over the delta at times of low sediment influx. These serve to subdivide the Hamilton into four formations, the Marcellus, Skaneateles, Ludlowville, and Moscow. The carbonate beds are from oldest to youngest, Mottville Member, Centerfield Member and Portland Point Member.

Combined Senecan and Chautauquan Series (Upper Devonian)

A32. The Upper Devonian sequence is a continuation of the Catskill Delta which was migrating westward during deposition. The same general Middle Devonian facies continue but persistent key beds are black shale tongues, not carbonate units, except for the lowest formation, the Tully Limestone. The black shales subdivide the Upper Devonian into a number of groups, totalling some 3,300 feet of thickness in southwest New York, 7,000 feet in south central New York, and 7,700 feet in the Catskills. The Upper Devonian is exposed throughout the southern tier region of the Allegheny Plateau, however the Chatauquan Series is present only in the western part, as it has been removed further eastward by erosion.

Senecan Series

A33. The Tully Limestone is the basal formation of the Senecan Series. It reaches a maximum thickness of 100 feet in east central New York and changes eastward into the Gilboa Shale and sandstone. The Tully is approximately 30 feet thick in the Syracuse region and is not present west of Canandaigua Lake due to its being removed by erosion prior to deposition of overlying units.

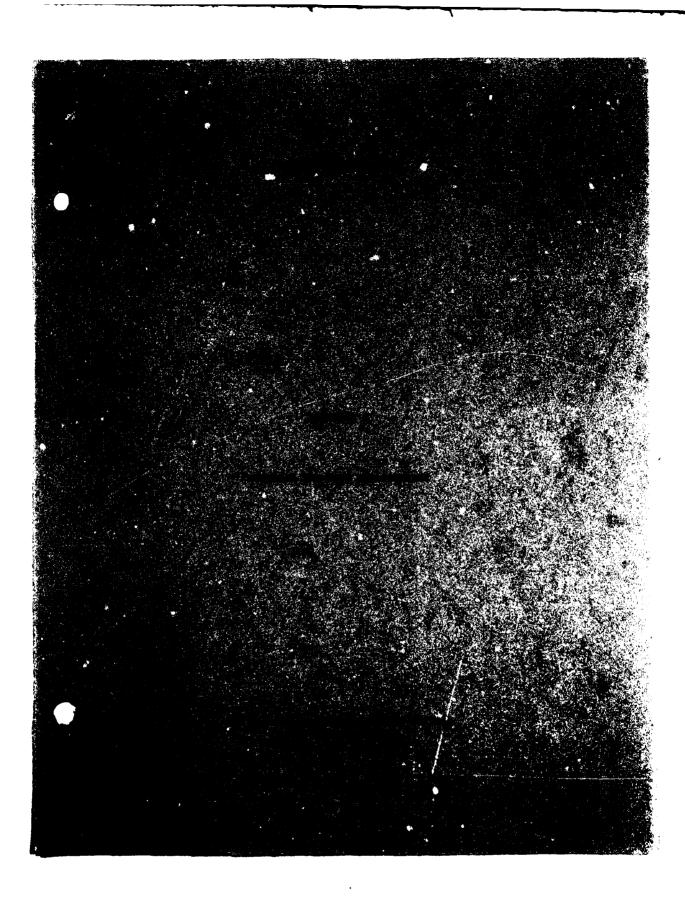
A34. Overlying the Tully is the Genesee, Sonyea and West Falls Groups. They are black shales in the west and red beds in the east, although each facies is found further westward in the West Falls Group.

Chautaquan Series

A35. The Chautauquan Series is made up of the Canadaway, Conneaut, and Conewango Groups. The Canadaway and Conneaut Groups contain several fresh water aquifers as well as the major oil producing sands of SW New York. They have been penetrated by numerous borings but records of many of the early wells have now been lost. Much of the oil is trapped along axis of the broad anticlinal structures that characterize the southwest portion of the Allegheny Plateau. If salt water is not present, the oil may also be found along the synclinal axes.

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APPENDIX B ONONDAGA DAM SITE STRATIGRAPHY

Bl. Older Silurian units are exposed to the north of the dam in Syracuse and younger Upper Devonian units cap the hills south of the dam, and form the valley sides and floor south of the Village of Tully.

B2. Table III presents the detailed stratigraphic column along the Syracuse Meridian in the vicinity of the dam. Figure 4 is a geologic map of the area surrounding the dam. A description of the stratigraphy summarized on Table III and Figure 4 follows.

Lower Devonian - Ulsterian Series

Helderberg Group

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B3. The Helderberg Group in the Onondaga Valley region is composed of two formations which, in ascending order, are the Rondout Formation (Chrysler Dolostone) and the Manilus Formation (Table III). The Rondout is poorly exposed in this area, however, a limited outcrop is exposed in the base of the rock cuts on Interstate 81, just south of the junction with Interstate 481.

B4. The Manilus Formation is subdivided from older to younger, into the Olney Limestone Member, Elmwood A, B and C Dolostone Member, Clark Reservation Limestone Member and the Jamesville Limestone Member.

> a. Olney Member - Named for Olney Station on the now abandoned Auburn and Syracuse Electric Railroad, is approximately 30 feet of fine-grained, even bedded limestones. The lower portion is thin-bedded; the upper portion is somewhat more thickly bedded, although both may have laminated beds. Fossils are neither abundant nor well diversified. The Olney underlies the spillway cut at the east end of the dam and is exposed in road cuts on Interstate 81 south of the junction with Interstate 481.

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b. Elmwood Member - Overlies the Olney Member and consists of a series of dolostones, waterlimes, and limestones. The name is derived from Elmwood Park near Onondaga Hill in the southwestern part of Syracuse. It is subdivided into 3 units known as Elmwood A, B, and C in ascending order. Elmwood A is a drab yellowish brown, thin-bedded and mud cracked waterlime, barren of fossils and about 6 feet thick. Elmwood B is a finegrained, blue limestone, 3 feet thick, locally known as the "Diamond blue". This is overlain by 4 feet of waterlime, similar to Elmwood A, and known as Elmwood C. The Elmwood is exposed on the west side of the Onondaga Valley in Sweet's Quarry and Split Rock Quarry, in the roadcuts along Interstate 81 south of Interstate 481 junction, at Clark Reservation State Park, and in the vicinity of the Village of Jamesville. It underlies the spillway rock cut at the east edge of the flood control structure.

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- c. Clark Reservation Limestone Member Lies above the Elmwood and consists of a bed of fine-grained, dark blue, white-weathering limestone characterized by a diagonal fracture system. The name comes from Clark Reservation State Park, one mile west of Jamesville, where the type section is located. It is approximately 5 feet thick and in addition to the type locality it is exposed at Onondaga Hill on the west side of the valley and in the rock cuts along Interstate 81 mentioned above. It is also exposed at the base of the exposure on the east side of U. S. 11 just north of the Nedrow interchange to Interstate 81.
- <u>d</u>. Jamesville Limestone Member Lies between the Clark Reservation and the Onondaga Limestone. It is approximately 25 feet of fine-grained, dark blue, brittle,

B-2

limestone in thin, even beds at its type exposure in Clark Reservation State Park. It makes up the lower 12 feet of the rock exposure in the spillway cut at the east edge of the dam. The lowest 2.5 feet of the Jamesville exposed at this location is a stromatroporoid biostrome which is a layer composed of enormous numbers of stromatroporoids, perhaps a type of fossil sponge. A complete section of the Jamesville is exposed on U.S. 11 just north of the Nedrow interchange to Interstate 81, and in the rock cuts on Interstate 81 south of the junction of Interstate 481.

Tristates Group

B5. The Oriskany Sandstone, the only formation of this group exposed in central New York, is not well developed in the vicinity of the dam and Onondaga Valley in general. A few inches of reddishorange, sandy layers at the base of the Onondaga Limestone represent this unit in the spillway rock cut.

Middle Devonian - Erian Series

Onondaga Limestone

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B6. The Middle Devonian sequence begins with the Onondaga Limestone which attains a thickness of approximately 70 feet in the Syracuse area. It is a series of light bluish-gray semi-crystalline limestones in even continuous layers from 1 inch to 2 feet in thickness, separated by thin seams of dark, calcareous shale. Flattened nodules of dark blue or black chert, sometimes in continuous sheets or beds are unevenly distributed throughout the Onondaga but mainly in the upper part.

B7. The Onondaga is comprised of five members as shown on Table III and discussed below:

B-3

- a. Edgecliff Member The lowest member of the Onondaga is distinguished by its coarser, more pronounced, crystalline texture and abundant fossils. The base of the Edgecliff is a disconformity, meaning it was deposited over an erosion surface that had removed previously deposited rock strata. The member is approximately 15 ft. thick and is exposed in the spillway cut.
- b. Nedrow Member Just above the Edgecliff, is characterized by thin-bedded, shaley limestone, marked in the lower part by abundant snail fossils. It is a medium gray, very fine-grained, argillaceous limestone containing little chert unevenly distributed at the top. It is approximately 10 ft. thick, is gradational with the overlying Moorehouse Member, and rests with a sharp contact on the Edgecliff. Both the Nedrow and the overlying Moorehouse Member occur in the spillway cut.
- <u>c</u>. The Moorehouse Member Consists of 20 to 25 feet of medium gray, very fine grained limestone beds which are separated in many places by thin shaley partings ranging from 2 inches to 5 feet thick. Dark gray chert is common in this member but varies in amount. It is more common in the upper part of the member and occurs as more or less continuous strata.
- d. Tioga Bentonite Member Occurs in the upper part of the Onondaga as a volcanic ash layer, 6 to 9 inches thick. It separates the underlying Moorehouse Limestone Member from the overlying Seneca Limestone Member. Neither the Tioga nor the Seneca Member were observed in the spillway cut.
- e. Seneca Member Although lithologically similar to the upper Moorehouse, is characterized by a slightly different fossil assemblage. Also chert is less common in the upper part of the Seneca Member and the limestone beds are darker gray. This Member is approximately 25 feet thick.

B-4

B8. The Onondaga Limestone is exposed in the upper 50 feet of the spillway cut on the east side of the dam. Other exposures in the vicinity of the dam include Hemlock Creek northeast of Indian Village, abandoned quarries along Quarry Road on the Indian Reservation, road cuts on U.S. 11 at the Nedrow interchange to Interstate 81, road cuts on Interstate 81 north of the Nedrow exit, along the west branch of Onondaga Creek 3 miles northwest of South Onondaga and near Cedarvale. Additional exposures occur along the northern margin of the Allegheny Plateau from Jamesville through southern Syracuse and the famous quarries at Split Rock.

Hamilton Group

B9. The Hamilton Group in the Tully Valley consists of approximately 1,100 feet of mostly silty shales and siltstones lying above the Onondaga Limestone and below the Tully Limestone. Three thin carbonate beds serve to define the formational boundaries. The succession of units, their lithologies and thicknesses are shown on Table III.

Bl0. The Hamilton is the rock unit exposed in the dam vicinity except for the spillway cut at the dam itself. The Hamilton forms the sides of the Onondaga Valley north and south of the dam and also the floor of the valley south of the dam (Figures 3 and 4).

Bll. The Hamilton rocks in the Tully Valley dip S25[°]W at approximately 48 feet per mile, based upon the Centerfield Member as a datum. Formations included in the Hamilton Group are described below, in ascending order.

Marcellus Formation

Bl2. The Marcellus Formation is divided into four members. The lower three are composed of black shales and limestone, while the upper member is dominantly a transition from black to gray shale. The lower contact of this formation with the Onondaga Limestone is fairly

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sharp throughout the area and is overlain by the Mottville Member of the Skaneateles Formation with a sharp contact.

- a. Union Springs Member In the vicinity of the dam, the Union Springs Member is 10 feet thick. It is an alternation of black, sooty, carbonaceous, calcareous shale with thin bands and lenses of black bituminous limestone. The limestone beds range from less than 1 inch to 6 inches in thickness. Pyrite is dissemminated throughout the Union Springs, many fossils being replaced by it. The contact with the overlying Cherry Valley Member is sharp.
- b. Cherry Valley Member The Cherry Valley Limestone is a hackly, black, bituminous limestone with an orange-red stain that fills cavities, stains the surface, and coats some of the fossils. It is approximately 3 ft. thick.
- <u>c</u>. Chittenango Member This member consists of about 80 feet of unfossiliferous, jet black, non-calcareous, bituminous shale that contains abundant pyrite. It is extremely thin bedded and displays excellent rectangular jointing. Wherever it is exposed it breaks into paper-thin flakes that are usually stained a rusty color. Calcareous, unfossiliferous, septarian concretions up to 6 or more feet in diameter are common in this member, as are also smaller non-septarian concretions. These shales grade upward into the overlying Cardiff.
- <u>d</u>. Cardiff Member This member is named for exposures around the Village of Cardiff in the Tully Valley. It is a noncalcareous, dark gray to grayish black, medium-bedded shale at its base and grades upward into a slightly coarser shale of medium dark gray color. Toward the top, particularly the uppermost 4 feet, small, round, calcareous, nonseptarian, unfossiliferous concretions are present. They range from 1 inch to 8 inches in diameter. As defined here,

the Cardiff is approximately 180 feet thick and is overlain by the Mottville Member of the Skaneateles Formation.

B13. The Marcellus Formation is exposed in the immediate vicinity of the dam. It makes up the lower slopes of the valley north of the dam, and south of the dam to a point south of Rattlesnake Gulf (Bare Mountain Ravine) approximately 6 miles south of dam. It is also exposed on the lower slopes of the valley along the west branch of Onondaga Creek westward through Pumpkin Hollow to the Village of Marcellus, the type section. Streams cascading down the sides of the Tully Valley notch ravines in which the Marcellus Formation is exposed in the lower reaches. Peppermill Gulf and Lords Hill Ravine near Ironsides are two excellent exposures as well as ravines flowing down the east side of the valley south of the dam. Road cuts on Route 11A south of the dam expose Marcellus to a point a few miles south of Cardiff. The lower part of the formation, the Union Springs and Cherry Valley Limestone Members, are best exposed north of the dam. Two localities would be Hemlock Creek and a road cut on Interstate 81 just north of the Nedrow exit on the east side of the highway.

Skaneateles Formation

Bl4. In the Tully Valley, the Skaneateles Formation, with the exception of the Mottville Member, represents the intertonguing of a black shale facies to the west with a silty shale facies more fully developed to the east. Three tongues of dark shale, which are the eastern elongations of the Levanna Member of Cayuga I we, extend into the silty facies and thereby provide the basis for subdividing this formation, above the Mottville, into three members. Each member begins with a dark, relatively unfossiliferous shale that passes upward into coarser, lighter colored, more fossiliferous beds. These are in turn sharply overlain by the next dark shale of the member above.

<u>a</u>. Mottville Member - In the Tully Valley, the Mottville Member is about 5 to 6 feet thick. It is highly fossiliferous, calcareous, medium gray shale containing one or two thin limestone beds.

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- b. Delphi Station Member At most localities this member is about 125 feet thick, except at Rainbow Creek where 145 feet is measured. It is a dark gray shale in its lower portions containing abundant irregularly shaped, calcareous, non-septarian, unfossiliferous concretions from 1 to 5 inches in diameter. The dark shales pass upward into slightly calcareous siltstones abounding in bivalves (clams). The contact with the overlying Pompey is sharp.
- <u>c</u>. Pompey Member The Pompey Member is 60 feet thick, with the exception of the exposure in Bare Mountain Ravine, where it is about 114 feet thick, probably due to faulting. It is a dark gray to grayish black shale at its base grading upward to silty shale and finally siltstone at the top. Concretions occur more or less throughout the section, but are especially conspicuous in the upper 5 feet. Most of them are small, rounded, non-septarian and calcareous, and less than 4 or 5 inches in diameter. Although some fossils may be found on the periphery of the concretions, few are fossiliferous internally. The upper 20 feet of the Pompey contains abundant pyrite as nodules and disseminated particles
- d. Butternut Member The characteristic feature of the lower and middle Butternut is that it is a "paper shale", i.e., it splits into thin flakes, a distinct change from the underlying Pompey Member. Upward, this lithology is gradually replaced by silty shales interbedded with thin siltstone bands and lenses. The lenses vary in size from 3 feet long with maximum thickness of 1 inch, to 24 feet long with a maximum thickness of 6 inches. The lenses are cross-bedded at their bases only. Higher in the section, the silty shale gradually gives way to silt-stone and the uppermost 4 or 5 feet of the Butternut is

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a calcareous flaggy siltstone. The contact with the overlying Centerfield Member is transitional and difficult to place. The Member is approximately 135 feet thick.

B15. The Skaneateles formation is exposed in virtually all the ravines mentioned above for the Marcellus, only in the middle and upper reaches of these ravines, usually above the first waterfall, which is held up by the Mottville Member. Therefore, it forms the middle slopes of the valley sides in the vicinity of the dam.

Ludlowville Formation

Bl6. The Ludlowville Formation is the most fossiliferous formation in the Tully Valley. It consists of intertonguing silty shales and siltstones and, on the basis of these, can be divided, above the Centerfield Member, into four members.

- <u>a</u>. Centerfield Member In the Tully Valley area, the Centerfield Member is a coarse siltstone about 25 feet thick. The lower and upper 10 feet are flaggy but the middle portion is massive. Only the middle portion is highly calcareous and fossiliferous. The siltstone beds in the flaggy parts are about 1 inch thick and at some localities crossbedded. The contact with the overlying Otisco is sharp.
- b. Otisco Member This unit is a soft, thinly bedded, slightly calcareous, silty, medium-gray to medium-darkgray shale, which is interbedded toward the top with thin siltstone beds. The contact with the Ivy Point Member, about 165 feet above the Centerfield, is sharp. This member is especially interesting for two coral biostromes which have been given submember status by Oliver (1951). The lower, designated the Staghorn Point Submember, is about 7 feet thick in the Tully Valley. The Staghorm Point occurs about 50 feet above the top of the Centerfield and rests on a massive

calcareous siltstone platform about 3 feet thick. The upper biostrome, named the Joshua Submember, varies from 0 to 55 feet in thickness in the Tully Valley. It covers some 40 square miles, less extensive than the Staghorn Point Submember.

- <u>c</u>. Ivy Point Member This member is about 36 to 37 feet thick and is dominantly a flaggy, slightly calcareous, gray siltstone, that weathers to a distinct yellowish-brown color. Most of the siltstone beds are approximately 1 inch thick. Spheriodal, calcareous, nonseptarian, unfossiliferous concretions 3 to 18 inches in diameter are present in the lower 20 feet.
- <u>d</u>. Spafford Member The Spafford Member consists of 27 feet of thin-bedded, gray to dark gray shale and silty shale sharply overlying the coarse Ivy Point Member. Although the Spafford becomes coarser upward, the upper 3 feet are very fine-grained.
- e. Owasco Member This is a calcareous, massive siltstone bed 2 feet thick.

Bl7. The Ludlowville Formation makes up the upper slopes of the Tully Valley and is best exposed in Lords Hill Ravine, and in Emerson Gulf and Fellow Falls Ravines south of Bare Mountain Ravine.

Moscow Formation

B18. Exposures of the lower 70 feet of the Moscow Formation are difficult to find in the Tully Valley. Even where this portion crops out in a ravine, the exposure is poor and incomplete. Therefore, this interval could not be examined thoroughly and much of the remarks pertaining to the Portland Point and lower and middle Windom Members were derived from the study of a single section. The upper 80 feet of this formation is well exposed as it occurs in high waterfalls capped by the resistant Upper Devonian Tully Limestone.

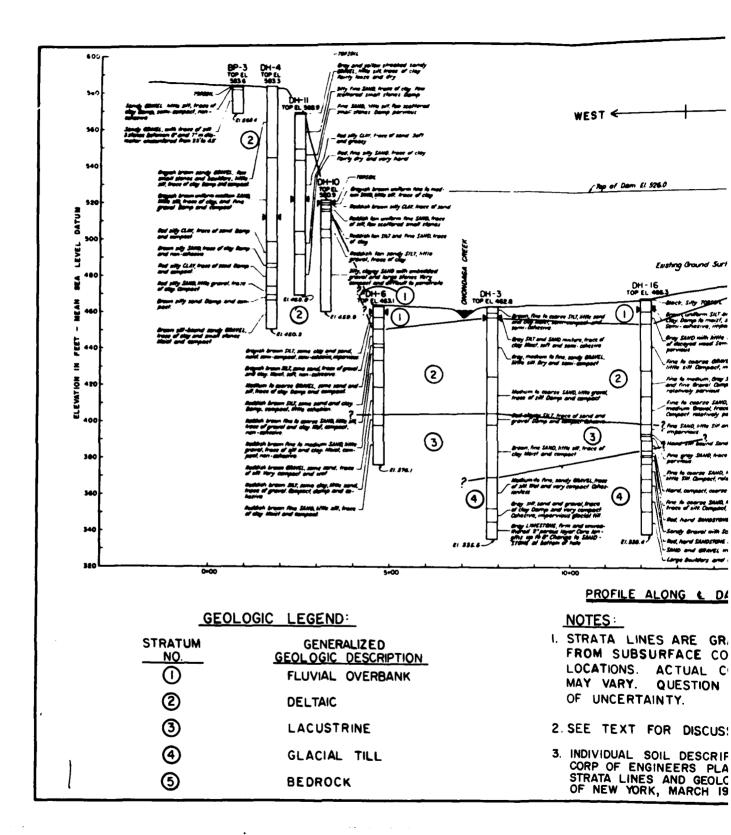
- <u>a</u>. Portland Point Member The upper and lower contacts of the Portland Point Member are sharp. A basal crinoidal, shelly limestone about 1 foot thick is succeeded by 11 feet of gray, silty shale, interbedded with thin, crinoidal, shelly limestone bands 2 inches thick or less and about 8 inches apart. The silty shales are not very fossiliferous.
- b. Windom Member The Windom Member is a thin-bedded, gray to medium-gray shale grading upward to medium-gray silty shale to a point 20 feet below the Tully. Here a sharp lithologic change to a medium-gray to dark gray or grayish black, non-calcareous, pyritiferous shale takes place and this is in turn sharply overlain by the Tully Limestone. This dark shale appears as a reddish orange zone beneath the Tully due to the weathering of the pyrite.

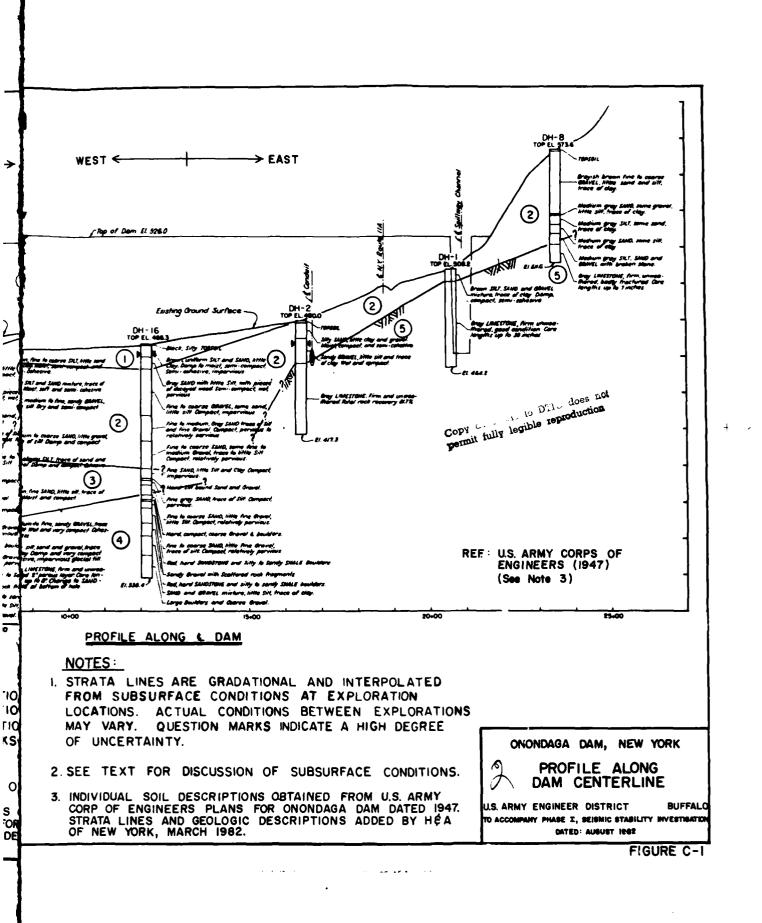
B19. The Windom is 163 feet thick in the Tully Valley. The Moscow Formation is exposed at the tops of the hills south of U.S. Route 20 such as Bare Mountain. It is well exposed in the vicinity of the Village of Tully in ravines and road cuts.

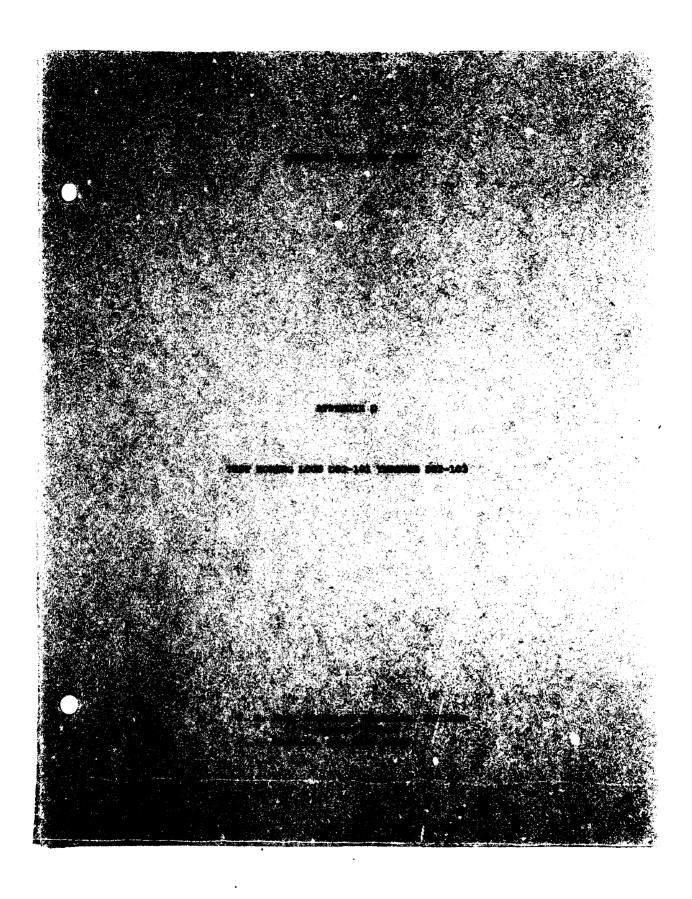
Upper Devonian - Senecan & Chautaquan Series

B20. The Upper Devonian sequence begins with the Tully Limestone exposed at the tops of the highest hills, at about El. 1,600 feet just south of the dam such as at Bare Mountain. However, the well developed and best exposed Upper Devonian does not occur until south of the village of Tully. The Upper Devonian is judged to be relatively far removed from the site and therefore no details of the stratigraphy will be given other than those shown on Table II and in Appendix A.



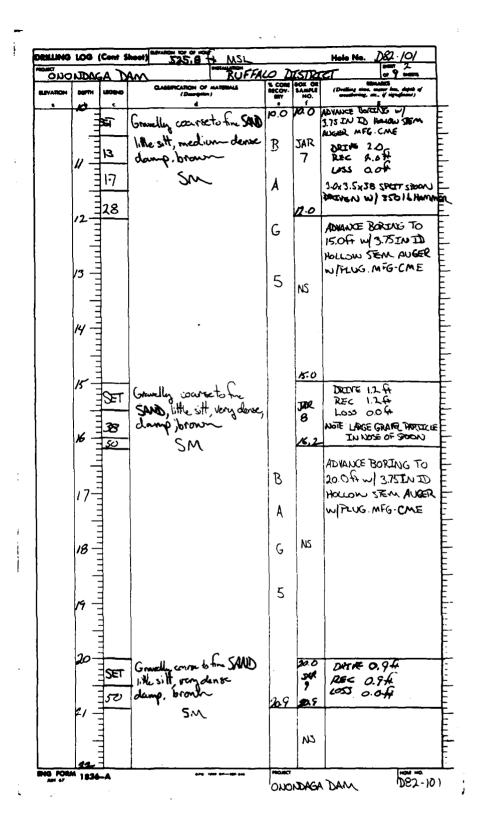


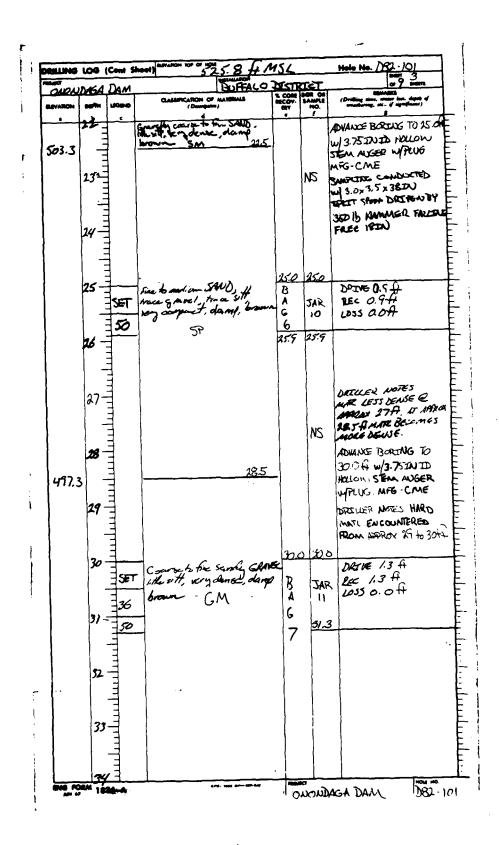




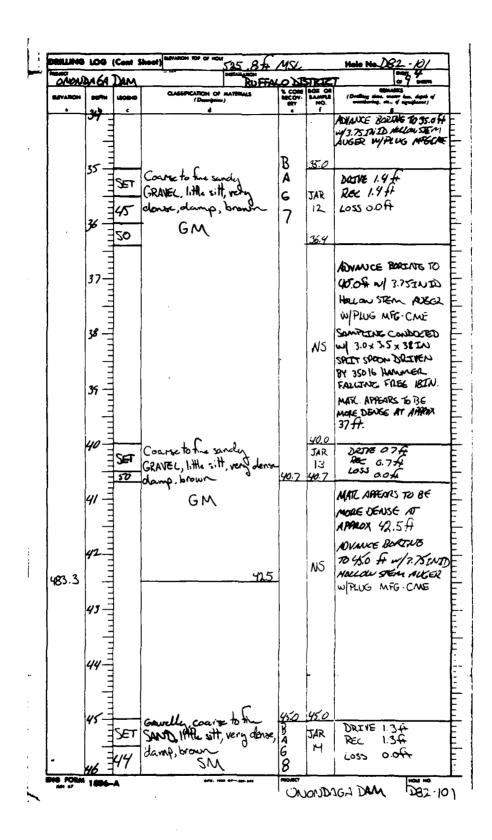
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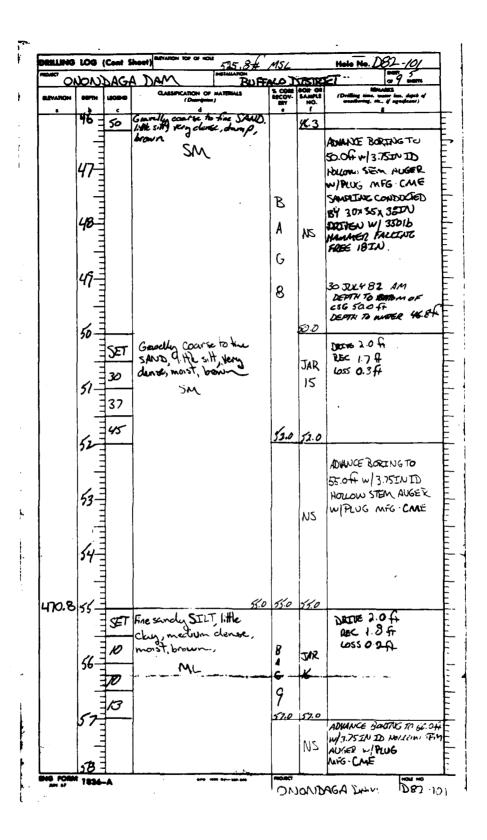
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450.0 13 = 21 Course to fine SAWD little 21 Course to fine SAWD little 31 Set 1.3 H 32 Set 5 M $760760 760$ 35 USS 1.3 H 449.8 76 $35 \text{ Course to fine secondary}$ 32 GRATEL trace sitt, B 33 Wey dense, wet, brown A 34 Wey dense, wet, brown A 35 32 78078.0 78.0 32 78078.0 79.0 78.0 79.0		•	. =	18		751		75.0	FALLING F	BE 3DTIN	F
449.8 20 449.8 20 $\frac{1}{32}$ $\frac{1}{32}$	950	87			c 1 f s	AND Little	4		Digne 3	,04	E
$\begin{array}{c ccccccccccccccccccccccccccccccccccc$			=1	21	Locarde Totine St	t dense	1	TAR	REC O	.7 ff	E
449.8 $76 = \frac{25}{59}$ SM 76.076.0 500 Barke 3.0x 3.5x 38" FIT STOND W) 38015 GRAVEL, truce s.tt. B Wenclense, wet, brown A GRAVEL, truce s.tt. B Wenclense, wet, brown A GRAVEL, truce s.tt. B Wenclense, wet, brown A GRAVEL truce s.tt. B HIMMARE FALLEND FIDE 11 LOSS 0.0 ft 11 LOSS 0.0 ft 12 Course to fue SAND ST Sitt, course to fue SAND ST SIT Sitt, course to fue SAND ST Sitt, course to fue SAND ST SIT Sitt, course to fue SAN ST ST SIT SITE ST SITE SITT SITE SITT SITE ST SITE SITT SITE SITE SITE SITE SITT SITE SITE SITE SITE SITT SITE SITE SITE SITE SITE SITT SITE SITE SITE SITE SITE SITT SITE SITE SITE SITE SITT SITE SITE SITE SITE SITE SITE SITT SITE SITE SITE SITE SITE SITE SITT SITE	1		-		grint house	,	1		1055 1.	34	F
449.8 78 SET Course to fine sandy GRAVEL, trace sitt, J 34 Key clanse, wet, brown GRAVEL, trace sitt, J 34 Key clanse, wet, brown G 37 JDTHE 2.044 LET SHOWN W/ JBD/G JDTHE 2.044 LET 2.044 LET 2.044 LET 2.044 LET 2.044 LET 3.0 KB/M SET Sitty course to fine SAND Many brown 53 Many, brown 53 SM 544 543 SM 544 544 544 545 Shop (Marsit, Kery dense, 81 498 SM 540 540 540 540 540 540 540 540			=1;	25		~					E
447.8 78 32 447.8 78 32 447.8 78 32 447.8 78 32 447.8 78 32 447.8 78 32 53 447.8 78 54 54 54 54 54 54 54 54 54 54	446	.87	(4/6.0	10.0	10-05- 3.0.	13.5138"	F
447.8 78 GRAVEL, trace s. It. 31 Very cleance, wet, brown 6 29 DIDLE 2.0 ft 10 IN. 11 2055 0.0 ft 11 2055 0.0 ft 11 2055 0.0 ft 12 2.0	' ''		- <u>-</u>	SET	Course to fine	- sandy			STATT SPEAK	1 1 2 2014	F
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144 SAND I ML SITH, very dense, 1445 8 80 SET Gravelly coarse to fine 144 SAND I ML sith, very dense, 144 SAND I ML sith, very dense, 144 SAND I ML sith, very dense, 147 SAND I ML sith, very dense, 148 SAND I ML sith, very dense, 148 SAND I ML sith, very dense, 149 SAND I ML sith, very dense, 149 SAND I ML sith, very dense, 140 SAND I ML sith ve	1447.	8 78	3-₹		<u> </u>		1.0.0	100	1 AUGUST 19	BilAM	Bin
445.880 SET Gravelly coarse to fine 144 SAND, 144 site years 144 SAN	1		-	ΞT	n 1	(, cont		{	DETTH TO BOTOM	- of CS6 78.0 Ft	HEL'
445.880 SET Gravelly coarse to fine 1445.880 SET Gravelly coarse to fine 1495.880 SET Gravelly coarse to fine 1400 SAFET HAMMER Fillows HO 1400 HO			-		Sitty course to	This Start			DETTH ID WATEL	R 47,044	
43 43 50 443 50 444 50 50 50 50 50 50 50 50 50 50			Ξ	57	little gover,	renjanik		SAR	and the second s	WARD ENGINE ME	H2
43 43 50 443 50 444 50 50 50 50 50 50 50 50 50 50	1	h	- 1		damp bon	\sim			BALLED 754 S	ET NW CSG 3%	
445.880 SET Gravelly coarse to fine BO. C DETRE 20 REC 0.5 COST 1.5 DETRE 19/0 × 201271/ IN SET Gravelly coarse to fine 48 SAND, I.M. sitt, very dense, 100 M SAFETY HAMMER FILLING HO JAR FREE 30IN SI DRIVE 2.074 REC 1.244 LOSS 0.874 E3 NO POINT 1000 MO	1	'		42	i `			1 30	INTO AT 77.71	DEVECASING TO	E
445.8 80 SET Gravelly course to fine 48 SAND, little sitt, very dense, 81	ł			75	S M		1		TREE 1 4 TN C	ANTING CLEANGU	4
445.8 80			ᅻ	1	500				DUT W/ THE CON	E ROLLEN BIT PRO	4 ″
943.080 SET Gravelly coarse to fine 948.5AND SAND 1.111 citt, very dense, 91 948.5AND 1.111 citt, very dense, 92 948.5AND 1.111 citt, very dense, 948.5AND 1.111 citt, very dense, 31 948.5AND Not, brown 31 948.5AND 83 948.5AND 1.111 citt, very dense, 949.5AND 1.111 citt, very dense,	lone	ماه	<u>,</u> =ť	71		80.0	<u> </u>	80.0	DET E 2 D KEC	0.5 6055 1.5	1
SET Gravelly coarse to fine Standato South Statute 4/8 SAND, 1.11 sitt, very dense, JAR 5/9 SM 5/9 SM 1005 0.874 240 SM 5/9 SM 1055 0.874 11 B 12 10 13 DRLVE 2.044 140 SA 159 SM 100 1.244 10 1.244 10 1.244 10 1.244 10 1.244	1222	· 08]				TET # 13/4 x 2	0x27% TN	1:
BI SAND (1. Mc sitt, very denic, JAR FREE 30TN) 579 Wot, brown 31 DRLVE 2.0 ft 579 SM Loss 0.8 ft B3 62 400 82.0 000 000 000 000 000 000 000 000 000	1		=	SET	Gow II. conce	tofine	1	1	KTANDARD SPLI	TSPOONWI	k
$\frac{61}{579} SM \qquad \qquad SI \qquad DELVE 2.0++ \\ Rec 1.2++ \\ Loss 0.8ft = 83 \\ Rec 1.2++ \\ Loss 0.8ft = 83 \\ Rec 1.2++ \\ R$	1	ł	-+		CALL CILL .H	men les			MON SAFETT N	AWNES FILLING	#0
$\frac{61}{579} SM \qquad \qquad SI \qquad DELVE 2.0++ \\ Rec 1.2++ \\ Loss 0.8ft = 83 \\ Rec 1.2++ \\ Loss 0.8ft = 83 \\ Rec 1.2++ \\ R$	1		. f	48 -	1.24ND (IMERI	, ver ig cuorse	5	1			E
100 MORM 1994 A		9	1		wet, brown			31			Ĕ
100 MORM 1994 A	1		1	59	C . A				REC	_	F.
HQ FORM 1994 A	[-		יאכ ן	•		1	Loss	0.84	E-3
NG FORM 1894 A THE ST THE ST PROPERTY HE HOLD THE HOLD THE ST PROPERTY AND A ST THE ST THE ST THE ST THE ST THE ST THE ST	1	(- E.	Lin	1					-	F
		k	<u>2</u>]'	-14	l			870			Ľ.
		XM 1	836-/	A		1000 07-000-300					

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83	SET Gravelly con SET Gravelly con SAND, little SAND,	rectofine sitt rem brown	•	51.0 51.0 32- 87.0	(Dulling the same in April 5 (Dulling the same in April 5 (Dulli	Ł
83	33 Gravelly can SanD, 1.4K, dense, wet, B SET Gravelly coar SAND, little si 28 clence, wet, 1 SAND, SM	eitt vern bronn	•	542.0 32	COSE & UNITED 65.14 DETHI CO WATED 65.14 DEDENCE ADVATIONS W/ CARE 75 TRUCKMALICI- ATTRUC SECTAL # 154115 ADMUCT EDERICS IN SOLUTIONS ANALYSIS JISON HALL FALLING ALTER AVIA MINISTING WITHOUT TO ADMITTING WITHOUT GALLAN ATT. DETUR 2.0	سأستهيس
83	33 Gravelly can SanD, 1.4K, dense, wet, B SET Gravelly coar SAND, little si 28 clence, wet, 1 SAND, SM	eitt vern bronn		JAQ 32	COSE & UNITED 65.14 DETHI CO WATED 65.14 DEDENCE ADVATIONS W/ CARE 75 TRUCKMALICI- ATTRUC SECTAL # 154115 ADMUCT EDERICS IN SOLUTIONS ANALYSIS JISON INTERNA FALLING ARTE ALLIN MARTING WATED FALCA TO LAMPLING FALCA TO	- - - -
83	33 Sand, 1944 dense, wet, SET Grandly cours SAND, little si 20 clence, wet, 1 SAN	eitt vern bronn		32	AND TS THICKMELTED ANTRAL SERIAL # 194115 ADMACE EXCENTS IN DELTITIONS NUCCESSING AUSSIO IN HAMMING FRALENCE ARTIC AUSSI MINIETING WITHLONG BALLER MINIETING WITHLONG BALLER MINIETING WITHLONG BALLER MINIETING WITHLONG BALLER	- - - -
83	B SET Grandly cour SAND, little si 28 clence, wet, 1 SM	to fine		32	ADMANCE RACEASE Nº SULVILISE MUCISENES AUSSO 16 PARAMA FALLING HERE ALLIN MERING WASHED PARAL TO MARTING W/TRILONG BALER TAT. DETTER 2.0	Ł
84 11 84	SAND, little si SAND, little si 28 clenne, wet, l SAN	n to fine			FRALING HASE 14100 TO TRUMES WAND PALOR TO THATTENG WITTLONG GALER	Ł
84	SAND, little si SAND, little si 28 clence, wet, l	to the		84.0	THE DETTE 2.0	47
84	SET Gravelly coar SAND, little si 18 denve, wet, 1 SM	to fine		84.0	10. DECKE 2.0	P''
8	SAND, little si 18 denic, wet, 1 SM	to fine			REC 0.8 LOSS 1.2	Ê
%	SAND, little si 18 denic, wet, 1 SM	It very			DRIVE 13/00 2.0x 27/23	AE.
					STANDARD SPEET SPEED	Fn
		otonin			W HO IS SAFETY HUMMEN	Ē
	16			33	DRIVE 2.0	E
86		:			REC 0.744	E.
86	14			86.0	655 1.3A	Ĕ
		24		00.0	C	╞
	ET Gravelly coar SAND, 9. the dense, wet,	sitt yerra			DRIVE 2.0 H REC 0.8 H	E
	8 dense, wet,	prown			LOSS 1.2 ft	29
87-	SM	•		JAR 34		E-
	5 300					E.
	~ ?>					Ē
1 195	30			<u>88.0</u>		Ē.
<u>ک</u> ا	ET Gravelly con	inseto time			DELVE 2.0 ft	F
	SAND, 1.ttle	sitt, rery	[JAR	REC O.» #	Ē₩
69-12	ET Gravelly con SAND, little 5 16 dense, wet, 16 SM	provv		35	653 1.5 A	E
	26 JM	-				F
				Ì		17
90 -	29			8.0		E
	SET Grandly con	arse to fine			DRIVE 1 Of	Ē
i -] -	- SAND Little	sitt, very	1		REC 0.5ft 1055 1.5ft	Æ
o, Ŧ.	34 dense, wet	prown		54R	NOTE : LANGE GRAVE C PIECE	sΕ
(P)	78 SM			36	OFTEN FRUND LODGED IN SPOON SAMPLES FROM	,Ę
-		`			78.0 TO 76.0 Ft	1
	29	92.0		720		F
433.8 72-	त			~~	· · · ·	+
	ET sitty convert	o fire SAND			DRIVE 2.04	E.
		renz denoe			REC 0.5 4	B
13-1-	- wet, brown			31R. 37	Loss 1.5 H	F
	<u> </u>	-				E
	7			a		Ē
ING FORM 1836-A	· · · · · · · · · · · · · · · · · · ·	1000 07100-040	PROJECT	<u> H.O</u>	HOM HO.	ŀ
APR 67		ł	ONX	ADAC	GA DAM D82-1) (C
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ation	DALA	25	m	······································	FAL	205	1021 09 mm	135
BEVATION		10000	CLASSIFICATION OF	NATERALS			Binshiti . (Drilling sime many fan Aget of manheing, an f simehener)	TOT I
	97 -	<u> </u>	ļe			940	Dare 13/8×2+171	F
·	<u> </u>	SET	Site course to	fire			COMMEN SET SERVICE	E
·]) =		SAND little gr	avel,		JAR	W/ 140/6 SAFET AND MORE	33
	8=	35	verydense. we	t, brown		38	IMMUCE BORENOUS ~/	E
	1° =	37				\sim	MCATING DETHEN N/ 30 & MANAGE FRICENT	F
	-		SM			l i	GEE 24IN	74
429.8	94 -	34		96.0		96.0	REF 24/IN DET = 2.0 ft REC 0.3 ft bass 1.7ft	E
-4.1.0		SET	Coursettine			[E
- {			GRAVEL, WHE	Han	(DEVE 20#	E 46
1	In E	14	wot, brown			39	RE 0.7#	E
	97-	A F				31	Loss 1.3 A	E
		25	GM		ł	ł		37
		17	}		}	00 0		Ĕ
	18-		10		}	98.0		ŧ-1
] =	SET	Coarse to fin GRAVEL, little dense, wet,	sandy	1	}	Drive 2.0 ft	33
			GRAVEL, 11th	L SIT		SAR	REC 0.8 fr	E
	199-	22	dense, wet,	brown		40	LOSS 1.2 F	E_
	'' Ξ	26	GM					En
	=		910					127
425.8		24		100.0		100.0		Ē
3. CAF	=	CET	S.H. I	fin SAND	1		~	F
] =	SET	Sitty course to	in second	}		DRIVE 2.0 ft	8)
		26	brown			SAR	Rec 0.6 #	Ĕ :
	101-		Sin			41	LOSS 1.4 F	F-
		23				l		Fac
[24				{		38
423.8	102-			102.0		62.0		E-
1	1 =	SET	Course to fine son GRAVEL, 1 He si	f.ch			DRIVE 1.34 Rec 044	E
			Very dense, w	it brown		SAR	Loss 0.9ft	Ers .
	h	25	G.M	-	ł	42	STOON REPUSAL 103.3-4	E_
422.8		100	APPROXEMATE TOP OF	Reck 103.3	N3.3	103.3	ADVAUCE CASTAR TO 103.24 W/35016 HAMMAR RULING FREE NETALUSED REI-	151
1.000	1 -		A ANTIN VITRESTON OLI	SEY MEMBER				E
			Limestone nodentali dense to finely cra flat lying few skylot watthankad, solid,	et lline	}	103.8	CLEAR PRESUMED OBSTRUCTION. SUTTINGS TUDICATED ROCK.	Ē
	104-		that lying ten stylot	dark aren		1	ATTEMPTED TO CORE N/ NX Devote THOSED BARRES (24 - D)	Ē
1			Frequent fractures in are mechanical in	THE COME U	,	PULL	Develo THE DESCRET (2 - D) BUT WAS UNABLE, DUE TO BENT CASENO AT BOTTOM OF DESCRET.	E
			caused by driling	. nature	}	1 1	ATTEMPTED TO CORE w/ BXM	Ē
	YOS-		। 19⊒ਹੱਥ `				NEW LOC REDARD CODE	E
			RAN 1.9ft REC 1.6ft	-			BARREN BROKE DIF. DRILLING TIME WAS 6 MIN /AT. NOLE WAR TRAMENATED AND COME BARREL	
400.4		•	UNACCONTAB	e cos 0.3fr		105.7	RECONSTRED AT A LATER DATE. The Aure C Row To Soltand	Ē
1	106 =		Bolton of BORIN				come 0324 were 65.26	E
AN 47	1836-	A	456. 100	6 64 108- 846	and and a		Dan D82:10/	
i					0,000			

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L PREJECT	INS LOS	-	DORTH CENTRAL	18. 884	BUFFALO MISTRICT OF AS WEETE							
	2 R1 1 - W	OLE	15379 85E MAD	- 14 191	1. T.S.				10E0			
		5. BEC	K 0000. 171000 1100	14. ELE	AL HUMBE VATION OF E HOLE VATION TO	5	14. 87- FALSE2					
			T FULLY PEDETRAZD		AL CORE I	MAPEC 1			1			
	PTN OF HOL		CLASSIFICATION OF NATE		M.	80X 01		NEW YOR	9			
ELEVATION 4	DEPTH LE	88NO	(Pressioned)		ALCON.	HO	Contraction of the second					
525.7		TC G G G G	course to the sus RAVEL, little sitt, surse, dry, bro GM	dig	<u>0.0</u>		Dec.	En Alect 1 135+ 38 w/350K E 182X 2.04 1.54	9/			
	E I	17			B	2.0	Loss	0.54				
	2 1 1 1 1 1 1 1 1 1 1 1 1		Course to fine s DANEL, little si lense, dry, bron	uly tt,	A G 1	SAR 2	REC	1.7 # 1.7 # 1.7 # 0.0 #				
	1		GM		į	37						
				4.0			NOWICE BORTON HALMU SEM AV	5 10 4.04 ml	Xind			
521.7	4 11 5		itte coarse to fine the gravels very brown SM	South Lenes	5.4	30R 3 3	DRIVE REC LOSS	1.4 1.4 0.0				
1,200					1	NS	ADMANCE BORT	NT. To 6.0+	iten			
ļ	ZE				60	60	WIRLIG MEG	-C.ME				
		ET (2 36	Snovelli coaroc SMD (Alle Sitt Lamp brown SM	to fire, denses	B A G 2	JAR 4	REC	2.0 f 2.0 f 0.0 f				
		2]	8.0	}					
	9 111	27	Francelle, costree SND, KHL eith Junse, damp, SM	te fre		JAR 5	DRIVO RAC LOSS	194				
	10.+	0				<u>99</u>	<u> </u>					
BIG PORM	1836 P	-	DITIONS ARE OBDLETE.		PROJECT		5A DAM	I MOL	210			

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DEILLING	100	(Cent 1	iheet) amount for an india 525.7 ft	WZI		Hole No. D82-102
Chan	N KA	<u>vin</u>	BUE		DIST	
BEVANON		10000	(Description)	T COME	SAMPLE NO.	(Drilling sime, water bas, depth of weathering, st., if significant)
	10	٠.	4	<u> </u>	10.0	NITE OSE
ļ		SET	Gravelly course to fine	B A G	JAR	Name 0.85 REC 0.85 Loss 0.077
] -	50	SAND, 9; the sitt, may	2	6	Loss 0.0ff
		<u> </u>	dense, damp. brown	10.8	10.8	NOWWE BORDOG TO 12.0 H
Ì	11-	3		ł		STR X TA BUDDI CEAN AUG
1		3	SM		INS	Sand Contract
		3		1		SELTSPORT IN / 350% MM
	12	1		12.0	12.0	FREENE FREE PETN
	10	5-		1		
		SET	Gravelle coarse to fire	}		
		37	Grandly coarse to time Sand, 1. The sitt, many	1	JAR	
	13-		dense damp, brown	1	7	^
1	[:	50		L	1	DRIVE 1.94
	-		24	B		REC 1.94
		50			13.9	LOSS 0.0 54
1	14-		Comelle coarse to.	A	13.9	
		Ser	Grovely coarse to fire SAND, little sitt	1		DRIVE 2 Off
		1	the shirt find shirt	G	L	REC 2.0ft
		48	very dense, doup,	Ŭ	542	Loss 0.0ft
	×-	1.10	brown		0	6AUG 1982
		48	SM	3	1	CANRA: DA SUIR HED TO
		37			1	139795
	16-	1		Ì	60	Notes A.44 Ber A.44
	1	SET	Gravelly coarse to The		398.9	SEEN REFESSION REPORT
		<u> </u>	SAND, "little citt,		K.Y	
1	1 :	7	very dense, drey,	1	1	ADVANCE BORENG TO
	17-	3	brown			18.07 W) 3.75 IN ID HOLOW STEM AUGER
		=	SM SM		NS	W PLUG MFG-CME
1	-	<u>-</u>			1	1
		7		1		
	18 -		Graveller coarse to the		8.0	Dorter 12 G
	-	Ser	Gravely course to the SAND, little sitt, very dense, damp, brown		JAR	DRIVE 1.3 ft REC 1.3 ft
		En E	SAND, MILL SHI, VETAL		10	REC 1.3++ Loss 0.0++
	19 -		and, any, and	1	1~	1031 0.011
		20	SN		13	10000 00 0 0000 0000
	-	1			1.	ADVANCE BORING TO 20.0 A W3.75 INTO
Ì.		1			NS	MILLOW STEM AUGUR W/
	20-		t . ^	· · · · · ·	20:0	Mintes CAS
F			Gravely coarse to the			DATAE TO H
	-		SAND, 1. the sitt, dance,			REC 1.3 ft LOSS 0.7 ft
	21-	1/2	SHOU, and suit ounder	1	SAR	Loss 0.7 ff
Ī		1	moist; brown	1	i u	
	_	21	SM			
		1 au				
503.7	122	34		120	22.0	1
AND POR	M 1834	HA .	676: 1000 07-300-540	ON	aui	IGA DAM DOZ-10

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UN6	100 (Cent S	hast) annen ter er nog	525.7	ff AC	sh	Hele No.T	<u>هم</u> ر	-102		
	AQA	GA	DAM	BUE	FALa		IRECT	-15			
ARCH		100040	GAMERICATION OF A	ATOMALS	% COM	SAMPLE	Hi (Deilling sine. maadaring, s	naditi maari kaa a Kaira	dapth of Second		
•	22	د .			<u>n.</u> 0	- f	Dunian matt	-	<u>.</u>	-1	
ר.נ	E	SET	9 ()	no		· · ·	Dunice Bar In ID Hould	N STE	A AUGER	ŧ i	
	-	, 1	Coarse to the	T anong	B	L	WITCHE SON	ic_ 2	5 6 199 117	FI	
•	3	24	GRAVEL, littles	T, dense,	AG	JAK	SPOON W!	35076	HAMING	Ē	
	23-		damp, brown		4	12	PALLING THE	EC ING.	710	EI	
		26	ĠM		11		REC 8	104		E	
			0.1				<i>2055</i> 0	.0 fi		FI	
7 .10	24	30		240	24.0	40				E	
	~ ′ =	SET	Medium to free Saus, St. gravel, and ism dense	, daugs how		335	DRIVE	1.74		E	
9.2			SM.	24.5	4		REC 1	:7f		EI	
		21	Grave ly coarse to fin	~, damp,		3AR 14	10350	ioff	•	E	
	25-		brown SM	•	{	רין				Fi	
o.g		30		25.4		25.4 Jar				E.	
-		50	Medium to fine SAUL very dense, clamp,), little sitt		15	Antiper Barrie	0 140	AL 4-12 78 1	F !	
S. 7	26-		SM S	46.0		NS	PHUCE BORDANS PHILOW STE PLUG MEG. G	MAUGO	x u/	EI	
11. (SET	Coarse to fine sender GA			26.0 Jar 16	DETAE	: 2.0	h fr	E	
19.2	_		GM	<u> </u>	26.5	26.5	REC	2.0	h (Fi	
	<u> </u>	3)	1.1.1	SAND.		-	Loss	0.0	s fi	E	
	27-		these sitts trace	-1.	B	302				E	
	Ξ	24	very danse, da	mp, bonn	۵	17				F	
	1 =	or	- Sp		G					Fi	
	12=	35	Nation to for	SAND	9	28.0				E I	
		চল			5	JAR	DRIVE	1.34		E	
	1 -	50	trace to the sitter great de	nse lamo	1	18	REC	1.3ft 0.0ft		F	
	ت مرا	47	boun	.,.,		0	Loss	0.01	•	E	
	(x-	50	5P		1	273	[ĒΙ	
	=	<u> </u>	4			<u>- 63</u>	ADVANCE BO	RING 1	10 30 OF	ŧ	
					Ì	NS	W/ HOLOW S	TEM /	UGER	E	
	30-	<u> </u>	Medum tof-	SAND		<u> </u>	N/PLUG MI			F.	
		SET	theme sitt terre	ament	}	30.0	Town	E 2.0	1.	F	
		<u> </u>	very durace then	, how		JAR	REC	2.(F	
		32	2"		30.8	b.8	. KCC Løss		5Å	E	1
44.9	31-	1	Course to have so	The have	B		ددس			E	l
		45	brown	, ₇		SAR 20				Ē	1
		44	GM		AG	í	i			ŧ.	1
	132-	<u> </u>		J		320	for the second			J÷.	i
		SET	Coarse to fine so GRAVEC, 1. He si dense, damp,	H	6	3AQ	DREN	e 1.3	ф.	E	1
		<u> </u>	GRAVEC, I. The SI	bioun		24	REC	1.7	s ti	E	
	33-	45	arenoc, oump,			1	Loss	6.0	oft	E	1
	-0-	50	GM		33.3	33.3			-	F	1
	_	1	1			1	ANHANC - BO	CAG NO	70 34.04	E	t
		1				INS	W/875:	LUG /	WG-CME	Έ.	•
il.7 • • • •	12/1 - * 1896	<u>.</u>		34.0	PROJECT	<u> </u>	448.0		HOLE NO.	_نل	1
NO1 67	1000	-			Or		Co Jan	C l	DOZ-K	っと	1

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-ON	anin	GA	Darn	BUFF	ALD	DIS	RICT	15
	-	utome	CLASSIFICATION OF	MATEMALS	1 COR	SOK OR SAMPLE NO.	(Drilling sus, and	13. 7 Ann. 4
100	34	د	d		37.0	1		
491.7	E	SET	1. 10.	• • • • • • • • •	B	340	DRIVE 1.3	s£
			Connects the	AND, IL	Ā	JAR.	REC 1.	34
1		56	Some gravel,	little sitt,	G	22		64
)	X -	<u> </u>	vendense, d	emp.	7			
1		Ś	Davor		35.3	35:3	And the Topologie	
1	7		SM			NS	ADVANCE BONEN a/ 3.75 IN SD AUGER +/ PL CO	~ <i>L</i>
	E,		-	36.0	04.0		AUGER #/ PLU	- ^
489.7	36-		Contor to fine &		36.0	36.0	DETVE 0 24	,
		SET	GRAVEL, little sit	t und have		JAR 23	DETVE USA REC 0.844 LOSS 0.044	
	E	50	GRAVEL, FREST	, von sone		36.8	1055 0.0++	
		<u>ب</u>	damp, brown Gun			0.0	ADWANCE BOATT	6
1	177	1	GAL	-	R		1 st we she can st	201
]				B	NS	MAGOST 1582 -	4~
	ΕI				A	ļ	DEPTH TO BOTH	100
	lm =					38.0	38.04	R
	38	5	Coarretofine		G		DRIVE 0.7	4
	1	SET	GRAVELS little 5	T, rely		JAR 24	REC 0.7 LOSS 0.0	Ē
	E	50	dense damp,	him		98.7	<u> </u>	
	20 7				8		10. 04 m/8.7	5
	39-		GM				Lange Star	<i></i>
						NS	FULIG - AIRG CI	116
		and a						
	Eaul					40.0	ł	
	40-	ст е	Coarse to fine	sandy		JIR	DAINE 0.54	
	=	SET		# 0		25 405	Rec 0.31	
	I I		GRAVEL, little dense, damps	sill rang			STOTIN REPUSAL	. 6
	In E		dense, damp,	brown."			185 BLOWS In <1	4
	4) =		GM'			NS	ANTALE DETMO	~
	E					63	NOR - NOPLU	6 (1.a
1								
	41-			. 1		(no	SAME AG MY S	e.
		SET	Coaros to time	sandy !		T IT	STOR MANAGE	<i>K</i> U
	E	Je /	GRANEL, little	sitt. yor		JAK	DELAE 1.01	ŧ
	EI		dense, damps	hand			KEC 0.3	••
	43-		the former		B	43.0		_
	E		GM		ע ו		ADVANCE BORIN 44 04 w1375	
	_])			.	NS	HULON STEM	
	E				A		W PLUG MFG	٠C
	44					44.0	NOTE AUGUTIN:	D.
	=	SET	Course to time	-sandy	~		To LHECK BIT, R	no. Eri
	-		CILANEL Little S	itt. very	6	_	BOTH DANCED	Ba
l l	1	41	dense, wit, b	on U	f		44.04 w/Howon AUSERS w/o PC	0 5
	45-1		·····	,			out increasing	TR:
-		38	GM	ĺ	8 1		ROLLER BIT AL	Sy
	-+			{		,	DRIVE 18	
	w F	50		j	, ,	<u>18</u>	10EC 1,3	
-	1836-/	ł				NS .	MALS BREDG TO	- 54

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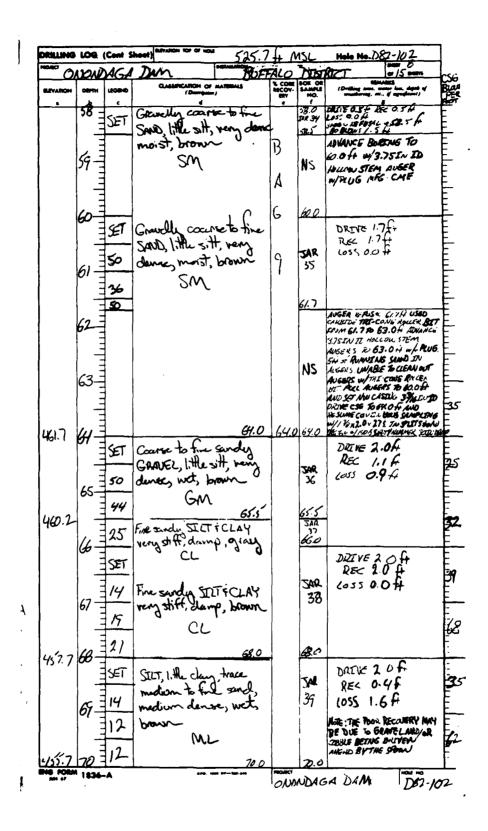
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	DE	I DA	M	525.	FAI	2015	TACT 0/5mm
BEVANON		10000	GASSIFICATION OF		% CORE	BOX OR	(Detiling sing, mater las, dath of
6	46	•	(Dampin	•)	NIV 4	NO. Í	unattering. etc., if significant)
	70 -	SEĨ	Course to fine	sandy		46.0	DRIVE 0.944
	-	201	GRAVEL, little	stt, dy		348	REC 0.5 FH
	Ξ	50	dorse, wet, bro	wn 188	46.5	46.9	1055 0.4 4
478.8	u7-		GN(<u></u>	<u>78./</u>	90.1	ANAWCE IS ARTAR TO 48.0
	Ξ						NJ 3.75 EN ED HALLAN PR
						NS	DATILER NOTES DEFFICI
				0	5	1.000	AUGERING
	48	┠	Gravelly coar SAND, 19the si dance, net,	- tatim	B	48.0	DRIVE 0.8
	Ξ	ET	Gravelly coar	H maa	A	JAR.	REC 0.3
		-	SAND, ITTLE S	II , reng	6	-	1055 0.0 ft
	116 -	30	danse, Net,	Diamic	9	48.8	ADMANCE BORING TO SO.C.
•	(19-		SW -		1	1	W/ 3.75 ENS) HOLLOW SEA
	=			•	1	NS	Mare 82 Am
			(1	1	DENTH TO BURGAN OF SEGUL
	ϕ^{-}			. 1 f		50.0	
	₽ =	ÆT	Gravelly coar SAND, 19the .	se is the		JAR	DRIVE 0,9 f
	_		SAND 14the si	TT, YONY		30	REC 0.9A
	=	50	dense, wet,	prov	1	<i>ź</i> 0.9	LOSS 0.0 ft
	分三		SM		1		ADWICE BORENG TO 52.00.
	<u> </u>		Sur	•	}	110	STEPTO MENT SEP . SEG.
	-					NS	MARUS AND CONTRACT SMELTAC M/ 25 + 3. 3x 36 SALT SPOUN DALISH 4/ 201 LANNER ALLENC FREE 18 5.
				-	}	510	Ammée ALLENG FEE 185
	52-		Gravelle coarr	. + C.		510 JAR 31	MEL OU
	=	50	Gravell court	E 10 TIME		12.4	(43 0.0
			Sania, little sit	meny	{	ĺ	Some ACTESN 252.4
	13=		dense. voct,				W/3.75TAD NOLL ON STEM
	53-		SM		ิจ	NS	ABER -/ AUG MIG
	_	1			B	ļ	cme
	1 _	1				-	ļ
	- 19			τC.	1	54.0	DATE - DG
	['] =	SET	Gowdy coar	e o The		3AR 32	DITVE 0844 REC 064
	-	50	SAND Mittle	sitt, very		1	1055 0 2A
	<i>"</i>	<u> </u>	Gendly coor SAND, 1, He dense, wet, 1	orown "	G	54.8	ADJUSTICE DATE IG TO JE
	<u> </u>	1	SM		-	1	12757 DALLES
	} =	1	2100	-		NS	ALE ALE ALE ALE
		}	}	-	0]	
	56 <u>-</u>	L		Jb	9	JAR JAR	
	Í	<u>شما</u>	Grovelly coa	ne to tire.		JAR	NEC 0.4 1
	_	ST .	Growelly coar SAND, little s.	tt verz	[33	LAS 0.2H
			dence, wet,	boon		1	ADVANCE BORING TO
	57-	ł			{	i -	58.04 ~ 3 MIN ID
	=	1	Sm Sm		ł	NS	HOLON STEM AUGE
	-	})		ļ	ſ	IN/PLUG
	1	1					NOTE: ALSERING DIFFICU
-	100	<u> </u>	i		Marti	1	L

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DRILLING	109	(Cent :	Sheet) Branch for or HO		5.21	MS	Hole No/)	22-102	
ONON	UDAC	A	Dam	BUF) DES	TREET	0/5 ums	22
BUVARON	-	1808-0	CASSINCATION O	ALATERIALS	L COM	SAMPLE	(Driffing sinte, a anathering, at	id ers ware ben, dapek of if agregicant (2015
455.7	70 -	and the second s	sitty fine SANC	trace 70.0		70.0	DREVE 2	off	È :
	-	SET	organics, woo	l'ishell		son.	REC 1	.o ft	ť
	Ξ	15	tamments, Sens	e, maist,		40	6055 0	.ofi	Ē
	7/-	!	grang SM						Ē
454.2	1 3	19	Contractofine	sandy 715	5	765			EX.
454.0		112	CEANEL to fine GEANEL , 1: 1818 S AND STATES , Stiff, Maist	1/21 CM- 7.7		1740 41 1747 -			B
453.7	72°-	14	fine synthe frame a	c sand, trace, "sankis, stiff 71.9	4	72.0	20016 0	-u	<u></u> ₽
1		SET		(4.9	}		DRIVE 2. REC 1.0	4	E
1		1.7	1			JAA	LOSS O.	WIED HITTH	F
	73-	17	1		}	112		JI. 91. Th	E-
1		28			}		MUING FRE	SOIL NAMMER	E.
}	-	1	-		}		148-2.0x 27	A IN STANARD	ER.
}	74-	18			1	74.0	KOK ARTIAL	HER ALCHINE	ter
	17		Cille course	to fine	{	1	DRIVE 2.	о Н	Ē
	-	SET	Sand tra	- a rend	1	ł	LOSS 0.	14	E.
		15	dence, net,	gray		SAR			E_
1	75-	15	1 .	2 4	(44	ĺ		E
Í	-	15	1 SM						195
		25)	76.0		76.0			È
449.7	76-	E	Gravelly course		1	1	Drive 2.	ofi .	F
		SET	SAND Tittle Si	tratense		1	LEC 1. LOSS 05	í í t	En
		15	wet a ray	, -)		LOSS OS	ζ++ ,	Ĕ^
	17-		1 °50		}	JAR 45			E
		1.15		-	ł	נד	2		4
		13							6 3
	78-			. t. f		78.0			╞
		SET	- Gravelly coar - SAND, little si	+ teace	ł	{	DAK	104	E.
			wood fing	ments		-m	REC C		EX.
1	79-	18	medium d	insk,"		342	LOSS /	,2#	<u> </u>
		16	wet. grav	й		40			1.
		-	T ŠM	-					78
445.7	80-	15	+	80.0	2	<u>æ.o</u>	<u> </u>		<u>+</u>
		र्झ	Sith coaree	tofine		JAR	DRIVE	2.0#	E.
		-	SAND little	grown),		47	REC (2.8 #	F 5
1	8)-	31	ren dense	- vet,			[032	1, Z. TT	E
ł	ſ,	33	bion			1	1		Ē.
	-	1	- SN	\sim		1			- [92
	82	12				81.0		M 082-	F

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ON	NOMO	AD		and the second	FFAL		TRICT	- 15 -		Dian Dea
REVATION	1	ucero	CLASSIFICATION OF I (Duription)		S COM	SAMPLE	(Drifting une. urbuduring,			COT :
	82	Gar	CII L	C		820	Drive	2.04		88
] _	SET	Sitty coarse T	5 The	} .		1ec	1.84		88
}		38	SAND, little of	found,	{	JAA	ددم	0.0		Ē
}	83-	30	very dense,	wet,	}	48				E
ł	-	1~	brown-							ЩĘ :
ļ		24	SW			94.0				Ē_
{	84-	SET	Sitty course	tofine			DRIVE REC	2.04		Ē.
	-	+	SAND, I.He z	iavel,	}		1055	0.944		EK.
	85-	30	dense, wet, 1	prov	}	JAR	CASING 33 WI350% N FREE 243	HINTO	DROMEN	Ē.
		19	Sm		}		4350% N FREE 245/	lunea ga N. Sampi	ING U	Ē.
{	-	}			1	} 1	CONDUCTED	WI-YEV	[.U¥11'2 (.S#=012	Ë,
439.7	86-	24		86.0		36.0	DUTION WI I	10 10 SAN	JOIN	Ę
(31.7		SET	Growelly cours	sto the				e 20 f	+	Ē,
		1.	Stap 1.# le sil	r, dense	1	JAR	REC.	1.1 f 0.9 f	t 4	1 47 E
	87-	12)	1		1	\$		• •		Ę
}		18	SM		ł					E,
1		15		^	}	000				E'
	88-	<u></u>	Gravelly coarse to	thank SAND		<u>88.0</u> 329	f		<u> </u>	Ē-
437.1		SET	little sill, dense.	88.5		5) 34R 52	Deary	E 201 1.24	+	E
}		22	Site mere un to the	m SAND		SAR 52	Los	50.8f	ł	Ē
436.7	81-	<u></u>	Gravelly coarge + SAND, 1. He sitt, M	Stine St.		87.0 Jac	1			E
}		16	SAND, little sitt, n	ery dense,	{	53				F
{		26	wet, brow ~ SM			90.0				Ë.
}	90-	SET		ьC	}				n D	E
	-	1	Gravella course	to the		502	REC	e 2.0' 1.4	<u>f</u>	25
}	91-	36	Gravelly course Savo, little sit	i read		54	1025		5 4	E
{		29	SM	-						E,
	-		300							F#
}	91-	jøe 1	4		}	20				Į.
[ইন্থ	{	. 0			12 AUG82 DEPTH TO B	AM.	4, 9 70	E.
{		129	Gravelly cour Sand Little sitt clanses wet, o	se to time		SAR		ASS2 41.	3	Ē
}	93-	1	SAND little sitt	, very		55	DRI	E 1.04	+	È-
{	{ _	<u>127</u>		2.29		1	Los	0.94	• 	Epe
431.7	h.,	PI	SM	94.0		940				F.
	M 1634		4		-	170	A DAM	HON	582-1	42
					ONP	NG G/	" UN	<i>۲</i> ۲	/84~/	v ~

ORILIN	0 LOG	(Cent :	Sheet) all ANON TOP OF HOLE	MS	6	Holo No. 282-102	
	NON	_	and the second second second	NOD	TSIR		င္ရ
REVATION		40000	CASSINCATION OF MATERIALS	T COM	BOX OR SAMPLE NO.	(Drilling size, orar bu, dapts of anadering, at., if significant)	-BL P
11217	94	<u> </u>		1 .	94.0		Ŕ
431.7	1	SET	Sitty course to fine SAND	1	JAR	DRIVE 2.0ft	E
}		<u></u> -	trace gravel, dense, wet		56	REC 1.14	8
1		21	gray to			Loss 0.9ft	E
430.7	95-		Medium to fine SAND, 11He sitt	4	<u>95.0</u> Jar 57		E
		25	make counce shad, dense	·	57		F
430.2	<u> </u>		wets gray SM 95.5	<u> </u>	95.5 JAR		Ę.
		25	Sitty fine SAND, trace	1	58		E
	96 -	╧	dense, wet, gray		96.0	BORE HOLE ADMANCED W/ NUL	÷
	1	SET	Sw 9		SAR	CASING 3 7/16 IN ID w	F.
]	Silve fine SAND, trace course	e	59	350 16 HAMMER FALLING	"B
1.	1 :	12)	Annalism sand denne	1	00-	FREE 24 IN SAMPING CONDUCTED W/ 13/8=20,21	ζŧ.
1422.7	-	<u>}~-</u>	wet, gray SM 970	-1	97.0	STANDARD SPLET SPOON	F
}		120	Correcto fine SAND, little	1	JAR	SAMPLER DRIVEN H/ HOID SAMPTY HAMMER FALLING	Ę,
		<u> ~~</u>	gravel, little sitt. dense,			DRIVE A OF REC 1.24	Ē
		120	wet, gray	1		Loss 0.8ft	Ē
	98-	<u> </u>	SM		9 <u>8.0</u>	(÷
1		SET	Course to Fine SAND, 1.+++e	1		DRIVE 2.0ft	E
		<u></u>				REC 1.17	Þ
1	0	20	sitt little gravel, donce,		JAR	Loss 0.9 ft	F
	199-		wet gray		61		F
ſ		18	Sm				E
							E
lune -		15	ko.c	.	100.0		F
425.7			Medium tofine SAND.	-	SAR		F
		SET	little sitt, dense, wet, going		62	DRIVE 2.05	199
		10-	SM	2		REC 1.44	F
424.7	1/10-	27	SIT 101.0	_	101.0	Loss 0.6ft	E
7^1/ /		29	Gravelly course to fine	7			F
1		101	SAND Attest very		522		F
ľ			dense, wet, gray	ł			Ē
1	101-	46	SM		102.0		-F-
423.4		7	Sity mealion to fine	_	101.0 JAN 101.3	Do- and	E
423.	1 -		SALD, medium dense, wet, grang Sn 102.6	6	3AR 65	DRITE 2.017	Ę
		71	Course to fine SAND,	7	1	REC 1.1tt	E
1	103	<u>+.'</u>	little sitt, little gravel.	1	JAR	Loss 0.9ft	F
1		17	medion dense, wet,		66		E
1		<u> </u>	dran E		}		3
i		15	3 1 SM		1		F
1	104-	<u>+</u>			104.0		÷
		SET	Course to five SAND, little sitt, little gravel		1	DRIVE 2.0ft	Ŀ
1	-	┼──			JAR	REC 1.4 Ft	E
1 ·		14	·SM	}	67		E
	los-	1	· ·		1	Loss 0.6 th	F
		81					E
ł			1			1	E
1	106	25			106.0		F
ING POR	^M 1836	-A	640 Had 07~10-145	ROACT		A DAM DB2-	
				ON	opun	DAMI UBI-	02
			• .				

$\begin{array}{c c c c c c c c c c c c c c c c c c c $		OND	GA .	CASSETCATION OF MATERIALS	FEAL	0 D1	STRUCT	OF AS SHEETS	PEA
$\begin{array}{c} 3 \text{ Set} & \text{Centre 1, bill sitt,} \\ 107 & 11 \\ 100 & 16 \\ 108 \\ 108 \\ 109 \\ 118 \\ 100 \\ 118 \\ 108 \\$	BLEVATION	DEPTH b	1400-10 c		BECOV-	1	weathering, a	waser las, depth of st., if squiftcast) 	
$\begin{array}{c ccccccccccccccccccccccccccccccccccc$		105	SET	Coarse to fine SAND,		106.0	DRIVE	2.0ff	Εİ
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$\begin{array}{c} 126 \\ 128 \\ 100 \\ 128 \\ 101 \\ 128 \\ 101 \\ 128 \\ 102 \\ 128 \\ 102 \\ 128 \\$		{	21	dense, wet, gray	ł	10			E
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117- GM NS NS 118.04 AND RESUMED SUMPLIANG. 117-			76762	dense wet, gray		14.0	USED TILL CO		-[""
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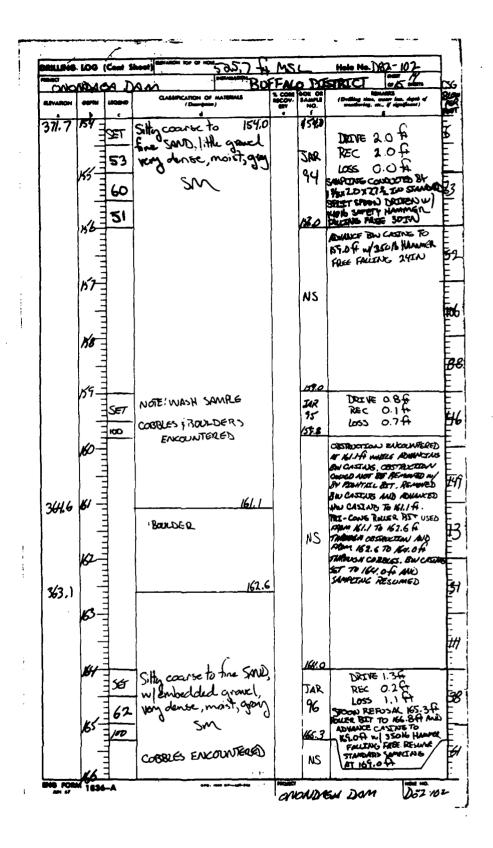
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)22	9	Eine sandu STI	Tteace 11		122.0 JAR	D /TT-1	E 2.0ft	
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		~ = 30	o dense, we	T, lect-born	\sim				713
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DRILLING	100	(Cont :		7 mm		Hole No.)	2.102	
ONON	DAG	A DA		WFFALO	DIST	RICT	als were	┛
BEVATION	1	LEGEND	CLASSIFICATION OF MATERIALS	% CORE RECOV- BEY	BOIL OR SAMPLE NO.	Dydling sins, or (Dydling sins, or (Dydling sins, or	ABCS were law, depth of . if segmificant)	
395.7	100		Fire sandy SILT, dense wet, rel-brown	50.0		DRIVE	2,04	
		SET	wet, red-brown		500	DEC 7	1.0++	
		121	mi		82	ions 0	.uff	
	131 -	<u>+</u>			0-			
		23				1		
		25		2.0	132.0			
373.7	132-		Sitty fine SAND, to fin SAND, little sitt, very dense, wet, red-box SM	e	152.0	DOTZ	105	
		SET	SAND 1. He with very			DRITE DEC LOSS	1.64	
		19	dense, wet rol-bon	m	SAU	Loss	9.4 क	
	133-		SM		83			
	1 -	34						
		39			134.0			
	134						20.	
	_	<u>= 7e1</u>	gilly fine SAND, very dense, wet, red-brow			DOIVE	2.07	
		27	dense, wet, red-brow	~	SAR	REC LOSS	1.5 ft	
	135-		SM		84	ROTAL ADVAN	SCED W/NW	
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		<u>76 E</u>				RDSH SOLDT CA CONDUCTED W/ I SAMPLER DATIE WALLOHT FALLI 30 IN	N BY 140 16 NG FREELY	
	136-	1			136.0	3010	- AF	
		SET	Sitty fire SOND, very clanac, wet red bron			DRIVE	6 =	
		<u>=</u> 25	dense, wet red bron	~	JAK 85	Rec 1 Loss O.	4 \$	
	137-	╡──	SM		రా			
	-	127						
		34			38.0			
	18-	<u> </u>	-		No.C	DRIVE 1	p.fr	
	-	SET	Sitte fine SAND dens	e		REC 2	.oft	
		23	web red. brown		JAR 82	6.	is ff	
	135-	-	$1 \leq M$		~			
	-	<u> 125</u>						
ļ		30			400			
	140-	-	-		100	0,2017 7	- ft	
	-	SET	Sthy fine SAVD, dense wet, red-brown w occassional sitt by	e Ha	Joe	DENE 2 Rec 1.	67	
		315	wet, red-brown w	ers	87	REC 1. 1035 0.	47	
	41-	-1	SM	נ]		
	-	<u>= 18</u>						
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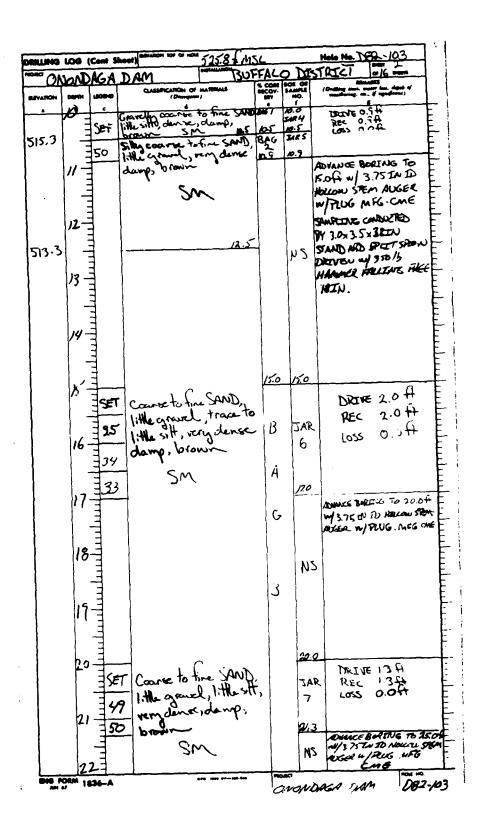


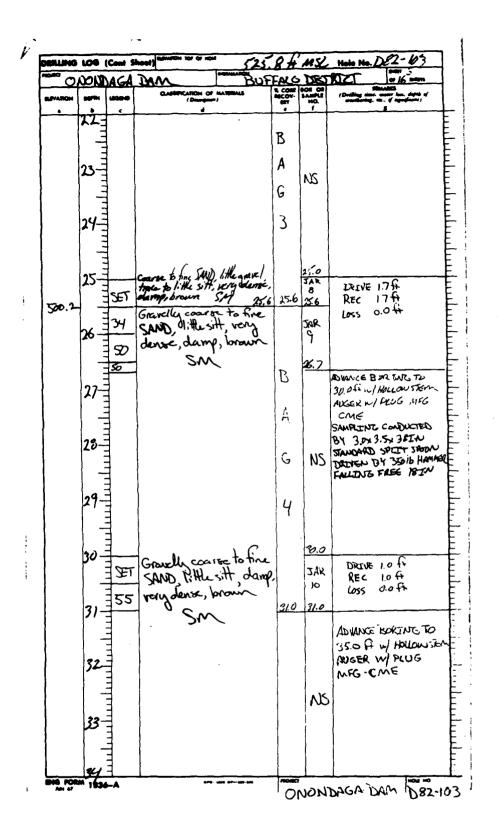
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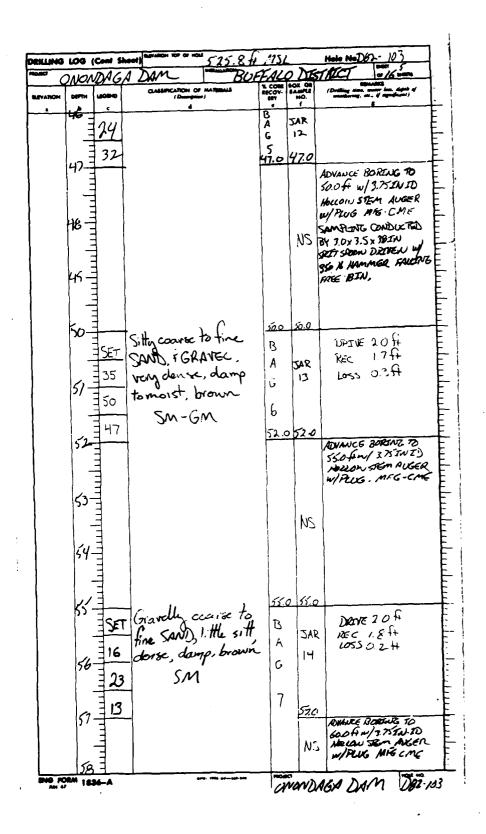
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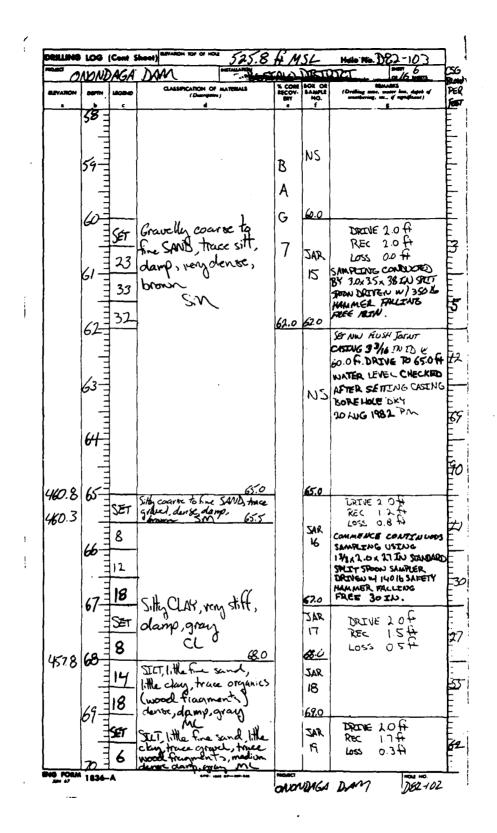
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GUTYARO	-	4005	CLASSIFICATION (Dum	QF MATERIALS	% COME RECOV- ERY	SAMPLE HO.	(Dritting sine, water in madering, etc., if an	n, dapet af professer)	
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424.3 101 25 course to fine SMUD little SET durine, ned brand SM 101.5 17 SET, little fine Sand, hard, 17 SET, little fine Sand, hard, 17 SET, little fine Sand, hard, 17 ML 17 ML 18 ML 19 ML 19 ML 19 ML 19 ML 19 ML 19 ML 19 SET durine, red brown 19 ML 19 ML 19 ML 19 SET sitty fine SAND. w/occassional SET sitty fine SAND. w/occassional 19 ML 19 SET sitty fine SAND. w/occassional 19 ML 19 SET sitty fine SAND. w/occassional 19 SET sitty fine SAND. w/occassional 10 SET sitty fine SAND. w/occassional 10 SET sitty fine SAND. w/occassional 10 SET sitty fine SAND. w/occassional 10 SET sitty fine SAND. w/occassional 10 SET sitty fine SAND. w/occassional 10 SET sitty fine SAND. w/occassional 10 SET sitty fine SAND. w/occassional 10 SET sitty fine SAND. w/occassional 10 SET sitty fine SAND. w/occassional 10 SET sitty fine SAND. w/occassional 10 SET sitty fine SAND. w/occassional 10 SET sitty fine SAND. w/occassional 10 SET sitty fine SAND. w/occassional 10 SET sitty fine SAND. w/occassional 10 SET sitty fine SAND. w/occassional 10 SET sitty fine SAND. w/occassiona	F								0.01]	100-	}
424.3 101 25 Convecto free JAND Little SET durine, red browley JANT 101.5 17 SET, Little free Sand, hard, 17 SET, Little free Sand, hard, 17 ML 17 ML 17 ML 17 ML 17 ML 17 ML 18 DET San SUD. w/scensional 18 ML 19 ML 19 ML 19 SET Jan SUD. w/scensional 19 ML 102 25 B3.0 19 ML 19 ML 19 ML 19 SET Sitty free SAND. w/scensional 19 ML 19 SET Sitty free SAND. w/scensional 19 ML 19 SET Sitty free SAND. w/scensional 19 SET Sitty free SAND. w/scensional 19 DET SCENS, wet, 19 DET SCENS, wet, 19 SET Sitty free Sand, treese 10 SET Sitty free Sand, treese 10 SET Sitty free Sand, treese 10 SET Sitty free Sand, treese 10 SET Sitty free Sand, treese 10 SET Sitty free Sand, treese 10 SET Sitty free Sand, treese 10 SET Sitty free Sand, treese 10 SET Sitty free Sand, treese 10 SET Sitty free Sand, treese 10 SET Sitty free Sand, treese 10 SET Sitty free Sand, treese 10 SET Sitty free Sand, treese 10 SET Sitty free Sand, treese 10 SET Sitty free Sand, treese 10 SET Sitty free Sand, treese 10 SET Sitty free Sand Streese 10 Set Streese 10 Set Streese 10 Set Streese 10 Set Streese 10 Set Streese 10 Set Streese 10 Set Streese 10 Set Streese 10 Set Streese 10 Set Streese 10 Set Streese 10 Set Streese 10 Set Streese 10 Set Streese 10 Set Streese 10 Set Streese 10 Set Streese 10 Set Strees	F						{			323		1
424.3 101-710 Convect to five SMUD little SET device, red brown SM 101.5 17 SET, little five Sand, hard, 17 wit, red brown 17 USS 0.7 ft 18 USS 0.7 ft 19 ML 12.8 104 25 B30 17 Sith, five SAND. w/occassional SET again STLT beens wet, 29 DRIVE 2 Off 18 USS 0.7 ft 19 DRIVE 2 Off 19 DRIVE 2 Off 10 DRIVE 2 DRIVE 2 DRIVE 10 DRIVE 2 DRIVE 2 DRIVE 10 DRIVE 2 DRIVE 10 DRIVE 2 DRIVE 10 DRIVE 2 DRIVE 10 DRIVE 2 DRIVE 10 DRIVE 2 DRIVE 10 DRIVE 2 DRIVE 10 DRIVE 2 DRIVE 10 DRIVE 2 DRIVE 10 DRIVE 2 DRIVE 10 DRIVE 2 DRIVE 10 DRIVE 10 DRIVE 2 DRIVE 10 DRIVE 1	E						1				-	{
424.3 SET durine, red brown SM 101.5 17 SIT, Hille fine samed, hard, 17 SIT, Hille fine samed, hard, 17 ML 17 ML 17 ML 17 ML 17 ML 17 ML 17 ML 18 LOSS 0.7 FF 18 LOSS 0.7 FF 18 LOSS 0.7 FF 19 ML 19 ML 10								HI MARY		325	1	{
424.3 17 Sett J. Hille fine send, hard, 17 wed, red. brown 19 ML 19 ML 128 103 25 1030 SET given STUD. w/occassional SET given STUD. w/occassional 19 ML 19 ML 1080 JAR 29 DRIVE 2 CH 19 DRIVE 2 CH 1080 REC 12FH 1080 REC 12FH 1080 REC 12FH 1080 REC 12FH 1080 REC 12FH 1080 REC 12FH 1080 NOTES BALLER BT 1080 NOTE	F	204	NE 1	NOT		37	t (- growed, w	sitt trace for		101	
42.8 102 17 ML 17 LOSS 0.7 FF 18 DELTE 2. CFF 19 DELTE 2. CFF 10 D	ľ	134			5	101.5	5	on on 10		1	2	inu
422.8 103 25 Billy fine SAND. w/occassional JAR 422.8 103 25 Billy fine SAND. w/occassional JAR JET chan JIIT reams, wet, 39 UNU REC 12Ft 421.8 104 25 M BALL BOOM 421.8 104 25 M BALL 420.3 JUT truce convect truce JAR LOSS 08FT ML 105 57 ML 420.3 JUT truce convector to fine 420.3 JUT t	ł	274			R	AZ	d,	. samel, he	SECT, little for	=	-1	17,7.
422.8 103 25 1030 1030 SET as an STIT second some of the second some		0.710	55 (Los:	1				wat, red by		In	1
422.8 103 2.5 103.0 1030 SET grans SILT serves with, 39 422.8 104 2.5 Silly fine SWD. w/occassional SET grans SILT serves with, 39 421.8 104 2.5 SM 104.0 104.0 REC 12.64 341 SILT, truce grand, truce 40 105 37 ML 105 0.84 105 5.5 ML 105 5.5 ML 105 5.5 ML 106 SET Sill (AV + truce and other 10 fine 10 0.55 ft) 106 SET Sill (AV + truce and other 10 fine 10 0.55 ft) 106 SET Sill (AV + truce and other 10 fine 10 0.55 ft) 105 5.5 JAR DET S 10.57 ft 105 5.5 JAR DET S					}	}	}					}
422.8 105 SET given SILT occurs, wet, 								-	1•\	-1	}	
422.8 105 SET given SILT verses, wet, 					30	D	0	Ŀ		725		l
421.8 104 25 Start to the same, there is the charge method by the same is the					ł		me	D. wlockas	Sith for se	ب مع لوم	8 10	422
421.8 104 25 SM 1040 DALO REC 12Ft 421.8 104 25 SM 1040 REC 12Ft 34 STIT truce gravel truce 105 ST ML NS Encourses balled at 420.3 Starter to true same structure 105 ST ML NS Encourses balled at 106 ST Bad free course brue 106 ST Course course brue 108 ST Bad free cour		- C					l	earny wet,	alas STLT	SET		
421.8 104 20 SM DHD The same frame 341 SILT three grownel traces 351 Silt three grownel traces 105 37 ML DSS 0.8 ft 105 37 ML DSS 0.8 ft 105 37 ML DSS 0.8 ft 105 ST ML DSS 0.8 ft 105 ST ML DSS 0.8 ft 106 SET Start trace and started 106 SET Start three control form 106 SET Start three control form 106 SET Start three on the form 106 SET Start three control form 106		70¥	ane y	DR	1	Ì		ed-brown	Lindens.		-	
421.8 107 34 SILT truce growel truce 508 105 37 ML 470.3 Sthe Clark true on the data 55.5 56.5		174	$\tilde{\sim}$		4.0	Ø	10	M	· · · · · ·	J 725		
420.3 Strange trace and better to the History of the total and total and total a		0.8++					. [snavel trace	STET toxe	· - · · ·	1.8 10	42
420.3 Sty Charles control bother CL 105.5 JAR DRIVE 15th 111.7 106 SET bould share control bother 41 Loss 0.3 Ft 100.5 March					10	14	ee.	nesand, t	coarse to	- 34		
470.3 Sthe CLAY truce on Institute 105:5 JAK UNY truce on Institute 106 SET Stand Annue course to fine 106 SET Stand, change, red. brown CL					v. 1		10-10	, and , in	Chez, have	T_		i l
470.3 Silve (LAN + Frice an indicated JAR DRIVE 1.5 Fr 106 Set Prove course to fine 41 REC 1.2 Ft 106 Set Prove change red brown (L 1050 0.3 Ft	-	Rales 81	NOTES	NULLEA	20	p.		L	1 ° N	<u>A31</u>	4	{ {
420.3 Sile (IN + truce an indicat JAR DRINE (:5) 106 SET Mare care to fine 41 REC 12ft 106 SET Mare hard change red brown (L 1050 0.3ft	40.7	Stores 1	Tered A	SHOULD B	50	1			}	۲ <u>ــــــــــــــــــــــــــــــــــــ</u>		1
106 - Read hards, comp, nor provent		54	IVE 1.5	DRI	TAR	- F	ا۔د ن	in an about the	Cil. (1.111 1	E	2	1.1
106 - Band half chill have been a set of the set of the		211	K 1.7	REC	41	`	e	court tot	Broset the	, 1 ~		} 44
	5.10							10.000, 7001- 2003	Lead her			

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· 0	NON	IDAG	ADAM	BOF	FALO	45	CET_	- de :	
	OEPTH	LEGEND	CASSIFICATION OF	MATERIALS	% CORE RECOV. ERY	BOK OR SAMPLE NO.	(Drilling in	REMARKS	-
.	ь	- C	f Duuripsin d	, 	etty e	но. (g, etc., if signific B	ant)
†1	06-		Sitty CLAY, trace	course to		JAR	DRIV	EISA	_
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}			rechbrown	, ,,	1		10SS	0.3 A	
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{I	107=			۱				VE 2,0 -	
		SET	Sitty CLAY, trace	course to			REC LOS		
[fine sandy har	d, damp,		342	BORING A	DANNCED	BY NO
		35	red-brown	•		42	CASTARS 7	1/6 INID	DODE
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ľ	=	SET	Sitz CLAY true	e coontr		JAR			
	-	<u>}</u>	Sitty CLAY true to fine sand, he	urel, damp.		43	REC LOSS	9.2 A	
)	-	125	red-brown	•)			-	
	110-	<u> </u>	CL			10.3	ļ		
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ł	-	1				ľ.,	ASTAN & ASTAN	ະ າ ກ ນ). <i>0</i> -6-	AND.
	m-Ξ	<u> </u>			1	JIII.O	PODNE COM	manus sam	A)//G
		SET	Silly CLAY, tro	ice course			DRIVE	52.0fz	
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]	-	43			1	1.15 A			
1	113				1	113.0	┝		
ľ	-	SET	Sitty CLAY, the	ice course	·		DRTIF	2.0 ft	
			to fine and	, hard,	1	1	REC	204	
	-	35	Sitty CLAY, the to fine sind damp, red-1	now	1	JAR	1-14	0.0ft	
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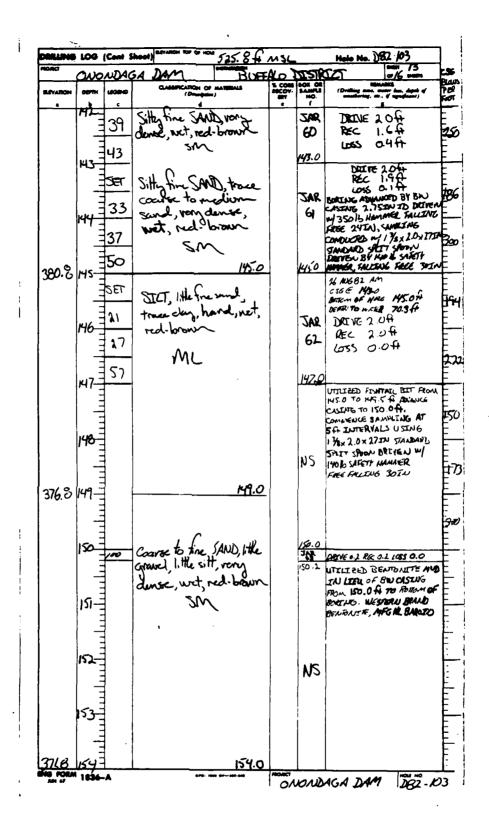
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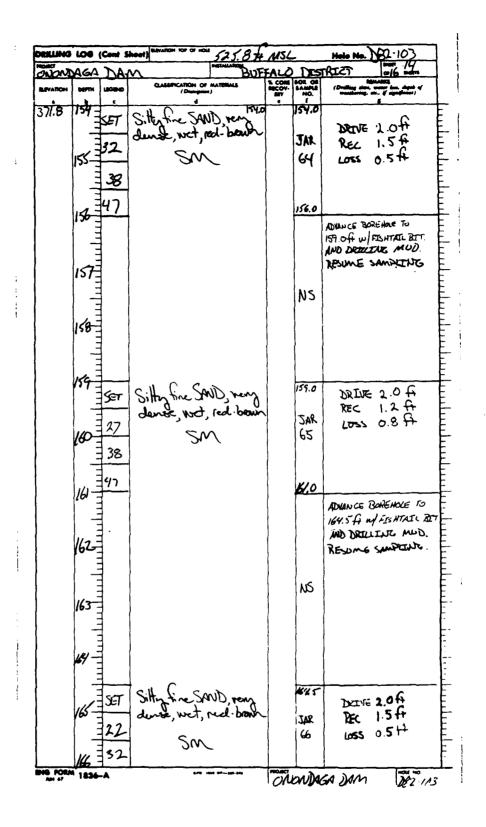
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141 OZ SET Sith fine SAVD, vory JAR DRIVE 2.0ft durse, wet, rub both JAR RE 1.6ft 149	1	-	En	ne r	١			CAN DITA)G		F
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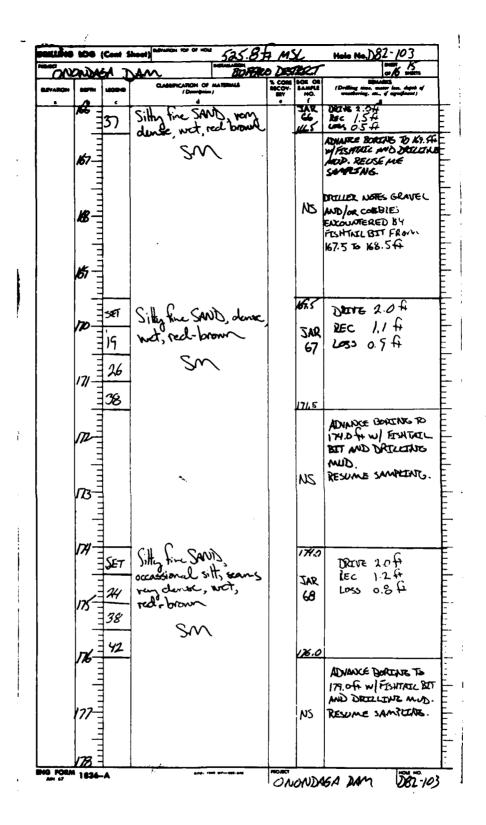
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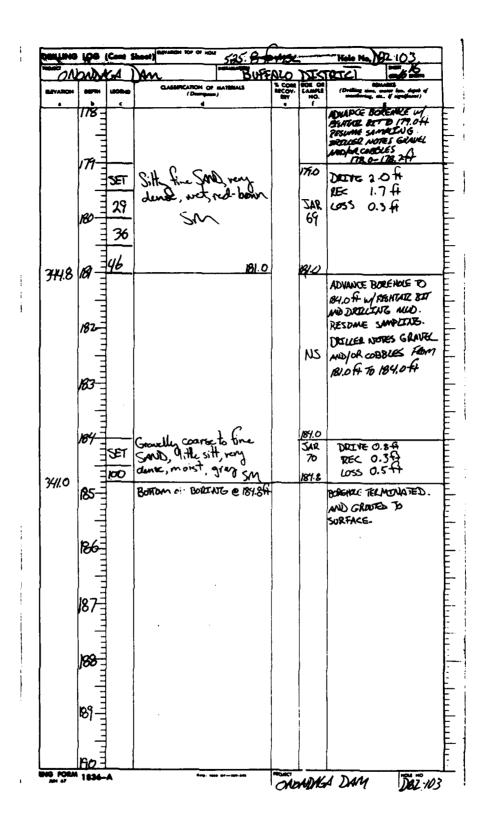
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	FILE NO 745003 SHEET 1 OF 1		UTHEN TESTS	Hydrometer H-1	Hydrometer H-2	Hydrometer H-3	Hydrometer H-4	Hydrometer H-5
	HEET	TION	ບັ					
	r w	CONSOLIDATION	MAX PAST PRESSURE TON/ SQ. FT					
		TEST	STRAIN %					
RESULTS	SU	UNCONFINED	COMPRESSIVE STRENGTH PSF		··			
- 1	12	-	WEIGHT I					
DRY T	c Inve Dam	BERG	PL PL	13.6	18.1	17.0	16.0	
LABORATORY TEST	Onondaga Dam Geologic Invi	ATTER BERG	F F	23.4	38.2 18.1	27.4 17.0	22.1 16.0	21.7 12.9
OF LAB	ic and G	NATURAL	CONTENT	19.6	19.3	20.6	23.1	13.6
SUMMARY	Seism		NO.	A-1	A-2	A-3	A-4	ม - ช
SUM			DEPTH (FEET)	60.0- 62.0	64.0 - 65.5	66.0- 68.0	69.0- 71.5	113.0
			DESCRIPTION	Brown, fine sandy SILT and CLAY, trace fine gravel, trace coarse to medium sandCL-	Brown, silty CLAY, little fine sandCL-	Brown, fine sandy SILT and CLAYCL-	Gray, fine sandy SIIT and CLAY, trace coarse to medium sandML-CL-	Red-brown, silty CLAY trace coarse to fine sandCL-
		SN (BOB	& SAMPLE NUMBER	D82-101/ S17	082_101/	D82-102/ S38	D82-103/ S19	D82-103/ 544

HALEY & ALDRICH, INC. CONSULTING SOIL ENGINEERS

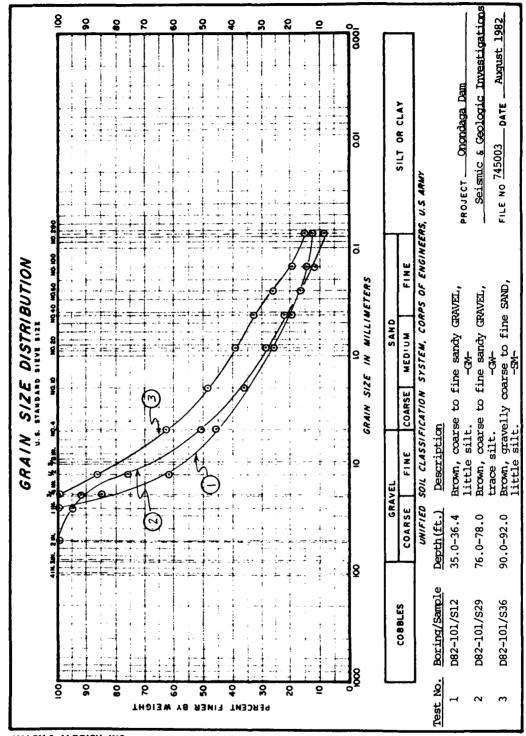
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FORM MBA FEB. 74 6

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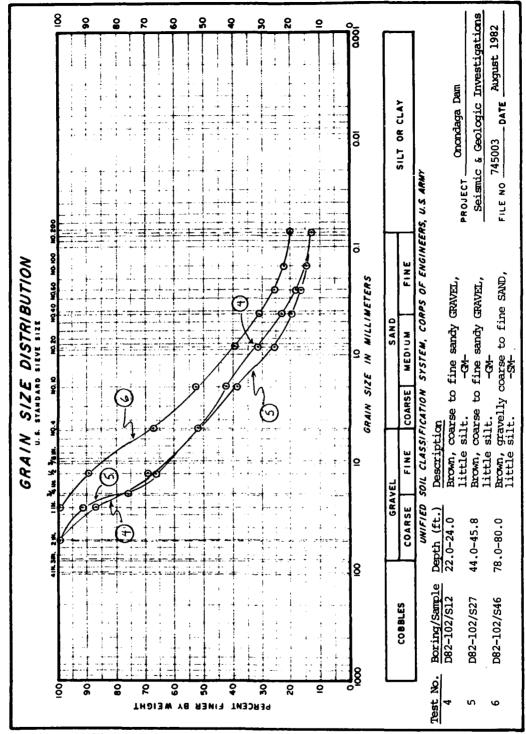
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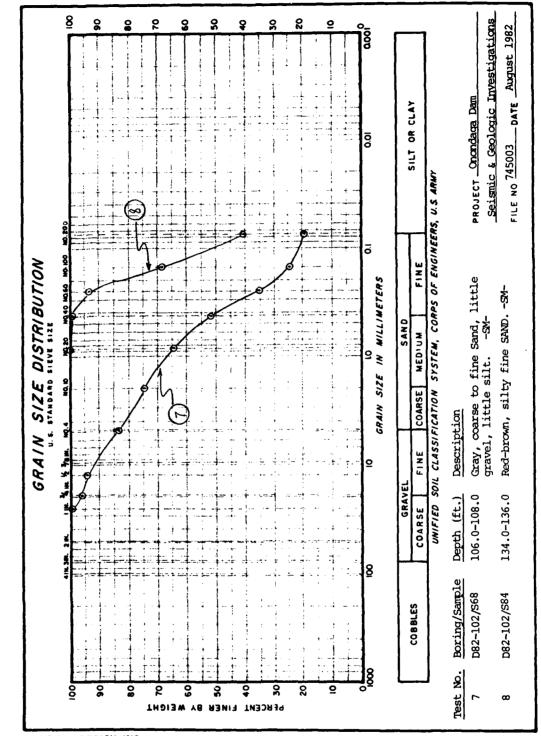


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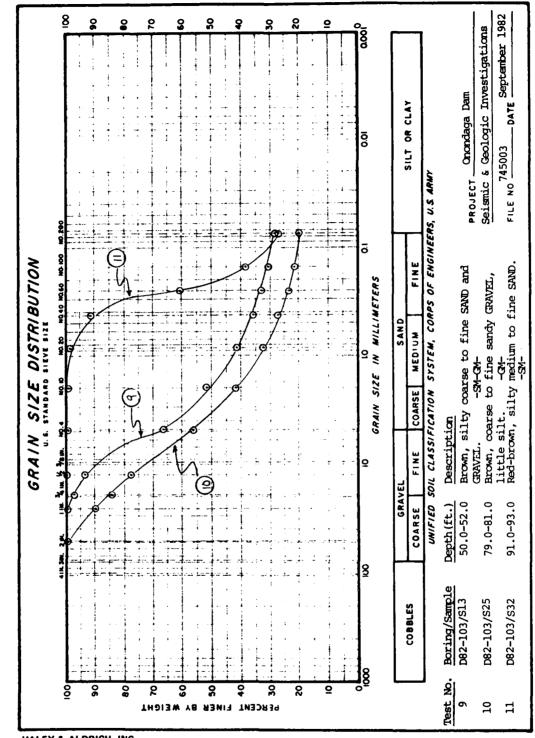


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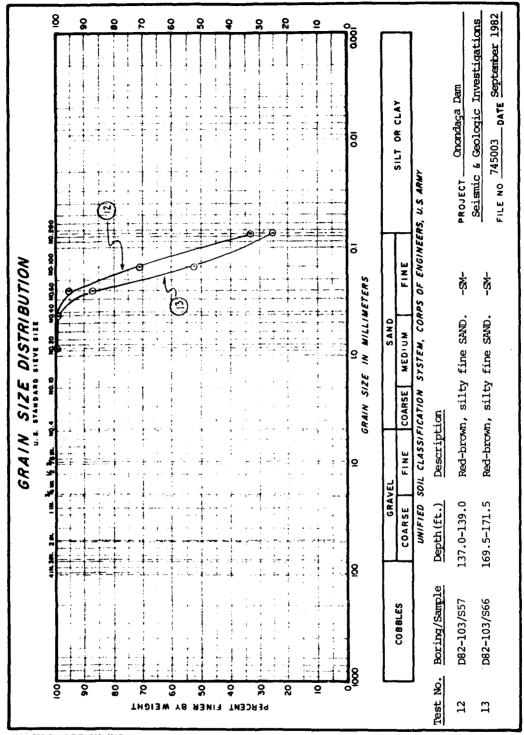


H&A FORM 40

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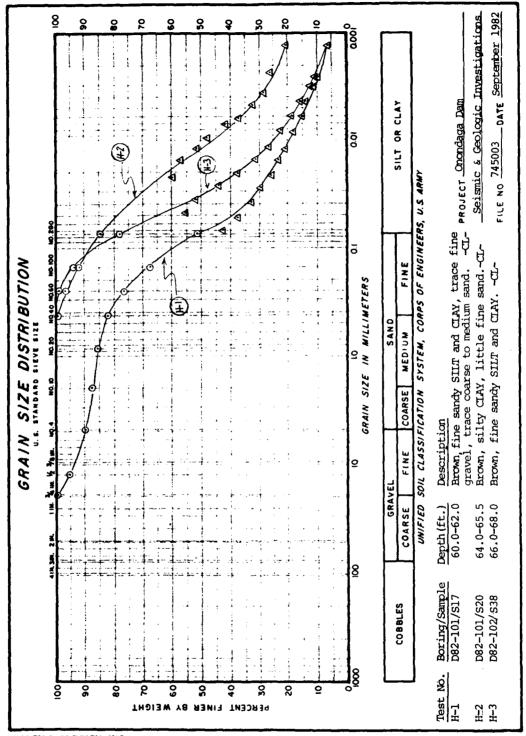


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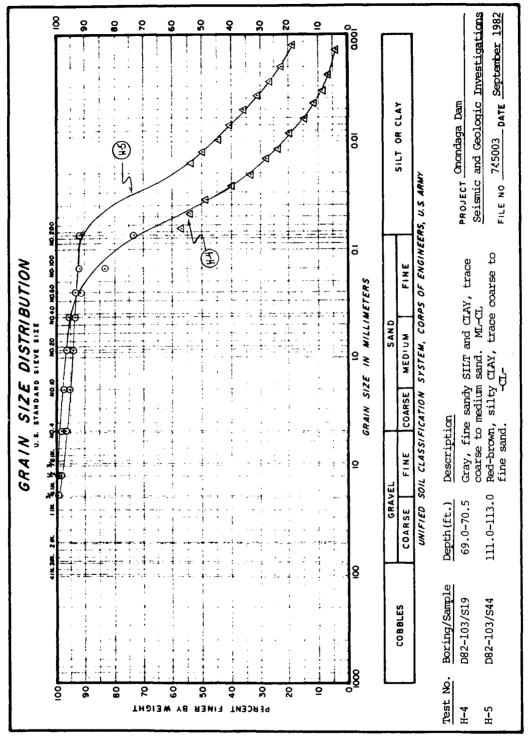


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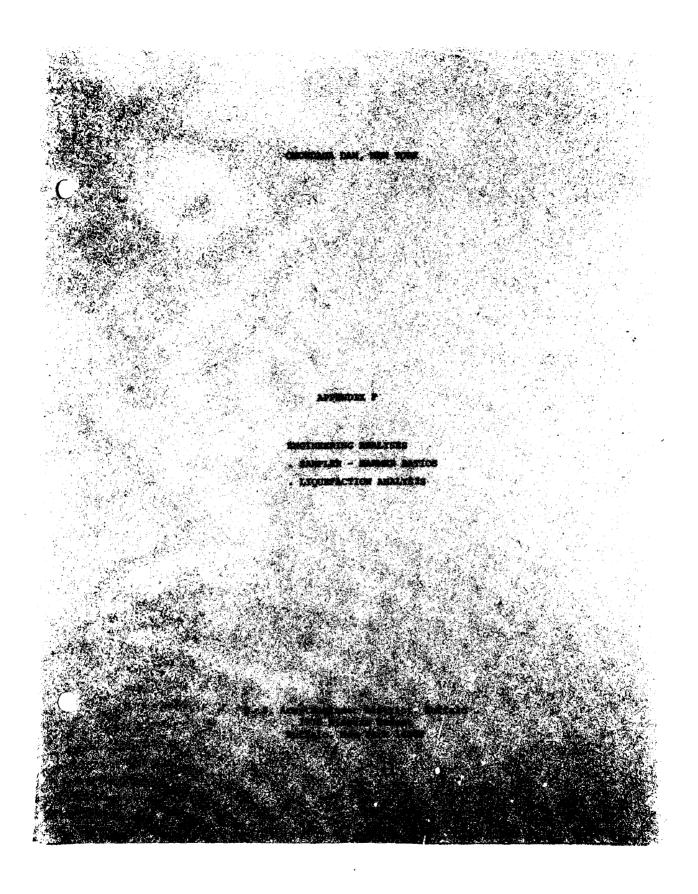
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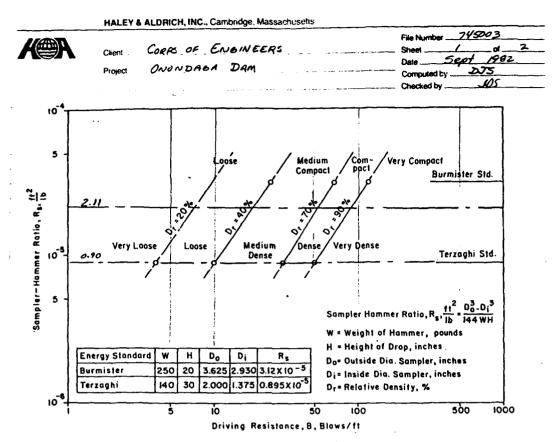
HBA FORM 40



HBA FORM 40

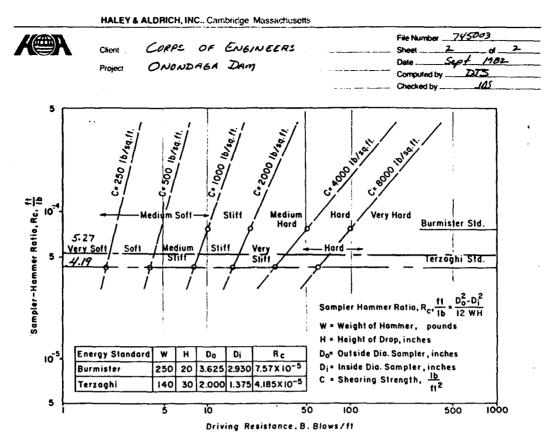


SAMPLER-HAMMER RATIOS



Sampler driving resistance vs compactness-cohesionless sands and silts.

SAMPLER	W (15)	H (in)	Do (in)	Di (in)	R3 (411/16)
Standard	140	30	2,000	/,375	0.90 * 10-5
Large	350	18	3.500	2.875	2.11 × 10-5



Sampler driving resistance vs consistency-cohesive soils.

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SAMPLER	W (15.)	H (in)	Do (in)	Di (in)	Rs (14./16)
Standard	140	30	2,000	1.375	4.19 * 10-5
Large	350	18	3,500	2.875	5-27 * 10 -5

LIQUEFACTION ANALYSIS

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i.

	HALEY & ALDRICH, INC., Cambridge, Ma	ssachusetts	
K	Chent	DAM	File Number 7450 - 03 Sheet SUMMARY of Date 9 SEP 81/L Computed by WAW Checked by 55

- SUMMARY OF LUQUEFACTION STUDIES FOR ONONDAGA DAM
- · SHEAR STRESS/CONFINING STRESS (C/J) RATIOS WITHE CALU-LATED AG: LEROM PROCEDURES BY SEED, 1976)

 $C/\overline{\sigma_{yo}} = 0.047 \pm 0.071$ BELOW TO E OF PAM $C/\overline{\sigma_{yo}} = 0.040 \pm 0.048$ FILLOW LEGT OF PAM

- FROM STUDIES BY SEED (1976) AND OTHERS, THE SPT-BLOW COUNT COORRECTED TO 1 THE OVERLAURDEN PRESSURE) NA MUST BE LESS THAN 4.3 TO 6.5 FOR LIQUEFACTION TO BE A CONCERN FOR SOLLS BELOW THE TOE OF THE DAM, AND NA MUST BE LESS THAN 3.6 TO 4.4 FUR SOLLS BELOW THE CREST OF THE DAM TO BE CONSURED FOTENTIALLY LIQUEFACE.
- DNE VALUE OF No EQUATE TO GOT WAS FOUND IN THE FOUND DATION SOLUS OF MORENSE DEC-101 AT 50 FT. BELOW THE DAM CREST MUSTER GOVEL, ALL OTHER VALUES OF No IN THE FOUNDATION SOLUS EQUALLED OF LE MEDED 9.
- SOUS AT THE DAM SITE ARE NOT SUSCEPTION TO LUDUEFACTION.

	C. O.		File Number 70	- 50-04
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	Project ONONDAC	TA DAM		DJSW
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and the second	FULLY DUFTO	man	OIT	
LRK CRK	EGGES PUE TO	BAR ON 19	26)	
	ή, τεροστη ή	BARON 197 pp. 682-	605	
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- L e			· - ·	
	> */H		5€ : ≈12,5'	
		H H	= 60'	
		7 1/2	20.2	
	× 1			
	1	0 5	DE SLOPEI = 1 X = 1	
7/H			× -	
	1	• A	SEVME MEC	0,3
			/	
	NEER OF EMBANK			
<i>y</i>) (DE OF 5MAANKIV	HEDI YHE (L		7
	JY/NGH	(8'H= 6001	30)= 7000 F	'SF)
(FT) Gerth	ULI GENTER	K=30°	TOE X-15°	TOF Q=30°
	00 0.8° (0.76)			
	, US -		0.0 (0.0)	0.0
90	57 0.78 (0.70)	0.61	0.07 (0.03)	0,10
	75 -		-	-
· · · · · · · · · · · · · · · · · · ·	,00 0.71 (0.62) ,25 -	0.5%	0.11 (0.12)	0.14
	·50 0.65 Lo.55) 0.45	0.14 (0.15)	0.17
165 2	.,75 -	-	•	
(S) 3	.00 0.61 (c.50)	0.49	0.16 (0.17)	<i>©.</i> 19
·	ATE A TO GUL AND	C		
2-PTM GE	AW The second se	SAM FROM	AF	
LOP OF	JNS /	PSF) (E)	DEV LESF	
<u> </u>		JE'ENTE,	55 chib.	65'EN13
Ċ	54 35	6,420(6190)	0	0
30	5005	5915(5700)	575	675
	-			
0 ف	44 35	5240 (5045)	8:53	1,015
90	4.0 a F	4650 (4400)	1,075	1070
-70	3935			1270
120	3575	4:230 (4070)	1,215 NT CE- 465 NAWLAL)	1,440
			NT CE- 465	.,

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HALEY & ALDRICH, INC., Cambridge, Massachusetts
$\begin{array}{c} \textbf{K} \hline $
* DYNAMIC SHEME STRESSES
C= 0.65 amox is to
at ground surface (g)
rd = depth reduction factor
ve total everyondan pressure
* REFER ATC-3 DOWMENT _ FIG. CI-8
· RATUODE "STREF GOLL" PETTIK TO "ROLK" PETTIK:
2.5 - 1.16 2.5 - 1.16
· NATURAL GLOUND SUK FACE ACCELERATION (PEAK HORIZONTAL) = (1.2)(0.05) = 0.06 g
amon = 0.06 g-
LNATURAL GUANNU) BELOW LIELT OF DANN
DENTALIET) TO COSE) TO CINE) TO COSE)

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1227 - AT (FT)	VO LPSF)	d	$\mathbb{C}(\mathbb{R}^{+})$	ra	て いやき
1	14.0	1.0	;		
5	12 C	1.0	مارژ		
10	17,00	0.20	50		
20	120 W C	5.93	4.7		
\$, O	ちゅうし	0.9	100		
4.0	1200	0,06	172		
い	12 32	0.4%	- 1	4.5	260
75	- 7.27	0.8	310	D. 2,	. 351
100 1	たいこ	0.9	4.10	0.0	410
:1-5 1	لآمة بالرفعا	6.0	5.0	, m	5-0
150 1	4,500	0.0	6-2	.,	

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HALEY & ALDRICH, INC., Cambridge, Massachusetts

03 7450 File Numb E 0. 2 ζ ONONDA4A DAM 4.0.0 1010 2010 10.01. 10.01. 10.01. 2 Ŧ REPRESENTS ACTUAL UNCORRECTED PENETRATION RESILTANCES FOR 31514 DIA, SPOON, 0.305 0.433 0.400 0.210 0.52**9** 0.506 0.498 0.512 0.400 0.518 0.431 0,441 ×2 0.052 (36) 20 0.024 (60) 32 0.053 39 * VALUES ADDED BY BUFFALD DIST BY CAL WIG183 ろろ 32 0.017 (20)12 θ Θ Ē 8 9 9 8 6 *> 0.04 6 0.050 0.046 0.056 0.068 6,04'L 2.067 0.07 240.0 0.071 100 706 c.045 0.045 0.047 18307 0.040 0,040 0.043 0.042 0.040 0.046 C. 04 0 4.4., 000 - 6,420 P.F., E. 6190 P.F. 000 000 000 3 \$ 197 6=2 ĩ 70 č n 12 v (2607 D 270 270 374 408 2692 tras 202 142 361 622 6326 3320 3400 -80 360 521 2502 2062 635 6600 1705 Rocht) 302 6423 2940 9370 0.700 9960 5629 2909 cruest 2.200 6150 2965 0/1/9 11/20 11/1/0 2200 305 610 250 205 dro LPGF) 6870 10,890 10.000 12,610 4505 10,235 0064 5400 5026 upert HI CAL 244-Ē - v 0+58V 10 0 10 **したい** によし 2/010 Honey Annas Money - e 0 0 -10 1000 -4-1-195 VOR 121 ₹ 20'd 61115 51.15 4W70 62.465 (600) 19-2 BORING 101 0 7,

