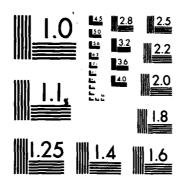
HD-A137 778 SPECTRAL ANALYSES OF HIGH-FREQUENCY PN SN PHASES FROM VERY SHALLOW FOCUS EARTHQUAKES(U) HAWAII INST OF GEOPHYSICS HOOLULU D A WALKER SEP 83 AFOSR-TR-84-0059 1/2 . UNCLASSIFIED F49620-81-C-0065 F/G 8/11 NL



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Air Force Office of Scientific Research

from

Daniel A. Walker

Hawaii Institute of Geophysics University of Hawaii Honolulu, Hawaii 96822

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The Wake Island Hydrophone Array has been successfully upgraded from				
a 3-channel slow-speed analog cassette system to an 11-channel computer controlled digital system for continuing research on Ocean P (Po) and				
Ocean S (So) phases, as well as normal, mantle-r	efracted P phases from			
underground nuclear explosions and earthquakes at	great distances.			

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Features of the upgraded system are a large dynamic range (events from at least 4.0 to 8.0 mb can be recorded without distortion), absolute timing accuracy to 1 msec, interchannel timing accuracy to 1 msec, digital recording for ease of processing, and a 40 Hz Nyquist frequency for recording frequencies actually observed in Po/So at large distances.

Software development for compression and efficient management of the data has been successfully completed with master tapes sent to DARPA's Center for Seismic Studies for use by interested investigators.

Research papers on Po/So phases have reported on the frequency content and propagation velocities at the Wake hydrophones and across a

1600 km long deep ocean hydrophone array. At a distance of about 18° (~2000 km) frequencies for Po/So are as high as 30 and 35 Hz, respectively;

at a distance of about 30° (~3,300 km), as high as 15 and 20 Hz, respectively. The travel time equations which successfully model all Northwestern Pacific shallow-focus Po/So first arrival data collected by the Hawaii Institute of Geophysics since 1963 at epicentral distances

greater than 12° are T = X/(7.96 \pm 0.05 km/sec) - (7.14 \pm 2.38 sec) and T = X/(4.57 \pm 0.04 km/sec) - (14.03 \pm 5.31 sec), respectively. Values for a frequency dependent Q are found to range from 625 \pm 469 at 2 Hz to 2106 \pm 473 at 13 Hz for Po and from 1401 \pm 296 at 5 Hz to 3953 \pm 863 at 15 Hz for So. This could alternatively be described as an average attenuation of -21.5 \pm 0.9 dB per 1000 km of travel path for both Po and So at all frequencies studied.

Regarding published research on normal-mantle refracted P phases, spectral comparisons for earthquakes and explosions have been completed and published. Expected differences between the spectral signatures of explosions and shallow focus earthquakes at great distances (the nuclear explosions being relatively stronger at high frequencies) are reported. The recording of a small explosion and improvement in its S/N ratio through some elementary enhancement techniques are also discussed.

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INTRODUCT ION

General Objectives

General objectives of our research in deep ocean seismology can be summarized as follows:

- (1) to determine the mechanism for the generation and propagation of Po/So phases; and
- (2) to utilize the low noise levels at high frequencies (i.e., > 2 Hz) of the deep oceans and the relatively large amplitudes observed in P at those high frequencies from events at great distances (i.e., >> 30°) to gain new insights into: (a) differences in spectra associated with source characeristics, and (b) the physical properties of the deep mantle.

In the past two years the principal means for achieving these objectives has been through the acquisition of data from hydrophones and ocean bottom seismometers located near Wake Island.

Specific Objectives

Specific objectives of an applied nature which we believe may be of interest to AFOSR are the following:

- (1) the relationship of Po/So spectra to focal depth and source parameters;
- (2) spectral comparisons of normal, mantle-refracted P phases at great distances from earthquakes and explosions;
- (3) estimates of the coherence of seismic phases recorded on the deep ocean floor; and
- (4) comparisons of S/N ratios for phases recorded on the Wake hydrophones, ocean bottom seismometers, and ocean sub-bottom seismometers.

The specific means by which AFOSR has contributed to our efforts to achieve these objectives has been major support for the upgrading of our Wake Island hydrophone array (partial support being derived from ACDA) and for the analysis of data acquired by: (a) the upgraded station, (b) the original three channel Wake system, and (c) other ocean bottom instrumentation in the Western Pacific.

Ressons for Interest

Reasons why these applied objectives may be of interest follow.

(1) For oceanic travel paths out to distances of about 3000 km, Po/So phases have signal-to-noise ratios generally at least ten times greater than the ratios of their respective normal, mantle-refracted P and S phases; and, in many instances, no P's or S's can be found in spite of

the presence of very strong Po's and So's. Therefore, the detection and discrimination of underground nuclear explosions at distances less than about 3000 km in an ocean environment requires an understanding of Po/So. Most important could be any possible relationship of Po/So spectra to focal depth and/or source parameters.

- (2) Normal, mantle-refracted P phases recorded at great distances (i.e., 60° to 90°, or approximately 6700 km to 10,000 km) have substantial amounts of energy at high frequencies (i.e., 2 Hz to 10 Hz). Such phases from earthquakes and explosions are well recorded in the low noise environment of the deep ocean. [Estimates of noise levels on the Wake bottom hydrophones in the 2 to 10 Hz range are comparable to those of the best continental sites.]
- (3) A major concern in the analysis of underground explosions by seismic methods is the reliability of yield estimates. Much of the scatter in yield estimates is due to inconsistencies between the waveforms from different stations, and even between the waveforms from elements of continental arrays with apertures the size of the Wake bottom array. These differences may be attributable to variations in the response of the continental crust underlying these arrays. Signals recorded in regions where the crust is very thin (i.e., under the deep ocean basins) may have higher coherencies—thereby providing a means for improved estimates of yield. The Wake array is an ideal tool for evaluating this hypothesis.
- (4) In recent years some members of the "detection and discrimination" community have demonstrated a substantial amount of interest in ocean sub-bottom instrumentation. The level of this interest may be best manifested in DARPA's commitment to develop a working sub-bottom system--the "Marine Seismic System" or MSS. Another sub-bottom instrument which has been developed (this at HIG through ONR sponsorship) is called the "Ocean Sub-Bottom Seismometer" (OSS). A reasonable presumption is that S/N ratios for ocean sub-bottom systems would be greater than for existing bottom systems (i.e., ocean bottom seismometers or ocean bottom hydrophones), resulting in a higher sensitivity (or lower detection threshold) for the sub-bottom instrument. A possible alternative to a sub-bottom instrument is an array of bottom instruments such as ocean bottom seismometers or ocean bottom hydrophones, with enough elements to make up the difference in S/N. Essential requirements to achieve an increase in S/N are incoherent noise and coherent signals in the frequency band of interest. Advantages of such an array over a sub-bottom instrument may be in cost, in ease of deployment, and in studies requiring an array. With the continued operation of the eleven element deep-ocean hydrophone array near Wake Island. these questions may be answered through further comparisons between signals recorded by the sub-bottom systems and those recorded by the Wake hydrophones.

PROGRESS

Upgrading of the Wake Island Hydrophone Array Recording System

Upgrading of the recording system for the Wake Hydrophone Array, from a 3channel slow-speed analog cassette system to an 11-channel, 16 bit, computercontrolled digital system, was accomplished by September, 1982. Some of the advantages of this new system are: (1) a large dynamic range to record, without distortion, events ranging from at least mb = 4.0 to mb = 8.0 (i.e., 16 bits = 96 dB); (2) absolute timing generally accurate to 1 msec (for ease in processing, no time correction needs to be applied to acheive this accuracy); (3) interchannel timing accurate to within 1 msec; (4) a digital recording format such that only a minimal amount of processing is necessary to convert the data to a widely useable format; (5) the capability to record all eleven available hydrophones; (6) the recording of frequencies at least as high as those already observed at Wake in Po and So (i.e., a 40 Hz Nyquist); (7) operation of the system so simple that it can be accomplished by the personnel at Wake who are untrained in computer hardware and software; (8) required servicing (i.e., changing tapes) no more than once per day; and (9) the capability of restarting automatically after power failures (which occur frequently at Wake).

The recording of data from all eleven available hydrophones produces four full-reel computer tapes per day requiring four tape drives to maintain a maximum of once per day servicing by an operator. Several problems which have occurred with the tape drives have made it necessary to record only eight of the eleven hydrophones (thus producing only three tapes per day) during most of the recording period. All other losses of data, including those caused by power-failures, amount to less than 3% of the total data collected.

Once the data on computer tape are received by HIG, an efficient compression and management of the data must be accomplished. Uncompressed, the data could amount to 1460 tapes per year, requiring a large space for storage and being inefficient to access for study, or to copy and transmit to other scientists. Software has been developed to compress this data to approximately 15% of its original volume and at the same time catalog all of the saved data for easy accessibility. Sections of data saved correspond to arrivals from events, and the raw data are also randomly sampled for future quantification of the temporal characteristics of ambient noise levels.

Compressed data tapes have been sent to the DARPA Center for Seismic Studies (CSS) for use by other scientists.

A complete description of the digital recording system at Wake, the data compression software used at HIG, and the tape format for data sent to CSS is contained in Appendix IX. A list of events for which the times of possible arrivals have been saved on the CSS compressed data tapes is contained in Appendix X.

Po/So Spectra vs. Focal Depth and Source Parameters

Although this was one of our original research objectives for AFOSR, no direct progress has been made in this area. Factors contributing to this lack of progress have been:

- (a) the re-direction of efforts towards topics more readily resolved. of more immediate interest, and/or of greater apparent importance;
- (b) anticipated improvements in instrumentation (i.e., the upgrading to a digital, rather than an analog system) which would greatly facilitate the data reduction required for this task; and
- (c) the small number of events with very shallow focal depths (<10 km).

Indirectly, a great deal of progress may have been made towards the achievement of this objective. [It is unlikely that relationships of spectra to focal depths and source parameters could be postulated in the absence of a generally acceptable model for the generation and propagation of Po/So phases.] Evidence of continuing progress in Po/So research may be found in: (a) recent publications ("Oceanic Pn/Sn: a qualitative explanation and reinterpretation of the T-phase" by D. Walker; <u>BIG Report 82-6</u>; and "Spectral characteristics of high-frequency Pn. Sn phases in the Western Pacific" by D. Walker, C. McCreery, and G. Sutton; op. cit.); (b) the formation of an Ocean P Alliance with the publication of an OPA newletter; and. (c) important advances apparent in analyses of data acquired by the ONR sponsored OBS Wake array experiment reported at the recent fall AGU meeting and now in draft form in preparation for publication. ("Po/So Phases: Propagation Velocity and Attenuation Across A 1600 km Long Deep Ocean Hydrophone Array" by D. A. Walker and C. S. McCreery).

We should note that with the upgraded system we have recorded two shallow focus events (15 km and 17 km depth) from the Marianas at distances of about 18° and one intraplate (and presumably very-shallow focus) event from approximately 12° to the west of Wake. The Po/So phases from these, and other, events will be useful in evaluating Po/So as a possible discriminant.

Spectral Comparison for Earthquakes and Explosions

Spectral comparisons between earthquakes and explosions have been completed and are the subject of a pubished paper ("Spectra of nuclear explosions, earthquakes, and noise from Wake Island bottom hydrophones" by C. McCreery, D. Walker, and G. Sutton, Geophys. Res. Lett., 10, 59-62, 1983.) The data examined show expected differences between the spectral signatures of explosions and shallow focus earthquakes at great distances—the nuclear explosions being relatively stronger at high-frequencies (or weaker at low-frequencies) than earthquakes. Also, a small explosion has been recorded with a significant S/N ratio on the Wake hydrophones. A unique aspect of hydrophone recordings is that the ocean surface reflection (recorded at a time after the main P corresponding to twice the water depth divided by the velocity of sound in water) can be used to increase S/N ratios.

Estimates of Coherence

Comprehensive estimates of coherence and possible improvements in yield estimates are a major topic for investigation in the coming year using data acquired from the expanded hydrophone array. Some preliminary estimates of coherence have already been made. These estimates are appended to this report.

Comparisons of S/N Ratios

In September of 1982, DARPA and ONR conducted tests (the "Downhole Experiment") in the Northwestern Pacific relating to sub-bottom seismic instrumentation. Successful deployments included OBS's and HIG's ocean sub-bottom seismometer (OSS). During the same time the expanded Wake system became operational. Preliminary comparisons of S/N ratios for the Wake system and the available downhole data have been made. They are appended to this report. Additional comparisons were (and are being) made after the OSS data was retrieved from its recording package on the ocean bottom in the summer of 1983. These comparisons will be the subject of a future report.

SUMMARY OF ACCOMPLISHMENTS

Major accomplishments supported in whole, or in part, by AFOSR since 3 March 1982 follow.

- (1) Successful upgrading of the Wake Hydrophone Array recording system from a three-channel analog cassette system to a sixteen-channel digital system. Some of the advantages of this new system are: (a) the ability to simultaneously record all of the active hydrophones in the array; (b) a 96 dB dynamic range; (c) millisecond accurate absolute and cross-channel timing; and (d) a digital format which allows immediate processing of the data.
- (2) Development of software to compress and manage the digital data collected at Wake. This data is compressed by saving intervals of data in which seismic phases of interest are known or suspected to be present, and by saving regular intervals of data for sampling the ambient noise levels. These compressed data are stored on magnetic tape files which are catalogued on paper and in computer files for easy reference.
- (3) Preparation of a report on items 1 and 2 above ("The Continuous Digital Data Collection System for the Wake Island Hydrophones" by C. S. McCreery).
- (4) Distribution of the Wake digital data through the DARPA Center for Seismic Studies (CSS). Compressed data tapes containing known or suspected seismic phases are being routinely sent to CSS for access by others who may have an interest in this data. Investigators at Rondout Associates Inc. have successfully accessed the data through CSS. A list of intervals, and the events and phases, which they represent, is contained in Appendix X. The event format of the data also is described in Appendix IX.

- (5) The observation and characterization of differences in spectral signature between P from explosions and P from shallow focus earthquakes recorded on the Wake bottom hydrophones. The observations reflect differences in source spectra between explosions and shallow focus earthquakes. For similar magnitudes, explosions had more energy at frequencies above 2.0 Hz and less energy at frequencies below 1.5 Hz.
- (6) The enhancement of S/N of an extremely small explosion by signal stacking to a level near that for perfectly coherent signals.
- (7) The publication of a report on items "4" and "5" above ("Spectra of nuclear explosions, earthquakes, and noise from Wake Island bottom hydrophones" by C. McCreery, G. Sutton, and D. Walker; op. cit.).
- (8) Continuing analyses of Po/So phases with additional spectrograms and discussions submitted for publication ("Spectral characteristics of high-frequency Pn. Sn phases in the Western Pacific" by D. Walker, C. McCreery, and G. Sutton; op. cit.).
- (9) The formation of an "Ocean P Alliance" with the publication of an OPA Newsletter to stimulate interest in, and research on, Po/So phases.
- (10) The evolution of a qualitative explanation for the propagation of Po/So phases consistent with many observations of these phases throughout the western, northern, and central Pacific.
- (11) The discovery of a probable relationship between Po/So and T phases.
- (12) Publication of a report discussing item "9" and "10" above ("Oceanic Pn/Sn: a qualitative explanation and reinterpretation of the T-phase" by D. Walker; <u>HIG Report 82-6</u>.)
- (13) Preliminary comparisons, with available data, of the Wake hydrophones to ocean bottom seismometers and the OSS. It should be noted that these are merely preliminary studies and in some cases the comparisons made are of an indirect nature. Nonetheless, it would appear that the Wake hydrophones may be comparable to the OSS in terms of S/N ratios, considering possible improvements through array processing. Direct comparisons to MSS are not possible since that system was not successfully deployed in the Pacific. Regarding the comparisons of Wake to ocean bottom seismographs. We have observed with data from an ONR sponsored OBS experiment near Wake Island in 1981 that the Wake hydrophones have higher S/N ratios than OBS's by about 10 dB over the range 1-20Hz.
- (14) The presentation of studies on the spectra of nuclear explosions and earthquakes, as well as spectral studies of Po/So. at the 1982 DARPA/AFOSR annual review.
- (15) Preparation of a report (Po/So Phases: Propagation Velocity and Attenuation Across a 1600 km Long Deep Ocean Hydrophone Array by D. Walker and C. McCreery) which quantifies the variations in velocity and spectra of Po/So phases across a long deep ocean array.

SIGNIFICANT TASKS REMAINING

- (1) Consistent with our efforts to determine the mechanism for the generation and propagation of Po/So phases, a comprehensive investigation of the relationship of the T phase to Po/So will be made. This study will include spectral studies of the entire T phase coda (i.e., forerunners as well as peak signals) and comparisons of these spectra to the spectra of Po/So phases.
- (2) Differences between the hydrophone/cable response (and possibly the lithosphere response) across the Wake array will be determined and removed from the data. This will be necessary to evaluate the stability of yield estimates from spectral or waveform data across .e array. Such an evaluation will be made.
- (3) Research on coherence across the Wake array will be completed with the coming year, and the significance of these studies in terms of possible improvements in S/N will be evaluated.
- (4) S/N comparisons to the Wake hydrophones will be made with the OSS data which was retrieved from its recording package on the ocean bottom this past summer.
- (5) Profiles of Po/So spectra as a function of focal depth will be made. If sufficient data is available for very shallow focus events, the possible significance of the relationships in terms of nuclear detection and/or discrimination will be evaluated.

The expanded Wake Island Hydrophone Array was successfully installed with funds provided primarily by AFOSR. However, since supplementary support of critical importance was provided by ACDA, all studies utilizing data acquired by the upgraded system should acknowledge both agencies. In the coming year ACDA funds will be used primarily on item #2 above, with the possibility that supplementary support for this task may be provided by AFOSR. On all other tasks AFOSR will be the principal supporter, with some supplementary funds provided by ACDA.

APPENDIX I

SPECTRA OF NUCLEAR EXPLOSIONS, EARTHQUAKES, AND NOISE FROM WAKE ISLAND BOTTOM HYDROPHONES

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Abstract. Spectral characteristics of P phases from 4 shallow focus earthquakes and 8 underground explosions, and of 52 samples of ocean bottom background noise, are examined by using tape recordings of ocean bottom hydrophones near Wake Island from July 1979 through March 1981. Significant differences are found between spectra of large shallow focus earthquakes and explosions (5.7 < mb < 6.3) observed at 61° to epicentral distance. For similar magnitudes. explosions were found to have less energy at frequencies below 1.5 Hz and more energy at frequencies above 2.0 Hz. Earthquakes were found to have a spectral slope of -28 dB/octave (relative to pressure) over the band 1 to 6 Hz. Explosions were found to have the same spectral slope over the band 2.2 to 6 Hz, but a different slope of -12 dB/octave over the band 1.1 to 2.2 Hz. High frequencies (>6 Hz) observed in the teleseismic P phases indicate high Q values for the deep mantle. Ambient noise levels on the ocean bottom near Wake are comparable to levels at the quietest continental sites for frequencies between 3 and 15 Hz. Also high levels of coherence (at least as high as 0.85) have been observed for P phases recorded on sensors with 40-km separation.

Introduction

In an earlier report (Walker, 1980), slowspeed paper recordings of hydrophones located near Wake Island were used in a study of P phases from underground nuclear explosions and earthquakes at comparable distances. That study was prompted by: (1) the work of Evernden (1977) and Evernden and Kohler (1979), which showed that P phases recorded in the 60° to 90° distance range from underground explosions were surprisingly rich in high-frequency energy (at least as high as 9 Hz); (2) the extreme sensitivity of the Wake hydrophones to high-frequency signals (Walker et al., 1978); and (3) the location of most known underground test sites in the 60° to 90° distance range from the Wake hydrophones. The major conclusion of Walker (1980) was that observable P phases were found for all Russian underground explosions with estimated yields in excess of 270 kilotons, whereas no such phases were found for earthquakes of comparable or greater magnitude at similar distances. Principal limitations of the investigation were that: (1) the slow-speed paper recordings were not suitable for detailed spectral analyses of either the recorded signals

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or the noise; and (2) the filtering was not optimized for the recording of distant earthquakes and explosions.

This report discusses the spectra of P phases from underground explosions and earthquakes as well as the spectra of ambient noise derived from recent tape recordings of the Wake Island ocean bottom hydrophone array. This array consists of six hydrophones on relatively flat ocean bottom near Wake at about 5.5-km depth. The hydrophones are located at the vertices and center of a pentagon roughly 40 km across and are cabled directly to Wake Island. Only three of the hydrophones could be recorded simultaneously on the recording system used. This system was a four-channel (three data, one time code) slowspeed cassette recorder, with automatic gainranging amplifiers following low-noise preamps connected to the differential outputs (via cabling) of the moving coil hydrophones on the ocean bottom. These recorded signals were used to compute absolute spectra of the seismic phases and background noise by the following steps: (1) digitization at 80 samples per sec after antialias filtering; (2) normalization of the automatic gain levels; (3) computation of contiguous 512-point FFT's (6.4 sec per FFT) and their corresponding power spectra; (4) averaging of those spectra over the time window of interest; (5) normalization of the spectral bandwidth from 0.156 Hz to 1.0 Hz; and (6) removal of the hydrophone/recording system/antialias response. The recording system and antialias response were determined in situ. The hydrophone response was taken from the Columbia University OBS Calibration Manual (Thanos, 1966), which describes the estimated response between 0.05 and 100 Hz of an equivalent hydrophone. More specific information on the Wake hydrophone/cable responses at these frequencies is not available because of the age of the array (about 20 years), its formerly classified status, and the original bandwidth of interest (>10 Hz) to those who installed the array.

Spectra of Underground Explosions and Natural Earthquakes

The earthquakes and explosions investigated in this study are listed in Table 1. These events were chosen because they all occurred within 60° to 90° epicentral distance, were shallow focus, had large signal/noise ratios, and did not exceed the dynamic range of the recording system. Figure 1 shows the pressure spectra of some of these events, as well as composite pressure spectra for the earthquake and explosion groups. For purposes of comparison, pressure and vertical

Table 1. Description of Events Used in Figure 1

No.	Date	Location	Distance (degrees)	h)دلا (kon)	Magnitude (mb)	Туре	Number of Hydrophones
1	07/24/79	S. of Java	65.7	31	6.3	Earthquake	3
2	08/04/79	E. Kazakh	73.2	Ó	6.1	Explosion	3
3	08/18/79	E. Kazakh	73.2	0	6.1	Explosion	3
4	09/24/79	Novaya Zemlya	76.7	Ō	5.7	Explosion	1
5	09/29/79	N. Sumatera	72.9	27	6.2	Earthquake	1
6	10/18/79	Novaya Zemlya	76.8	0	5.8	Explosion	1
7	10/28/79	E. Kazakh	73.2	Ō	6.0	Explosion	1
8	12/23/79	E. Kazakh	73.3	Ö	6.1	Explosion	ī
9	07/29/80	Nepal	76.4	18	6.1	Earthquake	2
10	09/14/80	E. Kazakh	73.2	0	6.2	Explosion	2
11	10/12/80	E. Kazakh	73.2	ō	5.9	Explosion	2
12	11/19/80	Sikkim	70.4	17	6.0	Earthquake	2

displacement may be related using the expression: $P = \omega \rho v A$, where P is pressure, ω is angular frequency, ρ is segmented density, v is the speed of sound in segmenter, and A is vertical

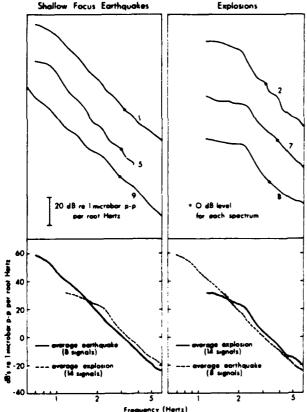


Figure 1. Sample spectra of P from some shallow focus earthquakes and nuclear explosions are shown in the upper portion of this figure.

Mumbers refer to the events as described in Table 1. The composite spectrum of each group, an average with ½ 1 standard deviation, is shown in the lower portion of the figure. Before standard deviations were computed, individual spectrums were normalized by subtracting the difference between their mean dB value over the range 1.5-3.0 Hz and the mean dB value for all spectra over the same frequency range.

displacement. This relationship holds for compressional energy arriving vertically from below the hydrophone. The following differences in the spectral signatures between explosions and shallow focus earthquakes are evident: (1) lack of energy in explosion P relative to earthquake P at frequencies below 1.5 Hz, (2) changes in spectral slope (corner frequency?) for explosions from -12 to -28 dB/octave (equivalent to -18 and -34 dB/octave in ground displacement) at about 2.2 Hz, and (3) greater energy in explosion P relative to earthquake P at frequencies above 2.0 Hz, despite smaller magnitude of the average explosion (6.03) than that of the average earthquake (6.16). Differences between the explosion and earthquake P phases observed at these frequencies is not surprising. The high frequencies observed and the actual shape of these curves have implications regarding Q along the travel path. Although these implications will not be discussed here, it is pointed out that the observation of frequencies in excess of 6 Hz (and up to 9 Hz) at distances greater than 60° implies a high Q at least along the deep mantle part of the P travel path.

Although this study found a common high-frequency slope (-28 dB/octave) for both earthquakes and explosions, at least one other investigator (Evernden, 1977) has found earthquakes to fall off at least f⁻¹ faster than explosions. This discrepancy might be explained by the small number of earthquakes used in the present study and the requirement in choosing those events of large signal/noise. Earthquakes having more high-frequency energy would have greater signal/noise because of much lower noise levels at high frequencies.

MTS events were not included in this study because they were rarely recorded with signal/noise greater than 2. Although signal levels near 1 Hz for NTS events are roughly equivalent to those observed for similar magnitude earthquakes and Soviet explosions, the signal fall-off above 1 Hz appears to be generally much greater for NTS. Therefore, little signal energy from NTS events is observed in the band above 2 Hz where lower ambient noise levels significantly enhanced the signal/noise of those events that were used. This rapid fall-off may be due to a recognized low Q effect in the source region (see for example Der et al., 1982).

Ocean Bottom Ambient Moise

The average background noise at the ocean bottom near Wake is shown in Figure 2 (labeled A). This average, along with its standard deviation, was determined from 52 samples of noise taken over 18 months of recording. Also plotted are an assortment of published noise curves for both ocean bottom and continental sites. When compared with noise levels from continental sites, the Wake ambient noise level could be described as: (1) high for frequencies between 0.2 Hz and 1.5 Hz; (2) average for frequencies between 1.5 Hz and 3.0 Hz; and (3) low for frequencies between 3.0 Hz and 15.0 Hz.

The observed low noise levels at higher frequencies affirm the ability of the Wake hydrophones to detect seismic signals at those frequencies. In addition, high-frequency phases recorded on the deep ocean bottom, which traverse only a few kilometers of homogeneous crust, may be less distorted than similar phases recorded on continents, which often traverse more than 40 km of crust. Consequently, coherence across the

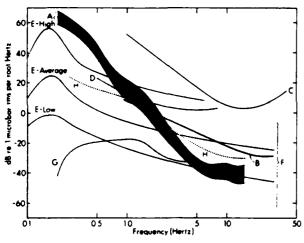
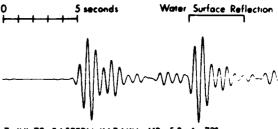


Figure 2. The average spectrum ± 1 standard deviation of 52 samples of background noise over 18 months from the Wake bottom hydrophones is labeled A. Also shown are some published noise curves for both ocean-bottom (B, C, D, and H) and continental (E, F, and G) environments, which have been converted from an assortment of units to the scale shown. B is a hypothetical "sample spectrum of deep-ses noise" (Urick, 1975; p. 188). C is a vertical seismometer measurement made in the Mariana Basin (Asada and Shimsmura, 1976). D is a vertical seismometer measurement made at 46-km depth between Hawaii and California (Bradner and Dodds, 1964). H is a noise curve for a hydrophone bottomed off Eleuthers Island at 1200-m depth (Nichols, 1981). E represents low, average, and high noise levels estimated from curves compiled by Brune and Oliver (1959). F is an area bounded by the limits of noise curves measured on vertical seismometers for 16 locations within the United States and Germany (Frantti et al., 1962). G is the noise curve for the Oyer subarray of the Norwegian seismic array measured during a period "when most of the Worth Atlantic Ocean was very quiet" (Bungum et al., 1971).



7 JUL 79 EASTERN KAZAKH MB: 58 A:73°



10 DEC 80 WESTERN SIBERIA MB = 4.6 4:77*

Figure 3. Sample time series of P, filtered to maximize signal/noise, from two nuclear explosions recorded on the Wake bottom hydrophones. The upper trace is from a single hydrophone and shows the direct arrival and its first water surface reflection. The lower trace is a composite of signals from two hydrophones with 40-km separation, obtained as follows: the filtered (1.5-5.0 Hz) time series from each hydrophone was inverted, shifted in time by the water surface reflection time, weighted to maximize the increase in signal/noise, and added to itself; the two resulting time series were then added with the appropriate propagation delay, and weighted to maximize the increase in signal/noise. Signal/noise was increased by 90% of the theoretical maximum with this method, indicating a high level of coherence between the signals added.

array appears to be high for telesiesmic P. These factors (low noise levels and a thin, homogeneous crust) have enabled the Wake array to acquire some impressive recordings of underground nuclear explosions. Shown in Figure 3 is an Eastern Kazakh explosion at 73° with a body wave magnitude of 5.9. Its signal/noise ratio is approximately 50/1. Also shown in Figure 3 is a Western Siberian explosion at 77° with a body wave magnitude of 4.6. This arrival is the weighted sum of signals from two of the hydrophones, as explained in the figure caption. Coherence between the two hydrophone signals in the band 1.5 to 5 Hz was measured at 0.85 for this arrival.

Summe TY

Significant differences are found between the spectra of P phases from explosions and from shallow focus earthquakes at 61° to 77° epicentral distance. Explosion spectra exhibit a change in spectral slope at about 2.2 Hz from -12 to -28 dB/octave relative to pressure. Earthquake spectra have a nearly constant slope of -28 dB/octave over the range of 1 to 6 Hz. High frequencies (>6 Hz) observed in these phases indicate a high Q in the deep mantle.

The ambient noise spectrum on the ocean bottom near Wake falls off at about -24 dB/octave over the range of 0.3 to 6 Hz. Between 3 and 15 Hz

the background noise levels are comparable to those at the quietest continental sites. Teleseismic P has been observed with a high level of coherence across a sensor separation of 40 km. The low level of ambient noise on the ocean floor at high frequencies and the high levels of coherence observed indicate that the ocean bottom may be an excellent observational regime for teleseismic P as well as other seismic phases rich in high frequencies.

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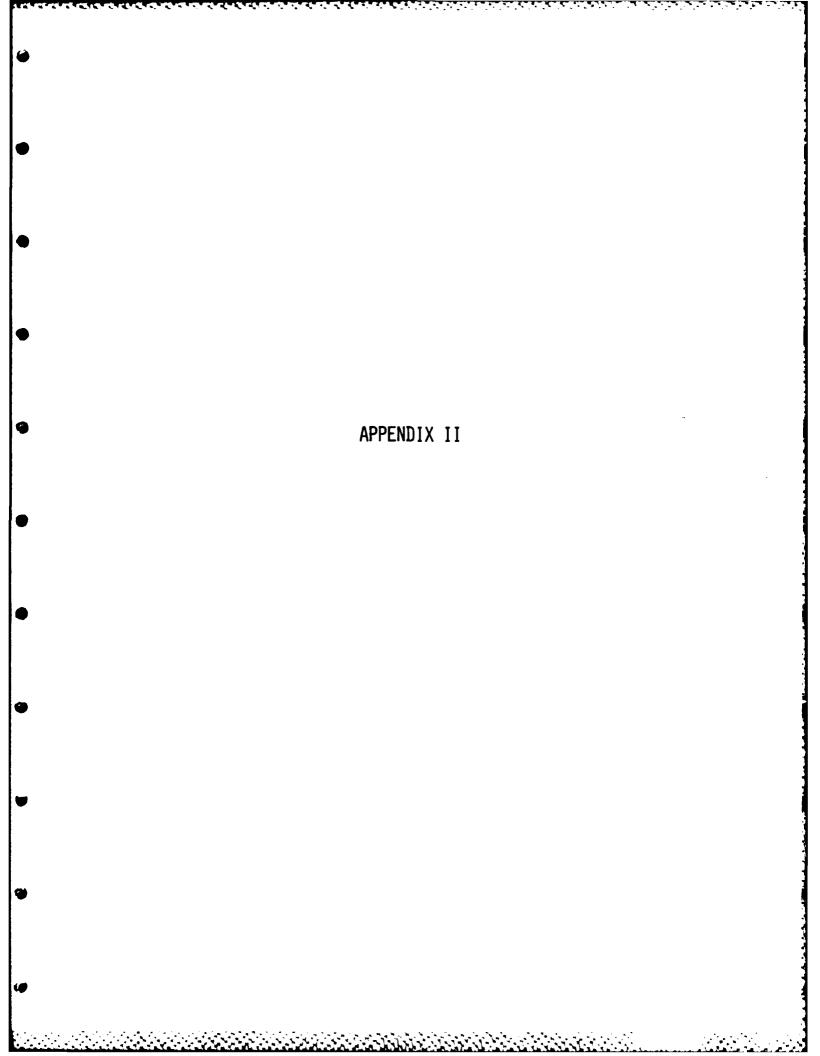
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SPECTRAL CHARACTERISTICS OF HIGH-FREQUENCY P_N , S_N PHASES

IN THE WESTERN PACIFIC

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Abstract. Pn and Sn phases from 25 selected earthquakes recorded since July of 1979 on ocean bottom hydrophones near Wake Island are used to complement and extend prior investigations of high-frequency Pn, Sn spectra in the Western Pacific. At a distance of about 18° ($\approx 2000 \text{ km}$), frequencies for Pn and Sn are as high as 30 and 35 Hz, respectively; at a distance of about 30° ($\simeq 3300$ km), as high as 15 and 20 Hz, respectively. Pn phases lose their high-frequency energy more rapidly than Sn phases do, yet Pn wavetrains are much longer than Sn wavetrains. Pn wavetrains of longer duration, more energy, and higher frequencies are found for travel paths primarily in the Northwestern Pacific Basin than for travel paths across the transition zone from the shallow Ontong-Java Plateau to the deep Northwestern Pacific Basin. Sn phases are extremely weak or absent for travel paths crossing this transition zone from the shallower Ontong-Java Plateau to the deeper Northwestern Pacific Basin, whereas Sn phases are well recorded for travel paths crossing the transition zone in the opposite direction. Although normal, mantle-refracted P phases are well recorded beyond about 21° (≈ 2300 km), available data indicate that detectable normal, mantle-refracted P phases may not exist at distances from about 17 to 21°.

Introduction

Recent investigations of high-frequency Pn, Sn in the Pacific [Walker, 1977; Walker et al., 1978; Sutton et al., 1978; Talandier and Bouchon, 1979; and McCreery, 1981] suggest that the real character of these phases is revealed at frequencies much higher than those traditionally associated with normal, mantle-refracted body waves at teleseismic distances (i.e., ~ 1 Hz). For example, in one investigation [Walker et al., 1978], frequencies as high as 12 and 15 Hz were found for the Pn and Sn phases, respectively, of an earthquake recorded at a distance of 28.3 (3147 km).

In this report we offer a more comprehensive analysis of the spectral characteristics of Pn, Sn using additional data recorded since July of 1979 on ocean bottom hydrophones near Wake Island. Only undistorted arrivals with signal/noise ratios of at least 3/1 were used in this investigation. Epicentral distances, origin

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times, depths, and magnitudes are given in Table 1; and locations of epicenters are shown in Figure 1.

Northwestern Pacific Basin Travel Paths

Spectrograms for some of the Pn, Sn phases having travel paths primarily under the deep Northwestern Pacific Basin (i.e., events 1 through 18) are shown in Figure 2. All reveal high frequencies, with values in excess of 20 Hz for both Pn and Sn at a distance of 18.0 (2000 km; event 2) and values of up to 15 and 20 Hz for Pn and Sn, respectively, at a distance of 29.4 (3270 km; event 17). (More detailed spectral analyses of the phases for event 2 at 18.0 indicate values as high as 30 and 35 Hz for Pn and Sn, respectively.)

Spectrograms for events 17 and 18 show the normal, mantle-refracted P phases as well as high-frequency Pn and Sn phases. Other events for which normal, mantle-refracted P phases have been clearly recorded are 11, 12, 13, 15, and 16. The fact that all of these events are at distances in excess of 21° is not coincidental, for it is only at these distances (the precise crossover depending, in part, on focal depth) that P phases begin to arrive ahead of the high-frequency Pn phase (Figure 3). With increasingly shorter distances, high-frequency Pn arrives increasingly ahead of the expected P.

Although it might seem reasonable to assume that P does arrive at distances less than about , but is masked by Pn, such an assumption should be tested. One test is to compare spectrums where all of the P's energy, or large portions of it, might be suspected of being present within the Pn coda (i.e., events 1 through 10) to the spectrums where only Pn is known to exist (i.e., events 13 through 18; 11 and 12 could not be used due to Pn clipping). Composite spectrums have been made for the two groups of Pn arrivals (i.e., Pn with P suspected, at distances from about 17° to 22°; and Pn with P known to be absent, at distances from about 26° to 33°), as well as for all P phases, at distances from about 22° to 33°, either clearly arriving well ahead of Pn (events 11, 12, 13, 15, 16, 17, and 18) or suspected of arriving close to, but shead of, Pn (events 9 and 10). These composites and the individual absolute spectrums from which they were derived are shown in Figure

Individual and composite P spectrums are obviously, and not unexpectedly, very different in character from individual and composite Pn

TABLE 1. Epicentral Distances, Origin Times, Depths, and Magnitudes of Events 1-25 in Figure 1

Event Number	Distance, deg.	Date	Time	Depth,	Magnitude mb
1	17.8	July 8, 1980	1704:15.1	54	4.8
2	18.0	July 11, 1980	0942:00.2	33	5.3
3	18.7	June 9, 1980	1923:33.3	33	5.6
4	19.0	Dec. 8, 1979	1258:55.2	51	5.5
5 6	19.8	Dec. 16, 1979	1050:48.0	96	5.0
6	20.1	March 26, 1980	0722:37.0	45	5.5
7	20.7	Nov. 1, 1980	0440:37.7	109	5.6
8	20.9	Dec. 17, 1979	0728:48.2	33	5.1
9	21.2	Jan. 15, 1980	0523:25.7	120	5.1
10	21.7	Nov. 29, 1979	1708:21.3	109	5.4
11	24.9	Dec. 11, 1979	1726:22.1	161	5.9
12	25.4	Dec. 19, 1980	2332:41.6	79	6.2
13	26.5	Oct. 20, 1980	0329:21.3	81	5.5
14	27.1	Oct. 28, 1979	0539:36.0	88	5.4
15	28.5	Feb. 23, 1980	0551:03.5	47	6.4
16	28.5	Jan. 1, 1981	1032:13.1	53	6.2
17	29.4	Nov. 26, 1980	2348:59.9	77	5.8
18	32.7	Aug. 22, 1979	1828:55.7	128	5.5
19	28.0	Feb. 12, 1980	0320:23.2	75	5.9
20	28.0	Aug. 13, 1979	0303:47.9	88	5.8
21	28.6	May 14, 1980	1126:00.6	57	6.1
22	28.8	Sept. 28, 1980	1825:59.7	68	6.0
23	30.4	Nov. 6, 1979	1138:31.5	30	6.0
24	31.1	Oct. 23, 1979	0951:06.7	22	6.1
25	31.1	Feb. 22, 1980	2115:42.1	68	5.9

spectrums, in that the P has larger signal-tonoise ratios at lower frequencies (i.e., 1-2 Hz)
than either Pn grouping (Figures 4 and 5), and
the composite P is richer in lows, and weaker in
highs, relative to the 26° to 33° Pn composite
(Figure 5a). In comparing the composite 17° to
22° Pn spectrum to the composite 26° to 33° Pn
spectrum (Figure 5b), we note that the 17° to 22°
Pn is similar in character to the 26° to 33° Pn
for frequencies higher than 2 Hz, but has lower
signal-to-noise ratios at frequencies less than 2
Hz (i.e., values fall below the "4 dB above
noise" requirement for plotting). This latter
observation is also apparent in the individual
spectrums (Figure 4). The fact that values for
the 17° to 22° Pn spectrum fall below the 4 dB
requirement for frequencies less that 2 Hz,
coupled with values for the 22° to 33° P spectrum

above the 4 dB level at those frequencies, suggests that detectable normal, mantle refracted P phases may not exist in the Pn codas of events at distances of 17° to 22° .

Another important, though not surprising, observation to be made from the composite plots (Figure 5b) is that Pn phases at great distances appear to be weaker at higher frequencies (\$10 Hz) than Pn phases at shorter distances.

Sn composite plots have also been made for the same events for which Pn composites have been made. These plots are shown in Figure 6. Although normal mantle-refracted P phases have been well recorded at great distances, normal mantle-refracted S phases from earthquakes have not been recorded by the Wake hydrophones. Presumably, this is due to the small pressure signal in the water resulting from S phases at

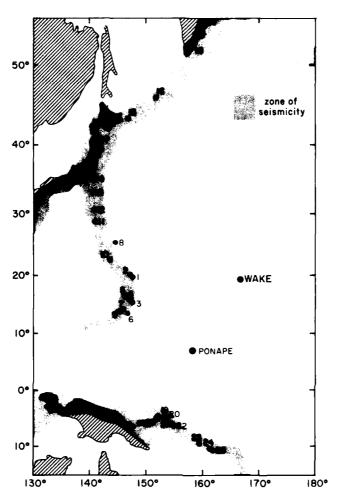


Fig. 1. Epicenter location map.

teleseismic distances. In addition, background noise levels are higher for those frequencies at which S would appear [McCreery et al., 1982].

In Figure 6a, Sn composite plots for the 17 to 22 and 26 to 33 distance ranges are compared to one another. Considering that the standard deviations of all of the composite plots presented in this paper are generally in the range of ± 3 dB, no significant differences are apparent in Figure 6a. In Figures 6b and 6c, Sn composite plots are compared to their respective Pn composite plots. Again, no statistically significant differences are indicated. In other words, Sn signal strength is generally comparable to Pn signal strength. This similarity of Pn and Sn spectrums was observed earlier for travel paths in the Western Pacific east of the Marianas [Ouchi, 1981].

Special mention should be made of the fact

that the composite Sn plot in Figure 6c is 4 dB above background noise at frequencies well above 8 Hz while the composite Pn plot is not. Also it should again be noted that the individual Pn's and Sn's used to formulate the composite plots were for the same earthquakes. These considerations suggest that Sn phases do not lose their high frequencies as rapidly as Pn phases.

It has been pointed out that some high frequencies observed elsewhere might be the result of instrumental nonlinearities [Sacks, 1980] and/or nonlinear seismic interactions in the vicinity of a receiving station [Nakamura and Koyama, 1982]. As the lower frequencies are of comparable amplitude for both Pn and Sn at both distance ranges considered (Figures 6b and 6c), it is unlikely that such nonlinearities could explain the data discussed here. We also note that P, which has higher average amplitudes than

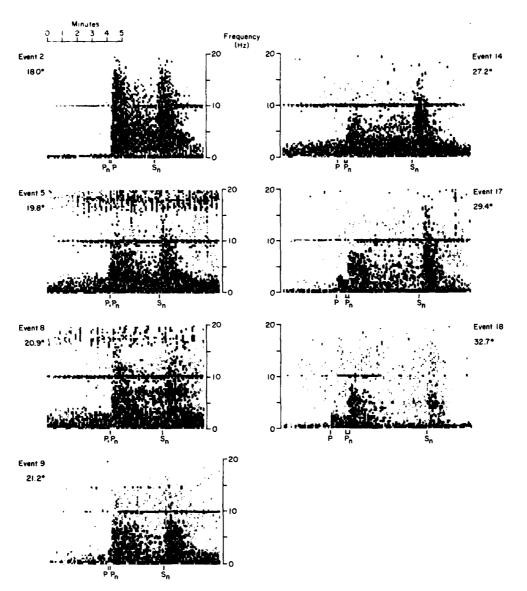


Fig. 2. Spectrograms for some earthquakes having travel paths to Wake under the Northwestern Pacific Basin. Expected times of arrivals are based on either the Jeffreys and Bullen [1958] tables for P or Pn/Sn travel time curves from Walker [1977]. These and succeeding spectrograms were made by dividing the time series into adjacent 512-point segments, Lanczos squared windowing the segments, and performing a fast Fourier transform (FFT) on each segment. In the horizontal direction, the width of each shaded block corresponds to one of the 512-point segments in the time series. In the vertical direction, each block is the average of two adjacent power spectral estimates out of the FFT. Only frequencies from 0 to 1/2 Nyquist are shown. The contour interval is 8 dB. The line at 10 Hz is due to time code cross talk. Instrument responses have not been removed.

Pn or Sn at the lower frequencies, has the most rapid falloff toward higher frequencies (Figure 5).

In all of these comparisons, another objection that could be made is that differences in source spectrums (and/or orientation of the source

relative to the recording station) were not considered. Although all of the events occurred within the subducting margin of the Northwestern Pacific and the Pn, Sn phases used were generated by earthquakes having focal depths of 128 km or less, differences in source spectrums might be

significant. We believe, however, that overall trends of the individual spectrums (Figure 4) used for the composite plots are similar (as opposed to specific details that may differ) and that such similarities could justify the general conclusions drawn from that data. We also note that source effects are minimized in those comparisons of composite Pn's and Sn's from the same earthquakes (Figures 6b and 6c).

Another interesting feature of Pn, Sn phases is that the Pn wavetrain is much longer than the Sn wavetrain (Figure 2). Spectral analyses indicate that energy is lost at all frequencies in the later arriving portions of these wavetrains, and that this loss is much greater in the Sn wavetrains than in the Pn wavetrains.

Ontong-Java Plateau Travel Paths

Spectrograms for some of the more interesting Pn phases with travel paths to Wake under the shallow Ontong-Java Plateau (as well as portions of the deep Northwestern Pacific Basin) are shown in Figure 7. (Refer also to Table 1, Figure 1, and Figure 8.) The most conspicuous feature of these spectrograms is that Sn phases are extremely weak, or absent, even though Pn phases are prominent.

Figure 9 compares the composite Pn spectrum for events having Ontong-Java Plateau travel paths (events 19 through 25) to the composite Pn spectrum for events at comparable distances having Northwestern Pacific Basin travel paths (events 15 through 18). The Northwestern Pacific Basin events appear to have more Pn energy at

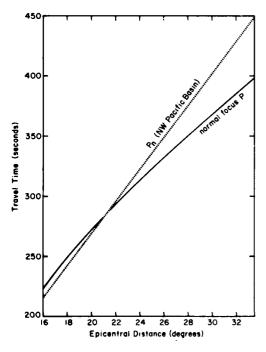


Fig. 3. Travel time curves for normal, mantlerefracted P phases and for Pn phases. P times are taken from Jeffreys and Bullen [1958], and Pn times are taken from Walker [1977].

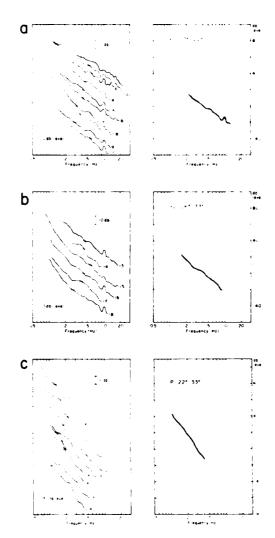


Fig. 4. Individual and composite spectrums for arrivals at Wake where (a) both Pn and P might be suspected of being present, at distances from about 17° to 22°, (b) only Pn is known to exist about 17° to 22°, (b) only Pn is known to exist at distances from about 26° to 33°, and (c) only P is known to exist, at distances from about 22° to 33°. The spectrums are power level spectrums in decibels relative to one microbar peak-to-peak pressure per root hertz. Values for individual spectrums were plotted only if they were at least 4 dB above background noise. The composite curves are simply the averages for the individual curves, with the condition that composite values were used for those frequencies lacking at most only one individual curve. Standard deviations of composite values are indicated by shading. Values for standard deviations in these and succeeding spectrums are generally in the range of \pm 3 dB. The peaks at 10 Hz are due to time code cross talk.

higher frequencies than the Ontong-Java Plateau events. However, because of the approximate + 3 dB standard deviation on each of these curves, this suggestion is not statistically significant. Comparing Figures 2 and 7, the duration of the Pn

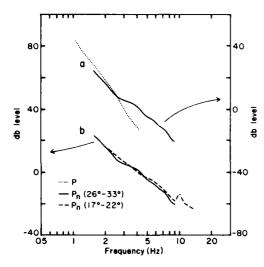


Fig. 5. Comparisons of composite spectrums: (a) P and the 26° to 33° Pn, and (b) the 17° to 22° Pn and the 26° to 33° Pn. Conclusions which can be drawn from these plots are (1) P and Pn spectrums are very different in character, (2) detectable P phases may not exist in the Pn codas of events at distances of 17° to 22°, and (3) Pn phases at great distances appear to be weaker at higher frequencies (~10 Hz) than Pn phases at shorter distances.

wavetrains appears to be greater for the Northwestern Pacific Basin events.

In such comparisons, the important question again arises of differences in source characteristics, in this instance for New Ireland-Solomon Island earthquakes and for Japan-Kuril Islands-Kamchatka earthquakes. It is not possible, however, to attribute the absence of Sn to differences in source characteristics, as Sn phases from New Ireland and the Solomons have been well recorded at Ponape on the northern margin of the Ontong-Java Plateau [Walker, 1977]. Examples of such phases are shown in Figure 10. Of the more than forty events from the New Ireland-Solomon Islands area recorded at Ponape, amplitudes of Sn phases are at least comparable to, and frequently larger than, those of their respective Pn phases.

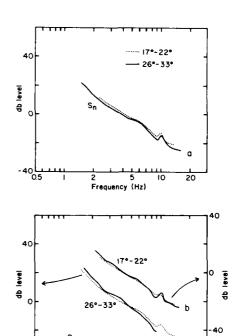
The absence of Sn at Wake would, therefore, appear to be a result of Sn's inability to propagate efficiently across the transition zone from the shallower Ontong-Java Plateau to the deeper Northwestern Pacific Basin, and would suggest that much of the energy in Sn travels through portions of the lithosphere involved in the transition. On the other hand, Sn's that have crossed this transition zone from the other direction (i.e., from earthquakes in the Marianas, Japan, the Ruriles, and Ramchatka) are well recorded at Ponape (Figure 10). Differences in the crustal structure of the Ontong-Java Plateau and the Northwestern Pacific Basin are indicated by the section profile [Hussong et al., 1979] shown in Figure 11.

Another comparison of spectrums at Wake for the two differing types of travel paths (Northwestern Pacific Basin and Ontong-Java Plateau travel paths) was made for the later arriving energy in the Pn wavetrains. For these comparisons, less energy at higher frequencies was present in those Pn's having travel paths that include the Ontong-Java Plateau. These deficiencies and the corresponding absence of Sn for paths across the transition zone from the shallow Ontong-Java Plateau to the deep Northwestern Pacific Basin suggest that the longer, stronger Pn phases observed for travel paths to Wake, primarily across the Northwestern Pacific Basin, may be the result of more efficient conversions of Pn to Sn (or Sn to Pn).

J)).

Concluding Remarks

The phenomenon of high-frequency Pn, Sn propagation is emerging as a major unresolved property of the oceanic crust and/or mantle. Others have described high-frequency Pn, Sn propagation as `a challenge remaining to the theoretician' [Richards, 1979] and as `the challenge to both explosion and earthquake



Frequency (Hz)

Fig. 6. Composite spectrums of Sn phases for those earthquakes with their Pn phases plotted in Figures 4a and 4b. Comparisons of spectrums are made (a) for the two Sn composites to one another, (b and c) and for each of the Sn composites to their respective Pn composite. A conclusion which can be drawn from these plots is that Sn signal strength is generally comparable to Pn signal strength except at high frequencies and large distances where Sn phases do not lose their energy as rapidly as Pn

phases.

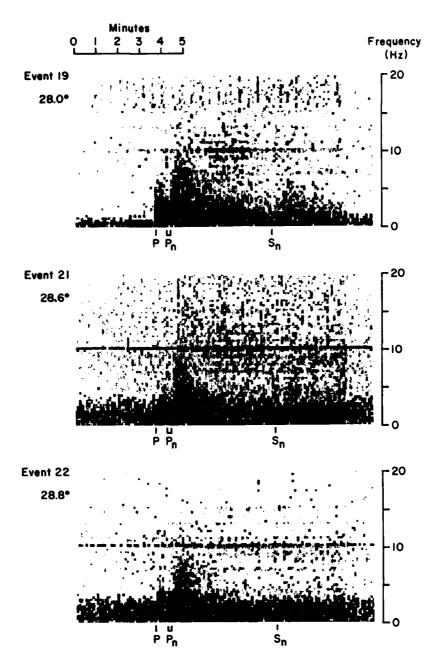


Fig. 7. Some spectrograms for earthquakes having portions of their travel paths to Wake under the Ontong-Java Plateau. In comparing these spectrograms to those of Figure 2, note the absence, or weakness, of Sn. Computational procedures are the same as used in Figure 2.

seismology for the coming decade' [Hirn et al., 1973]. These descriptions are supported not only by the unusual character of the phases but also by their probable occurrence throughout the world's oceans.

As important as recent efforts are to determine the mechanism of high-frequency Pn, Sn propagation [e.g., Stephens and Isacks, 1977; Menke and Richards, 1980; Sutton and Harvey, 1981; Gettrust and Frazer, 1981], we believe that

many essential characteristics of Pn, Sn phases (especially at very high frequencies) are not well known, and that accurate quantification of those characteristics through the acquisition of additional high-quality data is greatly needed. We hope that this report will further familiarize seismologists with high-frequency Pn, Sn propagation and will be viewed as a preliminary attempt to quantify, in a relative sense, some of the essential characteristics of these phases. .

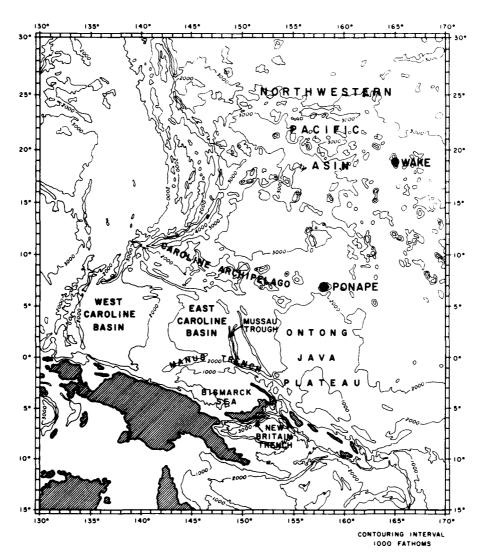


Fig. 8. Bathymetry map of the Northwestern Pacific area.

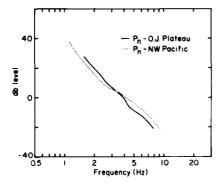


Fig. 9. Comparisons of composite Pn spectrums for events having travel paths across the Ontong-Java Plateau and the Northwestern Pacific Basin.

summary of principal observations contained in this report follows.

The apparent absence of normal, mantle-refracted P phases at distances less than about 21° ($\simeq 2300$ km).

Frequencies as high as 30 and 35 Hz for Pn and Sn, respectively, at 18.0° (2000 km)

Frequencies as high as 15 and 20 Hz for Pn and Sn, respectively, at 29.40 (3270 km).

With increasing distance (i.e., from about 20° to 30°), Sn phases not losing their high frequencies as rapidly as Pn phases do.

Pn wavetrains longer than Sn wavetrains.

The extreme weakness or absence of Sn phases for travel paths across the transition zone from the Ontong-Java Plateau to the Northwestern Pacific Basin and the presence of Sn phases for travel paths in the opposite direction across this transition zone.

Longer, more energetic Pn wavetrains for travel paths primarily in the Northwestern Pacific Basin than for travel paths across the transition zone from the Ontong-Java Plateau to the Northwestern Pacific Basin.

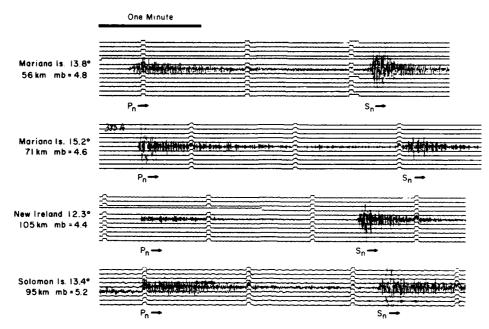


Fig. 10. Pn and Sn phases recorded at Ponape on the northern margin of the Ontong-Java Plateau.

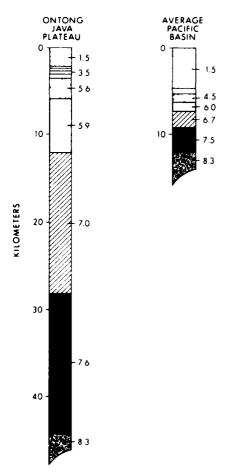


Fig. 11. Section profiles for the Ontong-Java Plateau and Pacific Basin as taken from Hussong et al. [1979].

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APPENDIX III

OCEANIC PN/SN PHASES:

A QUALITATIVE EXPLANATION

AND REINTERPRETATION

OF THE T-PHASE

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November 1982

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Charles E. Helsley

Director

Hawaii Institute of Geophysics

ABSTRACT

The combined effects of: (1) differing efficiencies between Pn and Sn energy transmission across the basement-sediment interface; (2) ocean surface reflections; (3) Pn to Sn conversions; and, (4) large lateral variations in the crust and upper mantle are used to formulate a working hypothesis which appears to explain, qualitatively, many observations of high-frequency Pn/Sn phases throughout the western, northern, and central Pacific. Also, the concept of Pn/Sn phases as sources of energy at the basement-sediment interface is suggested as a possible mechanism for T-phase generation through scattering or Stoneley wave generation.

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INTRODUCTION

High-frequency Pn/Sn phases were first observed nearly fifty years ago for travel paths in the Atlantic from the West Indies to the northern east coast of the United States (Leet et al., 1951, reported that these observations were made as early as 1935). The history of research on these phases since that time has been given by several authors, with some of the more recent being Molnar and Oliver (1969) and Walker (1977a).

My interest in the phenomenon began in 1963 with the recording of long-range, high-frequency Pn/Sn phases on hydrophones of what was then known as the Pacific Missile Range facility. My efforts at trying to understand these unusual phases (Sn wavetrains greater than Pn wavetrains with frequencies as high as 15 Hz at distances in excess of 3000 km; Walker et al., 1978) has continued since those initial observations. A reasonable, concise summary of my accomplishments prior to this report might be that no answers were found--only more questions. Such methodology can only be rationally tolerated for a length of time much shorter than my tenure on the case. So, for the past few years, it has been hoped that a working hypothesis could be found that would permit many, if not all, of the diverse observational pieces to be fitted together in at least a qualitative sense. Such a hypothesis could then serve as an appropriate starting point for comprehensive and detailed quantitative analyses leading to a generally acceptable model for the generation and propagation of long-range, high-frequency Pn/Sn phases. A working hypothesis has now emerged and is the subject of this report.

SN SIGNAL STRENGTH

One of the most interesting aspects of long-range, high-frequency Pn/Sn propagation in the northwestern Pacific is the strength of the Sn phase (Walker et al., in press). Relative to Pn, Sn often appears stronger at great distances—even at high frequencies (Fig. 1). A possible explanation for these observations begins with considerations of:
(1) the efficiencies of Pn/Sn energy transmission across the basement-unconsolidated sediment interface; (2) possible conversions of Pn and Sn at the basement-sediment interface; and, (3) ocean-surface reflections. Observations of conversions upward and across the basement-sediment interface, as well as ocean-surface reflections for short travel paths

off the coast of California are reported in Auld et al. (1969), where data indicate that the percentage of P (perhaps actually Pn) energy converted to S is greater than the percentage of S (perhaps actually Sn) energy converted to P. Examples of ocean-surface reflections are also given in Shimamura et al. (1975) and McCreery et al. (in press).

The very recording of "Pn/Sn" phases on ocean-bottom sediments is evidence that Pn and Sn energy does propagate, by some means, into the sediments. Energy losses above the basement-sediment interface will occur, i.e. not all of the energy passing into the sediments will be returned to the interface by way of reflections from the ocean surface. These losses could be much greater than those produced within that portion of the waveguide below the basement-sediment interface. Furthermore, if the percentage of Sn energy lost after propagation up through the basement-sediment interface, as S or converted P or both, is less than the percentage of Pn energy similarly lost up through this interface, then Sn would retain a greater percentage of its initial energy within that portion of the waveguide below the basement-sediment interface.

Under these assumptions Sn signal strength could increase with distance relative to Pn signal strength--even at high frequencies.

PN WAVETRAINS

In spite of Sn's greater signal strength, its wavetrain is considerably shorter than Pn's (Figs. 1 and 2). This seemingly peculiar observation can be explained by considering other types of conversions.

Any phases, originally Pn or Sn, that pass upward through the basement-sediment interface may eventually be reflected by the ocean surface (or the sediment-water interface) and returned to the basement-sediment interface to continue, somewhat weakened, in the waveguide as Pn or Sn or both. Throughout the travel path to the station, original Pn's returning to the interface and continuing in part as Sn's, as well as original Sn's returning to the interface and continuing in part as Pn's, would arrive between the main Pn and Sn phases. The net effect of such conversions would be Pn wavetrains greater in their duration than Sn wavetrains.

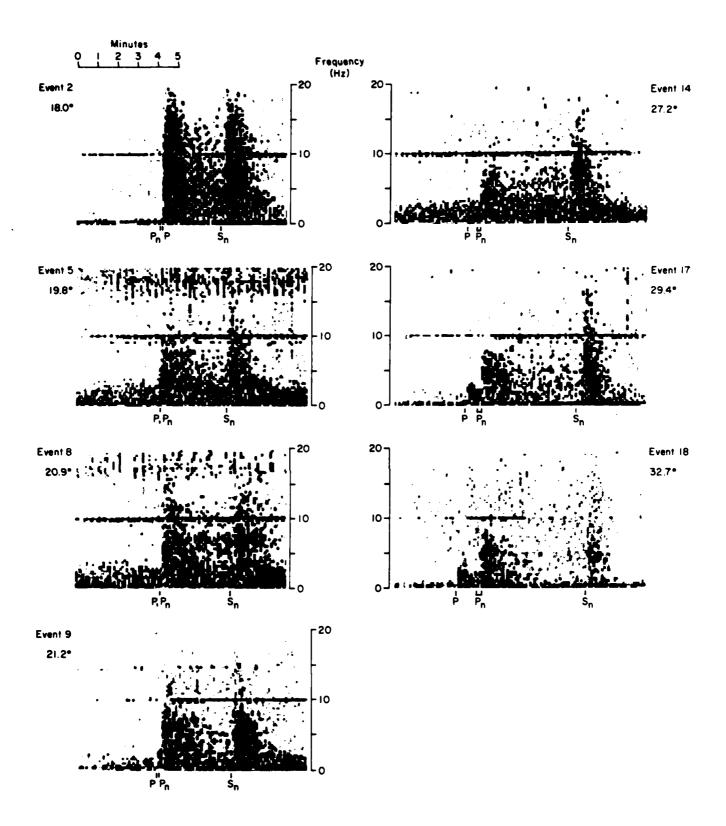


Fig. 1. Spectrograms for some earthquakes having travel paths to Wake under the Northwestern Pacific Basin. Expected times of arrivals are based on either the Jeffreys-Bullen tables (1958) for P or Pn/Sn travel time curves from Walker (1977a). The contour interval is 8 db. The line at 10 Hz is due to time code cross talk.

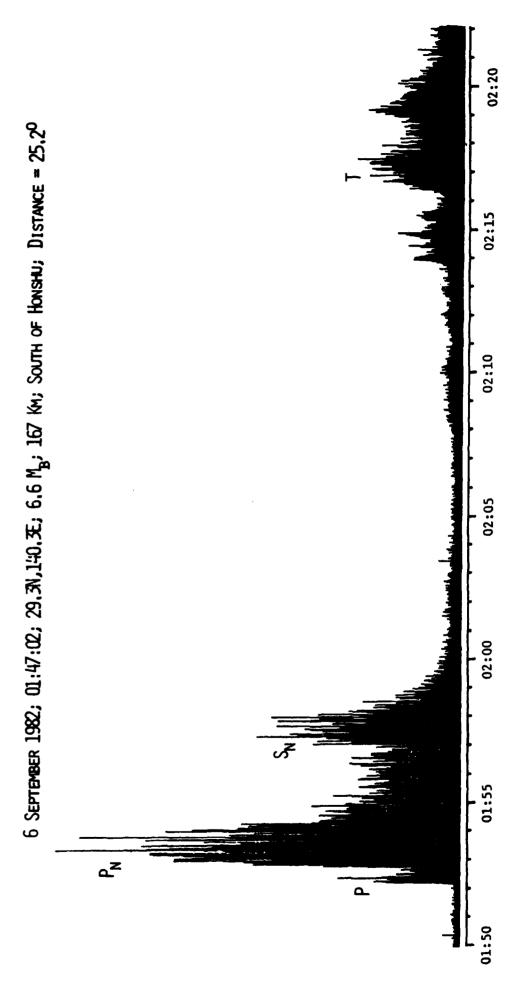


Fig. 2. Digitally rectified and compressed plot of P, Pn, Sn, and T phases for an earthquake south of Japan recorded by the Wake Island hydrophones.

OCEAN-SURFACE REFLECTIONS

Other potentially significant contributors to the Pn, as well as Sn, wavetrains are multiple ocean-surface reflections of all Pn/Sn variations observed on the ocean bottom that are capable of becoming compressional water waves at the sediment-water interface. Such reflections have already been proposed in synthetic Pn model studies (Gettrust and Frazier, 1981). An example of an ocean-surface reflection is shown in Figure 3.

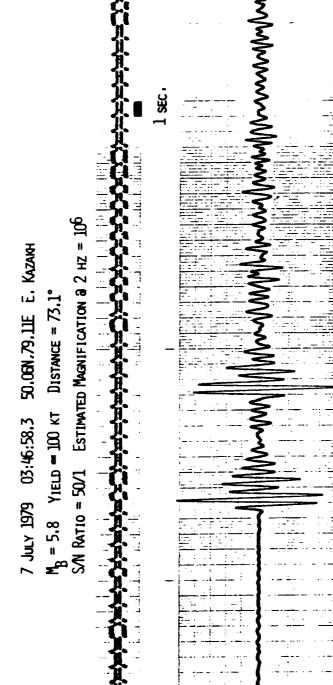
PN/SN PHASES AT GREAT DISTANCES

Another puzzling aspect of Pn/Sn propagation is the frequent absence or weakness of Sn, yet presence of Pn, at great distances (often more than 4000 km) throughout the North Pacific (Walker, 1977a and b) and Central Pacific (Talandier and Bouchon, 1979), in spite of stronger Sn's than Pn's for relatively homogeneous travel paths across the deep Northwestern Pacific Basin.

A possible explanation is that paths other than those across the relatively homogeneous deep Northwestern Pacific Basin are likely to encounter relatively large lateral changes in the crust and upper mantle. These changes would have to be of such a nature so as to reduce Sn signal strength without seriously affecting the Pn phase. Large lateral changes could be produced by plateaus, rises, ridge systems, island and seamount chains, fracture zones, transform faults, fossil arcs and trenches, and rafted continental fragments.

Specific examples of Sn's severely attenuated by large lateral variations are found in recordings of earthquakes from the Solomon Islands area on the Wake hydrophones. Although these events are at distances comparable to events from Japan and the Kurils which have strong Pn/Sn phases (Fig. 1), only their Pn's are well recorded (Fig. 4). Furthermore, this effect cannot be attributed to differences in source mechanisms because Sn's from the area are well recorded at Ponape (examples of such recordings are shown in Fig. 5; refer to Fig. 6 for the location of Ponape). A reasonable explanation would appear to be the large structural changes associated with the transition from the shallow Ontong Java Plateau to the deeper Northwestern Pacific Basin (Figs. 6 and 7). An additional factor may be the extension of the Caroline Archipelago through this region.

WAKE HYDROPHONE RECORDING OF UNDERGROUND EXPLOSION



SURFACE REFLECTION (NEARLY EXACT INVERSE OF FIRST FIVE PULSES)

Fig. 3. An example of ocean-surface relfections.

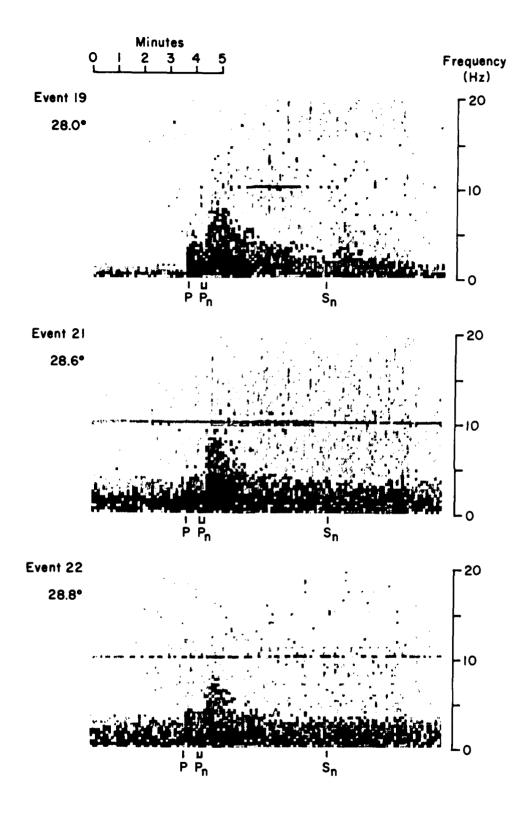
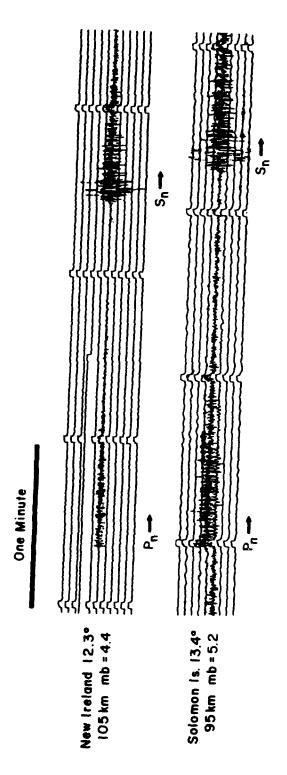


Fig. 4. Some spectrograms for earthquakes having portions of their travel paths to Wake under the Ontong-Java Plateau. Computational procedures are the same as used in Fig. 1.

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area, amplitudes of Sn phases are at least comparable to, and frequently larger than, Of more than forty events from the New Ireland-Solomon Islands Fig. 5. Examples of Pn and Sn phases recorded at Ponape on the northern margin of the those of their respective Pn phases.

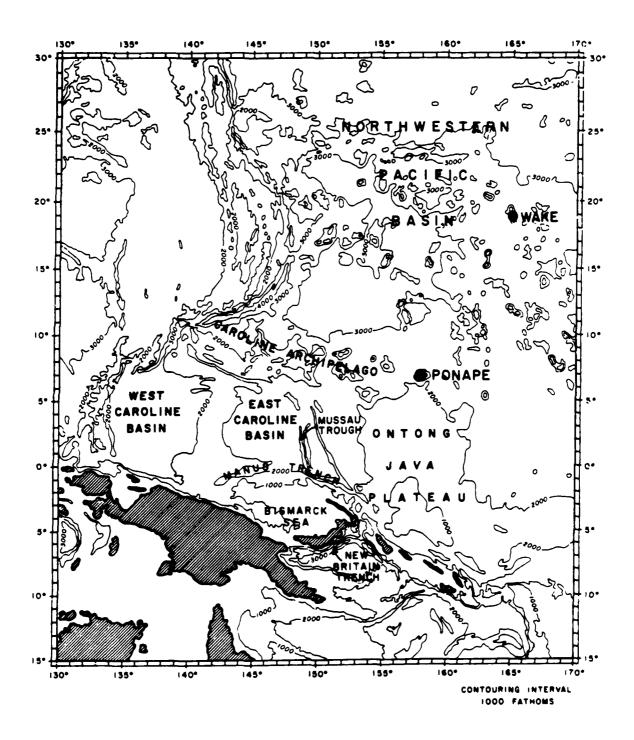


Fig. 6. Bathymetry map of the Northwestern Pacific area.

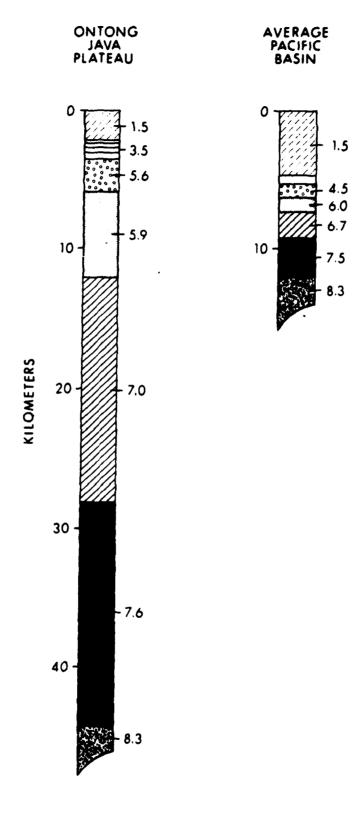


Fig. 7. Section profiles for the Ontong-Java Plateau and Pacific Basin (from Hussong et al. 1979).

T-PHASE MECHANISMS

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Tolstoy and Ewing (1950) recognized the importance of a sloping bottom in the production of T-phases, and Milne (1959) provided a specific mechanism involving multiple reflections between the surface of the ocean and its downward sloping bottom. Although such a mechanism could be significant for T-phases originating in areas with the appropriate downward sloping bottom, many strong T-phases have been observed from regions where the ocean floor is level or at a greater depth than adjacent areas in the direction of the receiver (Johnson et al., 1968; Duennebier, 1968). Included among these regions are the deep ocean floor off the coast of California and Oregon, the deep ocean floor south of the subducting margins of the Northern Pacific and east of the Western Pacific, and the East Pacific Rise.

Because of these observations, a mechanism other than downslope propagation is required. Some suggestions included
ocean surface scattering (Johnson et al., 1968), scattering
from the sea floor by fault scarps near the source (Johnson
and Norris, 1970), and coupling of Stoneley waves into the
SOFAR channel (Biot, 1952; Duennebier, 1968). All of these
proposed mechanisms, including downslope propagation,
presume that the energy of the T-phase comes ultimately from
P and, perhaps, S phases travelling upwards through the
crust to the ocean bottom near the T-phase source location;
however, with such a presumption, none of the proposed
mechanisms is capable of explaining T-phase forerunners. In
describing these forerunners, Johnson (1963) stated:

"The time of earliest perceptible arrival is probably primarily a function of magnitude as the signal emerges slowly from the ocean background noise...Early, low-level arrivals, undetected in most T-phase recordings, must be normal-mode ground waves or, at least, must have followed a ground path for a significant portion of their travel."

He also states that in one instance (a 7.0-Ms earthquake from the Kurils) the forerunners were so early that "the transformation from P or S waves to sound channel waves would have to occur at a distance of about 1, 15, 21° from the source toward the receiver." In describing a P, S, and T phase (Fig. 8; actually the P and S phases are Pn and Sn phases) from a large (6.2 mb) earthquake in the Marianas recorded on hydrophones near Enewetok Atoll at a distance of about 18°, Duennebier (1968) notes that "the T phase does not have a definite onset and that energy was continuously received at the hydrophone after the arrival of the P wave."

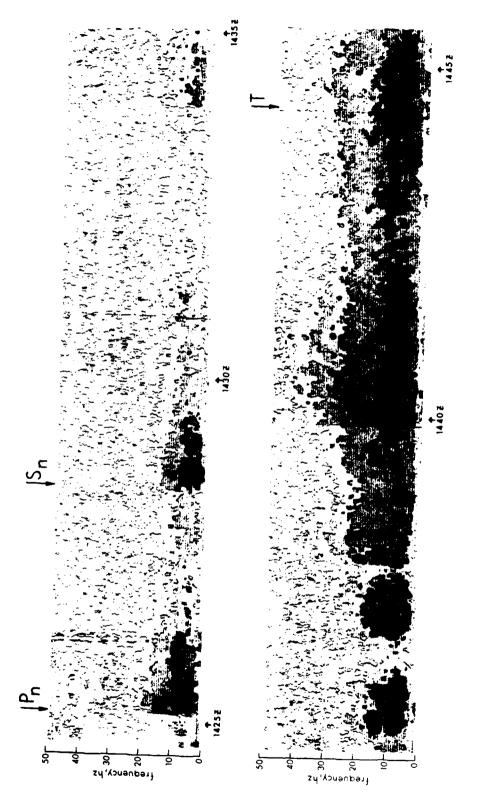
Both authors suggest that the apparent coupling into the SOFAR channel was due to normal, mantle P or S waves or both refracted by the Emperor Seamount Chain (Johnson, 1963) and by several groups of seamounts in the Northwestern Pacific Basin (Duennebier, 1968). In terms of frequency content and strength of signal, however, the energy from the long-range, high-frequency, guided Pn and Sn phases seems more likely to be coupled into the SOFAR channel than the mantle-refracted P and S phases which are extremely weak at high frequencies. At teleseismic distances frequencies of P and S generally do not exceed 3 or 4 Hz, while Pn, Sn, and T frequencies may be as high as 20 Hz (Figs. 8 and 9).

For large earthquakes sufficient energy could be contained in the Pn/Sn phases for coupling into the SOFAR channel throughout the travel path to the receiver. This coupling could occur by any of the mechanisms proposed for the generation of the T-phase (i.e., scattering or Stoneley waves or both), with seamount enhancement remaining as an important consideration. In effect the Pn/Sn phases would serve as potential sources of energy at the sediment-basement (or water-sediment) interface. As Pn/Sn energy declines with increasing distance, the amount of energy coupled into the SOFAR channel would also decline, thus producing a T-phase signal which would slowly emerge from the ocean background noise. A recently recorded example of such an emergent T-phase is shown in Figure 2. Energy arriving at 02:08 corresponds to a Pn path of 9.6° at 8.0 km/sec and a T-phase path of only 15.6° at 1.5 km/sec.

Finally, an additional explanation for Sn appearing to have relatively more energy than Pn with increasing distance could be that Pn energy is coupled more efficiently into the SOFAR channel than Sn energy is.

SUMMARY

Vast regions of the world's oceans are characterized by thinly layered, homogeneous crustal and uppermost mantle structure. This report suggests that comprehensive observations of high-frequency phases, often referred to as Pn/Sn, throughout the Western, Northern, and Central Pacific may be explained by a waveguide which extends upward from the uppermost mantle through the crust and, to some extent, into the sedimentary layers and the entire water column. Pn energy is more efficiently propagated upward into the sedimentary column than



not have a definite onset. Instead, energy appears to be continuously received at the hydro-Fig. 8. Sonogram of Pn, Sn, and T phases (after Duennebier 1968). Note that the T-phase does phone after the arrival of the P wave.

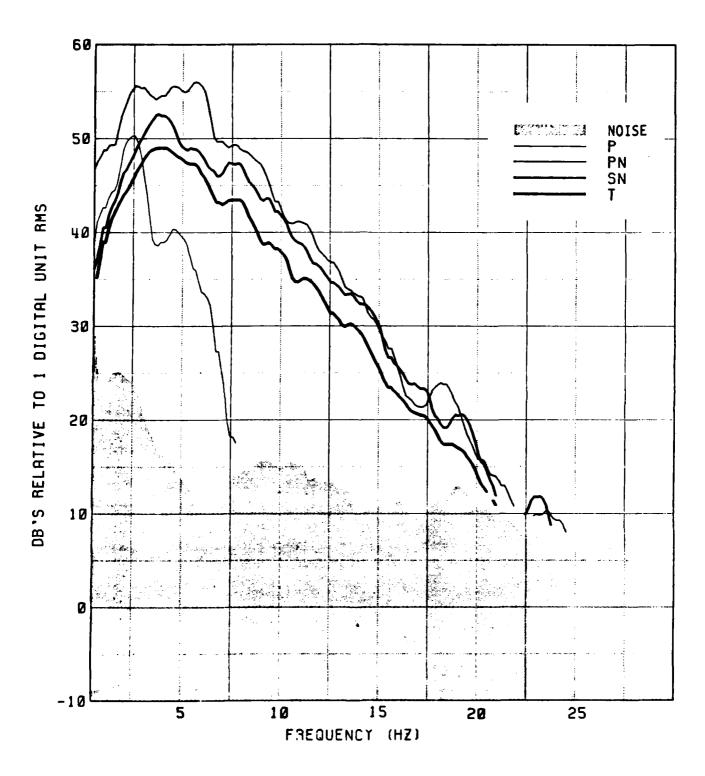


Fig. 9. Spectrums for the P, Pn, Sn, and T phases shown in Figure 2.

Sn energy, and, because of this, Pn is more rapidly attenuated than Sn. As distance increases, Pn's losses above the basementsediment interface produce a relative strengthening of the Sn phase, such that Sn eventually has more energy than Pn.

The observed data also suggests that the long Pn/Sn wavetrains may, in part, be the result of multiple ocean-surface reflections near the receiver.

Pn to Sn and Sn to Pn conversions are suggested to explain Pn wavetrains consistently longer than the wavetrains of Sn. These conversions could occur when Pn or Sn energy passes through the basement-sediment interface and is returned to that interface by way of reflections from the ocean surface to continue in the waveguide, at least in part, as Sn or Pn phases. Such conversions might be produced at any of the interfaces encountered by Pn or Sn, or their converted phases.

In conjunction with the scattering mechanisms and Stoneley wave propagation already proposed to explain T-phase observations, Pn/Sn phases as sources of energy at the sediment-basement interface could explain: (1) the generation of T-phases in regions not having the downward sloping ocean bottom required by the classical downslope mechanism of Milne (1959) and (2) T-phase forerunners extending well ahead of the peak arrivals--occasionally beginning perhaps as early as the Pn and/or Sn phases. Also, another explanation for Sn having more energy than Pn is provided if Pn energy is more efficiently transmitted than Sn energy into the SOFAR channel.

Finally, the observed data suggest that large lateral variations in the crust and upper mantle may produce significant reductions in Sn signal strength without seriously affecting the Pn phase.

CONCLUSIONS

Although many puzzling observations of long-range oceanic Pn/Sn phases may be resolved by the suggestions presented in this report, many of these suggestions can, and should, be tested quantitatively. These analyses could include the re-examination of existing data, new experiments, and theoretical modeling efforts leading to the generation of complete synthetic seismograms. Furthermore, one should not forget that the mechanisms for the generation and propagation of the main Pn and Sn phases are still not generally agreed

upon, nor have proposed models been matched in a comprehensive manner with observations. Such tasks are of utmost importance if the phenomenon is to finally achieve the status it deserves as a major geophysical feature and tool for mapping the crust and uppermost mantle of the world's oceans—this after nearly fifty years of being little more than an obscure curiosity.

Regarding the T-phase, it seems appropriate that the earth's most efficient acoustical waveguides, the SOFAR channel and the Pn/Sn waveguide, finally should be related; and in the hypothesis formulated in this report, the SOFAR channel is energized by leakage from the Pn/Sn waveguide. The recognition of this relationship could be an important, and perhaps critical, factor in arriving at a comprehensive understanding of these oceanic phases.

"In the bulletin of the Harvard Seismograph Station, under date of September 15, 1935 attention was directed to the unusual character of certain records from the vicinity of 17°N, 62°W. One of the nevel features was a short-period phase about 23 minutes after P. It has become known as T, for third, with P and S constituting the first and second groups of short-period waves of similar general appearance...Actually, many features of P and S are abnormal on this and later records from certain areas at this distance range, and work on that part of the problem is in progress, but the investigation of T has been undertaken first." (Leet et al., 1951)

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Special appreciation goes to Charles McCreery, George Sutton, and Loren Kroenke for constructive comments and discussions. Editorial assistance was provided by Barbara Jones.

SUPPLEMENTARY NOTE

It may be appropriate at this time to suggest that a new name be given to the high-frequency compressional and shear phases often observed at great distances in the world's oceans. The difficulty with the nomenclature used to date is that: (1) an, as yet, unsubstantiated relationship to the well known and much studied longer-period Pn/Sn phases of continents is inferred; and (2) the environmental feature most strongly linked to the observations is not cited. Thus, a more logical term would be "Ocean P" or "Ocean S" with the abbreviations being "Po/So." With this change, those unfamiliar with the phenomenon would not be as likely to make the false assumption that the phases are similar to continental Pn and Sn. Such assumptions in the past have been a major stumbling block in stimulating interest and support for "Po/So" research.

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The combined effects of: (1) differing efficiencies The combined effects of: (1) differing efficiencies between Pn and Sn energy transmission across the basement-sediment interface; (2) ocean surface reflections; (3) Pn to Sn conversions; and, (4) large lateral variations in the crust and upper mantle are used to formulate a working hypothesis which appears to explain, qualitatively, many observations of high frequency Pn/Sn phases throughout the western, northern, and central Pacific. Also, the concept of Pn/Sn phases as sources of energy at the basement-sediment interface is suggested as a possible mechanism for T-phase generation through scattering or Stoneley wave generation.

CORRECTIONS AND ADDITIONS

- pg. 12 line 5
 after "... (Duennebier, 1968)."
 add "Northrop (1972) also suggests a transformation from P to T
 at a distance of about 17° to explain the earliest
 - at a distance of about 17° to explain the earliest perceptible T phase arrivals on Oahu hydrophones from the 28 March 1964 Alaskan earthquake."
- pg. 12 line 27 (last sentence of 2nd paragraph)
 - "... path of 9.6° at 8.0 km/sec and a T-phase path of only 15.6°..."
 - should read "... path of 10.1° at 8.0 km/sec and a T-phase path of only 15.1°..."
- pg. 12 line 32 (last sentence of 3^{Td} paragraph)
 after "... than Sn energy is."
 add " Conversions of Pn or Sn to T near the receiver could also
 strengthen and extend the Pn and Sn wavetrains."
- pg. 15 line 28 (last sentence of 4th paragraph)
 after "... the SOFAR channel."
 add " Conversions of Pn or Sn to T near the receiver could
 strengthen and extend the Pn and Sn wavetrains."
- pg. 18 McCreery et al. has since been published in volume 10, p. 59-62, of Geophysical Research Letters, 1983.
- pg. 19 add "Northrop, J. 1972. T-phases, in The Great Alaskan Earthquake of 1964; Oceanography and Coastal Engineering, Nat. Acad. of Sci., Washington, D.C., p. 19-24."

Dan Walker

APPENDIX IV

A Preliminary Informal Comparison of Signal/Noise Capabilities Between The Wake Bottom Hydrophone Array, The Ocean Sub-Bottom Seismometer, And Ocean Bottom Seismometers

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Abstract. A comparison has been made between noise levels and signal/noise (S/N) ratios of the Wake Bottom Hydrophone Array (WBHA), the Ocean Sub-Bottom Seismometer (OSS), and Ocean Bottom Seismometers (OBS's) using data which has recently become available. The general purpose of this comparison is to evaluate WBHA as a possible alternative or complement to deep sea drillhole seismometers such as the Marine Seismic System (MSS) now being developed and tested for possible deployment in the Northwestern Pacific Basin. Absolute noise measurements show roughly comparable levels between WBHA and OSS, although the OSS data was scant. S/N ratios were found to be approximately 17 dB greater on the OSS than on a nearby OBS over the band 2-20 Hz, and 10 dB greater on WBHA than on a nearby OBS over the band 1-20 Hz. Improvements in WBHA S/N ratios of at least 3 dB and up to 11 dB, by stacking the water surface reflections and all 6 hydrophone signals, may be gained depending upon the level of coherence between signals. An evaluation of coherence levels has revealed apparent differences between the individual hydrophone/cable responses. Appropriate correction factors for these differences will have to be determined and then applied to the data before coherence levels are accurately known. The limited amount of data available for the noise and S/N ratio comparisons, and the indirect route by which the comparisons were made are arguments for attaching a certain amount of skepticism to the results. Additional OSS data which will become available after May 1983 should provide for a more direct and statistically stronger comparison.

Introduction. Considerable effort has been expended in recent years to develop a deep-ses borehole seismometer such as DARPA's Marine Seismic System (MSS) and the Hawaii Institute of Geophysics (HIG) Ocean Sub-Bottom Seismometer (OSS), which can be deployed down drillholes in the ocean bottom. Impetus for this effort has come from a desire to make high quality seismic observations on the vast portions of the earth covered by oceans. Although seismic observations in the oceans are capable of being made more or less routinely using Ocean Bottom Seismometers (OBS's) sitting on top of the sediments, it has often been presumed that an instrument located below the sediment-basalt interface would have significantly greater signal/noise (S/N) ratios, at least for upwardly arriving signals. This is because such an instrument should receive more energy from seismic signals arriving from below, and less energy from noise propagating downward through the water column and sediments.

Since July of 1979, HIG has monitored on a nearly continuous basis, an array of hydrophones located near Wake Island in the Northwestern Pacific Ocean. This array consists of six hydrophones on the ocean bottom at 5.5 km water depth in a 40 km, 2-dimensional grouping, and 3 pairs of hydrophones at 1 km water depth (SOFAR channel) spread over 300 km.

Recordings from only two or three of the hydrophones were made prior to September 1982; subsequently recordings from eight of the hydrophones have been made. The array was installed with the hydrophones hardwired to Wake Island more than 20 years ago for the purpose of recording transient, non-tectonic signals at frequencies above 10 Hz. Recordings made by HIG indicate that the hydrophones are also excellent seismic sensors in at least the band 0.5-20 Hz. The capability to evaluate the full potential of this array as a seismic tool has only been possible since the September

1982 upgrading. One aspect of this evaluation is to compare the noise levels and the S/N ratios between the Wake Bottom Hydrophone Array (WBHA), a sub-bottom seismometer (the OSS), and an OBS. Although data for a direct comparison does not yet exist, a somewhat indirect comparison has been made using data collected recently. That comparison is the subject of this report.

Absolute Background Noise Levels. In September 1982, during Leg 88 of the R/V Glomar Challenger, the OSS was successfully deployed off of the Kuril Islands, 378 m below the ocean floor and 20 m below the sediment-basalt interface. For a period of about 2.5 days, data was recorded from the OSS directly on the ship. Then a 2 month continuous recording package was dropped over the side for pickup sometime in May 1983. Briefly overlapping this 2.5 day period was the nearby deployment of several HIG OBS's. Noise levels measured at about the same time by the OSS vertical seismometer and a representative OBS vertical seismometer are shown in Figure 1. Although the general shape of the curves is similar, the difference in absolute levels is striking. Some of this difference is due to the impedance difference between the basalt and the top of the sediment. Smaller signals, as well as lower noise levels, would be associated with the higher impedance of the basalt. If normalized to the sediment impedance by a Voc correction, where ho is density and c is compressional velocity, the OSS noise curve moves up by about 9 dB to the position of the dashed curve. The remaining differences between these two noise curves may reflect the attenuation of downwardly propagating noise energy. Some of this noise is certainly generated by the research vessels which were operating overhead.

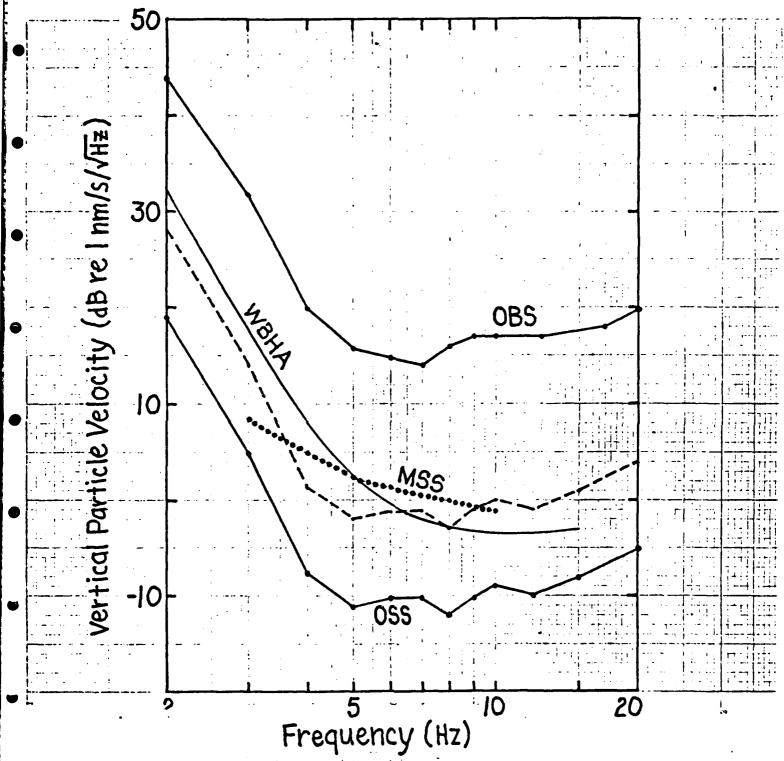


Figure 1. Background noise level measurements from WBHA (one hydrophone, 52 samples over 18 months), from an OBS and the OSS deployed in the NW Pacific Basin (vertical component-1 simultaneous sample) and the MSS in the Atlantic (vertical component-1 sample). The dashed curve is the OSS curve adjusted by 9 dB to normalize it to sediment impedance.

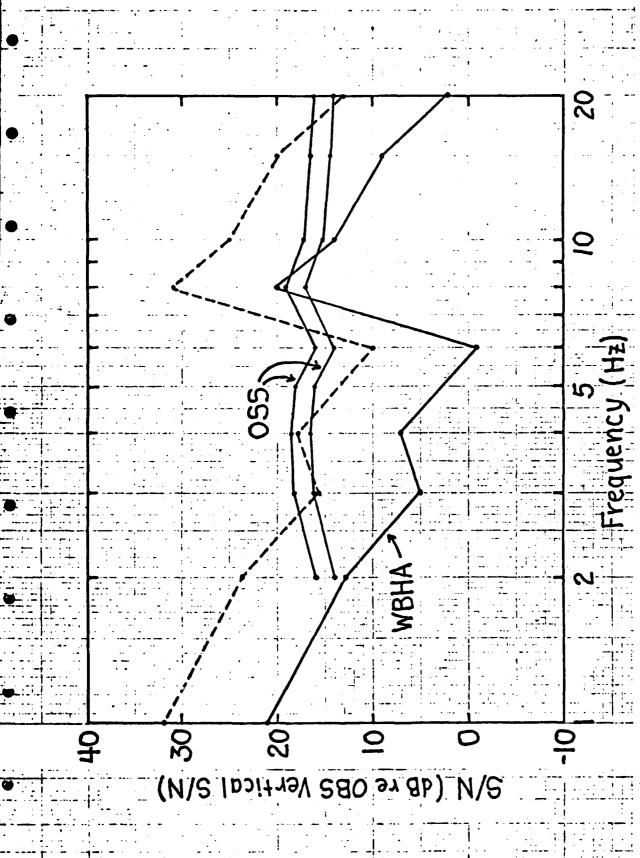
Also shown in Figure 1 is the average background noise for one of the Wake bottom hydrophones, determined from 52-2 minute random noise samples taken over a period of 18 months. Pressure has been converted to vertical particle velocity by the expression x=P/\rho c, where x is vertical particle velocity, P is the pressure, \rho is seawater density, and c is the speed of sound in seawater. Although the hydrophone noise is certainly not generated entirely by vertically arriving pressure waves (thus vertical particle velocity), this conversion is useful to determine an equivalent noise level for comparison.

For reference purposes, the noise curve for MSS in the Atlantic has also been plotted. It is not known if an impedance correction need be applied to this curve and so none has been.

The OBS and OSS noise curves may not be representative of average noise levels because they are only I sample and because of the ship generated noise. The relationship between these two curves however, is probably significant, and illustrates the decrease in noise level which can be achieved by a downhole instrument. The average noise level of the Wake bottom hydrophone is similar to that of the normalized OSS curve. These two curves are also comparable to the MSS curve for frequencies between about 4 and 10 Hz.

<u>Signal/Noise</u>. Although absolute noise levels are important for evaluating the usefullness of a seismic sensor, a more important parameter is S/N.

Ideally, a comparison should be made of an earthquake or explosion signal recorded simultaneously on the Wake hydrophones, an OBS, and a downhole instrument, all located in close proximity. Unfortunately, these data do



upper OSS curve is based strictly on noise data. The dashed curve is the WBHA curve described in the text. The lower OSS curve is based on shot S/N data, while the S/N measured on WBHA and on the OSS relative to S/N measured on nearby OBS's as raised by 11 dB to show the potential S/N by signal stacking. Figure 2.

not exist. With a few assumptions however, an indirect comparison can be made using data which do exist.

Within the 2.5 day period of OSS data previously mentioned is the record of a 126 pound shot at 80 km distance. It occurred, however, after the last of the nearby OBS's had been recovered. Fortunately, during the OBS deployment there was an 1800 pound shot also at 80 km distance from one of the OBS's. After making some assumptions and corrections, a comparison between ground arrivals on these records may yield information about the improvement in S/N between the OBS and OSS. The first correction is one for shot size, equal to the ratio of the shot weights to the power 0.66, or 15 dB. The spectral shape of the two shot records is very similar so a common shape was assumed. The second correction is for temporal changes in background noise which occured during the time between the shots. These effects were compensated for by assuming the same OBS/OSS noise ratios found for the simultaneous data discussed in the previous section. The resulting ratio of (OSS S/N)/(OBS S/N) is plotted in Figure 2.

Also shown is a curve which represents the possible gain in S/N of the OSS relative to the OBS based strictly on the simultaneous noise data. A correction for the impedance difference between the sediment and basalt has been applied. No assumptions were made about possible losses in upcoming signal levels between the OSS and OBS, and the effect of these would be to raise this curve by some amount. However, the near agreement in absolute level between this curve and the one based on the shot data may imply that these losses are not significant for the frequencies observed. (The parallel nature of the two curves is not significant but is due to having used the simultaneous noise data to correct for temporal differences in the

shot data.) On the average, these data support an improvement in S/N over the range 2-20 Hz of about 17 db.

A more direct comparison can be made between WBHA and an OBS. In August of 1981, HIG deployed an OBS array along a 1500 km line in the NW Pacific, and one of these OBS's was located within the bounds of WBHA. A strong Po (i.e., "Ocean P"; formerly called high-frequency Pn) signal from a large earthquake in the Kuril Islands was recorded simultaneously by this OBS and by a recording system connected to 3 of the hydrophones of WBHA. A comparison between the OBS vertical geophone and WBHA could not be made because that sensor was not working properly on the OBS. Consequently, the OBS hydrophone signal has been used. To correct for any difference in S/N between the OBS vertical and OBS hydrophone, a factor equal to the actual measured difference in S/N between the vertical geophone and hydrophone for this same signal recorded on another OBS at a closer epicentral distance was added to the S/N of the hydrophone. This factor averaged 1 dB over the range 1-20 Hz. Plotted in Figure 2 is the ratio between the average S/N of the 3 Wake hydrophones and the S/N measured on the OBS hydrophone with the previously described correction added. The somewhat wild behavior of this curve, such as the jump from -1 dB at 6 Hz to 20 dB at 8 Hz, is not entirely understood. It may be related to resonances in the OBS package which contribute to the OBS noise level, or it may be due to other factors which could not be identified by the analysis of only this one signal. For most frequencies, however, the WBHA data show an improvement in S/N, and this improvement averages roughly 10 dB over the range 1-20 Hz.

A natural question to ask is: "Why should there be any difference at all in S/N between the OBS hydrophone and the nearby WBHA hydrophones, since they are both lying at the top of the sediment column?" One explanation

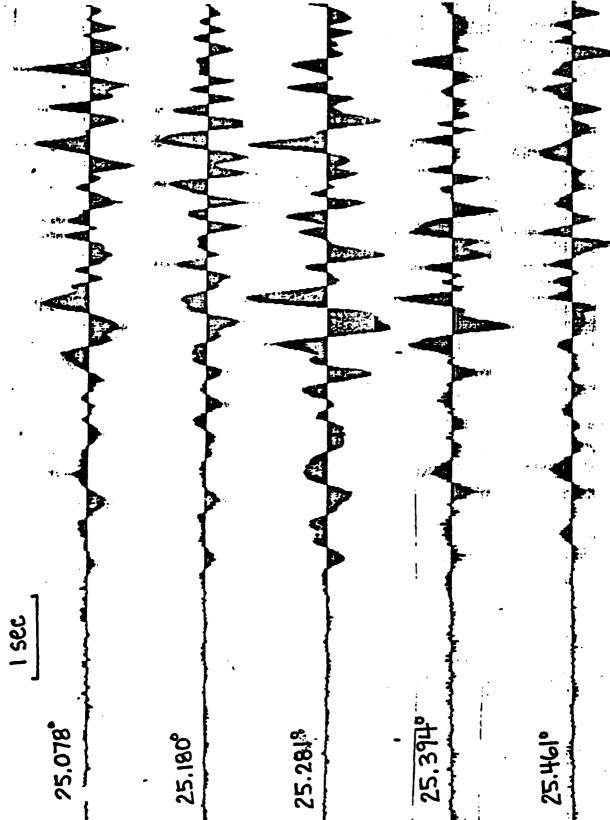
might be that the OBS used for comparison was particularly noisy due to noise sources in the water very close by or due to electrical noise in the amplification and recording system. Evidence which generally contradicts this hypothesis is that the noise levels measured on the other OBS's in the array were comparable. An alternate explanation may be related to the 20 year age of the array, and the possibility that the bottom hydrophones are now buried under a layer of sediment. Burial of the hydrophones would isolate them from noise generated by bottom water currents flowing around them, and also possibly from sediment surface phenomena such as Stonely waves. Although these explanations are not supported by direct evidence at this time, it would certainly be worthwhile to learn that shallow burial of a hydrophone under the sea floor sediment would greatly improve its S/N capabilities.

Coherence Across WBHA. An important aspect of current research at HIG is the investigation of coherence of teleseismic body waves across the 40 km aperture of WBHA. The impact of near receiver structure on these signals may be small due to the thin and relatively undisturbed nature of the oceanic lithosphere underlying WBHA. High coherency would have important implications for: (a) the planning of any future ocean seismic arrays; (b) studies of first arrivals requiring a great deal of timing precision; (c) studies requiring precise determination of seismic amplitudes (such as yield determination), and especially (d) studies of small signals where enhancement of S/N is desired (such as detection). Highly coherent signals may be added (with appropriate propagation delays) to yeild an increase in S/N equal to the square root of the number of sensors (for uncorrelated noise). At WBHA the theoretical increase in S/N would be equal to about 11

dB (3 dB of which comes from the first water surface reflection which is well established as having nearly perfect coherence with the initial arrival). Data from 3 WBHA hydrophones, collected prior to September 1982 on slow-speed cassettes, indicated high levels of coherence for a sampling of teleseismic P arrivals. Unfortunately, these data were not especially suited for this type of analysis because of timing inaccuracies and the suspect fidelity of the recording medium. Since September 1982, however, continuous data from 5 of the 6 bottom hydrophones have been collected digitally, essentially solving the timing and fidelity problem, and providing a format for a much speedier and more accurate analysis of the data. Software for the necessary quantitative analysis is still under development. However, "eyeball" checks for coherency have been made for a few of the digitally recorded teleseismic P phases with large S/N. The results are encouraging although not entirely conclusive.

Figure 3 illustrates the kind of broadband coherency which seems to be characteristic of the different P arrivals examined. Some features are "in phase" across the array while others are not. Certain pairs appear more coherent than other pairs. This relationship does not seem to be a function of epicentral distance or azimuth. Certainly a large improvement in S/N would not be expected from the stacking of these signals as they are.

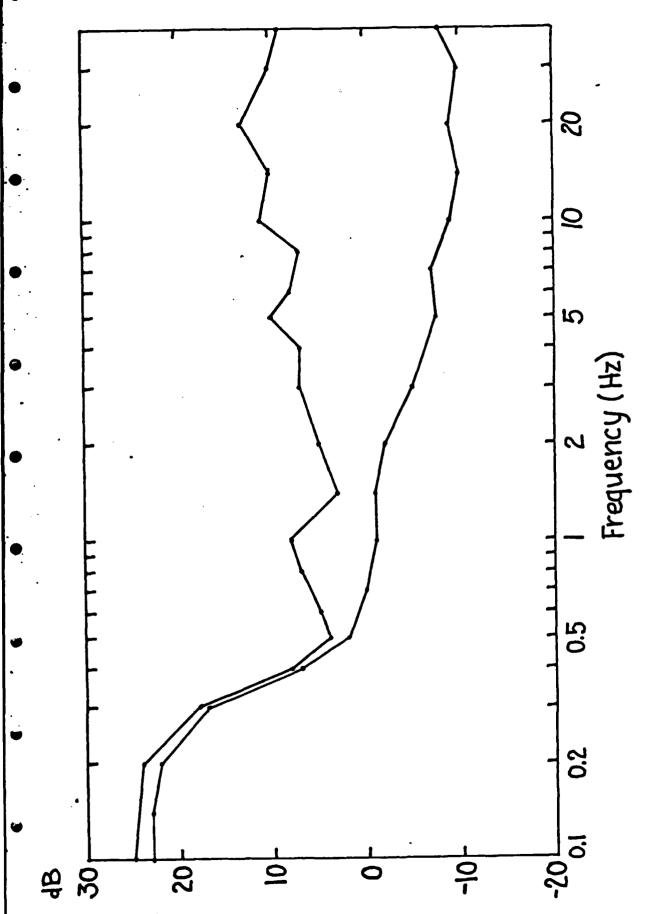
Possible reasons for the observed difference in signals may be classified into the following categories: (a) major differences in the coda exist for very small changes in takeoff angle (for Figure 3, much less than 1°); (b) variations in lithospheric structure directly beneath the array have a significant impact on the signals; and (c) the hydrophone/cable responses differ significantly across the array. If the



Unfiltered first arrivals of P across the Wake bottom hydrophones from an earthquake south of Honshu, Japan (6 Sept. 1981, h=167 km, m=6.6). Traces have been offset by travel time differences. Figure 3.

explanation is (a) then little, if any, improvement in coherence can be expected by massaging the data. If the explanation is (b) then significant improvements in coherence are possible, provided the lithospheric response does not vary greatly with small changes in azimuth or angle of incidence of incoming P arrivals. If the explanation is (c) then an even more significant improvement in coherence is possible. The technique which will be used to investigate (b) and (c) is basically a calculation of amplitude and phase differences between hydrophones for a given signal as a function of frequency. If (c) is true, then these differences should be consistent regardless of what kind of signal is being recorded - P, Po, So, T, or noise. If (b) is true, then the differences should be consistent for P and possibly Po and/or So. If both (b) and (c) are true, the effects should be seperable. These effects may then be removed by a deconvolution of the observed signals with the empirically determined impulse responses of the lithosphere and hydrophone/cable. Determination of these responses to a level of statistical signifigance will require the recording and analysis of a variety of signals at assorted azimuths and epicentral distances which have strong S/N. Evidence exists to suggest that at least (c) is true.

Figure 4 shows amplitude differences observed across WBHA for a 10 minute section of record which includes a P, Po, and So signal. The upper curve is the absolute value of the maximum difference in amplitude observed between the six hydrophone signals as a function of frequency. Energy below about 0.5 Hz is from microseisms while between 0.5 Hz and 20 Hz it is mostly an average of the P, Po, and So. The lower curve shows amplitude differences between two particular hydrophones. Although it may seem unusual to have combined energy from all of these different types of signals, the individual signals, P, Po, So, and noise, exhibit the same



Plot of differences between spectra across the Wake bottom hydrophone array for a 10 minute section of record containing P, Po and So signal and noise. The upper curve shows the The lower curve shows maximum value of variation in level between all of the hydrophones. the difference in level between two of the hydrophones. Figure 4.

relationships although with a somewhat greater scatter. This strongly suggests that there are at least differences in the hydrophone/cable responses which are removable. These differences might easily produce the incoherencies seen in Figure 3. Further and more careful analysis of other events will necessary to precisely determine the differences in amplitude and especially phase. The reward may be stronger coherency and thus an increase in S/N.

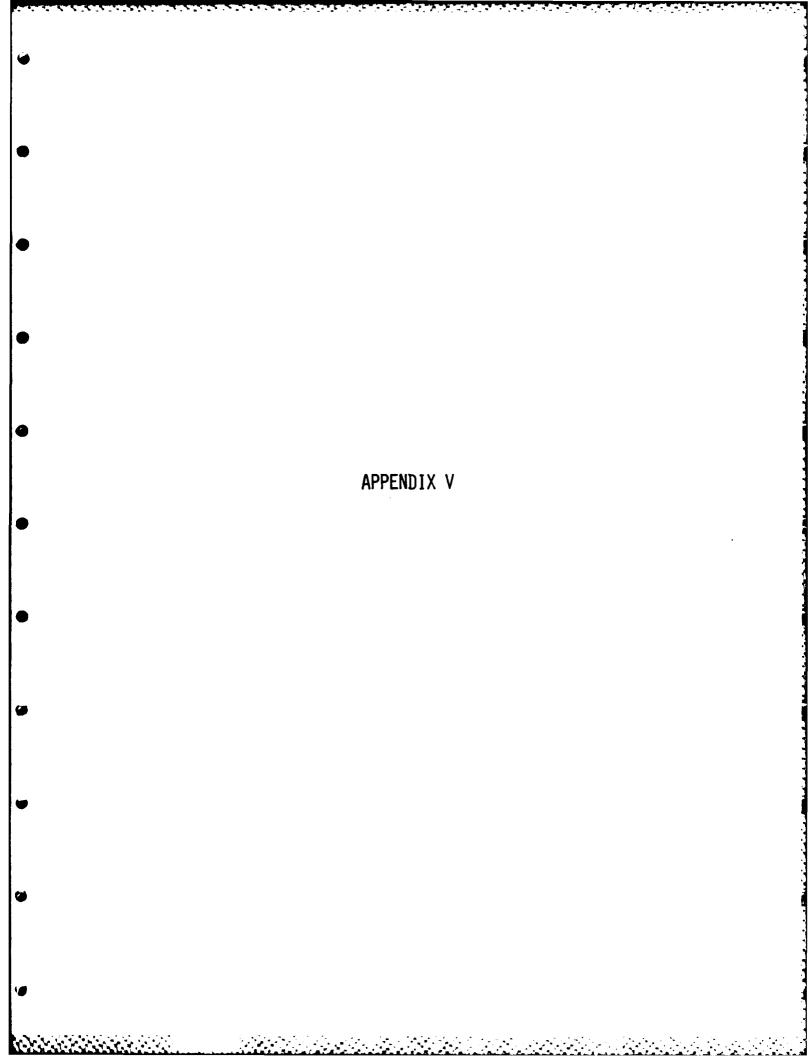
Conclusions. The data presented in this report indicate that WBHA, the OSS, and, by inference, the MSS, are seismic tools of comparable quality. Noise levels measured were similar for all three although the data were somewhat limited. S/N ratios measured for OSS were greater than that for individual WBHA hydrophones by an average of about 7 dB. Gains in S/N levels, through stacking of the WBHA signals, will reduce, if not eliminate or reverse, that difference. Precise estimation of the coherence across WBHA will not be made until recently discovered differences in the hydrophone/cable responses have been more accurately determined so that their effects may be removed from the signals.

Many of the comparisons in this report were made via rather indirect routes due to the limited amount of data. Additional data should become available sometime this summer after the OSS recording package has been picked up and the data returned, reduced, and distributed. More data from the August 1981 Wake OBS Experiment will be in a reduced form soon. Also, data from OBS's deployed during the OSS experiment by Oregon State University should be available later this year. All of these future data should be examined and added or compared to those data presented here to

further understand these results.

Acknowledgements. This preliminary and informal report was sponsored by the Air Force Office of Scientific Research, with supplementary funds provided by the U.S. Arms Control and Disarmament Agency.

Major upgrading of the Wake system in September of 1982, just prior to the Downhole Experiment, was made possible through the timely support of those agencies.



Po/So Phases: Propagation Velocity and Attenuation Across a 1600 Km Long Deep Ocean Hydrophone Array

Ъy

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Abstract. Po/So phases from numerous earthquakes along the margin of the Northwestern Pacific Basin were successfully recorded by a linear, 1500 km ocean bottom hydrophone (OBH) array deployed for two months near Wake Island. Data from ten shallow-focus (<100 km) events at 180 (2000 km) to 33° (3700 km) epicentral distance were used to compute propagation velocities for Po and So. The resulting travel-time equations T=X/(7.962 0.05 km/sec) - $(7.14 \pm 2.38 \text{ sec})$ and $T=X/(4.57 \pm 0.04 \text{ km/sec})$ - $(14.03 \pm 5.31 \text{ sec})$ sec) for Po and So, respectfully, successfully model all Northwestern Pacific Basin shallow-focus Po/So first arrival data collected by the Hawaii Institute of Geophysics since 1963 at epicentral distances greater than 12°. Also, arrivals from an event in the Kuril Islands (mb = 6.6, h = 45 km), which propagated down the axis of the array, were used to estimate the attenuation of Po and So between 25° (2800 km) and 33° (3700 km) epicentral distance. After a correction for cylindrical spreading, values for a frequency dependent Q are found to range from 625 ± 469 at 2 Hz to 2106 ± 473 at 13 Hz for Po and from 1401 ± 296 at 5 Hz to 3953 ± 863 at 15 Hz for So. These same data may be alternatively described as exhibiting an average attenuation of -21.5 ± 0.9 dB per 1000 km of travel path, which applies to both Po and So at all frequencies studied.

Introduction

Late in the summer of 1981, the Hawaii Institute of Geophysics (HIG) successfully deployed a 1500 km long linear array of twelve ocean bottom hydrophones (OBH's) near Wake Island (Fig. 1). Half of the instruments (indicated by open circles) started recording on 12 August and ended on 23 September. The remaining half started on 3 September and ended on 15 October. The total recording time was about 65 days, with all twelve instruments in operation from 3 September through 23 September. Of the twelve OBH's deployed, ten were successfully recovered, and nine of these had quality data throughout their operational period. Recorded concurrently were three bottom hydrophones of the Wake Hydrophone Array (WHA), a 40 km array of sensors located near, and cabled directly to, Wake Island. WHA data were used in the study of first arrivals, and were considered to be part of the OBH array data. All of the OBH and WHA instruments were in ocean depths between 5265 m and 5657 m. Of more than 130 events identified from seismic phases in the recordings, 108 were located by the National Earthquake Information Service (NEIS; Table 1 and Fig. 1). The primary purpose of the experiment was to acquire data of critical importance in understanding a phenomenon known as high-frequency Pn/Sn, long-range Pn/Sn, or Ph.f./Sh.f.; but referred to here as Po/So or Ocean P/Ocean S after Walker (1982).

Po/So phases were first observed in the North Atlantic and have been found throughout the North, Western, and Central Pacific. First arriving Po/So energy travels with a fairly constant apparent velocity (epicentral distance/travel time) of about 8.0 and 4.6 km/sec, respectively, while peak amplitude arrivals have apparent velocities of about 7.6 and 4.5 km/sec, respectively, which are comparable to basal crustal rates. At distances of about 18° (\cong 2000 km), observed frequencies of Po/So are as high as 30 and 35 Hz, respectively; and at distances of of about 30° (2 3300 km), as high as 15 and 20 Hz, respectively. The signal/noise ratios for Po/So phases are generally at least ten times greater than the ratios of their respective normal, mantle-refracted P and S phases; and in many instances no P's or S's can be found in spite of the presence of very strong Po's and So's. Aside from the SOFAR channel of the world's oceans, the Po/So waveguide appears the be the earth's most efficient acoustical waveguide. Also, it seems probable that the phenomenon is a dominant feature of all of the world's oceans and marginal seas.

Recent efforts have been made, using synthetics, to determine the mechanism of Po/So propagation (e.g., Stephens and Isacks, 1977; Menke and Richards, 1980; Sutton and Harvey, 1981; and Gettrust and Frazer, 1981), although many essential characteristics of these phases are still poorly known. Data collected before this experiment, composed almost entirely of events recorded at single stations and over an unevenly distributed range of epicentral distances, made difficult the precise determination of the travel-time curve between 0° and 40°, as well as the attenuation of these phases as a function of frequency and distance. These parameters need to be accurately determined so that a unique model can eventually be found.

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To exemplify this point consider Fig. 2: a plot of apparent velocity versus epicentral distance for Po and So having Northwestern Pacific travel paths and source focal depths of 100 km or less which is based on data collected and published (Walker, 1981) by HIG before this experiment. Two

Table 1. Events Recorded by the OBH Array (Data from NEIS Monthly Lists)

No.	Date		Time	Location	h	m	M	n	Po	Sc
1	Aug.	12	22:35	Solomon Is.	33	4.9	_	3	_	_
2		13	02:57	Tonga Is.	191	5.4	-	3	_	_
3		14	05:42	Volcano Is.	33	4.7	_	3	_	_
4		14	06:24	Molucca Passage	38	5.5	5.3	1	_	_
5		14	09:05	Honshu, Japan	58	_	_	2	-	_
6		15	10:30	Alaska	53	5.1	_	2	_	-
7		15	19:53	Philippine Is.	152	4.8	_	1	_	_
8		16	23:54	Kuril Is.	33	5.6	5.0	3	_	_
9		17	02:17	West Irian	34	5.7	5.8	3	_	-
10		17	17:07	Fiji Is.	383	5.5	_	2	-	-
11		18	05:29	Banda Sea	34	5.1	4.8	1	_	_
12		19	01:41	Kermadec Is.	185	5.1	_	1	_	_
13		19	03:01	Fiji Is.	507	4.8	_	1	-	_
14		19	06:06	Loyalty Is.	25	5.6	4.9	1	_	-
15		20	02:19	Santa Cruz Is.	71	5.0	- .	3	-	_
16		20	12:53	Mariana Is.	193	3.8	_	4	_	_
17		20	15:10	Kermadec Is.	347	4.9	_	1	-	_
18		21	14:29	Bonin Is.	509	4.6	-	1	_	-
19		23	01:59	Loyalty Is.	100	5.8	_	1	_	_
20		23	12:00	Kuril Is.	40	6.0	5.8	4	3	3
21		24	15:46	Aleutian Is.	56	5.2	_	2	_	_
22		25	06 : 56	Honshu, Japan	325	4.8	_	4	-	-
23		25	07:16	Tonga Is.	33	5.9	5.7	1	_	_
24		25	07:22	Tonga Is.	33	5.7	-	1	_	-
25		25	20:07	Mariana Is.	33	4.9	4.4	1	_	-
26		26	04:51	Mariana Is.	40	5.2	5.1	5	_	_
27		26	16:32	New Britain	74	5.7	_	6	_	-
28		26	18:56	Honshu, Japan	230	4.4	_	3	-	-
29		28	09:04	Alaska	71	5.1	_	3	-	-
30		30	11:36	Fiji Is.	609	5.4	-	6	_	-
31		31	06:14	Komandorsky Is.	33	4.7	3.8	1	_	_
32	Sept.	1	07:23	Tonga Is.	33	5.8	5.7	4	_	_
33	•	1	09:29	Samoa Is.	25	7.0	7.7	6	_	_
34		1	18:38	Tonga Is.	33	5.7	5.3	3	_	_
35		1	23:55	Tonga Is.	33	5.6	5.4	3	-	_
36		2	08:44	Samoa Is.	33	5.3	5.5	ì	-	_

Legend: h - event depth (km); m - body-wave magnitude; M - surface-wave magnitude; n - number of instruments in array which recorded body-wave phases P, Po, or So from this event; Po - number of Po arrivals from this event used in first-arrival study; So - number of So arrivals from this event used in first-arrival study.

No.	Date	Time	Location	h	m	М	n	Po	So
37	Sept. 2	09:24	Honshu, Japan	58	5.5	_	2	-	-
38	3	03:59	Hokkaido, Japan	52	4.7	-	4	-	_
39	3	04:29	Philippine Is.	93	5.8	-	4	_	-
40	3	05:35	Kuril Is.	45	6.6	6.6	10	6	8
41	3	19:39	Honshu, Japan	44	5.6	5.4	4	-	-
42	3	19:44	New Guinea	136	4.8	-	5	_	-
43	4	11:15	Philippine Is.	645	6.0	_	10	-	-
44	4	23:44	Solomon Is.	38	5.4	5.3	9	_	_
45	6	11:02	Loyalty Is.	31	5.9	6.2	1	_	-
46	7	15:11	Fiji Is.	231	5.2	-	3	-	_
47	7	16:20	Honshu, Japan	440	4.9	-	3	_	_
48	7	19:06	Honshu, Japan	33	5.8	5.5	7	6	5
49	7	20:07	Honshu, Japan	29	5.1	4.7	7	5	4
50	8	19:26	Kuril Is.	46	5.7	5.4	5	_	3
51	10	03:29	Solomon Is.	110	4.9	_	3	_	_
52	10	23:21	Mariana Is.	13	5.6	5.2	10	5	_
53	11	08:33	Fiji Is.	554	5.2	_	5	-	_
54	12	03:40	Fiji Is.	302	5.2	_	3	-	_
55	12	07:15	Kashmir	33	6.2	5.9	3	-	_
56	12	14:51	Hokkaido, Japan	111	4.9	_	7	-	_
57	12	16:14	Bonin Is.	33	4.6	-	2	_	-
58	13	01:20	Honshu, Japan	39	4.8	4.8	4	_	_
59	13	02:17	Eastern Kazakh	0	6.0	4.5	4	_	_
60	13	20:24	Honshu, Japan	87	4.9	-	3	-	_
61	14	15:08	Honshu, Japan	33	5.4	5.0	4	4	_
62	15	14:12	Banda Sea	102	5.9	-	2	-	_
63	15	20:43	Taiwan	167	4.9	-	3	_	_
64	17	06:19	Banda Sea	33	5.7	5.8	1	_	_
65	17	08:23	Loyalty Is.	30	5.7	6.6	2	_	-
66	17	12:42	Fiji Is.	356	5.2	-	10	_	_
67	17	21:12	Kamchatka	33	4.9	3.9	2	_	_
68	19	07:27	China	561	4.4	_	3	_	-
69	20	04:39	Ronshu, Japan	33	4.4	-	2	-	_
70	22	06:55	Mariana Is.	33	4.2	-	6	_	_
71	24	17:20	Bonin Is.	33	5.7	5.3	5	3	2
72	25	03:25	New Guinea	116	4.8	_	1	-	_
73	25	10:21	Bonin Is.	33	4.4	-	2	-	_
74	25	14:30	Kermadec Is.	45	5.9	5.9	5	-	_
75	25	15:01	Honshu, Japan	25	5.5	6.1	2	_	-
76	28	03:36	Honshu, Japan	31	5.5	5.3	5	_	_
77	28	17:56	Kermadec Is.	323	6.0	-	5	-	_
78	29	23:02	Tonga Is.	226	5.1	_	5		

Legend: h - event depth (km); m - body-wave magnitude; M - surface-wave magnitude; n - number of instruments in array which recorded body-wave phases P, Po, or So from this event; Po - number of Po arrivals from this event used in first-arrival study; So - number of So arrivals from this event used in first-arrival study.

No.	Date	Time	Location	h	D	M	n	Po	So
79	Sept. 30	07:04	New Guinea	123	5.4	-	4	-	_
80	30	23:03	Pacific Ocean	10	5.9	5.2	2	_	_
81	30	23 : 37	Hokkaido, Japan	52	4.9	4.9	2	-	-
82	Oct. 1	12:14	Novaya Zemlya	0	5.9	3.8	4	-	-
83	1	13:10	New Ireland	85	5.0	_	2	-	-
84	1	16:02	Kermadec Is.	33	5.6	5.1	5	_	_
85	1	17:04	Kuril Is.	33	5.9	5.7	5	4	4
86	1	19:00	S. Nevada	0	4.9	_	2	_	_
87	2	05:51	Savu Sea	109	4.6	-	4	-	_
88	2	15:13	Mariana Is.	113	5.0	_	5	-	_
89	3	07:21	Kuril Is.	75	5.1	_	4	_	_
90	3	08:26	Hokkaido, Japan	62	4.4	_	2	-	-
91	3	16:48	Philippine Is.	238	5.0	_	2	-	-
92	4	00:01	New Guinea	33	5.9	6.3	4	-	-
93	4	04:11	Honshu, Japan	38	5.2	5.0	4	-	2
94	4	10:18	Solomon Is.	58	5.1	_	2	-	_
95	4	10:27	Solomon Is.	24	5.7	5.4	5	-	_
96	5	04:28	China	534	4.7	-	2	-	_
97	6	07:40	Aleutian Is.	23	5.2	4.3	3	_	_
98	7	03:02	Fiji Is.	620	5.8	-	5	-	-
99	7	08:32	Solomon Is.	41	5.8	5.3	5	-	_
100	7	15:03	Samoa Is.	33	4.8	-	2	-	_
101	7	17:48	Philippine Is.	595	5.3	-	4	-	_
102	9	12:19	Solomon Is.	50	6.0	6.4	5	~	_
103	9	19:46	Mariana Is.	96	4.7	-	5	_	_
104	11	00:36	Minahassa Pen.	94	5.6	_	4	_	_
105	13	15:53	Kamchatka	112	5.3	-	3	_	_
106	14	12:29	Mariana Is.	205	5.0	_	5	-	_
107	14	20:09	Mariana Is.	131	4.9	-	5	-	_
108	15	01:47	Honshu, Japan	47	6.0	5.4	5	_	_

Legend: h - event depth (km); m - body-wave magnitude; M - surface-wave magnitude; n - number of instruments in array which recorded body-wave phases P, Po, or So from this event; Po - number of Po arrivals from this event used in first-arrival study; So - number of So arrivals from this event used in first-arrival study.

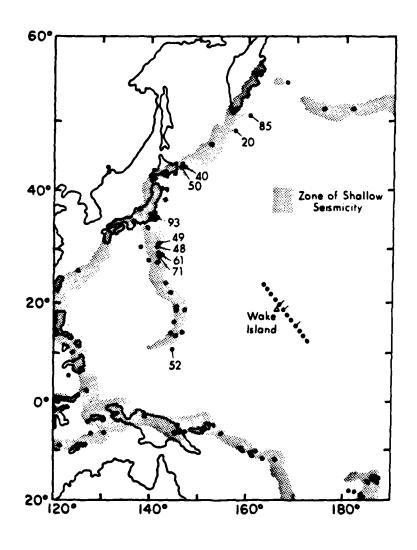
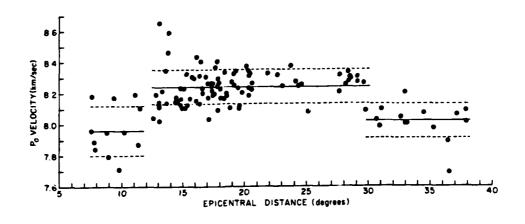


Figure 1. Map showing OBH locations and epicenters of earthquakes recorded by the array (see Table 1). Instruments operating from 12 August through 23 September are indicated by open circles. Those operating from 3 September through 15 October are indicated by closed circles. Three instruments which did not work properly are flagged. Event numbers (from Table 1) are indicated for epicenters of events specifically used in this study.



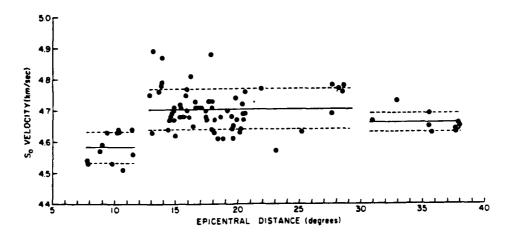


Figure 2. Apparent velocity (X/T) versus epicentral distance (X) for Po and So with Northwestern Pacific travel paths and focal depths of 100 km or less. Note the discontinuities at 12° and 30°. Solid and dashed lines represent the mean apparent velocity plus and minus one standard deviation over the three distance ranges. This plot is based on data compiled before the OBH experiment.

significant features of this plot are the lower apparent velocity values at distances less than 12° and greater than 30° in comparison to values within the 12° to 30° range. The large amount of scatter coupled with insufficient sampling near 12° and 30° , however, make the exact nature of these velocity transitions unknown. For the purpose of modeling, it is important to know if these transitions are smooth or abrupt, and might therefore represent a gradual increase in velocity with depth or a discontinuity in velocity at some depth.

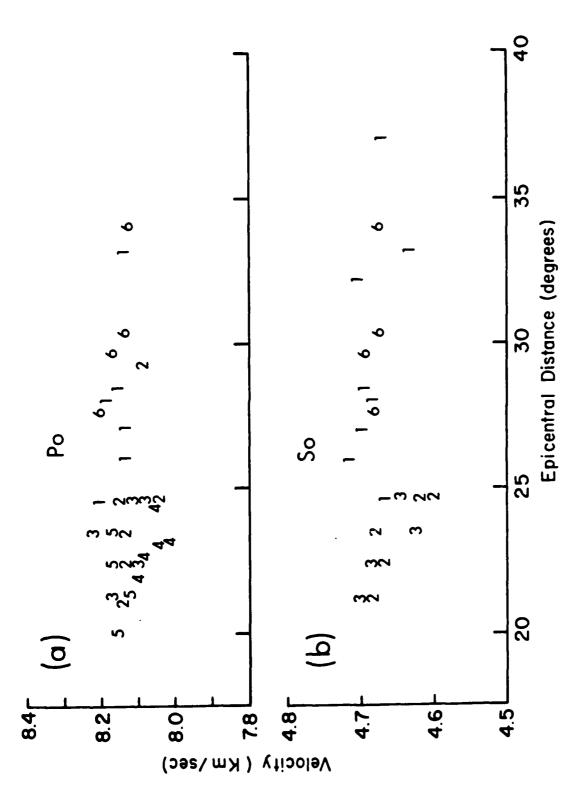
Similar difficulties arise when attempting to use older data to estimate the attenuation of Po/So energy as a function of travel-time (or distance). The high frequencies and large signal/noise ratios observed would indicate that their attenuation must be significantly smaller than that of normal, mantle-refracted P or S phases at similar distances. However, these observations might also be due in part to a frequency-dependent Q combined with low noise levels at the higher frequencies (> 4 Hz) where Po/So are most prominent. Unfortunately, the mix of data (i.e., different earthquakes recorded at only single stations) make it hard to separate attenuation from other factors which influence the observed amplitudes such as magnitude, focal depth, epicentral distance, source orientation, and possibly azimuth.

To help clarify these uncertainties, an experiment was designed to record Po/So across a linear, 1500 km array consisting of 12 OBH's. The data provided by such an experiment would permit determination of Po and So phase velocities independent of the source parameters, thus hopefully eliminating the major source of scatter in velocity estimates. Data from the experiment would also facilitate a nearly direct measure of attenuation as a function of frequency and travel-time (only a spreading term would have to be assumed in order to determine the apparent Q). In addition, any other changes in the Po/So coda would for the first time be observed at several distances along approximately the same azimuth.

The array was aimed towards northern Japan and the southern-most islands of the Kuril chain (Fig. 1), and was positioned to provide data across the 30° transition zone previously described. This target was chosen because of its long history of moderate-to-large earthquakes and several successful recordings of Po/So phases from this region on the hydrophone installation at Wake Island. Of the events recorded by the OBH array in the Marianas through Kuril portion of the circum-Pacific arc, the largest (a 6.6 mb; event 40 in Table 1) occurred in the target area while all of the instruments were recording.

Po/So Propagation and Apparent Velocity

To confirm the assumption that much of the scatter in Fig. 2 can be attributed to errors in the epicenters and/or origin times of these single station data points (i.e., no more than one station recorded any given earthquake), a similar plot has been constructed for some of the OBH data (Fig. 3). Although this data as a whole exhibits scatter of about the same magnitude as that shown in Fig. 2, there is significantly less scatter within each subset of data representing individual events at several OBH's. Thus, the larger source of scatter may be attributed to apparent velocity differences between events; and a strong possibility is that these

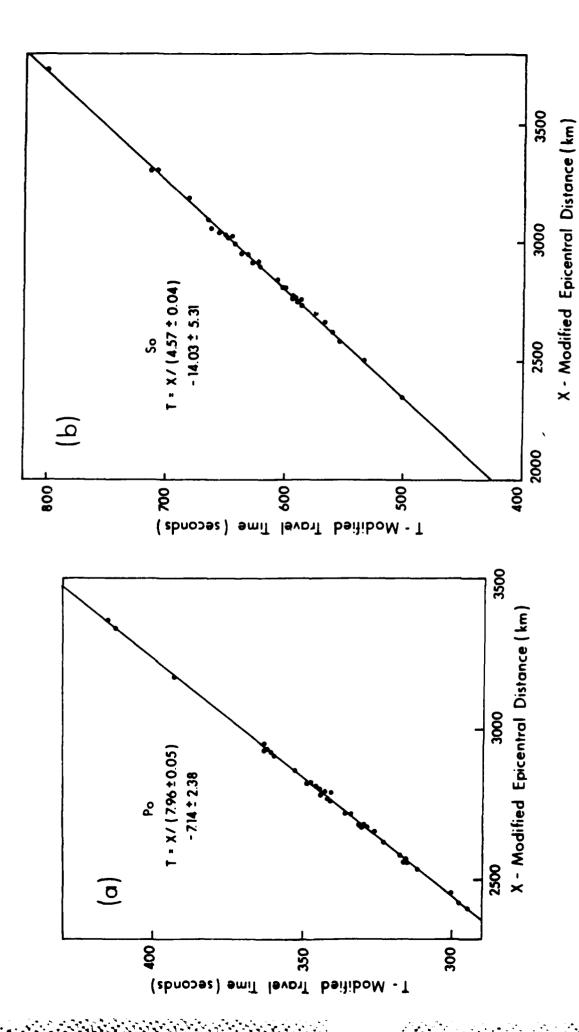


Apparent velocity (X/T) versus epicentral distance (X) for some arrivals recorded across the OBR array from events 40. 48. 49 52 61. and 85. respectively. Note that the scatter of each Data points numbered 1 through 6 represent each set of first of the Po and So data collected during the OBH experiment. subset is less than the scatter of the combined data. Figure 3.

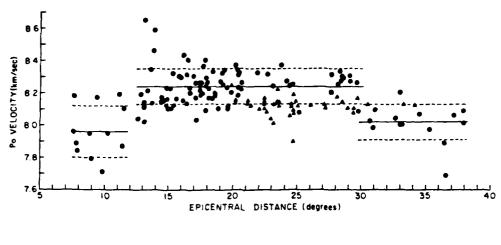
differences are a result of errors in NEIS origin times and hypocenter locations. The remaining scatter within each subset may be the result of: (1) uncertainty in the time of onset, (2) OBH location errors, (3) OBH clock correction errors, and/or (4) local differences in crustal and sediment structure between OBH's. Unfortunately, the magnitude of this remaining scatter is enough to cause an unacceptable level of uncertainty in the propagation velocity of Po or So for any given event. An uncertainty of 1 sec in travel-time over 500 km of the array at a phase velocity of 8.0 km/sec will produce an uncertainty of 0.13 km/sec in the determination of that phase velocity, as well as an uncertainty of 5.0 sec in the travel-time intercept for data taken at an average epicentral distance of 2500 km. To reduce this uncertainty, a method was sought to determine propagation velocity and travel-time intercept by combining the first arrival data of all the events in a way which would have the following important properties: (1) propagation velocity would be determined from the differences in travel time between OBH's for a given event and would be independent of the NEIS origin time published for the event; (2) events recorded on more instruments would be weighted more heavily; (3) events recorded over a larger range of epicentral distance would also be weighted more heavily; and (4) the travel-time - epicentral distance relationships within the data set of each individual event would be maintained. The following method satisfactorily meets those requirements.

Using only those events with focal depths of 100 km or less, having Northwestern Pacific Basin travel paths, and recorded with distinct onsets at two or more stations (see Table 1), the raw epicentral distance and travel time data were computed using NEIS epicenters. For each subset of data associated with a single event, the mean epicentral distance and travel time was computed and subtacted from the raw data for that event to yield zero-meaned data for each event, as well as for the data as a whole. From these combined data the slope of the travel-time line could be computed, but the intercept would be lost (it would be exactly zero). Consequently, the mean for all of the raw epicentral distance and traveltime data was computed and added to each point of the zero-meaned data set to restore the intercept. The travel-time lines computed from these data, and shown in Fig. 4, represent the best available estimates of Po/So velocities at distances between about 2400 and 3400 km. These values (i.e., inverse slopes) are 7.96 to 0.05 km/sec and 4.57 to 0.04 km/sec for Po and So, respectively. The large negative intercepts found, -7.14 2.38 sec and -14.03 2 5.31 sec for Po and So respectively, imply that the first arriving energy propagates at a higher velocity than indicated by the inverse slopes over some portion of the travel path nearer to the source than where the observations were made.

In Fig. 5, we have superimposed Po/So data from the Wake OBH experiment (solid triangles) on Fig. 2. It is obvious that these data do not support the discontinuity in apparent velocity at 30° which was a feature of the older data; but they appear, instead, to describe a smooth transition across this boundary. A re-examination of the older data, especially those data points near the offset, was made to determine if the cause of this discrepancy could be found. Consequently, the original computer cards used to calculate epicentral distances for data collected in 1963 and 1964 were found to contain systematic errors in the coordinates of some recording sites. The formerly classified status of these sites and



Modified travel-time plots for Po and So as described in the text. Figure 4.



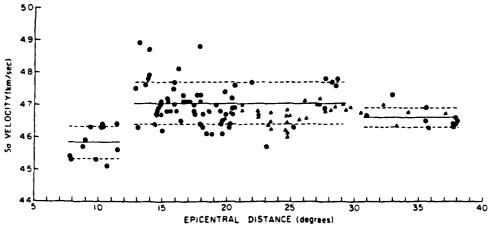


Figure 5. Po and So data from the OBH experiment (solid triangles) superimposed on Fig. 2a and Fig. 2b, respectively. Note that these data do not support a discontinuity in apparent velocity at 30°.

the circuitous route by which these coordinates had been previously obtained may be an explanation for these errors. The erroneous data have been plotted as open circles and the corrected data as solid circles (along with the rest of the older data) in Fig. 6. Also plotted in Fig. 6 (as open triangles) are some additional hydrophone data from the analog cassette recording system at Wake, collected since 1978. It is clear from

this comprehensive data set that an apparent velocity discontinuity at 30° is no longer supported by the data. Instead, a gradual decrease of

apparent velocity with distance is observed between about 12° and 38° for both Po and So. This characteristic of apparent velocity is entirely compatable with the inverse slope propagation velocity and intercept values computed from the OBH data and shown in Fig. 4.

To illustrate this compatability, the travel-time equations have been converted into relationships between apparent velocity and epicentral distance as follows:

$$T = X/V + I$$

Travel-time equation where T is travel-time, X is epicentral distance, V is propagation velocity, and I is the travel-time intercept.

$$A = X/T = X/(X/V + I)$$

Apparent velocity equation where A is the apparent velocity.

$$s_{A}^{2} = \frac{(x/v+1)^{2}v^{4}s_{x}^{2} + x^{4}s_{y}^{2} + x^{2}v^{4}(s_{1}^{2}+s_{0}^{2})}{v^{4}(x/v+1)^{4}}$$

Mean-squared error in

apparent velocity, S_A,
resulting from errors in:
the propagation velocity,
S_V; the travel-time
intercept, S_I; the
epicentral distance, S_X, due
to epicenter errors; and
the observed travel-time,
S_O, due to errors in origintime.

Plotted in Fig. 7 are A, and A $\stackrel{+}{=}$ S_A versus epicentral distance, X. Values for V, S_V, I, and S_I are those given in Fig. 4; S_X is 10 km; and S_O is 1 sec. Values used for S_X and S_O are reasonable for epicenter/origin-time data taken from the NEIS listings, and these lists have been used for all of the data in the figure. The figure shows that approximately one standard deviation of the data (ie. 68 percent for normally distributed

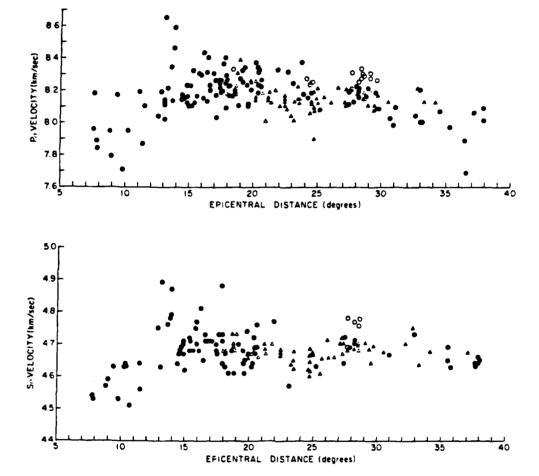


Figure 6. This figure is the same as Fig. 5 except that: (1) erroneous data from 1963 and 1964 (open circles) have been replotted as closed circles; and (2) addition hydrophone data from Wake. collected since 1978. are plotted as open triangles.

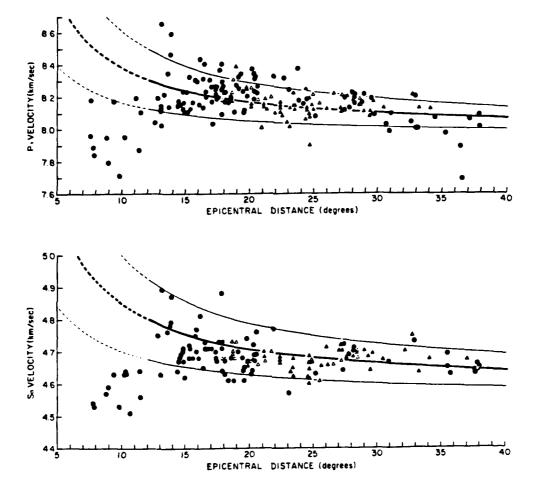


Figure 7. This figure contains the data from Fig. 6. representing all of the Po and So first arrival data from shallow-focus events with Northwestern Pacific Basin travel paths collected by HIG since 1963. Superimposed on these data are curves representing the Po and So travel-time equations (Fig. 4). with plus and minus one standard deviation. determined strictly from the OBH data.

Note that most of the data beyond 12° epicentral distance falls within the bounds described by these curves.

data) falls within the bounds of A S_A . It is concluded, therefore, that the first arrival times of Po and So phases from shallow-focus (i.e., focal depths less than 100 km) earthquakes occurring along the margin of the Northwestern Pacific Basin with travel paths greater than 12° epicentral distance across the Northwestern Pacific Basin, may be successfully modeled by the travel-time equations given in Fig. 4.

An important question which remains to be answered is: By what methods of generation and propagation are the first arrivals on both sides of the 12° apparent velocity discontinuity produced? In spite of those modeling efforts mentioned previously, which have successfully reproduced certain aspects of Po/So travel-time and coda, no model as yet exists which addresses this question or reproduces the observations which led to this question. A hypothesis proposed here, which seems to contain some credibility, is that energy observed at less than 12° has propagated from near the source to near the receiver entirely within the waveguide (i.e., up the descending lithosphere and across the plate), while first-arriving energy observed at greater than 120 has propagated along a higher velocity, or deeper, P or S type path over distances less than 120 before coupling into the waveguide at a distance greater than 120. (The term "waveguide" refers here to the structure within which Po/So energy propagates at the constant velocity value given in Fig. 4.) Evidence which supports this hypothesis is: (1) the apparent velocity data observed for Po and So at distances less than 120 could reasonably be fit by a zero-intercept linear travel-time equation with a propagation velocity equal to the propagation velocity found for the data beyond 120 (This would imply that the propagation velocity and thus the atructure of the waveguide are continuous across 120 - an intuitively pleasing result.); and (2) the apparent velocities for shallow-focus P and S at 120 (from Jeffreys and Bullen, 1958) are near those required for P and S to couple into the Po/So waveguide and produce the observed arrival times beyond 12°. Although a propagation mechanism of this type might fit the data quite well, a velocity model of the oceanic crust and upper mantle which produces this phenomena has not been found. It is probable that such a velocity model would contain lateral heterogeneities associated with the downgoing slab in order to propagate energy between 0° and 12° with an average velocity greater than the waveguide velocity, and still couple this energy into the waveguide. This requirement is not easy to satisfy (and may be impossible to satisfy) with a radially symmetric velocity model.

One test of this proposed hypothesis could be made by examining Po/So arrivals from earthquakes with depths greater than 100 km. Po/So energy has been observed and recorded by HIG for many earthquakes with focal depths greater than 100 km and up to 600 km. Energy which has propagated along a P or S type path before coupling into the waveguide should contain a depth dependence in travel-time, coda, and/or frequency content. The precise nature of such a dependance would serve to either confirm or deny the hypothesis, or to generate a new one.

Attenuation of Po and So

As stated earlier, one of the more remarkable features of Po/So is their high-frequency content; therefore, a major goal of Po/So research in recent years has been to quantify those properties within the earth in terms of Q which permit the efficient transmission of these frequencies over great distances. When a frequency-independent model for Q was fit to some older Po/So data with frequencies between about 2 and 10 Hz, the resulting Q values were greater than 5000 in most cases (Walker et al., 1978 and Sutton et al., 1978). These values are much higher than those generally found for upper mantle travel paths using P and S data at frequencies below 2 or 3 Hz. A frequency-dependent Q, however, might satisfy both sets of observations. Sutton et al. (1983) found two methods for extracting a relationship between Q and frequency from single-station Po/So data. They found that Q was approximately proportional to frequency over the range 2-9 Hz and that the Q of So was higher than the Q of Po for any given frequency.

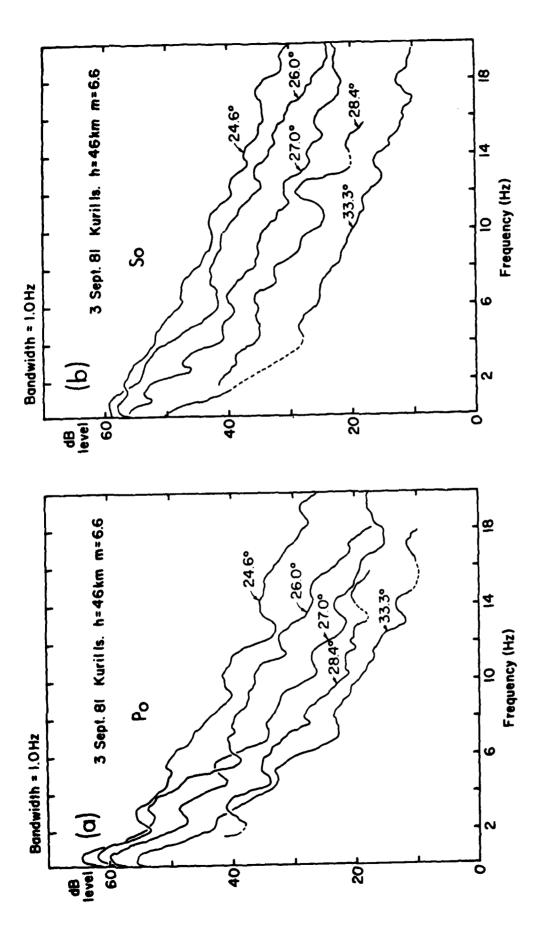
Data collected during the OBH experiment provided, for the first time, the opportunity to observe Po/So energy propagating across a large array of instruments. Although Po/So phases from numerous events were recorded, the magnitude 6.6, Kuril Is. event noted earlier (event 40, Table 1) was especially well suited for a study of attenuation. Its large magnitude and its location in the area targeted by the array produced Po/So phases which travelled down the axis of the 1500 km deployment line with sufficient energy to be observed with high signal/noise ratios on many of the instruments.

Fig. 8 shows the spectra of the Po and So coda from the Kuril event at five epicentral distances corresponding to the five OBH's that were used. These spectra have been computed from a series of contiguous 512 point FFT's taken on the time series data which was digitized at 80 samples per sec. The spectral energy has been summed over an apparent velocity window extending from 8.2-7.0 km/sec and from 4.7-4.0 km/sec for Po and So, respectively, and has also been summed over a 1 Hz band at any given frequency. Instrument response has not been removed; however, the response of each individual OBH is presumed to be identical so that intercomparison of the spectra is valid. Only signal levels which were at least 4 dB above background noise levels, at any given frequency, have been plotted. A correction for cylindrical spreading has also been applied to these spectra so that the losses observed in the figure represent attenuation due to anelasticity, scattering, tunneling, and/or any mechanism other than cylindrical spreading. Note that in general the energy level falls off at a constant rate with distance, regardless of frequency. Also note that the attenuation of So with distance is about the same as that for Po.

A general expression for the observed amplitude of seismic signals of this type is:

$$\Delta(f) = A_0(f)r^{-b}e^{-\pi fpq(f)[r-1]}$$

where A(f) is the observed signal amplitude at a given frequency, $A_0(f)$ is the source amplitude (r=1) at a given frequency, f is the given frequency,



ambient noise levels have been plotted. Dashed portions of the removed from the data. Only signal levels at least 4 dB above spectra are for continuity only. and do not represent actual velocity ranges 8.0 - 7.2 km/sec and 4.7 - 4.0 km/sec for Po observed at five epicentral distances corresponding to five The effect of cylindrical spreading. amounting to only 1.3 dB between 24.60 and 33.30. has been Spectral energy has been summed over the apparent Spectra of Po and So from a large event in the Kuril Is. data at those frequencies. and So. respectively. OBH's. Figure 8.

r is the epicentral distance, b is the spreading parameter (0.5 for cylindrical spreading; 1.0 for spherical spreading), p is slowness (i.e., travel-time/epicentral distance), and $q(f) = Q(f)^{-1}$ is the apparent anelastic attenuation coefficient.

Conversion of the observed amplitude into decibels gives:

$$20\log_{10}A(f) = 20\log_{10}A_0(f) - 20\log_{10}(r) - 20\pi fpq(f)[r-1]\log_{10}(e)$$

Rearranging terms and differentiating with respect to r gives:

$$\frac{d(20\log_{10}A(f) + 20\log_{10}(r))}{dr} = -20\pi fpq(f)\log_{10}(e)$$

Therefore:

$$Q(f) = q(f)^{-1} = \frac{-20\pi fp \log_{10}(e)}{\frac{d(20\log_{10}(A(f)r^b))}{dr}}$$

Each curve plotted in Fig. 8 represents $20\log_{10}[A(f)r^{0.5}]$ at a given r for a range of frequencies. At any given frequency, we know $20\log_{10}[A(f)r^{0.5}]$

at several r's and can thus determine $\frac{d\{20\log_{10}[A(f)r^{0.5}]\}}{dr}$ by simple linear regression. All other quantities in the equation for Q(f) are known. Q(f) has thus been found for data sampled in one Hertz bands, and these values are plotted in Fig. 9. Although cylindrical spreading (b = 0.5) has been assumed, the spreading term plays only a minor role in these calculations. A choice of spherical spreading would raise the values of Q(f) by no more than about 10 percent. Only frequencies where data exist at all five distances have been used. The standard deviations shown include contributions from the error in the determination of $d\{20\log_{10}[A(f)r^{0.5}]\}$, the error in slowness which is due to sampling the

data over a range of slownesses, and the error in frequency due to sampling over a 1 Hz band. This method gives values for Q(f) roughly proportional to frequency, and gives higher values of Q(f) for So relative to those for Po at any given frequency. These results are consistent with those found by Sutton et al. (1983) mentioned earlier, although the absolute levels for Q(f) are lower at any given frequency in this study. The relationship between the Q(f) values of Po and So, $Q_{\mathfrak{g}}(f) >= 1.4Q_{\mathfrak{q}}(f)$, is markedly

different from the relationship $t_{\beta}^{\pi} = 4t_{\alpha}^{\pi}$, which implies $Q_{\beta}(f) = 0.44Q_{\alpha}(f)$, found by others (see for example Cormier, 1982).

The nature of the Q(f) values suggests that a simpler mathematical model might be used to describe the attenuation of Po/So. This is because

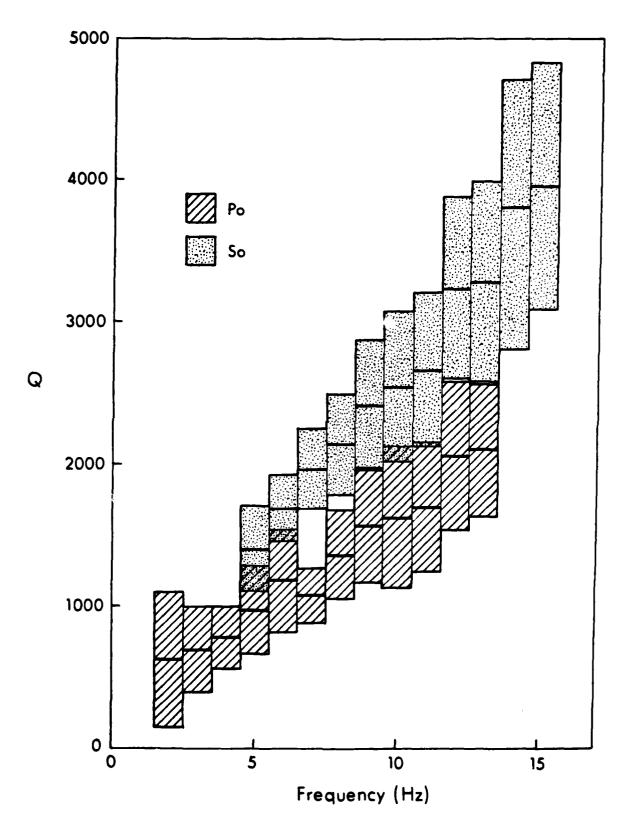


Figure 9. Values for Q(f) plus and minus one standard deviation as explained in the text are plotted versus frequency for both Po and So.

values for fpq(f) are approximately the same for any given frequency and slowness (ie., Po or So). Thus, the form of this model is: $A(f) = A_0(f)r^{-b}e^{-a(r-1)}$, where "a" is the coefficient of attenuation. In this expression, the change in signal amplitude over some propagation distance is not a function of frequency or slowness as in the previous expression containing "q". By a method similar to that described earlier it can be

shown that: $-20a\log_{10}(e) = \frac{d\{20\log_{10}[A(f)r^{0.5}]\}}{dr}$, where $-20a\log_{10}(e)$ equals the decibel change in amplitude. after a cylindrical spreading correction, per unit propagation distance. The right-hand side of the above equation may be evaluated by linear regression. using the set of values for $20\log_{10}[A(f)r^{0.5}]$, which can be computed from the observed values for A at a given r (and for a given frequency band), versus r. Values for $-20a\log_{10}(e)$ have been computed and tabulated in Table 2. The same data set used to determine the Q(f) values discussed previously has been used to compute values in the table. With the exception of two lower values for Po at 2 Hz and 3 Hz and the slightly higher average values for Po relative to So, it could be said that the data generally support the assumption that "a" is nearly a constant value for both Po and So at any of the frequencies examined.

All of the data can be combined by a zero-meaning method, similar to that described for combining the first arrival data. to yield a single value for -20alog₁₀(e). These modified data are shown in Fig. 10. and the regression line fit to them has a slope of -21.5 \$\pm\$ 0.9 dB per 1000 km of travel path. Also shown are regression lines fit to the subsets of Po and So data which have slopes of -19.7 ± 1.4 and -23.6 ± 1.1 dB per 1000 km. respectively. Differences between the Po and So values reflect those differences observed in Table 2. Note that all of the data lie at five distances corresponding to the five OBH's. Also note that the data points at each distance do not cluster about the central regression line. but tend to cluster about some value away from the line. It is possible that these shifts. amounting to a few dB at most. are due to differences between individual OBH responses. These differences may exist because of complexities in processing the slow-speed cassette tapes, and uncertainties in the absolute calibration of the individual hydrophones. Computation of new regression lines, under the assumption that the data of each OBH contain a constant but unknown bias. and under the condition that the sum of the squares of the biases are minimized, results in exactly the same regression lines shown in the figure with somewhat smaller variances. This result is due to the mathematics. and not to any unique property of the data.

Conclusions

Data collected by a 1500 km OBH array deployed for two months in the Northwestern Pacific Basin near Wake Island has provided important new information about the propagation of Po/So phases. Using this array. propagation velocities of first arrivals from shallow-focus (<100 km)

Table 2. Attenuation of Po and So as a Function of Frequency and Propagation Distance

f	Attenuation (dB/1000 km)						
(Hz)	Po	So					
2	-11.5 ± 8.4						
3	-15.4 ± 6.1						
4	-18.5 ± 4.7						
5	-18.4 ± 5.6	-22.4 ± 4.3					
6	-18.3 ± 5.4	-22.3 ± 2.6					
7	-23.4 ± 3.8	-22.3 ± 2.8					
8	-21.1 ± 4.6	-23.4 ± 3.4					
9	-20.6 ± 5.1	-23.4 ± 4.1					
10	-22.1 ± 6.7	-24.7 ± 4.8					
11	-23.4 ± 6.0	-25.9 ± 4.9					
12	-21.0 ± 5.2	-23.3 ± 4.3					
13	-22.2 ± 4.8	-24.8 ± 5.1					
14		-23.1 ± 5.2					
15		-23.8 + 5.0					

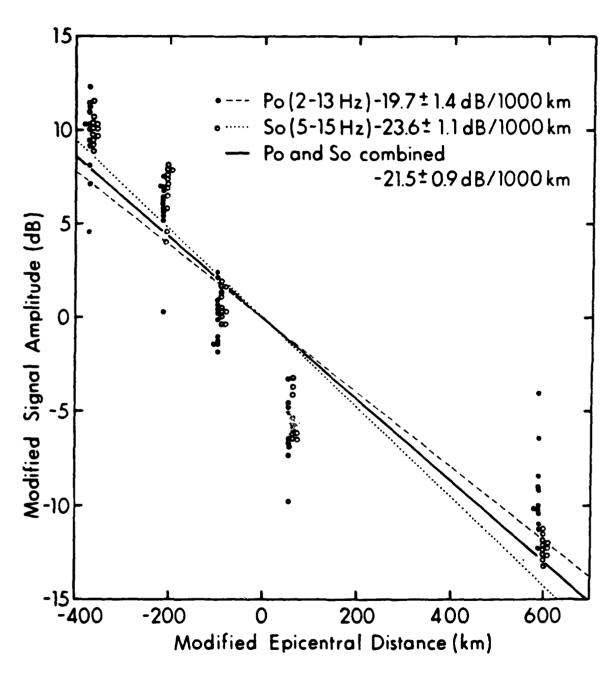


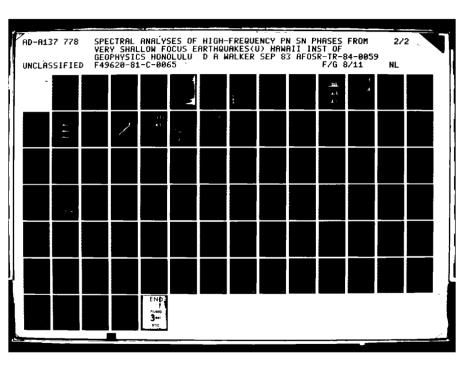
Figure 10. Spectral amplitudes at each frequency studied (Po: 2-13 Hz and So: 5-15 Hz) are plotted versus epicentral distance after modification by zero-meaning as explained in the text. The five data clusters represent the five OBH's used. Note that the data of both Po and So, at all frequencies, follow the same general trend, which is sitenuation as a function of distance. For reference, attenuation as a function of distance at each individual frequency in Po and So is listed in Table 2.

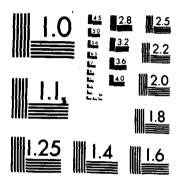
events have been measured by a more accurate method than was previously possible using single station data. The velocities found are 7.96 ± 0.05 km/sec and 4.57 ± 0.04 km/sec for Po and So, respectively.

measured at epicentral distances between 18⁰(2000 km) and 33⁰(3700 km). Associated with these propagation velocities in the linear travel-time equation are intercepts equal to -7.14 ± 2.38 sec and -14.03 ± 5.31 sec for Po and So. respectively. These large negative intercepts imply higher propagation velocities over some part of the travel path nearer to the source than where the data were recorded. Travel times of all Po/So first arrivals collected by HIG since 1963, from shallow-focus events with

Northwestern Pacific Basin travel paths greater than 12° epicentral distance. are successfully modeled by the linear travel-time equation: T=X/V + I, using those values for V (propagation velocity) and I (travel-time intercept) reported here. Data at epicentral distances less

than 12° (collected prior to this experiment) appear to follow a different Po/So travel-time branch than the one described by the OBH data. The exact nature of this branch is not clear due to deficiencies in the quantity and quality of this data. Additional insight into this transition may be found through the examination of Po/So arrivals from numerous deeper-focus events (up to 600 km focal depth) which have been recorded but not yet studied. The OBH data have also been used to directly measure the attenuation of Po and So. An event which occured in the Kuril Islands (mb=6.6, h=45 km) and generated Po and So arrivals which propagated almost exactly down the axis of the array, was used to quantify attenuation in terms of a frequency-dependent Q. Values found for Q(f) range from 625½ 469 at 2 Hz to 2106 ± 473 at 13 Hz for Po. and from 1401 ± 296 at 5 Hz to 3953 ± 863 at 15 Hz for So. This attenuation may be more simply described in terms of the single value -21.5 ± 0.9 decibels per 1000 km of travel path, which in general applies to all the frequencies studied in both Po and So.





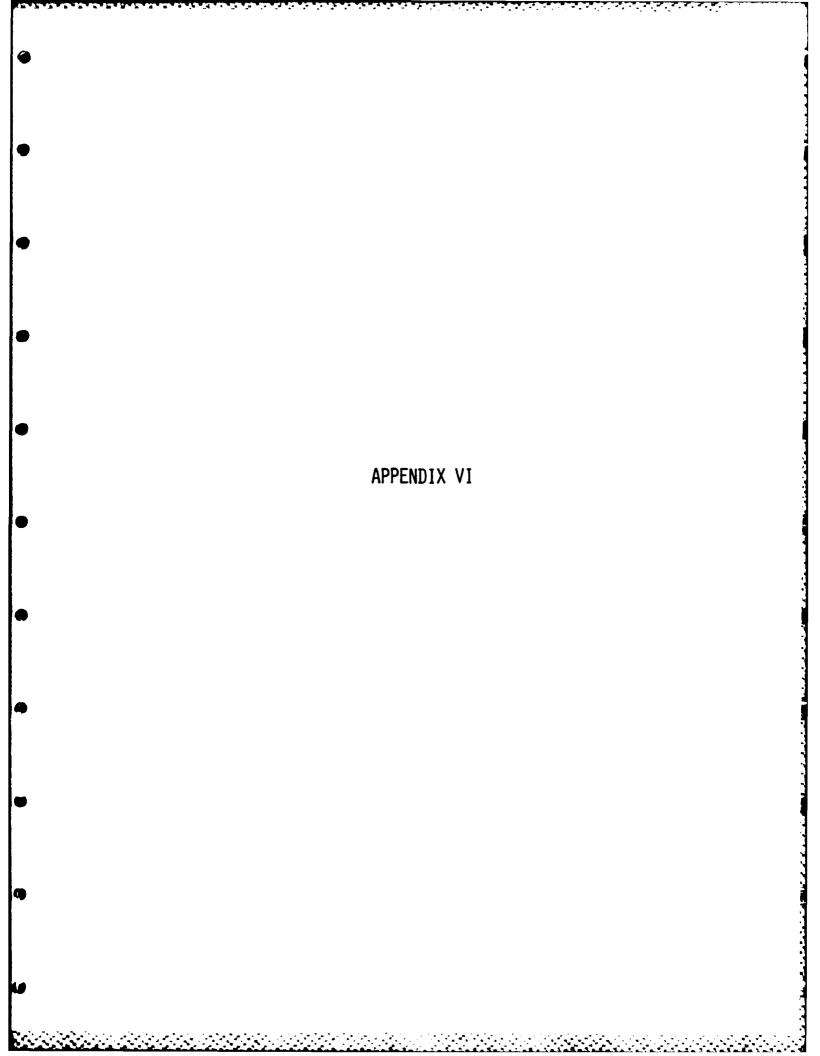
MICROCOPY RESOLUTION TEST CHART
NATIONAL BUREAU OF STANDARDS-1963-A

Acknowledgements. This research was supported by the Office of Naval Research (code 425GG). Funds for support of the hydrophone station at Wake Island were provided primarily by the Air Force Office of Scientific Research under Contract No. F-49620-81-C-0065. with supplementary support from the U.S. Arms Control and Disarmament Agency. The authors express thanks to Dave Byrne. Grant Blackinton. Bob Mitiguy. Dave Barrett. and Fred Duennebier for their help in modifying, testing, launching, and recovering the OBH's. Appreciation is also expressed to Al David and Kentron Corporation for their part in maintaining the recording station at Wake. The authors also thank George Sutton for reviewing this manuscript. and Rita Pujalet for editorial assistance. Hawaii Institute of Geophysics Contribution No. 0000.

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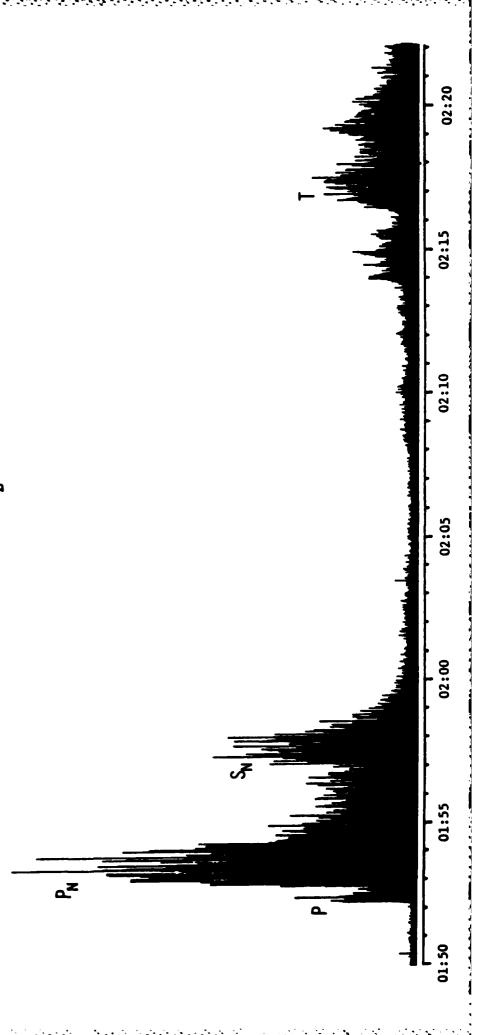
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OPA Newsletter of the Ocean P Alliance Number 1 September 15, 1982

6 September 1982; 01:47:02; 29,3N,140,3E; 6.6 M_B ; 167 KM; South of Honshu; Distance = 25,20



NEWSLETTER OF THE OCEAN P ALLIANCE

7.6 and 4.5 km/sec) comparable to basal crustal rates. At distances high as 30 and 35 Hz, respectively; and at distances of about 30° , as high as 15 and 20 Hz, respectively. The signal-to-noise ratios The purpose of OPA, the newsletter of the "Ocean P Alliance," This phase, Po, and its associated S phase, So, were first observed in the Morth Atlantic and have since been found throughout the travel with fairly constant apparent velocities of about 8.0 and 4.6 km/sec, respectively, while peak arrivals have velocities (about of about 180 (' 2000 km), observed frequencies of Po and So are as for Po/So phases are generally at least ten times greater than the ratios of their respective normal, mantle-refracted P and S phases; and, in many instances no P's and S's can be found in spite of the Morth, Western, and Central Pacific. First arriving Po/So phases is to stimulate interest in a seismic phase known as "Ocean P.

Aside from the SOFAR channel of the world's oceans, the Po/So waveguide (estimates of Q are as high as 20,000). Also it seems probable that the phenomenon is a dominant feature of all of the waveguide appears to be the earth's most efficient acoustical world's oceans and marginal seas.

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In spite of these remarkable and extensive observations, Po/So support that it deserves. Recent observations indicate the phenomenon is not just of interest in terms of basic science, but applied science as well. The applied aspects include the detection seas, and acoustical studies of ocean sediments and crust at high research has not yet received the general recognition, interest, and and discrimination of underground nuclear explosions along the crust and uppermost mantle of the world's oceans and marginal subduction zones and/or continental margins, large-scale mapping of

In view of the foregoing discussions, the intent of this newsletter is to promote interest in Po/So research by providing an additional, and more effective, forum for the exchange of observations and addressed to: Contributions or correspondence should be

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and do not reflect official positions of the Ocean P Alliance unless Views expressed in this publication are those of the authors only expressly stated.

Acknowledgments. The concept of a newsletter to improve communications among those interested in Po/So research grew those who prwided this encouragement. I would also like to express directly out of responses to my 30 March letter. I thank all of

Science Foundation, the Office of Naval Research, the Air Force Office of Scientific Research, and the U.S. Arms Control and my appreciation to a few visionary program managers at the National Disarmanent Agency .

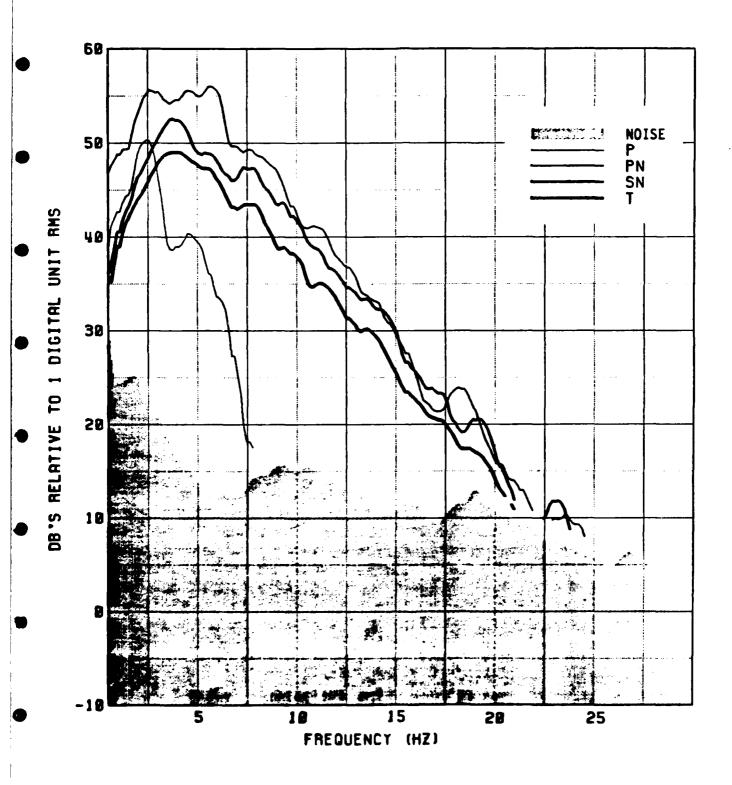
Dear Potential Ocean P Alliance Members:

importance to geophysics. (In the event that your may have misplaced the 30 March letter, some of the more significant observations are summarized in the masthead. I To share this concern and to obtain an objective assessment, I requested that you you concerning seismic phases variously referred to as highmy opinion that the level of interest in, and support for, research on these phases was far less than commensurate with their potential On 30 March 1982, I wrote a rather lengthy letter to some of frequency Pn/Sn, long-range Pn/Sn, or Ph.f./Sh.f. It was, and is, reciprocate and express your views to me.

encouraging -- so much so that you have before you a new means for the As you may have guessed, the response has been very exchange of observations and ideas on the phenomenon of "highfrequency Pn/Sn" propagation. The history of research on these phases since they were first recognized nearly fifty years ago could best be described as sporadic -- with only a small number (often zero) of researchers actively studying the phenomenon at any given time. Such efforts for those early years are understandable. Unfortunately, even with the advent of ocean-bottom hydrophones, digital systems which result much of what we publish is given little attention, or is totally ignored; and this situation is exacerbated by the large greatly facilitate spectral analyses, and rapidly evolving computer number of journals ("haystacks") in which these studies ("needles") techniques for modeling complex seismic phases, the number of researchers actively studying the phenomenon is still small. can be published. Rationale.

constituency and pushing for a solution. Legitimate means are increased communications, as well as continuing quality research. Thus, the timing now seems appropritate for building a larger Without such an effort at this time, I am fearful that research on the phenomenon will still be dangerously close to the "sporadic mode," making hard times and additional periods of dormancy very real possibilities.

I believe the assessment that high-frequency Pn/Sn propagation is "a challenge remaining to the theoretician," (Richards, 1979, Rev. Geophys. Space Phys., 17, 312-328) and is "the challenge to both explosion and earthquake seismology for the coming decade" (Hirn et al., 1973, 2, Geophys,, 39, 363-384). I also believe that major geophysical problems are increasingly rare and that such the opportunities for widespread participation in the solution of opportunities should be seized. Name Change. It may be advantageous at this time to suggest that a new name be given to the high-frequency compressional and shear .808. phases observed at great distances in the world's



difficulty with the nomenclature used to date is that: (1) an, as yet, unsubstantiated relationship to the well known longer-period Pm/5m phases of continents is inferred; and (2) the environmental festure most atrongly linked to the observations is not cited. Thus, a more logical term would be "Ocean P" or "Ocean S" with the abbreviations being "Po/So." With this change, those unfamiliar with the phenomenon would not be as likely to make the false assumption that the phases are similar to continental Pn and Sn. Such assumptions in the past have been a major atumbling block in atimulating interest and support for "Po/So" research.

Our Birthday. You may have noticed something special about the date of this inaugural newsletter--15 September 1982 commemorates the 47th anniversary of the first known published report of Po/Sophases.

"In the bulletin of the Harvard Seismograph Station, under date of <u>September 15, 1935</u>, attention was directed to the unusual character of certain records from the vicinity of 17° N, 62° N. One of the novel features was a short-period phase about 23 minutes after P. It has become known as T, for third, with P and S constituting the first and second groups of short-period waves of similar general appearance. The problem was discussed with Weston, and since that time those two stations have been working on it. Limeham published in 1940 the first description outside station bulletins." [Leet et al., 1951, <u>Bull. Seismol. Soc. Ams.</u>, 41, 123-141].

"At the 1939 meeting of the Eastern Section of the earthquakes occurring about 15° to 30° distance from Weston, Massachusetts. The difficulty of locating these quakes was stressed in the paper, due to the strange characteristics of the recordings. Observers of many stations thought of them as a series of locals. As has been mentioned, the records from this area are quite different from those we have recorded from their area are quite predominant characteristic is the multiplicity of their extremely short period. The so-called P-group may last as long as three or four minutes due to the multiplicity mentioned and frequently runs into the S-group: The second group lasts about the same length of time as the former. About 20 minutes after the P-group is a third unidentified group which we have labelled as I." [Linehan, 1940, EOS IXans. AGU, 21, 229-232].

"Actually, many features of P and S are abnormal on this and later records from certain areas at this distance range, and work on that part of the problem is in progress, but the investigation of T has been undertaken first." [Leet et al., op. cit.]

As I stated in my 30 March letter, although the T-phase was accurately identified in a relatively short time as compressional energy traveling in the sound channel of the world's oceans, one could best describe work on the other "part of the problem" (i.e.,

the "abnormal" P and S phases), 47 years later, as still being "in progress."

It is my hope that on the 50th anniversary of the first reported observations of Po and So, the need for a special newsletter would no longer exist.

Success. The minimum objective of the newsletter is nothing more than the stimulation of deserved interest in Po/So research. Achievements beyond the minimum objective will be strongly dependent on contributions and correspondence from members and their "lobbying efforts" with program managers.

Membership. The "Ocean P Alliance" is not at the moment a formal organization. If no useful purpose would be achieved by "formalization" then this current status should remain unchanged. Anyone wishing to be considered as a member of the Alliance is automatically "in." Anyone wishing to remain skeptical, but nonetheless interested, could consider themselves, if they wish as non-members and still receive the newsletter. I may at some later date directly ask you for your opinions on the newsletter.

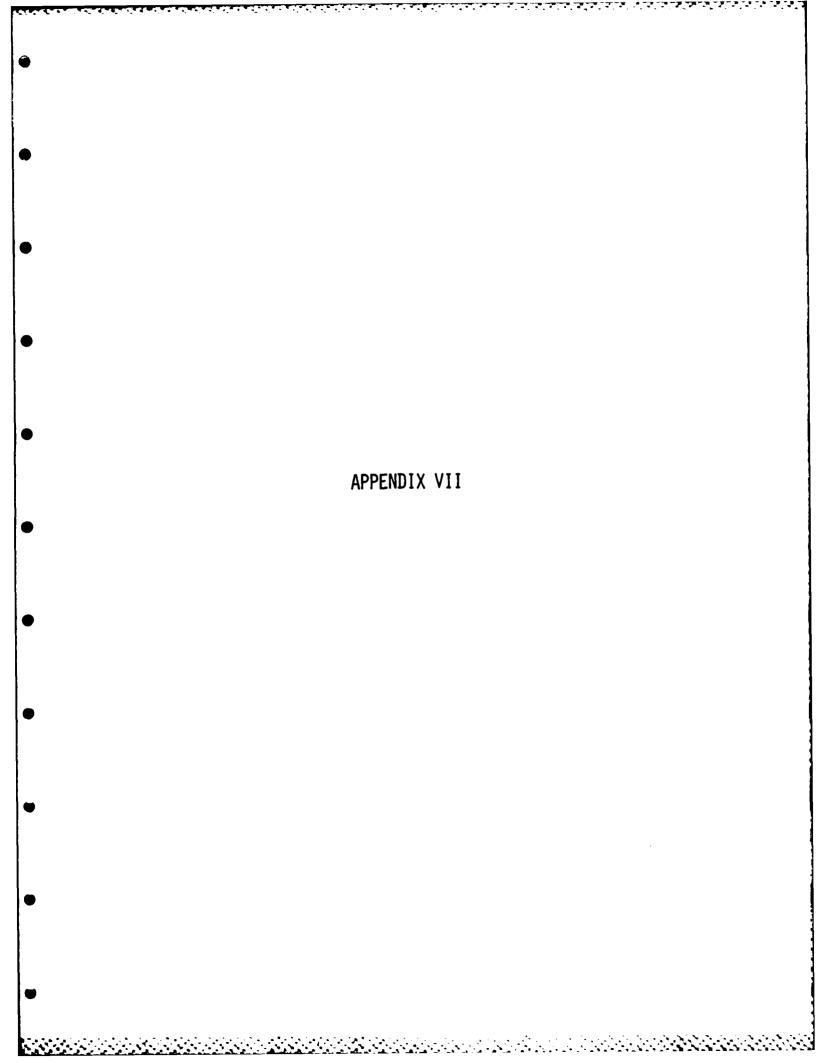
I would hope that those receiving the newsletter would provide:

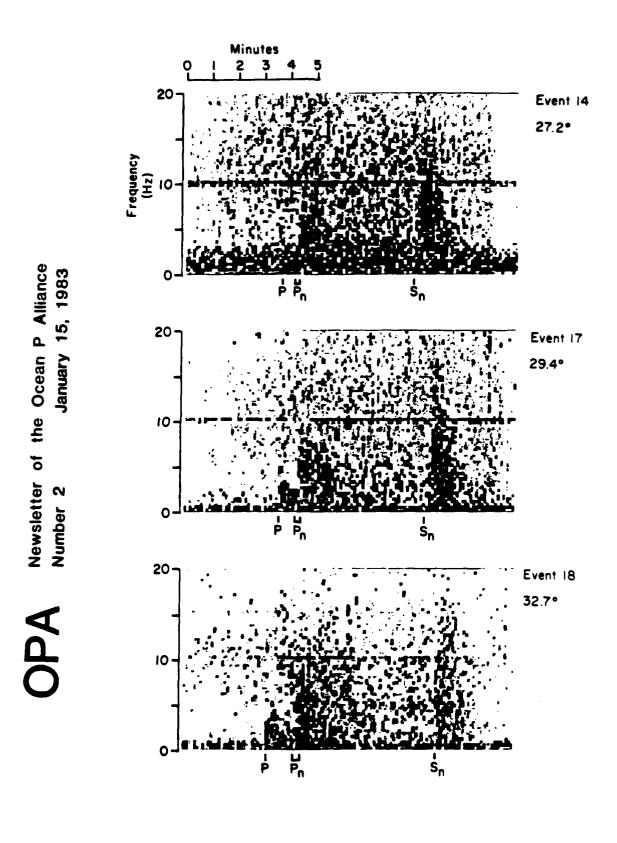
(I) The names and addresses of individuals who should be added to our mailing list, and (2) Po/So publications which should be added to our Po/So bibliography, as well as news of forthcoming publications. If you have some other newsworthy items or comments on topics discussed in OPA, please send them to me. I would also like to know if you have been mistakenly placed on the mailing list. Foxmat. We hope that you will save all issues of OPA for future reference. To that end we have a format conducive to storage in a 3-hole binder.

Future Topics. Items for future issues include the OPA mailing list, correspondence from members, a bibliography of Po/So research, more amazing Po/So recordings, news of upcoming publications, and interesting questions for debate.

Our Covers. To further commemorate the first reported observation of Po, So, and T phases in September of 1935 across the North Atlantic Basin, we have covered our newsletter with Po, So, and T phases which were generated in September of 1982 across the Northwestern Pacific Basin. In addition, the normal, mantle-refracted P is also shown. These phases were recorded on a twelve channel hydrophone array located near Wake Island. Do you see any similarities between these Pacific phases and those from the Atlantic discussed by Leet et al. (op.cit) and Linehan (op. cit.)? What significant observations can be made from these plots? For answers, "tune in" to the next issue of OPA!

Dan Walker





OPA

HENSLETTER OF THE OCEAN P ALLIANCE

The purpose of <u>OPA</u>, the neveletter of the "Ocean P Alliance," is to stimulate interest in a seismic phase known as "Ocean P." This phase, Po, and its associated S phase, So, were first observed in the North Atlantic and have since been found throughout the North, Western, and Central Pacific. First arriving Po/So phases travel with fairly constant apparent velocities of about 8.0 and 4.6 km/sec, respectively, while peak arrivals have velocities (about 7.6 and 4.5 km/sec) comparable to basal crustal rates. At distance of about 18° (* 2000 km), observed frequencies of Po and So are as high as 30 and 35 Mz, respectively; and at distances of about 30°, as high as 15 and 20 Mz, respectively. The signal-to-noise ratios for Po/So phases are generally at least ten times greater than the

ratios of their respective normal, mantle-refracted P and S phases; and, in many instances no P's and S's can be found in spite of the presence of very strong Po's and So's. Aside from the SOFAR channel of the world's oceans, the Po/So waveguide appears to be the earth's most efficient acoustical waveguide (estimates of Q are as high as 20,000). Also it seems probable that the phenomenon is a dominant feature of all of the world's oceans and marginal seas.

In spite of these remarkable and extensive observations, Po/So research has not yet received the general recognition, interest, and support that it deserves. Recent observations indicate the phenomenon is not just of interest in terms of basic science, but applied science as well. The applied aspects include the detection and discrimination of underground nuclear explosions along subduction some and/or continental margins, large-scale mapping of the crust and uppermost mantle of the world's oceans and marginal seas, and acoustical studies of ocean sediments and crust at high frequencies.

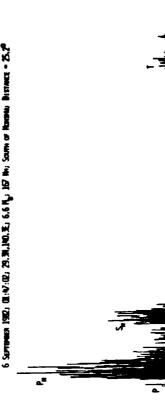
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Dan Walter, OPA Rm. 432, HIG, U. of H. 2525 Correa Rd. Ronolulu, HI 96822 Views expressed in this publication are those of the authors only and do not reflect official positions of the Ocean P Alliance unless expressly stated.

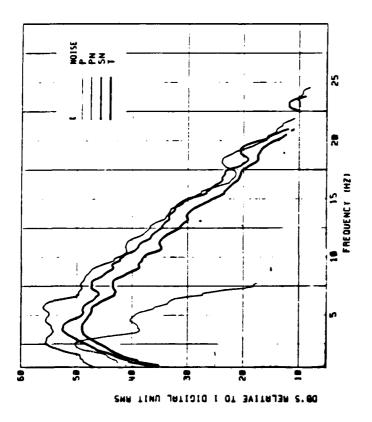
Dear Colleague:

As you may recall, the covers for our inaugural issue consisted of a digitally rectified and compressed plot of P. Po. So. and T phases for an earthquake south of Japas recorded by the Wake hydrophones (front cover) and spectrums for these same phases (back cover). For purposes of the discussion which will follow these figures are repeated here in a severely reduced form. In closing the inaugural issue, I asked whether you could see any similarities between these Pacific phases and those from the Atlantic discussed by Leet, Linehan, and Berger (1951) and Linehan (1940). The unusual characteristics mentioned by these authors are the similar general appearance of their so-called P (primary) and S (secondary) groups to the later arriving I (tertiary) group, the multiplicity of their extremely short period with the P frequently running into the S group, and S about as long as P.

Certainly all of these characteristics have been seen in the Pacific as exemplified in the figures shown here. The P, S, and T groups are or the same general appearance, the P group runs into the S group, in some respects the S group may be about as long as P, and the phases are of extremely short period. The striking and remarkable similarities of these phases are most evident in the spectrums, which also serve to quantify their short period content.



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The T Phase: A New Approach to Po/So Research. Especially intriguing is the detailed correlation of the T spectrum to the So spectrum. This and the general correlation of the T and Po spectrums suggests a possible new approach to Po/So research. In the future it would seem appropriate to consider the possible relationship of Po and So to T. [Other evidence suggesting a possible relationship of Po and So to T is contained in the report "Oceanic Pn/Sn Phases: A Qualitative Explanation and Reinterpretation of the T-Phase" (D. Walker, HIG Report 82-6, 1982). [Topics from this report are discussed in this newletter].

Spectrums for Atlantic Po. So. and I Phases: A Challenge for East and Gulf Coast Seismologists. I believe that Leet, Linehan, and Berger would be aurprised at just how short the periods were for the phases they observed. Unfortunately, the recordings and instrumentation which they had (in all probability 60 mm/minute analog recordings from standard 1 Hz instrumentation) did not permit them to see just bow high the frequencies were. In the deep ocean near Wake, forty years later, we have the advantages of lower noise in the 3 Hz rols Hz range than most continental stations, a better system response to those frequencies than conventional 1 Hz instrumentation, digital recordings, and computers which facilitate

the derivation of spectrums. It is nonetheless surprising that spectrums of Atlantic Po. So, and T phases have not been published. (If I am mistaken, I hope someone will write and provide references.) It would be interesting to compare these spectrums to those from the Pacific. I hope that some east or gulf coast seismologists will take the challenge and be the first to quantify the frequency content of Atlantic Po/So phases.

Continuing Japanese Research on Po/So. I was pleased to meet recently with Drs. Naguso and Kasahara and to review a preprint of their new paper, with Drs. Ouchi and Koresawa, on OSS observations of Po/So phases. The report contains important information on Po/So velocities. I wish them every success in efforts to publish this important paper.

<u>Po/So Bibliography.</u> One of the major objectives of the OPA Newsletter is to acquire a comprehensive and up-to-date bibliography of Po/80 research. A preliminary attempt to achieve this objective follows. I hope that readers aware of omissions will bring them to my attention so that they can be added to the listings.

Our Covers: Now you see "it". Now you don't! The covers are spectrograms from the Wake Island hydrophones. Expected times of arrivals are based on either the Jeffreys-Bullen tables for P or Po/So travel time curves. The contour interval is 8 db. The line at 10 Hz is due to time code cross talk.

What's "it"? The So - very atrong on the front cover, but weak or absent on the back cover. Phases shown on the front cover are typical of events from the Japan, Kuril, Kamchatka portion of the circum-Pacific arc; and, therefore, these phases have travel paths to Wake under the deep Northwestern Pacific Basin. Phases shown on the back cover are typical of events from the New Ireland and Solomon Islands area; and, therefore, these phases have travel paths to Wake under the shallow Ontong-Java Plateau as well as portions or the deep Northwestern Pacific Basin.

An obvious explanation is differences in source characteristics. Unfortunately, this seemingly reasonable explanation cannot be correct, since So phases from the New Ireland and Solomon Islands area are well recorded at Ponape on the northern margin of the Ontong-Java Plateau (roughly mid-way between the source locations and the receivers at Wake). (Of the more than forty events from the New Ireland and Solomon Island area recorded at Ponape during an approximate seventeen month experiment, amplitudes of So phases are at least comparable to, and frequently larger than, those of their respective Po phases.)

What, then, could be the explanation? Possible clues are: (1) the transition allower Ontong-Java Plateau to the deeper Northwei vit. it and (2) the frequent absence or weakness of So, presence or Po, at great distances (often more

than 4000 km) throughout the North Pacific (Walker, 1973 and b) and Central Pacific (Talandier and Bouchon, 1979), in spite of stronger So's than Po's for relatively homogeneous travel paths across the deep Northwestern Pacific Basin.

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A possible explanation will be given in the next issue of OPA. Please send any comments or suggestions on this or other items.

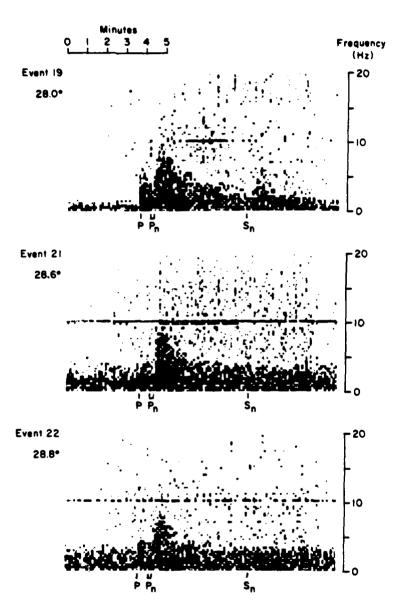
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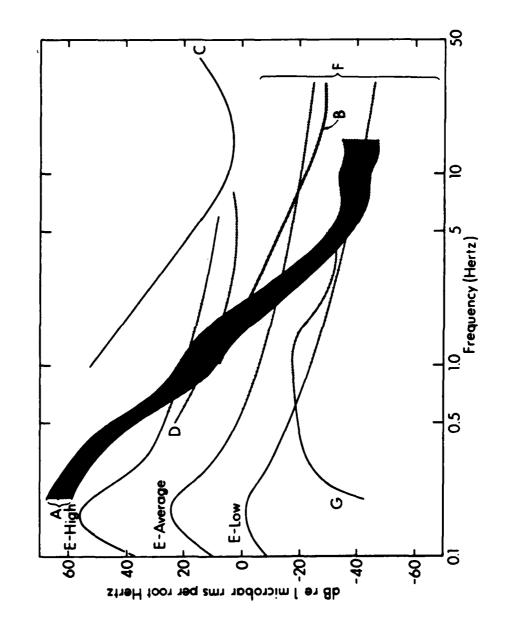
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APPENDIX VIII

OPA Newsletter of the Ocean P Alliance Number 3 May 15, 1983



OPA

REPORTETION OF THE OCEAN P ALLIANCE

The purpose of QLA, the meveletter of the "Ocean P Alliance," This phase, Fo, and its associated S phase known as "Ocean P." This phase, Fo, and its associated S phase, So, were first observed in the Morth, Wortern, and Central Pacific. First arriving Fo/So phases trravel with fairly constant apparent velocities of about S.O and 4.5 km/sec, respectively, while peak arrivals have velocities (about 7.6 and 4.5 km/sec) comparable to basel crustal rates. At distances of about 18° (~ 2000 km), observed frequencies of Po and So are as high as 30 and 35 Ms, respectively. The signal-to-noise ratios for Po/So phases are generally at least ten times greater than the ratios of their respective mortal, mantle-refracted P and S phases; and, is many instances no P's and S's can be found in spite of the presence of very strong Po's and So's.

Aside from the SOTAR channel of the world's oceans, the Po/Soveveguide appears to be the earth's most efficient acoustical waveguide (setimates of Q are as high as 20,000). Also it seems probable that the phenomenon is a dominant feature of all of the world's oceans and marginal seas.

In spite of these remarkable and extensive observations, Po/So research has not yet received the general recognition, interest, and support that it deserves. Recent observations indicate the phenomenon is not just of interest in terms of basic science, but applied science as well. The applied aspects include the detection and discrimination of underground nuclear explosions along subduction nomes and/or continental margins, large-scale mapping of the crust and uppermost mantle of the vorid's oceans and marginal seas, and accustical studies of ocean sediments and crust at high frequencies.

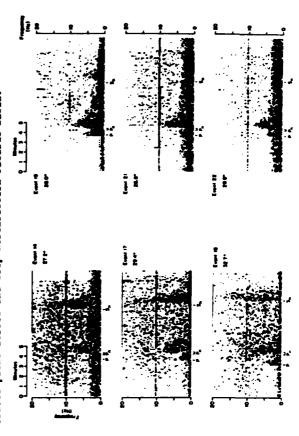
In view of the foregoing discussions, the intent of this newsletter is to promote interest in Po/So research by providing an additional, and more effective, forum for the exchange of observations and ideas. Contributions or correspondence should be addressed to:

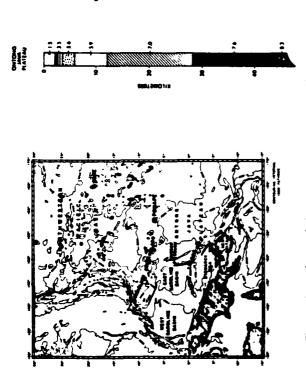
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Views expressed in this publication are those of the authors only and do not reflect official positions of the Ocean P Alliance unless expressly stated.

Dear Colleague:

spectrograms for phases with travel paths to the Wake Island As obvious explanation for the absence, or weakness, of So (labeled Sn) on the earthquakes in the Mew Ireland and Solomon Island arts (see with the amplitudes of 80 at least comparable to, and frequently recordings made during an earlier experiment at Ponspe, roughly midaccompanying map), indicate that this explanation can not be correct. Over an approximate seventeen month recording period, more than forty hydrophones under the deep Morthwestern Pacific Basin (front cover) events from the New Ireland and Solomon Island area were recorded, LOVEVET may between the receivers at Wake and the source locations of The covers of our last issue, shows here is reduced form, back cover is differences in source characteristics. and under the shallow Ontong Java Plateau (back cover). larger than, those of their respective Po phases. An observation which contributes to a possible explanation is the frequent absence or weakness of So, yet presence of Po, at great distances (often more than 4000 km) throughout the Morth Pacific (Walker, 1977s and b) and Central Pacific (Talandier and Bouchom, 1979), in spite of stronger So's than Po's for relatively bomogeneous travel paths across the deep Morthwestern Pacific Basin.





mestle. Observations require that the changes be such that 80 signal strength is reduced without seriously affecting the Po phase. Large arce and trenches, and rafted continental fragments. The transition from the shallow Ontong Java Plateau to the desper Morthwestern lew Ireland and Solomon Island area. [This interpretation has been bethe other than those across relatively bomogeneous ocean basins are likely to escounter large lateral changes in the crust and upper lateral changes could be produced by plateaus, rises, ridge systems, island and seamount chains, fracture somes, transform faults, fossil the Caroline Archipelago through this region, could be features responsible for the weak, or absent, So's at Wake from events in the takes from a recent attempt (Walker, 1982) to provide a qualitative Po/So - hoping that this would eventually lead to comprehensive and detailed quantitative analyses and, ultimately, a generally observations, then, is that facific Basin (see accompanying figures), as well as the extension of framework consistent with all of the diverse observational aspects of acceptable model for the generation and propagation of Po/80.] The explanation arising from these

Hactear Connections. In the masthead of this newsletter, reference is made to Po/So being of interest in terms of the detection and discrimination of underground nuclear explosions. It may be worthwhile at this time to be more explicit as to possible reasons for such interest. A. Obrions - Larre 8/H Batios for Po/80s. An examination of the covers of previous OPA neweletters is all that is needed to suggest that 8/H ratios for Po/80 may be larger than the ratios for manile refracted P (8 is not generally apparent for distances of up to about

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20°). Indeed, numerous observations from the deep ocean, as well as hard rock island sites, in the Western Pacific indicate the 8/H ratios for Po/So at their dominant frequencies (about 4 to 8 Hs) are generally from 5 to 10K greater than the 8/H ratios for P at its dominant frequencies (about 1.5 to 2.5 Hs). This dominance may extend out to distances of 3000 km, with P increasing in strength relative to Po/So beyond this range. Also, at distances less than about 20°, the travel times of Po phases are such that they may arrive ahead of the mantle refracted P, thus masking whatever P energy may be present.

Although some of the most entensive observations of Po/So phases have been in the Western Pacific, these phases have also been observed in other portions of the Pacific and the North Atlantic. The nature and extent of those observations suggest that the phenomenon could be a dominant feature of all the world's oceans and marginal seas.

Thus resional detection and discrimination of madereround amplosions with seismic instrumentation is an ocean environment (islands, the water column, bottom, or sub-bottom) may in large part be based on Po/Bo recordings.

As to wby detection and discrimination in an ocean environment should be important, one has only to look at a globe. As to whether, in fact, detection and discrimination in an ocean environment is important, one has only to consider recent investments in ocean subbottom seismometers. B. Leas Obvious - Le/Pa Comparisons. Just as Po/80 are the most prominent phases recorded at regional distances for oceanic travel paths, Lg and Pg are prominent phases recorded at regional distances for continental travel paths. Certainly Lg/Pg are of great importance in regional detection using continental stations (just as Po/80 abould be of great importance in regional detection using oceanic stations). Surprisingly, the much studied phenomenon of Lg/Pg suffers problems of understanding not unlike those associated with Po/80. In a recent publication (1983) Gupts and Blandford state:

"The mechanism for the generation of transverse component motion from explosions is still not clearly understood in spite of the large number of studies on the subject." [p. 571] and "... no convincing explanation has so far been offered for the generation of shortperiod (about 1 sec or less) transverse motion from explosions." [p. 572.]

In their report they propose a scattering mechanism for the generation of short period transverse motion (including Lg) from explosions.]

Regarding the utility of Lg/Pg in the discrimination of nuclear explosions, several theoretical studies suggest a strong dependence of Lg/Pg signal character on focal depth. Also, Cupta and Blandford note that "the observed differences in the spectra of shear waves

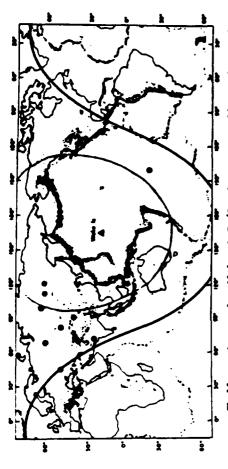
(i.e., ig.) from explosions and earthquakes seem to be large enough to be useful for source discrimination". [p. 589].

Could it be that Lg/Pg investigators might better understand certain aspects of Lg/Pg by turning their attention to Po/So? Or, conversely, might Po/So investigators better understand Po/So by turning their attention to Lg/Pg? Because of the thinness and basins relative to that under large portions of the deep ocean Po/So be viewed as less adulterated (and, therefore, less complex) forms of Lg/Pg? And, finally, could comparative studies of Lg/Pg and Po/So lead to important breakthroughs in detection and discrimination capabilties on continents and in the oceans?

In my opinion all of these questions are very important. If you share this view, I hope that you will bring the phenomenon of Po/So to the attention of your associates who may be working on Lg/Pg. I me happy to report that at least one organization (Rondout Associates Incorporated) has recognized the potential of comparative studies between Po/So and Lg/Pg in the form of a research proposal submitted to APOSR. [Indeed, most of my views on this topic were based on discussions presented in that proposal.]

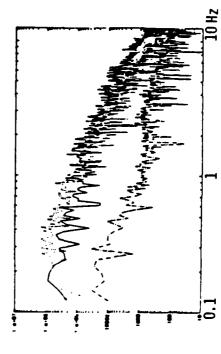
boundary is not firm since high frequency r energy has been observed at shorter distances); and, (d) ocean suface reflections can be used in signal enhancement (see back cover; a thorough description it is apparent from this plot that the noise levels near Wake for frequencies from about 3 to 15 Ms are comparable to, or better than, Novever, by some fortuitous accident of nature (my apologies to those of you who may be mew to the field). Po/So phases (as has already been mentioned) have atrong signals in the 4 to 8 Hz range. Further underground nuclear explosions recorded at great distances (i.e., at distances from about 60° to 90°) have suprisingly large amounts of energy at frequencies above 3 Az, with higher corner frequencies than earthquakes at comparable distances; (b) explosions with smaller yields are believed to be relatively richer at higher frequencies;(c) the Wake hydrophomes are ideally located in the 60° to 90° distance range from most of the known nuclear test sites (see accompanying figure: the shaded area represents the 60° to 90° distance range; solid circles represent underground nuclear test sites; the 90° boundary is firm since no P energy is observed beyond this distance; the 60 boundary is not firm since high frequency P energy has been plot of estimated ocean bottom background noise (shaded region "A") IA detailed little significance if there were no signals at these frequencies. compounding this accident, (a) mantle-refracted P phases from description of this figure is given at the end of the newsletter.] evels for some of the best continental sites. This would be of Our front cover is near Wake Island compared to other measurements. this figure is given at the end of the newsletter). Indirect - Heatle Refracted Explosion Plan

The connection, then, is that sites ideal for the study of Po/So may also be of great value in recording underground explosions at great distances. [Nore detailed discussions of hydrophone recorded explosion P's may be found in a recent publication by McCreery et al. (1983).]

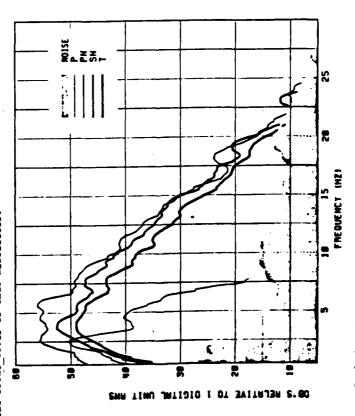


Challance Angustale Although Po/So phases were first observed meanly fifty years ago on the east coast from an earthquake in the West Indies, I noted in the last mewaletter that I had mever seen spectrums for Po/So phases from an Atlantic event. Therefore, in the last newsletter, east and gulf coast seismologists were "challenged" to spectrally analyze Po/So phases for Atlantic trevel paths.

Suprisingly, that challenge has already been met by Rondout Associates Incorporated (RAI), and they have generously permitted the publication of their spectrums in this mewaletter (see accompanying figure). The event occurred near Puerto Rico (h ~ 189km; mb ~ 5.7) and was recorded on RAI's Catakili (New York) Seismic Array (CBA) at a distance of 27.8. Shown are the ground velocity spectra for Po (solid line), So (dotted line), and background noise (dot-dash line).



Comparisons of those spectrums, those on the back cover of OPA No.1 (above here in reduced form), and the front cover of this meruletter are most interesting. For both sets of spectrums, Po's and Bo's are well above background noise out to about 10 Hz. The greater S/H ratios recorded on the ocean bottom (Wake hydrophoue) at frequencies above 10 Hz may be the result of stronger signals, the lower maise levels of the ocean bottom at those frequencies, and/or greater attenuation of the ocean bottom at those frequencies, and/or the Fwerto Rico to GMA path. The greater S/H ratios recorded on the continent (GMA) at frequencies below 1 Hz is almost certainly a result of the lower moise levels of continents at those values. Therefore, the S/H ratios indicated in the spectrums do appear to be consistent with the memiesters for continents and oceans shown on the front cover of this newsletter.



It is hoped that the publication of the Atlantic Po/So spectrums is just the beginning of many comparative studies with Pacific Po/So phases (e.g.; velocities, spectra, Q's, effects of continental paths, structural interpretations, etc.).

En/Ho. Bibliotraphy. Some additions to the bibliography published in the last issue of OPA can already be made. Item #1 is a tentative addition since this modeling effort is applicable only down to periods of 5 seconds and no direct references are made to the "Pu/8n" phases observed at much higher frequencies for oceanic travel paths (i.e. Po/8o). Item #3 was mistakenly omitted from the original list.

1. Nenke, W., and P. Richards, 1963, The horizontal propagation of P waves through scattering media: analog model studies relevant to long-range Pn propagation, Balla, Salangla, Sps. Amer., 73, 125-142.

22.2.2.2.2.2.2

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OUR. GOYELEA. The front cover shows the everage spectrum + 1 standard deviation of 52 semples of background noise over 18 months from the Wake bottom hydrophones (A). Also shown are some published noise deep-sea noise" (Urick, 1975; p. 188). C is a vertical seismometer measurement made in the Marians Basin (Asada and Shimmura, 1976). D is a vertical seismoneter measurement made at 4.6-km depth between for a hydrophone bottomed off Eleuthera Island at 1200-m depth (Michols, 1981). B represents low, average, and high moise levels estimated from curves compiled by Brune and Oliver (1959). F is an back cover shows a sample time series of P. filtered to maximise The upper trace is from a single hydrophone and shows series from each hydrophone was inverted, shifted in time by the water surface reflection time, weighted to maximise the facresse in signal/noise, and added to itself; the two resulting time series were then added with the appropriate propagation delay, and weighted to maximize the increase in signal/moise. Signal/noise was increased by 90% of the theoretical maximum with this method, indicating a high units to the scale shown. B is a hypothetical "sample spectrum of Baveit and California (Bradner and Dodds, 1964). H is a noise curve ares bounded by the limits of noise curves messured on vertical seismometers for 16 locations within the United States and Germany (Frantti et al., 1962). G is the moise curve for the Oyer subarray of the Norwegian seismic array measured during a period "when most of signal/noise, from two nuclear explosions recorded on the Wake bottom trace is a composite of signals from two hydrophones with 40-km separation, obtained as follows: the filtered (1.5-5.0 Ms) time curves for both ocean-bottom (B, C, D, and E) and continental (E, P, and G) environments, which have been converted from an assortment the North Atlantic Ocean was very quite" (Bungum et al., 1971). the direct arrival and its first water surface reflection. level of coherence between the signals added. bydrophones.

Both covers are discussed elsewhere in the newsletter. Interesting question posed by the front cover is this:

"Would instruments with responses comparable to those of the short period components of the World-Wide Seismograph Metwork (WMSR) be appropriate for the studies of regional and telessismic body phases recorded on, or in, the ocean bottom?"

Water Surface Reflection 5 seconds

JUL 79 MB = 5.8EASTERN KAZAKH

McCreery, C., D. Walker, and G. Sutton, 1983, Spectra of nuclear explosions, earthquakes, and noise from Wake Island bottom hydrophomes, Gmomkva, Res. Lett., 10, 59-62.

Hichols, R., 1981, Infrasonic ambient ocean noise measurements:

Eleuthers, Ja Acoustica Soca Ast. 69, 974-981.

Geophys. Res. 34, 5613-5619.

10 DEC 80 WESTERN SIBERIA **△=77°** MB = 4.6

Brune, J., and J. Oliver, 1959, The seismic noise of the earth's surface, Malla Raimpla Rosa Amera, 42, 349-353.

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short-period transverse motion from explosions, Bull, Saismol,

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Aseds, T., and H. Shimmura, 1976, Observations of earthquakes and explosions at the bottom of the western Pacific: Structure of

Also, since it is

on topics presented in the

scophysical journals, advice on any omissions in the Po/80 difficult to be aware of all papers published in the many differing

eveletter and to provide other items of interest.

You are invited to comment

bibliography would be appreciated. Again, special thanks to RAI for

tharing their Po/So spectrums with us.

leferences

Georgical of the Pacific Ocean Rasia and Re Markins, edited by G. R. Setton, M. R. Manghnani, and R. Moberly, Am. Geophys. Union Monograph 19, p. 135-153.

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a qualitative explanation and reinterpretation of the T-phases, Hawaii Inst. of Geophysics Walker, D., 1962, Oceanic Pu/Sn phases: Rept. HIG-82-6, 19 pp. APPENDIX IX

The Continuous Digital Data Collection System for the Wake Island Hydrophones

bу

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at

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ABSTRACT

A continuous digital data collection system was installed on Wake Island in August, 1982, to record seismic signals from an array of nearby hydrophones. Most of the hardware has been purchased "off-the-shelf" from various nationally known vendors. A few components have been built at the Hawaii Institute of Geophysics (HIG). Software for both the data collection and data reduction has been written at HIG. Some key features of the data collection system are: 1) 96 dB dynamic range (i.e., 16 bits), 2) up to 16 data channels, 3) accurate absolute timing (+ 1 msec, generally), 4) accurate interchannel timing, 5) power-failure recoverability, 6) up to 80 samples per second per channel (variable), 7) ease of operation (only four operator commands necessary), 8) operator intervention only once per day (to change up to 4 full reels of 9 track tape), and 9) common tape format (blocked data, 2 bytes per word, 2's complement notation). The preliminary data reduction consists of: 1) stripping off intervals of data for which seismic phases from known events are suspected of being present; and 2) stripping off randomly spaced 3-minute intervals (1 per hour on the average) as ambient noise samples. These tasks are accomplished by a series of programs run at HIG on a Harris H800 computer. This paper is primarily intended for the reader who wants to build a digital data acquisition system, or who wants to use the data collected by this particular system at Wake.

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WAKEUP Tape Format
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Event Data Entry
XSTRIP Parameters
Extracting the Data
Events Download to Strip Tapes
Random Interval Download to Interval Tapes
Checking Procedures
Strip Tape Copies

INTRODUCTION

In June, 1979, HIG began operation of a four-channel (3 data, 1 timecode), slow-speed, cassette recording system for continuous monitoring of hydrophones located near Wake Island. The seismic data collected from this system have provided much needed information about ambient noise in the deep ocean, Po/So propagation, and the spectra at high frequencies (i.e., > 2 Hz) of deep mantle P phases from earthquakes and explosions recorded at great distances. However, the recording system itself had limitations which prevented more quantitative types of data analysis. These limitations included: 1) a dynamic recording range of only about 40 dB - this caused many signals to be clipped; 2) poor timing due to clock drift - absolute timing was accurate to only ±0.3 sec.; 3) poor interchannel timing differences in skew between the recording and playback heads could easily result in a 0.1 sec. timing error between channels; 4) analog format - a sophisticated, time consuming and data degenerating process was necessary to convert the analog signals to digital time series for further processing; and 5) only three of at least eleven working hydrophones could be recorded.

To eliminate these problems, a system was sought with the following general capabilities and constraints: 1) a large dynamic range to record, without distortion, events ranging from at least mb = 4.0 to mb = 8.0 (i.e., at least 80 dB); 2) Absolute timing accurate to less than 0.1 sec (for ease in processing, no time correction should need to be applied to acheive this accuracy); 3) Interchannel timing accurate to within only a small change in phase of the highest frequency of interest (e.g., 0.002 sec for a TI/8 phase change at 30 Hz); 4) a digital recording format such that only a mip: as a smount of processing is necessary to convert the data to a widely

useable format; 5) the capability to record all eleven available hydrophones; 6) the recording of frequencies at least as high as those already observed at Wake in Po and So (i.e., 30 Hz); 7) operation of the system so simple that it can be accomplished by personnel untrained in computer hardware and software; 8) required servicing (i.e., changing tapes) no more than once per day; and 9) the capability of restarting automatically after power failures (which occur frequently at Wake). After exploring numerous options, a design was chosen which met the specifications outlined above, and also appeared to have the best chance for success in terms of the reliability and intercompatability of its components and the feasibility of writing the necessary data acquisition software. This system was purchased. assembled, softwared and tested between March and August, 1982, and was installed at Wake during the last week of August, 1982.

Because of the large volume of raw data which would be collected by this system (i.e., up to 1460 full reel, 9-track, computer tapes per year) a scheme was desired for compressing the data. Data types which were considered the highest priority for saving are: 1) seismic phases applicable to ongoing research topics, 2) ambient noise samples, and 3) seismic phases for future research topics. A system of programs was written to extract and save those intervals in the raw data which contain these priority events. These programs also manage the data for easy accessibility and transmission to other scientists.

RECORDING SYSTEM HARDWARE

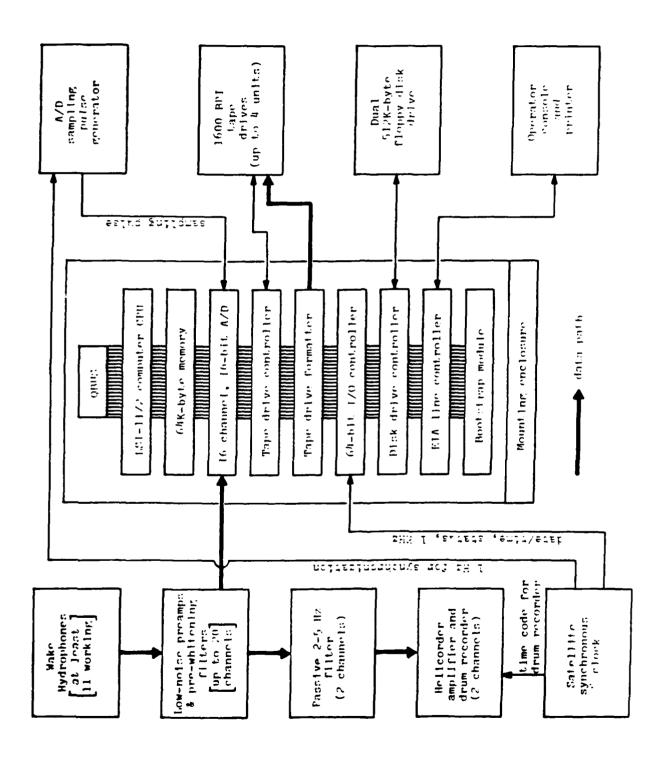
The hardware used for the digital recording system is outlined in the flow chart of Fig. 1. The hydrophone signals are first amplified and shaped by the preamps and pre-whitening filters. Two of these signals are filtered further and used to create a helicorder-style seismogram. All of the signals from the preamps are converted to discrete time series by the A/D (analog to digital) convertor. The digitized data is then assembled into blocks and output to the tape drives by the LSI-11/2 computer. The satellite clock provides a time code for the helicorder, a synchronization signal for the A/D sampling pulse generator, and the date and time in digital format for inclusion with the data blocks written to tape. A more detailed description of the function of each component in Fig. 1 is given below:

Wake Hydrophones

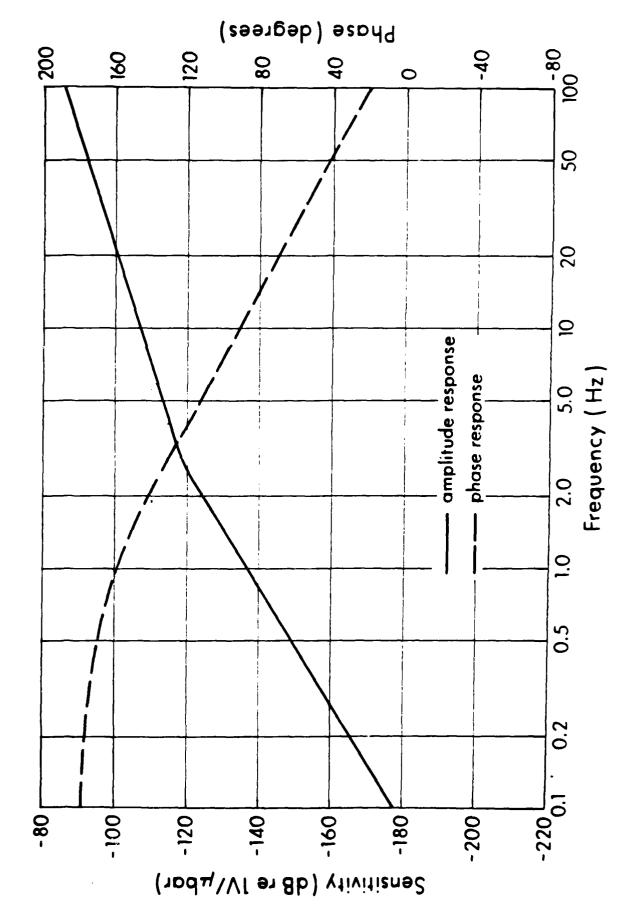
The hydrophones are passive, moving-coil type, which are connected via cable to Wake Island. The estimated response is shown in Fig. 2. The hydrophones are located in two arrays. One array contains six sensors on the ocean floor at 5.5 km depth, and has an aperture of 40 km. The other array consists of five hydrophone pairs at SOFAR depth (1 km) and has an aperture of 300 km. Of the ten hydrophones in the SOFAR array, at least five of them at three sites appear to be working.

Presmp/Pre-Whitening Filters (HIG built)

The preamps are designed to have very low intrinsic noise levels because the hydrophone outputs are very small [as small as 10 nVolts/(Hz)^{1/2} at 10 Hz]. The pre-whitening filters are designed to flatten the ambient



Each device on QBUS represents a single printed circuit board, and these boards are mounted in the A general diagram of the hardware used for the Wake data acquisition system. enclosure from top to bottom as shown. Figure 1.



The estimated response of the Wake hydrophones (taken from the Columbia University's OBS Calibration Manual by S. N. Thanos, 1966). Figure 2.

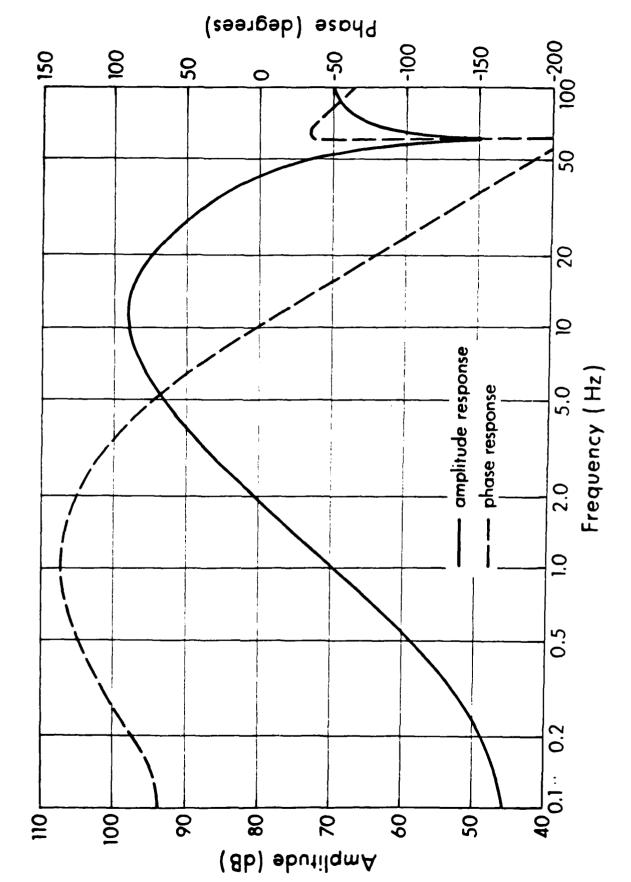
ocean-bottom noise spectrum out of the hydrophone (measured previously using analog data), for the purpose of maximizing the dynamic recording range at all frequencies. Anti-aliasing filters and a 60 Hz notch filter are also included. The output amplifier stage has 10 gain steps (0-9), each 3 dB apart, which can be manually set. The response of the preamp/pre-whitening filters is shown in Fig. 3 (for gain step 9).

A/D Convertor (Data Translation Model DT2784SE/DT5716-B

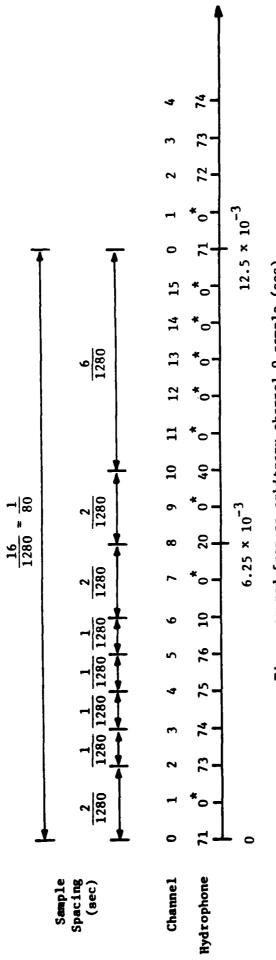
The A/D convertor is designed to fit into the backplane of the LSI-11/2 computer (i.e., into. QBUS), with a cable and connector for receiving the signals to be digitized. The resolution is 16 bits, giving decimal values between -32,768 and +32,767 for inputs between -10.0 and +10.0 Volts, respectively. Up to 16 single ended inputs are acceptable, and the maximum throughput rate is 2.5 kHz (i.e., a minimum of 0.4 msec between samples). Simultaneous sample and hold of all channels is not available on this board, so the 16 channels are sampled sequentially. Furthermore, the design of this device actually requires that all 16 channels be sampled sequentially; the data of unwanted channels being discarded by the software. The trigger which causes a sample to be taken is provided externally by the A/D sampling pulse generator. An example of the temporal spacing between data points which is created by this scheme is shown in Fig. 4. Digital output from the A/D convertor is loaded directly into the LSI-11/2 core memory via direct memory access (DMA).

Satellite Clock (Kinemetrics model 468 DC)

The satellite clock provides inexpensive accurate time at a remote site by monitoring signals from the GOES satellites (East or West). Accuracy is



The response of the preamp/pre-whitening filters at gain step 9 (of 0-9). Subtract 3 dB from the amplitude response for each gain step down from 9. Figure 3.



Time measured from an arbitrary channel 0 sample (sec)

convertor but later discarded by the WAKEUP software. This channel/hydrophone information numbered 0 (also starred, *) are for unwanted data channels which are sampled by the A/D is written into the header part of each data block. The configuration shown is one which An example of the temporal spacing between data channels created by the data acquisition hardware and software for a sampling rate of 80 samples/second/channel. Hydrophones has been in use since the system was installed. Figure 4.

better than ±1 msec when the satellite signal is being received. The time, in Julian days through milliseconds, is output in parallel BCD format through 45 pins of a 50 pin connector. This connector also contains 4 pins for the clock status (i.e., the worst case clock accuracy based on the amount of time since satellite synchronization was lost), and 1 pin with a 1 kHz square—wave used to determine when the other pins can be read (i.e., are not changing states). Also output (on BNC connectors) are a 1 Hz square—wave used to synchronize A/D sampling with the clock, and a slow code used for marking time on the helicorder record.

A/D Sampling Pulse Generator (HIG built)

This device is used to create precisely timed pulses (downward zero-crossings) which are synchronized with the satellite clock and trigger A/D conversions. The front panel has thumbwheel selectors for setting the number of channels of data to be sampled (always 16 in this system due to the A/D convertor; software discards unwanted channels), and the sampling rate (in samples/sec/channel). For 16 channels of data, sampled at 80 samples/sec/channel, the pulse generator produces a 1280 Hz square-wave. A/D synchronization with the clock is necessary so that the starting times of output data blocks, which are assembled for a constant number of A/D conversions, will also remain synchronized over long periods of time.

Mounting Enclosure (DEC model BAll-NE)

The mounting enclosure contains the power supply, backplane (QBUS), and cooling fan for the LSI-11/2 computer. The backplane has quad connectors for nine boards, all of which are occupied in this system configuration.

These nine boards, and the order in which they are placed from top to bottom in the enclosure, are noted in Fig. 1.

LSI-11/2 Computer (DEC model KD11-HA)

This is the central processing unit of the computer.

64 K-byte memory (DEC model MSV11-DD)

This is the core memory for the computer, consisting of 32K 2-byte words.

Bootstrap module (DEC model BDV11-AA)

This module brings up the computer from a down or power fail status, and causes the RT-11 operating system to be loaded from the disk.

64 Bit I/O Controller (DEC model DRV11-J)

This board is necessary to interface the LSI-11/2 with the 50 pin output of the satellite clock. The 64 bits of I/O correspond to four two-byte words which may be accessed by the software to read the clock.

Floppy Disk Drive Controller (DEC model RXV21-BA)

This board interfaces the disk drive to the computer.

Dual 512K Byte Floppy Disk Drive (DEC model RXV21-BA)

The floppy disks contain all of the software necessary for performing data acquisition with the hardware. Under the present configuration, one disk contains the DEC RT-11 operating system software, while the other disk contains the programs written at HIG for collecting the data and a

parameters file which describes the exact configuration to be run (i.e., number of channels of data, blocksize, etc.). It is possible, should one disk drive become faulty, to combine all of the programs and data on one disk, and run from a single drive. A special function of the disk drive is to record the number of the current tape drive to which the data are being written. In the case of a power failure, which causes the core memory of the LSI-11/2 to be lost, the current tape drive number is necessary to safely and efficiently resume recording.

EIA Line Controller (DEC model DLV11-J)

This device provides four ports (0 to 3) for serial I/O (RS-232C, RS-422, or RS-423) to the LSI-11/2. Port 3 is the only one in use under the current operating configuration and interfaces to the DEC LA38-GA which is serving as both a line printer and operator console at 300 baud. Separate ports for the line printer and a CRT operator console were used during program development to facilitate higher baud rates.

Console and Printer (DEC model LA38-GA)

This device is used for all operator input and output (I/O) with the system, and for all program messages. Since there is actually very little I/O to this device, its slow speed is not a problem. Also, there is a distinct advantage in having a hardcopy record of certain essential I/O, especially since the system is at a remote site. Sample I/O is shown in Fig. 5.

Tape Controller and Formatter (DATUM model 15221)

```
WAKE HYDROPHONE ARRAY CONTINUOUS DIGITAL SEISMIC RECORDING SYSTEM
USING DRIVE 2
                       DATE 1983:153 TIME/08:06:00:000 CLOCK/00
              10781
USING TAPE
DRIVE 2 END OF TAPE REACHED
        05820 BLOCKS WRITTEN
USING DRIVE 3 10782
                      DATE/1983:153 TIME/18:10:59:999 CLOCK/00
DRIVE 3 END OF TAPE REACHED
        05824 BLOCKS WRITTEN
DRIVE O OFFLINE
USING DRIVE 1 10783
                       DATE/1983:154 TIME/00:16:19:999 CLOCK/00
SWITCH DRIVES
DRIVE 1 05217 BLOCKS WRITTEN
WAKE HYDROPHONE ARRAY CONTINUOUS DIGITAL SEISMIC RECORDING SYSTEM
USING DRIVE 2 10784
                       DATE 1983:154 TIME/07:31:04:999 CLOCK/00
DRIVE 2 END OF TAPE REACHED
       05824 BLOCKS WRITTEN
USING DRIVE 3
                     DATE/1963:154 TIME/15:36:24:999 CLOCK/00
DRIVE 3 END OF TAPE REACHED
        05804 BLOCKS WRITTEN
DRIVE O OFFLINE
USING DRIVE 1
                       DATE/1983:154 TIME/23:40:04:999 CLOCK/00
RESET-WAIT FOR MINUTE
DRIVE 1 06065 BLOCKS WRITTEN
WARE HYDROPHONE ARRAY CONTINUOUS DIGITAL SEISMIC RECORDING SYSTEM
USING DRIVE 2
                       DATE 1983:155 TIME/08:11:00:000 CLOCK/00
DRIVE 2 END OF TAPE REACHED
       05822 BLOCKS WRITTEN
USING DRIVE 3
                      DATE/1983:155 TIME/16:16:10:000 CLOCK/00
             10788
USING TAPE
DRIVE 3 END OF TAPE REACHED
       05807 BLOCKS WRITTEN
DRIVE O OFFLINE
                       DATE/1983:156 TIME/00:20:05:000 CLOCK/00
USING DRIVE 1
             10789
USING TAPE
WITCH DRIVES
DRIVE 1 05063 BLOCKS WRITTEN
```

Figure 5. A sample of 3 consecutive days of I/O through the console at Wake. Operator input has been underlined (S, R, A). The operator writes-in the WAKEUP tape identification numbers.

The tape controller interfaces the tape drives with the LSI-11/2. Only one controller is necessary for the four tape drives used. The formatter encodes the blocks of data written to tape, and decodes the blocks read from the tape. Only one formatter is necessary for the four tape drives.

Tape Drives (DATUM model D451)

Four 1600 bpi tape drives (0-3) are svailable for writing out the digitized data. Data is written to one drive until its tape is full (or until a specified number of blocks have been written), then data is written to the next drive in the sequence until its tape is full, and so on. Four tape drives are required if tapes are changed only once per day, since one day's data (11 channels at 80 samples/sec/channel) fills four tapes. An important feature which these tape drives have, is that they re-tension the tape, advance the tape several inches for safety, and put themselves "on line" after a power failure. This is an important factor which allows the system to start collecting data again without operator intervention or the unnecessary loss of data.

2-5 Hz Filter (HIG built)

The two-channel 2-5 Hz bandpass filter is for filtering the two signals going to the drum recorder in a way which enhances signal to noise ratios of teleseismic P from underground nuclear tests. The filters are a two-pole passive RC design with the response shown in Fig. 6

Helicorder Amplifier and Drum Recorder (Teledyne Geotech models AR-311 and RV-301)

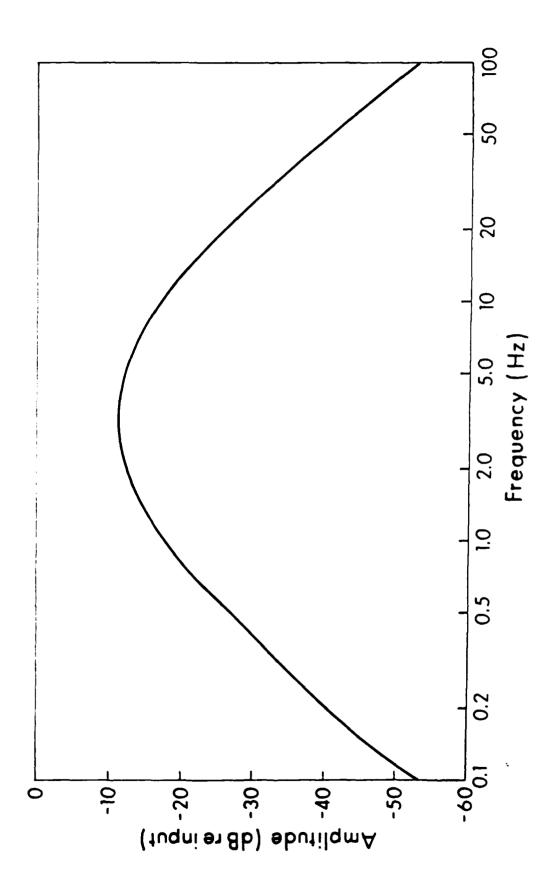


Figure 6. Response curve for the 2-5 Hz filter which precede, the helicorder.

This equipment is used to provide easily scanable visible records of the continuous seismic data. The recorder has two pens for the two input signals, and is geared so that one recording sheet lasts 24 hours.

Additionally, the amplifiers and drum recorder are powered by a trickle charged, automobile battery backup system to provide continuous operation through power failures. Time code is input from the satellite clock. A sample helicorder record is shown in Fig. 7.

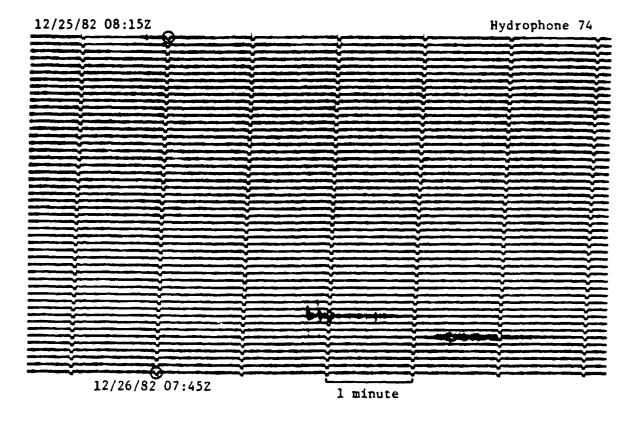
CONTINUOUS DATA ACQUISITION SOFTWARE

The data acquisition program, titled WAKEUP, is written in assembler code and executed under the RT-11 operating system. A simplified flow chart of this software is shown in Fig. 8. Some key features of program WAKEUP are: 1) it is controlled by an input parameters file which can be changed to satisfy differing data collection requirements; 2) it is extremely simple to run - only four, one-letter, operator commands (to be discussed later) are recognized; 3) it can restart itself after a power failure; and 4) it makes a hardcopy log of important program events.

WAKEUP Parameters

The WAKEUP running parameters can be displayed and changed by an interactive fortran program called PARAMS. A sample run of PARAMS is shown in Fig. 9. The parameters and their meanings are described below:

Digitization Rate. The value set here should be the same as the value set on the A/D sampling pulse generator for samples/sec/channel.



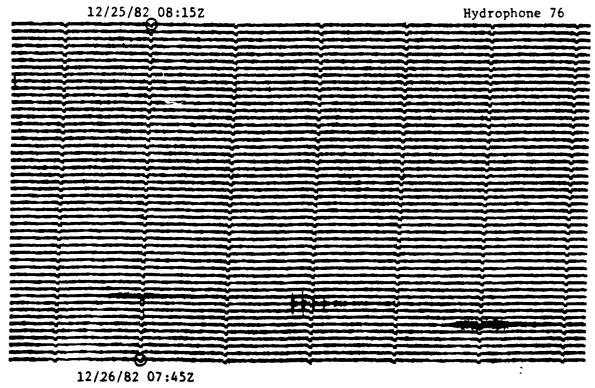
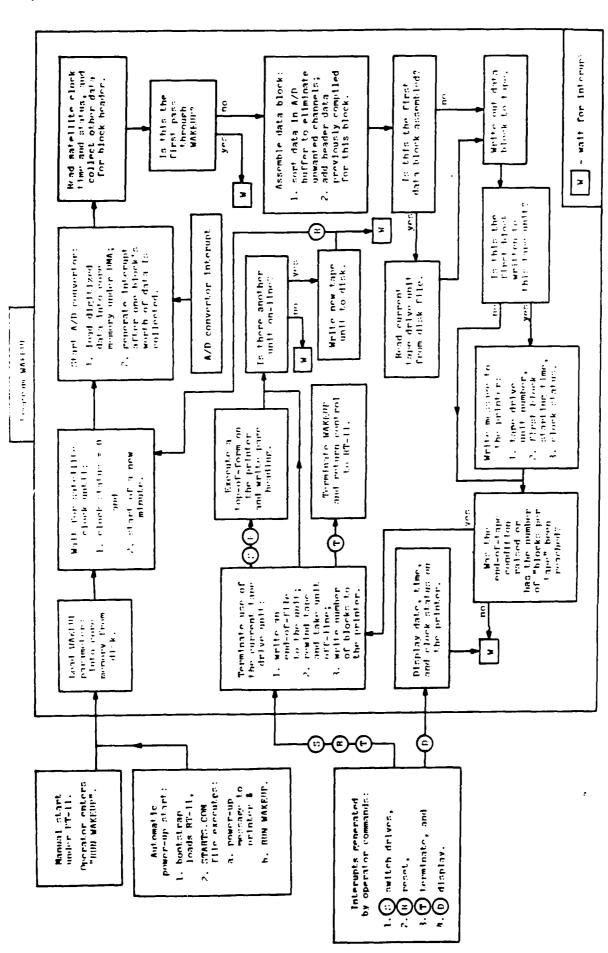


Figure 7. A section of helicorder record from Wake showing the P arrival from an Eastern Kazakh explosion (the impulsive arrival) and some other unidentified arrivals which are probably T-phases. There are 30 minutes between each trace.



0

are used; two of them provide a double buffer for DMA output from the A/D convertor, and the third has a high priority and is serviced immediately. The operator command interupts are low-priority The program Three data buffers The A/D convertor interupt Paths marked S, R, T, or D are Generalized logic flow-chart of program WAKEUP which controls the data acquisition. followed (only once) after the corresponding operator interupt is serviced. is started either manually, or automatically after a power failure. and are serviced when the CPU is in the wait (W) condition. is used for assembling the data block to be output to tape. Figure 8.

```
. BUN PARAMS
PARAMS - SETTING WAREUP PARAMETERS
 CURRENT PARAMETER VALUES
DIGITIZATION RATE: 80 SAMPLES/CHANNEL/SEC
SECONDS PER TAPE BLACK: 5
    EYD
         AMP
               GN
                   CH 
                      HYD AMP
                                           HYD
                                                              HYD
                                                                   AMP
                                                          CR
                                                                        GN
     71
                        72
                              8
                                       2
                                           73
                                                           3
                                                               74
                                                                     0
     75
           19
                                           10
                                                               11
                                                                          9
                                                                    16
     20
                9
                    9
                        21
                              3
                                  9
                                      10
                                                18
                                                                     0
           11
                                            40
                                                          11
                                                                0
                                                                         0
12
      0
           0
                0
                   13
                         0
                              0
                                  0
                                      14
                                             0
                                                  0
                                                      0
                                                                     0
BLOCKS PER TAPE: 32000
UNIT - DRIVE SERIAL NUMBER: 0-318 1-319 2-322 3-326
STARTING TAPE UNIT NUMBER: 0
TEAR: 1982 JULIAN DAY: 250
ENTER COMMAND (8 TO DISPLAY COMMANDS). 8
COMMAND - FUNCTION
 0 - CHANGE DIGITIZATION RATE.
 1 - CHANGE SECONDS PER BLOCK.
 2 - CHANGE A/D CHANNEL ASSIGNMENTS.
 3 - CHANGE BLOCKS PER TAPE.
 4 - CHANGE TAPE UNIT SERIAL NUMBER.
 5 - CHANGE STARTING UNIT NUMBER.
 6 - CHANGE YEAR/JULIAN DAY.
 7 - DISPLAY CURRENT PARAMETER VALUES.
 8 - DISPLAY COMMAND FUNCTIONS.
 9 - UPDATE AND DISPLAY WEPARM AND RCOVRY.
10 - CHANGES COMPLETE. UPDATE AND EXIT.
ENTER COMMAND (8 TO DISPLAY COMMANDS). 10
DK:WKPARM.BIN
         400
                   4400
                               71
                                          69
                                                                    73
 99
          74
                     9
                                         199
                                                                    10
                               75
                                                   76
                                                            19
139
          11
                   169
                               20
                                         119
                                                   21
                                                            39
                                                                     40
189
           0
                                0
                                                                     0
                                           0
                                                    0
                                                             0
                            6400
  0
                                       32000
           Ð
                     n
                                                  318
                                                           319
                                                                   322
326
                     0
                                                    0
                                                             0
                                                                     0
DK: RCOVRY. BIN
                   250
  0
        1982
                                                    0
           0
                     0
                               0
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  0
           0
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           0
                     0
                                                    0
DE:WEPARM.BIN AND DE:RCOVEY.BIN UPDATED.
STOP --
```

Figure 9. A sample run of program PARAMS which is used to change (or view) the running parameters of program WAKEUP. These parameters are stored on disk files DK:WKPARM.BIN and DK:RCOVRY.BIN. Operator commands have been underlined.

Seconds Per Tape Block. This value, multiplied by the sampling rate and by the number of active data channels, will determine the total number of samples per block. The number of samples per block should not exceed 5000 due to size limitations in the core memory containing the buffers. Very small blocksizes should also be svoided, since the time to write a very small block to tape may exceed the time necessary to collect the data block.

Channel Number (CH). This is not a changeable parameter, but merely indicates which channel (0-15), of the A/D convertor, is being referred to.

Hydrophone (HYD). This parameter indicates which hydrophone (71-76 for the bottom array, and 10, 11, 20, 21, 40, and 41 for the SOFAR array) is connected to the A/D channel. If 0 is specified, then the channel is considered inactive. To compute the number of active channels, simply count the number of channels with non-zero values for this parameter (11 for the example in the figure). It is not necessary that the active channels be contiguous as shown in the figure.

Amplifier (AMP). This parameter indicates which of the preamp/pre-whitening filters (1-20) the hydrophone signal is connected to.

Gain (GN). This parameter indicates the gain step (0-9) manually set on the final stage of the preamp/pre-whitening filters.

Blocks Per Tape. This parameter will set the number of blocks written to any tape drive before the next tape drive is used. When the value is set to a large number, as shown in the example, the program will write to a drive until the end-of-tape mark is sensed before switching drives.

Estimates of how many blocks may fit on a tape can be computed by dividing the block length in bytes by 1600 bytes per inch, adding the length of the preamble, postamble and inter-record gap (~0.6 inches), and dividing into the tape length (28800 inches for a full-reel).

Unit-Drive Serial Number. This information is used to indicate which physical tape drive is connected to the tape drive unit (0-3) recognized by the computer.

Starting Tape Unit Number. This parameter is used to determine which tape drive to write to at the start of program WAKEUP. It is continually updated by WAKEUP to facilitate a smooth recovery after a power failure.

Year and Julian Day - These parameters are necessary because the satellite clock does not output the year. WAKEUP continually reads the Julian day from the parameters file and then updates it with the value ouput by the satellite clock. When the satellite clock Julian day is less than the Julian day in the parameters file (i.e., at the beginning of a new year), then the year is incremented by one.

Next in the example shown in Fig. 9, the user has typed 8 to display all the commands. No changes were necessary so a 10 was then typed to exit PARAMS.

The values in the parameters file (actually two files, DK:WKPARM.BIN and DK:RCOVRY.BIN) are displayed and control returned to the system.

Current WAKEUP Configuration

The WAKEUP configuration, set up by PARAMS, which has been in operation from 8 September 82 to the present (16 June 83) has the following key features: 1) data is digitized at a rate of 80 samples/sec/channel; 2) there are 8 active channels which are (given in the order in which they are multiplexed) 71, 73, 74, 75, 76, 10, 20, and 40 as previously described in Fig. 4; 3) each block contains 5 seconds of data which produces a block size of 3300 2-byte words (i.e., 3200 words of data plus a 100 word block header) and 4) the number of blocks per tape is large so that tapes are always written to their end-of-tape marks. Although it had been intended to digitize and record 11 hydrophones, which would produce 4 tapes per day, assoyted problems with the tape drives have made it necessary to only record 8 hydrophones, which produces 3 tapes per day.

WAKEUP Operator Commands

Because Wake Island is such a remote place, and because those who must operate the system are untrained in computer hardware or software, program WAKEUP has been designed to handle many unusual situations without operator intervention. Only four operator commands are necessary during continuous data collection under WAKEUP, and these are described below:

Switch Drives(S). This command is used to force WAKEUP to write to the next available tape drive. This command is normally executed once per day when the operator comes in to change tapes. First he changes tapes on

those drives which have tapes full of data, and allows WAKEUP to continue writing onto the current drive. When fresh tapes are mounted, he executes the S, which causes WAKEUP to start writing to a fresh tape, and then he changes tapes on that last drive. A sample of S is shown in Fig. 5.

Reset(R). The reset command is nearly identical to the switch drives command, except that WAKEUP will wait for the satellite clock to indicate a change in the minute before it starts the A/D convertor collecting data for the next block. It is desirable to have data blocks which start on an integral second, (or integral minute) because it makes the measurement of time on plots, made later on during data reduction, somewhat simpler. Under ideal conditions, once the block beginnings are synchronized with the minute (i.e., if the number of seconds per block is 5, then all blocks should begin at 0, 5, 10, 15, 20, 25, 30, 35, 40, 45, 50, or 55 seconds after the minute) they should stay that way forever. Unfortunately, the satellite clock sometimes loses synchronization with the satellite for a while and begins to drift. When resynchronization occurs, the accumulated error gets propagated into the starting times of the blocks. When the operator notices this condition, he can correct it by using reset instead of switch drives. A sample of R is shown in Fig. 5.

Terminate (T). This command is used to terminate WAKEUP and return control to the RT-11 operating system.

Display (D). This command is used to display the date, time, and clock status at the beginning of the last data block collected.

Power Fail Recovery

Probably the most harmful abnormal condition, in terms of potential loss of data, from which WAKEUP can recover without operator intervention is a power failure. Temporary periods without power are fairly common on Wake and can be due to storms, construction, and generator problems. Since an operator is only present once per day, an irrecoverable system crash due to a short power failure could cause the loss of the entire day's data.

Fortunately, with the help of some unique features in the Datum tape drives (described previously), the DEC LSI-11/2 hardware, and the RT-11 operating system, it was possible to write the WAKEUP program with the capability to resume collecting data after a power failure, without wasting tape (which could cause the system to run out of tape before the operator arrived) and without writing over old data. An essential element of this recovery capability is that the particular tape drive being written to is stored on a disk file. When power comes back after a failure, WAKEUP reads this file and then resumes writing to the proper drive.

WAKEUP LOR

All operator input as well as RT-11 and WAKEUP messages and operator prompts go through the printer/console to create a log of events. A sample of this log is shown in Fig. 5, and has already been referred to several times. This log is essential for keeping track of normal system operations as well as problems which occur from time to time. The operator is

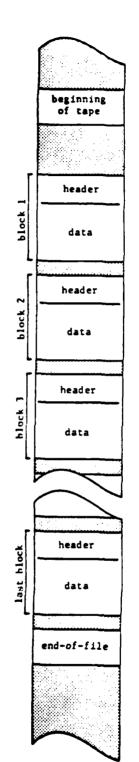
responsible for filling in the identification numbers of tapes mounted on each drive.

WAKEUP Tape Format

The WAKEUP tapes are numbered sequentially beginning with 10001. A detailed description of the format of the tapes and the blocks contained within the tapes is given in Fig. 10. The date and time given in the header are for the first data point in the block.

PRELIMINARY DATA REDUCTION SOFTWARE - WAKE HYDROPHONE INFORMATION PROCESSING SYSTEM (WHIPS)

Because of the large amount of data which is collected by this system (i.e., 50 billion bytes per year or 1500 tapes per year for 11 channels at 80 samples/sec/channel) it was realized that some sort of data compression scheme would have to be implemented in order to efficiently store, manage, retrieve, and distribute those data of current and future interest. These data have been defined in two general categories: 1) seismic phases such as P. S, Po, So, and T, and 2) ambient background noise. Software to sample these categories has been written and its structure is diagramed in the flow chart of Fig. 11. Intervals of data are saved when seismic phases from particular events are suspected of being present; and, in addition, 3-minute intervals of data selected at random times, with an average of one interval per hour, are also saved in order to sample the ambient noise. Three minutes of data is long enough to view several cycles at the longest periods thus far observed on the hydrophones - 20 seconds, and also long enough to provide reasonable statistics on the higher frequencies (1-20 Hz) where most of the oceanic, short-period seismic data are present. By sampling at



WAKEUP Tape Description

General Format: 1600 BPI

2 bytes per word
IBM format
2's complement notation

	All Blocks				
header: wor	ds 1-100 ds 101-end				
Word(s)	Description				
1	Year (e.g., 1983)				
2	Julian day (1-366)				
	Hour				
4]	Hinute				
5	Second				
6	Millisecond				
7 Digitization rate (samples/sec/channel)					
8 Number of samples/channel/block					
9 Total number of samples/block					
10-41 A/D channel information - 16 channels, 2 words per channel 1st word - hydrophone I.D. number (0 if channel not used)					
	2nd word - hundreds and tens digit is the amplifier I.D. number, units digit is the amplifier gain ste				
42	Tape unit (0-3)				
43	Tape drive serial number				
44 Clock status					
45 A/D buffer size					
46-100	Not used				
101-end	Digitized hydrophone data (multiplexed-order of hydrophones is the same as the order of the active channels described in words 10-41)				

Figure 10. The WAKEUP tape format.

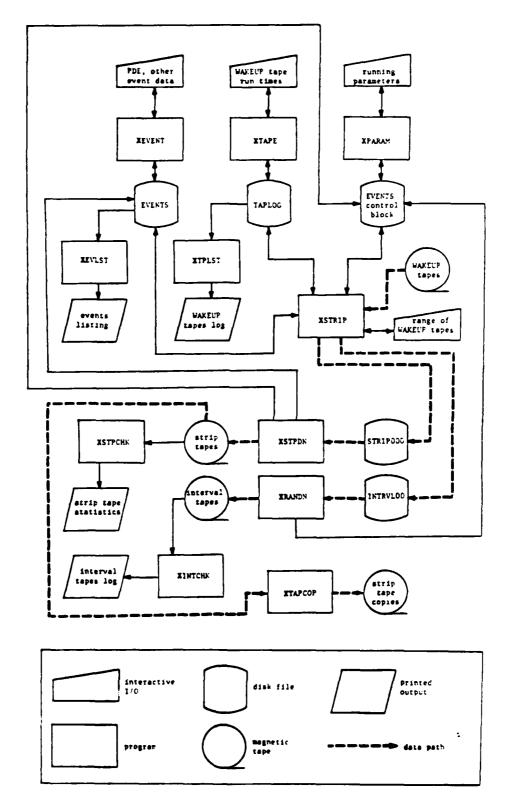


Figure 11. Generalized flow chart of the programs and files used in the preliminary reduction of the Wake digital data.

random time intervals, periodic noise of artificial nature is hopefully avoided. This system of programs and files which performs processing of the WAKEUP data is called WHIPS for Wake Hydrophone Information Processing System, and is run on a Harris model H800 computer.

WAKEUP Log Entry (XTAPE)

An initial step in the WHIPS procedure is to create or update file TAPLOG, which contains an identification number, starting time, block count, and status of each tape created by WAKEUP. This information is entered from the WAKEUP log (Fig. 5), using interactive program XTAPE. The data is necessary, as input to program STRIP, for calculating the time window to be processed for the compression of a given range of WAKEUP tape. Some sample data contained in TAPLOG, and printed out using program XTPLST is shown in Fig. 12. The total time represented on an individual tape can be found by multiplying the number of blocks by the number of seconds per block (which is 5 seconds per block for the tapes shown). The tape status indicates whether the tape is available for current or further processing (ACTIVE), or has been recycled (RECYCL).

Event Data Entry (XEVENT)

Another initial step in the WHIPS procedure is to create or update file EVENTS, which contains data about each earthquake for which an interval of data is to be saved. A sample of data contained in this file has been listed using program XEVLST and is shown in Fig. 13. Hypocenter/origin-time data are taken mostly from the Preliminary Determination of Epicenter (PDE) lists, published by the National Earthquake Information Service (NEIS), although other sources of data are used to complement the PDE's when

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WAKE HYDROPHONE INFORMATION PROCESSING SYSTEM

WAKEUP TAPE LOG LIST

		~-ST.	ART	T	IME			
TAPE#	YR	JLN	HR	MN	SC	MIL	#BLKS	STATUS
18561	83	8 76	Ø5	27	25	888	5783	ACTIVE
18562	83	<i>8</i> 76	13	29	2Ø	000	5791	ACTIVE
18563	83	<i>8</i> 77	Ø5	12	4.8	888	5766	ACTIVE
18564	83	Ø77	13	13	10	200	5824	ACTIVE
18565	83	Ø78	Ø5	15	15	888	5922	ACTIVE
18566	83	Ø78	13	28	45	888	5829	ACTIVE
18567	83	Ø79	Ø 4	34	45	888	5862	ACTIVE
18568	83	Ø79	12	43	15	888	5796	ACTIVE
18569	83	Ø8Ø	Ø5	17	25	888	5857	ACTIVE
18578	83	Ø8Ø	13	25	3₿	888	5885	ACTIVE
10571	83	Ø81	Ø6	15	50	888	5874	ACTIVE
18572	83	Ø81	14	25	2Ø	<i>888</i>	5827	ACTIVE
10573	83	Ø81	22	3 <i>Ø</i>	55	888	4895	ACTIVE
18574	83	Ø82	Ø5	18	5₿	888	586Ø	ACTIVE
10575	B3	Ø82	13	27	18	888	5787	ACTIVE
10576	83	Ø82	21	29	25	888	56#5	ACTIVE
18577	B 3	883	Ø 5	16	3Ø	888	5755	ACTIVE
18578	83	₿ ₿3	13	16	Ø 5	888	58#5	ACTIVE
18579	83	Ø83	21	19	5Ø	888	5696	ACTIVE
10580	83	Ø84	Ø5	14	3₿	888	5856	ACTIVE

Figure 12. A sample of the WAKEUP tape information contained within WHIPS file TAPLOG which has been printed using program XTPLST. A detailed description of this printout is contained in the text.

WAKE HYDROPHONE INFORMATION PROCESSING SYSTEM (WHIPS) EVENT LISTING

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KEV: INFORMATION SOURCES - PDE-WEIS PDECARD, MON-WEIS MONTHLY LIST, ILL-ICS LIST, HEL-HELICORDER, OTH-OTHER EVENT TYPE - FO-FARTH-DUALL, NX-N-HILLS RY-LOSION, SX-SCHENTHIC EXPLOSION, OF-OTHER, UN-UNFNOWN PHASES - A-P, B-PO, C-S, D-LO, E-T, F-OTHER

1																			
- ON-		: :	#0-0#	#151#	= 5		**COORDINAT.S**	IMAT. See	orest orest OCATION ************************************	90		-MAGNI-					• •	STRIP	
					;	:		5		E .			2	<u>د</u> د		PHASE SEE IN IT ROAL COMP	^	1A7E/11LE	
-	8		~			27:59:47.1	4.8265	151.8128	NIV IPFLAND RIGION	Ξ	-	6	104	-	O A B D	27:68-23:15		200427-11	_
Z.	æ	~	_			11:24:48.4	2.37 iN	126.8001	MOLDICA PASSAGE	, <u>,</u>		. 7		-		95:11-H2:11	12 21	20012/ 12	
E 1	•	_	_		٣.	13:45:49.8	5.22.7	1531.1531	NEW TREEAND PEGION	Ī	<u>.</u> د			_	ARDE	13:46-14:27		78842/ 17	~
- :	Œ :	~ ,			5	16:24:13.3	18.H.75	111.1679		9	5.6	3		-		16:26-16:37		28842/ 14	_
	¥ 3					=		177.67.10	_	F 1 9	-	E		-		17:46-17:57	12 21	28842/ 15	
1 1		• •					2.5	15 1.69HI	_	7		•	Ja 4.	1.0	1 ABDE	84:87-84:48		31-120WHZ	
2 2	-		- ~			#/:44:1/.5	71.66d.	75.2.2	10NGA	`		3		_	«	A7:47-47:59		ZHB42/ 17	
3	=		. ~				7	2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2	BARDA SEA	~ ~		9		~	«	BH: BH-BB: 12	~ C	2#842/ 17	
3 3						A . 3 . 5 7 . 5 1		175. 45.31	_	=	٠,	s :		_	⋖ ·	15:31-15:45		28842/ 18	_
=	, 3	. –			-	A 01.175.10	1 20 / N	700.07	ANDREADS 15.	Ξ.		. ·		_		15:53-16:84		200427 19	
	=					78:17:47		101	151. ALE		•			_		Ø1:34-Ø1:5Ø	17 2	28842/-2	_
88	3					Pb #9: 15.6	77 T Y	154 61.25	COLOMON ICLANDS	~ `	· .	8. S				7#:19-28:35		208427 21	
7 I HE	•	_	_			H6: 4F: 17. 1	6.6635	154 5591	CONTRACTOR TOTAL	• 3			104 7.	2:	ABOR	25:04-81:48	7	77 // 1000	
6815	×	_	_			88:26:55.2	6.6275			· ~	ی ا					57:74-18:44 57:44-18:44	7 7	27 /24887	
9 9	3	~	~		7 1	12:11:23.9	37.8114	71.4871	R BORDER RE	GION 1.		3		-		17:18-12:36		20802/ 20	
186	æ :	۰,	_		¥5	14:84:58.1	6.7B2S	154.5281	SOLUMON ISLANUS	چ	Š	\$			ABDE .	14:86-14:47	12 21	20043/ B1	
		٠,	٠,			15:8/:81.2	3.375	177.4911	RFGION	~	•	8. E				12:91-/8:91		20043/ 0 5	٠.
			- 6			61:54:36.3	37.224N	138,1756	MEAR W COAST OF HON', HU JAPAN	7H 7:	•	6. 8 t		_		21:56-22:11		2M843/ 83	_
								M. M. G.	LUCAL PO. SO. T	į	3	# . O. E		_		BF: 22-88:51		20043/ 64	_
	9 2					67:54:48.4	2 23 2	7.7.	SOLUMON ISLANDS	S .	<u>.</u>	Z :		_	1 ABD	81:49-69:49		20043/ 85	
2.3	=	, ,					7.0133		CAVA	,r.	•			-		87:38-87:41	12 21	288437 86	
1280	Œ	, ~	•			18:47:48.5	4 2215	15.1 7365	MEC - DEC AND BLC - DE	<i>.</i> :	٠.							2PR43/ 67	
6 H . 5	=	_				12:24:35.2	10.8145	: -	COLOMON IN AND							18:48-11:29		88 /EVAN/	
68.3		-				96:98:98.6	21.0	19119	PROBABLE MARIANAS PO SO 1				2 2	2 5	0 2 2	10:21-52:21	\ \ !	20 /2000/ 2000/	
1288	Œ	_	_			87:44:B1.2	6.5555	130.0v61	;		•			٠.		60 - 10 - 70 - 34 67 - 47 - 67 - 50		10000	
826	•	~	~			3:66		151.777	NIV BRITAIN REGION	•		6		-	ARDE	•		C	
6 2 H G		_,	٠.			19:33:54.1		124.6HJT	TIMOR	7. T	-	3		-	. <	19:37-19:49	13 21	28843/ 13	
	9 3	٠.	٦.				٦.	155.1189	HON ISLANDS	-	<u>.</u>	E .	10.1	Ξ	1 AB0E	12:22-21:12		20043/ 14	
2 8 8	. 4					17.45.47.4	N T T T T		M.V. IRAN-USSP BORDER REGION		-	2	-	_	4	2:		288437 15	_
643	•	· ~			ري . ت	14:57:49.3	1.1678	100.00	MOUNTED TO THE MOUNTED THE MOUNTED TO THE MOUNTED TO THE MOUNTED T	~ •	· .	= •				12:47-13:82		. A643/ 16	
6 834	æ	3	~	26 01	OH5	17:19:17.5	4.6315	14.8.13	PAPUA NEW CITTLES	3	٠,			<u> </u>	4 4	15:21-18:51	2	/	
8 35	90 (m (~ ·		ž.	IB: 42: 46.8	42.847N	141.744	HUKLATED JAPAN REGION	2		. z		_		15 - 51 - 17 · d ·		01 /5788	
25.00	9 3	٦.				8.80:22:02	NIBE /E	115.45.7	SOUTH FIN PLIVADA	•	<i>.</i> :	±.		z	< <	78:8-18:37		1844/ 81	
BH 3R	30	. ~					34. / him		CINCHAL PPOVINCE CHIMA	ς.	•	ر ح		-	⋖	F.:48-#3:8#		: P844/ 82	٠.
BB 3.9	8	_	-			18:32:08.1	A BOHS	15.2 5.6.5	MIN THE PARTY OF T	- :	•		-		A80	61:68-88:18		PP44/ 83	_
9119	Œ	_	_			211:05:07.6	6,89.45	14.1.641		,					A 400 .			78684/ B4	. .
7 =	x	~	_			Ë	26. JR6N	127.311.	RVUKYU 1 JANDS	٠.						C: 07-90:47		10 / 10 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1	
	Φ.	_				\sim	1 9H5H	146. 35.40	SOUTH OF HARTANA ISLANDS	-					(4	22.24.21.61	2 2	74420	
	x :	. .				84:53:69.4	34.44.38		FYU, HU JAPAN	_	-	-		_		M4:55-85:11			_
	D 8						6.984	٠	SOLUMON ISLANDS	-		-		_		37:57-73:38		PB 44 89	
	* 3	٠.	٠,				E / L	12		•	•	-		-	⋖ -	MIT 53-11 : 84		10844/ 18	_
984	æ	. ~	-			2	6.6375	154 56.0	FINISHIZENIA TOTAL BURDER REG	REGION		· :		-		PIC : 5.3-1.7:05	13 27	20044/ 11	_
9199	20	_	~		5	16:13:45.2	39. RB4N	75.45.01	STREET STREET		•					1. : 64-12: : 18	12 51	20044/ 12	
5760	œ	_	_		2	14:55:49.1	5.9195	111.7.7.01	PAP IA NIV CHILLA		ي ا	. <u>-</u>		-	, , ,	16:21-16:3.	2	EL /9994?	
F 2 F	20	~	~	-	£.	•	34.281N	135.81.11	MIA'S CHA'T HE SOUTH HONSHU	HO Dir	-	: =	-	. <u>.</u>				``	
																	;	,	

A sample of event information contained within WHIPS file EVENTS, which has been printed using A detailed description of this printout is contained in the text. Figure 13.



necessary. A general procedure has been formulated for determining which events from the PDE lists might produce seismic energy observable at Wake, based on experience gained from scanning the analog data collected since 1979. An outline of event types and their corresponding phases is presented in Fig. 14. Intervals to save for these events and phases are computed automatically by program XEVENT, at the time the hypocenter data are entered, and may be altered manually if necessary. The only seismic energy from events which is deliberately discarded is that of T-phases from smaller magnitude and unknown events. T-phase energy is generally so abundant (it is estimated that at least 10 times as many circum-Pacific events may be detected using T-phases than may be detected using other seismic phases) that very little data compression would be possible if all of these T-phase data were saved. This procedure is weighted towards saving more data than necessary (i.e., many of these events may not have observable energy at Wake, but this is hard to determine without scanning all the channels in particular frequency bands) so that a) unusual events are not discarded, b) data from small events may be enhanced with array processing, and c) data from events which are not of current interest but may be of future interest are saved. The helicorder records from Wake are also scanned for seismic phases such as P, Po, S, and So which do not correspond to known events, and intervals are saved to capture these phases. Two such intervals, corresponding to events 820 and 826, appear in Fig. 13. The final column of the listing shown in Fig. 13 gives the strip tape and file number which contains the corresponding interval of data. Some data intervals are contiguous or overlap and have therefore been saved in a single file (for example: events 807 and 808, and events 813 and 814). Other events have not yet been extracted from the WAKEUP tapes and are noted

	Source Region(s)	Magnitude	Depth(km)	~△	Phases
1.	Nuclear Explosions Anywhere	all	a11	0°-35°	P.Po.So.T
		all	all	> 35°	P
2.	Mariana Is., Bonin Is., Volcano Is., Honshu,	m _b < 5.0	<100	18°-25°	P.Po.So
	and Gilbert Islands	m _b > 5.0	<100	18°-25°	P.Po.So.T
		all	>100	18°-25°	P.Po.So.T
3.	Kuril Is., Kamchatka, Aleutian Is., New Guinea,	m _b > 5.0	<100	25°-35°	P.Po.So.T
	New Britain, Solomon Is., and Santa Cruz Islands	all	>100	25°-35°	P.Po.So.T
4.	Asian Continent	m _b > 4.5	all	30°-90°	P
5.	Mid Pacific Plate Events	all	a 11	0°-35°	P.Po.So.T
		m _b > 4.5	all	>35°	P,Po.So.T
6.	Not Regions 2, 3, 4, or 5.	m > 5.0	all	<90°	P
7.	Mid Atlantic Antipode	al1	all	340°-360°	P
8.	Anywhere	M > 6.0	all	a 11	P
		m _b > 6.5	a11	340°-360°	P

Figure 14. The general criteria used for deciding which phases (if any) to save for a known event. Time intervals which contain those phases are computed automatically by XEVENT or may be entered manually.

by a tape and file number of 0 (such as event 850). Some events, such as 801. 806, and 811, could not be extracted because there was no WAKEUP data available (in this case, due to a tape drive failure at WAKE). A file for these events, containing only header and trailer records, exists on the strip tape indicated, and the file number has been flagged as negative on the XEVLST listing.

XSTRIP Parameters (XPARAM)

The last initial step in the WHIPS procedure is to set up the running parameters for program XSTRIP which are stored in the header block of the EVENTS file. These parameters may be viewed and changed using program XPARAM, and a sample output from this program is shown in Fig. 15. Many of these parameters (D, E, F, G, H, and I) are continually updated by programs XSTRIP, XSTPDN, and XRANDN; and these parameters may be changed manually when necessary. Other parameters (A, B, and C) must always be entered and changed manually as the situation warrants. The "WAKEUP tape record length" (A) is equal to the WAKEUP sampling rate times the number of channels, times the number of seconds per block, plus 100 words for the block header (i.e., $[80 \times 8 \times 5] + 100 = 3300$). The "strip/interval disk record length" (A) is the block length used for disk files STRIP000 and INTRVL00; and it must be a multiple of 112 words (Harris sector size = 112 words) which is greater than or equal to the WAKEUP tape record length. The "maximum records ..." (B) is a conservative estimate of how many blocks of data (in this case, 3300 word blocks), will fit onto a tape created by XSTPDN or XRANDN. This parameter is necessary for XSTPDN to organize its output data so that individual intervals which are saved do not cross XSTPDN tape boundaries. The "seconds per WAKEUP tape block" (C) is used to convert terms which refer to a

WHIPS-->DEFINE SYSTEM PARAMETERS.

A.	WAKEUP TAPE RECORD LENGTH	IS		3388
	STRIP/INTERVAL DISC RECOR	LENGTH I	S	336 <i>8</i>
В.	MAXIMUM RECORDS ON EVENT/	INTERVAL T	APE IS	558 <i>8</i>
С.	SECONDS PER WAKEUP TAPE BI	LOCK IS		5
D.	NEXT WAKEUP TAPE TO STRIP	15		18588
	LAST WAKEUP TAPE TO STRIP	15		Ø
Ε.	LAST RECORD ON STRIP DISC	FILE IS		Ø
F.	CURRENT STRIP TAPE NUMBER	IS		28844
	LAST FILE WRITTEN'IS			5
	NUMBER OF TAPE RECORDS IS			1222
G.	LAST RECORD ON INTERVAL D	ISC FILE I	S	Ø
н.	CURRENT INTERVAL TAPE NUM	BER 1S		388 29
	NUMBER OF TAPE RECORDS IS			3856
I.	RANDOM INTERVAL 2626	666 888888	2	626668179995

ENTER A-J, U (UPDATE), OR X (EXIT).

Figure 1. A part of the interactive I/O generated by WHIPS program XPARAM which is used to view or change the running parameters of program XSTRIP. More detailed explanation of these parameters is given in the text.

particular block on a WAKEUP tape into absolute times (for example - to calculated the ending time of a WAKEUP tape, given its starting time and total number of blocks). The "next" and "last WAKEUP tape to strip" (D) refers to the range of WAKEUP tapes which are to be processed during a particular run of XSTRIP. These parameters are continually updated during the XSTRIP run to facilitate recovery in case of a system crash. At the completion of the run, the last tape which was processed becomes the "next WAKEUP tape to strip", and the "last WAKEUP tape to strip" is set to 0, as shown in the figure. The "last record on the strip disk file" (E) allows XSTRIP to begin writing into STRIP000 where it last left off. This parameter is set to 0 after a successful downloading of the data using XSTPDN, as is shown in the figure. The "current strip tape", "last file", and "number of tape records" (F), allows XSTPDN to download the data onto tape from the point on the tape at which it left off after the last run. G and H serve the same function for the random interval data as did E and F for the stripped events data. The "random interval" (I) gives the 3 minute time range, in century-milliseconds, of the next random interval to be stripped.

Extracting the data (XSTRIP)

XSTRIP is the program which extracts and stores data of events and of random intervals on disk for later downloading. The only operator input required is specification of the range of WAKEUP tapes to be processed.

Normally, this processing takes place in contiguous chronological order. After the total time interval represented by the WAKEUP tapes has been computed from the TAPLOG file, a search of the EVENTS file is made to determine which event intervals, not yet extracted, lie entirely within the

total time interval. From these individual event intervals, grand intervals are computed which combine the times of overlapping event intervals. As the WAKEUP tapes are read one by one (the operator is prompted to mount each tape), the grand intervals are extracted and stored on disk file STRIP000. Concurrently, the random, three-minute intervals are being extracted and stored on disk file INTRVL00. The timing of these intervals is computed "on the fly" and the number of minutes of data between the beginnings of one 3minute interval and the next is computed using a random variable with a uniform distribution between 3 (i.e., no overlap) and 117, giving an average separation of 60 minutes. The parameters list at the beginning of the EVENTS file is updated after the extraction of each grand interval and each random interval. This facilitates a smooth and efficient recovery after an XSTRIP abortion of any kind and provides information for later downloading of the data to tape. XSTRIP is complete after the last WAKEUP tape has been run. It should be noted that a typical XSTRIP run at HIG processes from 10 to 20 WAKEUP tapes, and requires approximately 20 million words of disk storage for the STRIP000 and INTRVL00 files.

Events Download to Strip Tapes (XSTPDN)

The STRIPOOD disk file containing the event data is downloaded to tape using program XSTPDN. This program first accesses the running parameters in the EVENT file to find out which tape to download to, and how many files to advance into that tape before writing. As each grand interval in the STRIPOOD file is downloaded to tape, the running parameters are updated, and the EVENTS file is updated with the strip tape number and file for those events contained in the grand interval. If the grand interval contains too many records to completely fit on the strip tape (based on "B" in the

running parameters). then an end-of-file is written to the strip tape and a new strip tape is mounted following a prompt to the operator. Strip tapes are numbered sequentially beginning with 20001. The XSTPDN run is complete when all the grand intervals have been downloaded, and at that time the "last record on strip disk file" (E) in the running parameters is set to 0 as shown in Fig. 15. The format of the strip tape, and format of the blocks contained on the strip tape is given in Fig. 16.

Random Interval Download to Interval Tapes (XRANDN)

The INTRVLOO disk file, containing the 3-minute randomly spaced ambient noise samples, is downloaded to tape using program XRANDN. This program first accesses the running parameters to determine which interval tape to download to. and how many records from the INTRVLOO file to tape until the maximum number of records on tape ("B" in the running parameters) is reached. There are no end-of-files between the 3-minute intervals, however, an end-of-file is written after the last record on the interval tape. Interval tapes are numbered sequentially beginning with 30001. The XRANDN run is complete when all of the records in INTRVLOO have been downloaded, at which time item "G" in the running partameters is set to 0 as shown in Fig. 15. The format of the interval tape, and the format of the blocks contained on the interval tape is given in Fig. 17.

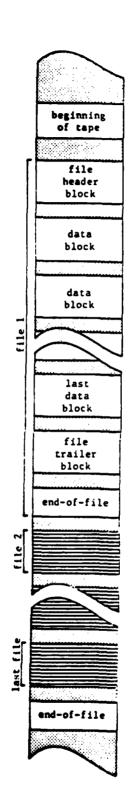
Checking Procedures (XSTPCHK and XINTCHK)

As a safety measure, and as a way of accumulating some preliminary.

statistics on the data, all strip and interval tapes are checked using

programs XSTPCHK and XINTCHK. respectively. The outputs from these programs

are viewed before the STRIPOOO and INTRVLOO files are eliminated from disk



WHIPS Strip Tape Description

General Format: 1600 BPI

2 bytes per word IBM format

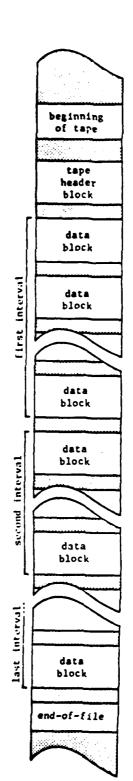
2's complement notation

File Header Block						
Word(s)	Description					
1	Always 0 to identify block as a header/trailer					
2	Strip tape I.D. number					
3	Strip file number					
4-56 Not used						
57	Estimated number of data block in this file					
58-62	Not used					
63	Total number of events (n) represented in this file					
64-100	Not used					
101-(101+n)	Event number(s)					
(102+n)-end	Not used					

		Data	Block	L						
The format of of the WAKEUP	the strip tape tape blocks.	data blo	cks is	exactly	the	same	88	the	format	

	File Trailer Block	
Word(s)	Description	
1-3	Same as header block words 1-3	
4-56	Not used	
57	Actual number of data blocks in file	
58	Number of words per block	
59-63	Same as header block words 59-63	
64-65	WAKEUP tape I.D. number(s)	
66	Number of WAKEUP tape read errors	
67-86	Block numbers with WAKEUP tape read errors	
87-end	Same as header block words 87-end	

Figure 16. Format for the WHIPS strip tapes which contain the event data.



WHIPS Interval Tape Description

General Format: 1600 BPI

2 bytes per word

IBM format

2's complement notation

Word(s)	Description	
1 2 3-end	Always 0 to identify block as a tape header Interval tape I.D. number Not used	

Data Block

The format of the interval tape data blocks is exactly the same as the format of the WAKEUP tape blocks.

Figure 17. Format for the WHIPS interval tapes which contain randomly spaced 3-minute intervals of data (1 per hour on the average) which are collected as noise samples.

and the WAKEUP tapes recycled. A sample output from XSTPCHK is shown in Fig. 18. Most of the lines of output are self-explanatory. The "error blocks" are for blocks of data which had tape read errors coming off of the WAKEUP tapes during XSTRIP. The "clipped data" lines show the number of data points in a particular block, and for a particular channel, which are equal to -32,768 or 32,767, the maximum and minimum values output by the A/D convertor at Wake. The clipped data in this file is on two SOFAR hydrophones and corresponds to the T-phase arrival from event 824 (see Fig. 11 for event 824). The table of values gives information about each multiplexed in the file which is described as follows: 1) "ch" - multiplexed channel number; 2) "hydr" - hydrophone number; 3) "amp" - preamp/prewhitening filter identification number; 4) "gain" - gain step of the final stage of the preamp/pre-whitening filter (see previous section on the recording system); 5) "avgval" - the average value of this channel (non-zero due to small but constant DC biases in the amplifiers); 6) "stddev" - the standard deviation of each channel (note the much higher values for the noisier SOFAR hydrophones); 7) "0/blk" - average number of zeros per block (note that channel 3 has most zeros due to its near zero average value and small standard deviation); 8) "0/blk/std" - the standard deviation of the number of zeros per block; and 9) "# clips" - the total number of clipped data points for this channel in this file. The "worst case clock error code" indicates the maximum error between the satellite clock time written in each block header and the actual coordinated Universal Time. The codes are as follows:

^{0 -} less than + 1 msec

^{1 -} less than + 5 msec

```
TAPE FILE NUMBER:
WHIPS STRIP TAPE: 20043 WHIPS FILE:
                                       8
NUMBER OF EVENTS IN THIS FILE:
EVENT NUMBER(S):
                 824
NUMBER OF ERRORS READING WAKEUP TAPE(S):
ERROR BLOCKS:
NUMBER OF MULTIPLEXED DATA CHANNELS:
SAMPLES PER SECOND PER CHANNEL: BO
NUMBER OF SECONDS PER BLOCK: 5,000
TIME-FIRST DATA BLOCK, 1983 03 24 (083) 10:48:00.000
RECORD: 393 CLIPPED DATA CH/HPTS:
                                    7/
                                        2
RECORD: 395
            CLIPPED DATA CHAMPTS:
                                    7/
RECORD: 396 CLIPPED DATA CH/#PTS:
                                    7/
RECORD: 400 CLIPPED DATA CH/#PTS:
                                    B/ 10
RECORD: 401 CLIPPED DATA CH/#PTS:
                                    8/ 12
RECORD: 402
            CLIPPED DATA CH/#PTS:
                                    8/ 48
RECORD: 403 CLIPPED DATA CH/#PTS:
                                    8/ 51
RECORD: 404
            CLIPPED DATA CH/#PTS:
                                    8/ 10
RECORD: 405 CLIPPED DATA CHIMPTS:
                                    8/8
TIME-LAST DATA BLOCK: 1983 03 24 (083) 11:29:55.000
RECORD
        506 IS A TRATLER RECORD.
CH HYDR AMP GAIN AVGVAL STDDEV O/BLK O/BLK/STD
   71
              9
                  396.4
                          35.2
                               0.00
                                       0.000
                                                    O
                 -430. B
2
   73
                                       0.000
              9
                          66.4
                               0.00
                                                    0
3
   74
         0
              9
                   32. 2
                          58.8 4.38
                                       2.202
                                                   0
   75
4
         19
              9
                  391.7
                          88. 6 0. 00
                                       0.000
                                                   0
5
   76
              9
         1
                 109.0 238.0 1.07
                                       1.005
   10
         13
              9
                  129. 3 1363. 5 0. 59
                                       0.762
                                                   0
7
   20
         11
                   32.6 1388.4 1.95
                                       1.682
                                                   14
8
  40
         18
              9 -189, 9 2266, 2 0, 22
                                       0.473
                                                 139
THE WORST CASE CLOCK ERROR CODE:
FILE:
TOTAL RECORDS IN THIS FILE (INCL. HEADER): 506
```

Figure 18. Sample output from program XSTPCHK. Detailed explanation of this figure is given in the text.

- 3 less than + 50 msec
- 7 less than + 500 msec
- $15 more than \pm 500 msec$

A sample output from XINTCHK is shown in Fig. 19. Much less computation is done on the interval data, and the XINTCHK output serves more as a simple log of what has been saved.

Strip Tape Copies (XTAPCOP)

As a final step in the preliminary processing, the strip tapes are copied using program XTAPCOP, and the copies are sent to the DARPA Center for Seismic Studies (CSS) in Rosslyn, Virginia. Copies of interval tapes are not currently sent to CSS but may be at some time in the future.

```
RECORD
          1 IS A HEADER RECORD.
#BLKS: 21
                 22 INT: 1983 84 28 (118) 82:18:15-82:28:88 #SEC: 185.888
#BLKS: 36
            23-
                 58 INT: 1983 84 28 (118) 83:38:88-83:41:88 #SEC: 188.888
#BLKS: 36
            59-
                 94
                    INT: 1983 64 26 (118) 85:28:88-85:23:88 #SEC: 188.888
#BLKS: 36
            95- 130 INT: 1983 04 20 (110) 06:47:00-06:50:00 #SEC: 180.000
#BLKS: 36
           131- 166 INT: 1983 84 28 (118) 87:83:88-87:86:88 #SEC: 188.888
                    INT: 1983 Ø4 2Ø (11Ø) Ø8:Ø7:ØØ-Ø8:1Ø:ØØ #SEC: 18Ø.ØØØ
●BLKS:
      36
           167- 282
           203- 238
                    INT: 1983 84 28 (118) 89:57:88-18:88:88 #SEC: 188.888
#BLKS: 36
●BLKS: 36
           239- 274 INT: 1963 04 28 (110) 11:18:00-11:21:00 #SEC: 180.000
                    INT: 1983 84 28 (118) 11:41:88-11:44:88 #SEC: 188.888
#BLKS:
      36
           275- 31Ø
                    INT: 1903 84 28 (118) 12:28:88-12:31:88 #SEC: 188.888
#BLKS: 36
           311- 346
           347- 382
#BLKS: 36
                    INT: 1983 84 28 (118) 13:13:88-13:16:88 #SEC: 188.888
#BLKS: 36
           383- 418
                    INT: 1983 04 20 (110) 15:83:88-15:86:88 #SEC: 188.888
                    INT: 1983 84 28 (110) 15:54:00-15:57:00 #SEC: 180.000
#BLKS: 36
           419- 454
#B∟KS: 36
           455- 495 INT: 1963 04 26 (116) 16:22:80-16:25:80 #SEC: 180.000
#BLKS:
      36
           491- 526
                         1983 84 28 (118) 18:23:88-18:26:88 #SEC: 188.888
                    INT:
                    INT: 1963 84 28 (118) 19:38:88-19:33:88 #SEC: 188.888
#BLKS: 36
           527- 562
#BLKS:
      36
           563- 598 INT: 1983 04 20 (110) 28:48:88-28:51:88 #SEC: 188.888
#BLKS: 36
           599- 634
                    INT: 1983 84 28 (118) 21:48:88-21:43:88 #SEC: 188.888
           635- 67#
                    INT: 1983 84 28 (118) 22:29:88-22:32:88 #SEC: 188.888
#BLKS: 36
#BLKS:
      3€
           671- 786
                    INT: 1983 84 28 (118) 23:18:88-23:13:88 #SEC: 188.888
           787- 742
#ELKS: 36
                    18T: 1983 84 28 (118) 23:55:88-23:58:88 #SEC: 188.888
#bLKS: 36
           743- 778
                    INT: 1983 84 21 (111) 81:87:88-81:18:88 #SEC: 188.888
#6LKS: 36
           779- 814 INT: 1983 04 21 (111) 01:33:00-01:36:00 #SEC: 180.000
#BLKS: 36
           815- 85Ø
                    INT: 1983 84 21 (111) 82:26:88-82:29:88 #SEC: 188.888
           851- 886
#BLKS: 36
                    INT: 1983 Ø4 21 (111) Ø4:15:00-84:18:00 #SEC: 180.000
#BLKS: 36
           887- 922 INT: 1983 84 21 (111) 85:84:88-85:87:88 #SEC: 188.888
#BLKS: 36
           923- 958
                    INT: 1983 #4 21 (111) #6:13:##-#6:16:## #SEC:
                                                                    180.000
#BLKS: 36
           959- 994 INT: 1983 Ø4 21 (111) Ø7:27:ØØ-Ø7:3Ø:ØØ #SEC: 18Ø.ØØ
#BLKS: 36
           995-1838 INT: 1983 84 21 (111) 88:13:88-88:16:88 #SEC: 188.888
#BLKS: 36
          1831-1866 INT: 1983 84 21 (111) 18:84:88-18:87:88 #SEC: 188.888
#BLKS: 36
         1867-1182 INT: 1983 84 21 (111) 12:84:88-12:87:88 #SEC: 188.888
#BLKS: 36
         1103-1138 INT: 1983 84 21 (111) 12:22:88-12:25:88 #SEC: 188.888
#BLKS: 36
         1139-1174 INT: 1983 Ø4 21 (111) 12:27:00-12:30:00 #SEC:
                                                                    180.000
#BLKS: 36 1175-1210 INT: 1983 04 21 (111) 13:15:00-13:18:00 #SEC: 180.000
#BLKS: 36 1211-1246 INT: 1983 Ø4 21 (111) 13:48:00-13:51:00 #SEC: 180.000 #BLKS: 72 1247-1318 INT: 1983 Ø4 21 (111) 14:47:00-14:53:00 #SEC: 360.000
●BLKS: 36 1319-1354 INT: 1983 Ø4 21 (111) 15:19:8Ø-15:22:8Ø ●SEC: 18Ø.ØØØ
       36
         1355-1390 INT: 1983 04 21 (111) 16:41:00-16:44:00 #SEC: 180.000
#BLKS: 36 1391-1426 INT: 1983 04 21 (111) 18:07:00-18:10:00 #SEC: 180.000
#BLKS: 36 1427-1462 INT: 1983 Ø4 21 (111) 18:53:ØØ-18:56:ØØ #SEC: 18Ø.ØØØ
#BLKS:
       36
         1463-1498 INT:
                         1983 Ø4 21 (111) 19:27:00-19:30:00 #SEC: 180.000
#BLKS: 36 1499-1534 1NT: 1983 04 21 (111) 19:38:00-19:41:00 #SEC: 180.000
#BLKS: 36 1535-1578 INT: 1983 84 21 (111) 21:82:88-21:85:88 #SEC: 188.888
#BLKS:
          1571-1606
                    INT:
                         1983 Ø4 21 (111) 22:Ø3:ØØ-22:Ø6:ØØ #SEC:
                                                                    190.000
#BLKS: 36 1687-1642 INT: 1983 84 21 (111) 22:18:88-22:21:88 #SEC: 188,888
#BLKS: 36 1643-1678 INT: 1983 Ø4 22 (112) ØØ:Ø8:Ø8-ØØ:11:ØØ #SEC: 18Ø.ØØØ
#BLKS: 36 1679-1714 1NT: 1983 Ø4 22 (112) ØØ:18:ØØ-ØØ:21:ØØ #SEC: 18Ø.ØØØ
```

Figure 19. Sample output from program XINTCHK showing some of the 3-minute intervals saved on a WHIPS interval tape. The first interval shown, blocks 2-22, contains only 105 seconds because part of that interval was written to the previous interval tape. Another interval, blocks 1247-1318, contains 360 seconds and represents two contiguous 3-minute intervals.

ACKNOWLEDGEMENTS

This system has been sponsored primarily by the Air Force Office of Scientific Research with supplemental funds provided by the U.S. Arms

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George Sutton, Paul Jubinski, Grant Blackinton and Joe Gettrust for their respective contributions towards the design and implementation of this system. Firmin Oliveira successfully programmed most of the WAKEUP and WHIPS software in less time than seemed possible. Al David, Bonnie Jose, and Kentron Corporation have provided excellent day to day operation of the system at Wake.

APPENDIX X

KEY: INFORMATION SQURCES - PDE-NEIS PDECARD, MON-NEIS MONTHLY LIST, ICS-ICS LIST, HEL-HELICORDER. OTH-OTHER EVENT TYPE - EG-EARTHOUAKE, NX-NUCLEAR EXPLOSION, SX-SCIENTIFIC EXPLOSION, OT-OTHER, UN-UNKNOWN PHASES - A-P. B-PO, C-S, D-SO, E-T, F-OTHER

EVNT ******ORIGIN TIME******	**COORDINATES***	**************************************	DEP KM=	"MAGNI" IN	EV *****	***SAVED******* S**INTERVAL***MNS	**STRIP** TAPE/FILE
### ### ### ### ### ### ### ### ### ##	38.299N 137.678E 3.386S 177.566E 55.597S 182.453W 43.563N 146.686E 15.493N 147.571E 15.518N 147.593E 15.638N 147.593E 15.638N 147.648E 3.437S 177.658E 2.447N 126.918E 2.628S 176.544W 26.883S 176.218W 15.611N 147.818E 2.628S 176.544W 26.883S 176.218W 15.611N 147.818E 2.727S 177.914W 26.952S 176.4787 26.952S 176.4787 27.27S 177.914W 26.952S 176.4787 27.27S 177.914W 37.526N 142.782E 3.442S 177.667E 3.442S 177.667E 3.442S 177.667E 3.442S 177.667E 3.442S 177.672E 3.442S 177.682E 2.822N 143.7498W 27.247S 179.615W 27.247S 179.615W 27.247S 179.615W 27.247S 179.615W 27.247S 177.672E 3.568S 177.682E 2.822N 143.743E 3.526S 177.684E 15.335S 173.283W 3.518S 177.684E	SOUTH OF HONSHU, JAPAN GILBERT ISLANDS REGION SOUTHERN PACIFIC OCEAN. KURIL ISLANDS MARIANA ISLANDS REGION MARIANA ISLANDS REGION MARIANA ISLANDS REGION MARIANA ISLANDS REGION MARIANA ISLANDS GILBERT ISLANDS GILBERT ISLANDS GILBERT ISLANDS SOUTH OF FIJI ISLANDS 49 DEGS. SOUTH OF FIJI ISLANDS A9 DEGS. MARIANA ISLANDS MOKKAIDO, JAPAN FIJI ISLANDS, ALEUTIAN ISLANDS. FOX ISLANDS, ALEUTIAN ISLANDS. FOX ISLANDS, ALEUTIAN ISLANDS. FOX ISLANDS, ALEUTIAN ISLANDS. FIJI ISLANDS. MARIANA ISLANDS. GILBERT ISLANDS. MARIANA ISLANDS. MARIANA ISLANDS. KERMADEC ISLANDS. SOUTH OF MARIANA ISLANDS SOUTH OF HARIANA ISLANDS SOUTH OF HARIANA ISLANDS SOUTH OF HARIANA ISLANDS TONGA ISLANDS. NEW BRITAIN. GILBERT ISLANDS. NEW BRITAIN.	481 333 333 333 333 333 333 333 333 333 3	5.4 #.3 POI 5.4 #.3 POI 5.4 #.3 POI 6.7 #.5 POI 5.2 #.5 POI 5.3 #.5 POI 5.5 #.6 POI 5.6 #.6 POI 5.7 #.6 POI 5.8 #.6 POI 5.8 #.6 POI 5.9 #.6 POI 5.9 #.6 POI 5.1 #.6 POI 5.1 #.6 POI 5.2 #.6 POI 5.3 #.6 POI 5.4 #.6 POI 5.5 #.6 POI 5.6 #.6 POI 5.7 #.6 POI 5.8 #.6 POI 5.9 *	EQ ABDE	## 3: ## 6- ## 3: 5# 43 ## 5: 44- ## 6: 22 39 ## 6: 44- ## 6: 22 39 ## 6: 44- ## 6: 22 39 ## 6: 44- ## 6: 22 31 ## 6: 44- ## 6: 22 31 ## 6: 44- ## 6: 42 31 ## 6: 44- ## 6: 42- ## 6: 43- ## 6: 44- ## 6: 43- ## 6: 44- ## 6: 43- ## 6: 44- ## 6: 43- ## 6: 44- ## 6: 43- ## 6: 44- ## 6: 43- ## 6: 44- ## 6: 43- ## 6: 44- ## 6: 43-	28981/ 86 28981/ 87 28981/ 87 28981/ 88 28981/ 89 28981/ 19 28981/ 19 28981/ 19 28981/ 19 28981/ 19 28982/ 83 28982/ 83 28982/ 86 28982/ 87 28982/ 88 28983/ 87 28983/ 89
##51 82 #9 12 255 23:81:88.5 ##52 82 #9 12 255 23:16:58.8 ###53 82 #9 12 255 23:33:23:38:54 ###54 82 #9 12 255 23:33:23:38:54 ###55 82 #9 12 255 23:33:32:3 ###54 82 #9 12 255 23:33:32:3 ###55 82 #9 12 255 ###5:37:1 ###56 82 #9 17 26# ##2:57:48.1 ###59 82 #9 17 26# ##2:57:48.1 ###59 82 #9 17 26# ##3:28:25.6 ###61 82 #9 17 26# 13:28:25.6 ###61 82 #9 17 26# 13:28:25.6 ###61 82 #9 17 26# 28:21:43:1.7 ###58 82 #9 17 26# 28:21:43:1.8 ###62 82 #9 18 261 13:28:16.6 ###61 82 #9 18 261 13:28:31.8 ###63 82 #9 18 261 13:28:31.8 ###63 82 #9 18 261 13:28:31.8 ###64 82 #9 18 261 13:28:31.8 ###66 82 #9 18 261 13:28:31.8 ###68 82 #9 18 261 21:28:33.8 ###68 82 #9 18 261 21:28:31.8 ###68 82 #9 18 261 21:28:31.6 ###68 82 #9 18 261 21:28:31.6 ###68 82 #9 28 263 13:48:32.5 ###68 82 #9 28 263 13:48:32.5 ###68 82 #9 28 263 13:48:32.5 ###68 82 #9 28 263 13:48:32.5 ###68 82 #9 28 263 13:48:32.5 ###68 82 #9 28 265 13:356:46.6 ###78 82 #9 28 265 13:356:46.6 ###78 82 #9 28 265 13:356:31.7 ###78 82 #9 21 264 ##:28:231.7 ###78 82 #9 21 264 ##:28:231.7 ###78 82 #9 21 264 ##:28:231.7 ###78 82 #9 21 264 ##:28:31.7 ###78 82 #9 21 264 ##:28:31.7 ###78 82 #9 21 265 ##:17:32.3 ###78 82 #9 21 265 ##:17:32.3 ###78 82 #9 22 265 ##:17:32.3 ###78 82 #9 23 266 ##:231.3 ###878 82 #9 23 266 ##:231.3 ###888 82 #9 23 266 ##:28:46.7 ###888 82 #9 23 266 ##:28:46.7	44.276N 168.826E 44.296N 168.891E 44.316N 168.981E 44.316N 168.981E 44.316N 168.982E 44.359N 161.837E 11.151S 162.212E 17.725S 172.756W 3.489S 177.688E 23.446S 179.947W 11.852S 179.6174 23.426S 179.6174 23.426S 179.6174 26.924S 175.6174 26.924S 175.7477E 26.924S 175.934W 26.878S 175.934W 26.878S 175.934W 27.116S 143.723E 26.858S 176.461W 27.116S 176.663W 27.116S 176.663W 27.116S 176.663W 27.116S 176.673W 27.116S 176.674E 28.388W 171.314E 39.388W 171.314E 39.388W 171.314E 39.388W 171.314E 39.388W 171.314E 39.388W 171.314E 39.388W 126.7716 33.378 177.684E 37.389W 126.7716 33.398 129.7795E 38.3898 129.7795E 38.3898 129.7795E 38.3988 178.612F 38.3988 178.612F 38.3988 178.612F 38.3988 178.612F 38.3988 178.612F	KURILS DOWNHOLE - 248# SHOT SOLOMON ISLANDS TONGA IS. REG. GILBERT IS. REG. S. OF FIJI IS. S. OF SUMBAWA IS. S. OF FIJI IS. S. OF FOR SUMBAWA IS. S. OF FIJI IS. S. OF FOR SUMBAWA IS. REG. DENTRECASTEAUX IS. REG. DENTRECASTEA	######################################	### ### ### ### ### ### ### ### ### ##	SXXXAAABDD E EEQAAABDD E EEQAAABDDD E EEQAABDDD E EEQAAABDDD E EEQAAABDDD E EEQAABDDD E EEQAAABDDD E EEQAABDDD E EEQAABDDD E EEQAABDDD E EEQAABDDD E EEQAABDD E EEQAABDDD E EEQAABDDD E EEQAABDDD E EEQAABDDD E EEQAABDDD E EEQAABDD E EEQAAABDD E EEQAABDD E EEQAABD E EEQAABDD E EQAABDD E EEQAABDD E EQAABDD E EQAABDD E EQAABDD E EQAABDD E EQAABDD E EQAABDD E EQAABD E EQAABDD E EQAABD E EQAABDD E EQAABDD E EQAABDD E EQAABD E EQAABD	22:45-23:24 48 23:81-23:48 48 23:81-23:48 48 23:31-23:48 48 23:33-88:12 48 82:59-83:44 48 82:59-83:43 45 84:24-48:43 12 89:41-18:89 29 13:37-13:57 21 18:43-15:54 12 28:24-28:35 12 14:25-15:84 48 15:84-15:15 12 16:21-18:48 28 21:24-21:35 12 13:81-14:48 13 14:85-14:24 28:24-28:35 12 13:81-81-81 14:85-14:24 28:24-28:35 12 13:81-88:32 12 83:81-88:32 15 12:88-12:14 13:84-83:15 12 17:57-18:88 12 84:13-84:21 12:88-12:15 12 13:81-81:16:85 12 12:88-12:28 13:81-88:32 15 12:88-12:28 13:81-88:32 15 12:88-12:28 13:81-88:32 15 12:88-12:28 13:81-88:32 15 12:88-12:28 13:81-88:32 15 12:88-12:28 13:81-88:32 15 12:88-12:28 13:81-88:32 15 12:88-12:28 13:81-88:32 15 12:88-12:28 13:81-88:32 15 12:88-12:28 13:81-88:32 15 12:88-12:28 13:81-88:32 15 12:88-12:28 13:81-88:32 15 12:88-12:28 13:81-88:32 15 13:81-88:32 15 13:81-88:32 15 13:81-88:32 16 13:81-88:32 17 13:81-88:32 12 13:81-88:32	28882/ 12 28883/ 81 28883/ 81 28883/ 81 28883/ 81 28885/ 81 28885/ 81 28885/ 85 28885/ 87 28885/ 87 28885/ 87 28885/ 87 28885/ 87 28885/ 87 28885/ 87 28885/ 87 28885/ 87 28885/ 87 28885/ 87 28885/ 87 28885/ 87 28885/ 87 28885/ 87 28885/ 87 28885/ 87 28885/ 87 28885/ 88 28885/

KEY: INFORMATION SOURCES - PDE-NEIS PDECARD, MON-NEIS MONTHLY LIST, ICS-ICS LIST, HEL-HELICORDER, OTH-OTHER EVENT TYPE - EG-EARTHQUAKE, NX-NUCLEAR EXPLOSION, SX-SCIENTIFIC EXPLOSION, OT-OTHER, UN-UNKNOWN PHASES - A-P. B-PO, C-S, D-SO, E-T, F-OTHER

		, D-SO, E-T, F-OTHER		
EVNT *****ORIGIN TIME****** *NO* VR*MO*DA*JUL*HR*MN*SECS	**COORDINATES***	*********DESCRIPTION*******	DEP "MAGNI" INF KM" BDY"SRF SRC	EV *******SAVED****** **STRIP** TP PHASES**INTERVAL***MNS TAPE/FILE
### ### ### ### ### ### ### ### ### ##	39.278	TAJIK-SINKIANG BORDER REGION. PHILIPPINE IS. REGION. GILBERT IS. REGION. SOUTH OF FIJI IS. MARIANA IS. REGION. TAJIK SSR VOLCANO ISLANDS REGION. SOUTHERN NEVADA SOLOMON IS. SOLOMON IS. SOLOMON IS. SOLOMON IS. SOLOMON IS. TAJIK-SINKIANG BORDER REGION. SONOMON IS. SOLOMON	33 4.8 4.8 PDE 54 5.8 S.8 PDE 23 4.9 S.8 PDE 48 6.8 6.1 PDE 33 5.1 S.8 PDE 33 5.4 4.9 PDE 33 5.4 4.8 PDE 183 4.9 S.8 PDE 183 4.9 S.8 PDE 183 4.9 S.8 PDE 28 5.6 5.5 PDE 152 5.5 S.8 4.5 PDE 28 5.6 5.5 PDE 152 5.5 S.8 PDE 28 5.6 5.5 PDE 152 5.5 S.8 PDE 28 5.6 5.5 PDE 152 5.5 S.8 PDE 152 6.5 S.8 PDE 152 6.7 S.8 PDE 152 6.7 S.8 PDE 153 6.7 S.8 PDE 154 5.5 S.8 PDE 155 6.8 PDE 156 5.7 S.8 PDE 157 S.8 PDE 158 5.7 S.8 PDE 159 5.7 S.8 P	EQ A
#157 82 1# 1# 283 2#:36:28.6 #158 82 1# 11 284 #2:14:58.1 #159 82 1# 11 284 #2:5:35.2 #16# 82 1# 11 284 #2:5:35.2 #16# 82 1# 12 285 #2:49:38.9 #161 82 1# 12 285 #2:49:39:57.3 #163 82 1# 13 286 18:52:49.3 #163 82 1# 14 287 21:59:44.4 #166 82 1# 14 287 21:59:44.4 #166 82 1# 14 287 21:59:39.9 #169 82 1# 16 289 #2:53:39.9 #168 82 1# 16 289 #3:#2:59.3 #17# 82 1# 16 289 #3:#2:59.3 #17# 82 1# 16 289 #3:#2:59.3	44.1/1N 143.8821 77.2225 129.6218 77.3286 154.5321 62.5515 166.8666 27.4395 175.4410 37.2887 69.7844 21.1865 175.6370 14.156N 139.8371 38.449N 133.5811 5.849N 139.8371 38.449N 133.5811 5.849N 132.5811 6.748N 46.2471 64.727N 46.1621 64.727N 46.1621 64.727N 46.1621 64.727N 46.2381 64.748N 48.2381 64.748	OFFERST COST OF WORSDU, SAPA BANDA SEA NOVAVA ZEMLYA BALENNY IS. REGION KERMADEC IS. REGION KERMADEC IS. REGION TONGA IS. LUZON. PHILIPPINE IS. BONIN IS. REGION SEA OF JAPAN HINDANAO, PHILIPPINE IS. NORTH PACIFIC OCEAN. EAST PAPUA NEW GUINEA REGION NEAR S. COAST OF HONSHU, JAPAN SOUTHWESTERN USSR SOUTHMESTERN USSR SOUTHMESTERN USSR SOUTHMESTERN USSR SOUTHMESTERN USSR SOUTHMESTERN USSR SOUTHOE NORTH AFGHANISTAN-USSR BORDER REGION: KURIL IS. SOUTH OF MARIANA IS. TONGA IS. TONGA IS. TONGA IS. TONGA IS. TONGA IS. SOUTHOF MONSHU, JAPAN MINDANAO. PHILIPPINE IS. LIBBERT IS. REG. LEYTE. PHILIPPINE IS. HEAR W. COAST HONSHU, JAPAN HOLICCA SEA SOUTHMEST OF SUMATERA HINDINAO, PHILIPPINE IS. HINDINAO, PHILIPPINE IS.	28 3.8 3.8 PDE E 8 5.6 3.6 PDE N 18 5.8 4.9 PDE E 45 5.2 #.8 PDE E 33 5.1 #.8 PDE E 33 5.1 #.8 PDE E 455 5.1 #.8 PDE E 466 4.3 #.8 PDE E 416 4.3 #.8 PDE E 416 5.3 #.8 PDE E 5.3 8.8 PDE E 5.3 8.8 PDE E 8 5.2 3.8 PDE E 8 5.2 3.8 PDE M 8 5.2 3.8 PDE M 8 5.2 3.8 PDE M 8 5.3 3.1 PDE M 8 5.4 3.1 PDE M	Q ABDE

KEY: INFORMATION SOURCES - PDE-NEIS PDECARD, MON-NEIS MONTHLY LIST, ICS-ICS LIST, MEL-MELICORDER, OTH-OTHER EVENT TYPE - EQ-EARTHQUAKE, NX-NUCLEAR EXPLOSION, SX-SCIENTIFIC EXPLOSION, OT-OTHER, UN-UNKNOWN PHASES - A-P, B-FO, C-S, D-SO, E-T, F-OTHER

PHASES	- A-P. B-FO. C-S. D-SO. E-T. F-OTHER	
-NO- YR-MO-DA-JUL-HR-MN-SECS	:LATLON	DEP "MAGNI" INF EV """""""""""""""""""""""""""""""""""
### 192 18 27 300 18:38:13:9 ### 22:2 82 10 27 388 15:36:36.3 ### 213 82 10 27 388 15:36:36.3 ### 213 82 10 27 388 15:36:36.3 ### 213 82 10 27 388 17:49:15.8 ### 214 82 10 28 301 18:38:35.2 ### 215 82 18 28 301 18:38:35.2 ### 216 82 18 28 301 18:38:39.2 ### 217 82 18 29 302 81:42:15.5 ### 218 82 18 29 302 81:41:44.9 ### 218 82 18 29 302 81:41:44.4 ### 218 82 18 29 302 83:47:25.4 ### 228 82 18 29 302 83:47:25.4 ### 228 82 18 29 302 83:47:56.6 ### 221 82 18 31 384 82:48:26.6 ### 222 82 18 31 384 82:48:26.6 ### 222 82 18 31 384 82:48:26.8 ### 222 82 18 31 384 82:48:26.8 ### 223 82 18 31 384 82:48:26.8 ### 224 82 18 31 384 82:48:26.8 ### 225 82 10 31 384 82:48:26.8 ### 226 82 10 31 384 82:48:26.8 ### 227 82 18 31 384 83:38:41.2 ### 228 82 18 31 384 83:38:41.2 ### 228 82 18 31 384 83:38:41.2 ### 238 82 11 ## 385 88:47:55.2 ### 238 82 11 ## 385 88:47:55.2 ### 238 82 11 ## 236 87:48:27.7 ### 239 82 11 ## 236 87:48:27.7 ### 239 82 11 ## 236 87:48:27.7 ### 239 82 11 ## 236 87:48:27.7 ### 239 82 11 ## 236 87:48:27.7	22.484A 12: 596: PHILIPPINE IS. REGION 13.838N 195.897E YUNNAN PROVINCE, CHINA 13.838N 194.887E SEA OF JAPAN 7.5925 189.144E JAVA 46.44TN 144.867E SEA OF JAPAN 4.5695 189.144E JAVA 4.5695 152.462E NEW BRITAIN REGION 15.9985 173.768W TONGA ISLANDS 6.851S 132.426E BANDA SEA 8.114S 187.172E JAVA 12.178S 167.432E SANTA CRUZ ISLANDS 11.6325 177.722E JOUTH OF SUMBAWA ISLAND 11.6325 177.722E SOUTH OF SUMBAWA ISLAND 18.6625 189.888E VANJATU ISLANDS 6.6425 138.5988 2ANDA SEA 2.944N 96.126E NORTHERN SUMATERA 135.921N 82.518E TIBET 9.672N 126.895E MINDANAO, PHILIPPINE IS. 6.8675 185.548E SUNCA STRAIT 31.397N 141.755E SOUTH OF FIJI IS. 12.397N 125.668E SAMAR, PHILIPPINE IS. 55.494S 124.485W EASTER IS. CORDILLERA 3.2345 139.692E WEST JRIAN 25.167S 179.724E SOUTH OF FIJI IS.	18 4.8 5.3 PDE EQ A 18 8.8 PBE EQ A 22-35-22:46 12 298:1/ 12 18 4.8 5.3 PDE EQ A 22-35-22:46 12 298:1/ 14 18 4.8 5.3 PDE EQ A 12:33-12:44 12 298:1/ 14 18 8.8 P.8 HEL _Q ABDE 12:31-12:49 12 298:1/ 15 18 4.6 5.8 PDE EQ ABDE 13 4.8 8.8 PDE EQ ABDE 13 5.6 8.8 PDE EQ A 11 5.4 5.2 PDE EQ A 13 5.2 5.1 PDE EQ A 23 5.2 5.1 PDE EQ A 23 5.2 5.1 PDE EQ A 25 14.8 8.8 PDE EQ A 25 14.8 8.8 PDE EQ A 26 4.4 8.8 PDE EQ A 27 5.5 1.8 PDE EQ A 28 4.4 4.8 F.8 PDE EQ A 28 4.4 4.8 F.8 PDE EQ A 28 5.1 8.8 PDE EQ A 28 6.8 PDE EQ A 29 6.9 PDE EQ A 20 6.9 PDE EQ A
### ### ### ### ### ### ### ### ### ##	\$5.848H 165.687E KOMANDORSKY IS. REGION 3.6165 177.768E GILBERT IS. REGION 4.827H 127.835E TALAUD IS. 21.383S 178.734W FIJI IS. REGION 7.236N 94.437E HICOBAR IS. REGION 3.447S 177.728E GILBERT IS. REGION 3.447S 177.728E GILBERT IS. REGION 3.447S 177.728E GILBERT IS. REGION 41.171H 149.532E KURIL IS. 33.983H 137.878E MEAR S. COAST OF MONSHU JAPAN 15.454S 176.822W FIJI IS. REGION 6.677S 181.671E SOUTHWEST OF SUMATERA 44.237H 149.515E KURIL IS. 44.138H 149.488E KURIL IS. 44.138H 149.488E KURIL IS. 44.138H 149.595I KURIL IS. 6.198S 146.369E EAST PAPUA NEW GUINEA REG. 27.431S 176.735W KERMADEC IS REGION 1.478H 126.459E MOLUCCA PASSAGE 7.824H 116.832W SOUTHERN NEVADA 43.753H 158.875E FURIL IS. REGION 1.7879S 176.386W FIJI IS. REGION 17.879S 176.386W FIJI IS. REGION 17.879S 176.386W FIJI IS. REGION 13.385S 146.835E SISMARCK SEA 18.558S 167.241E VANUATU ISLANDS 6.872S 185.355S SUNCA STRAIT 44.944H 148.924E KURIL ISLANDS 6.872S 185.355S SUNCA STRAIT 44.934H 129.498W OFF COAST OF OREGON 11.978H 143.812E SOUTH OF MARIAMA IS. 22.499H 143.7488 SOUTH OF MARIAMA IS. 52.977H 158.697E HEAR EAST COAST OF KAMCHATKA 9.419S 122.584E SOUT MO F MARIAMA IS. 52.979H 158.697E HEAR EAST COAST OF KAMCHATKA 9.419S 122.584E SOUTH OF MARIAMA IS. 52.978H 124.985E NORTHEAST OF TAIWAN 124.985E 108.886F SOUTH OF MARIAMA IS. 53.887S 168.886F SOUTH MATLANTIC RIDGE 28.668H 138.878E NORTHEAST OF TAIWAN 28.1989H 29.9895W CENTRAL MID-ATLANTIC RIDGE 28.668H 138.878E NORTHEAST OF TAIWAN 14.628S 168.886F SOUTHWESTERN RYUNVU ISLANDS 5.767S 152.367L NEW BRITAIN REGION 1.2789 147.167E EAST PAPUA NEW GUINEA REGION 1.2789 74.753W PERU 1.283H 121.902E MINAMASSSA PENINSULA 1.283H 121.902E MINAMASSSA PENINSULA 1.283H 121.902E MINAMASSSA PENINSULA 1.5894S 147.5752 REGION JAPAN REGION 1.57652 REGION JAPAN REGION 1.57	33 5.2 8.8 PDE E2 ABDE 15:42-17:32 51 28213/ 17 33 5.8 8.8 PDE E3 ABDE 1:489-17:27 39 28213/ 17 161 5.4 8.8 PDE E3 ABDE 1:38-18:49 12 28213/ 18 556 5.2 8.8 PDE E3 A 87:49-81:88 12 28213/ 18 33 5.2 8.8 PDE E3 ABDE 8:125-18:84 48 28213/ 21 33 5.1 8.8 PDE E3 ABDE 8:125-18:84 48 28213/ 21 33 5.1 8.8 PDE E3 ABDE 8:137-81:28 48 28213/ 21 33 5.1 8.8 PDE E3 ABDE 8:137-81:28 44 28214/ 81

KEY: INFORMATION SOURCES - PDE-NEIS PDECARD, MON-NEIS MONTHLY LIST, ICS-ICS LIST HEL-HELI.ORDEP, OTH-OTHER EVENT TYPE - EQ-CARTHQUAKE, NX-NUCLEAR EXPLOSION, SX-SCIENTIFIC EX-LOSION, O'-OTHER, UN-UNKNOWN PHASES - A-P, B-PO, C-S, D-SO, E-T, F-OTHER

EVAT ******ORIGIN TIME************************************	**COORDINATES***	**************************************	DEP *MA 41* KM- BDY-SRF	INF EV *****	STRIP.
### 48 2 11 22 326 #1:97:59.# ### 81#5 62 11 22 326 #1:97:59.# ### 81#5 62 11 22 326 #1:97:59.# ### 81#6 82 11 22 326 #5:32:51.2 ### 81#6 82 11 22 326 #5:32:51.2 ### 81#8 82 11 23 327 ### 81#8 ### 82 11 23 327 ### 81#8 ### 82 11 23 327 ### 81#8 ### 82 11 23 327 ### 81#8 ### 82 11 25 329 ### 81#5 ## 82 11 25 329 ### 81#5 ## 82 11 25 329 ### 81#5 ## 82 11 25 329 ### 81#5 ## 82 11 26 33# ### 81#5 ## 82 11 26 33# ### 81#8 ## 82 11 26 33# ### 81#8 ## 82 11 27 33# ## 81#8 ## 82 12 11 28 332 11 18 ## 86 9 ## 82 11 27 33# ## 81#8 ## 82 11 27 33# ## 81#8 ## 82 11 27 33# ## 81#8 ## 82 11 27 33# ## 81#8 ## 82 11 28 332 11 18 ## 86 9 ## 82 11 27 33# ## 81#8 ## 82 12 11 38 33# ## 81#8 ## 82 12 11 38 33# ## 81#8 ## 82 12 12 12 33 34 14 18 3 18 18 18 18 18 18 18 18 18 18 18 18 18	39,712N 77,718E 55.667H 163.221E 23.7485 176.864W 32.2565 179.273W 32.2635 178.484W 3.7445 187.788E 23.9555 175.478W 3.3875 177.639L 36.732N 71.474E 5.826N 122.63E 11.857H 142.647E 5.826N 125.812E 52.8845 144.228W 35.7645 144.389W 55.7645 144.389W 56.5876 174H 147.79E 23.5445 175.286W 39.818N 148.39E 23.5445 175.286W 39.818N 148.39E 24.4365 178.853W 39.818N 148.39E 24.4365 178.853W 39.8648 178.	CENTRAL MID-ATLANTIC RIDGE FOX IS ALEUTIAN IS. OFF EAST COAST OF HONSHU JAPAN	32 4 - 1 - 2 - 2 - 2 - 2 - 2 - 2 - 2 - 2 - 2	PDE E A A BDE ECA BDE ECA A BDE ECA	##:15-##:166 ## 2 ## 2 ## 2 ## 2 ## 2 ## 2 ## 2 #
#351 82 12 #6 34# 14:#31:31.6 #352 82 12 #7 341 #5:43:49.6 #353 82 12 #7 341 #5:43:49.6 #353 82 12 #8 342 #2:41:34.1 #354 82 12 #8 342 #2:41:34.1 #355 82 12 #8 342 #2:41:34.1 #355 82 12 #8 342 #3:54:12.6 #256 82 12 #8 342 18:54:12.6 #257 82 12 #8 342 18:54:12.6 #358 82 12 #9 343 #1:41:37.6 #358 82 12 #9 343 #1:41:37.6 #358 82 12 #9 343 #1:41:37.6 #356 82 12 #9 343 #1:52:#39:48.3 #361 82 12 #9 343 #1:52:#39:48.3 #361 82 12 #9 343 18:39:48.3 #365 82 12 #9 343 18:39:48.3 #366 82 12 1# 344 15:22:#39:48.3 #367 82 12 1# 344 15:22:#39.9 #365 #8 #8 #8 #8 #8 #8 #8 #8 #8 #8 #8 #8 #8	36. 321N 141.262E 36. #25N 114.252E 36. #25N 114.252E 36. #25N 114.6 #25E 14.235N 145.126E 31.4725 177.577E #.632N 119.918E 41.3645 #7.626V 29. #295 112.655W 31.3625 168.527E 47.833N 155.246E 47.833N 155.246E 47.833N 155.246E 47.833N 156.872W #.#828 #.#828E #.#828 #.#828E 11.3775 119.813E 17.3775 119.813E 17.812N 145.616E 63.2755 61.237W 14.756N 44.2518E 11.3775 178.624E 28.5355 169.887E 28.3755 61.237W 14.756N 44.2518E 28.5355 169.887E 28.6355 169.887E 28.6525 177.8888E 28.6765 178.887E 28.1575 178.321W 36.242M 69.896E 36.356N 69.595E 11.689N 92.993E 13.5825 177.716E 26.6715 176.277W 13.5825 177.716E 26.6715 176.277W 17.528 125.634E 25.6755 176.574W 6.8375 178.457W 32.4785 178.3784E 32.4785 178.457W 32.4785 178.478W 32.47885 178.478W 32.47885	NEAR EAST COAST OF MONSHU, JAP SOUTHERN NEVADA KUPIL ISLANDS MARIAMA ISLANDS GILBERT ISLANDS REGION MINAMASSA PENINSULA WEST CHILE RISE EASTER ISLAND REGION LOVALTY ISLANDS REGION LOVALTY ISLANDS REGION LOVALTY ISLANDS REGION KURIL ISLANDS REGION SOUTHERN NEVADA SCUTH OF FIJI ISLANDS PROBABLE MARIANAS PO SO T SOUTH OF FIJI ISLANDS PROBABLE MARIANAS PO SO T SOUTH OF FIJI ISLANDS SOUTH SHETLAND ISLANDS WESTERN ARABIAN FENINSULA FIJI IS. REGION VANUATU IS TOWAS IS. REGION VANUATU IS. FIJI IS. REGION MINDU KUSH REGION HINDU KUSH REGION HINDU KUSH REGION MINDU KUSH REGION MOLUCCA PASSAGE HINDU KUSH REGION GILBERT IS. REGION SOUTH OF FIJI ISLANDS SOUTH OF FIGURALERA SOUTH OF FIGURALERA SOUTH OF FIGURALERA SOUTH OF TONGA ISLANDS SOUTH OF TONGA ISLANDS COUTH OF		PODE ED ON ARA ARDE DE LOCALE ED ON ARA ARA DE LOCALE ED ON ARA ARA DE LOCALE ED ON ARA ARA ARA ARA ARA ARA ARA ARA ARA AR	14-84-14-10 16 28810/ 28

PHASES	- A-P, B-PO, C-S,	D=SO, E=T, F=OTHER		
"NO" YR"MO"DA"JUL "HR"MN"SECS	**LAT****LON***	LOCATION	KM* BDY*SRF SRC TP PHASE	SP":NTERVAL ***MNS TAPE/FILE
## ## ## ## ## ## ## ## ## ## ## ## ##	16.1725 172.589w 24.2395 175.787w 29.221N 81.333E 24.7445 175.776w 23.6287 175.822w 23.9995 175.899w 37.172N 71.743E 23.2515 179.895E 24.4795 176.218W 23.8125 175.849w 24.5875 176.898w 24.1885 175.816w 23.7785 175.826w 23.7887 175.826w 24.4655 175.1738w 24.4655 175.1738w 24.4655 175.1738w 24.3757 177.8921 25.251N 173.331E 25.252N 151.2288 25.252N 151.2288 25.252N 151.2288 25.3955 175.174w 3.6375 177.8921 36.8787 79.8899 41.811N 61.693E 45.252N 151.2886 3.3885 177.7961 19.886N 145.826E 3.3885 177.7961 19.886N 145.826E 3.3885 177.7961 19.886N 145.826E 3.3882N 139.2451 33.822N 139.2451 33.822N 139.2451 33.822N 139.2451 33.822N 139.2451	NEAR ISLANDS. ALEUTIAN ISLANDS POSSIBLY NEAR WAKE ISLAND WEST IRIAN FLOTES ISLAND REGION BUP-HA-INDIA BORDER REGION KUPIL ISLANDS TONGA ISLANDS REGION GLIBERT ISLANDS REGION EASTERN KAZAKH SSR. UZSEK SSR PAPUA NEW GUINEA NEAR EAST COAST OF HONSHU JAPA MARIANA ISLANDS REGION GLIBERT ISLANDS REGION SOLOMON ISLANDS PROBABLE MARIANAS PO.SO.T NEAR EAST. COAST OF HONSHU JAPAN SOUTH OF KOMSHU JAPAN REAR S. COAST OF HONSHU JAPAN SOUTH OF KOMSHU JAPAN REAR S. COAST OF HONSHU JAPAN SOUTH OF HONSHU JAPAN SOUTH OF HONSHU JAPAN SOUTH OF HONSHU JAPAN	33 5.2 8.8 PDE EO A 33 4.7 8.8 PDE EO A 33 4.7 8.8 PDE EO A 33 5.3 5.7 PDE EO A 33 5.3 5.7 PDE EC A 33 5.1 8.8 PDE EC A 573 5.5 P.8 PDE EC A 573 5.5 P.8 PDE EC A 33 5.4 5.4 PDE EC A 33 5.3 5.1 PDE EC A 33 5.3 5.1 PDE EC A 33 5.2 5.1 PDE EC A 56 5.2 R.8 PDE EC A 56 5.2 R.8 PDE EC A 56 5.2 R.8 PDE EC A 57 5.2 R.8 PDE EC A 33 5.8 R.8 PDE EC A 33 5.8 R.8 PDE EC A 33 5.8 R.8 PDE EC A 33 4.9 R.8 PDE EC ABDE 34 5.7 R.8 PDE EC ABDE 35 5.8 R.8 PDE EC ABDE 36 5.8 R.8 PDE EC ABDE 37 5.1 R.8 PDE EC ABDE 38 5.8 R.8 PDE EC ABDE	#6:13-R6:24 12 2022/1 14 12:16-12:27 12 2022/1 14 12:16-12:27 12 2022/1 15 10:16-12:27 12 2022/1 15 10:16-12:27 12 2022/1 15 10:16-12:27 12 2022/1 15 10:16-12:27 12 2022/1 15 10:16-12:27 12 2022/1 17 16:47-8:56 12 2022/1 17 16:47-8:56 12 2022/1 17 16:47-8:56 12 2022/1 18 10:16-12:26 13 2022/1 18 10:17-72:29 13 2022/1 18 10:17-72:29 13 2022/1 18 10:17-72:29 13 2022/1 18 10:17-72:29 13 2022/1 18 10:17-72:29 13 2022/1 18 10:17-72:29 13 2022/1 18 10:17-72:29 13 2022/1 18 10:17-12:29 12 2022/1 18 10:17-12:20 12 2022/1 18 10:18-12:26 18 1222/1 18 10:18-12:26 18 1222/1 18 10:18-12:26 18 1222/1 18 10:19-12:26 18 1222/1 18 10:19-12:26 18 1222/1 18 10:19-12:26 18 1222/1 18 10:19-12:26 18 1222/1 18 10:19-12:26 18 1222/1 18 10:19-12:26 18 1222/1 18 10:19-12:26 18 1222/1 18 10:19-12:26 18 1222/1 18 10:19-12:26 18 1222/1 18 10:19-12:26 18 1222/1 18 10:19-12:26 18 1222/1 18 10:19-12:26 18 1222/1 18 10:19-12:26 18 1222/1 18 10:19-12:26 18 1222/1 18 10:19-12:26 18 1222/1 18 10:19-12:26 18 12 1222/1 18 10:19-12:26 18 12 1222/1 18 10:19-12:26 18 12 1222/1 18 10:19-12:26 18 12 1222/1 18 10:19-12:26 18 12 1222/1 18 10:19-12:26 18 12 1222/1 18 10:19-12:26 18 12 1222/1 18 10:19-12:26 18 12 1222/1 18 10:19-12:26 18 12 1222/1 18 10:19-12:26 18 12 1222/1 18 10:19-12:26 18 12 1222/1 18 10:19-12:26 18 12 1222/1 18 10:19-12:26 18 12 1222/1 18 10:19-12:26 18 12 1222/1 18 10:19-12:26 18 12 1222/1 18 10:19-12:26 18 12 1222/1 18 10:19-12:26 18 12 1222/1 18 10:19-12:26 18 18 1222/1 18 10:19-12:26 18 18 18 18 18 18 18 18 18 18 18 18 18
## 12	33.578N 139.546E 32.355N 127.86EE 22.355N 127.86EE 25.577N 145.7821 19.958N 121.423E 23.685S 175.7824 33.635N 175.7824 33.635N 175.7824 33.635N 177.713E 23.384S 175.6354 18.194S 178.56354 18.194S 178.56354 17.717S 139.296E 17.717S 139.296E 14.866N 128.323E 33.782N 139.295E 14.866N 128.323E 33.782N 139.295E 24.154S 174.255E 24.154S 174.255E 24.154S 174.25E 24.267N 142.25E	SOUTH OF MONSHU JAPAN SOUTH OF MONSHU JAPAN SOUTH OF MONSHU JAPAN BURMA-CHINA BORDER REGION YUNNAN PROVINCE, CHINA MARIANA ISLANDS REGION TONGA ISLANDS REGION TIBET-INDIA BORDER REGION FIJI ISLANDS REGION SOUTH OF MONSHU JAPAN SOUTH OF MONSHU JAPAN SOUTH OF MONSHU JAPAN SOUTH OF MONSHU JAPAN EVER EAST COAST OF MONSHU JAPA PAKISTAN SOUTH OF MONSHU JAPAN SOUTH OF TONGA TSLANDS REAR EAST COAST OF MONSHU JAPA PAKISTAN SCUTH OF TONGA TSLANDS GLUBERT ISLANDS SCUTH OF TONGA TSLANDS GLUBERT ISLANDS GLUBERT ISLANDS GLUBERT ISLANDS REGION MINDANAO PHILIPPINE ISLANDS PROEABLE MARIANAS PO.SO.T PERUBOLIVIA BORDER REGION MINDANAO PHILIPPINE ISLANDS SOUTHERN ALASIA SANTA CRUZ ISLANDS SOUTHERN ALASIA SOUTHERN SUMATERA MOLUCCA PASSAGE FIJI ISLANDS REGION MINDANAO PHILIPPINE ISLANDS SOUTHERN SUMATERA SOUTHERN SUMATERA OUTHERN SUMATERA MOLUCCA PASSAGE FIJI ISLANDS REGION MINDANAO PHILIPPINE ISLANDS SOUTHERN SUMATERA MOLUCCA PASSAGE FIJI ISLANDS REGION MINDANAO PHILIPPINE ISLANDS SOUTHERN SUMATERA MOLUCCA PASSAGE FIJI ISLANDS REGION MINDANAO PHILIPPINE ISLANDS SOUTHERN SUMATERA MOLUCCA PASSAGE FIJI SLANDS REGION MINDANAO PHILIPPINE ISLANDS SOUTHERN SUMATERA MOLUCCA PASSAGE FIJI SLANDS REGION MINDANAO PHILIPPINE ISLANDS SOUTHERN SUMATERA MOLUCCA PASSAGE FIJI SLANDS REGION MINDANAO PHILIPPINE ISLANDS SOUTHERN SUMATERA MOLUCCA PRILIPPINE ISLANDS NEGGOS PHILIPPINE ISLANDS NEGGOS PHILIPPINE ISLANDS NEGGOS PHILIPPINE ISLANDS PROCABLE BOWIN ISL PO.SO.T	15 5.2 8.8 PDE EQ ABDE 28 5.9 6.1 PDE EQ ABDE 26 5.2 8.8 PDE EQ A 33 5.5 8.8 PDE EQ A 33 5.9 5.9 PDE EQ A 33 5.3 5.6 PDE EQ A 33 5.3 5.7 PDE EQ ABDE 33 5.1 8.8 PDE EQ ABDE 33 5.1 8.8 PDE EQ ABDE 33 5.8 8.8 PDE EQ A 419 5.1 8.8 PDE EQ A 419 5.1 8.8 PDE EQ A 28 5.5 5.3 PDE EQ ABDE 33 5.8 8.8 PDE EQ ABDE 33 5.8 8.8 PDE EQ ABDE 33 5.8 8.8 PDE EQ ABDE	15:15-15:57 42 282247 12 15:29-15:48 12 282247 12 15:59-17:41 43 288247 13

KEY: INFORMATION SOURCES - PDE-NEIS PDECARD, MON-NEIS MONTHLY LIST, ICS-ICS LIST, HEL-HELICORDER, OTH-OTHER EVENT TYPE - EO-EARTHQUAKE, NX-NUCLEAR EXPLOSION, SX-SCIENTIFIC EXPLOSION, OT-OTHER, UN-UNKNOWN PHASES - A-P. B-PO, C-S, D-SO, E-T, F-OTHER

PHASES	- A-P. B-PO. C-3	, D-SO. E-T. F-OTHER	THE CANADALON OF COMM.	
2332*NM*RH*JUL*AC*OM*RY *OM*	**LAT****LON***	•=====DESCRIPTION	KM" BDY"SRF SRC TP PHASE	S**INTERVAL***MNS TAPE/FILE
## 1950	26.86.35 176.6532 3.1685 177.6756 17.3585 178.117 36.423N 7.86516 3.826N 7.86516 3.1645 146.2555 55.265N 163.1262 55.265N 163.1262 55.265N 163.1262 55.285N 163.2951 35.92N 142.6876 35.92N 142.6876 35.92N 143.6876 35.92N 143.6876 35.92N 143.3977 37.4818 133.7481 9.628N 133.7481 9.628N 122.3386 3.4175 177.5976 27.2995 63.3967 3.4175 177.5976 27.2995 63.3967 3.4175 177.5976 3.455N 177.5976 3.455N 122.3386 3.4175 177.5976 3.455N 122.3386 3.4175 177.5976 56.988N 122.3386 54.988N 122.3386 54.958N 148.1566 54.988N 148.1566 54.988N 148.1567 66.988N 148	SOUTH OF FIJI ISLANDS GILBERT ISLANDS REGION FIJI ISLANDS REGION FIJI ISLANDS REGION FROBABLE BONN' ISLANDS TONGA ISLANDS BISMARCK SEA OFF EAST COAST OF KAMCHATKA HOKKAIDO JAPAN REGION NEAR S. COAST OF KAMCHATKA HOKKAIDO JAPAN REGION NEAR S. COAST OF KAMCHATKA WEST IRIAN REGION HEROS PHILIPPINE ISLANDS SANTIAGO DEL ESTERO PROV. ARGE GILBERT ISLANDS REGION BURMA-INDIA BORDER REGION BURMA-INDIA BORDER REGION BURMA-INDIA BORDER REGION BURMA-INDIA BORDER REGION HEW BRITAIN REGION KEAR EAST COAST OF KAMCHATKA SOUTH OF TONGA ISLANDS BURMA SOUTH OF TONGA ISLANDS WENTAL TONGA TO KAMCHATKA SOUTH OF TONGA ISLANDS BURMA HOW BOLL TO COAST OF KAMCHATKA SOUTH FOR TONGA TO KAMCHATKA SOUTH FOR TONGA TO SOUTH. HONSHU OFF EAST COAST OF KAMCHATKA TONGA PHILIPPINE ISLANDS KODIAK ISLANDS KODIAK ISLANDS KODIAK ISLANDS REGION EAST PAPUA NEW GUINEA REGION	33 5.3 5.3 PDE EO A D 33 4.9 8.8 PDE EO A BDE 33 6.1 6.3 PDE EO A BDE 33 6.1 6.3 PDE EO A BDE 33 5.8 8.8 PMLL EO BDE 33 5.8 P.8 PDE EO A BDE 261 4.7 8.8 PDE EO A BDE 261 4.7 8.8 PDE EO A BDE 33 5.2 5.1 PDE EO A BDE 33 5.3 5.7 PDE EO A BDE 33 5.4 6.8 PDE EO A BDE 33 5.3 5.7 PDE EO A BDE 33 4.9 8.8 PDE EO A BDE 33 5.5 5.9 PDE EO A BDE 33 5.5 6.8 PDE EO A BDE 33 5.7 S.8 PDE EO A BDE	19:59-28:18 12 28226/84 28:36-28:47 12 28226/85 89:28-89:51 12 28226/86 18:28-19:83 12 28226/86 18:28-19:83 12 28226/86 18:28-19:83 12 28226/86 18:35-86:28 16 28226/86 18:35-86:28 16 28226/86 18:35-86:28 16 28226/86 11:88-11:37 32 28226/11 11:88-11:37 32 28226/11 11:88-11:37 36 28226/11 11:24-12:18 55 22/26/11 11:24-12:18 55 22/26/11 11:24-12:18 55 22/26/11 11:24-12:18 55 22/26/11 11:25-13:31 47 28227/81 18:38-19:21 44 28227/82 12:86-21:55 53 28227/83 18:38-19:21 44 28227/82 12:86-21:55 53 28227/83 82:88-21:55 53 28227/83 82:88-21:56 53 28227/83 83:34-32-81:49 12 28227/83 83:34-32-81:49 12 28227/18 18:38-19:39 13 28237/18 84:32-81:49 12 28227/18 84:32-81:49 18 28227/18 84:22-81:49 18 28227/18 18:24-12:39 12 28227/18 18:24-12:39 12 28227/18 18:24-12:39 12 28227/18 18:24-12:39 12 28227/18 18:24-12:39 12 28227/18 18:24-12:39 12 28227/18 18:24-12:39 12 28227/18 18:24-12:39 12 28227/18 18:24-12:38 12 28227/18 18:24-12:38 12 28227/18 18:24-12:28 12 28227/18 18:24-12:28 12 28227/18 18:24-12:28 12 28227/18 18:24-12:28 12 28227/18 18:24-12:28 12 28227/18 18:24-12:28 12 28228/81 18:24-12:28 12 28228/81 18:24-12:28 12 28228/81 18:24-12:28 12 28228/81 18:24-12:28 12 28228/81 13:13-13:24 12 28228/81 13:13-13:24 12 28228/81 13:13-13:24 12 28228/81 13:13-13:24 12 28228/81 11:55-12:89 13 28228/11
#551 83 #1 10 #18 22:51:56.3 #552 83 #1 19 #19 11:5#:1#.3 #553 83 #1 19 #19 12:49:39.1 #5554 83 #1 19 #19 12:49:39.1 #5555 83 #1 19 #19 12:29:39.1 #5555 83 #1 19 #19 22:22:#6.9 #5556 83 #1 2# #2# 12# #22 12:47:6.9 #5556 83 #1 2# #2# 12# #22 12:47:6.9 #555 83 #1 2# #2# 12# #23:47:6.9 #5559 83 #1 2# #2# 12# #55:47:42.8 #568 #3 #1 2# #21 #5:#4:42.8 #568 #3 #1 2# #21 #5:#4:42.8 #568 #3 #1 2# #21 #5:#4:42.8 #568 #3 #1 2# #21 #5:#4:42.8 #568 #3 #1 2# #21 #5:#4:42.8 #568 #3 #1 2# #22 #1:24:33.9 #564 #3 #1 2 #22 #1:27:43.4 #5656 #3 #1 22 #22 #1:27:43.4 #5656 #3 #1 22 #22 #1:27:43.4 #5656 #3 #1 22 #22 #1:27:43.4 #5656 #3 #1 22 #22 #1:27:43.4 #5656 #3 #1 22 #22 #1:27:43.4 #5656 #3 #1 22 #22 #1:27:43.4 #5656 #3 #1 22 #22 #1:27:43.4 #5656 #3 #1 22 #22 #1:27:43.4 #5656 #3 #1 22 #22 #1:27:43.4 #5656 #3 #1 22 #22 #1:27:43.4 #5656 #3 #1 22 #22 #1:27:43.4 #5656 #3 #1 22 #22 #1:27:43.4 #5656 #3 #1 22 #22 #1:27:43.4 #5656 #3 #1 22 #22 #1:27:43.4 #5656 #3 #1 24 #24 #1:31.36.9 #577# #3 #1 24 #24 #8:17:36.6 #574 #8 ## ## ## ## ## ## ## ## ## ## ## ##	36.368N 78.446E 18.618S 166.142E 25.259N 91.853E 25.259N 168.762E 25.741S 154.379E 7.741S 154.379E 7.7516S 128.4781 28.751N 142.379E 42.462N 124.373E 42.631 124.559E 42.462N 143.474E 3.556S 177.662E 6.654S 141.274E 3.556S 177.662E 6.753S 183.861E 26.299S 178.743E 6.753S 183.861E 26.299S 178.743E 6.753S 183.861E 26.299S 178.743E 143.8781 128.3781 128.3781 128.3781 128.3781 128.3781 128.3781 128.3781 128.3781 128.3781 128.3781 128.3781 128.3781 128.3781 128.3781 128.3781 128.3781 128.3781 128.3781 148.3781 148.3781 157.2581 157.3781 148.3781 157.2581 157.3781 148.3781 157.2581 157.3781 148.3781 157.2581 157.3781 148.3781 157.2581 157.3781 157.3781 157.3781 148.3781 157.2581 157.3781 148.3781 157.3781 157.3781 148.3781 157.3781 137.4781 148.3774 1488 1488 1488 1488 1488 1488 1488 1	MINDU KUSH REGION SANTA CRUZ ISLANDS INDIA-BANGLADESH BORDER REGION KAMCHATKA SOLOMON ISLANDS BANDA SEA EONIN ISLANDS REGION ANDREANOT ISL., ALEUTIAN ISL. CELEBES SEA KOKKAIDO JAPAN REGION PAPUA NEW GUINEA GILBERT ISLANDS REGION SOUTHWEST OF SUMATERA SOUTHERN FACIFIC OCEAN SOUTHERN PACIFIC OCEAN SOUTHERN PACIFIC OCEAN OANACA MEXICO POSSIBLE NUCLEAR TEST RAT ISL., ALEUTIAN ISL. WIAR EAST COAST OF KAMCHATKA RAT ISL., ALEUTIAN ISL. SOUTH OF HONSHU JAPAN ANCAMAN ISLANDS WEST IRIAN FOR HONSHU JAPAN FOR HONSHU JAPAN SOLOMON ISLANDS WEST IRIAN FOR HONSHU JAPAN FOR HONSHU FOR HONSHU JAPAN FOR HONSHU J	5.8 8.8 PDL EQ ABDE 15.3 5.8 8.8 PDE EQ ABDE 15.3 5.8 8.8 PDE EQ ABDE 69 4.5 8.8 PDE EQ ABDE 33 5.4 8.2 PDE EQ ABDE 33 5.4 5.2 PDE EQ ABDE 33 5.4 5.2 PDE EQ ABDE 33 5.4 5.5 PDE EQ ABDE 33 5.8 8.8 PDE EQ ABDE 33 5.8 8.8 PDE EQ ABDE 162 4.8 8.8 PDE EQ ABDE 162 4.8 8.8 PDE EQ ABDE 163 5.6 8.8 PDE EQ ABDE 18 5.6 8.8 PDE EQ ABDE 18 5.6 8.8 PDE EQ ABDE 33 5.4 5.7 PDE EQ ABDE 33 5.4 5.7 PDE EQ ABDE 33 4.7 8.8 PDE EQ ABDE 33 4.7 8.8 PDE EQ ABDE 33 4.9 8.8 PDE EQ ABDE 33 4.9 8.8 PDE EQ ABDE 33 4.9 8.8 PDE EQ ABDE 33 5.1 8.8 PDE EQ ABDE	22:59-23:18

KEY: INFORMATION SOURCES - PDE-MEIS PDECARD, MON-MEIS MONTHLY LIST, ICS-ICS LIST, MEL-MELICORDER, OTH-OTHER EVENT TYPE - EG-EARTHQUAKE, NX-MUCLEAR EXPLOSION, SX-SCIENTIFIC EXPLOSION, OT-OTHER, UM-UNKNOWN PMASES - A-P, 8-PO, C-S, D-SO, E-T, F-OTHER

EVAT ******ORIGIN TIME******	**COORDINATES***		DEP "MAGNI" INF EV **** KM" BDY"SRF SRC TP PHAS	****SAVED******* **STRIP** ES**INTERVAL***MNS TAPE/FILE
### ### ### ### ### ### ### ### ### ##	18.7485 165.813E 20.735M 139.318E 21.87385 177.869E 3.8385 177.869E 36.585M 141.488E 48.214M 146.177E 48.214M 146.177E 17.5155 179.886E 5.428M 94.942E 18.3585 161.231E 33.369M 148.796E 7.3975 128.7361 3.369M 148.796E 3.4865 161.231E 37.156M 135.888E 27.832M 92.878E 4.3965 153.888E 27.832M 92.878E 4.3965 152.848E 27.832M 92.878E 4.513M 127.259E 4.3965 153.889E 6.525 144.482E 27.832M 92.878E 1.513M 127.259E 4.3965 177.678E 3.3755 144.862E 27.832M 92.878E 4.3965 157.67E 3.3755 144.862E 27.832M 92.878E 4.5966 153.899E 6.525 164.866E 15.3965 177.678E 8.3978 177.678E 8.3978 177.8661 8.2915 161.8661 26.428M 155.627C 7.841S 129.885E 8.3978 161.8661 26.428M 155.627C 7.841S 129.885E 8.3978 161.8661 26.428M 155.627C 7.841S 129.885E 8.382M 155.233W 18.55.627C 18.558 167.667E 29.7885 177.8661 19.358M 155.233W 18.55.627C 19.538M 155.627C 19.548M 126.488E 19.569M 128.488E 19.569M 128.488E 19.569M 128.488E 19.569M 128.488E	SANTA CRUZ ISLANDS BONIN ISLANDS REGION GILBERT ISLANDS REGION FIDI ISLANDS NEAR E. COAST OF MONSHU JAPAN SEA OF ONMOTSK SOLOMON ISLANDS SOUTH OF MONSHU JAPAN BANDA SEA GILEERT ISLANDS REGION PRODABLE MARIANAS PO, SO, T WEST ISLANDS REGION NEAR N COAST OF PAPUA NEW GUIN INDIA-CHINA BORDER REGION HALMAHERA NEW BRITAIN REGION TIMOR NEW IRELAND REGION SOLOMON ISLANDS GILBERT ISLANDS REGION GILBERT ISLANDS KURLI ISLANDS SOUTH OF FIJI ISLANDS KURLI SILANDS BANDA SEA POSSIBLE NUCLEAR TEST KERMADEC ISLANDS NEAF S. COAST OF S. HONSHU RYUKYU JILANDS MAUAII OFF EAST COAST OF S. HONSHU RYUKYU JILANDS MAUAII OFF EAST COAST OF S. HONSHU RYUKYU JILANDS MAUAII OFF EAST COAST OF SUMATERA FIJI ISLANDS REGION PHILIPPINE ISLANDS REGION PHOLIPPINE ISLANDS REGION	79 5.6 8.8 PDE EQ ABDE 443 5.1 8.8 PDE EQ ABDE 33 4.9 8.8 PDE EQ ABDE 634 4.8 8.8 PDE EQ ABDE 634 4.9 8.8 PDE EQ ABDE 479 4.9 8.8 PDE EQ ABDE 479 4.9 8.8 PDE EQ ABDE 638 4.8 8.8 PDE EQ ABDE 638 4.8 8.8 PDE EQ ABDE 638 4.8 8.8 PDE EQ ABDE 638 5.3 9.8 PDE EQ ABDE 638 5.3 9.8 PDE EQ ABDE 63 5.3 8.8 PDE EQ ABDE 63 5.3 8.8 PDE EQ ABDE 63 5.1 8.2 PDE EQ ABDE 63 5.1 8.2 PDE EQ ABDE 63 5.3 8.8 PDE EQ ABDE 63 5.1 8.2 PDE EQ ABDE 63 5.1 8.8 PDE EQ ABDE 63 5.8 8.8 PDE EQ ABDE 64 5.8 8.8 PDE EQ ABDE 65 5.8 PDE EQ ABDE 65 5.8 PDE EQ ABDE 65 5.8 PDE EQ ABDE 6	14:24-15:## 45 2##32/ 13 15:#6-15:46 12:#8-22:16 15:#6-15:46 18:#8-88:2# 13 2#232/ 13 ##:#8-88:2# 13 2#232/ 15 ##2:45-80:27 43 2#232/ 15 ##2:45-80:27 43 2#233/ #1 1#:#9-18:#1 13 2#233/ #1 1#:#9-18:#1 13 2#233/ #1 1#:#9-18:#1 13 2#233/ #1 13:31-13:46 16 2#233/ #2 ##1:32-81:43 16 2#233/ #3 2#2:46-23:27 42 2#233/ #3 22:46-23:27 42 2#233/ #3 22:46-23:27 42 2#233/ #3 22:46-23:27 42 2#233/ #3 22:46-23:27 42 2#233/ #3 22:46-23:27 42 2#233/ #3 #5:35-85:47 42 2#233/ #3 #7:#8-17:18 12 2#233/ #3 15:48-16:36 49 2#233/ #3 #7:#8-17:18 12 2#233/ 13 15:23-16:#8 49 2#233/ 13 15:23-16:#8 12 2#233/ 13 15:23-16:#8 12 2#233/ 13 15:23-16:#8 49 2#233/ 13 15:23-16:#8 49 2#233/ 13 15:23-16:#8 49 2#233/ 13 15:23-16:#8 49 2#233/ 13 15:23-16:#8 49 2#233/ 13 15:23-16:#8 49 2#233/ 13 15:23-16:#8 49 2#233/ 13 15:23-16:#8 49 2#233/ 13 15:23-16:#8 49 2#233/ 13 15:23-16:#8 49 2#233/ 13 15:23-16:#8 49 2#233/ 13 15:23-16:#8 49 2#233/ 13 15:23-16:#8 49 2#233/ #3 #################################
#6552 83 #2 12 #43 #8:47:18 8 #553 83 #2 12 #43 11:28:47:1 #554 83 #2 12 #43 11:28:47:1 #554 83 #2 12 #43 2:58:24.5 #555 83 #2 13 #44 #8:48:13:2 #656 83 #2 13 #44 #8:52:52:2 #657 83 #2 13 #44 #8:152:52:2 #657 83 #2 13 #44 #8:152:52:3 #658 83 #2 13 #44 #8:152:3 #11.8 #668 83 #2 13 #44 #8:23:11.8 #666 83 #2 13 #44 #3:23:11.8 #666 83 #2 13 #44 #3:23:11.8 #666 83 #2 13 #44 \$15:23:11.8 #666 83 #2 14 #45 #8:23:16.6 #666 83 #2 14 #45 #8:23:16.6 #666 83 #2 14 #45 #8:23:18.3 #666 83 #2 14 #45 #8:18:33.3 #666 83 #2 14 #45 #8:18:33.3 #666 83 #2 14 #45 #8:18:33.3 #666 83 #2 14 #45 #8:18:33.3	28.784S 178.2834 36.298M 71.853E 5.618N 125.555E 23.789N 185.5897E 39.988N 75.897E 39.988N 75.897E 39.988N 75.897E 39.9862N 75.275E 13.836N 75.128E 13.836N 75.128E 13.836N 75.128E 13.836N 126.521E 16.479N 126.215E 16.479N 126.215E 16.479N 126.215E 16.479N 126.215E 16.479N 179.71E 16.236S 161.388E 16.489E 16.489N 179.571E 18.236S 161.388E 18.888N 175.571E 18.236S 161.388E 18.388N 18.88E 18.38E	FIJI ISLANDS REGION AFGHANISTAN-USSR BORDER REGION MINDANAO PHILIPPINE ISLANDS WINNAAN PROVINCE CHINA SOUTHERN SINKIANG PROV CHINA MARIANAI SLANDS KIRGHIZ-SINKIANG BORDER REGION MINDANAO PHILIPPINE ISLANDS MOLUCCA PASSAGE SOUTH OF ALASKA LOYALTY ISLANDS REGION SOUTH OF ALASKA LOYALTY ISLANDS REGION SOUTH OF ALASKA LOYALTY ISLANDS MOLUCCA PASSAGE SOUTH OF ALASKA LOYALTY ISLANDS MOLUCCA PIJI ISLANDS MOLUCAND ISLANDS NORTHERN SINKIANG PROV CHINA PROCEABLE BONIN ISL. PO SO T TONGA ISLANDS WOLCAND ISLANDS WOLCAND ISLANDS WOLCAND ISLANDS WOLCAND ISLANDS WOLCAND ISLANDS WOLCAND ISLANDS MINDANAO PHILIPPINE ISLANDS WOLTH OF MARIAND REGION MINDANAO PHILIPPINE ISLANDS WOLTH OF SURBAWA ISLAND MINDANAO PHILIPPINE ISLANDS WINDI ISLANDS SOUTH OF MARIANA ISLANDS MINDANAO PHILIPPINE ISLANDS WINDI SURBAWA ISLANDS MINDANO PHILIPPINE ISLANDS MINDAND PHILIPPIN	38 5.8 6.8 PDE EQ AE 33 5.8 8.8 PDE EQ A 33 5.6 6.2 PDE EQ A 33 5.6 6.2 PDE EQ AE 33 5.6 6.2 PDE EQ AE 33 5.2 8.8 PDE EQ A 33 5.8 8.8 PDE EQ A 33 5.6 8.8 PDE EQ ABDE 33 5.8 4.9 PDE EQ A 33 5.8 4.9 PDE EQ A 33 5.8 5.8 PDE EQ ABDE 36 5.2 8.8 PDE EQ ABDE 36 5.2 8.8 PDE EQ AE 37 5.8 8.8 PDE EQ AE 37 5.8 8.8 PDE EQ AE 38 5.8 8.8 PDE EQ AE 38 5.8 8.8 PDE EQ AE 38 5.8 8.8 PDE EQ AE	#7:38-87:41 12 28836/ 81 14:26-14:37 12 28836/ 82 15:13-15:24 12 28836/ 83

KEY: IMPORMATION SOURCES - PDE-NEIS PDECARD, MON-NEIS MONTHLY LIST, ICS-ICS LIST, MEL-MELICORDER, OTH-OTHER EVENT TYPE - EQ-EARTHQUAKE, NX-NUCLEAR EXPLOSION, SX-SCIENTIFIC EXPLOSION, OT-OTHER, UN-UNKNOWN PHASES - A-P, B-PO, C-S, D-SO, E-T, F-OTHER

EVN: ******ORIGIN TIME****** *NO" YR*MO*DA*JUL*HR*MN*SECS	**CODED! NATES***	**************************************	0EP **	AGNI" INF DY"SRF SRC	EV PHASE	***SAVED******** S**INTERVAL***MHS	**STRIP** TAPE/FILE
# 3 8 3 8 2 5 8 5 6 22:49:52.1 # 3 8 3 8 2 6 8 5 7 8 3:81:36:4 # 3 8 8 2 6 8 5 7 8 5:83:81:26.4 # 3 8 8 2 2 6 8 5 7 8 5:83:81:26.4 # 3 8 3 8 2 2 6 8 5 7 8 5:83:81:26.4 # 3 8 3 8 2 2 6 8 5 7 8 5:83:81:26.4 # 3 8 3 8 2 2 6 8 5 7 8 5:83:81:26.4 # 3 8 3 8 2 2 6 8 5 7 8 5:83:81:26.4 # 3 8 3 8 2 2 6 8 5 7 7 8 5:83:81:26.4 # 3 8 3 8 2 2 6 8 5 7 7 8 5:83:88:88:88:88:88:88:88:88:88:88:88:88:	35. 040M 79.8382 7.3735 107.1165 5.4235 146.8262 18.755 9.552V 22.418M 121.234E 31.931N 131.678E 31.931N 131.678E 31.2175 136.853E 18.8382 172.286W 49.231N 155.6175 58.5245 162.796E 8.8382 8.973N 72.815E 36.899N 141.9262 35.072M 139.948E 32.638M 75.5677 44.164M 148.839E 48.824M 72.1532 24.164M 148.839E 48.824M 72.1532 24.164M 148.839E 18.938 142.219E 48.8382 142.219E 48.8382 142.219E 18.8382 142.219E 18.8382 8.8382 142.219E 18.8382 8.8382 142.219E 18.8382 8.8382 8.8382 142.219E 18.8382 8.8382 8.8382 142.2486 141.433N 139.873E 6.1525 188.3392 8.8382 13.9538 143.3252 14.2655 14.3192 8.8382 8.8382 143.392 8.3928N 85.5972 14.7638 143.3252 14.2655 14.3192 8.8382 8.8382 14.398 8.8382 14.398 8.8382 14.398 8.8382 14.398 8.8382 14.398 143.3252 14.2655 14.3192 8.8382 14.398 143.3252 14.2655 14.3192 8.8382 14.398 143.3252 14.2655 14.3192 8.8382 14.398 143.3252 14.2655 14.3192 8.8382 14.398	EASTERN RASHMIR JAVA EAST PAPUA NEW GUINEA REGION NOPTHERN CHILE TAIVAN REGION KYUSHU JAPAN WEST IRIAN TONGA ISLANDS REGION KURIL ISLANDS AUCKLAND ISLANDS REGION MINKNOWN P PHASE AFGRANISTAN-USSP BORDER REGION NEAP S COAST OF HONSHU JAPAN NEAP S COAST OF HONSHU JAPAN KASHMIR-TIBET BORDER REGION KURIL ISLANDS KIRCHIZ SSR BONIN ISLANDS REGION PAPUA NEW GUINEA TUPPMEN SSR SOUTH OF HONSHU JAPAN INDIA-CHINA BORDER REGION SOUTHERN SUMATERA PROJABLE MARIANAS FORDIN ISLANDS SOUTHERN SUMATERA PROJABLE MARIANAS SOUTHERN SINKIANG PROV CHINA HOKKAIDO JAPAN REGION WEST CARDIINE ISLANDS SOUTHERN SINKIANG PROV CHINA HOKKAIDO JAPAN REGION OFF E COAST OF HONSHU JAPAN OFFIERN SINKIANG PROV CHINA HOKKAIDO JAPAN REGION OFF E COAST OF HONSHU JAPAN SOUTH ATLANTIC RIDGE PROBABLE MARIANAS SOUTH ATLANTIC RIDGE PROBABLE MARIANAS SOUTH BRONDER REGION OFF E COAST OF HONSHU JAPAN OFFIERN SINKIANG PROV CHINA HOKKAIDO JAPAN REGION OFF E COAST OF HONSHU JAPAN SOUTH ATLANTIC RIDGE PROBABLE MARIANAS SOUTH BRONDER REGION OFFIERN HONSHU JAPAN SOUTHERN HONSHU JAPAN SOUTHERN HONSHU JAPAN OFFIERST SILANDS EAST PAPUA NEW GUINEA REGION MARIANA ISLANDS EAST PAPUA NEW GUINEA REGION MARIANA ISLANDS GILBERT ISLANDS REGION MARIANA ISLANDS GANDA SEA	38 4 5 6 5 6 5 6 5 6 5 6 5 6 5 6 5 6 5 6 5	9 1 2 4 5 6 7 5 6 7 5 7 5 7 5 7 5 7 5 7 5 7 5 7	EQUAASS DE E EEQUAAS DE EEQUAAS D	22:45-22:58 46 22:59-23:18 28 82:88-82:28 13 83:83-83:51 49 85:34-85:51 18 86:84-86:15 12 87:12-97:57 16 11:49-12:88 12 12:15-16:34 12 12:15-12:58 14 28:15-28:26 12 18:12-18:26 12 18:12-18:26 12 18:12-18:26 12 18:12-18:27 13 28:48-28:59 12 89:83-89:41 39 22:19-23:87 49 89:83-89:41 39 13:28-13:39 12 19:32-19:44 13 81:58-88:12-88 88:12-88:27 16 88:12-88:28 12 88:12-88:28 13 16:44-17:28 15 17:15-28:19 18 18:38-18:19 38	28837/ 22 28837/ 23 28837/ 25 28837/ 26 28837/ 26 28838/ 81 28838/ 82 28828/ 83 28828/ 85
#751 #3 #3 1# #69 #8:27:49.4 #752 #3 #3 1# #69 #8:27:49.4 #753 #3 #3 1# #69 #8:46:86.# #753 #3 #3 1# #69 #8:46:86.# #755 #3 #3 11 #7# #2:24:49.3 #755 #3 #3 11 #7# #2:24:49.3 #755 #3 #3 11 #7# #2:24:49.3 #755 #3 #3 11 #7# #2:24:49.3 #756 #3 #3 12 #71 #2:42:34:42.5 #757 #3 #3 12 #71 #2:43:16:44:44.6 #758 #3 #3 12 #71 #2:#8:15.# #758 #3 #3 12 #71 #2:#8:15.# #758 #3 #3 12 #71 #2:#8:15.# #758 #3 #3 12 #71 #2:#8:15.# #768 #3 #3 12 #71 #2:#8:15.# #768 #3 #3 12 #71 #2:#8:15.# #768 #3 #3 12 #71 #2:#8:15.# #768 #3 #3 12 #71 #2:#8:15.# #768 #3 #3 12 #71 #2:#8:15.# #768 #3 #3 12 #71 #2:#8:15.# #768 #3 #3 12 #71 #2:#8:15.# #768 #3 #3 12 #71 #2:#8:15.# #768 #3 #3 13 #72 #2:54:18.# #769 #3 #3 13 #72 #2:54:18.# #769 #3 #3 13 #72 #2:54:18.# #778 #3 #3 #3 13 #72 #2:54:18.# #778 #3 #3 #3 #3 #72 #2:54:18.# #777 #3 #3 #3 #3 #72 #2:54:18.# #777 #3 #3 #3 #3 #72 #2:54:18.# #777 #3 #3 #3 #3 #72 #2:54:18.# #777 #3 #3 #3 #3 #72 #2:54:18.# #777 #3 #3 #3 #3 #72 #2:32:3#:14.6 #778 #3 #3 #3 #3 #72 #2:32:3#:14.6 #779 #3 #3 #3 #4 #73 #2:12:2:44.9 #779 #3 #3 #3 #4 #73 #2:12:2:44.9 #779 #3 #3 #3 #4 #73 #2:2:2:2:4.9 #779 #3 #3 #3 #4 #73 #2:2:2:2:4.9 #779 #3 #3 #3 #4 #73 #2:2:2:2:4.9 #779 #3 #3 #3 #4 #73 #2:2:2:2:4.9 #779 #3 #3 #3 #5 #74 #2:2:27:53.2 #778 #3 #3 #3 #5 #74 #2:2:27:53.2 #778 #3 #3 #3 #5 #74 #2:2:27:53.2 #778 #3 #3 #3 #5 #74 #2:2:27:53.2 #779 #3 #3 #3 #5 #74 #2:2:27:53.2 #779 #3 #3 #3 #5 #74 #2:2:27:53.2 #779 #3 #3 #3 #5 #74 #2:2:27:24.9 #779 #3 #3 #3 #5 #74 #2:2:27:24.9 #779 #3 #3 #3 #5 #74 #2:2:27:12.4	43.811N 147.441E 14.6455 157.282E 5.411N 126.768E 18.249S 13.438W 6.992S 147.374E 5.372N 126.623E 36.435N 78.959E 4.1837S 127.918E 5.1828 127.918E 5.1828 127.918E 5.1828 127.918E 5.1828 127.918E 6.671E 18.898S 166.128E 8.671N 93.889E	KURIL ISLANDS VANUATU ISLANDS VANUATU ISLANDS MINCANAO PHILIPPINE ISLANDS ASCENSION ISLAND REGION EAST PAPUA NEW GUINEA REGION MINDANAO PHILIPPINE ISLANDS MINDANAO PHILIPPINE ISLANDS MINDANAO SEA BANDA SEA BANDA SEA FOX ISL. ALEUTIAN ISL. MINDANAO PHILIPPINE ISLANDS WANYATU ISLANDS BANDA SEA SICHUAN PROVINCE CHINA BANDA SEA NICOBAR ISLANDS REGION MASIAN ISLANDS REGION EAST PAPUA NEW GUINEA REGION KAMCHATKA MOKAIDO JAPAN REGION KAMCHATKA MEAR E COAST OF KAMCHATKA MOKAIDO JAPAN REGION BANDA SEA FOX ISL. ALEUTIAN ISL. TURKMEN SSR TURYMEN SSR T	42 6 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5	2 5.8 PDE 3 8.8 PDE 4 5.6 PDE 9 8 8 PDE 9 8 8 PDE 2 8 8 PDE 5 8 8 PDE 6 5 PDE 8 6 8 PDE 6 5 9 PDE 6 5 9 PDE 8 8 8 PDE	EO ABDE EO A ABDE	## 129-#1:12 44 #5:48-#5:59 12 12:#5:48-#5:59 12 12:#5:48-#5:59 12 13:#5:52-#3:58 47 23:#5:52-#3:155 89 #7:#2-#3:155 59 #7:#2-#3:155 59 #7:#2-#7:36 13 #8:52-#3:155 59 #7:#2-#7:36 13 #8:52-#3:45 13 #8:13-18:#5 13 #8:13-18:#5 13 #8:13-18:#7 13 21:11-21:23 13 21:11-21:23 13 21:11-21:23 13 21:11-21:23 13 21:11-21:23 13 21:11-21:23 13 21:11-21:23 13 21:11-21:23 13 21:11-21:23 13 21:12-12:23 13 #8:28-28:28 16 1#:#7-18:24 17 28:28-28-28:38 12 12:21-12:32 12 12:21-12:32 12 12:21-12:32 12 12:31-22:42 13 12:31-25:42 13 12:31-25:42 13 12:31-25:42 13 12:31-25:42 13 12:31-25:42 13 13:31-25:42 13 13:31-25:42 13 13:31-25:42 13 13:31-25:42 13 13:31-25:42 13 13:31-25:42 13 13:31-25:42 13 13:31-25:42 13 13:31-25:42 13 13:31-25:42 13 13:31-25:42 13 13:31-25:42 13 13:31-35:31 13 33:31-35:31 13 33:31-35:31 13 33:31-35:31 13 33:31-35:31 13 33:31-35:31 13 33:31-35:31 13 33:31-35:31 13 33:31-35:31 13 33:31-35:31 13 33:31-35:31 13 33:31-35:31 13 33:31-31 13 33:31 13 33:31 13 33:31 13 33:31 1	28848/-87 22048/89 8/88 8/88 8/88 28848/-18 28848/-18 28848/-11 22848/-11 22848/-11 22848/-11 22848/-11 22848/-12 22848/-12 22848/-12 22848/-12 22848/-12 22848/-13 22848/-13 22848/-13 22848/-13 22848/-18 22

KEY: INFORMATION SOURCES - PDE=NEIS PDECARD, MON-NEIS MONTHLY LIST, ICS=ICS LIST, NEL=NELICORDER, OTH-OTHER EVENT TYPE - EO-EARTHQUAKE, NX-NUCLEAR EXPLOSION, SX-SCIENTIFIC EXPLOSION, OT-OTHER, UN-UNKNOWN PHASES - A-P. B-PO, C-S, D-SO, E-T, F-OTHER

PHASES - APP, B=PO, C=S			
EVAT *****ORIGIN TIME***** **COORDINATES*** *********************************	********LOCATION**********	DEP "MAGNI" INF EV PP"	****SAVED******* **STRIP**
and Akamoana Johankamu 2562 antwice and Course	DESCRIPTION CONTRACTOR	KH- BUY-SKI SKC II YAKS	ra- intender - mas investite
### ### ### ### ### ### ### ### ### ##	NEW IRELAND REGION	119 4.8 8.8 PDE EQ ABD	23:88-23:15 16 28842/-11 11:28-11:39 12 28842/ 12
#8#3 #3 #3 2# #79 13:45:49.8 4.7275 153.125E	NEW IRELAND REGION	86 5.9 8.8 PDE EQ ABDE	13:46-14:27 42 28842/ 13
#884 83 83 28 879 16:24:13.3 18.8275 177.762W	FIJI ISLANDS REGION	632 5.8 8.8 POE EO A	16:26-16:37 12 20242/ 14
#8#5 83 #3 2# #79 17:44:#7.1 18.8#45 177.679V	FIJI ISLANDS REGION	634 4.9 8.8 PDE ED A	17:46-17:57 12 22242/ 15 84:87-84:48 42 28242/-16
##F7 #3 #3 21 ##F #7:44:17.6 21.6685 175.321W	TONGA ISLANDS	78 6.3 8.8 PDE EO A	#7:47-#7:59 13 28P42/ 17
#6#8 63 #3 21 #8# #7:57:18.4 7.3425 128.92#E	BANDA SEA	143 5.6 8.8 PDE EQ A	#8:E8-88:12 13 2EE42/ 17
#8#9 83 #3 21 #8# 15:23:4#.7 36.5#2N 7#.593E	HINDU KUSH REGION	166 4.2 8.8 PDE EG A	15:31-15:42 12 22842/ 18 15:53-16:84 12 22842/ 19
#811 #3 #3 22 #81 #1:32:28.6 51.297N 178.481W	ANDPEANOF ISL. ALEUTIAN ISL.	33 4.9 8.8 PDE EO ABD	#1:34-81:58 17 28842/-28
#812 83 #3 22 #81 28:17:87.5 53.211N 162.259E	OFF EAST COAST OF KAMCHATKA	33 4.7 #.# PDE EG ABO	28:19-28:35 17 28842/ 21
#813 83 #3 23 #82 #6:#9:29.6 6.6185 154.612E #814 83 #3 23 #82 #6:48:32.1 6.6635 154.859E	SOLOMON ISLANDS	41 5.7 6.2 PDE EG ABDE 55 5.2 8.8 PDE EG ABDE	#6:1#-#6:52 43 28242 / 22 #6:41-#7:23 43 28242 / 22
#815 83 #3 23 #82 #8:26:55.2 6.6275 154.5855	SCHALZI MONCHOZ	35 5.4 5.6 PDE EQ ABDE	#8:28-#9:#9 42 22242/ 23
#816 83 #3 23 #82 12:11:23.9 37.#11N 71.487E	AFGHANISTAN-USSR BORDER REGION	122 5.2 8.8 PDE EG A 68 5.2 8.8 PDE EG ABDE	12:18-12:38 13 2P842/ 24 14:86-14:47 42 28843/ 81
#817 83 #3 23 #82 14:#4:52.1 6.7#25 154.5282 #818 83 #3 23 #62 16:#7:21.2 3.3755 177.4912	GILBERT ISLANDS REGION	27 4.8 8.8 PDE EO ABD	16:87-16:21 15 2EF43/ 82
#819 83 #3 23 #82 21:54:36.3 37.224N 138.125E	NEAR W COAST OF HONSHU JAPAN	54 4.6 8.8 PDE EO ABD	21:56-22:11 16 28843/ 83
## 818 ## 83 ## 32 ## 82 16:#7:21.2 3.3755 177.491E ## 819 ## 83 ## 32 ## 82 21:54:36.3 37.2244 138.125E ## 82 ## ## 82 ## ## 82 ## 83 ## 83 ## 82 ## 82 ## 82 ## 83 ##	LOCAL PO. SO. T	151 4.9 M.M PDE EO ASD	#8:22-88:51 3E 2EE43/ 84 87:83-87:18 16 2EE43/ 85
#822 83 #3 24 #63 E7:24:12.5 7.6155 1#7.173E	JAVA	57 4.9 8.8 PDE EQ A	#7:3E-#7:41 12 28E43/ E6
#823 83 #3 24 #63 #8:44:53.9 6.6#2\$ 154.564E	SOLOMON ISLANDS	52 5.1 8.8 PDE EQ ABDE	#8:46-#9:27 42 28843/ #7 18:48-11:29 42 28843/ #8
#824 UJ #3 24 DU3 12:4/:4r.5 4./215 153.3361 #6?5 83 #3 24 #83 12:24:35.2 1#.8145 163.9926	SOLOMON ISLANDS	74 4.9 8.8 PDE EQ ABD	12:25-12:41 17 20043/ 89
8926 88 88 88 888 88:88:88.8 8.882N 8.82E	PROBABLE MARIANAS PO. SO, T	# B.# B.B HEL EO BDE	2F:1E-2E:39 3E 2E243/ 1E
#627 83 #3 25 264 #7:44:21.2 6.5555 132.##65 #828 83 #3 25 #84 13:22:14.2 6.4945 151.722E	BANDA SEA NEW RRITAIN REGION	45 5.3 8.8 PDE EQ ABDE	#7:47-#7:58 12 28#43/ 11 13:#1-13:44 44 28#43/ 12
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#838 83 83 25 884 21:41:54.8 6.49FS 155.118E	SOLOMON ISLANDS	123 5.1 #.# PDE EQ ABDE	21:42-22:24
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#828 83 #3 25 #84 19:31:92:14.2 6.4945 151.7222	PAPUA NEW GUINEA	BR 4.4 B.B POE EC ABD	17:21-17:37 17 28843/ 18 18:44-19:88 17 28843/ 19
#835 83 #3 26 #85 25:28:#8.# 37.3#1N 116.46F4	SOUTHERN NEVADA	# 5.1 #.# PDE NX A	28:26-28:37 12 28844/ 81
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#847 #3 #3 #3 3# #89 12:#2:52.3 6.6375 184.5665 #848 #3 #3 3# #89 12:#2:52.3 6.6375 184.5665 #848 #3 #3 3# #89 18:55:49.1 5.9195 142.22# #85# #3 #3 31 #9# 13:24:45.6 34.281M 135.8336	NEAR S COAST OF SOUTH HONSHU	38. 4.9 B.B POE EO ABDE	13:25-14:11 47 28844/ 15
8851 83 84 81 891 81:84:59.8 18.534N 145.8865 8852 88 88 82 888 88:88:82.8 8.228N 8.8865 8853 83 84 82 892 84:18:23.5 3.1235 135.2876 8854 83 84 82 892 84:18:23.5 3.1235 135.2876	MARIANA ISLANDS	102 5.1 8.8 PDE EQ ABDE	#1:#4-#1:36 33 2##44/ 16 2#:43-21:12 3# 2##44/ 17
8052 88 88 82 888 88:88:82.8 8.828N 8.882E	FROBABLE BONIN PO. SO. T	33 5.8 8.8 PDE EQ A	84:28-84:32 13 28044/ 18
#853 83 #4 #2 #92 #4:18:23.5 3,1235 135.2672 #854 83 #4 #2 #92 #5:#9:18.3 54.#99N 16#.314E #855 83 #4 #2 #92 #6:49:#1.9 24.#115 176.196V #856 83 #4 #2 #92 21:#4:3#.6 6,3275 154.553E	NEAR EAST COAST OF KAMCHATKA	92 4.9 #.# PDE EQ ABD	#5:11-#5:26 16 2R844/ 19
#855 83 #4 #2 #92 #6:49:#1.9 24.7115 176.196V	SOUTH OF FIJI ISLANDS	AN 5.4 5.N PDE EQ A	#6:52-#7:#4 13 2##44/ 2# 21:#5-21:2# 16 2##44/ 21
#856 83 #4 #2 #92 21:#4:3#.6 6.3275 154.5336 #857 83 #4 #2 #92 21:55:44.6 36.393N 7#.7346 #858 83 #4 #3 #93 22:5#:#57.0 0.731N 83.1154.584 #859 83 #4 #3 #93 1#:14:53.# #671N 124.1816 #868 83 #4 #3 #93 19:14:#5.2 \$1.997N 179.2556	HINDU KUSH REGION	215 4.8 8.8 PDE EQ A	22:#3-22:14 12 28845/ #1
#858 #3 #4 #3 #93 #2:5#:#E.7 #.731N #3.115V	COSTA RICA	33 6.5 7.2 PDE EQ AE	#2:59-#5:#8 13# 2##45/ #2 1#:18-1#:29 12 2##45/ #3
#859 83 #4 #3 #93 19:14:53.8 #.671N 124-181E #868 83 #4 #3 #93 19:14:#5.2 \$1.997N 179.255E #861 83 #4 #3 #93 19:26:24.5 \$1.8#9N 176.919W #862 83 #4 #4 #94 #8:45:#7.9 38.517N 78.334E	MINAHASSA PERINSULA BAT ISL. ALFUTIAN ISL.	116 5.6 F.S PDE EQ ABDE	19:16-28:84 49 26245/ 84
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#862 #3 #4 #4 #94 #8:45:#7.9 38.517N 78.334E	AFGHANISTAN-USSR BORDER REGION	33 4.9 8.8 PDE EQ A	##:52-#1:#4 13 2##45/ #5 #2:58-#3:#9 12 2##45/ #6
#863 83 84 84 894 #2:51:34.9 6.736N 94.811E #864 83 84 84 894 #3:83:35.8 5.752N 94.798E		74 6 0 8 6 801 10 4	#3:18-#3:21 12 28845/ 86
8855 88 82 88 888 82:88:88.8 8.882N 8.8821	PROBABLE BONIN'S PO. SO, T	B B. B B. B HEL EQ BDE	#5:25-#5:54 38 28845/ 87 #7:#5-#7:16 12 28845/ 88
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#868 83 84 84 894 19:48:84.1 7.8575 129.379E	BANDA SEA	100 3.8 B.B FUL EU M	19:51-28:82 12 28845/ 89 23:14-23:59 46 22845/ 18
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#871 83 #4 #5 #95 #2:55:24.6 41.631N 7#.763E	KIRGHIZ SSR	33 5.1 #.# PDE EQ A	#1:#3-#1:14 12 22#45/ 11
#872 83 #4 #5 #95 #6:5#:33.5 4#.#62N 75.285E	KIRGHIZ-XINJIANG BORDER REGION	33 5.4 5.6 PDE EQ A 33 4.8 8.8 PDE EQ A	#6:57-#7:E9 13 2E046/ B1 #7:14-#7:26 13 2E046/ BD
#873 83 #4 #5 #95 #7:#7:23.5 39.874N 75.449E #874 83 #4 #5 #95 #7:4#:38.3 5.438N 126.698E		41 5.1 8.8 POE EQ A	#7:43-#7:54 12 2#246/ #3
#875 83 #4 #5 #95 21:39:#8.# 26.489N 126.488E	RYUNYU ISLANDS	121 4.7 B.B PDE EO ABD	21:41-21:58 18 28845/ 84 88:51-81:87 17 28846/ 85
#876 83 84 86 896 88:58:23.7 6.1715 149.8655 8877 83 84 86 896 82:88:11.8 36.438N 71.477E	NEW BRITAIN REGION AFGHANISTAN-USSR BORDER REGION	53 4.8 4.7 PDE EQ ABD 181 4.7 8.8 PDE EQ A	#2:15-#2:27 13 28246/ #6
#878 #3 #4 #6 #96 12:3#:36.6 5.5485 153.56#£	NEW IRELAND REGION	SU 4.5 P.P PUL LU ADV	1#:31-1#:46 16 2PP46/ #7
#879 83 84 86 896 12:49:28.4 39.974N 75.129E #888 83 84 86 896 18:88:28.2 21.6145 179.391W	ZOUTHERN XINDIANG CHINA	33 4.7 8.8 PDE EQ A 636 4.9 8.8 PDE EQ A	18:18-18:21 12 22446/ 29
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###2 #3 #4 #6 #96 22:18:47.5 38.846N 71.488E	AFGHANISTAN-USSK BORDER REGION	33 5.2 4.7 PDE EQ A 33 5.8 8.8 PDE EQ A	22:26-22:37 12 28R46/-11 #5:18-85:22 13 22846/ 12
#883 83 84 87 897 85:84:22.3 5.9965 188.7522 8864 83 84 87 897 22:56:52.3 54.512N 168.929E	NEAR EAST COAST OF KAMCHATKA	33 5.1 #.# PDE EQ ABDE	22:59-23:48 58 28846/-13
#985 83 #4 #8 #98 #2:28:25.5 11.448N 57.5#9E	ARASIAN SEA	1# 5.8 6.2 PDE EQ F	#2:38-84:29 128 28846/-14 #6:28-86:39 12 28846/ 15
#886 83 84 88 898 86:21:24.8 5.648N 94.719E 8887 83 84 88 898 14:18:53.3 49.636N 155.286E	NURIHERN SUMATERA KURIL ISLANDS	76 4.8 8.8 PDE EQ A 82 5.8 8.8 PDE EQ ABDE	14:12-14:58 47 22245/ 16
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889 83 84 18 188 82:28:36.3 17.6365 174.648V 8098 83 84 18 188 85:38:47.7 48.826N 77.868E	TONGA ISLANDS	33 5.4 5.3 PDE EQ A	#5:46-#5:57 12 22846/ 19
##91 B3 #4 1# 1#3 #5:51:37.6 22.1285 179.7#9¥	2001H O. 1121 12FW402	888 3.3 P.S PUL LU A	#5:54-86:#5 17 28845/ 19
#992 #3 #4 1# 1## 12:19:37.5 2.3#2N 125.7#5E	MOLUCCA PASSAGE	94 5.8 8.8 PDE EQ A 33 5.3 5.3 PDE EQ ABDE	12:22-12:33 12 20045/ 28 18:38-19:88 39 20246/ 21
8853 83 84 18 188 18:29:35.4 38.785N 141.952C 8894 63 84 11 181 83:18:18.3 18.449N 62.786W	SOUTH OF HUMSHU JAPAN	43 6.8 5.8 PDE EQ A	#8:27-#8:45 28 2P246/ 22
#895 83 #4 11 1#1 14:32:54.9 4#.315N \$3.#83E	TUR MEN SSR	33 4.8 8.8 PDE EQ A	14:41-14:52 12 2P846/ 23 15:36-16:19 44 2P846/ 24
#896 83 #4 11 1#1 15:34:55.6 44.3#.N 147.775E #897 83 #4 11 1#1 17:#3:45.9 35.48#5 179.383W	KURIL ISLANDS	98 5.7 8.8 PDE EQ ABDE 84 5.7 8.8 PDE EQ A	17:#8-17:19 12 CP#46/ 25
##98 93 #4 12 1#2 #3:41:#5.2 49.815N 78.227E	EASTERY PAZAKM SSR	# 4.9 #.# PDE NX A	#3:48-#3:59 12 22846/ 26
#29° #3 #4 12 1#2 12:#7:54.6 4.8655 79.1834	PERU-ICUATOR BURDER REGION	187 6.5 # ' PDE EQ A	12:17-12:36 2# 22246/ 27 16:#2-16:13 12 2#846/ 28
#9#2 93 #4 12 1#2 16:##:56.4 12.1695 167.#BBE	SMITTER OF JE TOURNESS		

KEY: INFORMATION SOURCES - PDE-NEIS PDECARD. MON-NEIS MONTHLY LIST, ICS-ICS LIST, MEL-HELICORDER, OTH-OTHER EVENT TYPE - EQ-EARTHQUAKE, NX-NUCLEAR EXPLOSION, SX-SCIENTIFIC EXPLOSION, OT-OTHER, UN-UNKNOWN PHASES - A-P, B-PU, C-S, D-SO, E-T, F-OTHER

PHASES	- A-P. B-PU. C-S	, D=SO, E=T, F=OTHER				

## 14 18 ## 14 18 ## 18 18 18 18 18 18 18 18 18 18 18 18 18	18. 4315 165. 2286 16. 3630 71. 112E 35. 9530 769. 9631 22. 1265 175. 5032 18 8. 88810 12. 1265 175. 5032 18 8. 88810 12. 1265 175. 6114 6. 5822 154. 99. 1462 13. 1385 121. 971E 13. 1385 121. 971E 13. 1385 121. 971E 13. 1385 127. 6284 14. 7550 142. 8450 1.7750	NEW IRELAND REGION SANTA CRUZ ISLANDS MAPIANA ISLANDS MOPT-WEST OF KURIL ISLANDS BUPIJA MARIANA ISLANDS SANIA CRUZ ISLANDS SANIA CRUZ ISLANDS AFGMANISTAN-USSR BORDER REGION HINDU KUSH REGION TONIA ISLANDS REGICN PROCABLE MARIANAS PO. SO. T SOUTHERN NEVADA MOLUCCA PASSAGE TONIA ISLANDS SOUTHEAST ASIA MINDORO PHILIPPINE ISLANDS NEAR EAST COAST OF KAMCHATKA SOUTH OF JAVA TALAUD ISLANDS HOKNAIDO JAPAN REGION NORTHERN SUMATERA VANUATU ISLANDS BURMA TONCA ISLANDS SOUTHERS TABLE SOUTHERS FOR THE SECTION KURIL ISLANDS SOUTHERN IRAN FLORES ISLAND REGION OFF EAST COAST OF KAMCHATKA VANUATU ISLANDS KURIL ISLANDS SOUTHERN IRAN FLORES ISLAND REGION OFF EAST COAST OF KAMCHATKA VANUATU ISLANDS KURIL ISLANDS SOUTH OF HONSHU JAPAN BANDA SEA KURIL ISLANDS SOUTH OF HONSHU JAPAN BANDA SEA KURIL ISLANDS MOULCCA PASSAGE PROSABLE MARIANA ISLANDS MARIANA ISLANDS MARIANA ISLANDS MOLUCCA PASSAGE PROSABLE MARIANAS PO., SO. T SCUTHEAST ASIA UZBEK SSR	333 2244 333 33 446 454 34 46 45 45 45 45 45 45 45 45 45 45 45 45 45	5.8 4.4 PDE 4.6 B PDE 4.6 B PDE 5.7 B PDE 5.8 B B PDE 5.8 B B PDE 5.8 B B PDE 5.8 B B PDE 6.8 B B PDE 6.8	CETEGENES E E E E E E E E E E E E E E E E E E	### ### ### ### ### ### ### ### ### ##
## 951 83 ## 22 112 13:53:80.# ## 952 ## ## 22 112 18:28:24.5 ## 953 83 ## 22 112 18:28:24.5 ## 953 83 ## 23 113 ##:28:56.3 ## 955 83 ## 23 113 ##:28:56.3 ## 955 83 ## 23 113 ##:28:13.# ## 955 83 ## 23 113 ##:28:13.# ## 955 83 ## 23 113 ##:28:13.# ## 955 83 ## 23 113 ##:28:13.# ## 955 83 ## 23 113 ##:28:13.# ## 955 83 ## 24 114 ##:3:26:39.# ## 955 83 ## 24 114 ##:3:26:39.# ## 956 83 ## 24 114 ##:3:26:39.# ## 956 83 ## 24 114 ##:3:26:39.# ## 956 83 ## 24 114 ##:3:26:39.# ## 956 83 ## 24 114 ##:3:56:33.# ## 956 83 ## 24 114 ##:3:56:33.# ## 956 83 ## 25 115 ##:55:66.33.# ## 956 83 ## 25 115 ##:55:56.# ## 956 83 ## 25 115 ##:55:56.# ## 956 83 ## 25 115 ##:55:56.# ## 957 83 ## 25 115 14:##:3:47.# ## 957 83 ## 26 116 11:13:14.5 ## 957 83 ## 26 116 11:13:14.5 ## 957 83 ## 26 116 11:13:14.5 ## 957 83 ## 26 116 11:13:14.5 ## 957 83 ## 26 116 11:13:14.5 ## 957 83 ## 26 116 11:13:14.5 ## 957 83 ## 26 116 11:13:14.5 ## 957 83 ## 26 116 11:13:14.5 ## 957 83 ## 26 116 11:13:14.5 ## 957 83 ## 26 116 11:13:14.5 ## 957 83 ## 26 116 11:13:14.5 ## 957 83 ## 26 116 11:13:13:13.5 ## 957 83 ## 26 116 11:13:13:13.5 ## 957 83 ## 26 116 11:13:13:13.5 ## 957 83 ## 26 116 16:18:2:17.5 ## 957 83 ## 26 116 16:18:2:17.5 ## 957 83 ## 26 116 16:18:2:17.5 ## 957 83 ## 26 116 16:18:2:17.5 ## 957 83 ## 26 116 16:18:2:17.5 ## 957 83 ## 26 116 16:18:2:17.5 ## 957 83 ## 26 116 16:18:2:17.5 ## 957 83 ## 26 116 16:18:2:17.5 ## 957 83 ## 26 116 16:18:2:17.5 ## 957 83 ## 26 116 16:18:2:17.5 ## 957 83 ## 26 116 16:18:2:17.5 ## 957 83 ## 26 116 16:18:2:17.5 ## 957 83 ## 26 116 16:18:2:17.5 ## 957 83 ## 26 116 16:18:2:17.5 ## 957 83 ## 26 116 18:18:37.# ## 957 83 ## 27 117 ## 16:28:19.3 ## 957 83 ## 27 117 ## 16:28:19.3 ## 957 83 ## 27 117 ## 16:28:19.3 ## 957 83 ## 27 117 ## 16:28:19.3 ## 957 83 ## 28 118 ## 86:58:#6.# ## 958 #	37.112N 116. #22W # .888W # .2.28E	SOUTHERN NEVADA PROEABLE BONIN 1SL. PO.SO.T NORTHERN KIDIANG CHINA NEW BRITAIN REGION PAKISTAN SOUTH OF SUMBAWA ISLAND PAKISTAN EASTER ISLAND CORDILLERA FIJI ISLANDS REGION SOUTH OF FIJI ISLANDS NORAR S COAST OF HONSHU JAPAN SOUTH OF HARIANA ISLANDS SOLOMON ISLANDS TONGA ISLANDS TONGA ISLANDS TONGA ISLANDS AMMALIN ISLAND KAMCHATKA BAST COAST OF KAMCHATKA SOLOMON ISLANDS SAMMALIN ISLAND KAMCHATKA EAST PAPUA NEW GUINEA LUZON PHILIPPINE ISLANDS KAMCHATKA VANUATU ISLANDS TAIWAN NEW BRITAIN REGION HINDU KUSH REGION HINDU KUSH REGION TONGA ISLANDS KAMCHATKA E. PAPUA NEW GUINEA REGION NEW BRITAIN REGION HINDU KUSH REGION HOKAIDO JAPAN REGION HOKKAIDO JAPAN HOKLI ISLANDS HOKLI ISLANDS HOKLI ISLANDS HOKLI ISLANDS HOKLI ISLANDS	# # # # # # # # # # # # # # # # # # #	8 8.8 PDE 1.8 8.8 HEL 1.1 4.1 PDE 1.8 8.8 PDE 1.8 4.3 PDE 1.8 8.6 PDE 1.7 8.8 PDE 1.8 8.6 PDE 1.8 8.6 PDE 1.8 8.6 PDE	NY BAGAAAAAAAAAAAAAAAAAAAAAAAAAAAAAAAAAAA	13:59-14:18 12 28249/ 12 14:32-15:81 39 28245/ 13 18:34-18:46 12 20249/ 14 88:32-81:14 43 20249/ 15 84:22-84:33 12 28249/ 16 89:25-89:36 12 28249/ 17 18:58-19:89 12 28249/ 18 23:42-88:16 35 8/ 82 83:329-83:48 12 8/ 82 84:49-85:84 16 8/ 82 89:28-89:48 32 28249/ 19 89:28-89:48 32 28249/ 19

KEY: INFORMATION SOURCES - PDE=NEIS PDECARD, MON=NEIS MONTHLY LIST, ICS=ICS LIST, MEL=HELICORDER. OTH=OTHER EVENT TYPE - EG=EARTHQUAKE, NX=NUCLEAR EXPLOSION, SX=SCIENTIFIC EXPLOSION, OT=OTHER, UN=UNKNOWN PHASES - A=P, B=PO, C=S, D=SO, E=T, F=OTHER

EVAT *****ORIGIN TIME*****	**COORDINATES***	**************************************	DEP "MAGNI" INF EV """ KM" BDV"SRF SRC TP PHAS	****SAVED******* **STRIP** ES**INTERVAL***MNS TAPE/FILE
1881 83 85 83 123 82:38:15.4 1847 63 85 83 123 123 18:39:39.3 1843 83 85 83 123 18:37:13.4 1844 83 85 83 123 18:37:13.4 1845 83 85 84 124 14:29:38:3 1847 83 85 85 125 84:43:55.4 1842 83 85 85 125 85:49:45.2 1842 83 85 85 125 85:49:45.2 1842 83 85 85 125 85:29:89.8 1818 83 85 85 125 15:29:89.8 1818 83 85 85 125 15:29:89.8 1811 83 85 86 126 18:24:18.5 1814 83 85 86 126 18:24:18.5 1814 83 85 87 127 86:11:24.1 1815 83 85 87 127 86:11:24.1 1816 83 85 87 127 86:11:38.8 1818 83 85 87 127 23:58:01.9 1818 83 85 87 127 23:58:01.9 1818 83 85 87 127 23:58:31.6 1819 83 85 86 126 18:24:18.5 1819 83 85 87 127 23:58:31.6 1819 83 85 87 127 23:58:31.6 1818 83 85 87 127 23:58:31.6 1819 83 85 87 127 23:58:31.6 1821 83 85 87 127 23:58:31.6 1821 83 85 87 127 23:58:31.6 1821 83 85 87 127 23:58:31.6 1821 83 85 87 127 33:58:31.6 1821 83 85 87 127 33:58:31.6 1821 83 85 87 127 33:58:31.6 1821 83 85 87 127 33:58:31.6 1821 83 85 87 127 33:58:31.6 1822 83 85 18 138 18:27:38.6 1822 83 85 18 138 18:27:38.6 1822 83 85 18 138 18:27:38.6 1823 83 85 18 138 18:27:38.6 1824 83 85 18 138 18:27:38.6 1825 83 85 18 138 18:27:38.6 1826 83 85 18 138 18:27:38.6 1823 83 85 18 138 18:27:38.6 1832 83 85 11 131 87:41:44.5 1833 83 85 11 131 87:41:44.5 1834 83 85 11 131 13:85:38:37 1835 83 85 11 131 13:85:38:37 1836 83 85 11 131 13:85:38:37 1837 83 85 11 131 13:85:38:37 1838 83 85 11 131 13:85:38:37 1838 83	46. 489N 153. 492E 28. 2125 16. 345W 46. 297W 152. 536E 3. 885N 122. 523E 9. 443S 143. 678E 42. 654N 142. 979E 33. 823S 172. 353E 38. 762N 72. 354E 3. 187S 135. 165E 9. 7812N 116. 989W 43. 187N 145. 934E 9. 782S 174. 152W 15. 439N 121. 691E 6. 848S 147. 415. 934E 9. 782S 159. 584E 33. 484W 144. 884E 9. 782S 159. 584E 34. 462S 173. 3482W 19. 983N 189. 453W 24. 422N 189. 925W 19. 993N 189. 453W 24. 422N 182. 829E 37. 211N 141. 817E 2. 382N 128. 346E 27. 467S 173. 461W 27. 877S 173. 353E	KURIL ISLANDS FIJI ISLANDS REGION KURIL ISLANDS TALAUD ISLANDS EAST PAPUA NEW GUINEA REGION MOXIMIDO JAPAN REGION SOUTH OF KERMADEC ISLANDS TAJIK SSR WEST IRIAN REGION SOUTHERN NEVADA KURIL ISLANDS TONGA ISLANDS LUZON PPILIPPINE ISLANDS EAST PAPUA NEW GUINEA REGION SOLOMON ISLANDS SOUTH OF HONSHU JAPAN TAILAN HINDU KUSH REGION KURIL ISLANDS PROMONE REGION KURIL ISLANDS PANAMA-COSTA RICA BORDER REGION KURIL ISLANDS PANAMA-COSTA RICA BORDER REGION KURIL ISLANDS PANAMA-COSTA RICA BORDER REGION KURIL ISLANDS MEVILLA GIGEDO ISLANDS REGION KURIL ISLANDS REGION NEV BRITAIN REGION NEW BRITAIN REGION NEW BRITAIN REGION HUZON PHILIPPINE ISLANDS HINDU KUSH REGION MALMAHERA UNKNOUN PHASE NEAR E COAST OF HONSHU JAPAN SOUTH OF MARIANA ISLANDS	33 5.1 8.8 PDE EQ ABDE 593 5.5 8.8 PDE EQ A 33 5.2 4.7 PDE EQ ABDE 64 5.8 4.9 PDE EQ ABDE 15 5.8 4.9 PDE EQ ABDE 167 4.7 8.8 PDE EQ ABDE 187 4.7 8.8 PDE EQ ABDE 187 4.7 8.8 PDE EQ ABDE 23 5.4 8.8 PDE EQ ABDE 23 5.2 4.7 PDE EQ ABDE 23 5.6 5.6 PDE EQ ABDE 23 5.2 8.8 PDE EQ ABDE 23 5.6 8.8 PDE EQ ABDE 24 4.8 8.8 PDE EQ ABDE 25 5.8 8.8 PDE EQ ABDE 26 4.9 8.8 PDE EQ ABDE 27 5.5 6.2 PDE EQ ABDE 28 5.5 6.2 PDE EQ ABDE 28 5.5 6.2 PDE EQ ABDE 29 5.5 6.2 PDE EQ ABDE 20 5.5 6.2 PDE EQ ABDE 20 5.5 6.2 PDE EQ ABDE 21 8 5.5 6.2 PDE EQ ABDE 23 5.7 5.4 PDE EQ ABDE 23 5.7 5.4 PDE EQ ABDE 24 6.8 8.8 PDE EQ ABDE 25 5.7 6.9 PDE EQ ABDE 26 5.9 8.8 PDE EQ A	#2:39-#3:22
1851 83 85 18 138 88:21:31.4 1852 83 85 18 138 18:45:24.8 1853 83 85 19 139 17:48:37.8 1855 83 85 19 139 17:48:57.8 1855 83 85 19 139 17:48:57.8 1855 83 85 28 148 86:89:18.8 1856 83 85 28 148 86:89:18.8 1856 83 85 28 148 86:89:18.8 1856 83 85 21 141 89:55:44.7 1868 83 85 21 141 89:55:44.7 1868 83 85 21 141 89:55:44.7 1868 83 85 21 141 89:55:44.5 1863 83 85 21 141 89:55:44.5 1864 83 85 21 141 18:32:44.5 1865 83 85 21 141 18:32:44.5 1866 83 85 21 141 18:32:44.5 1867 83 85 21 141 18:32:44.5 1868 83 85 22 142 12:15:19:4 1864 83 85 22 142 12:15:19:4 1866 83 85 22 142 17:13:18.9 1867 83 85 22 142 17:13:18.9 1868 83 85 22 142 17:13:18.9 1868 83 85 22 142 17:13:18.9 1869 83 85 22 142 17:13:18.9 1869 83 85 22 142 17:13:18.9 1874 83 85 24 144 87:89:89:87.7 1872 83 85 24 144 87:89:88.8 1871 83 85 24 144 87:39:88.3 1871 83 85 24 144 87:39:88.3 1871 83 85 24 144 87:39:88.3 1871 83 85 26 146 83:85:87.9 1874 83 85 26 146 81:33:34:14 1878 83 85 26 146 81:33:34:14 1879 83 85 26 146 81:33:34:14 1879 83 85 26 146 81:33:34:14 1879 83 85 26 146 13:13:23:2 1881 83 85 26 146 12:27:14.6 1881 83 85 26 146 13:13:23:2 1881 83 85 26 146 13:13:23:2 1881 83 85 26 146 13:13:23:2 1881 83 85 26 146 13:13:23:2 1881 83 85 26 146 13:13:23:2 1881 83 85 26 146 13:13:23:2 1881 83 85 26 146 13:13:23:2 1881 83 85 26 146 13:13:23:2 1889 83 85 26 146 12:27:14.6 1889 83 85 26 146 15:88:88.8 1889 83 85 27 147 22:55:39.2 1899 83 85 26 146 15:28:38.8 1898 83 85 27 147 13:35:43.1 1899 83 85 26 146 15:28:38.8 1899 83 85 26 146 15:28:38.8 1899 83 85 27 147 22:55:39.2 1899 83 85 28 148 86:14:41.1 1899 83 85 28 148 86:14:41.1 1899 83 85 28 148 86:14:41.1	# .592N 121. 739E 8.359S 121. 272E 28.172N 137. 122E 28.172N 137. 122E 51.432N 174. #778V 53.962N 161. 299E 4.7765 153. 524E 53.116N 132. 481W 29.291N 132. 481W 29.291N 132. 482E 25.315S 179. 319E 25.315S 179. 319E 25.315S 179. 319E 25.315N 124. 765E 29.897N 140. 3775E 29.897N 123. 765E 29.898N 123. 765E 21. 356S 178. 294E 37. 984N 14. 945E 24. 431N 14. 295E 24. 431N 14. 295E 24. 431N 14. 295E 25. 685N 162. 481W 26. 914N 142. #38E 21. 356S 178. 541W 24. 431N 143. 295E 24. 431N 143. 295E 24. 431N 143. 295E 24. 431N 148. 298E 21. 356S 178. 541W 22. 417N 187. 838E 48. 378N 139. 936W 48. 378N 139. 839E 48. 138N 139. 897E 22. 417N 139. 839E 48. 139N 139. 987E 48. 284N 138. 998E 27. 1834 139. 988E 48. 139N 139. 988E 48. 139N 139. 988E 48. 139N 139. 988E 48. 139N 139. 988E 49. 931N 131. 4597 48. 193N 139. 1839 49. 193N 139. 1839 55. 52. N 12. 1935 55. 52. N 12. 1935 58. 888 1N 12. 1335	MINAHASSA PENINSULA FLORES ISLAND REGION RYUCYU ISLANDS ANDREANOF ISL. ALEUTIAN ISL. OFF EAST COAST OF KANCHATKA NEW IRELAND REGION ANDPEANOF ISL. ALEUTIAN ISL. SOUTHEAST OF SHIKKU JAPAN SOUTH OF FIJI ISLANDS NEAR E COAST OF HONSHU JAPAN MINDANAO PHILIPPINE ISLANDS NEAR E COAST OF HONSHU JAPAN MINDANAO PHILIPPINE ISLANDS NEAR E COAST OF HONSHU JAPAN YANUATU ISLANDS KURIL ISLANDS KURIL ISLANDS KURIL ISLANDS KURIL ISLANDS KURIL ISLANDS REGION NEAR E COAST OF HONSHU JAPAN NEAR E COAST OF HONSHU JAPAN NEAR E COAST OF HONSHU JAPAN NEAR COAST OF HONSHU JAPAN NEAR W COAST OF HONSHU JAPAN NEAR W COAST OF HONSHU JAPAN NEAR W COAST OF HONSHU JAPAN NOWHEN SEA OF JAPAN SOUTHERN SEA OF JAPAN NOK-AIDO JAPAN REGION EASTERN SEA OF JAPAN SOUTHERN SEA OF JAPAN NOK-AIDO JAPAN REGION EASTERN SEA OF JAPAN SOUTHERN SEA OF JAPAN NOK-AIDO JAPAN REGION EASTERN SEA OF JAPAN SOUTHERN SEA OF JAPAN NOKA-SICO JAPAN REGION EASTERN SEA OF JAPAN SOUTHERN SEA OF JAPAN NOKA-SICO JAPAN REGION EASTERN SEA OF JAPAN NOKA-SICO JAPAN REGION EASTERN SEA OF JAPAN SOUTHERN SEA OF JAPAN NOR WE WOAST OF HONSHU JAPAN NEAP W COAST	131 5.3 8.8 PDE EQ A 289 4.9 8.8 PDE EQ A 33 4.8 8.2 PDE EQ ABD 39 5.8 8.2 PDE EQ ABD 33 4.4 8.8 PDE EQ ABD 233 4.4 8.8 PDE EQ ABD 233 4.9 8.8 PDE EQ ABD 234 4.9 8.8 PDE EQ ABD 558 4.9 8.8 PDE EQ ABD 31 5.1 4.5 PDE EQ ABD 31 5.1 4.5 PDE EQ ABD 33 4.5 8.8 PDE EQ A 33 4.5 8.8 PDE EQ ABD 33 5.2 8.8 PDE EQ ABD 33 5.2 8.8 PDE EQ ABD 33 5.2 8.8 PDE EQ ABD 33 4.4 8.8 PDE EQ ABD	#8:25-#8:36 12 2#855/ 15 18:49-19:8# 12 2#855/ 17 #8:56-18:13 18 2#85/ 18 17:58-18:4# 51 2#855/ 27 #4:46-#5:81 16 2#855/ 21 #6:11-#6:28 18 2#855/ 23 #9:56-18:1# 16 2#855/ 23 #9:58-18:1# 13 2#855/ 23 #9:58-18:1# 13 2#855/ 23 #9:58-18:1# 13 2#855/ 23 #9:58-18:1# 13 2#856/ #1 18:#9-18:2# 12 2#856/ #1 18:#9-18:2# 12 2#856/ #5 11:#48-11:59 12 2#856/ #5 12:#4-17:15 12 2#856/ #5 17:#4-17:15 12 2#856/ #5 17:#4-17:15 12 2#856/ #5 17:#4-17:15 12 2#856/ #5 17:#4-17:15 12 2#856/ #5 17:#4-17:15 14 2#856/ #5 17:14-17:29 16 2#856/ #5 17:14-17:29 16 2#856/ #5 17:15-72#:13 17 2#856/ #5 #8:#9-#8-#8-#8-#8##########

KEY: INFORMATION SOURCES - PDE-NEIS PDECARD, MON-NEIS MONTHLY LIST, ICS-ICS LIST, MEL-HELICORDER, DTH-OTHER EVENT TYPE - EQ-EARTHDUAYE, NX-NUCLEAR EXPLOSION, SX-SCIENTIFIC EXPLOSION, OT-OTHER, UN-UNKNOWN PHASES - A-P. B-PJ, C-S, D-SO, E-T, F-OTHER

EVNT ******ORIGIN TIME****** *NO* YR*MO*DA*JUL*NR*MN*SECS	**COORDINATES***	**************************************	DEP =MAGNI* INF EV ** KM- BDV*SRF SRC TP PE	INTERVALOR TAPE/FILE
1186 83 85 29 149 28:55:57.7 1183 83 85 29 149 28:55:57.7 1183 83 85 29 149 28:55:57.7 1183 83 85 29 149 22:81:49.4 1119 83 85 38 158 82:25:28.5 1111 83 85 38 158 82:25:28.5 1111 83 85 38 158 82:23:24.1 1112 83 85 38 158 12:32:44.1 1113 83 85 38 158 12:32:44.1 1114 83 85 38 158 158 13:31:44.8 1115 83 85 31 151 88:17:33:8 1116 83 85 31 151 88:17:33:41.8 1117 83 85 31 151 88:17:33:41.8 1117 83 85 31 151 88:17:33:41.8 1119 83 85 31 151 151 88:31:41.8 1119 83 85 31 151 151 88:31:41.8 1117 83 85 31 151 151 88:32:41.1 1129 83 85 31 151 21:23:23:3 1122 83 86 81 152 28:23:23:3 1122 83 86 81 152 81:32:32:3 1124 83 86 81 152 81:32:32:3 1127 83 86 81 152 81:33:59.8 1126 83 86 81 152 81:33:59.8 1127 83 86 81 152 18:58:42.8 1129 83 86 81 152 18:58:42.8 1132 83 86 81 152 18:58:42.8 1133 83 86 88 88 888 888:88:28.8 1133 83 86 88 155 13:23:59.8 1131 88 88 88 88 888:88:28.8 1132 83 86 83 154 21:51:50.2 1134 83 86 83 154 21:51:52.4 1137 83 86 83 154 21:51:52.4 1138 83 86 84 155 11:77:48.2 1138 83 86 84 155 11:77:48.3 1139 83 86 84 155 11:77:48.3 1131 83 86 83 154 21:51:52.4 1138 83 86 84 155 11:77:48.3 1139 83 86 84 155 11:77:48.3 1141 83 83 86 84 155 11:77:48.3 1139 83 86 84 155 11:37:48.3 1141 83 83 86 84 155 11:37:48.3 1141 83 83 86 84 155 11:37:48.3 1144 83 82 85 85 156 85:54:48.2	4.227k 122.622E 49.255h 153.375E 36.865h 141.937k 48.442N 125.449W 48.442N 125.449W 48.467l 137.381E 15.5755 174.977W 41.865h 139.163E 49.742h 73.21JE 1.894h 133.381E 5.5755 174.977W 41.865h 139.853I 5.892S 143.572E 4.7685 153.258E 4.7685 153.258E 48.711h 139.473E 48.795h 133.478E 48.795h 134.879E 48.795h 135.171E 13.7624 123.7784 48.888h 872 48.399h 133.132E 15.722N 173.822W 43.919h 88.576E 17.8375 174.614W 48.888h 8888h 872 48.595b 179.585E 51.724N 173.822W 51.785h 179.585E 51.785h 179.585E 51.785h 179.585E 51.785h 179.585E 51.164h 177.1648	CELEBES SEA KURIL ISLANDS HEAR E COAST OF HONSHU JAPAN OFF COAST OF HONTH CALIFORNIA EASTERN SEA OF JAPAN HOP-AIDO JAPAN PEGIUN HEAP W COAST OF HONSHU JAPAN TONGA ISLANDS HOW	28 5.1 8.8 PDE EQ AL 33 5.6 4.7 PDE EQ AL 33 5.6 4.7 PDE EQ AL 33 4.6 8.8 PDE EQ AL 33 4.6 8.8 PDE EQ AL 35 5.4 8.8 PDE EQ AL 36 5.4 8.8 PDE EQ AL 37 5.5 8.8 PDE EQ AL 38 7.5 8.8 PDE EQ AL 38 7.5 8.8 PDE EQ AL 38 7.7 8.8 PDE EQ AL 38 3.4 7 8.8 PDE EQ AL 38 3.4 7 8.8 PDE EQ AL 25 4.9 8.8 PDE EQ AL 25 5.3 8.8 PDE EQ AL 25 5.3 8.8 PDE EQ AL 25 5.3 8.8 PDE EQ AL 25 8.3 8.8 PDE EQ AL 26 8.5 8.8 PDE EQ AL 27 8.8 PDE EQ AL 28 3.3 4.7 8.8 PDE EQ AL 38 3.8 PDE EQ AL 38 3.8 PDE EQ AL 39 3.8 PDE EQ AL 39 3.8 PDE EQ AL 30 3.8 PDE EQ AL 31 8.8 PDE EQ AL 32 3.8 8.8 PDE EQ AL 33 5.8 8.8 PDE EQ AL 34 5.8 8.8 PDE EQ AL 35 5.2 8.8 PDE EQ AL 36 5.2 8.8 PDE EQ AL 36 5.2 8.8 PDE EQ AL 37 5.2 8.8 PDE EQ AL 38 5.8 8.8 PDE EQ AL 38 5.8 8.8 PDE EQ AL 38 5.8 8.8 PDE EQ AL 39 5.8 8.8 PDE EQ AL 30 5.8 8.8 PDE EQ AL 31 5.8 8.8 PDE EQ AL 32 5.8 8.8 PDE EQ AL 33 5.8 8.8 PDE EQ AL 34 5.8 8.8 PDE EQ AL 35 5.8 8.8 PDE EQ AL 36 5.2 8.8 PDE EQ AL 37 5.2 8.8 PDE EQ AL 38 5.4 8.8 PDE EQ AL 38 5.4 8.8 PDE EQ AL 38 5.4 8.8 PDE EQ AL 58 5.4 8	## ## ## ## ## ## ## ## ## ## ## ## ##
1151 83 86 89 168 84:36:58.6 1152 83 85 89 168 18:23:31.9 1153 83 86 89 168 12:49:84.1 1155 83 86 89 168 12:49:84.1 1155 83 86 89 168 12:49:84.1 1155 83 86 89 168 12:49:84.1 1156 83 86 89 168 12:49:84.1 1158 83 86 89 168 18:46:81.8 1158 83 86 89 168 18:46:81.8 1158 83 86 89 168 12:26:56.9 1161 83 85 18 161 82:13:23.8 1162 83 86 81 161 82:13:27.6 1163 83 86 18 161 87:28.7 1164 83 86 18 161 22:37.7 1165 88 88 88 888 888 88:28:38.1 1167 83 86 11 162 28:38:31.1 1167 83 86 11 162 28:38:31.1 1167 83 86 11 162 28:38:31.1	48.814N 139.178E 3.8765 123.6465 48.193N 139.8546 48.193N 139.8546 48.193N 139.8546 48.193N 139.8546 48.193N 139.8546 48.193N 137.7516 51.441N 174.161W 5.9555 122.6586 51.441N 174.161W 5.9555 122.6586 51.441N 174.148W 75.565N 122.6526 48.244N 139.8586 16.275N 145.6786 24.2715 176.283W 8.888N 127.6528 48.244N 127.162.8388 48.488N 127.4888 48.485N 143.8856 44.185N 143.8856 44.185N 143.8856 45.147N 152.6538 48.485N 143.8856 44.189N 92.8176 53.5885 143.6356 24.189N 92.8176 53.5885 143.6366 24.189N 92.8176 25.587N 127.6886 24.189N 92.8176 25.587N 127.6886 24.189N 93.8176 25.587N 127.6886 24.189N 93.8176 25.587N 127.6886 26.189N 129.4781 26.5865N 129.189N 26.485N 129	NEAP W COAST OF HONSHU JAPAN CERAM NEAR W COAST OF HONSHU JAPAN NEAR W COAST OF HONSHU JAPAN NEAR W COAST OF HONSHU JAPAN EASTERN SEA OF JAPAN SOUTHERN NEVADA NOBERANOF ISL. ALEUTIAN ISL. SULAWESI ANDREANOF ISL. ALEUTIAN ISL. LAPTEV SEA SULAWESI ANDREANOF ISL. ALEUTIAN ISL. LAPTEV SEA SOUTH OF FIJI ISLANDS SOUTH OF FIJI ISLANDS UNKNOWN PO MOLUCCA PASSAGE CENTPAL CALIFORNIA SULAWESI NEAP W COAST OF HONSHU JAPAN EASTERN KAZAKH SER SANTA CRUZ ISLANDS REGION HALMAHERA KURIL ISLANDS OFF E COAST OF HONSHU JAPAN FUNIL ISLANDS OFF E COAST OF HONSHU JAPAN FUNIL ISLANDS OFF E COAST OF HONSHU JAPAN PAKISTAN FLORES ISLAND REGION WEST OF MACQUARIE ISLAND OINGMAI PROVINCE CHINA TONGA ISLANDS PROSABLE MARIANDS PO, SO. T. PHILIPPINE ISLANDS REGION UZBEK SER WEST CHILE RISE HONKAIDO JAPAN REGION OFF E COAST OF KAMCHATKA BANDA SEA SOUTH OF MARIANA ISLANDS SOLOMON ISLANDS SOLOMON ISLANDS FYUSHU JAPAN EASTERN SEA OF JAPAN	33 4.9 8.8 PDE EO AB 114 5.2 9.8 PDE EO AB 33 5.4 8.2 PDE EO AB 33 5.3 5.6 PDE EC AB 33 5.8 9.8 PDE EC AB 33 5.8 9.8 PDE EC AB 28 6.2 5.8 PDE EC AB 28 6.2 5.8 PDE EC AB 28 6.2 5.8 PDE EC AB 33 5.7 S.8 PDE EC AB 34 5.4 PDE EC AB 35 5.5 5.2 PDE EC AB 33 5.1 8.8 PDE EC AB 33 5.7 8.8 PDE EC AB 33 5.8 8.8 PDE EC AB 35 5.8 8.8 PDE EC AB 36 5.1 8.8 PDE EC AB 37 5.8 8.8 PDE EC AB 38 5.8 8.8 PDE EC AB 39 5.1 8.8 PDE EC AB 39 5.1 8.8 PDE EC AB 39 5.1 8.8 PDE EC AB 30 5.8 8.8 PDE EC AB 31 5.8 8.8 PDE EC AB 32 5.8 8.8 PDE EC AB	##:38-##:54:54 ##:38-##:54:55 ##:38-##:51:32 ##:38-##:51:32 ##:38-##:51:33 ##:38-

KEY: INFORMATION SOURCES - PDE-NEIS PDECARD, MON-NEIS MONTHLY LIST, ICS-ICS LIST, HEL-HELICORDER, OTH-OTHER EVENT TYPE - EQ-EARTHQUAKE, NX-NUCLEAR EXPLOSION, SX-SCIENTIFIC EXPLOSION, OT-OTHER, UN-UNKNOWN PHASES - A-P, B-PD, C-S, D-SO, E-T, F-OTHER

THV3	***	O1	IGIN	TIME *****	**COORD	INATES ***	**********LOCATION************************************	DEP	*MAGNI*	INF	EV ****	***SAVED****	•••	**STRIP*	••
	••			n× 111 3543			A STATE OF SCHIPTION	Km-	BUT-SKF	SKL	IF PHASE	STPINIERVAL	MNS	TAPE / FIL	. E
12#1	83	86 21	172	#6:25:27.6	41.315N	139.1365	HOKKAIDO JAPAN REGION	13	6.6 6.0	PDE	FO ABOS	#6:27-£7:13			
12#2	93	#6 21	172	#6:46:5#.1	41.385N	139.345E	HOKKAIDO JAPAN REGION		5.4 8.8			\$6:48-87:34	44	22003/ B B/ B	
1286	83	#6 21	172	87:84:22.2	41.358N	139.349E	HOFFAIDO JAPAN REGION		5.5 8.8			#7:#6-#7:52		8 / B	
12#4	8 2	#6 21	172	87:13:51.4	41.328N	139.27#	HOKKAIDO JAPAN REGION		5.4 8.8			87:15-P2:8;		8/8	
1245	83	#6 21	172	#9:12:15.3	41.300N	139.596E	HOFKAIDO JAPAN REGION		4.6 8.8				17	8/ 2	
1286	83	#6 21	172	18:16:85.7	41.38BN	139.421E	HOKKAIDO JAPAN REGION		4.8 8.8				17	8/2	
1287	83	#6 Z1	172	1#:26:28.8	5.4835	151.Ø37E	NEW BRITAIN REGION	185	5.2 8.8	PDE	EQ ABDE		44	8/2	
12#8	83	#E 21	172	18:38:46.7	24.845N	122.853E		33	4.9 8.8	PDE	EO A	18:41-18:53	13	8/8	
1249	83	#6 21	172	11:29:55.5	41.735N	139.1688	TAIWAN REGION HOKKAIDO JAPAN REGION	28	5.2 8.8	PDE	EO ABDE	11:3:-12:18	48	8/8	
1218	83	86 ZI	172	11:32:12.7	7.1225	129.983E	BANDA SEA	111	4.8 8.8	PDE	EQ A		12	2/ 8	
1211	83	86 21	172	11:36:57.5	15.4365	173.451W	TONGA ISLANDS	33	5.1 B.B	PDE	EQ A	11:39-11:51	13	8/ 2	
1212	8.3	96 SI	172	14:31:50 9	24.877N	122.467E	TAIWAN REGION		5.3 4.7			14:74-14:46	13	8/8	
1213	83	#6 Z1	172	14:48:05.5	24.253N	122.488E	TAIWAN REGION	24	5.8 6.3	PDE	EQ A	14:51-15:82	12	8/8	ē
1214	83	#6 Z1	172	15:49:41.9	6.26BN	125.335E	MINDANAO PHILIPPINE ISLANDS	342	8.8 8.8	PDE	EO A	15:57-16:83	12	8/8	
1215	83	86 21	172	16:82:21.3	41.521N	139.488E	HORKAIDO JAPAN REGION		4.8 8.8			16:84-16:28	17	8/ 2	12
210	8.3	80 51	172	17:06:52.1	29.785N	129.3628	RYUKYU ISLANDS		5.9 8.8			17:26-17:56	51	2/2	12
1217	8.3	86 ZZ	173	#4:34:28.8	41.378N	139.277E	HOKKAIDO JAPAN REGION		5.8 B.8			B4:36-25:22	47	8/ 8	e
1218		Bo 22	173	#4:36:34.3	41.847N	139.161E	HOKKAIDO JAPAN REGION		5.3 8.8				47	8/8	Ø
1219	83	#6 22	173	85 39:33.9	41.455N	139.3816	HOFKAIDO JAPAN REGION		4.8 8.5				17	2/2	2
1421	8.3	No 22	1/3	14:32:87.4	5F.315N	142.257E	SAKHALIN ISLAND		4.8 8.8			14:34-14:51	18	8 / €	2
1777	8.3	No 55	1/3	15:28:23.4	41.348N	139.224E	HOKRAIDO JAPAN REGION		5.1 B.B			15:22-16: <i>B</i> a	47	₽. ₽	8
1222	83	80 22	173	28:15:46.8	42.833N	138.974E	EASTERN SEA OF JAPAN		5.3 4.1			28:17-21:83	47	8 / E	ď
1223	6.3	40 22	1/3	21:38:27.8	41.165N	139.335E	HOKKAIDO JAPAN REGION		4.4 8.8			21:48-21:56	17	E / E.	2
1225	83	P6 23	174	#1:45:F3.5	51.366N	179.824W	ANDREANOF 13L. ALEUTIAN ISL.		4.8 8.8			#1:47-#2:#3	17	B / 2	
1225	9.7	PO 23	1/4	11:86.38.8	41.556N	141.9561	HONKAIDO JAPAN REGION		4.9 8.8			11:#2-11:17	16	8/8	
1272	83	4: 33	174	12:25:18.4	51.763N	139.5991	SOUTH OF AUSTRALIA	1.0	5.5 5.6			12:86-12:19	12	8 / B	
12.	83	#0 23	1/4	19:48:23.1	26.7625	177.171W	SOUTH OF FIJI ISLANDS	54	5.8 4.B			19:44-19:55	12	8 / 8	
1270		FO 24	1/5	#2:#9:28.8	4.6695	182.562E	SOUTHERN SUMATERA SOUTHEAST ASIA	33	5.1 B.E	PDE	OA	#2:15-#2:26	12	8/ 8	
1270	# J	86 24	175	#0:06:16 7	21.7239	185.3816	SOUTHEAST ASIA	18	6.1 6.5	PDE I	QF	#7:16-#8:47	98	8 / E	
123	-3	06 24	175	#7.#6:46./	24.19/N	122.4391	TAIWAN REGION		6.8 8.8			#9:#9-#9:21	13	8/8	
1277	93	86 24	175	12:24:14.6	45.155N	146.9216	KURIL ISLANDS		4.9 8.8			89:88-89:24	17	8/ 21	
1227	93	85 24	175	16:40:14.6	34.435N	99.6891	AFGHANISTAN		4.7 8.8				12	2 / 21	
EOF	- 3	#U 24	./5	19.33:28.9	23.993N	122.6988	TAIWAN REGION	33	5.8 8.8	PDE	Q A	19:56-28:85	12	8/81	e

1234 #3 #6 25 176 #2:33:47.3 29.898S 177.351V		33 5.1 #.# PDE EQ A	#2:38-#2:49 ? 2EE649
-1235 83 #6 25 176 1#:#3:17.6 32.8845 178.798V		46 5.6 5.6 PDE EQ A	18:27-12:19 \3 22254-11
1236 83 #6 25 176 11:59:45.6 18.8585 177.5899	FIDE IS ANDS REGION	595 5.1 #.# POE EQ A	12:21-12:13 13 222:4-12
123" 83 #6 25 176 13:5#:25.3 32.277N 137,6838		377 5.1 8.8 FDE EQ ABDE	13:51-14:34 44 28264-14
1238 83 86 25 176 88:48:43.8 1.332N 99,918			
		2#5 4.9 8.8 PDE EQ A	#8:46-#8:57 12 22264-1#
- 1239 83 86 25 176 1+:12:54.9 23.183N 123.343E		16 5.4 8.8 PDE EQ A	14:15-14:27 13 28264-14
- 1242 83 86 25 176 15:84:89.9 22.837S 177.465V		274 5.5 8.8 PDE EQ A	15:87-15:18 12 22254-13
124: 83 #6 25 176 19:1#:#2.2 32.8175 178.679	SOUTH OF CERMADEC ISLANDS	33 5.5 8.8 PDE EQ A	19:14-19:26 13 22264-15
1242 83 86 25 176 19:48:55.3 24.88IN 122.55EE		42 5.5 5.8 PDE EO A	19:44-19:55 12 22064-16
1243 83 86 25 176 23:14:56.8 7.2915 186.8178		56 5.2 8.8 PDE EQ A	23:21-23:32 12 28864-17
1244 53 64 24 123 63 124 36 6 7,2713 180,8176	UMV#	30 5.2 B.B PUL EQ A	
1244 63 #6 26 177 #2:31:16.2 23.##5N 93.7588 1245 03 #6 26 177 #5.1#:2#.1 44.9315 167.2778 1246 93 #6 26 177 13:56:54.5 41.381N 139.3868	BURMA-INDIA BURDER REGION	73 4.7 8.8 PDE EQ A	#2:37-#2:48 12 22254-18
1245 83 #6 26 177 #5:1#:2#.1 44.931S 167.277E	SOUTH ISLAND NEW ZEALAND	33 5.1 8.8 PDE EG A	#5:16-85:27 12 28264-**
- 1246 #3 #6 26 177 13:56:54.5 41.381N 139.3868	HGKKAIDO JAPAN REGION	29 4.9 B.B PDE EC ABD	13:58-14:14 17 28864-28
1247 83 86 26 177 15:54:23.5 17.8555 167.4786	VANUATU ISLANDS	29 4.9 8.8 PDE EC ABD 45 5.2 5.2 PDE EQ A 281 4.9 8.8 PDE EQ ABDE	15:56-16:68 13 28254-21
1248 83 86 26 177 17:45:11,1 28.893N 139.894E	BONTH TOLANDS BESTON	281 4.9 8.8 PDE EQ ABDE	17:45-18:25 41 28865-1
1249 03 86 27 170 88:25:13.2 18.1885 167.666E	MANUATU ANDE		
		58 S.# #.# PDE EQ A	#6:27-#6:39 13 22255- 2
1258 03 86 27 170 14:18:85.5 23.0475 179.9554		548 5.3 8.8 PDE EQ A	14:12-14:24 13 20055- 3
125, 83 #6 28 179 #3:25:16.6 6#.25#N 141,272W		15 5.9 5.4 PDE EQ ABDE	#3:38-R4:48 71 202654
1252 83 #6 28 179 17:45:58.2 21.8155 138.95£	TUAMOTU ARCHIPELAGO REGION	# 5.4 #.# PDE NX A	17:52-18:83 12 22825- 5
1253 83 86 28 179 21:55:46.1 22.7345 174.7824	TONGA ISLANDS REGION	33 5.2 4.9 PDE EQ A	21:59-22:18 12 20265- 6
1254 83 86 28 179 23:59:33.4 43.874N 84.875E		33 5.1 4.5 PDE EQ A	88:26-82:17 12 32255- 7
1255 83 86 29 188 85:11:57.3 48.198N 139.822E			
		33 5.1 #.# PDE EQ ABDE	#5:13-#5:59 47 22465- B
1256 B3 86 29 188 18:55:48.6 36.632H 148.794E		183 4.6 8.8 P.DE EQ ABD	18:56-11:11 16 22265- 9
1257 83 86 29 188 12:51:37.9 6.9165 129.821E		193 4.2 8.8 PDE EQ A	12:55-13:26 12 222:5-12
1256 80 80 88 888 88:88.88.8 8.882N 8 808;	PROBABLE BONIN ISL, PO.SO.T	# #.# #.# HEL EQ BDE	16:14-16:53 4# 20066-11
1259 83 85 29 188 2218:22.6 41.277N 139.358F	HORKAICO JAPAN REGION	33 4.9 #.# PDE EQ ABD	22:28-22:36 17 222:5-12
1262 83 86 38 181 86:53:28.7 18.871N 121.3131 1261 83 86 38 181 186:53:28.7 18.871N 121.3131 1261 83 86 38 181 186:53:28.7 18.871N 121.3131 1262 83 86 38 181 14:11:23:2.2 18.5865 174.841N 1263 83 86 38 181 14:11:23:23.2 18.5865 174.841N 1265 83 86 38 181 18:24:51.5 44.482N 149.555 1266 83 86 38 181 18:24:51.5 44.482N 149.555 1265 83 86 38 181 23:83:58.1 6.9105 155.2281 1265 83 86 38 181 23:83:58.1 6.9105 155.2281 1266 83 87 81 182 23:23:16:51.8 44.428N 149.4985 1266 83 87 88 182 23:23:16:7 44.523N 149.4985 1268 83 87 88 182 36:35:56.8 22.2620 148.4648 1268 83 87 88 182 86:35:55.5 48.256N 149.5865 1278 83 87 88 182 86:35:55.5 5.4 2.265N 149.5865	LUZON PHILIPPINE ISLANDS	75 5 1 # # PDF FO A	#6:56-87.87 12 28765-13
1261 82 86 38 101 13-30-84 6 44 8229 147 0400	MIDTI ICIANCE	42 6 6 6 4 BDE CO ABDE	13:48-14:23 44 28865-14
1367 03 06 36 101 14-11-22 3 10 7106 124 041	RONIL ISLANDS	ar s.o s.a rut tu mout	14:14-14:25 12 20265-14
1262 03 00 3F 101 14:11:23.2 10.2405 1/4.841W	TUNGA .SLANUS	2/2 4.8 B.B FUL EU A	
1203 83 86 38 181 17:37:49.8 2.529N 128.265E	HALMAHERA	49 5.6 5.4 PDE EQ A	17:48-17:52 13 28265-15
1264 83 86 38 181 18:24:51.5 44.482N 149.519E	KURIL ISLANDS	33 5.3 5.1 PDE EQ ABDE	18:26-19:88 43 28266- 1
1255 83 26 38 181 23:83:58.1 6.9105 155,220E	SOLOMON ISLANDS	97 5.1 8.8 PDE EQ ABDE	23:#5-23:46 42 22266- 2
1266 83 86 38 181 23:18:51.8 44.42RN 149 49RF	KURTL IS: ANDS	AR S. I M. P POF FO ARDE	23:28-88:82 43 28866- 2
1267 83 86 38 181 23:42:16 9 44 318N 149 5860	KURTI TS: ANDS	33 5 # # # PDE ED ARDE	23:43-88:26 44 28265- 2
1268 82 87 81 182 82:27:11 7 44 5824 140 4186	PUBLIC TO ANDO	6# 5 4 # # POC EO ABDE	82:24-83:87 44 288.6- 3
1260 00 07 01 102 05.25.11.7 44.3534 149.4101	RUMIL ISLANUS	DD 3.4 B.B FUT TO ABUE	
1203 03 E7 MI 102 MB: 33:50.8 32.202N 149.404E	SOUTH OF HUNSHO SAPAR	70 4.8 B.B PUL EU ABU	#6.36-86'E) 16 272:6- 4
12/# 83 M/ #1 182 #8:53:55.5 44.256N 1/9.58DE	RURIL ISLANDS	33 S.W 4.W PDE EQ ABDE	#8:55-#9: ' 43 22756- 5
		33 5.2 4.7 PDE EQ ABDE	18:48-11:38 43 ZRPEE- E
1272 83 87 81 182 22:83:42.8 36.898N 141,857E	NEAR E COAST OF HONSHU JAPAN	53 5.5 #.# PDE EQ ABDE	22:84-22:47 44 288:6- 7
1273 83 87 81 182 23:32:58.4 39.961N 139.816E	NEAR W COAST OF HONSHU JAPAN	37 4.6 8.8 POE EQ ABD	23:33-23:49 17 222:5- 8
1274 83 #7 #2 183 #5:32:42.6 4#.#684 138.965E		33 5.1 8.8 PDE EQ ABDE	#5:34-86:28 47 22857- 1
	UNKAOWN PO	2 8.8 8.8 HEL EQ 8	#7:18-87:47 3# 2###3- 2
12/6 83 87 82 183 89:34:85.6 5.7244 94.7798		98 5.7 8.8 PDE EQ A	#9:4#-#9:51 12 22#67- 3
1277 83 87 82 183 11:42:21 9 39.385N 64.381E		33 4.9 8.8 PDE EQ A	11:58-12:8: 12 282:7- 4
1278 83 87 83 184 82:49:26.8 28.157N 122.418E		211 6.1 8.8 PDE EQ A	#2:52-#3:#3 12 2P2e7- 5
1279 83 87 83 184 86:11:35.7 5.5895 154.3185		485 5.4 B.B PDE EQ ABOE	#6:12-86 53 42 27767- 6
128# 83 27 #3 184 #7 #5:26.1 41.453N 139.347E	HCFRAIDO JAPAN REGION	33 5.8 8.8 PDE EQ ABOE	#7:#7-07.53 47 20057- 7
1261 85 87 83 184 88:81:22.8 13.8935 167,9546	ZONAJZE UTALINAV	33 4.9 8.8 PDE EQ A	#8:#3-28 14 12 27761- 8
1282 83 87 83 184 17:14:23.2 9.650N 83.644W		33 5.7 6.1 PDE EQ AE	17:23-19:31 129 22/:7- 9
1283 83 87 83 104 17:40:81.4 24.823N 141.812E			17:48-18:26 39 28867- 9
1447 02 04 02 144 14:40:01'4 54'9534 141'BISE	AOTCHUD : 25 MAD2 MEG 104	292 4.5 8.8 PDE EQ ABDE	11.100-10-50 33 TEEP . A

KEY: INFORMATION SOURCES - PDE-NEIS PDECARD, MON-NEIS MONTHLY LIST, ICS-ICS LIST, HEL-HELICORDER, ÖTH-OTHER EVENT TYPE - EO-EARTHQUAYE, NX-NUCLEAR EXPLOSION, SX-SCIENTIFIC EXPLOSION, OT-OTHER, UN-UNKNOWN PHASES - A-P. B-PO, C-S. D-SO, E-T, F-OTHER

THO TREMOTE TIME	**COORDINATES***	LOCATION	DEP "MAGNI" INF EV ***** KM" BDY"SRF SRC TP PHASE	P**SAVED******* *STRIP** S**INTERVAL***MNS TAPE-FIL
1284 83 87 83 184 28:24:41.5 1285 83 87 84 185 28:51:19.6 1287 83 87 84 185 28:51:19.6 1287 83 87 84 185 28:51:19.6 1288 83 87 84 185 28:51:19.6 1288 83 87 85 186 85:51:4 1299 83 87 85 186 85:51:4 1299 83 87 85 186 89:57:38.8 1291 83 87 85 186 11:11:41:1 1292 83 87 85 186 11:49:87:38.8 1293 83 87 85 186 12:55:55:5.8 1294 83 87 85 186 12:55:55:5.8 1294 83 87 85 186 12:55:55:5.8 1295 83 87 86 187 84:54:36:5 1296 83 87 86 187 87:47:12.8 1299 83 87 86 187 87:47:12.8 1299 83 87 86 187 87:47:12.8 1299 83 87 86 187 87:47:12.8 1299 83 87 86 187 87:47:12.8 1299 83 87 86 187 87:19:31.7 1382 88 88 88 88 88 88 88:88:88:88.8 1387 83 87 87 188 13:35:35:55 1383 83 87 87 188 13:35:35:55 1385 83 87 87 188 13:33:55.5 1386 83 87 87 188 13:33:55.5 1386 83 87 87 188 13:33:55.5 1386 83 87 87 188 16:52:5:48.8 1387 83 87 88 189 13:38:12.1 1389 83 87 88 189 13:38:12.1 1318 83 87 87 188 16:52:5:48.8 1387 83 87 88 189 13:38:12.1 1318 83 87 87 188 16:52:5:48.8 1387 83 87 88 189 13:38:12.1 1318 83 87 87 189 189 18:59:13.7 1318 83 87 87 189 15:18:189.8 1311 83 87 87 189 15:18:189.8 1312 83 87 18 191 84:89:57.8 1313 83 87 18 191 84:89:57.8 1314 83 87 18 191 84:89:57.8 1315 83 87 18 191 84:89:57.8 1316 83 87 17 18 191 84:89:57.8 1317 83 87 18 191 84:89:57.8 1318 83 87 18 191 84:89:57.8 1319 83 87 11 192 12:56:28.1 1319 83 87 11 192 12:56:28.1 1322 83 87 12 193 14:556:28.1 1323 83 87 12 193 14:556:28.1 1324 83 87 12 193 14:56:68.8 1327 83 87 12 193 14:56:68.8 1327 83 87 12 193 14:56:88.8 1327 83 87 12 193 14:56:88.8 1327 83 87 12 193 14:56:88.8 1327 83 87 15 196 83:18:83.1 1328 83 87 15 196 83:18:83.1 1329 83 87 15 196 83:18:83.1 1329 83 87 15 196 83:18:83.1 1321 83 87 15 196 83:18:83.1 1322 83 87 15 196 83:18:83.1 1323 83 87 15 196 83:18:83.1 1324 83 87 15 196 83:18:83.1 1325 83 87 15 196 83:18:83.1 1326 83 87 15 196 83:18:83.1 1327 83 87 15 196 83:18:33.1 1328 83 87 15 196 83:18:33.1 1329 83 87 15 196 83:18:33.1 1321 83 87 15 196 83:18:33.1 1322 83 87 15 196 83:18:33.1	24.372N 122.499E 55.569S 27.942W 49.753N 155.619E 4.415N 126.736E 55.696S 123.453W 26.431N 126.947E 24.2015 174.894W 22.5515 171.815E 22.718S 174.8931E 22.713S 171.8422W 23.5515 171.815E 23.718S 171.822W 23.812N 144.759E 11.6325 166.347E 17.715N 98.148E 8.882N 8.882E 22.633S 171.862E 12.2633 173.862E 13.3638 128.434E 18.8253 178.826E 13.37N 53.286E 51.336N 33.298E 13.37N 53.286E 51.336N 32.393E 12.4638 139.45E 61.8338 133.41.895E 61.8338 126.866E 8.895S 126.866E 8.895S 126.866E 8.895S 126.866E 8.895S 126.866E 8.895S 126.866E 8.895S 126.8362E 12.2888N 126.5362E 13.895N 126.5362E	TAIVAN REGION SOUTH SANDUICH ISLANDS REGION KURIL ISLANDS EASTER ISLAND CORDILLERA RYUKYU ISLANDS SOUTH OF TONGA ISLANDS LOVALTY ISLANDS REGION TONGA ISLANDS MARIANA ISLANDS MARIANA ISLANDS SANTA CPUZ ISLANDS NORTHERN SUMATERA UNKNOWN PO, SO, T LOVALTY ISLANDS REGION FIDI ISLANDS REGION LOVALTY ISLANDS PROBABLE MARIANAS PO, SO, T EMST PAPUA NEW GUINEA REGION EUROPEAN USSR EUROPEAN USSR EUROPEAN USSR EUROPEAN USSR EUROPEAN USSR EUROPEAN USSR RYUKYU ISLANDS PROBABLE MARIANAS PO, SO, T HOKKAIDO JAPAN REGION SOUTH SHETLAND ISLANDS KIRCHIZ-XINJIANG BORCER REGION NEAR E COAST OF HONSHU JAPAN SOUTHERN ALASKA HINDA KUSH REGION JAVA TONGA ISLANDS MINDANAO PHILIPPINE ISLANDS WANJATJ ISLANDS MINDANAO PHILIPPINE ISLANDS	65 5.2 8.8 PDE EQ A 33 5.5 4.6 PDE EQ A 48 5.8 8.8 PDE EQ ABDE 188 5.1 8.8 PDE EQ ABDE 188 5.1 8.8 PDE EQ ABDE 185 5.3 6.8 PDE EQ AB 33 5.2 5.1 PDE EQ AB 33 5.5 8.8 PDE EQ AB 33 5.5 8.8 PDE EQ AB 33 5.4 8.8 PDE EQ AB 33 5.7 5.5 PDE EQ A 33 5.7 5.5 PDE EQ A 33 5.7 5.5 PDE EQ AB 33 5.7 8.8 PDE EQ ABDE 33 5.2 8.8 PDE EQ ABDE 33 5.2 4.8 PDE EQ ABDE 33 5.2 8.8 PDE EQ ABDE 33 5.2 8.8 PDE EQ ABDE 33 5.3 8.8 PDE EQ ABDE 33 5.3 8.8 PDE EQ ABDE 33 5.5 4.8 PDE EQ ABDE 33 5.5 4.8 PDE EQ ABDE 33 5.5 4.8 PDE EQ ABDE 33 5.5 8.8 PDE EQ ABDE 33 5.5 8.8 PDE EQ ABDE 33 5.5 8.8 PDE EQ ABDE 33 5.6 8.8 PDE EQ ABDE 33 5.8 8.8 PDE EQ ABDE 34 5.5 8.8 PDE EQ ABDE 35 5.8 8.8 PDE EQ ABDE 36 5.8 8.8 PDE EQ ABDE 37 6.1 6.3 PDE EQ ABDE 38 5.8 8.8 PDE EQ AB	28:27-28:39 11:42-12:81 28:28-22:138 27:88-22:14 12:28-8-8 22:88-22:14 12:28-8-8 22:88-22:14 12:28-8-5 38:57-89:46 58:61-789:46 18:81-18:12 12:28-8-8 11:14-12:11 56:28-8-8 11:52-12:83 12:28-8-9 16:46-17:85 28:28-8-9 16:46-17:85 28:28-8-9 16:46-17:85 28:28-8-9 16:48-8-8:89 17:28-8-8 18:28-8-8:18 18:28-8-8 18:38-8 18:38-8
1335 83 87 16 197 14:89:47. 1336 83 87 16 197 15:58:87. 1337 83 87 16 197 17:19:22. 1338 83 87 16 197 18:19:22. 1348 83 87 16 197 8:17:11. 1339 83 87 16 197 8:15:33:48. 1341 83 87 17 198 81:56:52. 1342 83 87 17 198 81:56:52. 1343 83 87 17 198 81:56:52. 1344 83 87 17 198 81:56:52. 1344 83 87 17 198 86:29:14:46. 1345 83 87 17 198 86:29:14:26. 1346 83 87 17 198 18:21:54. 1347 83 87 17 198 18:21:54. 1348 83 87 17 198 18:21:55. 1348 83 87 17 198 18:21:54. 1352 83 87 17 198 19:81:41:5. 1353 83 87 17 198 19:84:19. 1353 83 87 17 198 19:84:19. 1355 83 87 17 198 19:84:19. 1355 83 87 17 198 19:84:19. 1355 83 87 17 198 19:84:19. 1355 83 87 17 198 19:84:19. 1355 83 87 17 198 19:84:19. 1355 83 87 18 199 81:35:19. 1356 83 87 18 199 81:35:19. 1357 83 87 18 199 87:27:56. 1368 83 87 18 199 87:27:56. 1368 83 87 18 199 87:27:56. 1368 83 87 18 199 87:27:56.	5	E MINAHASSA PENINSULA E FIJI I SLANDS E EASTERN SEA OF JAPAN E MINAHASSA PENINSULA E JAVA E MINAHASSA PENINSULA E KURIL ISLANDS E MINAHASSA PENINSULA E MINAHASSA PENINSULA E MINAHASSA PENINSULA I FIJI I SLANDS REGION E MINAHASSA PENINSULA E BANDA SEA E MINAHASSA PENINSULA I SOUTH OF FIJI ISLANDS E MINAHASSA PENINSULA I MINAHASSA PENINSULA I SOUTH OF FIJI ISLANDS E MINAHASSA PENINSULA I MINAHASSA PENINSULA I SOUTH OF FIJI ISLANDS I MINAHASSA PENINSULA I MINAHASSA PENINSULA I MINAHASSA PENINSULA MINAHASSA P	71 5.2 8.8 POE EO A 246 4.8 9.8 POE EO A 246 4.8 9.8 POE EO A 138 4.9 8.8 POE EO A 54 5.8 8.8 POE EO A 54 5.8 8.8 POE EO A 56 5.8 8.8 POE EO A 56 5.3 4.1 POE EO A 56 5.3 8.8 POE EO A 57 5.3 4.1 POE EO A 58 5.3 8.8 POE EO A 58 5.3 8.8 POE EO A 58 5.8 8.8 POE EO A 59 5.8 8.8 POE EO A 59 5.4 8.8 POE EO A 59 5.4 8.8 POE EO A 59 5.4 8.8 POE EO A 50 5.4 8.8 POE EO A 50 5.4 8.8 POE EO A 50 5.8 8.8 POE EO A	11:28-11:48 13 28871-11 14:13-14:24 12 28871-12 15:52-16:84 13 28371-13 17:28-18:80 49 28371-14 18:21-18:30 12 28271-15 33:34-23:45 12 28271-15 81:37-81:48 12 28271-16 81:37-81:48 12 28271-16 81:37-81:48 12 28271-17 85:26-85:37 12 28271-18 85:26-85:37 12 28271-18 85:26-85:37 12 28271-19 85:48-85:59 12 28271-22 18:18-18:29 12 28271-22 18:18-18:29 12 28271-22 18:18-18:29 12 28271-22 18:18-18:29 12 28271-22 18:18-18:29 12 28271-22 18:18-18:29 12 28271-22 18:18-18:29 12 28271-22 18:13-21:31 13 28272-2 18:13-21:35 13 28272-2 19:88-19:19 12 28272-3 21:13-21:25 13 28272-5 22:14-22:25 12 28272-6 22:18-22:12 13 28272-7 88:13-81:38 12 28272-8 88:112-85:54 43 28272-18 88:13-81:38-18:54 12 28272-18 88:13-81:513-18:24 12 28272-18 18:14-18:55 12 28272-18 18:14-18:55 12 28272-18 18:14-18:55 12 28272-18 18:14-18:55 12 28272-18 18:14-18:55 12 28272-18 18:14-18:55 12 28272-18 18:14-18:55 12 28272-18 18:14-18:55 12 28272-18 18:14-18:55 12 28272-18 18:14-18:55 12 28272-18 18:14-18:55 12 28272-28 18:18-18:26-18 12 28272-18 18:18-18:25 12 28272-18 18:18-18:25 12 28272-18 18:18-18:25 12 28272-25 11:46-11:57 12 28272-28 11:33-21:82 30 28272-25 11:32-11:33 12 28272-28 23:18-22:18 31 22272-28 23:18-22:19 13 22272-3 23:18-22:19 13 22272-3 23:18-23:11 12 22273-3 28:33-21:82 30 28273-3 23:18-23:11 12 22273-3 28:33-21:82 30 28273-3 23:18-23:11 12 22273-3 23:18-23:

KEY: INFORMATION SOURCES - PDE=NEIS PDECARD, MON-NEIS MONTHLY LIST, ICS=ICS LIST, HEL=HELICORDER, OTH=OTHER EVENT TYPE - EQ=EARTHOUAKE, NX-NUCLEAR EXPLOSION, SX=SCIENTIFIC EXPLOSION, OT=OTHER, UN=UNKNOWN PMASES - A=P, B=PO, C=S, D=SO, E=T, F=OTHER

EVNT *	*****ORIGIN TIME***** R*MO*DA*JUL*HR*MN*SECS	COORDINATES	**************************************	DEP "MAGNI" INF EV **** KM" BDY"SRF SRC TP PHAS	ESSAINTERVALSSONNS TAPE-FIL
1384 8: 1386 8: 1386 8: 1388 8: 1388 8: 1388 8: 1399 8: 1399 6: 1399 6: 1399 6: 1399 6: 1399 8: 1399 8: 1399 8: 1399 8: 1482 8: 1482 8: 1483 8: 1483 8: 1484 8	3 #7 22 2#3 #2:25:4#, 3 #7 22 2#3 #2:36:43:4 3 #7 22 2#3 #2:36:43:4 3 #7 22 2#3 #2:36:36:14:4 3 #7 22 2#3 #2:36:36:14:4 3 #7 22 2#3 #2:36:36:14:4 3 #7 22 2#3 #8:56:13:4 3 #7 23 2#4 #8:56:13:4 3 #7 23 2#4 #8:56:13:4 3 #7 23 2#4 #8:56:13:4 3 #7 24 2#5 #5:59:35:25:5 3 #7 24 2#5 #5:59:35:25:5 3 #7 24 2#5 #5:59:35:33:3 3 #7 24 2#5 #5:59:36:44:1 3 #7 24 2#5 #5:59:33:3 3 #7 24 2#5 #5:59:35:33:3 3 #7 24 2#5 #5:59:36:33:3 3 #7 24 2#5 #5:59:36:33:3 3 #7 24 2#5 #5:59:36:33:3 3 #7 24 2#5 #5:59:36:33:3 3 #7 24 2#5 #5:59:36:33:3 3 #7 24 2#5 #5:59:36:33:3 3 #7 24 2#5 #5:59:36:33:3 3 #7 24 2#5 #5:59:36:33:3 3 #7 24 2#5 #5:59:36:33:3 3 #7 24 2#5 #5:59:36:33:3 3 #7 24 2#5 #5:59:36:33:38:38:3 3 #7 24 2#5 #5:59:36:33:38:38:3 3 #7 24 2#5 #5:59:36:33:38:38:3 3 #7 24 2#5 #5:59:36:33:38:38:3 3 #7 24 2#5 #5:59:36:33:38:38:3 3 #7 24 2#5 #5:59:36:36:36:36:36:36:36:36:36:36:36:36:36:	## ## ## ## ## ## ## ## ## ## ## ## ##	MINAHASSA PENINSULA MER EAST COAST OF KAMCHATKA MER WA COAST OF HONSHU JAPAN FIJI ISLANDS REGION MEAP EAST COAST OF KAMCHATKA FLORES ISLANDS REGION MINAHASSA PENINSULA MINAHASSA PENINSULA MINAHASSA PENINSULA MINAHASSA PENINSULA MUFIL ISLANDS MEAP NA COAST OF MOSHU JAPAN MINAHASSA PENINSULA MUFIL ISLANDS MEAR NA COAST OF MOSHU MINAHASSA PENINSULA MURIANA ISLANDS MEAR NA COAST OF MOSHU MINAHASSA PENINSULA	193 5.4 8.8 PDE EQ ABDE 7 6.8 5.7 PDE EQ ABDE 7 6.8 5.7 PDE EQ ABDE 7 6.8 5.7 PDE EQ ABDE 2 5.1 8.8 PDE EQ A 46 5.2 8.8 PDE EQ A 46 5.2 8.8 PDE EQ A 41 4.9 8.8 PDE EQ A 41 4.9 8.8 PDE EQ A 33 5.8 8.8 PDE EQ A 33 5.8 8.8 PDE EQ ABDE 65 4.6 8.8 PDE EQ ABDE 65 4.6 8.8 PDE EQ ABDE 198 5.2 8.8 PDE EQ ABDE 198 5.5 5.4 PDE EQ ABDE 198 5.5 5.4 PDE EQ ABDE 198 5.5 8.8 PDE EQ ABDE 198 5.4 8.8 PDE EQ ABDE 199 5.4 8.8 PDE	## 2:29-#2:4# 12 28873- 8 #2:38-#3:26 49 28873- 8 #2:38-#3:26 49 28873- 8 #2:38-#3:26 49 28873- 8 #2:38-#3:26 49 28873- 1 #8:38-#3:26 49 28873- 1 #8:38-#6:48-#6:51 12 2873-19 #8:38-#6:48-#6:12 12 2873-19 #8:38-#6:12 12 2873-13 #8:56-#1:52 12 2873-13 #8:56-#1:12 12 2873-13 #8:56-#1:12 12 2873-13 #8:56-#1:12 12 2873-13 #8:56-#1:12 12 2873-13 #8:56-#1:13 39 47 28723-15 #8:141-12:23 43 28723-15 #8:141-12:23 43 28723-16 #8:148-#6:42 43 28723-16 #8:12-#6:84 13 28723-17 11:27-11:16 #4 13 28724-19 12:31-12:42 12 2874-5 #8:34-#2:56 13 28724-7 #8:34-#2:56 13 28724-7 #8:34-#2:56 13 28724-7 #8:34-#2:56 13 28724-7 #8:34-#2:56 13 28724-7 #8:34-#2:56 13 28724-7 #8:33-#2:35 14 28724-7 #8:33-#2:35 14 28724-7 #8:33-#2:35 14 28724-7 #8:33-#2:35 14 28724-7 #8:33-#2:35 14 28724-19 #8:12-#9:#2 12 28725-11 #8:12-#9:#2 12 28725-11 #8:13-#9:#2 12 28725-12 #8:14-#9:#2 12 28725-12 #8:14-#9:#2 12 28725-13
14356 98333333333333333333333333333333333333	80 83 215 18:17:44.3 80 83 215 22:29:44.6 80 83 215 23:25:27.6 80 83 215 23:25:27.6 80 83 215 23:25:27.6 80 83 215 23:25:27.6 80 845 217 86:33:24.7 80 85 217 86:23:25:42.7 80 85 217 86:23:23.3 80 85 217 86:23:23.3 80 85 217 86:23:26:49.3 80 85 217 86:23:26:49.3 80 85 218 22:26:49.3 80 86 218 22:26:49.3 80 86 218 22:37:40.5 80 80 218 22:37:25.3 80 80 218 22:37:25.3 80 80 218 22:37:36.6 80 80 22 80:36:37.9 80 80 22 80:36:36.8 80 80 80 80:36.8 80 80 80:36.8 80 80 80:36.8 80 80 80:36.8 80 80 80:36.8 80 80 80:36.8 80 80 80:36.8 80 80 80:36.8 80 80 80:36.8 80 80 80:36.8 80 80 80:36.8 80 80 80:36.8 80 80 80:36.8 80 80 80:36.8 80 80 80:36.8 80 80 80:36.8 80 80 80:36.8 80 80:36.8 80 80:36.8 80 80:36.8 80 80:36.8 80 80:36.8 80 80:36.8 80 80:36.8 80 80:36.8 80 80:36.8 80 80:36.8 80 80:36.8 80 80:36.8 80 80:36.8 80 80:36.8 80 80:36.8 80 80:36.8 80 80 80:36.8 80 80 80:36.8 80 80 80:36.8 80 80 80:36.8 80 80 80:36.8 80 80 80:36.8 80 80 80 80:36.8 80 80 80 80 80 80 80 80 80 80 80 80 80 8	7. 3685 167.793E 22.2885 7. 463E 21.669N 143.848E 21.669N 143.854E 8.2745 128.278E 52.967N 159.778E 17.2165 167.791E 3.5895 62.171M 17.2165 167.9277 7.7275 123.548E 16.134N 93.886W 44.413N 149.188E 44.413N 149.188E 6.5565 138.127E 6.3365 149.236E 31.112N 148.548E 35.466N 138.911E 27.2215 13.332W 7.12215 13.332W 8.3995 132.598E 56.668N 138.911E 27.5215 13.332W 8.3995 132.598E 18.8875 118.4795 18.8875 118.795 18.8875 118.795 18.8875 118.795 18.8875 118.795 18.8875 118.795 18.8875 118.795 14.968N 152.362W 25.955N 123.754E 18.8875 118.775E 14.968N 152.375E 14.968N 152.375E 14.968N 152.375E 14.968N 152.375E 14.968N 152.375E 14.968N 153.189E 12.9985 189.832W 17.2815 123.755E 14.968N 141.167E 14.8915 168.886E 17.896N 145.374E 14.8915 166.886E 17.896N 145.374E 14.8915 166.886E 18.758N 145.374E 14.8915 166.886E	VANUATU ISLANDS S ATLANTIC OCEAN GOOD ANTIPODE MARIANA ISLANDS REGION TIMOR SEA OFF EAST COAST OF KAMCHATKA VANUATU ISLANDS WESTERN BRAZIL VANUATU ISLANDS WESTERN BRAZIL VANUATU ISLANDS MESTERN BRAZIL VANUATU ISLANDS BANDA SEA CHIAPAS MEXICO AEGEAN SEA KURIL ISLANDS BANDA SEA NEW BRITAIN REGION SOUTH OF HONSHU JAPAN HONSHU JAPAN SOUTH ATLANTIC RIDGE BANDA SEA WEST IRIAN REGION SOUTH FOI SILANDS NORTHEAST OF TAIWAN SOUTH FIJIJ ISLANDS NORTHEAST OF TAIWAN SOUTH FIJIJ ISLANDS HOKABJOO JAPAN REGION SOUTH FOI SILANDS HOKABJOO JAPAN REGION TIMOR KURIL ISLANDS SOUTH OF MAPIANA ISLANDS HINAHASSA PENINSULA HEAR EAST COAST OF KAMCHATKA WEAR EAST COAST OF HONSHU JAPAN SOUTH OF MAPIANA ISLANDS HINAHASSA PENINSULA HEAR EAST COAST OF KAMCHATKA VARIANA ISLANDS VANUATU ISLANDS	# 4.2 #.8 PDE NX A 49 5.4 5.5 PDE EQ ABDE 1# 5.1 #.8 PDE EQ ABDE 1# 5.1 #.8 PDE EQ ABDE 3#5 5.4 #.8 PDE EQ ABDE 3#5 5.4 #.8 PDE EQ ABDE 3#5 5.4 8.8 PDE EQ ABDE 23 5.3 5.4 4.7 PDE EQ ABDE 23 5.3 5.7 PDE EQ ABDE 23 5.5 5.3 PDE EQ A 23 5.5 5.2 PDE EQ ABDE 24 4.8 #.8 PDE EQ ABDE 27 4.8 #.8 PDE EQ ABDE 18 6.5 #.8 PDE EQ ABDE 18 6.5 #.8 PDE EQ ABDE 127 4.8 #.8 PDE EQ ABDE 127 4.8 #.8 PDE EQ ABDE 133 5.8 5.3 PDE EQ ABDE 135 5.8 #.8 PDE EQ ABDE 33 5.1 #.8 PDE EQ ABDE 33 5.1 #.8 PDE EQ ABDE 33 5.4 5.4 PDE EQ A 33 5.1 #.8 PDE EQ ABDE 33 5.4 5.4 PDE EQ A 39 5.4 5.4 PDE EQ A 39 5.7 #.8 PDE EQ ABDE 18 5.2 \$.8 #.8 PDE EQ ABDE 18 5.2 \$.8 #.8 PDE EQ ABDE 18 5.3 #.8 PDE EQ ABDE 18 5.4 5.4 PDE EQ ABDE 18 5.5 #.8 PDE EQ ABDE 26 5.2 #.8 PDE EQ ABDE 26 5.2 #.8 PDE EQ ABDE 27 5.8 #.8 PDE EQ ABDE 28 5.8 #.8 PDE EQ ABDE 29 5.8 #.8 PDE EQ ABDE 21 5.5 #.8 PDE EQ ABDE 26 5.2 #.8 PDE EQ ABDE 26 5.2 #.8 PDE EQ ABDE 26 5.2 #.8 PDE EQ ABDE 27 5.8 #.8 PDE EQ ABDE 28 5.8 #.8 PDE EQ ABDE 28 5.8 #.8 PDE EQ ABDE 29 5.8 #.8 PDE EQ ABDE 26 5.2 #.8 PDE EQ ABDE 26 5.2 #.8 PDE EQ ABDE 27 5.4 5.4 #.8 PDE EQ ABDE 28 5.8 #.8 PDE EQ ABDE 28 5.8 #.8 PDE EQ ABDE 29 5.8 #.8 PDE EQ ABDE 21 5.5 #.8 PDE EQ ABDE 21 5.5 #.8 PDE EQ ABDE 23 5.6 #.8 PDE EQ ABDE 24 5.6 #.8 PDE EQ ABDE 25 #.8 #.8 PDE EQ ABDE 26 5.7 #.8 PDE EQ ABDE 27 5.8 #.8 PDE EQ ABDE 28 5.8 #.8 PDE EQ ABDE 29 5.8 #.8 PDE EQ ABDE 29 5.8 #.8 PDE EQ ABDE 20 5.8 #.8 PDE EQ ABDE 21 5.8 #.8 PDE EQ ABDE 23 5.8 #.8 PDE EQ ABDE 23 5.8 #.8 PDE	13:39-13:58

KEY: INFORMATION SOURCES - PDE-NEIS PDECARD, MON-NEIS MONTHLY LIST, ICS-ICS LIST, MEL-HELICORDER, OTH-OTHER EVENT TYPE - EQ-EARTHQUAKE, NX-NUCLEAR EXPLOSION, SX-SCIENTIFIC EXPLOSION, OT-OTHER, UN-UNKNOWN PHASES - A-P. 8-PO. C-S. D-SO. E-T, F-OTHER

								LOCATION								
1484	e3 #	9 10	238	#8:38:	10.2	19.4165	172.779V	TONGA ISLANDS REGION NEW IRELAND REGION NEW IRELAND REGION NOVAWA ZEMLYA PROBABLE MARIAMAS PO SO T NEW IRELAND REGION MOLUCCA PASSAGE PROBABLE BONIN ISL. PO SO T SUMBAWA ISLAND REGION BONIN ISLANDS REGION KURIL ISLANDS REGION KURIL ISLANDS REGION TAJIK-XINJIANG BORDER REGION TAJIK-XINJIANG BORDER REGION TALAUD ISLANDS TALAUD ISLANDS TALAUD ISLANDS TOLOW FOR THE STANDS TOLOW FOR THE STANDS TOLOW TO HONSHU JAPAN VANUATU ISLANDS TANIMBAR ISLANDS REGION CENTRAL MID-ATLANTIC RIDGE NEAR EAST COAST OF KAMCHATKA VANUATU ISLANDS	33	5.3 5.	# PDE	EO	A	#8:33-#8:44	12	28878- 9
1485	13 5	8 18	230	#9:42:	38.1	3.8355	151.293E	NEW IRELAND REGION	**	5 7 8	# PDE	- 50	ABDE	16:86-16:47	42	202/8-18
1486	83 8	8 18	238	16:85	48.3	3./485	151.344L	NEW INCLAND REGION	33	5.0 4	1 905	NY	A	16:17-16:28	12	200 10-11
1400	***		235	*******		# # # # # # # # # # # # # # # # # # #	# ##05	PROBABLE MARIANAS PO SO T	Ĩ	8.8 8	# HF	FO	RDE	17:37-18:26	36	22276-12
1480			235	10.46	49 6	3 7865	151 198F	NEW IRFLAND REGION	33	5.2 4	9 PDE	ĒĞ	Ā	19:47-19:59	13	20278-13
1498			231	14:35	34 3	2.175N	127.8785	MOLUCCA PASSAGE	55	5.2 4.	3 PDE	EQ	Ä	14:38-14:49	12	20076-14
1491	## #			88:88	98.8	# BORN	8 . 8 D A E	PROBABLE BONIN ISL. PO SO T	B	8.8 8.	8 HEL	EQ	BDE	88:54-81:38	37	28278-15
1492	** *	8 82	866	80:89	83.8	8.858N	3.0020	PROBABLE JAPAN PO SO T	8	8.8 8.	# HEL	EQ	BDE	#5:24-86:85	42	28279- 1
1493	83 #	8 28	232	#6:19:	28.2	8.558\$	117.5698	SUMBAWA ISLAND REGION	134	5.5 #.	# PDE	EQ	A	#6:24-26:35	12	20279- 2
1494	83 #	8 28	535	13:58:	31.7	27.86@N	141.887E	BONIN ISLANDS REGION	35	5.8 5.	5 PDE	EO	ABDE	13:29-13:46	38	28879- 3
1495	83 8	8 28	232	14:48:	30.2	5#.865N	156.16PE	KURIL ISLANDS	128	5.8 8.	# POE	EO	ABBE	14:58-15:37	48	22279- 4
1496	83 B	8 25	232	17:15:	86.6	18.1175	179.734W	FIJI ISLANDS REGION	659	4.7 8.	# PDE	EO	•	17:17-17:28	12	20279- 5
1497	83 #	8 Z.E	232	17:15:	24.4	39.328N	73.611E	TAJIK-XINJIANG BORDER REGION	33	4.9 4.	4 PUL	FO	•	17:22-17:34	13	29279- 5
1498	73 #	B 21	233	#6:18:	55.0	4.519N	127.7896	TALAUD ISLANDS	40	5 P E	# PUL		2	#0.21-80.33 #0.27-80.40	13	20270 3
1500	83 8	8 21	233	##: J4:	45.8	3.7147	120.025	ELOCES SEV	641	5 1 8	8 PDE	FO	7	17:25-17:36	12	22270- 0
1588	83 8	0 21	233	10.50	13.0	7.3043	112 6524	FACTER ICIAND DECION	18	5.4 5	5 90.	FO	Â	19:86-19:18	iż	20070-0
1582		0 21	233	22.57	71 4	22 4815	177 1884	SOUTH OF FIGH ISLANDS	169	5.5 #.	# PDE	ĒŌ	Â	23:88-23:12	13	8- 8
15#3	83 8	8 22	234	#4:5R:	11.9	31.888N	141.7185	SOUTH OF HONSHU JAPAN	33	4.5 8.	# PDE	ĒŌ	ABD	#4:59-#5:13	15	ā- ā
1584	83 8	8 22	234	#5:53:	26.8	14.8275	166.9698	VANUATU ISLANDS	112	5.5 8.	# PDE	EQ	ABDE	#5:55-#6:43	49	Ø- B
1525	83 #	B 22	234	£9:#9:	42.3	7.8195	138.811E	TANIMBAR ISLANDS REGION	33	5.8 8.	Ø PDE	EQ	A	89:13-29:24	12	2006P- 1
1586	83 #	8 22	234	12:38:	21.4	7.288N	34.422W	CENTRAL MID-ATLANTIC RIDGE	18	5.1 #.	Ø PDE	EQ	A	12:47-13:27	21	2 - E3535
1597	83 #	8 22	234	12:39:	86.7	53.861N	168.493E	NEAR EAST COAST OF KAMCHATKA	33	5.4 8.	# PDE	EG	ABDE	12:41-13:32	58	28788- 2
15#8	83 #	B 22	234	23:88:	42.5	16.2895	167.9#3E	VANUATU ISLANDS	185	5.1 #.	# PDE	EQ	ABDE	23:82-23:52	51	20065- 3
1589	83 #	8 22	234	23:16:	38.8	3.5875	146.644E	BISMARCK SEA	29	5.4 5.	9 PDE	F 0	ABUE	23:18-52:21	44	2P2EP- 3
1515	83 N	8 23	235	11:15:	16.9	5.91/5	158.9961	NEW BELLWIN KERION	126	5 1 4	8 6UE	50	ABUE	12:18-12:29	12	20000- B
1211	93 #	0 23	235	12:12:	10.9	E 0130	151 8685	NEW BRITAIN RECION	61	6 3 5	B POF	60	ARDE	23:41-00:23	43	20080- 6
1512	87 A	8 24	235	12.36	28 0	48 385N	124 7674	NEAR COAST OF NORTHERN CALLE	3.6	5.5 5.	7 PDE	ΕĞ	ABDE	13:42-14:59	78	20202- 7
1514	83 8	8 24	236	28:19:	51.3	3.6271	122.258E	NEAR EAST COAST OF KAMCHATKA VANUATU ISLANDS BISMARCK SEA NEW BRITAIN REGION BURMA NEW BRITAIN REGION NEAR COAST OF NORTHERN CALIF CELEBES SEA NEW BRITAIN REGION NEW BRITAIN REGION NEW BRITAIN REGION SOUTHERN XINJIANG CHINA KYUSHU JAPAN PROBABLE BONIN ISL. PO SO T NEAR N COAST OF WEST IRIAN TALAUD ISLANDS PROBABLE MARIANAS PO SO T FIJI ISLANDS PROBABLE MARIANAS PO SO T FIJI ISLANDS PROBABLE MARIANAS PO SO T FIJI ISLANDS REGION HINDANAO PHILIPPINE ISLANDS	621	4.8 8.	# PDE	ĒŌ	A	28:22-28:33	12	288888
1515	83 #	8 25	237	88:12:	50.0	5.4275	158.919E	NEW BRITAIN REGION	112	5.1 #.	# PDE	ΕQ	ABDE	##:14-88:56	43	28:689
1516	83 8	8 25	237	#7:#3:	86.3	5.9455	151.008E	NEW BRITAIN REGION	48	5.5 8.	B PDE	ΕQ	ABDE	#7:B4-37:47	44	26665-18
1517	83 &	8 25	237	11:85:	30.8	39.152N	74.852E	SOUTHERN XINJIANG CHINA	33	5.8 4.	9 PDE	ΕO	A	11:13-11:24	2.1	20056-11
1518	83 %	8 25	237	28:23:	32.6	33.498N	131.438E	KYUSHU JAPAN	121	b.1 #.	8 PDE	FO	ABUL	28:25-21:14	58	26762-12
1519	88 8	8 88	888	88:88:	23.8	8.888N	B.888E	PROBABLE BONIN ISL. PO SO T		W.B B.	Ø MEL		BUE	#1:13-#1:52	40	20261- 1
1528	93 7	8 26	238	#2:34:	34.0	2.84/5	139.89/1	TALADO TELANDE	23	5 2 8			7	88.21-20.22	12	20161- 2
1521	83 5	0 26	238	10.20	22.1	3.0841	160 5075	VANHATH ISLANDS	188	5.3 #	# PDF	FÖ	ARDE	18:23-19:17	55	202514
1522	88 8	a 4a	888	88:88	CA A	a agan	A ACAF	PROBABLE MARIANAS PO SO T		B.S B.	Ø HEL	ĒŌ	BDE	#1:49-#2:29	41	28261- 5
1524	83 8	8 27	239	84:14:	28.8	28.6485	178.823V	FIJ1 ISLANDS REGION	628	4.7 8.	Ø PDE	EQ	A	84:16-84:28	13	28281- 6
1525	83 B	B 27	239	11:58:	29.8	8.174N	126.8858	MINDANAO PHILIPPINE ISLANDS	55	5.2 £.	Ø PDE	EQ	A	11:53-12:84	12	22251- 7
1526	83 B	8 27	239	12:31:	48.9	85.517N	90.603E	NORTH OF SEVERNAVA ZEMLVA	18	4.7 4.	8 PDE	EQ	A	12:38-12:49	12	2 <i>02</i> 518
1527	83 B	8 27	239	18:49:	58.3	8.186N	125.962E	FIJI ISLANDS REGION MINDANAO PHILIPPINE ISLANDS NORTH OF SEVERNAVA ZEMLVA MINDANAO PHILIPPINE ISLANDS KURIL ISLANDS KURIL ISLANDS BISHARCK SEA CENTRAL CALIFORNIA PROBABLE KURILS MINDANAO PHILIPPINE ISLANDS	33	5.5 6.	ø PDE	EQ	ABDE	18:58-19:49	68	20081- 9
1528	83 <i>8</i>	8 28	248	11:30:	14.7	46.197N	151.544E	KURIL ISLANDS	75	5.5	e PDE	EO	ABDE	11:31-12:15	45	2728:-18
1529	83 B	8 28	248	13:19:	32.5	44.972N	148.771E	KURIL ISLANDS	33	5.8 8.	# PUL	FO	ABDE	13:28-14:84	45	20281-44
1538	83 #	8 29	241	#5:19:	24.5	3.4385	148.891E	BISMARÇK SEA	33	5.6 5.	# PDE	FO	ABUE	18:14-11:27	43	262212 20072-1
1531	83 5	8 29	241	18:18:	31.5	35.83BN	121.353W	BRODARIC KURZIC	- :	5.3 B.	8 HE.		BUC	18:12-14:54	42	20007- 1
1532	02 6	. 20	241	15:26:	12.6	0 561N	126 2275	MINDANAO BUILIBBINE TELANDE	49	5.3 4	1 PDF	10	¥ .	15:39-15:58	12	8- 8
1933	•	0 29	241	13.30.	13.0	9.3014	120.33/2	HINDRING PRICEPPINE ISLANDS	٠,	•.• •.			-		••	
1624	e2 =		242	# E • E 7 •	e1 =	6 36P¢	138.4855	BANDA SEA SAMDA ISLANDS REGION BURMA-INDIA BORDER REGION SOLDMON ISLANDS SOUTH OF MARIANA ISLANDS UNIMAK ISLAND REGION PROBABLE MAPIANAS SOUTHERN NEVADA NEAP E COAST OF MONSHU JAPAN ANDREANDE ISL. ALEUTIAN ISL.	136	5.1 .	S PDE	£Ο	A	86:88-26:11	12	8- 8
1875	83 8	8 38 8 38	247	#S:50:	16.6	16.6835	172.#75	SAMOA ISLANDS REGION	35	6.1 5.	7 PDE	EQ	ABD	#8:53-£9:11	19	8- 6
1536	13 ž	8 38	242	18:39:	27.3	25.871N	94.7226	BURMA-INDIA BORDER REGION	63	5.6 #	# PDE	EO	A	18:45-12:56 22:17-23:88 19:48-28:15	12	8- 8
1537	83 8	8 3 <i>E</i>	242	22:15:	59.6	9.558\$	158.81BE	SOLOMON ISLANDS	76	5.8 8.	B PDE	ΕQ	ABDE	ZZ:17-23:88	44	B- 8
1538	83 #	8 31	243	19:48:	11.2	12.514N	144.238E	SOUTH OF MARIANA ISLANDS	33	4.3 3.	9 PDE	EQ	ABDE	19:48-28:15	36	8- 8
1539	83 B	8 31	243	55:58:	87.4	53.583N	163.622W	UNIMAK ISLAND REGION	33	3.1 5	B PUE	FU	000	22:23-22:34 #9:21-£9:54	24	8- 8
1548	88 8	8 88	888	88:82:	22.3	#.888N	# . 888E	PROBABLE MARIANAS	-	5 A F	B BDL	NY	A	14:86-14:17	12	8- 8 8- 8
1541	83 /	9 #1	744	14:28:	82.S	37.273N	116.355W	MEAD E COACT OF MONCHII JABAN	68	5.1 6	PDF	ËÔ	ABDE	#3:#6-#3:49	44	0- 8
1542	83 #	9 #2	245	#4:#0	39.4	50.024N	175.5174	ANDREANOF ISL. ALEUTIAN ISL.	1:2	4.7 8.	PDE	ĒĢ	ABD	#4:11-#4:28	18	8- B
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