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FOREWORD

The purpose of rain erosion testing on the High Speed Test Track (HSTT) at Holloman Air Force Base is to study the erosive effects of extended supersonic or hypersonic flight through rain clouds on material samples and components of aerospace systems.

The mechanism used to generate the simulated rain environment is a series of adjustable spray nozzles spaced along the side of the track adjusted to discharge over the flight path. The rain rate is controlled by the water pressure and drop size distribution is controlled by a combination of nozzle adjustments and pressure.

For the selection of materials to be effective, it is essential that the component designer understand the relationship of the simulated rainfield characteristics and those of natural rainfields. This report is the first attempt to quantify these relationships, and provide the designer with a tool to assess the differences.

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PROJECT DELUGE: A PROPOSED MATHEMATICAL MODEL FOR THE HOLLOMAN AFB TEST TRACK SIMULATED RAINFIELD, VOLUME II

1.0 BACKGROUND

1.1 During the performance of Project Deluge it became apparent that a requirement existed for a means to define the Test Track's simulated rainfield capabilities to achieve the parameters most often used by prospective customers to identify their requirements, i.e. accumulated rain rate (ARR) and mass median diameter (MMD).

1.2 Prior research of natural rain has demonstrated that the drop size distribution of raindrops in a rainfield can be defined by an exponential expression of the form (Ref 1)

 $N_d = No EXP (- \Lambda D)$ where: $N_d = number of drops of unit size range per unit volume (N/m³)$ $<math>N_o = Value of N_d$ when D=0D = Drop diamter (mm)

 Λ = Variable which is dependent upon rain rate or intensity of precipitation The relationship of Λ with rain rate is defined by the expression:

 $\Lambda = aR^b$ (mm⁻¹) where: a and b are constants R = rain rate (mm/hr)

1.3 The following methods were used by past experimenters to derive the relationships noted above:

a. Rain rate was determined by rain gauge measurements

b. Drop size distributions were determined by using flour beds to catch rain and calculating the size from sieves or,

c. Measuring the impressions of drops on blotting paper

1.4 A parallel methodology has been employed to formulate the equations for the Test Track simulated rainfield. The major difference is the use of a

mass spectrometer to count and catalog the raindrops by unit size. The substitution of the spectrometer for flour/blotting paper will enhance the accuracy in determining the statistics for the drop size distribution and the resulting equations.

2.0 <u>DATA BASE</u> The data used in this report was accumulated as part of Project Deluge during November 1981. The mass spectrometer used to measure and count the drop sizes was the Illinois Institute of Technology Research Institute (IITRI) Counter specifically designed and constructed for use in the Test Track's simulated rainfield. The use of this counter is discussed in Volume 1, along with the test conditions, raw data formating and data output from the counter. (Reference AD-TR-83-42, The Development of a Holloman High Speed Test Track Heavy Rainfield)

2.1 The basic parameters measured during the experiment were:

a. Rain rate - using standard rain accumulation type gauges (ARR)(in/hr).

b. Drop size and number of each - IITRI Counter. Drop size interval width of 0.6mm, centered around diameters of 0.6, 1.2, 1.8, 2.4, 3.0, 3.6, 4.2, and 4.8mm.

2.2 Data parameters calculated and provided by Project Deluge are:

- a. Number of drops per unit size per cubic meter (N/m^3)
- b. Mass Median Diameter (MMD or Do), (mm).
- c. Liquid Water Content (LWC), (grams/m³).
- d. Equivalent Rain Rate (ERR), (in/hr).

3.0 DATA ANALYSIS AND EVALUATION

3.1 The data of paragraph 2.0 above was subjected to regression analysis to determine the best fit relationships between the different variables. The computer program selected for this function provides the capability of evaluating seven regression models simultaneously for each set of data. The resulting equations providing statistically significant relationships are provided below with the result of previous experimenters.

3.2 The matrix format provides a means of comparing previous experimenters results with natural rainfall and the results obtained with Project Deluge.

TABLE 1

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	Type of	°	۲ [.]	°a	D LMC
Source	Precipitation	(mm ⁻¹ m ⁻³)	(mm ⁻¹)	(mm)	(g/m ⁻³)
Marshall-Palmer	Ottawa - Summer	8000	4.1R ^{-0.21}	0.9R ^{0.21}	0.072R ^{0.88} (a)
Jones (1956)	Thunderstorm	389R ¹ .02	2.5R ⁻ 0.05	1_48R ^{0.05}	0.052R ^{0.97} (b)
Joss et al (1970)	Thunderstorm	1400	3.0R ^{-0.21}	1.22R ^{0.21}	1
Project Deluge (1981) _(c)	Simulated Rainfield	R/(0.117+2.29X10 ⁻⁵ R)	3.346R ^{-0.125}	0.916R ^{0.154}	0.0789R

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NOTES: Rain Rate (R) is in units of mm/hr for all cases.

- (a) Marshall-Palmer as reported by Atlas.(Ref 2)
- Sissenwine developed this relationship by integrating the Jones' equation over volume. (Ref 1) (P
- The mean rain rate measured by gauges over all positions within the simulated rainfield was 29.1 inches per hour. (c)

3.3 Other Established Relationship: Dyer (Ref 1) $Do = 3.67 \Lambda$ -1.00 Project Deluge $Do = 3.6779 \Lambda$

Coefficient of Determination = 0.875

3.3.1 Assessment of Statistical Parameters: Coefficient of Correlation - The hypothesis

$$H_{0}: \qquad \rho = 0$$
$$H_{i}: \qquad \rho > 0$$

was tested using the student's t test, where the test value is defined by:

$$t^{*} = \frac{r \sqrt{N-2}}{\sqrt{1-r^{2}}}$$
 with acceptance criterion

and with rejection criterion

$$t > t$$
, q , N-2

By manipulating the equation for t^* , it can be shown that for a sample size of 81, the following minimum values of r is required for rejection of H₀: at the levels of significant (\triangleleft) indicated.

a	0.10	0.05	6.02	0.01
ρ	0.1820	0.2153	0.2532	0.2783

Based on this test, all equations selected have coefficients of correlation significantly different from zero, i.e. $H_0: \rho = 0$ is rejected implying the relationship between the variables is not a chance occurrence.

3.3.2 Regression Coefficients: The alternate hypothesis

$$H_0: a_1 = 0$$
 $H1: a_1 \neq 0$
or $H_0: b_1 = 0$ $H1: b_1 \neq 0$

was tested using the Student's t test, where the test value is defined by either:

$$t^{*} = \frac{a_{1} - A_{1}}{S_{yx}/Sx}$$
 where $A_{1} = 0$

or: $t = \frac{b_1 - B_1}{S_{yx}/Sx} \sqrt{N-2}$

where $B_1 = 0$

For both cases, the acceptance criterion is

and the rejection criterion is

Based on this test, the regression coefficients for all selected equations are significantly different from zero at all levels of significance.

4.0 <u>COMPARISON OF SIMULATED RAINFIELD MODEL (TENTATIVE) TO KNOWN NATURAL</u> RAINFIELD MODELS.

4.1 <u>vs Marshall-Palmer Model</u>: This model was postulated from observations of summer rains in Ottawa, Canada. Evaluation of each parameter provides the following:

(a) $N_o (mm^{-1}m^{-3})$ - Defined as the number of drops per unit size per unit volume when droplet diameter is zero. Mathematically,

$$\underset{D \neq 0}{\text{Limit } N_{D}} = N_{0}$$

For the Marshall-Palmer model this is defined to be a constant value of 8000. The best relationship derived from the simulated rainfield is:

$$N_0 = R \left[0.117 + 2.288 \times 10^{-5} R \right]^{-1}$$

This equation implies that as rain rate (R) goes from zero to infinity. $\rm N_{_{O}}$ goes from zero to 43706.3. Mathematically given by:

(1) Lim N _o = Lim	R
₽~70 R ~ 0	$0.117 + 2.288 \times 10^{-5}$ R
	0
=	$0.117 + 2.288 \times 10^{-5}$ (0)
Lim N _o = 0	
R ~ 0	
(2) Lim N _o = Lim	R
R→∞ R → ∞	$0.117 + 2.288X 10^{-5} R$
	1
r ∠1m R → ∞	$\frac{0.117}{R}$ + 2.288X10 ⁻⁵
2	۱
	$0 + 2.288 \times 10^{-5}$
Lin N _C =	43706.3

Conclusion: There is a significant difference in both the value and characteristics of $N_{\rm O}$ between the two models.

(b) LAMDA, (Λ , mm⁻¹), A variable which is dependent upon rain rate. Generally stated mathematically as:

$$\mathbf{\Lambda} = \mathbf{a} \mathbf{R}^{\mathbf{D}}$$

For the Marshall-Palmer model the values of a and b are 4.1 and -0.21, respectively. The calculated values of a and b for the simulated rainfield are 3.346 and -0.125 with standard error of estimates of 1.146 and 0.021, respectively. Using these values the indicated confidence limits were calculated to be:

(1-~) %	a = 3.346	b = -0.125
99.00	0.389 ≤ a ≤6.302	-0.1794 ≤b ≤ -0.0711
95.00	1.100 ≤ a ≤ 5.592	-0.1664≦ b≤-0.0841
90.00	1.461 ≤ a ≤ 5.231	-0.1598 ≤ b ≤ -0.0907
Z _c	0.658	4.04

Conclusions: The mathematical relationship between Λ and R is confirmed to be a power function. While there is no statistically significant difference in the scaling parameter (a) values, there is a very significant statistical difference in the shaping parameter's (b) values.

(c) Mass Median Diameter--Do(mm), defined as the droplet - diameter which divides the Liquid Water Content (LWC) into two parts. Generally defined mathematically by the equation:

 $D_0 = a_1 R^{+b}$

For the Marshall-Palmer model, the values of a_1 and b are 0.90 and 0.21, respectively. The calculated values of a_1 and b (simulated rainfield) are 0.916 and 0.154 with standard error of estimates of 1.143 and 0.0206, respectively.

(]-∝)%	a ₁ = 0.916	b = 0.154
99.0	-2.033 ≤ a ₁ ≤ 3.8659	$0.100 \le b \le 0207$
95.0	$-1.324 \leq a_1 \leq 3.157$	0.1132 ≤ b ≤ 0.194
90.0	$-0.9643 \le a_1 \le 2.797$	0.120 ≤ b ≤ 0.1875
Z _c	0.0143	2.734

Again the indicated confidence limits are:

Conclusions: Again, the basic relationship between the variables is confirmed. Also, there is no statistically significant difference in the values of a_1 , but there is a statistically significant difference in the b values. However, note that the critical Z value is less than 3.0, i.e. the Marshall-Palmer's specified b value is less than 3 σ from the value determined for the simulated rainfield.

(d) Overall conclusions: The expected relationships for both Λ and D_0 as a function of rain rate were verified. The significant differences in the regression coefficients and the form of N₀, simply implies that the simulated rainfield as configured does not duplicate the droplet diameter distribution of an Ottawa summer rain.

4.2 <u>Vs Jones (1956) Model</u>: This model was postulated from observations of thunderstorms. Definitions of the parameters are the same as previously provided. Comparing the simulated rainfields to Jones' provides the following information:

(a) $N_0 (mm^{-1}m^{-3})$. Jones values is $N_0 = 3.89R^{1.02}vs$ $N_0 = R [0.117 + 2.288X10^{-5}R]^{-1}$

For both situations $N_0 = f(R)$, with only the form of the relationship differing. If the simulated rainfield data is fitted to a power function to place

it into the same form as the Jones' observation, the following equation results

0.3**959** N_o = 1925.73R

The confidence limits for the regression coefficients for the the power function are:

(1-~)%	a ₁ = 1925.73	b = 0.3959
99.0	$1816.35 \le a \le 2035.11$	$0.2024 \le b \le 0.5893$
95.0	$1842.64 \le a \le 2008.82$	$0.2489 \le b \le 0.5428$
90.0	$185.99 \leq a \leq 1995.47$	0.2725 ≤ b≤ 0.5192
Zc	36.2	8.3

Conclusions: Both of the regression coefficients are statistically significantly different from those identified by Jones. Also, although both situations identify N_0 as a f(R), the forms of the relations are different.

(b) Λ (mm⁻¹). The Jones' determined relationship is $\Lambda \approx 2.5R^{-0.05}$

vs Track's $3.346R^{-0.125}$. The table of confidence limits provided in 4.1(b) above is applicable with the same conclusions. The critical Z scores for a and b, are 0.74 and 3.6, respectively.

(c) $D_{(mm)}$. Jones 1.48R^{0.05} vs Track 0.916R^{0.154}

The Table of Confidence Limits provided in 4.1(c) above is applicable and the conclusions are the same. The critical Z scores are 0.493 and 5.05 for a and b respectively.

(d) Liquid Water Content, $\Sigma LWC (g/m^3)$. Defines the liquid water content of the rainfield per unit volume. Using the Jones' equations Dyer defined the following relationship for LWC as a function of rain rate:

(1) $\sum LWC = 0.052R^{0.97}$

Independently, Atlas, using the Marshall-Palmer equations derived this relationship:

(2) $\sum LWC = 0.072R^{0.88}$

The relationship between these two variables observed during DELUGE is:

(3) $\sum LWC = 0.07889 R^{0.868}$ (Regressing R on $\sum LWC$) The confidence limits for the derived Deluge values are:

(1-α)%	a = 0.07889	h = 0.868
99.0	$-4.458 \le a \le 4.616$	$0.644 \le b \le 1.0926$
95.0	$-3.368 \leq a \leq 3.526$	0.698 ≤ b ≤ 1.0387
90.0	-2.814 ≤a ≤ 2.972	0.725 ≤ b≤ 1.0114
Z Critical	0.004 0.015	0.138 (Atlas) 1.169 (Dyer)

Conclusion: There are no statistically significant differences between the coefficients for the relationships described by the equation:

 $\int LWC = aR^{t}$

As determined by archer Atlas or Dyer and those experimentally determined by Track.

(c) Overall Conclusions: The expected relationships for all of the bandwiders were verified. The differences in the coefficients are most probably attributable to differences in the rain rates observed. Deluge rain rates of 29 inches per hour would be observed only on rare occasion in nature. 4.3 <u>Vs Joss et al (1970)</u>. This model is based or the Marshall-raiser model. The differences observations and is very similar to the Marshall-raiser model. The differences between the two models (Joss vs Marshall) are the values cited for N_o (both constants) and values for scaling parameters for both Λ and D_o as f(R). The comments and conclusions drawn for the comparison of Marshall-Palmer vs Track are applicable.

4.4 <u>Vs Dyer</u>. In her investigation of the other experimenters' work Dyer developed the following relationship between Λ and D_0 , i.e.:

$$D_0 \Lambda = 3.67$$

or $D_2 = 3.67 \Lambda^{-1}$

This relationship appears to be consistent regardless of either the characteristics of the rainfall, i.e., thunderstorm, summer shower, etc. or the investigator. This relationship also holds for the simulated rainfield. The values derived via regressions are:

 $D_0 = 3.6779 \Lambda^{-0.996}$

It can be readily shown that no statistical difference exists between the equations.

4.5 <u>Proposed Model for Mini-Rainfield</u>. The following equations are tentatively suggested for use in describing the mini-rainfield when it is configured as specified in the Project Deluge report:

General Equation:
$$N_0 = N_0 EXP$$
 (~ ΛD)
 $N_0 = R [0.117 + 2.288X10^{-5}R]^{-1}$
 $D_0 = 0.916R^{0.154}$
 $D_0 = 3.6779\Lambda^{-1}$

Rain rate (R) is the measured value for rain rate in mm/hr.

NOTE: (1) An alternate equation for N_0 is: N_0 1925.7R^{0.3959}

(2) \sum LWC may be estimated using this equation: \sum LWC = 0.07889R^{0.868} (g/m³)

RECOMMENDATIONS FOR ADDITIONAL TESTS:

In some instances where statistically significant differences exist, most can be attributed to the fact that the rain rates produced for Deluge, i.e. 25-30 inches per hour are almost never found in nature.

While it has been shown by the experimenters referenced, that the parameters of interest are dependent upon rain rate for the rain rate intensities observed, the sensitivity of the indicated parametric relationships to changes in rain rate has not been identified.

The simulated rainfield at Test Track has the capability to produce a wide spectrum of rain rate intensities to allow such a sensitivity analysis.

The following additional tests are planned:

(1) To extend the applicability of the proposed model from a "specific condition" of the mini-rainfield to a general statement of the mini-rainfield capabilities, additional measurements are required with the water pressure values set at values incrementally spaced between the maximum and minimum allowed.

(2) Because the Knollenberg Probe has the inherent capability to provide better resolution of droplets diameter, that it be used to acquire future data or the data for each set-up be acquired with both the Probe and the Counter.

(3) All of the above tests will be repeated on the Track so that the effects of out-of-door conditions can be incorporated into the final model equations.

The selection of the final model for the Test Track simulated rainfield facility will be delayed until after a comparison is made between the two counters and the above planned tests are completed.

PFERRICE BOCUMENTS+

1. Extremes of Hydrometers at Altitude for MIL-STD-210B, Supplement Drop Size Distributions. AFCRI-TR-73-0008, 9 Jan 73.

2. Shorter Contribution (Optical Extinction by Rainfall), Atlas, David, Journal of Meteorology, Vol 10, Dec 1953.

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