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**TECHNICAL REPORT GL-83-1** 

# CAVITY DETECTION AND DELINEATION RESEARCH

**Report 4** 

# **MICROGRAVIMETRIC SURVEY:** MANATEE SPRINGS SITE. FLORIDA

- by 🚊 wain K. Butler, Charlie B. Whitten, and Fred L. Smith

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Report 4 of a Series

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# Technical Report GL-83-1

# CAVITY DETECTION AND DELINEATION RESEARCH

	Title	Author
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station spacing, and with the long dimension of t	he area approximately per-
pendicular to the known trend of the main cavity.	The main cavity is about
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20. ABSTRACT (Continued).

- contrast of -1.3 g/cm<sup>3</sup>, the gravity anomaly is calculated to be -35  $-\mu$ Gal \* with a width at half maximum of 205 ft.

The microgravimetric survey results clearly indicate a broad negative anomaly coincident with the location and trend of the cavity system across the survey area. The anomaly magnitude and width are consistent with those calculated from the known depth and dimensions of the main cavity. In addition, a small, closed negative anomaly feature, superimposed on the broad negative feature due to the main cavity, satisfactorily delineated a small secondary cavity feature which was discovered and mapped by cave divers.



#### PREFACE

This investigation was performed by personnel of the Earthquake Engineering and Geophysics Division (EEGD) and the Engineering Geology and Rock Mechanics Division (EGRMD), Geotechnical Laboratory (GL), U. S. Army Engineer Waterways Experiment Station (WES), for the Office, Chief of Engineers (OCE), U. S. Army, during the period May 1980 to January 1981. The investigation was part of CWIS Work Unit 31150, "Remote Delineation of Cavities and Discontinuities in Rock." Mr. Paul R. Fisher, OCE, was Technical Monitor for this work.

This report was prepared by Mr. Dwain K. Butler, EEGD, and Messrs. Charlie B. Whitten and Fred L. Smith, EGRMD, under the supervision of Dr. Arley G. Franklin, Chief, EEGD, and Dr. Don C. Banks, Chief, EGRMD, and under the general supervision of Dr. William F. Marcuson III, Chief, GL. In addition, Mr. Rodney N. Walters, EEGD, assisted with the fieldwork. This investigation was closely coordinated with ongoing work under Project MA 61102AT22, Work Unit 002/Q6, "Analytical and Data Processing Techniques for Interpretation of Geophysical Properties."

Special appreciation is expressed to Mr. Sheck Exley and Dr. John Zumrick of the National Speleological Society for assistance in obtaining maps of the Manatee Springs Cave System and arranging for cave divers to conduct detailed cave mapping during the fieldwork. Also, the courtesy extended by MAJ Ellison E. Hardee of the Florida Department of Natural Resources in granting permission to use the Manatee Springs Site and the assistance of CPT Cecil Dykes, Park Superintendent, throughout the fieldwork are gratefully acknowledged.

COL Nelson P. Conover, CE, and COL Tilford C. Creel, CE, were Commanders and Directors of the WES during this investigation. Mr. Fred R. Brown was Technical Director.

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## CONVERSION FACTORS, U. S. CUSTOMARY TO METRIC (SI) UNITS OF MEASUREMENT

U. S. customary units of measurement used in this report may be converted to metric (SI) units as follows:

Multiply	Ву	To Obtain
cubic feet per second	0.02831685	cubic metres per second
feet	0.3048	metres
gallons (U. S. liquid) per minute	3.785412	cubic decimetres per minute
inches	2.54	centimetres
miles (U. S. statute)	1.609347	kilometres

#### CAVITY DETECTION AND DELINEATION RESEARCH

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MICROGRAVIMETRIC SURVEY: MANATEE SPRINGS SITE, FLORIDA

PART I: INTRODUCTION

#### Background

1. Early in 1979, two field test sites in karst regions were selected for assessment and evaluation of geophysical methodologies for cavity detection and delineation. The first of these sites, Medford Cave, in Marion County, Fla., has a relatively shallow (approximately 10 to 30 ft\* (3 to 9 m) to top) air-filled cavity system. Manatee Springs, in Levy County, Fla., the second site, has a deeper (approximately 100 ft (30 m) to top) water-filled cavity system. In addition to the known and mapped cavity systems, the sites have additional geologic complexities in common with other sites in karst regions.

2. The field programs at the two sites were not equivalent in scope. An extensive program was carried out at the Medford site, including 19 geophysical techniques. Report 1 in this series (Butler 1983) discusses the scope of the program at the Medford site in more detail. The program at the Manatee Springs site was more limited in scope due partially to fiscal and time constraints in the research program. Also it was desired to limit the program at the Manatee Springs site to those techniques showing the greatest potential and to techniques for which understanding would be considerably advanced by application at the deeper, water-filled site.

3. This report documents the results of a microgravimetric survey conducted at the Manatee Springs site. Detection and delineation

 <sup>\*</sup> A table of factors for converting U. S. customary units of measurement to metric (SI) units is presented on page 3. The customary unit for gravitational acceleration in microgravimetric surveys is the µGal, which is defined in the text. Also grams per cubic centimetre is used as the unit for density.

of the Manatee Springs cavity system represents a more difficult objective for microgravimetric surveying than Medford Cave, since the density contrast is less than for the air-filled Medford Cave system and the depths are considerably greater than for the Medford Cave system. Maximum residual gravity anomalies at the Medford Cave site were about -70  $\mu$ Gal,\* while calculations based on the known cave system depth and dimensions indicate that the maximum residual gravity anomaly at Manatee Springs should be about -35  $\mu$ Gal (Butler 1980, 1983). Since the error of relative gravity determinations in a microgravimetric survey lies in the range of ±5 to ±10  $\mu$ Gal (Butler 1980) and there is likely a certain level of lithological gravity noise at the Manatee Springs site (as was present at the Medford Cave site (Butler 1983)), the Manatee Springs site will allow further definition of the range of applicability of microgravimetry for geotechnical objectives.

4. In addition to the microgravimetric survey, the following geophysical techniques were also applied at the Manatee Springs site:

- a. Crosshole acoustic surveys.
- b. Crosshole radar surveys.
- <u>c</u>. Acoustic resonance (surface survey with subsurface sonar source).
- d. Spontaneous-potential (SP) survey (surface).

Conventional borehole geophysical logging suites were also obtained in two boreholes at the site. The results of these geophysical investigations at the Manatee Springs site will be the subjects of other reports in this series.

#### Purpose

5. Results of geophysical investigations at the Medford Cave site indicate that microgravimetry is one of the best methods for shallow cavity detection and delineation. The purpose of the present work is to determine the capability of microgravimetry to detect the deeper

\* 1 Gal  $\equiv$  0.01 m/sec<sup>2</sup>.

Manatee Springs cavity system. A secondary purpose of this work is to investigate procedures for determining the local regional gravity field, and the effect on the residual gravity anomaly map, qualitatively and quantitatively, of using different local regional fields.

## Scope

6. Part II of this report will discuss the Manatee Springs site, the site grid system and topographic survey, and the cave system. Part III will discuss the site geology and drilling program. Results of the microgravimetric survey and correlations with known site conditions will be presented in Part IV. Finally, a summary and conclusions will be presented in Part V.

## PART II: THE MANATEE SPRINGS SITE AND CAVE SYSTEM

#### Location and History

7. Manatee Springs is located in the Manatee Springs State Park, which is in Levy County, Fla., about 6 miles west of Chiefland, Fla. The site chosen for the geophysical surveys is near the spring boil,\* which is located about 1500 ft from the east bank of the Suwannee River and 23 miles upstream from the Gulf of Mexico. Figure 1 shows the loca tion of the site on a road map and Figure 2 is a portion of the Manatee Springs 7.5-minute Quadrangle Sheet (U. S. Geological Survey 1954).

8. The spring has been known and used by man for centuries, originally for the crystal-clear drinking water and abundant fish and game and more recently as a popular site for picnicking and swimming. William Bartran, a naturalist, visited the spring in 1774 and provided the first known description of the area. The spring was named for the manatee, an endangered aquatic mammal, which is occasionally seen in the area.

#### Cave System

9. The Manatee Springs cave system extends several miles to the southeast of the spring boil, and through 1980, 10,040 ft had been mapped by the Cave Diving Section of the National Speleological Society. Figure 3 is a map of a portion of the cave system (Exley 1979). Depths to the top of the cave system are in the range of 80 to 100 ft (25 to 30 m) in the vicinity chosen for the geophysical surveys, and the cavity has a cross section 18 ft in height and greater than 40 ft in width at one location in the survey vicinity.

10. Depth and cross section of the cave system are actually quite variable. Also, there are numerous smaller branching cave features along the length of the system. Sinkholes and even smaller

\* Flow from the spring is 81,820 gallons/min.



Figure 1. Road map showing location of the Manatee Springs site





#### Site Grid Layout and Topographic Survey

11. The site chosen for the microgravimetric survey is about midway between the spring boil and the Catfish Hotel (sink), as shown in Figure 4. The long axis of the rectangular topographic survey area was



Figure 3. Map of a portion of the Manatee Springs cave system (Exley 1979)

chosen to be perpendicular to the local trend of the cavity system, and was originally 120 by 400 ft.\* Discovery of a small branching cavity feature by cave divers led to the extension of the survey area to the southeast to include the small cavity feature (Figure 5) and thus accommodate other geophysical tests planned to investigate detectability of the smaller feature.

<sup>\*</sup> Microgravimetric survey area - 100 by 400 ft (Figure 4); topographic survey area - 120 by 400 ft (Figure 6); extended grid area (Figure 5) was not topographically surveyed.







12. A basic site grid dimension of 20 ft was chosen (illustrated in Figure 5). Orientation of the long axis of the survey area was established at N45°E. Grid locations were determined using steel tapes and chaining pins in the usual manner. At each grid location, a 2- by 2- by 6-in. stake was driven flush with the ground surface. At every alternate grid location, an offset reference stake labeled with coordinates was placed. The sense of the survey coordinates is explained in Figures 4 and 5.

13. Elevations of the tops of the 2- by 2-in. stakes were determined to  $\pm 0.01$  ft. Elevations were referenced to a U. S. Geological Survey benchmark (14.00 ft mean sea level (msl))\* located at (47,250) of the survey grid, and the closure error was 0.01 ft. Figure 6 is the topographic map resulting from the elevation survey. Elevation variations over the site are very small, with the <u>maximum</u> elevation difference in the survey grid being approximately 7.0 ft. Except for the semiconcentric contours about the point (120,320) with an elevation of 8.5 ft msl, which reflects the presence of a small sinkhole with center approximately 100 ft beyond the survey area, the topography resembles an inclined plane, decreasing in elevation from 15.6 ft msl at (120,0) to 10.0 ft msl at (0,400). The grid layout, elevation survey, and mapping of the plan location of the cavity system required four man-days (two days for two-man crew).

\* United States Geological Survey Benchmark TT28BB5.



## PART III: GEOLOGY OF THE MANATEE SPRINGS SITE

#### Physiography

14. The site is in Vernon's Terraced Coastal Lowlands Subdivision which is within the Floridian Section of the Coastal Plain Province (Vernon 1951). This section is a very youthful, recently emergent, terraced coastal plain with karst topography. The Terraced Coastal Lowlands is divided into four Pleistocene marine terraces based on elevation above msl--the top elevation of the Coharie is 220 ft msl; the Okefenokee, 150 ft msl; the Wicomico, 100 ft msl; and the Pamlico, 25 ft msl. The Manatee Springs site ranges in elevation from about 10 to 15 ft and is within the Pamlico marine terrace which extends inland from the coast.

15. The Pamlico surface in the Manatee Springs area consists of a broad shelf of Tertiary limestones covered with a thin veneer of Pleistocene sands. Erosion of the Tertiary limestones prior to deposition of the sands created a rugged surface with considerable relief. The ground surface at the site varies less than 2 ft between any of the boreholes, but the top of the limestone varies in depth from 6.6 ft at boring S-1 to 20 ft at boring C-4 (Figure 7). The top of limestone at C-5 varies 9.5 ft vertically in a 2.5-ft horizontal distance (see boring log C-5 in Appendix A).

## Field Exploration Program

16. The number of borings was initially limited to a maximum of 10 by the Florida State Park Service. Only six borings were actually drilled in the initial program. Borings C-1 through C-5 are for testing downhole geophysical instruments, and S-1 was an exploratory boring used to locate borings C-2 through C-5 (Figure 7). The boring locations, diameters, and depths were selected to accommodate test equipment requirements and are discussed in the following paragraphs.

17. Prior to the fieldwork, a map of the Manatee Springs cave,



prepared by cave divers, was used to project the cave outline to the surface at the site, and a 20-ft grid was established. The survey was used to position the borings. The cave outline was used for the drilling of two exploratory borings; the first boring was to penetrate the cave, and the second boring was located approximately 50 ft to the east of boring 1. Divers found boring 1 was drilled through a relatively thin angular section of limestone projecting out from the west cave wall. Boring 2 penetrated the cave near its east wall and was used by the divers as a fixed point for distance and direction measurements to locate borings C-2 through C-5. Boring 2 was later designated boring S-1. Boring 1 was later designated boring C-1 and was to be used for testing. Boring logs are presented in Appendix A.

18. The divers discovered a small side cave extending approximately 600 ft to the southwest of boring C-1. The small cave was more suitable for testing purposes than the larger Manatee Springs cave, so borings C-2 through C-5 were located astride the small cave. The small cave is between borings C-2 and C-3.

19. Borings C-2 and C-3 were cored with 3-in. core barrels and then reamed with a 6- and 6-1/2-in. rock bit, respectively. Boring C-1 was drilled with a 4-in. rock bit, and borings C-4, C-5, and S-1 were drilled with a 6-1/2-in. rock bit.

20. The depth requirement for borings C-1 through C-5 was 165 ft, but the borings were drilled to 170 to 180 ft to allow room for cuttings and possible spalling material to settle. Boring S-1 penetrated the roof of the Manatee Springs cave at a depth of 88.0 ft and was not drilled any deeper. The floor of the cave was at a depth of 105.6 ft.

21. Cavity openings, varying from 1.1 ft in boring C-2 to 7.0 ft in boring C-4, were detected in borings C-2, C-3, and C-4. Divers confirmed that the cavities were not locally connected with the small cave. Boring C-4 was grouted with 155 sacks of cement and reamed to 6-1/2 in. to prevent instruments from getting hung up in the openings. Borings C-2, C-3, and C-4 were not affected by the grouting, indicating that the cavities were probably not locally interconnected.

## Stratigraphy

22. The geologic units encountered were the Pamlico Formation of Pleistocene age and the Ocala, Williston, Inglis, and Avon Park Formations of Eocene age (Figure 8). The Ocala Formation has been renamed the Crystal River Formation in more recent publications, but the term Ocala Formation will be used in this text to maintain continuity with previous work. Lithologic descriptions of the Eocene limestones are limited to the core samples from borings C-2 and C-3 and cuttings from borings S-1, C-1, C-4, and C-5.

Era	System	Series	Group	FORMATION
	Quaternary	Pleistocene		Pamlico Formation
zoic		Eocene	cson	Ocala Formation (Crystal River Formation) Williston Formation
Ceno	iary	Upper	Jacl	Inglis Formation
	Tert	Middle Eocene	Claiborne	Avon Park Formation

# Figure 8. Geologic formations at the Manatee Springs site

## Pamlico Formation

23. The Pleistocene Pamlico Formation is a fine, well-sorted, white to yellow sand, varying from 6.6 to 20 ft thick (Figures 9a and 9b). The sands were deposited in a marine environment and later reworked into a Pleistocene marine terrace as sea level regressed. The terrace sands have been locally reworked by the Suwannee River. Vernon (1951) classified the sands at the site as Suwannee River sandbars. <u>Ocala Formation</u>

24. The Ocala Formation is the upper limestone in the Jackson Group and is typically a soft to moderately hard, white to pink, massive, friable, fossiliferous limestone (Figures 9a and 9b). The fossils consist mainly of gastropods, pelecypods, and foraminifera, and form a coquina in many places. The Ocala is up to 63.4 ft thick at the site. The upper 2 to 10 ft of the Ocala is a weathered zone consisting of limestone fragments in a sandy clay to clay matrix. The Ocala grades conformably downward into the Williston Formation. Changes in fauna and the granular texture of the Williston were used to identify the Ocala-Williston contact.

## Williston Formation

25. The Williston Formation is a grey, moderately hard, friable, fossiliferous limestone and is harder than the Ocala and softer than the underlying Inglis. The abundance of foraminifera gives the Williston a granular appearance. The Williston is up to 41.5 ft thick at the site. The lower 16 to 17 ft is partially silicified with scattered zones of poorly silicified to unsilicified limestone (Figures 9a and 9b). The silicified zone was encountered at a depth of 15.7 ft in boring C-2 and 125 ft in C-3 and can be easily identified in the subsurface by the increased drilling time. The Williston-Inglis contact is conformable and can be recognized by the change from a silicified limestone to a porous dolomitic limestone.

#### Inglis Formation

26. The Inglis Formation is the base of the Jackson Group and is a hard, brown, porous, partially dolomitic limestone with scattered zones of hard, grey, fossiliferous limestone. The dolomitic limestone



a. Section A-A'

Figure 9. Geologic sections (Continued)



Figure 9. (Concluded)

is porous but poorly permeable and ranges from soft and friable to hard and indurated. Laminated dolomitic clasts eroded from the Avon Park are incorporated in the base of the Inglis. The Inglis-Avon Park contact is unconformable with a soil zone on the eroded surface of the Avon Park. The thickness of the Inglis varies from 26.6 ft at boring C-2 to 19.2 ft at boring C-3.

#### Avon Park Formation

27. The Avon Park Formation is a brown to grey, porous, partially dolomitized, moderately hard limestone with carbonaceous plant remains scattered throughout. A few large fossil casts and molds of gastropods and pelecypods and a burrow occur at the top of the unit (Figures 9a and 9b). Only 23 ft of the Avon Park was cored at the site; however, Vernon (1951) described it as being several hundred feet thick in this area.

## Groundwater and Solution Features

28. The high permeability of the Pamlico sand and the Eocene limestones has produced a very poorly developed surface drainage system which has resulted in almost all of the rainfall being absorbed directly into the groundwater system. The groundwater elevation at the site is controlled by the fluctuation of the Suwannee River which was above flood stage at the beginning of drilling operations on 5 May 1980. The water table at the site was at 7.6 ft on 5 May 1980. As the flooding receded, the water table dropped to 9.5 ft on 15 May 1980 and 11.5 ft on 16 June 1980.

29. The groundwater is highly charged with organic and carbonic acids and plays an important part in the development of the caverns in the area (Vernon 1951). The permeability of the Ocala and Williston limestones allows for rather free movement of groundwater through these units. The poor permeability of the Inglis has restricted groundwater flow, thus concentrating the groundwater in the overlying more permeable lower Williston. This concentration of goundwater in the Williston has resulted in the development of a large complex cavern system and

probably in the irregular silicification of the lower Williston. Cave divers have explored over 10,000 ft of the Manatee Springs cave and report that it probably extends much further. The Manatee Springs cave has a variable height and width, and has numerous smaller branching caves. Data from the divers and boring S-1 show Manatee Springs cave is 17.6 ft high and over 40 ft wide at boring S-1. The small cave is approximately 4 ft high and 8 ft wide between borings C-2 and C-3, but the height and width are very variable along the length of the cave.

30. The depth of the Manatee Springs cave decreases from 105.6 at boring S-1 to approximately 55 ft (relative to the borings) at the Manatee Springs boil, a distance of approximately 250 ft to the north. The measured surface discharge from the spring has varied from 110 ft<sup>3</sup>/sec to 238 ft<sup>3</sup>/sec (Rosenau et al. 1977). The spring pool is approximately 100 ft wide and 55 ft deep.

31. The only other solution features detected during the drilling operations were in the Ocala. Boring S-1 intersected a vertical solution pipe 1 to 1.5 ft in diameter. The solution pipe collapsed as a drill rig was being moved by the boring, creating a hole approximately 5 ft in diameter at the surface. Boring S-1 and the funnel-shaped collapsed area were filled with a sand-cement mixture to prevent further deterioration or enlargement by rainfall. Some minor weathering along fractures was found in the core samples, but it is only recognizable by the discoloration of the limestone.

32. The collapse of caverns has created several sinks in the Manatee Springs area. Three sinks, varying from approximately 75 to 200 ft in diameter, are located along the course of the main underground cave, and are within 500 ft of the study site. Catfish Cabin, the largest of the sinks, is approximately 75 ft southeast of boring C-4 and is connected to the Manatee Springs cave by an opening up to 10 ft in diameter.

#### PART IV: MICROGRAVIMETRIC SURVEY

#### Background

33. As noted in Part I, detection and delineation of the Manatee Springs cave system by microgravimetry represents a more difficult problem than the delineation of the Medford Cave system. The difficulty arises from two aspects of the situation: (a) the maximum value of the gravity anomaly at Manatee Springs should be smaller by a factor of two (-35  $\mu$ Gal compared to -70  $\mu$ Gal); and (b) the greater depth of the Manatee Springs cave system will cause the spatial wavelength or width of the gravity anomaly to be much larger than the gravity anomalies at the Medford site (see Butler 1980, 1983). Low amplitude, long wavelength anomalies are difficult to discern in cases where the dimensions of the survey area are comparable to the wavelength.

34. For site surveys with areas comparable to those at the Medford and Manatee Springs sites, it is possible that long wavelength anomalies will be interpreted as part of the local regional component. This fact can be illustrated by considering a model calculation of the anomaly caused by a structure with dimensions and density contrast appropriate for the Manatee Springs cave system. The depths and dimensions of the main cave near boring S-1 (see Part III) are used to model the cave as a long horizontal cylinder with rectangular cross section and a density contrast of -1.3  $g/cm^3$  (simulating a water-filled cavity in limestone of density 2.3 g/cm<sup>3</sup>). The gravity anomaly is calculated along a 400-ftlong profile line perpendicular to the cylinder axis, with the 200-ft position directly above the axis. The geometry is illustrated and the gravity anomaly profile is shown in Figure 10. Magnitude of the anomaly is -36  $\mu$ Gal; and the wavelength, defined to be the width at half-maximum (Butler 1980), is 205 ft. However, the anomaly still has an amplitude of  $-7 \mu$ Gal at the ends of the profile line.

35. Unless there is considerable lithological noise (Butler 1983) or there are other superimposed gravity anomalies, an anomaly such as shown in Figure 10 should be detected by a carefully conducted





microgravimetric survey. The primary difficulty is that the size and depth of the feature causing the gravity anomaly may be underestimated due to clipping of the gravity profile by the selected regional field or because profile lines are too short to define the complete anomaly.

## Survey and Data Processing Procedures

## Station layout and survey programs

36. Stations occupied during the microgravimetric survey are

illustrated in Figure 11. The survey consisted of 126 stations on the 20 ft grid with about 30 percent of the stations reoccupied at least once during the survey. Station (0,200) was selected as base station for the survey. All stations occupied between successive reoccupations of the base station are referred to as a program. Programs were designed to occupy stations in a "zigzag" or "leapfrog" fashion in order to distribute random errors and to prevent any cumulative errors from combining to produce fictitious anomalies or elongated anomalies such as can result from a long, continuous program of station occupations. Three typical programs (E, I, and N) are indicated by the dashed lines in Figure 11. Each program was planned to include at least one repeat station occupation and to be short enough that the base station can be reoccupied in less than 1 hr. In fact, the mean program duration was 40 min. The field gravity measurements required a total of 40 manhr (20 hr each for a two-man crew).

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37. The survey was conducted with LaCoste and Romberg Model D-25 gravimeter. This instrument has a range of about 200 mGal and the smallest dial division equals approximately 1  $\mu$ Gal.\* All field measurements were obtained using the top-mounted galvanometer (capacitance electronic readout option) rather than the eyepiece to obtain a null indication. With this procedure, the reading sensitivity is 1 to 2  $\mu$ Gal. A concave leveling baseplate was used. At each station, the baseplate was pressed to make contact with the elevation survey point and the bull's-eye level bubble centered. The gravimeter was placed on the baseplate and shifted to obtain approximate level prior to final precision level. Microgravimetric survey procedures are discussed in detail in Butler (1980).

#### Base station drift

38. Reoccupations of the base station are used to correct the gravity data for time variations caused by gravity meter drift and earth-tide gravity variations, with the correction referred to as the drift correction. A drift curve for the survey is shown in Figure 12.

\* The meter factor for Model D-25 is 1.1980.



- Gravity Station
- Reoccupied Station and Number of Occupations

- 126 Stations
- 37 Reoccupied Stations (29 percent)
- <u>I</u>\_Program I

27



It is assumed that the drift was the same at all stations in the survey grid and linear interpolation is used to compute the correction for stations in each program.

## Data corrections

39. Required corrections to gravity data are thoroughly discussed in Butler (1980) and will only be summarized here. Briefly, the data corrections are necessary because of latitude and elevation differences between the gravity stations and the effects of surrounding topography. The corrections applied to the drift-corrected data for the Manatee Springs microgravimetric survey are summarized below:

- <u>a</u>. Latitude correction. The latitude correction to be applied to each station gravity value is given by  $\pm 0.81$ × sin  $2\phi \times \Delta L \ \mu Gal/m$ , where  $\phi$  is the latitude of the site ( $\phi \simeq 29.5^{\circ}$ ) and  $\Delta L$  is the north-south distance in metres from a reference station. Using station (0,0) as the reference station,  $\Delta L = \Delta Y \times \cos 45^{\circ}$ , where  $\Delta Y$  is the distance along the NE-SW grid lines from a line passing due east through the survey grid from (0,0). The positive sign is used if the station is south of the line, and the negative sign is used if the station is north of the line. The final expression is  $\pm 0.49 \ \Delta Y$  for  $\Delta Y$  in metres or  $\pm 0.15 \ \Delta Y$  for  $\Delta Y$  in feet, with the correction given in  $\mu$ Gal.
- <u>b</u>. <u>Free-air correction</u>. Three corrections (free-air, Bouguer, and terrain) are used to compensate for elevation differences between stations. The free-air correction accounts for the normal free-air vertical gravity gradient and is given by  $\pm 308.55 \times \Delta h \ \mu Gal$  for  $\Delta h$  in metres or  $\pm 94.04 \times \Delta h \ \mu Gal$  for  $\Delta h$  in feet, where  $\Delta h$  is elevation difference of a station relative to a reference elevation. For the Manatee Springs site, msl is used as a reference elevation.
- <u>c</u>. Bouguer correction. The Bouguer correction accounts for the fact that gravity values in a survey are affected by differing masses of material beneath the stations due solely to elevation differences. The correction is calculated using  $\pm 41.91 \times \rho \times \Delta h \ \mu Gal$  for h in metres or  $\pm 12.77 \times \rho \times \Delta h \ \mu Gal$  for  $\Delta h$  in feet, where  $\rho$  is the density of the near-surface material (1.8 g/cm<sup>3</sup> was used for the Manatee Springs site).
- d. <u>Terrain correction</u>. The terrain correction compensates station gravity values for nearby topographic variations which are not included in the Bouguer corrections.

Because of the nature of the problem, the terrain correction is always added (Butler 1980, Telford et al. 1976). Terrain corrections were determined using a terrain template and correction curves as described in Butler (1980). As a consequence of the small topographic variation in the survey area, no station required a correction greater than 5  $\mu$ Gal, and the majority of the stations required corrections <2 µGal. The spring boil sink and the Catfish Cabin sink were not of concern since they were a sufficient distance from the survey area; effects due to these sinks would be small (due to distance, low density contrast due to being water-filled, and the relatively small elevation differences) and tend to be roughly equal over the survey area. Stations along a line from (100,260) to (100,340) had terrain corrections of <5 µGal due to the small water-filled sink centered about 100 ft to the southeast of this area. Generally, the effects of terrain corrections as small as for this survey area can be ignored.

#### Results

#### Bouguer gravity map

40. Station gravity values which have been correced for drift, latitude differences, and elevation differences (free-air, Bouguer, and terrain corrections) are called the Bouguer gravity values for the station. Figure 13 is the resulting Bouguer gravity contour map for the Manatee Springs site. Two prominent features or trends are evident from examinations of the Bouguer gravity values along profile lines: (a) along the NW-SE profile lines, the gravity values decrease from SE to NW; and (b) along the SW-NE profile lines, the gravity values decrease from the SW reaching minimum values near the center and then increase to the NE.

41. The Bouguer gravity values reflect density variations in the subsurface caused by geologic structures or other lateral variations with varying scales. An important step in analyzing the results of a gravity survey is to separate out a "regional" component of the gravity field, presumably caused by deeper sources or structures, leaving a "residual" component which consists primarily of anomalies caused by structures of interest. Clearly the scale of the regional will depend



on the scale of the survey (Telford et al. 1976, Butler 1980). For small-scale surveys, microgravimetric surveys, like the Medford Cave and the Manatee Springs survey, the regional component can usually be satisfactorily approximated as a constant or as a planar surface. This procedure is called the regional-residual separation, and two methods of accomplishing it will be investigated here: (a) a planar regional will be determined by inspection; and (b) a planar regional, as well as regionals of higher orders, will be determined by polynomial surface fitting of the data.

#### Bouguer anomaly (residual) map; planar regional field determined by inspection

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42. The observed general trend of the Bouguer gravity values to decrease along profile lines from SE to NW, discussed previously, is used to specify a planar regional variation. Using mean values of Bouguer gravity along the bounding SW-NE survey lines, a linear variation of 0.22  $\mu$ Gal/ft (0.72  $\mu$ Gal/m) is determined, decreasing SE to NW. Clearly this procedure is subjective in nature; the effect of determining a best-fit planar regional field in a least-squares sense will be examined later.

43. Subtracting this planar regional field from the Bouguer gravity map (Figure 13) gives the residual or Bouguer anomaly map shown in Figure 14. Several anomaly features appear in the residual gravity map. The most prominent feature is a broad negative anomaly region (A) trending from SW to NE across the center of the map. Obvious correlation of this negative feature with the known location of the main cavity is demonstrated by the dashed outline of the cavity system in Figure 14. The apparent width of the negative feature is considerably greater than the width of the cavity itself, as expected from examination of Figure 10. Also, the magnitude of the negative anomaly, -20 to -30  $\mu$ Gal, is similar to that predicted by the model study of Figure 10; however, if the concentric positive anomaly values (particularly evident to the SW of the negative feature) are considered part of the anomaly caused by the cavity, the anomaly magnitude becomes -50 to -60  $\mu$ Gal. This larger





anomaly magnitude is more consistent with the observation that anomalies associated with cavities in karst regions are frequently larger by factors of two than that predicted on the basis of the cavity dimensions alone (Butler 1980, Neumann 1977). Two areas within the broad negative region (A' and A") to the NE of the mapped cavity location exhibit closure, indicating anomalous features in addition to the known cavity system.

44. Other features of the anomaly map (Figure 14) are the positive area in the SW area of the map which attains 40  $\mu$ Gal magnitude (B). A positive anomaly area exhibiting closure (C) is adjacent to the central negative feature to the NE and attains magnitude slightly in excess of 10  $\mu$ Gal. Another positive anomaly (D) occurs in the NE corner, attaining 30  $\mu$ Gal magnitude at (80,400). A negative region (E) with a magnitude of approximately -10  $\mu$ Gal separates anomalies C and D.

45. A negative feature (F) occurs as a small southwestern trending extension of the central negative feature (A), with a magnitude of approximately -10  $\mu$ Gal. This feature is interesting due to its very close correlation with the small secondary cavity feature mapped in this area (Figure 5) shown dashed in Figure 14. All of the other anomaly features can be expected to change only slightly in detail with a different choice for a planar regional; however, a narrow, elongated anomaly such as F could be altered considerably by the choice of planar regional. This association of F with the small cavity feature will be discussed further after presentation of the regionals determined by polynomial surface fitting.

## Bouguer anomaly (residual) map; regional field determined by polynomial surface fitting

46. A more objective procedure for determining the regional cavity field is by fitting a low-order polynomial surface to the Bouguer gravity values (Coons, Woolard, and Hershey 1967, Nettleton 1971). The Bouguer gravity values  $g_B$  are considered to be a function of the station coordinates (X,Y), i.e.,  $g_B(X,Y)$ . For example, the equation for a planar surface can be written as  $g_B(X,Y) = C_1X + C_2Y + C_3$ , where the constants  $C_1$ ,  $C_2$ , and  $C_3$  are determined so that the plane best fits the gravity data in a least-squares sense. In principle, surfaces of increasingly higher order can be fit to the Bouguer gravity data. Subtracting any polynomial surface from the Bouguer data results in residual gravity values relative to that surface; residual gravity maps will correspond to shallower anomalous features as higher order polynomial surfaces are subtracted.

47. Figure 15 contains polynomial surface fits to the Bouguer gravity data (Figure 13) through fourth order. The plane determined by inspection dips from right to left through the grid compared to the best-fit planar surface which dips from lower right to upper left. It is worthy of note, however, that the right-to-left gradient through the grid is the <u>same</u> for both surfaces (~0.23  $\mu$ Gal/ft). The second-order surface (as well as the third and fourth) show the second prominent feature discussed earlier in regard to the Bouguer gravity values, i.e., that the Bouguer gravity values along SW-NE profile lines decrease from the SW, reaching minimum values near the center, and then increase to the NE. Higher order polynomial surfaces (not shown) will account for increasingly smaller scale features of the Bouguer gravity values, until ultimately a surface is generated for which the residuals will approach zero.

48. Residual ε avity maps will be presented only for the firstand second-order polynomial regionals of Figure 15. The residual gravity map referenced to the first-order polynomial regional is given in Figure 16. Comparison of Figures 14 and 16 reveals that all the major details are similar. All of the lettered anomalies in Figure 14 can be identified in Figure 16. Of course, absolute values of anomalies have changed, but the important relative anomaly values are the same. The most obvious visual difference between Figures 14 and 16 is the shifting of the zero contour separating the anomalous regions A anu F from B in Figure 14 further to the SW in Figure 16. Also, the pronounced "outlining" of the small secondary cavity feature by a continuous zero contour line (see anomaly F in Figure 14) is no longer present in Figure 16, although the presence of a relative negative anomaly apparently





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Figure 16. First-order residual gravity anomaly map

associated with the feature is still evident. Thus, observations made previously regarding Figure 14 and its correlation with the known cavity system are unaffected by the choice of the planar regional field (by inspection or by least-squares surface fitting).

49. Using the second-order polynomial surface of Figure 15 as a regional field results in the residual map shown in Figure 17. A comparison of Figure 17 with Figures 14 and 16 shows that anomalies A, B, and F are removed or significantly altered in character by subtracting the second-order regional field. The obvious interpretation is that the long wavelength anomaly caused by the known cavity system has been accounted for by the second-order regional. Within the area occupied by anomaly A (including A' and A") in Figure 14, negative anomalies apparently corresponding to A' and A" are present in Figure 17. Also, negative anomalies labeled A"' and A"" are identified in an area of otherwise very small negative values in Figure 17. Presumably, these four negative anomaly features are caused by conditions shallower than the main cavity system. For example, as noted in Part III, a wheel of the drill rig fell in a vertical solution pipe about 1.5 ft in diameter and 80 ft deep, creating a hole about 5 ft in diameter at the surface, while drilling boring S-1. Anomaly A"" in Figure 17 seems to be caused by this feature. This association suggests similar causes for anomalies such as E' and E"" in Figure 17.

#### Verification Drilling

50. Because of time and fiscal constraints, only a very limited number of verification borings was possible. These borings, E-1 through E-6 shown in Figure 17, were placed subsequent to the microgravimetric survey and several cross-borehole surveys in the vicinity of the small secondary cavity feature. These borings were not logged in detail; cuttings were monitored and ease of drilling noted to indicate soft or hard zones or tool drops.

51. Simplified boring logs for these verification borings are given in Figure 18. Boring E-1 was placed to verify the mapped location



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> Second-order residual gavity anomaly map Figure 17.

•5 S1			E-2		E-3
â	lty sand w/original aterial	0-1	Silty sand w/original material	0-1	Silty sand w/orig material
-15 Sil	lty sand w/clay	1-8	White, silty sand	1-14	White. silty sand
6.5 Re:	sidual clay w/LS	8-9	Residual clay w/LS	14-15.5	Residual clay w/L
105 142	odules	0 175	nodules	16 6	nodules
	mestone, sort, nite	677-6	Limestone, Soit, White		Limescone, soit, white
-107 Car	vity	114-114.2	Tool drop		
E-1	*		E-5		E-6
.5 Sti	lty sand w/original aterial	0-10	Silty sand w/original material	0-1	Silty sand w/or g. material
-5.5 Whi	ite, silty sand	10-13	Sandy clay s/LS nodules	1-13	White, silty sand
-50 Lin	mestone, soft, ofte	13-96 80 5-05 0	Limestone, soft, white	13-16	Residual clay w/L
			CAVILY	101 101	
				CU1-01	Limestone, white, soft
				90-105	Noticeably softer
					material, possibl
					sand- or clay-
					filled pocket
				101-101.3	Tool drop
	Figure 18.	Simplified bo (depth j	ring logs for verification   .ntervals in feet)	borings	

and depth of the secondary cavity feature between borings C-2 and C-3. Figure 19 shows the location, depths, and geometry of the secondary cavity feature as confirmed by boring E-1 and cave divers. The secondary cavity feature is somewhat shallower and larger than originally thought. Boring E-2 was placed to investigate an anomalous response indication on a crosshole electromagnetic survey between borings C-3 and C-4. Borings E-5 and E-6 were placed to verify the mapped location of the main cavity as it crosses under the northeastern side of the survey grid. Finally, borings E-3 and E-4 were placed to investigate the cause of anomalies C and E (Figure 14).

52. Boring E-5, placed at (10,210), intercepted the main cavity, and the dashed cavity location in Figure 14 reflects the proper borehole location relative to the cavity as verified by cave divers. Boring E-6 was originally planned to investigate anomaly A' at (0,260) and the possibility that the main cavity might turn and pass out of the survey area at this location, but it was relocated due to presence of large trees at the desired location. As seen from Figures 14 and 17, boring E-6, as finally located, is at the periphery of anomaly A'. From the boring logs in Figure 18, two features are apparent: (a) 3 ft of residual clay was intercepted by boring E-6 but not by boring E-5; and (b) a soft zone was intercepted in boring E-6 from approximately 90- to 105-ft depth, possibly a clay- or sand-filled pocket. Feature (a) above suggests that anomaly A' itself may be due to a clay-filled depression or pocket which extends to the west in the top of the limestone. The positive anomaly centered at (20,200) in Figure 17 (near boring E-5) may be a limestone high or pinnacle. The broader negative area around A' in Figure 17 and NE of the main cavity location may be due to the clay- or sand-filled pocket below boring E-6, which, due to its depth range, may be a filled cavity connected to the main cavity.

53. Results of borings E-3 and E-4 present a paradox which cannot be readily resolved. On the basis of the nature of the anomalies C and E' in Figure 17 (Figures 14 and 16), it could be expected that the anomalies associated with C are caused by limestone highs (pinnacles) or limestone with greater than typical density contrast while E' would





be a clay pocket, vertical pipe, or other form of shallow cavity. If the feature causing E' is a vertical pipe similar to the one encountered near S-1, it is easy to imagine how a single borehole could miss the feature. The top of rock at boring E-3 is at 15.5-ft depth, which is typical for the site. At E-4, the top of rock is at 5.5 ft. Thus, except for the possibility of missing a vertical pipe with boring E-4, the results seem inconsistent. The possibility of denser limestone at boring E-3 cannot be verified since the boreholes were not cored.

#### PART V: SUMMARY AND CONCLUSIONS

54. This report presents the results of a microgravimetric survey over a known cavity system at the Manatee Springs site, Fla. The cavity system is water-filled and is at a nominal depth of 90 ft to top, with cavity height ranging from 6.5 ft at boring E-5 to 17.6 ft at S-1 and nominal width of 40 ft. A smaller cavity feature (15 ft in vertical dimensions with a highly irregular cross section ranging up to 10 ft maximum horizontal dimension) branching at about 90 deg to the main cavity was discovered by cave divers during work at the site.

55. Detection of the cavity system by microgravimetric methods was expected to be difficult due to the predicted low amplitude (~-35 μGal) and long spatial wavelength (~200 ft) of the anomaly caused by the cavity system. The survey was planned primarily for the detection of the anomaly caused by the main cavity system. Thus, while the 20-ft station spacing used in the survey is certainly adequate to delineate\* larger solution features, 6-ft effective diameter and larger, at depths from 2 to 4 times the effective diameter (depending on geometry),\*\* small solution features such as the small-scale, shallow limestone pinnacles and clay pockets encountered at the Medford site likely would not be delineated (Butler 1980, 1983). Small solution features as well as other small-scale density variations will appear as "lithological gravity noise" on the gravity maps.

56. Results of the microgravimetric survey clearly indicated the location and trend of the main cavity system across the survey area. Anomaly magnitude and width were consistent with predictions based on known geometry and depth of the main cavity system. Also a closed

<sup>\* &</sup>lt;u>Delineation</u> refers to mapping of the details of the anomaly caused by a subsurface feature. <u>Detection</u> of the anomaly caused by a subsurface feature can be accomplished, however, without this detailed mapping.

<sup>\*\*</sup> Actually, these size/depth rules of thumb will generally be conservative since gravity anomalies will often be greater than that calculated based on measurable dimensions due to secondary effects around the solution features.

negative anomaly superimposed on the broad negative caused by the main cavity satisfactorily delineated the small secondary cavity feature. Mapped locations of the main and secondary cavities were verified by borings and cave divers. These two anomalies were apparent despite a fairly high level of lithological noise.

57. Several other anomalies were apparent on the gravity maps. One small negative anomaly can be correlated to a vertical solution pipe discovered during drilling operations. Only a very limited verification drilling program was possible at the site (six borings) and these were not cored. Of the total of 12 borings at the site, the gravity data were consistent with the subsurface conditions revealed by all but two of the borings. These two borings were in the NE half of the survey area away from the area above the main cavity system. Two gravity anomalies (one positive and one negative), which these borings were to investigate, indicated the possibility of a limestone pinnacle or extended high rock area for the positive anomaly and the possibility of a clay pocket, very shallow cavity, or vertical solution pipe for the negative anomaly. The borings confirmed neither of these possibilities and, in fact, indicated conditions somewhat contradictory to the stated possibilities. Narrowly missing a small-diameter vertical pipe could explain the contradiction for the case of the negative anomaly, but a similar explanation for the contradiction between gravity anomaly indication and boring results is not possible for the positive anomaly, in the absence of core data for the boring.

58. Approximately 12 man-days were required for establishing the site grid system, conducting the elevation survey, conducting the microgravimetric survey, processing data, and producing a hand-drawn residual gravity contour map (Figure 14). At this stage, i.e., after the residual gravity contour map is produced, recommendations regarding site suitability for a given use could be made, assessment of the possibility of solution features at the site could be made, and/or a site drilling program layout could be recommended. For comparison, a twoman rotary rig drill crew (no geologist or inspector) could only produce approximately six 100-ft-deep NX core drill holes for the same

investment of man-time. Six borings at random could do little to characterize a 100- by 400-ft site in a karst region. Microgravimetric surveys coupled with selective drilling are seen to be a time- and cost-effective site investigation procedure for karst areas.

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	170						water tabl	e at 10.7 ft
	+ =		Bottom of hole at	173.0	<u>†∙</u> —		on to duite	
	-	]						
	=	1						
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	·	T 6		III ALL	A 1128	·	Hole No. 0-2
DRILL	ING LO	<u>د</u>	wotechnical Lab		WEG		OF ? SHEETS
anatee	Spring	s Site		10. 3128 11 DATE	AND TYPE	OF BIT :	SHOWN (THE - MEL)
LOCATION	Coordina	nee er Ste	tion)	12. MANI	FACTURE	R'S DESIG	
omilling obile D	AGENCY Distrie	t Corp	s of Engineers	Fici	ling 1	.0 <u>)</u>	
HOLE NO. and No ma	(As shows	an drawa	ng sute	aunt	DEN SAMPL	ES TAKE	n
NAME OF	Juni Creating			14. TOT	AL NUMBER	OUND WA	ORES 13
DIRECTIO	N OF HOL	C			EHOLE		RTED COMPLETED
	CAL []'	-	DEG. FROM VERT.	17. ELE	VATION TO	P OF HOL	<u>2 May 80 - 1 20 May 50 -</u> F <b>E</b>
THICKNES	S OF OVE	ROURDE	■ <u>15.8 m</u>	18 TOT	AL CORE R	ECOVERY	FOR BORING 38
TOTAL DE	PTH OF	HOLE	170.0 ft	19. SIGN	ATURE OF	INSPECT	0 <b>R</b>
EVATION	<b>рертн</b> Г.	LEGEND	CLASSIFICATION OF MATERIA (Description)	1.5	1 CORE RECOV- ERV	BOX OR SAMPLE NO.	REMARKS (Driffing time, water loss, depth of mealburing, etc., if eignificant)
	" –		sand, fine to very fine	· ,			8" ID plastic casing
F			white to yellowish br becoming silty with d	own, euth			set to 20 ft
4			secontin, arrey writh 0				
lic		.					
[ແມ]	10 -				ole		
			13.8		Sami		
			Ls residue - sandy clay with is nodules		2		
	=		19.0				
	20 =		Ls, white to creamy,		Run 1		limonitic stained join
	=		granular, slightly o ted, fossiliferous	emen- verv	942		RQD = 94%
			weak and soft, pasty			Box 1	
			gastropods, pelecype	ds,	Run 2		RQD = 94%
	30 -		THE DESTOIDT OF STREET				joints
	=				100%		
				diam.	L	3	
	=				Run 3		RQD = 96%
	<sup>40</sup> =				985		
	=						
						3	
£	50 _				3 <b>un</b> 4		RQD = 80%
ala	ЃТ Т				100%		
00			-concentrated formit	zone	1.000		
				20110			
					Run 5	1,	8QD = 20 <b>%</b>
	<sup>60</sup> –				99%		
	=						
							3QD = 24%
	70 -				362	r,	
	l' T						Watertable 7.4 ft on
					Run 7		RQD = 495
					953		
	-03						losing water return
			Eg. olightly commenter	<u> </u>	inan 6	6	7QD = 23%
			granular, massive, gr	nyish	975		
ខ្ម			brown to white, vuryy				
ton.			maroforams with some	ρ <b>τ</b>	Pan 9		RqD -= 53\$
1131	- <sup>20</sup>		pelecypois and gastro	poda	107	7	
111	] =				1.00 <i>0</i>		
.•				ici-	L		
	100 =		fied is, brown hard, fossil., hard and coff	zones.	$\frac{\text{R}_{10}}{315}$	8	
G FORM	1836	Pervio			PROJECT		HOLE NO.

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AILLING	100	(CONT 3				Hole No. C-2
<b>osc</b> Manatee	Sprin	gs Sit	e WES			SHEET 2 CH 2 SHEETS
LEVATION	DEPTH It	LEGEND	CLASSIFICATION OF MATERIALS (Den riphon)	% CORE RECOV- ERY	BOX OR SAMPLE NO.	REMARKS (Drilling time: water luss, depth of weathering: etc. if significant)
•	100 _	- C	d	e Run 10		<b>K</b>
	_		irregularly silicified ls.	Run IO		
æ				31 %		
u u				100,		RQD = 33% - Run 13
sto				Run 12	8	800 - 624 - Pup 14
11	110			16%		lost allivater return
111	-	1		Run 1		core barrel stuck in
-		<u> </u>		629		hole; overreamed with
	1 1		115.7	0.5%		4" steel casing to
	=	1	Ls, hard, partially dolomite,	Run 1		retrieve barrel
	120	1	grey to brown, porous,	025		
	_		fossiliferous. friable	32.0		Ls has appearance of
			zones.	Run 1	9	vesicular basalt
	1			92%		ROD = 245 - Run 15
÷	_			Run 1		
G	130	1		1004	]	ven = 013
Lis	-	1		1000	LI	DOD - 514
Dg1	-	<u>t</u> '		Run 1	T I	KQD = 51%
H	-			96%	10	
		1	Laminated dolomite clasts	Run 1	B I	RQD = 60%
(	140	1	in base	024		
	_		142.3	76.0	Į i	505 off
	_		Ls, brown, friable, hard	Run 1	P	RQD = 38%
	-	1	and soft zones, indurated	100%	11	
	_		partially dolomitized,	Run 20	b - 1	RQD = 60%
	150 -	1	scattered carbonaceous	100%		•
			materials	100%		202
i	_	1		Run 2		NAD = 302
:				100%	12	
4				Run 21		RQD = 30%
č,	Te0 –			82		
R				<u> </u>		
8				Run 2	βĺ	RQD = 40%
Āv				047	13	
	111			No		
	170		Bottom of hole 170 0	- core		
1			BOLLOW OI NOTE TIO'O			
	11					
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DRILL	ING LO	G G	vision Potechnical Lab	INSTALL	WES			OF 2 SHEET
PROJECT				10. \$1ZE		OF BIT	3" core and 6	rock bit
lanatee	Sprin	<u>zs Site</u>	<u> </u>	11. BAY		EVATION	- THOWN ( THE - MEL.)	
				12. MAN	UFACTURE	R'S DESIG	NATION OF DRILL	
DRILLING	AGENCY	t Corr	of Engineers	Fail	ling 150	20		
HOLE NO.	(As show	- on draws	ne title	13. TOTAL NO. OF OVER- DISTURBED UNDISTURBES BURDEN SAMPLES TAKEN				
NAME OF	DRILLER		C-3	14. TOTAL HUMBER CORE BOXES ID				
arl Moc	n			18. ELE	VATION GR		TEN 9.5 ft on	15 May
DIRECTION		.E Nelined		16. DAT	E HOLE	13	May 80	22 May 80
<u></u>			18.1.0	17. ELE	VATION TO	POFHO	E	
THICKNES	S OF OVE	ABURDE	N 18.1 ft	18. TOT	AL CORE P	ECOVER	FOR BORING	
TOTAL DE	PTH OF	HOLE	<u>151.9_ft</u>	19. SIGN	ATURE OF	INSPECT	OR	
			CLASSIFICATION OF MATERIA		SCORE	BOX OR	REMAI	RKS
	ft	CEGEND C	(Description)		ERY	HO.	(Drilling time, wet, weathering, etc., 1	it eignificent)
			sand, fine, white to ta	an,			8" plastic I	U casing
	=		silt and clay increase	9			set to 18.1	ft; casing
Ē	_		with depth		1		19.0 ft on 1	eset to 4 Hav
0	=							··- <b>v</b>
lic	10				nfe nfe		Run 1 was ca	ve-in
5	Ξ				t a			
					o fa			
	-		clay tan to brown, vit	h Is	ਿੰਬੋ			
			nodulles		<u> </u>	19.0	100 - 00e	
	20 —		'Ls, pink to gray, hard	,	Run 2		lost circula	tion at
	=		weakly cemented, fria massive, fossiliferou	ble, s	88%	Bort	22 ft; cemen	ted hole,
	_		(gastropods, pelecypo	is		DOX 1	have water r	eturn
	=		and macro-forams)			20 5		
	30				Run 3		RQD = 35%	
					10.4			
	-	1			~~,//	Box		
	-				Dum h	2	losing more	water
	40-				Run 4		RQD = 18%	– Run 4
	-				26%		ROD = 87%	- Run 5
	_						cemented aro	und outside
	-		LS, SOIT, White to cre friable, massive sc	amy, attere	Run 5	1	of casing to	stop wash
	50 -		chert fragments, fos	sili,-	100%	49.6		
-	́ =		ferous as above		Run 6		Add revert t	o stôp
() g	=				100%	Box	water loss ROD = $35\%$ -	Run 6
یے۔ اینے					100%	3		
č Š	=				1			
ala Ri	60 —					60.4		
B.1 Oct	Ξ				Run 7		3QD = 43%	
уяt	-				86=	HOV		
L C	-					4		
.,	=							
	70		1		Run 8		RQD = 255	
	-				45%			
			-76 5			76.0		
			Ls, hard, cream to bi	nk.	Run 9		RQD = 39%	
			well cemented, massiv	e,		Bar	÷	
	- 00	┝	as above		845	5		
	=		13, grey, hard, massiv	e,	I .			
	-	1	friable, abundance of			g., .		
E	=	}	macroforams gives gra	nular ronada	Bun 10	- 01.7	RQD = 575	
2	90 -	]	appearance, some dast	robod3	1000			
<b>*</b> 2	=	1			1.015			
11		\$			1	'		
	=	1			Run 11		891 = 26 <b>%</b>	

RILLING	LOG	(Cont S	heet)		··		Hole No. C-3	
ionici Har	natee S	Springs	: Site	INSTALLATION	ES		SHEET 2	
<b>LEVATION</b>	DEPTH	LEGEND	CLASSIFICATION OF (Decreption	MATERIALS	% CORE	BOX OR SAMPLE	REMARKS (Drilling time, water luss, depth of weathering, etc., of upperform)	4
		ſ			e	<u>(</u>		
						вох 6	void from 106.0 to	
	_	Void .	106.0			106.0	106.5: lost all water	-
	11		Ls, irregularly	silicified,	Bun 12		ROD = 6%	•
ė	110	k/	soft zones. fos	nard and siliferous				
g	-	vòiá	(gastropods, pe	lecypods,	17%	Box	void from 109.0 to 1	14.5
stc		$\angle$	macroforams), f	riable,		7		
11	-		20001 VC			126.3	ROD = 38%	
FM	=				Run 1			
	120 _				70%			
	=		115.7		1			
			partially dolom	, porous,				
			few gastropods	and pele-	Run 1	Bor	RQD = 44%	
,	130		cypods, some gr	anular	604	8		
8 1	-	1	2018:0		027	'		
611	_		3				core barrel stuck in	n
In	=		clasts	u te	Run 1	138.0	hole; over reamed	~
	140	1			67%		to retrieve barrel.	5
		<u>├</u>	142.3		4	Berr	RQD = 48%	
	_	1	Ls, brown to grey	, friable,	ļ	DOX		
		[	partially dolon	utized.	Run 10	5	RQD = 69%	
	150-		carbonaceous ma	terial	925			
	. I		scattered in up	oper section	, , , , ,	153.0		
							1	
ė	=	1			Bup 17	Į	RQD = 52%	
¥	160				- un 11	Box		
Par	100				0.0	10		
ä	Ξ				927			
Av	-				no		with 6" rock bit	ft
			Bottom of hole 17	0.0 ft	core			
					· · · · · · · · · · · · · · · · · · ·			
	=							
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							Hole No	-			
DRILL	ING LO	c C	Geotechnical Lab	INSTALL	ATION UPPE		OF , SHEETS				
PROJECT				10. SIZE AND TYPE OF BIT ' Pock bit							
Location	(Coordin	<u>zs 311</u> atos or \$14	2								
DRILLING AGENCY					12 HANUFACTURER'S DESIGNATION OF DALL						
Mobile .	istri Ae enem	et Cor n <del>ondram</del>	ps of Engineers nemme	13. TOTAL NO. OF OVER- DISTURSED UNDISTURSED							
AND NO ME	HILLER			14. TOT		R CORE E	IOXES				
Carl Mo	on in	1 Earb	ert Wens	18. ELEN	ATION GP	IOUND WA	TER 11.0' on 16 June				
VENTIC	AL []	NCLINED	GES. PROM VERT.	IS DATE HOLE							
-	-		■ <u>30.0 ft</u>	17. ELEVATION TOP OF HOLE							
DEPTH DR	ILLED H	ITO ROCK	150.0 ft	18. SIGN	ATURE OF	INSPECT	OR CONTROL CONT	1			
, TOTAL DE	PTH OF	HOLE	LOD. ) Pt.	L	S CORE	BOX OR	REMARKS	1			
e	ft b	LEGEND	(Description)		ERY	NO.	(Delling time, mater loss, depth of meadlaring, etc., if eignificant)	Ļ			
	Ŭ =		sand, fine, white	!			8" ID plastic casing	E			
	Ξ						set to ji leet	E			
	=							F			
į	10 _	]						E			
- -								E			
lic	-	1 '						Þ			
5	_	1						F			
	-	1						F			
	20-				ŀ	}	casing bent at	E			
	=		ciay, sandy, 15 nodules	3			20 ft joint	F			
	_	1						þ			
ļ	=	1 '				1		þ			
	-	1						Þ			
	30	1	Ls, white, fossiliferou	12				E			
1	Ξ			i		1		E			
	_	1						Þ			
	=	1						F			
	40 <del>-</del>						lost all water at 40 ft;	F			
je i	-		hard ledge				cemented hole with 9	E			
		\$				1	Water return	þ			
3	-	1						F			
ð	50 _	1			les		1	E			
		1			dua			E			
	Ξ	1	J		ທີ່	J		E			
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	60 <del>-</del>	1				1		E			
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111							1	E			
5	=	1						E			
	100 -	1	Cilletted williaton f	`m		1	Drilling much harder	F			
as Foom	18 34		US EDITIONS ARE OBSOLETE		TOJECT		HOLE NO				

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	100	(Cour 3							
May	natee	Spring	Site		STALLATION	WES			SHEET 2
	DEPTH	LEGEND	CLASSIFIC	ATION OF M. (Decemption)	ATERIALS	% CORE RECOV-	BOX OR SAMPLE	REMA (Drilling time, wa weathering, etc.	MRKS Her luss, depth of if ugnificant)
	110-	Ý X	101.0 cavity 108.0 l11.8 cavity 114.8 hard dril alternatin zones to b	drilling ag hard a pottom of	and soft f hole	٤	f	120' - cen with 30 se cement mix	mented hole teks of ted with
. Inglis th.	130					No Samples		1/2 sack mud; cemer to 105.0'	drilling t set up
Avon Park fm.	160							165.0 ft hole with cement, se 105.0 ft, hole with 85 sacks o up to 102.	cemented 40 sacks o t up to cemented additional ement, set 2 ft
			Bottom of h	ole				Note: hole Water table on 16 June	is crooked at 11.0 ft

D <b>P</b> 41 4	INGLO	c   0"	VISION	INSTAL	LATION		Hole Ho.	C-S			
PROJECT		• <u> </u>	<u>destechnicai hab</u>	10. \$12	10. SIZE AND TYPE OF BIT (1-1/2" aroun bit						
LOCATION	Ser Lite I (Coordin	rs wite noe or Sea	(ion)	-11.07		EVATION	SKOWN ( <b>TBN <del>(</del>F MRL</b> )				
DRILLING	AGENCY		<u> </u>	12. MAI	UPACTURE		SNATION OF DRILL	· · · ·			
bile :	<u>11.31.91</u>	)r:	<u>s of Chytineyrs</u>	12. 10	12. YOTAL NO. OF OVER- DISTURSED UNDISTURSED						
and Nie ma			C-5		14. TOTAL NUMBER CORE BOXES						
NAME OF	DAILLEA N. 2 1 1			18. EL	EVATION G	IOUND WA	TER 11. ft (	n 10 June			
DIRECTIO				16. DA	TE HOLE	= <b>- 18та</b> 27	ATED IC	SO May 80			
THICKNES	S OF OVE	REUROE	17.)		EVATION TO	P OF HO					
-	ILLED IN	TO ROCK	1,3.0 rt	18. 10	TAL CORE P	INSPECT	Y FOR BORING	0			
TOTAL DE	PTH OF	HOLE	170.0 ft	1	1						
LEVATION	DEPTH I't	LEGEND	CLASSIFICATION OF MATE (Description) d	RIALS	RECOV-	SAMPLE NO.	(Drilling time, wet weathering, etc.,	it eignificanti			
÷			sand, white, fine				8" ID plast:	ic casing			
tn.	-						set to 27.0				
			7.0 it		-		hole starte	12.5 ft			
	10 _		clay, sandy with 1s :	nodules			west drille	1 to 26.5 f			
						l	lost 8" reel	ing rock; k bit in			
	_						hole; fille	l with sand			
			17.0"  Ls, creamy to whit	ce.			and 1 sack moved to pro	cement and esent			
	20 _		soft, fossilifere	ous			location				
	30 =				1						
	]					1	Lost all wat	er return			
÷	_=	1			1	1					
							Cemented hol	e with 4			
Cal	10 E				1	[	sacks cement	; 60 to 80%			
ð							cementing an	d using			
						l	drilling mud	~			
					1	1					
	50 -		hard zone		1						
					1						
					1						
			very soft aone		1						
	<u> </u>										
	Ξ										
					I						
	Ξ					1					
	70 <del></del>		hard conc of 71 and	71. 6+							
	Ξ		nanaa na oronna 100 jil oliilli j	14 16		l					
					1		207 Water	return			
	Ē				1						
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	<u>1 03</u>	L	· · · · · · · · · · · · · · · · · · ·		I REGISCE	L	L				

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RILLING	LOG	(Cont S	heet)				Hole No.	0-5
Manat	ee Spr	ings S	Site	INSTALLATION	WES			SHEET 2
LEVATION	PERTH TETTH	LEGEND	CLASSIFICATION OF	MATERIALS	* CORE RECOV- ERY	BOX OR SAMPLE NO.	REMA (Drolling time, w. weathering, etc.	AKS Her luss, depth of of Hymsficant)
	100_	<u>с</u>						·
			top of silicifi	ed			verv bard d	rilling
ron 20			Williston at 1	04 ft	i i			
	-				1			
- <u>-</u>	110-							
-	-							
:_?	_						lost all wat	er at 114.0
	=						of cement; 1	ost all
	120		alternating hard	and soft	1		water return	after
	-		zones to bottom	or note		ļ	drilling thr	ough cement
	-							
å								
lis	130	1						
Ing	_	1	1					
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1_ 2_	140							
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Par					[	Í		
g	100 =							
Av	_							
	-						water table	at 11.5
	=		Bottom of hole				commented hol	a with 8
	170						sacks of cem	ent to help
	-						hold it open	; cement se
	-						questionable	how good
	=						top 25 ft wa	s set up
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In accordance with letter from DAEN-RDC, DAEN-ASI dated 22 July 1977, Subject: Facsimile Catalog Cards for Laboratory Technical Publications, a facsimile catalog card in Library of Congress MARC format is reproduced below.

Butler, Dwain K. Cavity detection and delineation research : Report 4 : Microgravimetric survey : Manatee Springs Site, Florida / by Dwain K. Butler, Charlie B. Whitten, and Fred L. Smith (Geotechnical Laboratory, U.S. Army Engineer Waterways Experiment Station). -- Vicksburg, Miss. : The Station ; Springfield, Va. : available from NTIS, 1983. 47, [12] p. : ill. ; 27 cm. -- (Technical report ; GL-83-1, Report 4) Cover title. "March 1983." "Prepared for Office, Chief of Engineers, U.S. Army under CWIS Work Unit 31150." Bibliography: p. 47. 1. Geophysical research. 2. Gravimeter (Geophysical instruments). 3. Manatee Springs (Fla.) I. Whitten, Charlie B. II. Smith, Fred L. III. United States. Army. Corps of Engineers. Office of the Chief of Engineers. IV. U.S. Army Engineer Waterways Experiment Station.

Butler, Dwain K. Cavity detection and delineation research : ... 1983. (Card 2) Geotechnical Laboratory. V. Title VI. Series: Technical report (U.S. Army Engineer Waterways Experiment Station) ; GL-83-1, Report 4. TA7.W34 no.GL-83-1 Report 4

