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TECHNICAL REPORT GL-83-1

CAVITY DETECTION AND DELINEATION RESEARCH

Report 4

MICROGRAVIMETRIC SURVEY: MANATEE SPRINGS SITE. FLORIDA

- by 🚊 wain K. Butler, Charlie B. Whitten, and Fred L. Smith

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Report 4 of a Series

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Technical Report GL-83-1

CAVITY DETECTION AND DELINEATION RESEARCH

·····	Title	Author	
Report 1:	Microgravimetric and Magnetic Surveys: Medford Cave Site, Florida	Dwain K. Butler	
Report 2:	Seismic Methodology: Mediord Cave Site, Florida	Joseph R. Curro, Jr.	·· ·
Report 3:	Acoustic Resonance and Self-Potential Applications: Medford Cave and Manatee Springs Sites, Florida	Stafford S. Cooper	
Report 4:	Microgravimetric Survey: Manatee Springs Site, Florida		بر ا
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	Electromagnetic (Radar) Techniques Applied to Cavity Detection	Higher Proceed, Jr.	
	Cavity Detection		14
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20. ABSTRACT (Continued).

- contrast of -1.3 g/cm³, the gravity anomaly is calculated to be -35 μ Gal \cdot with a width at half maximum of 205 ft.

The microgravimetric survey results clearly indicate a broad negative anomaly coincident with the location and trend of the cavity system across the survey area. The anomaly magnitude and width are consistent with those calculated from the known depth and dimensions of the main cavity. In addition, a small, closed negative anomaly feature, superimposed on the broad negative feature due to the main cavity, satisfactorily delineated a small secondary cavity feature which was discovered and mapped by cave divers.



PREFACE

This investigation was performed by personnel of the Earthquake Engineering and Geophysics Division (EEGD) and the Engineering Geology and Rock Mechanics Division (EGRMD), Geotechnical Laboratory (GL), U. S. Army Engineer Waterways Experiment Station (WES), for the Office, Chief of Engineers (OCE), U. S. Army, during the period May 1980 to January 1981. The investigation was part of CWIS Work Unit 31150, "Remote Delineation of Cavities and Discontinuities in Rock." Mr. Paul R. Fisher, OCE, was Technical Monitor for this work.

This report was prepared by Mr. Dwain K. Butler, EEGD, and Messrs. Charlie B. Whitten and Fred L. Smith, EGRMD, under the supervision of Dr. Arley G. Franklin, Chief, EEGD, and Dr. Don C. Banks, Chief, EGRMD, and under the general supervision of Dr. William F. Marcuson III, Chief, GL. In addition, Mr. Rodney N. Walters, EEGD, assisted with the fieldwork. This investigation was closely coordinated with ongoing work under Project MA 61102AT22, Work Unit 002/Q6, "Analytical and Data Processing Techniques for Interpretation of Geophysical Properties."

Special appreciation is expressed to Mr. Sheck Exley and Dr. John Zumrick of the National Speleological Society for assistance in obtaining maps of the Manatee Springs Cave System and arranging for cave divers to conduct detailed cave mapping during the fieldwork. Also, the courtesy extended by MAJ Ellison E. Hardee of the Florida Department of Natural Resources in granting permission to use the Manatee Springs Site and the assistance of CPT Cecil Dykes, Park Superintendent, throughout the fieldwork are gratefully acknowledged.

COL Nelson P. Conover, CE, and COL Tilford C. Creel, CE, were Commanders and Directors of the WES during this investigation. Mr. Fred R. Brown was Technical Director.

CONTENTS

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	Page
PREFACE	. 1
CONVERSION FACTORS, U. S. CUSTOMARY TO METRIC (SI)	
UNITS OF MEASUREMENT	. 3
PART I: INTRODUCTION	. 4
Background	. 4
Purpose	
Scope	
PART II: THE MANATEE SPRINGS SITE AND CAVE SYSTEM	. 7
Location and History	
Cave System	
PART III: GEOLOGY OF THE MANATEE SPRINGS SITE	. 15
Physiography	. 15
Field Exploration Program	
Stratigraphy	
Groundwater and Solution Features	. 22
PART IV: MICROGRAVIMETRIC SURVEY	. 24
Background	. 24
Survey and Data Processing Procedures	
Results	. 30
Verification Drilling	. 38
PART V: SUMMARY AND CONCLUSIONS	. 44
REFERENCES	. 47
APPENDIX A: BORING LOGS	. A1

CONVERSION FACTORS, U. S. CUSTOMARY TO METRIC (SI) UNITS OF MEASUREMENT

U. S. customary units of measurement used in this report may be converted to metric (SI) units as follows:

Multiply	By	To Obtain	
cubic feet per second	0.02831685	cubic metres per second	
feet	0.3048	metres	
gallons (U. S. liquid) per minute	3.785412	cubic decimetres per minute	
inches	2.54	centimetres	
miles (U. S. statute)	1.609347	kilometres	

CAVITY DETECTION AND DELINEATION RESEARCH

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MICROGRAVIMETRIC SURVEY: MANATEE SPRINGS SITE, FLORIDA

PART I: INTRODUCTION

Background

1. Early in 1979, two field test sites in karst regions were selected for assessment and evaluation of geophysical methodologies for cavity detection and delineation. The first of these sites, Medford Cave, in Marion County, Fla., has a relatively shallow (approximately 10 to 30 ft* (3 to 9 m) to top) air-filled cavity system. Manatee Springs, in Levy County, Fla., the second site, has a deeper (approximately 100 ft (30 m) to top) water-filled cavity system. In addition to the known and mapped cavity systems, the sites have additional geologic complexities in common with other sites in karst regions.

2. The field programs at the two sites were not equivalent in scope. An extensive program was carried out at the Medford site, including 19 geophysical techniques. Report 1 in this series (Butler 1983) discusses the scope of the program at the Medford site in more detail. The program at the Manatee Springs site was more limited in scope due partially to fiscal and time constraints in the research program. Also it was desired to limit the program at the Manatee Springs site to those techniques showing the greatest potential and to techniques for which understanding would be considerably advanced by application at the deeper, water-filled site.

3. This report documents the results of a microgravimetric survey conducted at the Manatee Springs site. Detection and delineation

 ^{*} A table of factors for converting U. S. customary units of measurement to metric (SI) units is presented on page 3. The customary unit for gravitational acceleration in microgravimetric surveys is the µGal, which is defined in the text. Also grams per cubic centimetre is used as the unit for density.

of the Manatee Springs cavity system represents a more difficult objective for microgravimetric surveying than Medford Cave, since the density contrast is less than for the air-filled Medford Cave system and the depths are considerably greater than for the Medford Cave system. Maximum residual gravity anomalies at the Medford Cave site were about -70 μ Gal,* while calculations based on the known cave system depth and dimensions indicate that the maximum residual gravity anomaly at Manatee Springs should be about -35 μ Gal (Butler 1980, 1983). Since the error of relative gravity determinations in a microgravimetric survey lies in the range of ±5 to ±10 μ Gal (Butler 1980) and there is likely a certain level of lithological gravity noise at the Manatee Springs site (as was present at the Medford Cave site (Butler 1983)), the Manatee Springs site will allow further definition of the range of applicability of microgravimetry for geotechnical objectives.

4. In addition to the microgravimetric survey, the following geophysical techniques were also applied at the Manatee Springs site:

- a. Crosshole acoustic surveys.
- b. Crosshole radar surveys.
- <u>c</u>. Acoustic resonance (surface survey with subsurface sonar source).
- d. Spontaneous-potential (SP) survey (surface).

Conventional borehole geophysical logging suites were also obtained in two boreholes at the site. The results of these geophysical investigations at the Manatee Springs site will be the subjects of other reports in this series.

Purpose

5. Results of geophysical investigations at the Medford Cave site indicate that microgravimetry is one of the best methods for shallow cavity detection and delineation. The purpose of the present work is to determine the capability of microgravimetry to detect the deeper

* 1 Gal \equiv 0.01 m/sec².

Manatee Springs cavity system. A secondary purpose of this work is to investigate procedures for determining the local regional gravity field, and the effect on the residual gravity anomaly map, qualitatively and quantitatively, of using different local regional fields.

Scope

6. Part II of this report will discuss the Manatee Springs site, the site grid system and topographic survey, and the cave system. Part III will discuss the site geology and drilling program. Results of the microgravimetric survey and correlations with known site conditions will be presented in Part IV. Finally, a summary and conclusions will be presented in Part V.

PART II: THE MANATEE SPRINGS SITE AND CAVE SYSTEM

Location and History

7. Manatee Springs is located in the Manatee Springs State Park, which is in Levy County, Fla., about 6 miles west of Chiefland, Fla. The site chosen for the geophysical surveys is near the spring boil,* which is located about 1500 ft from the east bank of the Suwannee River and 23 miles upstream from the Gulf of Mexico. Figure 1 shows the loca tion of the site on a road map and Figure 2 is a portion of the Manatee Springs 7.5-minute Quadrangle Sheet (U. S. Geological Survey 1954).

8. The spring has been known and used by man for centuries, originally for the crystal-clear drinking water and abundant fish and game and more recently as a popular site for picnicking and swimming. William Bartran, a naturalist, visited the spring in 1774 and provided the first known description of the area. The spring was named for the manatee, an endangered aquatic mammal, which is occasionally seen in the area.

Cave System

9. The Manatee Springs cave system extends several miles to the southeast of the spring boil, and through 1980, 10,040 ft had been mapped by the Cave Diving Section of the National Speleological Society. Figure 3 is a map of a portion of the cave system (Exley 1979). Depths to the top of the cave system are in the range of 80 to 100 ft (25 to 30 m) in the vicinity chosen for the geophysical surveys, and the cavity has a cross section 18 ft in height and greater than 40 ft in width at one location in the survey vicinity.

10. Depth and cross section of the cave system are actually quite variable. Also, there are numerous smaller branching cave features along the length of the system. Sinkholes and even smaller

* Flow from the spring is 81,820 gallons/min.



Figure 1. Road map showing location of the Manatee Springs site





Site Grid Layout and Topographic Survey

11. The site chosen for the microgravimetric survey is about midway between the spring boil and the Catfish Hotel (sink), as shown in Figure 4. The long axis of the rectangular topographic survey area was



Figure 3. Map of a portion of the Manatee Springs cave system (Exley 1979)

chosen to be perpendicular to the local trend of the cavity system, and was originally 120 by 400 ft.* Discovery of a small branching cavity feature by cave divers led to the extension of the survey area to the southeast to include the small cavity feature (Figure 5) and thus accommodate other geophysical tests planned to investigate detectability of the smaller feature.

^{*} Microgravimetric survey area - 100 by 400 ft (Figure 4); topographic survey area - 120 by 400 ft (Figure 6); extended grid area (Figure 5) was not topographically surveyed.







12. A basic site grid dimension of 20 ft was chosen (illustrated in Figure 5). Orientation of the long axis of the survey area was established at N45°E. Grid locations were determined using steel tapes and chaining pins in the usual manner. At each grid location, a 2- by 2- by 6-in. stake was driven flush with the ground surface. At every alternate grid location, an offset reference stake labeled with coordinates was placed. The sense of the survey coordinates is explained in Figures 4 and 5.

13. Elevations of the tops of the 2- by 2-in. stakes were determined to ± 0.01 ft. Elevations were referenced to a U. S. Geological Survey benchmark (14.00 ft mean sea level (msl))* located at (47,250) of the survey grid, and the closure error was 0.01 ft. Figure 6 is the topographic map resulting from the elevation survey. Elevation variations over the site are very small, with the <u>maximum</u> elevation difference in the survey grid being approximately 7.0 ft. Except for the semiconcentric contours about the point (120,320) with an elevation of 8.5 ft msl, which reflects the presence of a small sinkhole with center approximately 100 ft beyond the survey area, the topography resembles an inclined plane, decreasing in elevation from 15.6 ft msl at (120,0) to 10.0 ft msl at (0,400). The grid layout, elevation survey, and mapping of the plan location of the cavity system required four man-days (two days for two-man crew).

* United States Geological Survey Benchmark TT28BB5.



PART III: GEOLOGY OF THE MANATEE SPRINGS SITE

Physiography

14. The site is in Vernon's Terraced Coastal Lowlands Subdivision which is within the Floridian Section of the Coastal Plain Province (Vernon 1951). This section is a very youthful, recently emergent, terraced coastal plain with karst topography. The Terraced Coastal Lowlands is divided into four Pleistocene marine terraces based on elevation above msl--the top elevation of the Coharie is 220 ft msl; the Okefenokee, 150 ft msl; the Wicomico, 100 ft msl; and the Pamlico, 25 ft msl. The Manatee Springs site ranges in elevation from about 10 to 15 ft and is within the Pamlico marine terrace which extends inland from the coast.

15. The Pamlico surface in the Manatee Springs area consists of a broad shelf of Tertiary limestones covered with a thin veneer of Pleistocene sands. Erosion of the Tertiary limestones prior to deposition of the sands created a rugged surface with considerable relief. The ground surface at the site varies less than 2 ft between any of the boreholes, but the top of the limestone varies in depth from 6.6 ft at boring S-1 to 20 ft at boring C-4 (Figure 7). The top of limestone at C-5 varies 9.5 ft vertically in a 2.5-ft horizontal distance (see boring log C-5 in Appendix A).

Field Exploration Program

16. The number of borings was initially limited to a maximum of 10 by the Florida State Park Service. Only six borings were actually drilled in the initial program. Borings C-1 through C-5 are for testing downhole geophysical instruments, and S-1 was an exploratory boring used to locate borings C-2 through C-5 (Figure 7). The boring locations, diameters, and depths were selected to accommodate test equipment requirements and are discussed in the following paragraphs.

17. Prior to the fieldwork, a map of the Manatee Springs cave,



prepared by cave divers, was used to project the cave outline to the surface at the site, and a 20-ft grid was established. The survey was used to position the borings. The cave outline was used for the drilling of two exploratory borings; the first boring was to penetrate the cave, and the second boring was located approximately 50 ft to the east of boring 1. Divers found boring 1 was drilled through a relatively thin angular section of limestone projecting out from the west cave wall. Boring 2 penetrated the cave near its east wall and was used by the divers as a fixed point for distance and direction measurements to locate borings C-2 through C-5. Boring 2 was later designated boring S-1. Boring 1 was later designated boring C-1 and was to be used for testing. Boring logs are presented in Appendix A.

18. The divers discovered a small side cave extending approximately 600 ft to the southwest of boring C-1. The small cave was more suitable for testing purposes than the larger Manatee Springs cave, so borings C-2 through C-5 were located astride the small cave. The small cave is between borings C-2 and C-3.

19. Borings C-2 and C-3 were cored with 3-in. core barrels and then reamed with a 6- and 6-1/2-in. rock bit, respectively. Boring C-1 was drilled with a 4-in. rock bit, and borings C-4, C-5, and S-1 were drilled with a 6-1/2-in. rock bit.

20. The depth requirement for borings C-1 through C-5 was 165 ft, but the borings were drilled to 170 to 180 ft to allow room for cuttings and possible spalling material to settle. Boring S-1 penetrated the roof of the Manatee Springs cave at a depth of 88.0 ft and was not drilled any deeper. The floor of the cave was at a depth of 105.6 ft.

21. Cavity openings, varying from 1.1 ft in boring C-2 to 7.0 ft in boring C-4, were detected in borings C-2, C-3, and C-4. Divers confirmed that the cavities were not locally connected with the small cave. Boring C-4 was grouted with 155 sacks of cement and reamed to 6-1/2 in. to prevent instruments from getting hung up in the openings. Borings C-2, C-3, and C-4 were not affected by the grouting, indicating that the cavities were probably not locally interconnected.

Stratigraphy

22. The geologic units encountered were the Pamlico Formation of Pleistocene age and the Ocala, Williston, Inglis, and Avon Park Formations of Eocene age (Figure 8). The Ocala Formation has been renamed the Crystal River Formation in more recent publications, but the term Ocala Formation will be used in this text to maintain continuity with previous work. Lithologic descriptions of the Eocene limestones are limited to the core samples from borings C-2 and C-3 and cuttings from borings S-1, C-1, C-4, and C-5.

Era	System	Series	Group	FORMATION					
Quaternary	Pleistocene		Pamlico Formation						
Cenozoic Tertiary				ле	ne	ne	he	ы	Ocala Formation (Crystal River Formation)
		Upper Eocene	upper cuce Jackson	Williston Formation					
	iary			Inglis Formation					
	Tert	Tert	Middle Eocene	Claiborne	Avon Park Formation				

Figure 8. Geologic formations at the Manatee Springs site

Pamlico Formation

23. The Pleistocene Pamlico Formation is a fine, well-sorted, white to yellow sand, varying from 6.6 to 20 ft thick (Figures 9a and 9b). The sands were deposited in a marine environment and later reworked into a Pleistocene marine terrace as sea level regressed. The terrace sands have been locally reworked by the Suwannee River. Vernon (1951) classified the sands at the site as Suwannee River sandbars. <u>Ocala Formation</u>

24. The Ocala Formation is the upper limestone in the Jackson Group and is typically a soft to moderately hard, white to pink, massive, friable, fossiliferous limestone (Figures 9a and 9b). The fossils consist mainly of gastropods, pelecypods, and foraminifera, and form a coquina in many places. The Ocala is up to 63.4 ft thick at the site. The upper 2 to 10 ft of the Ocala is a weathered zone consisting of limestone fragments in a sandy clay to clay matrix. The Ocala grades conformably downward into the Williston Formation. Changes in fauna and the granular texture of the Williston were used to identify the Ocala-Williston contact.

Williston Formation

25. The Williston Formation is a grey, moderately hard, friable, fossiliferous limestone and is harder than the Ocala and softer than the underlying Inglis. The abundance of foraminifera gives the Williston a granular appearance. The Williston is up to 41.5 ft thick at the site. The lower 16 to 17 ft is partially silicified with scattered zones of poorly silicified to unsilicified limestone (Figures 9a and 9b). The silicified zone was encountered at a depth of 15.7 ft in boring C-2 and 125 ft in C-3 and can be easily identified in the subsurface by the increased drilling time. The Williston-Inglis contact is conformable and can be recognized by the change from a silicified limestone to a porous dolomitic limestone.

Inglis Formation

26. The Inglis Formation is the base of the Jackson Group and is a hard, brown, porous, partially dolomitic limestone with scattered zones of hard, grey, fossiliferous limestone. The dolomitic limestone



a. Section A-A'

Figure 9. Geologic sections (Continued)



Figure 9. (Concluded)

is porous but poorly permeable and ranges from soft and friable to hard and indurated. Laminated dolomitic clasts eroded from the Avon Park are incorporated in the base of the Inglis. The Inglis-Avon Park contact is unconformable with a soil zone on the eroded surface of the Avon Park. The thickness of the Inglis varies from 26.6 ft at boring C-2 to 19.2 ft at boring C-3.

Avon Park Formation

27. The Avon Park Formation is a brown to grey, porous, partially dolomitized, moderately hard limestone with carbonaceous plant remains scattered throughout. A few large fossil casts and molds of gastropods and pelecypods and a burrow occur at the top of the unit (Figures 9a and 9b). Only 23 ft of the Avon Park was cored at the site; however, Vernon (1951) described it as being several hundred feet thick in this area.

Groundwater and Solution Features

28. The high permeability of the Pamlico sand and the Eocene limestones has produced a very poorly developed surface drainage system which has resulted in almost all of the rainfall being absorbed directly into the groundwater system. The groundwater elevation at the site is controlled by the fluctuation of the Suwannee River which was above flood stage at the beginning of drilling operations on 5 May 1980. The water table at the site was at 7.6 ft on 5 May 1980. As the flooding receded, the water table dropped to 9.5 ft on 15 May 1980 and 11.5 ft on 16 June 1980.

29. The groundwater is highly charged with organic and carbonic acids and plays an important part in the development of the caverns in the area (Vernon 1951). The permeability of the Ocala and Williston limestones allows for rather free movement of groundwater through these units. The poor permeability of the Inglis has restricted groundwater flow, thus concentrating the groundwater in the overlying more permeable lower Williston. This concentration of goundwater in the Williston has resulted in the development of a large complex cavern system and

probably in the irregular silicification of the lower Williston. Cave divers have explored over 10,000 ft of the Manatee Springs cave and report that it probably extends much further. The Manatee Springs cave has a variable height and width, and has numerous smaller branching caves. Data from the divers and boring S-1 show Manatee Springs cave is 17.6 ft high and over 40 ft wide at boring S-1. The small cave is approximately 4 ft high and 8 ft wide between borings C-2 and C-3, but the height and width are very variable along the length of the cave.

30. The depth of the Manatee Springs cave decreases from 105.6 at boring S-1 to approximately 55 ft (relative to the borings) at the Manatee Springs boil, a distance of approximately 250 ft to the north. The measured surface discharge from the spring has varied from 110 ft³/sec to 238 ft³/sec (Rosenau et al. 1977). The spring pool is approximately 100 ft wide and 55 ft deep.

31. The only other solution features detected during the drilling operations were in the Ocala. Boring S-1 intersected a vertical solution pipe 1 to 1.5 ft in diameter. The solution pipe collapsed as a drill rig was being moved by the boring, creating a hole approximately 5 ft in diameter at the surface. Boring S-1 and the funnel-shaped collapsed area were filled with a sand-cement mixture to prevent further deterioration or enlargement by rainfall. Some minor weathering along fractures was found in the core samples, but it is only recognizable by the discoloration of the limestone.

32. The collapse of caverns has created several sinks in the Manatee Springs area. Three sinks, varying from approximately 75 to 200 ft in diameter, are located along the course of the main underground cave, and are within 500 ft of the study site. Catfish Cabin, the largest of the sinks, is approximately 75 ft southeast of boring C-4 and is connected to the Manatee Springs cave by an opening up to 10 ft in diameter.

PART IV: MICROGRAVIMETRIC SURVEY

Background

33. As noted in Part I, detection and delineation of the Manatee Springs cave system by microgravimetry represents a more difficult problem than the delineation of the Medford Cave system. The difficulty arises from two aspects of the situation: (a) the maximum value of the gravity anomaly at Manatee Springs should be smaller by a factor of two (-35 μ Gal compared to -70 μ Gal); and (b) the greater depth of the Manatee Springs cave system will cause the spatial wavelength or width of the gravity anomaly to be much larger than the gravity anomalies at the Medford site (see Butler 1980, 1983). Low amplitude, long wavelength anomalies are difficult to discern in cases where the dimensions of the survey area are comparable to the wavelength.

34. For site surveys with areas comparable to those at the Medford and Manatee Springs sites, it is possible that long wavelength anomalies will be interpreted as part of the local regional component. This fact can be illustrated by considering a model calculation of the anomaly caused by a structure with dimensions and density contrast appropriate for the Manatee Springs cave system. The depths and dimensions of the main cave near boring S-1 (see Part III) are used to model the cave as a long horizontal cylinder with rectangular cross section and a density contrast of -1.3 g/cm^3 (simulating a water-filled cavity in limestone of density 2.3 g/cm³). The gravity anomaly is calculated along a 400-ftlong profile line perpendicular to the cylinder axis, with the 200-ft position directly above the axis. The geometry is illustrated and the gravity anomaly profile is shown in Figure 10. Magnitude of the anomaly is -36 μ Gal; and the wavelength, defined to be the width at half-maximum (Butler 1980), is 205 ft. However, the anomaly still has an amplitude of -7μ Gal at the ends of the profile line.

35. Unless there is considerable lithological noise (Butler 1983) or there are other superimposed gravity anomalies, an anomaly such as shown in Figure 10 should be detected by a carefully conducted





microgravimetric survey. The primary difficulty is that the size and depth of the feature causing the gravity anomaly may be underestimated due to clipping of the gravity profile by the selected regional field or because profile lines are too short to define the complete anomaly.

Survey and Data Processing Procedures

Station layout and survey programs

36. Stations occupied during the microgravimetric survey are

illustrated in Figure 11. The survey consisted of 126 stations on the 20 ft grid with about 30 percent of the stations reoccupied at least once during the survey. Station (0,200) was selected as base station for the survey. All stations occupied between successive reoccupations of the base station are referred to as a program. Programs were designed to occupy stations in a "zigzag" or "leapfrog" fashion in order to distribute random errors and to prevent any cumulative errors from combining to produce fictitious anomalies or elongated anomalies such as can result from a long, continuous program of station occupations. Three typical programs (E, I, and N) are indicated by the dashed lines in Figure 11. Each program was planned to include at least one repeat station occupation and to be short enough that the base station can be reoccupied in less than 1 hr. In fact, the mean program duration was 40 min. The field gravity measurements required a total of 40 manhr (20 hr each for a two-man crew).

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37. The survey was conducted with LaCoste and Romberg Model D-25 gravimeter. This instrument has a range of about 200 mGal and the smallest dial division equals approximately 1 μ Gal.* All field measurements were obtained using the top-mounted galvanometer (capacitance electronic readout option) rather than the eyepiece to obtain a null indication. With this procedure, the reading sensitivity is 1 to 2 μ Gal. A concave leveling baseplate was used. At each station, the baseplate was pressed to make contact with the elevation survey point and the bull's-eye level bubble centered. The gravimeter was placed on the baseplate and shifted to obtain approximate level prior to final precision level. Microgravimetric survey procedures are discussed in detail in Butler (1980).

Base station drift

38. Reoccupations of the base station are used to correct the gravity data for time variations caused by gravity meter drift and earth-tide gravity variations, with the correction referred to as the drift correction. A drift curve for the survey is shown in Figure 12.

* The meter factor for Model D-25 is 1.1980.



- Gravity Station
- Reoccupied Station and Number of Occupations

- 126 Stations
- 37 Reoccupied Stations (29 percent)
- <u>I</u>_Program I

27



It is assumed that the drift was the same at all stations in the survey grid and linear interpolation is used to compute the correction for stations in each program.

Data corrections

39. Required corrections to gravity data are thoroughly discussed in Butler (1980) and will only be summarized here. Briefly, the data corrections are necessary because of latitude and elevation differences between the gravity stations and the effects of surrounding topography. The corrections applied to the drift-corrected data for the Manatee Springs microgravimetric survey are summarized below:

- <u>a</u>. Latitude correction. The latitude correction to be applied to each station gravity value is given by ± 0.81 × sin $2\phi \times \Delta L \ \mu Gal/m$, where ϕ is the latitude of the site ($\phi \simeq 29.5^{\circ}$) and ΔL is the north-south distance in metres from a reference station. Using station (0,0) as the reference station, $\Delta L = \Delta Y \times \cos 45^{\circ}$, where ΔY is the distance along the NE-SW grid lines from a line passing due east through the survey grid from (0,0). The positive sign is used if the station is south of the line, and the negative sign is used if the station is north of the line. The final expression is $\pm 0.49 \ \Delta Y$ for ΔY in metres or $\pm 0.15 \ \Delta Y$ for ΔY in feet, with the correction given in μ Gal.
- <u>b</u>. <u>Free-air correction</u>. Three corrections (free-air, Bouguer, and terrain) are used to compensate for elevation differences between stations. The free-air correction accounts for the normal free-air vertical gravity gradient and is given by $\pm 308.55 \times \Delta h \ \mu Gal$ for Δh in metres or $\pm 94.04 \times \Delta h \ \mu Gal$ for Δh in feet, where Δh is elevation difference of a station relative to a reference elevation. For the Manatee Springs site, msl is used as a reference elevation.
- <u>c</u>. Bouguer correction. The Bouguer correction accounts for the fact that gravity values in a survey are affected by differing masses of material beneath the stations due solely to elevation differences. The correction is calculated using $\pm 41.91 \times \rho \times \Delta h \ \mu Gal$ for h in metres or $\pm 12.77 \times \rho \times \Delta h \ \mu Gal$ for Δh in feet, where ρ is the density of the near-surface material (1.8 g/cm³ was used for the Manatee Springs site).
- d. <u>Terrain correction</u>. The terrain correction compensates station gravity values for nearby topographic variations which are not included in the Bouguer corrections.

Because of the nature of the problem, the terrain correction is always added (Butler 1980, Telford et al. 1976). Terrain corrections were determined using a terrain template and correction curves as described in Butler (1980). As a consequence of the small topographic variation in the survey area, no station required a correction greater than 5 μ Gal, and the majority of the stations required corrections <2 µGal. The spring boil sink and the Catfish Cabin sink were not of concern since they were a sufficient distance from the survey area; effects due to these sinks would be small (due to distance, low density contrast due to being water-filled, and the relatively small elevation differences) and tend to be roughly equal over the survey area. Stations along a line from (100,260) to (100,340) had terrain corrections of <5 µGal due to the small water-filled sink centered about 100 ft to the southeast of this area. Generally, the effects of terrain corrections as small as for this survey area can be ignored.

Results

Bouguer gravity map

40. Station gravity values which have been correced for drift, latitude differences, and elevation differences (free-air, Bouguer, and terrain corrections) are called the Bouguer gravity values for the station. Figure 13 is the resulting Bouguer gravity contour map for the Manatee Springs site. Two prominent features or trends are evident from examinations of the Bouguer gravity values along profile lines: (a) along the NW-SE profile lines, the gravity values decrease from SE to NW; and (b) along the SW-NE profile lines, the gravity values decrease from the SW reaching minimum values near the center and then increase to the NE.

41. The Bouguer gravity values reflect density variations in the subsurface caused by geologic structures or other lateral variations with varying scales. An important step in analyzing the results of a gravity survey is to separate out a "regional" component of the gravity field, presumably caused by deeper sources or structures, leaving a "residual" component which consists primarily of anomalies caused by structures of interest. Clearly the scale of the regional will depend


on the scale of the survey (Telford et al. 1976, Butler 1980). For small-scale surveys, microgravimetric surveys, like the Medford Cave and the Manatee Springs survey, the regional component can usually be satisfactorily approximated as a constant or as a planar surface. This procedure is called the regional-residual separation, and two methods of accomplishing it will be investigated here: (a) a planar regional will be determined by inspection; and (b) a planar regional, as well as regionals of higher orders, will be determined by polynomial surface fitting of the data.

Bouguer anomaly (residual) map; planar regional field determined by inspection

42. The observed general trend of the Bouguer gravity values to decrease along profile lines from SE to NW, discussed previously, is used to specify a planar regional variation. Using mean values of Bouguer gravity along the bounding SW-NE survey lines, a linear variation of 0.22 μ Gal/ft (0.72 μ Gal/m) is determined, decreasing SE to NW. Clearly this procedure is subjective in nature; the effect of determining a best-fit planar regional field in a least-squares sense will be examined later.

43. Subtracting this planar regional field from the Bouguer gravity map (Figure 13) gives the residual or Bouguer anomaly map shown in Figure 14. Several anomaly features appear in the residual gravity map. The most prominent feature is a broad negative anomaly region (A) trending from SW to NE across the center of the map. Obvious correlation of this negative feature with the known location of the main cavity is demonstrated by the dashed outline of the cavity system in Figure 14. The apparent width of the negative feature is considerably greater than the width of the cavity itself, as expected from examination of Figure 10. Also, the magnitude of the negative anomaly, -20 to -30 μ Gal, is similar to that predicted by the model study of Figure 10; however, if the concentric positive anomaly values (particularly evident to the SW of the negative feature) are considered part of the anomaly caused by the cavity, the anomaly magnitude becomes -50 to -60 μ Gal. This larger





anomaly magnitude is more consistent with the observation that anomalies associated with cavities in karst regions are frequently larger by factors of two than that predicted on the basis of the cavity dimensions alone (Butler 1980, Neumann 1977). Two areas within the broad negative region (A' and A") to the NE of the mapped cavity location exhibit closure, indicating anomalous features in addition to the known cavity system.

44. Other features of the anomaly map (Figure 14) are the positive area in the SW area of the map which attains 40 μ Gal magnitude (B). A positive anomaly area exhibiting closure (C) is adjacent to the central negative feature to the NE and attains magnitude slightly in excess of 10 μ Gal. Another positive anomaly (D) occurs in the NE corner, attaining 30 μ Gal magnitude at (80,400). A negative region (E) with a magnitude of approximately -10 μ Gal separates anomalies C and D.

45. A negative feature (F) occurs as a small southwestern trending extension of the central negative feature (A), with a magnitude of approximately -10 μ Gal. This feature is interesting due to its very close correlation with the small secondary cavity feature mapped in this area (Figure 5) shown dashed in Figure 14. All of the other anomaly features can be expected to change only slightly in detail with a different choice for a planar regional; however, a narrow, elongated anomaly such as F could be altered considerably by the choice of planar regional. This association of F with the small cavity feature will be discussed further after presentation of the regionals determined by polynomial surface fitting.

Bouguer anomaly (residual) map; regional field determined by polynomial surface fitting

46. A more objective procedure for determining the regional cavity field is by fitting a low-order polynomial surface to the Bouguer gravity values (Coons, Woolard, and Hershey 1967, Nettleton 1971). The Bouguer gravity values g_B are considered to be a function of the station coordinates (X,Y), i.e., $g_B(X,Y)$. For example, the equation for a planar surface can be written as $g_B(X,Y) = C_1X + C_2Y + C_3$, where the constants C_1 , C_2 , and C_3 are determined so that the plane best fits the gravity data in a least-squares sense. In principle, surfaces of increasingly higher order can be fit to the Bouguer gravity data. Subtracting any polynomial surface from the Bouguer data results in residual gravity values relative to that surface; residual gravity maps will correspond to shallower anomalous features as higher order polynomial surfaces are subtracted.

47. Figure 15 contains polynomial surface fits to the Bouguer gravity data (Figure 13) through fourth order. The plane determined by inspection dips from right to left through the grid compared to the best-fit planar surface which dips from lower right to upper left. It is worthy of note, however, that the right-to-left gradient through the grid is the <u>same</u> for both surfaces (~0.23 μ Gal/ft). The second-order surface (as well as the third and fourth) show the second prominent feature discussed earlier in regard to the Bouguer gravity values, i.e., that the Bouguer gravity values along SW-NE profile lines decrease from the SW, reaching minimum values near the center, and then increase to the NE. Higher order polynomial surfaces (not shown) will account for increasingly smaller scale features of the Bouguer gravity values, until ultimately a surface is generated for which the residuals will approach zero.

48. Residual ε avity maps will be presented only for the firstand second-order polynomial regionals of Figure 15. The residual gravity map referenced to the first-order polynomial regional is given in Figure 16. Comparison of Figures 14 and 16 reveals that all the major details are similar. All of the lettered anomalies in Figure 14 can be identified in Figure 16. Of course, absolute values of anomalies have changed, but the important relative anomaly values are the same. The most obvious visual difference between Figures 14 and 16 is the shifting of the zero contour separating the anomalous regions A anu F from B in Figure 14 further to the SW in Figure 16. Also, the pronounced "outlining" of the small secondary cavity feature by a continuous zero contour line (see anomaly F in Figure 14) is no longer present in Figure 16, although the presence of a relative negative anomaly apparently





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Figure 16. First-order residual gravity anomaly map

associated with the feature is still evident. Thus, observations made previously regarding Figure 14 and its correlation with the known cavity system are unaffected by the choice of the planar regional field (by inspection or by least-squares surface fitting).

49. Using the second-order polynomial surface of Figure 15 as a regional field results in the residual map shown in Figure 17. A comparison of Figure 17 with Figures 14 and 16 shows that anomalies A, B, and F are removed or significantly altered in character by subtracting the second-order regional field. The obvious interpretation is that the long wavelength anomaly caused by the known cavity system has been accounted for by the second-order regional. Within the area occupied by anomaly A (including A' and A") in Figure 14, negative anomalies apparently corresponding to A' and A" are present in Figure 17. Also, negative anomalies labeled A"' and A"" are identified in an area of otherwise very small negative values in Figure 17. Presumably, these four negative anomaly features are caused by conditions shallower than the main cavity system. For example, as noted in Part III, a wheel of the drill rig fell in a vertical solution pipe about 1.5 ft in diameter and 80 ft deep, creating a hole about 5 ft in diameter at the surface, while drilling boring S-1. Anomaly A"" in Figure 17 seems to be caused by this feature. This association suggests similar causes for anomalies such as E' and E"" in Figure 17.

Verification Drilling

50. Because of time and fiscal constraints, only a very limited number of verification borings was possible. These borings, E-1 through E-6 shown in Figure 17, were placed subsequent to the microgravimetric survey and several cross-borehole surveys in the vicinity of the small secondary cavity feature. These borings were not logged in detail; cuttings were monitored and ease of drilling noted to indicate soft or hard zones or tool drops.

51. Simplified boring logs for these verification borings are given in Figure 18. Boring E-1 was placed to verify the mapped location



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> Second-order residual gavity anomaly map Figure 17.

E-3	Silty sand w/orig. material White, silty sand .5 Residual clay w/LS nodules Limestone, soft, white	E-6	 0-1 Silty sand w/or g. material 1-13 White, silty sand 13-16 Residual clay w/LS nodules 16-105 Limestone, white, soft 90-105 Noticeably softer material, possibly sand- or clay- filled pocket 101-101.3 Tool drop
	0-1 1-14 14-15.5 15.5-50		0-1 1-13 13-16 16-105 90-105 90-105 101-101 orings
E-2	 0-1 Silty sand w/original material 1-8 White, silty sand 8-9 Residual clay w/LS nodules 9-125 Limestone, soft, white 114-114.2 Tool drop 	E-5	 0-10 Silty sand w/original 0-10 Silty sand w/original 10-13 Sandy clay s/LS nodules 13-96 Limestone, soft, white 13-95.9 Cavity 16- 90- 90- 101- Simplified boring logs for verification borings (depth intervals in feet)
T-3	Silty sand w/original material Silty sand w/clay Residual clay w/LS nodules Limestone, soft, white Cavity	E-4	Silty sand w/original material White, silty sand Limestone, soft, white Figure 18. Si
	0-1.5 1.5-15 15-16.5 16.5-125 91.5-107		0-1.5 1.5-5.5 5.5-50

and depth of the secondary cavity feature between borings C-2 and C-3. Figure 19 shows the location, depths, and geometry of the secondary cavity feature as confirmed by boring E-1 and cave divers. The secondary cavity feature is somewhat shallower and larger than originally thought. Boring E-2 was placed to investigate an anomalous response indication on a crosshole electromagnetic survey between borings C-3 and C-4. Borings E-5 and E-6 were placed to verify the mapped location of the main cavity as it crosses under the northeastern side of the survey grid. Finally, borings E-3 and E-4 were placed to investigate the cause of anomalies C and E (Figure 14).

52. Boring E-5, placed at (10,210), intercepted the main cavity, and the dashed cavity location in Figure 14 reflects the proper borehole location relative to the cavity as verified by cave divers. Boring E-6 was originally planned to investigate anomaly A' at (0,260) and the possibility that the main cavity might turn and pass out of the survey area at this location, but it was relocated due to presence of large trees at the desired location. As seen from Figures 14 and 17, boring E-6, as finally located, is at the periphery of anomaly A'. From the boring logs in Figure 18, two features are apparent: (a) 3 ft of residual clay was intercepted by boring E-6 but not by boring E-5; and (b) a soft zone was intercepted in boring E-6 from approximately 90- to 105-ft depth, possibly a clay- or sand-filled pocket. Feature (a) above suggests that anomaly A' itself may be due to a clay-filled depression or pocket which extends to the west in the top of the limestone. The positive anomaly centered at (20,200) in Figure 17 (near boring E-5) may be a limestone high or pinnacle. The broader negative area around A' in Figure 17 and NE of the main cavity location may be due to the clay- or sand-filled pocket below boring E-6, which, due to its depth range, may be a filled cavity connected to the main cavity.

53. Results of borings E-3 and E-4 present a paradox which cannot be readily resolved. On the basis of the nature of the anomalies C and E' in Figure 17 (Figures 14 and 16), it could be expected that the anomalies associated with C are caused by limestone highs (pinnacles) or limestone with greater than typical density contrast while E' would





be a clay pocket, vertical pipe, or other form of shallow cavity. If the feature causing E' is a vertical pipe similar to the one encountered near S-1, it is easy to imagine how a single borehole could miss the feature. The top of rock at boring E-3 is at 15.5-ft depth, which is typical for the site. At E-4, the top of rock is at 5.5 ft. Thus, except for the possibility of missing a vertical pipe with boring E-4, the results seem inconsistent. The possibility of denser limestone at boring E-3 cannot be verified since the boreholes were not cored.

PART V: SUMMARY AND CONCLUSIONS

54. This report presents the results of a microgravimetric survey over a known cavity system at the Manatee Springs site, Fla. The cavity system is water-filled and is at a nominal depth of 90 ft to top, with cavity height ranging from 6.5 ft at boring E-5 to 17.6 ft at S-1 and nominal width of 40 ft. A smaller cavity feature (15 ft in vertical dimensions with a highly irregular cross section ranging up to 10 ft maximum horizontal dimension) branching at about 90 deg to the main cavity was discovered by cave divers during work at the site.

55. Detection of the cavity system by microgravimetric methods was expected to be difficult due to the predicted low amplitude (~-35 μGal) and long spatial wavelength (~200 ft) of the anomaly caused by the cavity system. The survey was planned primarily for the detection of the anomaly caused by the main cavity system. Thus, while the 20-ft station spacing used in the survey is certainly adequate to delineate* larger solution features, 6-ft effective diameter and larger, at depths from 2 to 4 times the effective diameter (depending on geometry),** small solution features such as the small-scale, shallow limestone pinnacles and clay pockets encountered at the Medford site likely would not be delineated (Butler 1980, 1983). Small solution features as well as other small-scale density variations will appear as "lithological gravity noise" on the gravity maps.

56. Results of the microgravimetric survey clearly indicated the location and trend of the main cavity system across the survey area. Anomaly magnitude and width were consistent with predictions based on known geometry and depth of the main cavity system. Also a closed

^{*} Delineation refers to mapping of the details of the anomaly caused by a subsurface feature. Detection of the anomaly caused by a subsurface feature can be accomplished, however, without this detailed mapping.

^{**} Actually, these size/depth rules of thumb will generally be conservative since gravity anomalies will often be greater than that calculated based on measurable dimensions due to secondary effects around the solution features.

negative anomaly superimposed on the broad negative caused by the main cavity satisfactorily delineated the small secondary cavity feature. Mapped locations of the main and secondary cavities were verified by borings and cave divers. These two anomalies were apparent despite a fairly high level of lithological noise.

57. Several other anomalies were apparent on the gravity maps. One small negative anomaly can be correlated to a vertical solution pipe discovered during drilling operations. Only a very limited verification drilling program was possible at the site (six borings) and these were not cored. Of the total of 12 borings at the site, the gravity data were consistent with the subsurface conditions revealed by all but two of the borings. These two borings were in the NE half of the survey area away from the area above the main cavity system. Two gravity anomalies (one positive and one negative), which these borings were to investigate, indicated the possibility of a limestone pinnacle or extended high rock area for the positive anomaly and the possibility of a clay pocket, very shallow cavity, or vertical solution pipe for the negative anomaly. The borings confirmed neither of these possibilities and, in fact, indicated conditions somewhat contradictory to the stated possibilities. Narrowly missing a small-diameter vertical pipe could explain the contradiction for the case of the negative anomaly, but a similar explanation for the contradiction between gravity anomaly indication and boring results is not possible for the positive anomaly, in the absence of core data for the boring.

58. Approximately 12 man-days were required for establishing the site grid system, conducting the elevation survey, conducting the microgravimetric survey, processing data, and producing a hand-drawn residual gravity contour map (Figure 14). At this stage, i.e., after the residual gravity contour map is produced, recommendations regarding site suitability for a given use could be made, assessment of the possibility of solution features at the site could be made, and/or a site drilling program layout could be recommended. For comparison, a twoman rotary rig drill crew (no geologist or inspector) could only produce approximately six 100-ft-deep NX core drill holes for the same

investment of man-time. Six borings at random could do little to characterize a 100- by 400-ft site in a karst region. Microgravimetric surveys coupled with selective drilling are seen to be a time- and cost-effective site investigation procedure for karst areas.

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DEPTH DP	HLLED I		151.1 ft		AL CORE P		FOR BORING	38	-
TOTAL DE	PTH OF	HOLE	170.0 ft						
	DEPTH	LEGEND	CLASSIFICATION OF MATERIA	1.5	S CORE RECOV- ERY	BOX OR SAMPLE NO	REMA (Driffing time, wat) weathering, etc.,	RKS or loss, depth of if sidesilicant)	
•	ft	e	4		•	Ĩ			_
	=	1	sand, fine to very fine white a vellowish br				-8" ID plasti set to 20 fi		
بر	-	1	<pre>white .o yellowish br becoming silty with d</pre>			1	art to 20 M		
		}		r		1			
icc	-]							
Pamlico	10 -	1			le l				
Pr		1	13.3		Sampl				
			Ls residue - sandy clay						
:	=	1	with 1s nodules		101				
]	19.0						
	20 =		Ls, white to creamy,		Run 1		limonitic s	tained join	it
	-		granular, slightly o		94%		RQD = 94%		
	_	1	ted, fossiliferous, weak and soft, pasty		.94 /0	Box 1			
	-	1	gastropods, pelecypo		Run 2		40D = 94%		
	30 -	1 '	and macroforams.		hun 2	1	weathered ze	one with	
	, so =	1			100%		joints		
	-	3			100%				
		3		diam.	4	2			
I	_	1			Run 3	1	RQD = 96%		
	40 -	1							
	=	1			98%				
		1							
e l	-	3			L	3			
Ę	50 _	-			⊇un 4		RQD = 80%		
Ocala	ľ =	4			100%				
00	=		autorite de la Constit		100%				
	- 1	1	-concentrated fossil	zone					
	=	1			Run 5	4	20D - 20 4		
	60 -]					80D = 50%		
	=				99%				
	<u> </u>	1			l				
	=	1			Bun 6		3QD = 24%		
		1			363	5			
	70	1			, n 1, 3	2	Yntertable 1	1.4 ft on	
	=]			Bun 7		14 May	• • • • • 744	
							RQD = 493		
	1 =				955				
	80-	1			Run 8		losing water	return	
	=	1	Ls, slightly cemented			6	RQD = 23%		
	=	1	granular, massive, gr		975				
Ę	-]	brown to white, vuggy	•		┝━━━ ─-			
	-]	harel, predominance a macroforams with some	of	Pun ()		RQD-= 53≸		
williston.	20	1	pelecypois and gastro				0407 - 218		
111		4			.00%	7			
		1							
		1		ici-					
	100 =	1	fied 1s, brown hard, fessil., hard and coff	zones.	P-10-10	8			
		<u> </u>	US EDITIONS ARE OBSOLETE.						

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RILLING	LÕG	(Cont 3					Hole No.	C-2
Manatee	Sprin	gs Sit	e	WES				SHEET 2
	1		CLASSIFICATION OF	· · · · · · · · · · · · · · · · · · ·	% CORE	BOX OR	REMA	AKS
ELEVATION	DEPTH ft	LEGEND	(Dewraption		RECOV-	SAMPLE NO.	(Drilling lime. u. usaibering, etc.	ster luss, depth of . if significant)
	ft 100	<u>ر</u>			e Dum 16			<u> </u>
) Ξ		irregularly silid	ified ls.	Run 10	1		
å	-				31%			
E .	1 =				100,		RQD = 33% -	Run 13
sto	=				Run 12	8	RQD = 62% -	Pup 1h
Williston	<u>110</u>				16%		lost alliwat	
12	=				Run 1	1	core barrel	
_					63%		hole; overre	
	1 -		Ls, hard, partial		ļ		4" steel cas retrieve bas	
	-		grey to brown,		Run 1	•	letilete ba	
	120		vuggy, hard and		92%			
			fossiliferous,	friable	Run 1		Ls has appea	
			zones.		92%	9	vesicular be	
					<u> </u>		RQD = 24% -	Run 15
Ę	130				Run 1	۳ ا	RQD = 61\$	
1s	=				100%	4		
Inglis fm.	=				Run 1	†	RQD = 51%	
Ir	-				96%	10		
	=		Laminated dolom	ite clasts	Run 1	8	RQD = 60%	
	140-		in base		í .			
	=	<u> </u>	142.3		92%	ł i		
	_		Ls, brown, friable	e, hard	Run 1	9	RQD = 38%	
			and soft zone	s, indurated	100%	111		
			partially dol	omitized,	Run 2		RQD = 60%	
	150 -		scattered car materials	uonaceous	100%	\vdash		
	=				Run 2	1	RQD = 38%	
	_					1 1		
į.	1 =				100%	12	RQD = 30%	
r K	160				Run 2	P (
Avon Park fm.	=				82%	├		
u l	=				Run 2	B	RQD = 40%	
Avo					84%	13		
					No			
	170				c <u>ore</u>			
			Bottom of hole 17	0.0				
	1 =				1			
	-							
					l			
	-							
	1 3							
					l			
					1			
					Í			
	-							
-								
	-							
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IG FORM			6P0 19		MORCI			HOLE NO

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			715100	INSTALL			Hole No.	SHERT 1.
PROJECT	ING LO	G Ge	otechnical Lab		WES			OF 2 SHEETS
lanatee	Intin	a Site	, ,				3" core and 6" THOWN (THN @ MAL)	rock bit
LOCATION	(Coordin	sten er Sta	tion)					
DRILLING	AGENCY	_					MATION OF DRILL	
Nobile D	istri	et Corr	os of Engineers		Ling 15		DISTURBED	
HOLE NO.	(Ao shom nbasi	n on drawb		BUR	DEN SAMPL	ES TAKE	N	
NAME OF C			C-3		AL NUMBE			
Carl Moc				16. ELE	VATION GR		TER 9.5 ft on	15 May
DIRECTIO				H. DAT	E HOLE			2 May 80
				17. ELE	VATION TO			
THICKNES							FOR BORING 71	
TOTAL DE			<u>151.9.ft</u> 170.0 ft	19. SIGN	ATURE OF	INSPECT	OR C	
		· · · · ·		L.S.	SCORE	BOX OR	AEMAR	K\$
LEVATION	ft	LEGEND	CLASSIFICATION OF MATERIA		S CORE RECOV- ERY	SAMPLE NO.	(Dettiing time, wate weathering, etc.,	r Jose, depth of It eignificant)
		- <u>-</u>			· ·		R ^H plantin T	
	-		sand, fine, white to ta silt and clay increase				8" plastic II set to 18.1 f	
			with depth	-			crooked so re	
£ (F		1		19.0 ft on 12	
8					อ นี			-
Pamlico	10				sarrple taugei		Run 1 was cav	re-in
5	=				8			
<u>е</u> ,	-					4		
	_		clay tan to brown, with		no fligt	l I		
	=		nodules	. 13	^{ند} ا	19.0		
	20		-19.0		Run 2	2.2.2	RQD = 88%	
			'Ls, pink to gray, hard.		1		lost circula	ion at
	=		weakly cemented, friat massive, fossiliferous	5	88%	Box 1	22 ft; cement	
			(gastropods, pelecypod]	DON 1	have water re	eturn
	=		and macro-forams)			20 5		
	30				Run 3	29.5	RQD = 35%	
	JU							
	_				44%			
						Box		
						2		
	40 -				Run 4	I I	losing more	vater
1					26%		RQD = 18%	- Run 4
	=		•••		20%		RQD = 87%	- Run 5
	_	1			<u> </u>		cemented arou	und outside
	=		Ls, soft, white to creat friable, massive, sca		Run 5		of casing to	stop wash
,			chert fragments, fos		100%	49.6		
	50		ferous as above				Add revert to	o stôp
$\widehat{}$	=				Run 6	Box	water loss	
Ę	_	1			100%	3	RQD = 35% - 1	Run 6
Ę	=							
r f	=				i	ł		
Ocala fm. stal River fm	60 -				L -	60.4		
00	=	1			Run 7		30D = 43%	
уяt	-	£			86%	Вох		
(Cry	-	ł I			0.00	вож Ц		
, j	-	1						
	70-	4			Run 8	1	10D - 255	
	=	1					HQD = 255	
	=	ł			45%			
	-	1	76.5		ļ	76.0		
	=	1	Ls, hard, cream to pi		Run 9		RQD = 39%	
	80	1	well cemented, massive fossil, as above	2,		Вох		
	³³ =	<u> </u>	31.5		84.7	5		
	=		15, grey, hard, massiv	e,		Į		
	-	1	friable, abundance of		[
Ę		1	macroforams gives gra-		Run 10	87.7	RQD = 575	
	an	1	appearance, some gast	ropoda		1		
101	90]	and pelecypods		¥า05			
Williston		}			l			
3	-	Ŧ						
3	=	1			Pon 11	1	EQ1 = 26%	
	100 -	1			57	[Still losing	water
	1836				PROJECT			HOLE NO

	100	(Com 3	iteet) ELEVATION TOP OF HOL				Hole No. C	-3
onci Mar	natee	Springs	3 Site	INSTALLATION	ies			SHEET 2
	DEPTH	LEGEND	CLASSIFICATION OF	MATERIALS	% CORE	SAMPLE	REMA (Drilling time, w	AKS ever luss, depth of
	ft	c	(Descriptus d	•,	ERY	NO. f	weathering, etc.,	of upupiant)
	100 _				1	Box	·	·
	-	Voia -	106.0			6	void from 10	
		-1010	Ls, irregularly	silicified,			106.5; lost	all water
đ	110	k ,	brown to gray,	hard and	Run 13	2	RQD = 6%	
	110	Voia	soft zones, fos (gastropods, pe		17%	Box	void from 10	19.0 to 114.
Williston	-	\bigtriangleup	macroforams), f	riable,		7		,
1			massiv e			126.3	RQD = 38%	
3	=	1			Run 1	\$		
	120 _				70%			
	=	<u> </u>	115.7				1	
]	Ls, grey to brown partially dolon		L			
	1.20 -	1	few gastropods		Run 1	Box	RQD = 44%	
ė	130	1	cypods, some gi zones	anular	62%	8		
18	=	1			02,0			
Inglis fm.		1	laminated dolor	nite		138.0	core barrel hole; over	
E I		1	clasts		1		with 4" ste	
	140				67%		to retrieve	barrel.
	-	[Ls, brown to grey	. friable.	1	Box	RQD = 48%	
	-	1	hard and soft	cones,		9	RQD = 69%	
	-		partially dolor carbonaceous ma		Run 1	\$	NQD - 094	
	150-	1	scattered in u		92%	153.0		
]				153.0	İ.	
2		1				ł	RQD = 52%	
Avon Partk fm.			[Run 17	Box	(
5	160				1	10		
g	=				92%			
¥					no		hole reamed with 6" rock	
			Bottom of hole 1	70.0 ft	core			
	170						1	
		1						
		1						
	=	1						
		1						
	-							
	_	1			1	1	1	
	=	1						
	-]	ļ			ļ	ļ	
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	-	1						
	1 3							
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	-	1				1	1	
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	-	1			1	i I		
]						
	- 1	1	ĺ		1			
]						

R

-	ING LO	e 1	VILION	INSTALL				INCEY 1
PROJECT		- . [(Geotechnical Lab	10. SIZE	AND TYPE	-	" rock bit	OF SHEETS
	Sprin	zs Site	<u></u>	11. BATU	HI FON EL	EVATION	SHOWN (THE	,
				12 MANU	FACTURE	A'S DESIG	NATION OF DRILL	
Mobile			ps of Engineers		.ng 150		DISTURSED	UNDISTURBED
HOLE NO.	(As show		ne milo		NEN BANPI	OVER-	N	
NAME OF	ORILLER		,i,,,			R CORE B		
Carl Mo Directio			ert Wens	+	<u></u>			16 June
E VERTI			OES. FROM VERT.	H. DATI	I HOLE			' June 50
. THICRNES	S OF OVE		10.0 ft	the second se		P OF HO		
. DEPTH DA	HLLED H	TO ROCK	150.0 ft			INSPECT	POR BORING	<u>`</u>
TOTAL DE	PTH OF	HOLE	<u>100.0 ft</u>	<u> </u>		100× 00	AEWA	
LEVATION	ft b	LEGEND	CLASSIFICATION OF MATERIA (Description)	ALS	S CORE RECOV- ERY	BOX OR SAMPLE NO.	(Detting time, main mailering, etc.	er loss, depth of , il significant
			sand, fine, white				8" ID plast	
	=						set to 31 f	eet
ė	10							
2	=							
Pumilico fon	=							
Ptam	=							
	=							
	20 =		clay, sandy, 15 nodules	s			casing bent 20 ft joint	
	=		··· ·····				0	
	30		Ls, white, fossiliferout	us				
	I Ξ				ĺ	1		
	1 =							
	40-		hard Julia				lost all wat	
Ē			hard ledge				cemented hol sacks cement	
Ocala fm		1					water return	
0c.s	1 3	ł			ci a			
	50	1			Samples			
	=	1			3 8 S			
					Ň			
	=							
	60 -					1		
	=							
			hard ledge					
	=		THE REAL POINT					
	70					ļ		
	=							
· ?—? —	80 -	1	Joft zone at 81 ft					
		1						
:		1						
Williston Mm.	=	1				{		
nct	₽° —	1						
13(=	1						
111	- 1							
28		1						
	1836	1	Lilleffied williaton f	`m.			Drilling much	harier.

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WILLING.	LOG	(Cont S	heet)				Hole No.	C-4
TORON		Spring		INSTALLATION	WES			SHEET 2 OF 2 SHEETS
		pring		OF MATERIALS	* CORF	BOX OR	REI	AARKS
ELEVATION	DEPTH	LEGEND		ne of materials	RECOV	SAMPLE NO.	(Drilling time.) weathering, et	water luss, depth of 1 . if ugstificant)
	108	<u> </u>		_d		<u> </u>		<u>R</u>
	-	$\overline{7}$	101.0					
	-	I V	cavity					
8	-	$I \land I$	_					
5		<u>r </u>	108.0					
s Ç	110-	1	iard dr	illing				
Williston fm	=		111.8					
5		\sim	cavity 114.8			1		
		1	hard drilli					
-1-?		1	1	-			120' - ce	mented hole
	120-]	alternating zones to bot	hard and soft			with 30 s	acks of
		}	201288 00 000	CON OI NOIE			cement mi	
	-	1	1		1	:		drilling ent set up
ė		1					to 105.0'	
-	130-	1				1		
Inglis tha.	= =	1				1		
2 U	-]			i			
~]	}		1			
	-	1						
- ? ? -	140	1			Samples			
		1				1	ļ	
		1	1		v.	1		
	=	1	ļ		No			
	=	1						
	150-	1						
	-	}					Ì	
	-	1						
	=	1						
	160	1			Ì		1	
ė	=	1]		
ч.]					165.0 ft	- cemented
E E	=		ļ					40 sacks of
Avon Park	=	1					cement, s	
<u><u></u></u>	170-	1						cemented additional
	1 -	1				}		cement, set
]					up to 102	
	_							
	180	·						
	180	1	Bottom of hol	e				
		1						
	_	}	1					
-	-						Note: hole	is crooked
	-						1	
		1	ļ					at 11.0 ft
		1				1	on 16 June	
		ļ						
	-	1				I		
	-							
	-	}			İ			
						[
	-							
	-	1	ł		1	1		
			1					
	~							
NG FORM	<u> </u>	<u> </u>	L		PROJECT	<u> </u>		HOLE NO
	1836-	- A		CPD 1961 07-124 143			prings Site	C-4

		100	VISION	INSTAL	ATION		SHEET 1
DRILL	ING LO	6	destechnical Lab		TEC		OF SHEETS
untee				10. SIZE	UN FOR EL	EVATION	HAWN (THIN - MIL)
LOCATION				12 101-			
DAILLING				- Mada	ing 150	າດ	
HOLE NO.	(Ae aheu		es of instinears	12. TOT	AL NO. OF	OVER- LES TAKE	
NAME OF		_	C=5	14. TOT		A CORE B	OXES
orbert	$\mathcal{O}_{X} \subset \mathcal{O}_{X}$			18. ELE	VATION G		111 11 511 10 6115
DIRECTIO			DES. FROM VERT.	16. DAT	E HOLE	27	May 80 30 May 80
THICKNES					VATION TO		the second second second second second second second second second second second second second second second s
DEPTH OF					AL CORE P		V FOR BORING 0
TOTAL DE	PTH OF	HOLE	170.0 ft	1			
EVATION		LEGEND	CLASSIFICATION OF MATERIA (Description)	LS	S CORE	BOX OR SAMPLE NO.	REMARKS (Drilling time, mater lose, depth of weathering, etc., if significant)
•	1°4.	· ·	<u>المعامل المعامل u>		— •	Ĩ	• • • • • • • • • • • • • • • • •
<u>.</u>			sand, white, fine				8" ID plastic casing set to 27.0'
artic Tha							
<u> </u>	-		7.0 it		1		hole started 2.5 ft
	10 _		clay, sandy with 1s not	lules			west drilled to 26.5 f
	=					Į	before hitting rock; lost 8" rock bit in
	_	ł	17.01				hole; filled with sand
	=		17.0' Ls, creamy to white.				and 1 sack cement and moved to present
	20		soft, fossiliferous				location
	=						
	=						
					1		
					1		
	30						Lost all water return
:] =				1		
đ	-						Cemented hole with 4
Jeala					1		sacks cement; 60 to 80%
oc	40						water return after
					1		cementing and using drilling mud
							-
	-		hand some		1		
	50		hard zone				
	-						
			very soft aone				
	=						
	60 <u>-</u>						
		ł					
	70 -						
	=		hard cone of 71 and 74	ft			
							207 water return
	=	;					ilo, water return
? - ? -					ſ		
	=						
	1, 3						
.:	<u></u>						
ب. •		1					
		1					
÷	_ =	1					
	18 36	1	L		PROJECT	l	HOLE NO

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RILLING	LOG	(Cont 1	inet) ELEVATION TOP OF HOL				Hole No.	<u>c</u>
Manat	ee Spr	ings S	Site	INSTALLATION	ÆS			SHEET 2
LEVATION	н	LEGEND	CLASSIFICATION OF (Description		& CORE RECOV- ERY	BOX OR SAMPLE NO.	REM. (Drolling time. u weathering. etc	ABKS
•	100	<u>د</u>	d		·	<u>(</u>		R
			top of silicifi	ed			very hard d	rilling
ton			Williston at 1					
i e								
hilliston fm.	110-							
· _?							lost all wat ft; cemented	
							of cement; 1	
	120		alternating hard				water return drilling thr	
			201168 00 00 0000	OI NOTE			dritting on	ougn cement
÷								
4								
Inglis fm.	130							
Inf					;			
			ļ				1	
?_ ?_	140							
	-		1			ł	}	
	150							
	=							
ė								
Avon Park fm.	1		[
Pa	160-							
NOD								
A							water table	at 11.5
							on 16 June	
	170	···	Bottom of hole			 	cemented hol	
	1						sacks of cen hold it oper	
							up to 120 ft	but
							questionable top 25 ft wa	
								D GCO up
	-							
	-							
1	-	I			i			
	-							
	-							
					1			
	i							
	1							
	-				1			
]					_			
G FORM	1836-	A		1 DI - 18+ 8+1	PROJECT	-	- prings Site	C-5

In accordance with letter from DAEN-RDC, DAEN-ASI dated 22 July 1977, Subject: Facsimile Catalog Cards for Laboratory Technical Publications, a facsimile catalog card in Library of Congress MARC format is reproduced below.

Butler, Dwain K. Cavity detection and delineation research : Report 4 : Microgravimetric survey : Manatee Springs Site, Florida / by Dwain K. Butler, Charlie B. Whitten, and Fred L. Smith (Geotechnical Laboratory, U.S. Army Engineer Waterways Experiment Station). -- Vicksburg, Miss. : The Station ; Springfield, Va. : available from NTIS, 1983. 47, [12] p. : ill. ; 27 cm. -- (Technical report ; GL-83-1, Report 4) Cover title. "March 1983." "Prepared for Office, Chief of Engineers, U.S. Army under CWIS Work Unit 31150." Bibliography: p. 47. 1. Geophysical research. 2. Gravimeter (Geophysical instruments). 3. Manatee Springs (Fla.) I. Whitten, Charlie B. II. Smith, Fred L. III. United States. Army. Corps of Engineers. Office of the Chief of Engineers. IV. U.S. Army Engineer Waterways Experiment Station.

Butler, Dwain K. Cavity detection and delineation research : ... 1983. (Card 2) Geotechnical Laboratory. V. Title VI. Series: Technical report (U.S. Army Engineer Waterways Experiment Station) ; GL-83-1, Report 4. TA7.W34 no.GL-83-1 Report 4

