**NORSAR** 

ROYAL NORWEGIAN COUNCIL-FOR-SCIENTIFIC AND INDUSTRIAL RESEARCH

Scientific Report 1-78/79

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9 FINAL TECHNICAL SUMMARY

1 April 30 September 1978

D. Rieber-Mohn (ed.)

The Norwegian Seismic Array (NORSAR). Phase 3.

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<sup>1</sup>This report describes the operation and research activities at the Norwegian Seismic Array (NORSAR) for the time period from 1 April to 30 September 1978. In general, the array operation is characterized as stable, with little change from the previous reporting period. There has been a slight improvement in the performance of the data recording and online detection processing relative to the previous reporting period (uptime increased to 93.7% from 91.6%). The Special Processing System (SPS) is still the

major cause of the breaks in operation (75 out of 139), the longest one lasting for more than 4 days. Statistics from short sample intervals indicate that the number of Online detections average about 200 per day with the present threshold setting, while the On-line Event Processor processes about 25 events per day for transmission to the SDAC. The average number of analyst-retrieved and accepted events has been 10.8 per day during the period. k

The work load for the data center has been too large to handle within the regular working hours. Long jobs must be run during evenings, and additional part time help has been employed, in order to catch up with Data Retention processing outside working hours. NORSAR personnel have gradually taken over more responsibility for maintenance and error corrections of hardware (tape drives, SPS, etc.). The performance of the array's communications circuits has been satisfactory during the period, with three outages lasting for more than one hour and affecting more than one subarray simultaneously. One case of attenuation distortion has been discovered on the ARPANET SDAC communications circuit. No changes have been made to the TIP or its connections within the reporting period. There have been few modifications to the Online Detection Processor, the most important being implementation of logic to stop the system when erroneous timing is discovered, and a change in the selection of the channels that are automatically transmitted to SDAC with each processed event. An off-line interactive bulletin editing system has been implemented

In the beginning of the reporting period, a Teledyne-Geotech S-500 seismometer was tested out within the array configuration (subarray 06C). It was removed on 12 April. Also, on 5 April, a Teledyne-Geotech S-13 threeaxis seismometer was installed at subarray OlA. The components were connected to standard NORSAR channel equipment. The relatively large number of maintenance visits in this period (average 9.3 per subarray), is mainly due to preventive maintenance work such as painting, amplifier replacements and instrument adjustments.

The research work at NORSAR in the period is described in eight subsections in Chapter VI. They cover evaluation of the MORSAR Detection and Location capabilities and other research conducted under NTNF's contract with ARFA, as well as projects sponsored by Norwegian authorities. In the evaluation of detection capability in Section VI.1 it is found that the average monthly number of events reported in the edited NORSAR seismic bulletin now is reduced to about 60% of the level when 22 subarrays were in operation, or to 65% if one compensates for the increased system downtime. This corresponds to a reduction in the 50% incremental detectability threshold of about 0.2 m units, which agrees well with theoretical expectations. The event location capability of the present NORSAR array is evaluated in Section VI.2, and the median location difference between NORSAR and USGS epicentral solutions is found to be 230 km in the teleseismic distance range. Section VI.3 presents results on our continued seismic magnitude studies, in particular regarding the M<sub>s</sub>:m<sub>b</sub> relationship. Sections VI.4, VI.5 and VI.6 present initial results from a detection and discrimination study based upon near-field observations. The data base for these studies will be greatly expanded during the forthcoming year. A study of the general question of to what extent tectonic features observable at the surface have counterparts in the deeper parts of the lithosphere is presented in Section VI.7. Section VI.8 presents initial data analysis results from the Stiegler's

Corge Seismic Network installed by NTNF/NORGAR for the purpose of seismic risk studies in Tanzania. Finally, Section VI.9 presents results from the continued investigation of the seismicity of Svalbard.

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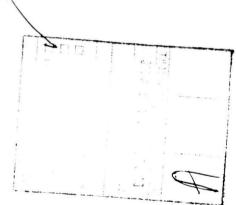
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#### I. SUMMARY

This report describes the operation and research activities at the Norwegian Seismic Array (NORSAR) for the time period from 1 April to 30 September 1978. In general, the array operation is characterized as stable, with little change from the previous reporting period. There has been a slight improvement in the performance of the data recording and online detection processing relative to the previous reporting period (uptime increased to 93.7% from 91.6%). The Special Processing System (SPS) is still the major cause of the breaks in operation (75 out of 139), the longest one lasting for more than 4 days. Statistics from short sample intervals indicate that the number of Online detections average about 200 per day with the present threshold setting, while the On-line Event Processor processes about 25 events per day for transmission to the SDAC. The average number of analyst-retrieved and accepted events has been 10.8 per day during the period.

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#### II OPERATION OF ALL SYSTEMS

#### II.1 Detection Processor (DP) Operation

There have been 139 breaks in the otherwise continuous operation of the NORSAR Online DP system within the current 6-month reporting interval. The uptime percentage is 93.7%, which is a slight improvement over the 91.6% reported for the previous interval (October 77-March 78). Fig. II.1.1 and the accompanying Table II.1.1 both show the daily DP downtime for the days between 1 April and 30 September 1978. The monthly recording times and up percentages are given in Table II.1.2. As can be seen from Table II.1.1, the longest break occurred from mid-day 18 May to mid-day 22 May and lasted more than 96 hours. A power break brought the SPS down and acted as a catalyst to bring forward an inherent SPS hardware error, which subsequently was found and corrected. In addition, the stops related to the SPS alone were 75, so that obviously this component maintains its status as the weakest link in the system. The breaks can be grouped as follows:

a)	SPS malfunctioning	:	75
b)	Error on the Multiplexor Channel	:	16
c)	Stops related to possible program errors	:	12
d)	Maintenance stops	:	12
e)	Power jumps and breaks	:	10
f)	Hardware problems	:	5
g)	Magnetic tape drive problems	:	4
h)	Stops related to system operation	:	3
i)	TOD error stops	:	2

The varying performance of the SPS is reflected in the fact that the stops caused by this component occur in 'bursts', with long quiet periods in between, when the SPS performs normally. (See also comments elsewhere in this report.)

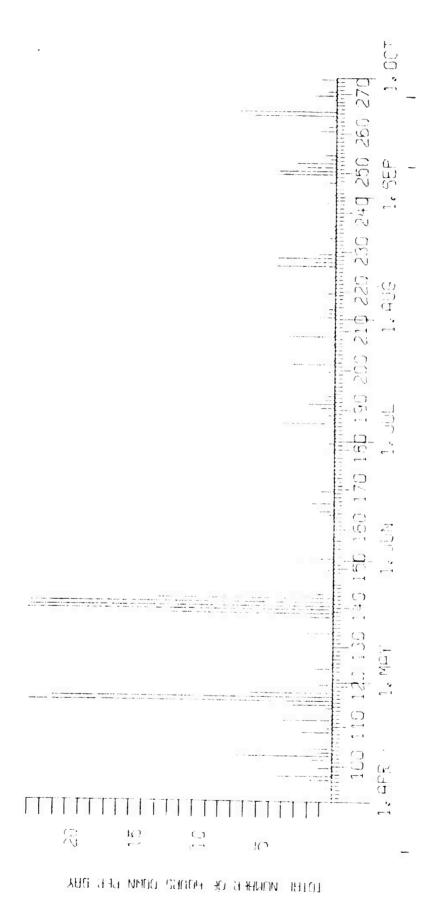
The number of stops caused by the error on the multiplexor channel (category b) is exactly twice the corresponding number for the last reporting period. However, as can be seen from Table II.1.1, 12 of the 16

stops in this category occurred within daily working hours, so that the system was restarted promptly. The probable cause of this type of error is the use of the printer, which seems to put a too heavy load on the multiplexor channel. Neither the printer nor the operator console are used, however, outside working hours.

The relatively high number of program errors reflects partly the strain on the 360-based Online system when the SPS is behaving abnormally. Any such situation causing all the available queue storage to be used will, for instance, show itself as a program error in the queue block leasing routine. The continuous effort of NORSAR personnel to maintain a system with optimum performance is reflected in the number of maintenance stops (category d), which shows an increase from last reporting period. This is partly caused by the more intensive care necessary for the SPS subsystem.

The total downtime for this period was 274 hours 51 minutes. The mean-time-between-failures (MTBF) was 1.2 days, as compared with 1.5 days for the previous reporting period.

The average Detection Rates (number of detections per day) for the NORSAR Online System's Detection task, and the average Event Rates (number of events per day) for the Online Event Processor (OEP) subtask in the same system are given in Table II.1.3. For practical reasons, we have not been able to provide comprehensive statistics during this reporting period; instead we present the data for short time intervals in the beginning, middle and end of the period. The cases of Coherent and Incoherent Detection/Event processing have been separated. As can be seen, the Event Rates are significantly less than the corresponding Detection Rates. This is to be expected, since the OEP has somewhat higher signalto-noise ratio thresholds for its event candidates than the corresponding detection thresholds (3.6 and 2.4 versus 3.16 and 1.6, respectively). This difference is sufficient to eliminate most of the noise detections from consideration. The table also seems to indicate that the total number of OEP evens is relatively constant and independent of the Detection Rates (about 24 events per day). However, insufficient statistics prevent any firm conclusions to be drawn, even when, as in this case, the sampled intervals are spread out in time. As an example of the variability of daily Detection and Event Rates, Fig. II.1.2 shows these parameters for (A) a 5-day interval in July and (B) for a 5-day interval in September.



g. II.1.1 Online System Downtime, April - September 1978

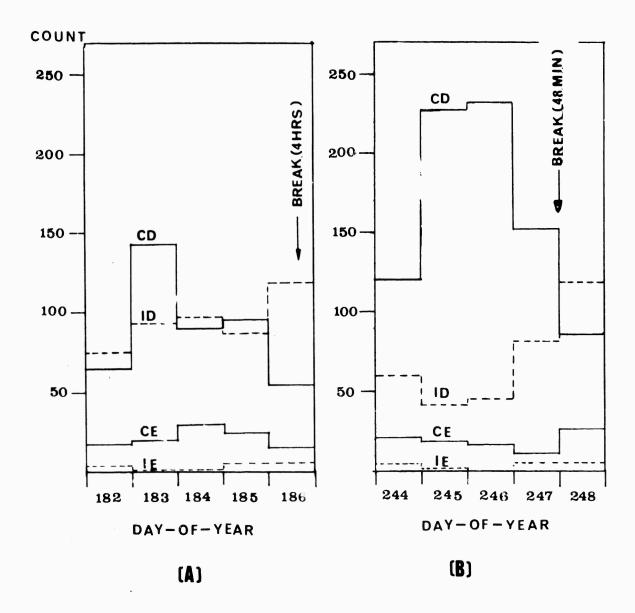


Fig. II.1.2 Variability in the coherent and incoherent detection and event rates. (A) 5 days in July; (B) 5 days in September.

CD: Coherent Detections

CE: Coherent Events

ID: Incoherent Detections

IE: Incoherent Events.

LIST	OF BRE	AKS	IN DP	PROCESSING THE LAST HALF-YEAR
UAY	START		STUP	COMMENTS
94	11	53	11	57 ERRUNEOUS SPS TIMING
94		20	15	46 SPS STOP
96		36	10	O SPS NO DETECTIONS
96		44		6 ERRONEOUS SPS TIMING
96	15	46	16	29 SPS STOP
96		20	21	29 ERRUNEOUS SPS TIMING
97		42	10	12 SPS CHECK
97	14	<b>3</b> 8	<b>.</b> 5	7 SPS STOP
97		35	21	55 SPS STOP
99		14		29 MPX/LATE
99			20	26 SPS STOP
99	20	56		6 ERRONEOUS SPS TIME
101	5		6	
			12	
101			16	
	6			57 SPSZEOC TESTS 38 SPS CHECK
		26		51 SPS STOP
				4 SPS STOP
	18			37 MAINTENANCE
	20	45		
		9		15 SPS/EUC TESTS
104	12		12	37 SPS STOP
104			13	22 SPS STOP
	14		14	36 SPS STOP
104	15	37	16	46 SPS STOP
107		10	8	35 ERRONEOUS SPS TIME
108	11	10		26 SPS STOP
£08	21	38		56 SPS STOP
109	9	2.	9	35 SPS STOP
.09	16	i	16	13 MAINTENANCE
1.0	12	35	12	45 SPS STOP
111	7	32	10	13 SPS MAINTENAFILE
111	20	55	22	10 SPS STOP
1.2	11	26	11	42 ERRONEOUS SPS TIME
114	9	13	9	49 SPS MAINTENANCE
114	22	12	22	21 ERRUNEOUS SPS TIME
115	13	4.3	14	12 POWER BREAK
115	14	2 C	14	31 SPS DOWN WHEN ECC UP
115	19	50	20	50 SPS STOP
115	21	30	22	33 SPS STOP
115	23	36	24	O SPS STOP
116	0	O	¿ <b>4</b>	O SPS STOP

TABLE II.1.1

Sheet 1 of 4

LIST	OF BREAK	S IN DP	PRUCESSING THE LAST HALF-YEAR
DAY	START	STOP	COMMENTS
117 118 118 119	0 0 8 2 23 30 0 0	13 24	
120 124	11 13 0 .5	12	23 MPX/LATE 41 SPS STOP
125 133 136	6 41 12 5	7 8 12	
137 138 138		7 24	
139 140 141	0 0	24 24	O SPS ERROR O SPS ERROR O SPS ERROR
142 142 142	0 0 13 40 14 37	13 15	22 SPS ERROR 54 SPS ERROR 31 SPS ERROR
142 143 143	19 17 0 13 8 19	0 9	46 SPS 3 SPS
143 145 146	9 . 6 7 53 23 25	8 è	25 MAINTENANCE O MPX/LATE 33 POSSIBLE PROGRAM ERRUR
140 149 151	2 11 14 1 15 i	14	56 SPS STOP 28 SPS STOP 52 POW: F. BREAK
152 152 157	6 5 8 1 2 7 3 7 3 7 3 7 3 7 3 7 3 7 3 7 3 7 3 7	9	26 SPS STOP 16 SPS STOP 13 SPS STOP
163 163 164	12 6 12 5 7 46	13	31 MPX/LATE 3 MPX/LATE 51 POSSIBLE PROGRAM ERROR
164 164 166	10 50 13 16 3.	10 7 14	58 MPX/LATE 5 MAINTENANCE 20 SPS STOP
169 179 181	2 3 2 2 3 4 4 5 5 5 5 6 5 6 5 6 5 6 6 6 6 6 6 6 6		40 SPS STOP 27 POSSIBLE PROGRAM ERROR 33 MPX/LATE
184 186 187	9 21 17 59 12 54	7 9 5 21	35 MPX/LATE 55 1052 ERROR, A TO B 5 B TO A
188 188 188	9 59 11 48 13 48	10	24 POSSIBLE PROGRAM ERROR 6 POSSIBLE PROGRAM ERROR 54 FOD ERROR
190 191 191	6 50 11 20 11 50	7	42 SPS STOP 34 MAINTENANCE, A TO B 21 MAINTENANCE
191	15 4		10 HARDWARE ERROR

TABLE II.1.1

DAY START STOP COMMENTS	LIST	OF BREAKS	IN Do bb	RUCESSING THE LAST HALF-YEAR
193 10 53 11 36 8 TO A 199 20 2 20 13 PUSSIBLE PROGRAM ERROR 199 22 12 22 34 MPX/LATE 201 14 5 17 29 POWER BRAK 202 10 23 10 41 MPX/LATE 206 14 13 14 34 POSSIBLE PROGRAM ERROR 208 10 5d 14 35 POWER BREAK/PROG. MAINT 209 8 52 d 58 ARPANET HANGUP 213 20 34 21 48 POWER BREAK 214 6 14 6 44 SPS/EDC TESTS 215 10 35 11 31 DISK CONTROLLER 216 6 4. 7 1 SPS/EDC TESTS 219 7 13 7 55 POWER BREAK 221 12 22 12 27 SPS/EDC TESTS 223 17 53 18 12 SPS/EDC TESTS 224 17 53 18 12 SPS/EDC TESTS 225 10 20 12 51 MAINTENANCE 227 12 6 13 21 PUSSIBLE PROGRAM ERROR 228 3 54 7 41 POSSIBLE PROGRAM ERROR 229 8 10 9 33 MAINTENANCE 233 6 26 6 6 6 MPX/LATE 234 7 37 7 45 SPS STOP 248 C 0 C 1 SPS STOP 249 10 5 10 39 SPS STOP 249 10 5 10 39 SPS STOP 249 11 35 12 44 SPS STOP 249 10 5 10 39 SPS STOP 249 11 35 12 44 SPS STOP 249 20 15 20 43 SPS STOP 249 17 47 20 9 SPS STOP 249 20 15 20 43 SPS STOP 250 19 58 21 11 SPS STOP 250 19 58 21 11 SPS STOP 251 12 55 13 17 SPS STOP 252 11 28 12 40 SPS STOP 251 12 55 13 17 SPS STOP 252 11 28 12 40 SPS STOP 251 12 55 13 17 SPS STOP 252 11 28 12 40 SPS STOP 251 12 55 13 17 SPS STOP 251 12 55 13 17 SPS STOP 252 11 28 12 40 SPS STOP 252 11 28 12 40 SPS STOP 253 11 28 12 40 SPS STOP 251 12 55 13 17 SPS STOP 252 11 28 12 40 SPS STOP 253 11 28 12 40 SPS STOP 255 11 28 55 TOP 255 11 28 52 TOP 255 11 28 55 TOP	DAY	START	STOP	COMMENTS
221 12 22 12 27 SPS/EUC TESTS  223 17 53 18 12 SPS/EUC TESTS  226 6 47 8 48 POWLR BREAK  226 10 20 12 51 MAINTENANCE  227 12 6 13 21 POSSIBLE PRUGRAM ERROR  227 23 40 24 0 POSSIBLE PRUGRAM ERROR  228 3 54 7 41 POSSIBLE PRUGRAM ERROR  229 8 15 9 33 MAINTENANCE  233 6 26 6 56 MPX/LATE  233 6 26 6 56 MPX/LATE  234 7 37 7 45 SPS STUP  244 7 37 7 45 SPS STUP  247 23 13 24 0 SPS STUP  248 0 0 0 0 SPS STUP  249 10 5 10 39 SPS STUP  249 10 5 10 39 SPS STUP  249 11 35 12 44 SPS STUP  249 20 15 20 43 SPS STUP  250 4 0 4 54 SPS STUP  250 19 6 19 46 SPS STUP  250 19 58 21 11 SPS STUP  250 23 48 74 0 SPS STUP  251 10 5 10 38 SPS STUP  252 11 28 12 11 SPS STUP  253 19 58 21 11 SPS STUP  255 10 13 17 SPS STUP  256 27 SPS STUP  257 12 55 13 17 SPS STUP  258 STUP  259 27 SPS STUP  250 27 SPS STUP  250 28 SPS STUP  251 12 SPS STUP  252 11 28 12 11 SPS STUP  255 11 28 12 11 SPS STUP  257 11 28 12 11 SPS STUP  257 11 28 12 11 SPS STUP	192 193 199 199 201 202 206 208 209 213 214 215 216	12 55 10 53 20 2 22 12 14 5 10 23 14 13 10 5d 8 52 20 34 6 14 10 35 6 41	13 46 11 36 20 13 22 34 17 29 10 41 14 35 8 58 21 48 6 44 11 31	MAINTENANCE B TO A B PUSSIBLE PROGRAM ERROR MPX/LATE POWER BRAK MPX/LATE POSSIBLE PROGRAM ERROR POWER BREAK/PROG. MAINT B ARPANET HANGUP B POWER BREAK SPS/EOC TESTS L DISK CONTROLLER L SPS/EOC TESTS
233 6 26 6 56 MPX/LATE 235 7 35 7 41 MPX/LATE 236 13 46 14 4 DISK HARDWARE ERRUR 244 7 37 7 45 SPS STUP 247 23 13 24 0 SPS STUP 248 0 0 0 1 SPS STUP 249 10 5 10 39 SPS STUP 249 11 35 12 44 SPS STUP 249 17 47 20 9 SPS STUP 249 20 15 20 43 SPS STUP 249 20 15 20 43 SPS STUP 250 9 7 10 8 TAPE CONTROLLER FAILURE 250 10 13 10 38 SPS STUP 250 19 58 21 11 SPS STUP 250 23 48 74 0 SPS STUP 251 0 0 1 10 SPS STUP 251 11 28 12 11 SPS STUP 252 6 32 7 22 SPS STUP 252 11 28 12 40 SPS STUP 252 21 33 22 17 SPS STUP	221 223 226 226 227 227 227 228 228	12 22 17 53 6 47 10 20 12 6 23 40 0 0 3 54	12 27 18 12 8 48 12 51 13 21 24 0 0 57 7 41	7 SPS/EUC TESTS 2 SPS/EUC TESTS 3 POWLR BREAK L MAINTENANCE L PUSSIBLE PRUGRAM ERRUR 7 POSSIBLE PROGRAM ERRUR L POSSIBLE PROGRAM ERRUR L POSSIBLE PROGRAM ERRUR
249 11 35 12 44 SPS STOP 249 17 47 20 9 SPS STOP 249 20 15 20 43 SPS STOP 250 4 0 4 54 SPS STOP 250 9 7 10 8 TAPE CONTROLLER FAILURE 250 10 13 10 38 SPS STOP 250 19 6 19 46 SPS STOP 250 19 58 21 11 SPS STOP 250 23 48 74 0 SPS STOP 251 0 0 1 10 SPS STOP 251 11 28 12 11 SPS STOP 252 6 32 7 22 SPS STOP 252 11 28 12 40 SPS STOP 252 21 33 22 17 SPS STOP	233 235 236 244 247 248	6 26 7 35 13 46 7 37 23 13 0 0	6 56 7 41 14 4 7 45 24 0	MPX/LATE MPX/LATE DISK HARDWARE ERRUR SPS STUP SPS STUP SPS STUP
6.16 (1) 91 (4 U NY NIIP	249 249 250 250 250 250 250 251 251 251 252 252	11 35 17 47 20 15 4 0 9 7 10 13 19 6 19 58 23 48 0 0 11 28 12 55 6 32 11 28	12 44 20 9 20 43 4 54 10 8 10 38 19 46 21 11 7 12 13 17 7 22 12 40	SPS STOP

LIST	OF	BREAKS	IN	UP	PROCESSING	THE	LAST	HALF-YEAR
------	----	--------	----	----	------------	-----	------	-----------

DAY	STAR	ſ	STUP		COMMENTS
254	17	36	18	28	SPS STOP
260	10	13	11	26	POWER FAILURE
261	13	12	13	37	MPX/LATE
261	13	40	1.3	52	TAPE DRIVE FAILURE
263	8	13	8	37	MPX/LATE
264	9	3	9	27	MPX/LATE
264	17	51	24	0	TAPL DRIVE/CONTRULLER F
265	. 0	, i	7	48	TAPE DRIVE/CONTROLLER F
268	13	4 i	14	0	EOC. POWER DOWN
269	4	44	6	22	TOD STOP
270	1	33	2	15	SPS STOP
273	7	37	ಕ	49	TAPL DRIVE FAILURE

TABLE II.1.1

Sheet 4 of 4

MONTH	DP UPTIME (Hrs)	DP UPTIME	NO. OF DP BREAKS	NO. OF DAYS WITH BREAKS	DP MTBF* (Days)
APR	635.1	88.2	45	21	0.6
MAY	633.4	85.1	18	16	1.4
JUN	712.9	99.0	12	8	2.3
JUL	725.8	97.5	19	14	1.5
AUG	727.0	97.7	16	14	1.8
SEP	683.0	94.9	29	18	0.9
THE TOT PERIOD	AL 4117.2	93.7	139	91	1.2

<sup>\*</sup> Mean-time-between-failures = (Total uptime/No. of Up Intervals).

TABLE II.1.2
Online System Performance
April-September 1978

		ОН.	3	4	3
AVERAGE OEP EVENT RATE (EVENT/DAY)		INCOH			
		сон.	18	24	21
		INCOH. COH.	53 18	110	81
AVERAGE	DETECTION RATE (DET./DAY	COH: INCO	103	93	150
	NO. OF DETECTOR OF EVENTS RATE (DET		13	21	20
INCOHERENT	AVG. NO. OF NO. LTA DETECTIONS O		220	651	489
	AVG. LTA		737	797	622
LNE	OF P EVENTS		75	141	124
COHERENT	AVG NO. OF LTA DETECTIONS		426	547	905
	L		631	547	702
INTERVAL	UPTIME (HRS)		66	142	145
INTERVAL			92/11.26- 96/15.40	181/08.13- 187/12.21	243/08.04- 249/09.59

TABLE II.1.3

Detection and Event Statistics for the Online (DP) System, Selected Intervals

#### II.2 Event Processor Operation

The operation of the Event Processor was, after a one-year break, resumed as of 1 October 1977. Some statistics for the present reporting period are given in Table II.2.1, where it can be seen that an average of 10.8 events are reported in the NORSAR bulletin per day during the period.

Based on the one year of data now available after the automatic NORSAR event processor was implemented, an analysis of the detection and location capability of the array has been undertaken. Preliminary results from this analysis are given in Section VI.1 and VI.2.

H. Bungum

P. Engebretsen

	Teleseismic	Core Phases	Sum	Daily
Apr 78	275	41	316	10.5
May 78	286	51	337	10.9
Jun 78	336	51	337	12.9
Jul 78	305	56	361	11.6
Aug 78	241	60	301	9.7
Sep 78	201	75	276	9.2
Apr-Sep 78	1644	334	1978	10.8

Table II.2.1

# 11.3 NORSAR Data Processing Center (NDPC) Operation

Data Center

It has now become clear that one shift of computer operation is not sufficient to handle the combined load of operations and research activities at NORSAR. It is often necessary to run jobs that take an hour or more in the evening. A previous employee has therefore been hired on an hourly basis to run the Data Retention program at night and on the weekends, as this work has failen far behind schedule. Also during the summer extra help was hired to fill in during vacations.

The DP uptime for the period is 93.7%, which is 2.1% better than last period. There have been two major breakdowns on the SPS, and those breakdowns account for more than half the downtime. The number of stops has not gone down though.

After the reduction in IBM's maintenance contract, the servicing of 10 out of the 15 tape units is now being taken care of by NORSAR personnel. NORSAR personnel have also been engaged in maintenance and problem solving for the SPS, EOC and 360B computer.

#### J. Torstveit

#### Array Communications Circuits

Outages when groups of circuits have been affected simultaneously have been few this period. Twenty-one outages of this kind were observed, of which three lasted more than 1 hour. They are as follows:

April	1 outage	Lasted 1 hour
May	6 outages	Of which 1 lasted approx. 2 hours 25.5)
June	4 outages	Of which 1 lasted approx. 1.5 hours 7.6)
July	2 outages	Short
August	5 outages	Short
September	3 outages	Short.

On the other hand, we have experienced single subarray outages of rather long duration, and quite a few have been affected. Reasons have been: cable damages, equalizer trouble, level outside tolerances and reasons not stated. Subarrays particularly affected are:

```
01A Week 14
                 8.0%
     Week 15
                12.5%
 **
     Week 25
                 0.8%
01B Week 15
                13.1%
 11
     Week 25
                 0.8%
02B
     Week 38
                15.5%
02C
     Week 28
                45.2%
 11
     Week 30
                 8.0%
 11
     Week 31
                11.9%
 **
     Week 34
                 1.4%
     Week 18
03C
                 8.0%
 **
     Week 23
                35.7%
 **
     Week 25
                 2.4%
04C
     Week 18
                 6.5%
 **
     Week 23
                21.4%
 11
     Week 31
                12.8%
06C
     Week 25
                 0.6%
 **
     Week 27
                21.4%
     Week 28
                 7.1%
     Week 33
                27.4%
```

All modems, either located in the CTV's or at the Data Center, have been most reliable. However, in Week 35 a separation filter (AHS-card) had to be replaced in the CTV modem at OlB. Prior to the replacement the output level had been too low.

Table II.3.1 shows outages/degraded performance related to communication circuits.

	N N		1			•••••						
	½ YEAR	007	5.1	3.1	3.4	16.5	10.6	7.6	10.3	,	3.5	
	AVER. > 20	ì	2.2	1.0	6.0	1.8	4.5	1.6	0.7	0	0.1	
	SEP (4) (430.9)	>200	1	ī	14.9	ı	3.9	ı	1	2.3	+	02B
		^	0.7 0.6	9.0	9.0	0.8	6.0	1.1	8.0	0.8		
	AUG (5)	>20 >200	0.7	1.2	0.7	27.3	0.5	1.8	29.6	8.8 0.8	0,0/0	1 0
						2.0	6.0	0.5 4.6	1.4	1.7	5	70
		7200		5		9.70	0.2	0.5	78.7	15.9 0.9 12.6 1.7	020/060	0.4
1	^			÷ v			1.7	1.1	1	0.9		
	JUN (4) (5.6-2.7) 20 >200		1.7	7 0 0	ί α	r	27.10	0.70		15.9	03C 03C/04C	$\frac{1.2}{3.3}$
-	^	2 0	6.0	0.4	3.0	7 61 7.7	9	0.2		2.8	030	1.2
	(1.5-4.6) $(20 > 200)$		0.5	0.7	2.6	7.7	8.4	0.5	,	7:7		
	^	0.7		1.5	2.9	4.5	1.5	1.1	000			
APR (7.)	(3-30.4) 20 >200	26.3	15.1	0.7	3.2	0.7	0.7	0.7	6.8		01A 01A/01B	7.1
APR	>20	8.4	2.3	0.7	9.0	0.2	7.0	0.2	1.8		01A 01 0.8	
Sub-	array	01A	018	02B	02C	030	04C	290	AVER.		LESS	
	!											

TABLE II.3.1

The ARPA Subnetwork (i.e., TIP to TIP incl. modems, lines and interfaces)

The London Communications Circuit

This circuit has had a high degree of reliability most of the time. In the beginning of August, however, the 'Marginal Circuit' indicator was frequently initiated, but as 'Good Data' was on simultaneously, the error rate was within specifications. On 3 August, the carrier was lost and remained so until 7 August when the 'Co Slow' action among the British BPO employees was terminated.

The SDAC Communications Circuit

In August and September this circuit was subjected to repeated measurements and tests. According to the Network Control Center (NCC) short breaks in the data stream have been observed frequently. Several attempts have been made to isolate the fault, first of all by means of the looping facilities in the modems on both sides (i.e., DC Bus Back, Audio Bus Back and Modem Check). Also agencies such as ITT (in the USA), NTA (Norwegian Telegraph Administration) and the Codex modem representative (in Oslo) have been engaged. In addition, a representative from BBN (Bolt Beranek and Newman) has been involved. Measurements carried out on the line between NORSAR and SDAC 17 September proved attentuation distortion caused by irregularities in the US.

The Terminal Interface Massage Processor (TIP)

Preventive maintenance (PM) has been carried out according to the schedule. The teletype (TTY) is also regularly checked and maintained. In the period thunderstorms and power outages have caused a few problems in connection with the system restart. On 13 and 19 July lightning resulted in damaged cards in the Host Interface no. 3 and the Distant Host Driver. On 27 July, after a drop in the main supply, the machine was impossible to restart, and was down approx. 11 hours. A cross-patch setup in London when the NORSAR TIP was down delayed the system restart. Also on 28 July the TIP caused problems. Several tests/restarts under

restart due to main supply outage. When NCC tried to reload, it stopped in the same location each time. On 15 August a BBN representative arrived. The system resumed operation after some outages. On 17 September a BBN representative arrived in connection with problems in the subnetwork. At the same time he modified the IMP ID card, word 4 (this card defines the layout of an IMP, and word 4 now defines modem line speed, instead of satellite interfaces).

#### TIP Connections

No changes have been made to the IMP portion of the TIP, or the TIP port (LIU) connections since the last reporting period.

O.A. Hansen

#### III. IMPROVEMENTS AND MODIFICATIONS

#### III.1 Detection Processor

There have been few, but important, modifications to this system within the reporting period, as described below:

- As indicated in Table II.1.1, there were several stops caused by discovery of wrong timing (i.e., discrepancy between SPS internal time and TOD time) early in this reporting period. This was caused by SPS malfunctioning. However, on such occasions, the only thing done by the system was to write out an error message. Additional logic nas now been inserted into the DP system, to make it go down gracefully whenever such a discrepancy between the SPS internal time and the TOD time is detected. In this way registration of data corrupted with erroneous time is prevented. This modification was implemented 28 April.
- The parameters in the system holding the value of the subarray beam number (TALESA) and the single sensor channel number (TALESS) to be transmitted to SDAC for an Online Event was changed 18 May. The new channels transmitted with an Online Event are the subarray beam from 03C (5) and the single sensor 02B00 (13). This change was effected by a CORE card in the initialization deck. Future changes should thus be easy to perform.
- The Online Event number (EPX) system was changed 26 May. The philosophy is that Online stops should not cause gaps or jumps in the sequence of EPX numbers. After modification, the system now updates the EPX counter and uses the updated value every time an event is declared. The updated value is also queued, to be written to the disk. During initialization, the value on disk is read and inserted into the EPX counter as a starting value, thus maintaining continuity in the EPX numbers. The only situation that will cause problems is when the system goes down between updating the EPX counter and the actual writing of the updated value to disk. This will lead to duplicate EPXes (before and after the break). However, the first of the events with identical EPX will then probably be incompletely transmitted and registered on tape.

- On 27 July a programming error was found and corrected in the ARPANET message generating module (PNRSAD). Although this error did not influence the effective transmission of ARPANET data, since it occurred only when the local IMP node went 'dead', it prevented the system from staying up on such occasions, when an event was declared.

#### III.2 Event Processor

The new AUTOEP processing system, which reads Online Event Processor (OEP) results off the Detection Log tape, and performs further processing on these data, has been further improved and modified during the reporting period. Some of the changes have been done in order to adapt and interface it to the Interactive NORSAR Bulletin Editing System (INBES). The AUTOEP-INBES systems now handle all automatic and manual refinement/modification of the OEP solutions (see Fig. III.2.1).

The AUTOEP system reads the OEP results from the Detection Log tape, does solution refinement and computes event parameters. It produces Detection/Event Lists, Event Plots and a first version of an Event Bulletin. A reformatted version of this bulletin, adapted for telex transmission, is also produced. The Event Bulletin is finally written to the Disk Bulletin File. The INBES system allows the analyst to interactively review and modify entries in the bulletin, using the 2260 Display Station as his tool. If an event parameter is modified by the analyst, then automatic re-computation of related parameters will be performed. The final result is an Edited Bulletin, which has been reviewed by the analyst. This bulletin may be extracted (printed & punched) at regular intervals (say, each week).

The following changes have been made to the AUTOEP system:

- A subroutine was added (EDIT) to re-format the bulletin lines, so they conform to the Disk Bulletin File format. (8 May).
- Event statistics showed that the AUTOEP automatic onset selection on the average was 0.6 seconds too late. As an ad-hoc solution, the code was therefore changed to subtract 6 from the sample number returned from the onset selection subroutine. (12 September).

- Code was inserted to check for bad input data (i.e., flat trace),
   so that division by zero and plotter problems were prevented.
   (21 September).
- Clipping of traces with amplitudes too large for the allocated interval on the plot was implemented, as an extra security measure against cases of 'gallopping plotter pen'. (21 September).
- The area allocated for each trace was enlarged, so as to increase the amplitude resolution. This was wanted by the analysts. (21 September).

D. Rieber-Mohn

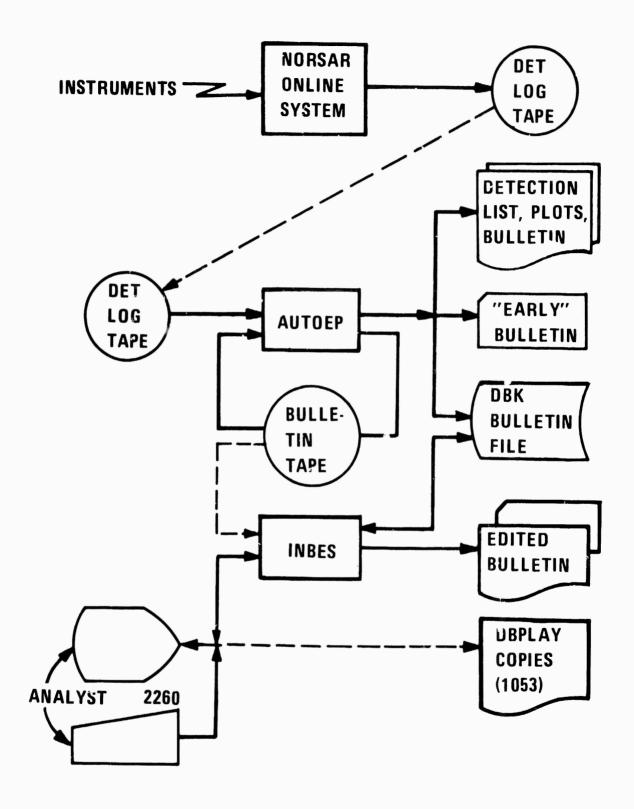


Fig. III.2.1 Data flow in the AUTOEP-INDES system

#### III.3 Array Instrumentation and Facilities

The test out of a Teledyne Geotech S-500 seismometer initiated 15 March 1978 on 06C channel 02 and described in the last Semiannual Technical Summary continued from 30 March to 12 April placed in NS horizontal position. A technical report (Larsen, 1978) has been written describing the experiment, whereof the frequency response measurement is copied in Fig. III.3.1.

As of 5 April 1978 a three-axis seismometer system was installed on channels 01 (vertical), 02 (NS horizontal) and 03 (EW horizontal) at subarray 01A in the long period vault, position latitude 60°50'39.2" longitude 10°53'11.5", elevation 426 meters. The seismometers are Teledyne Geotech S-13 short period seismometers, connected to NORSAR standard channel equipment. Frequency response curve for the vertical channel is given in Fig. III.3.2 and in Table III.3.1 the corresponding numbers are given, including the numbers for NS and EW channels. The calculation of equivalent e. ch motion and channel resolution is as follows:

Input calibration voltage  $E_i$  = 20 VPP Calibration network resistance  $R_n$  = 50 K $\Omega$  Calibration coil resistance  $R_c$  =  $\frac{23 \ \Omega}{E_i}$  Calibration coil current  $I_c$  =  $\frac{R_c + R_c}{R_c + R_c} \approx \frac{400 \ \mu A}{R_c}$ 

Calibration coil motor constant  $G_c = 0.1975 \text{ N/A}$  (for vertical seismometer 0.1973).

Equivalent earth motion at 1.0 Hz in microns:

$$y = \frac{G_c \cdot i \cdot 10^6}{(\pi^2 \cdot f^2 \cdot m)}$$

Where

y = equivalent earth motion in microns, peak-to-peak

G = calibration coil motor constant, newtons/ampere

i = current through the calibration coil, amperes peak-to-peak

f = frequency of calibration signal

m = weight of mass i kilograms.

$$y = \frac{0.1975 \cdot 400 \cdot 10^{-6} \cdot 10^{6}}{4 \cdot 9.8696 \cdot 1 \cdot 5} \approx \frac{.400 \text{ } \mu\text{M P-P}}{}$$

Channel resolution at 1.0 Hz:

$$\frac{\text{Equivalent earth motion}}{\text{Quantum units (QU) P-P at channel output}} = \frac{.400}{9360} = \frac{.42.73 \text{ PM/QU.}}{2.73 \text{ PM/QU.}}$$

A.K. Nilsen

#### Reference

Larsen, P.W. (1978): Test of Teledyne-Geotech S-500 Seismometer, NORSAR Internal Report 1-78/79, in press.

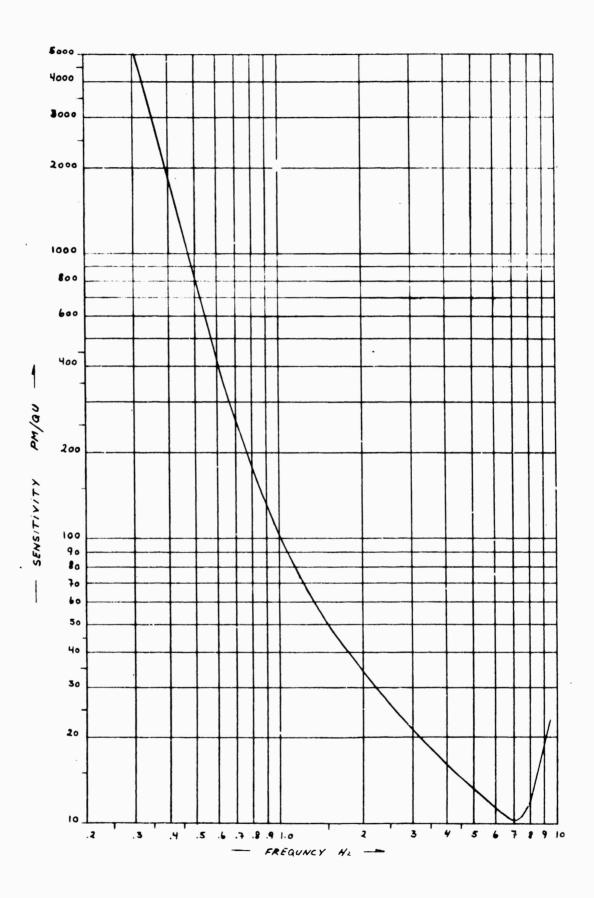


Fig. III.3.1 S-500 frequency response.

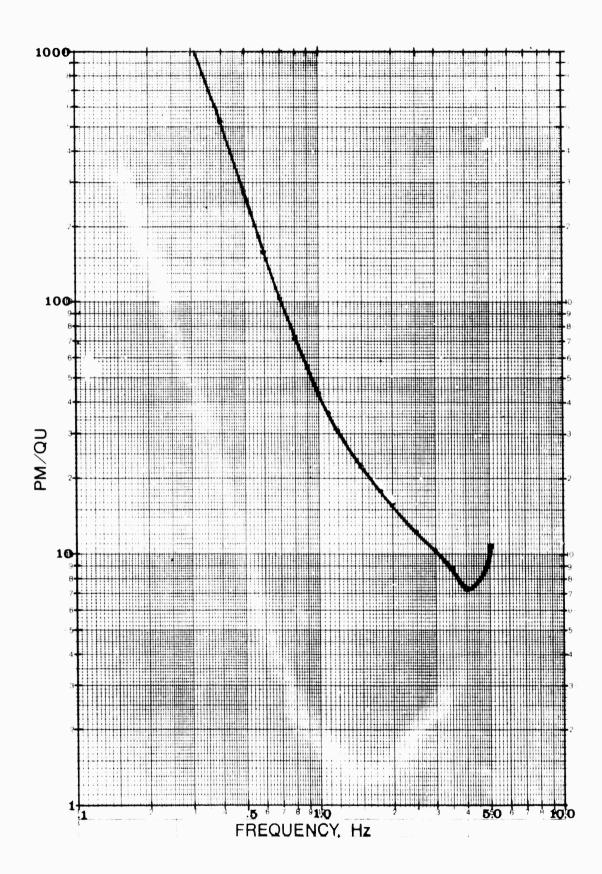


Fig. III.3.2 S-13 V frequency response.

Frequency	Output of ADC, PM/QU					
Hz	01A01 (V)	01A 02 (NS)	01A 03 (EW)			
0.1	48777	48826	40689			
0.2	5081	4694	4069			
0.3	1355	1292	1356			
0.4	525.6	544.9	544.9			
0.5	278.0	271.3	271.3			
0.6	157.55	157.71	157.71			
0.7	103.69	99.65	101.68			
0.8	71.90	69.36	71.97			
0.9	54.74	53.82	53.82			
1.0	42.73	42.73	42.73			
1.1	35.99	36.03	35.40			
1.2	30.79	30.82	30.82			
1.5	22.12	21.70	22.14			
2.0	15.63	15.65	15.65			
2.5	12.19	12.21	12.21			
3.0	10.42	10.43	10.05			
3.5	8.66	8.66	8.66			
4.0	7.26	7.27	7.63			
4.75	8.32	8.32	9.02			
5.0	10.84	10.85	10.85			

TABLE III.3.1 Frequency response of three-axis seismometer at 01  $\!\Lambda$  as of 26 July 1978

# IV. MAINTENANCE ACTIVITY

A brief review of the maintenance activity at the subarrays by the field technicians as a result of the remote array monitoring and routine inspection is given. The main preventive work in the period is dryout and painting of the floor and long period tanks in some of the long period vaults and replacement of seismometer amplifiers due to decaying battery power.

# Maintenance Visits

Fig. IV.1 shows the number of visits to the subarrays in the period, in average each subarray has been visited 9.3 times. The large number of visits to 06C are due to cable breakages. The relatively high number of visits is mainly explained by the preventive work mentioned above.

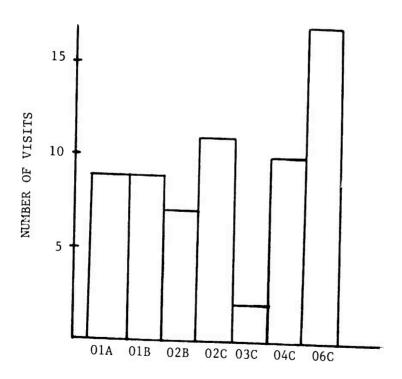


Fig. IV.1 Number of visits to the NORSAR subarrays in the period 1 April to 30 September 1978.

#### - 29 -

### Preventive Maintenance Projects

The preventive maintenance work in the array is described in Table IV.1. The adjustments are corrections of characteristics within the tolerance limits.

Unit	Action		No. of Actions
LTA	Adjustment of DC offset	SP	15
	l	LP	22
	Adjustment of channel gain	SP	5
		LP	44
	Adjustment of CMR	LP	3
Seism. Amplifier	RA-5 replaced due to decayir battery power	ng	15
Seis-	MP adjustment (in field)		9
mometer	FP adjustment (in field)	7	
SLEM	RSA/ADC adjustment	1	
Facilities	Dryout and painting of LPV a including replacement of RCI	; 3 	
CTV/LPV Access Roads	ess		

TABLE IV.1

Preventive Maintenance Work in the Period

1 April to 30 September 1978

# Disclosed Malfunctions on Instrumentation and Electronics

Table IV.2 gives the number of accomplished adjustments and replacements of field equipment in the array with the exception of those mentioned in Table IV.1.

Unit	Characteristic	SP Repl. Adj.	LP Repl. Adj.
Seis- mometer	Damping MP (field) FP (field)	3	10 7
Seism. Ampl. RA-5/ Ithaco	Balance Gain	1	1
LTA	DCO Ch. Gain CMR	4 4 1	1 2

#### TABLE IV.2

Total number of required adjustments and replacements in the NORSAR data channels and SLEM electronics

(1 April - 30 September 1978)

# Malfunction of Rectifiers, Power Loss, Cable Breakages

There has been no malfunction of the rectifiers in the period, or power loss requiring action of the field technicians. The number of cable breakages was 15, requiring 15 days' work.

# Array Status

The status of the array characteristics is similar to previous periods with little change. The SP array average DC offset was -3.4 millivolts averaging four minutes of quiet background noise. The LP DC offset was -0.5 millivolts over a ten-minute period.

A. Kr. Nilsen

#### ABBREVIATIONS

A	ADC	-	Analog to digital converter
(	CIIR	-	Common mode rejection
(	CTV	-	Central Terminal Vault
1	DC	-	Direct current
]	DCO	-	DC offset
]	FP	-	Free period
1	LP	-	Long period
1	LPV	-	Long period vault
į	LTA	-	Line termination amplifier
]	MP	-	Mass position
	PM	-	Pico meter
(	QU	-	Quantum units
	RCD	-	Remote centering device
	RSA	-	Range switching amplifier
	SLEM	-	Short and long period electronics modules
	SP	_	Short period.

# V. DOCUMENTATION DEVELOPED

# V.1 Reports, Papers

- Bungum, H. (1978): Re-analyzation of three focal-mechanism solutions for earthquakes from Jan Mayen, Iceland and Svalbard, Tectonophysics, in press.
- Gjøystdal, H. (1978): Semiannual Technical Summary, 1 October 77 30 April 1978, NORSAR Scientific Report 2-77/78.
- Haddon, R.A.W. (1978): Scattering of seismic body waves by small random inhomogeneities in the earth, NORSAR Scientific Report No.3-77/78.
- Larsen, P.W. (1978): Test of Teledyne-Geotech S-500 seismometer, NORSAR Internal Report No. 1-78/79.
- Rieber-Mohn, D. (1978): The Interactive NORSAR Bulletin Editing System User Guide and Documentation, NORSAR Internal Report No. 3-77/78.

  L.B. Tronrud

# V.2 Program Documentation

Documentation N/PD-93 has been completed within this period. N/PD-93 is a subroutine that reads tapes produced by the DHR-1632 recording system.

D. Rieber-Mohn

# VI. SUMMARY OF SPECIAL TECHNICAL REPORTS/PAPERS PREPARED

# VI.1 Evaluation of the Current NORSAR Detection Capabilities

One year of analyzed data (Oct 77-Sep 78) is now available after the NORSAR array was reduced in size. This data base is considered sufficient to conduct a preliminary analysis of the capabilities of the new NORSAR configuration. In the present study, the current event detection capabilities have been estimated both by comparison with the old 22 subarray system and by recurrence analysis of the magnitude-frequency relationship of the reported earthquakes.

Table VI.1.1 and Fig. VI.1.1 show the monthly average number of reported events during the last four years of 22 subarrays operation as compared to the most rount year. All of these numbers are based upon events reported in the analyst-reviewed NORSAR seismic bulletin, and thus represent real seismic events with minimal occurrence of false alarms. Apart from one month, March 1978, during which a large earthquake sequence occurred, the picture is quite stable, and the average monthly number of reported events is now about 60% of what it was before the reduction from 22 to 7 subarrays. The ratio is about 65% if one compensates for the increased system downtime after the reconfiguration. From Table VI.1.1 it is further seen that the estimated degradation does not change significantly if one deletes all months that contain large earthquake sequences.

Assuming that the b-value of the magnitude-frequency recurrence relationship is independent of time, it is easily seen that the change in 507 incremental detectability threshold  $\Delta m_b$  corresponding to the ratio R of detected events (in per cent) can be expressed as (Pirhonen et al, 1976)

$$\Delta m_b = -\frac{1}{b} \cdot \log_{10} (\frac{R}{100})$$

Assuming b=0.9, R=65 then gives  $\Delta m_b^{}=0.21$ . This may be compared to the theoretical reduction in beamforming gain  $\Delta G$  for 7 versus 22 subarrays, which is  $10 \cdot \log_{10} 22/7 = 5.0$  dB or  $0.25 \, m_b^{}$  units. Thus the observed performance relative to the old configuration corresponds closely to what could be expected. It appears that the increased automation in the bulletin generation procedure has not significantly affected the

detection performance, as far as the number of reported events is concerned.

We now turn to the problem of obtaining an independent measure of the current NORSAR detectability (i.e., a measure not relative to previous capabilities). Our progress so far has been rather limited, because the most reliable estimation method, namely, that of checking detections against an independent reference station (Ringdal, 1975) has not been possible to use after the only reference system of sufficiently high capability, the SDAC/LASA bulletins, has been discontinued. Therefore, we have resorted to the recurrence technique, analyzing the available one year of data (Oct 77 - Sep 78) in a way identical with what was done by Berteussen et al (1976) in their final evaluation of the NORSAR detectability before the reduction in array size. Only the results from the least squares cumulative method will be presented here, as shown in Table VI.1.2. The method is illustrated in Fig. VI.1.2, which shows the combined teleseismic data (Region 14). The results in Table VI.1.2 must be considered relatively uncertain for most regions, due to the limited data base. Nonetheless, we may take note of the 90% cumulative thresholds for regions such as Central Asia (3.6) and Japan-Kamchatka (4.0). It is also evident that the performance has decreased somewhat relative to that of the NORSAR system during 1972-75 (Berteussen et al, 1976), with a degradation varying in the range 0-0.3 m, units. The uncertainties inherent in the estimation method should, however, not be forgotten, and in general we consider the number of reported events to be a more reliable indicator of the array performance than the results from the recurrence analysis.

Our future plans include developing new methods for a more reliable direct estimation of the detectability of the NORSAR array, in particular at regional and near-regional distances. As more data are accumulated, it should also be possible to obtain better estimates of event detectability in selected seismic regions.

- H. Bungum
- F. Ringdal

# References

- Berteussen, K.-A., H. Bungum and F. Ringdal (1976): Re-evaluation of the NORSAR detection and location capabilities. Scientific Report No. 3-75/76, 20 June 1976.
- Pirhonen, S.E., F. Ringdal and K.-A. Berteussen (1976): Event detectability of seismograph stations in Fennoscandia. Phys. Earth Planet. Inter., 12, 329-342.
- Ringdal, F. (1975): On the estimation of seismic detection thresholds.

  Bull. Seism. Soc. Amer., 65, 1631-1642.

(a)	) 4 yrs	(b) 1 yr	(c) Ratio (R)	(d) log R	(e) Swarms Kemoved
	72/76	77/78	(%)		(%)
Oct	د 483	380	79	-0.10	79
Nov	430	214	50	-0.30	49
Dec	505	235	47	-0.33	54
Jan	658	210	32	-0.49	48
Feb	499	263	53	-0.28	53
Mar	513	855	167	0.22	-
Apr	555	316	.57	-0.24	60
May	550	337	61	··0.21	65
Jun	769	387	50	-0.30	7 <i>5</i>
Ju1	659	361	50	-0.26	59
Aug	692	301	43	-0.37	57
Sep	442	276	62	-0.21	62
Average	563	345	61	-0.24	60
Average Compen- sated f DP Down Time	576	373	65	-0.22	64
Average	DP upti	me 1972/75:	97.7%		
Average	DP upti	me 1977/78:	92.7%		

# TABLE VI.1.1

Monthly averages of the number of NORSAR-reported events (a) for the four years Oct 72 - Sep 76, (b) for the year Oct 77 - Sep 78, (c) the ratio R (%) between the numbers, (d)  $\log_{10}$  R, (e) the ratio R modified by deleting months during which significant earthquake swarms occurred.

Region	Area of Coverage	Events 1977/78	90% Cumulative	Down from 1972/75
1	Aleutians-Alaska	214	4.0	0.3
2	Western North America	36	-	-
3	Central America	62	4.4	0.1
4	Mid-Atlantic Ridge	52	3 . 9	0.1
5	Mediterranean-Middle East	259	3.7	0.1
6	Tran-Western Russia	147	3.7	0
7	Central Asia	276	3.6	0.1
8	Southern-Eastern Asia	167	3.9	0.3
9	Ryukuo-Philippines	325	4.5	0
10	Japan-Kamchatka	1255	4.0	0.2
11	New Guinea-Hebrides	105	4.6	0.1
12	Fijl-Kermadec	6°5	4.1	0.2
13	South America	31	-	-
14	Distance range $30^{\circ}$ - $90^{\circ}$	2815	3.9	0
15	Distance range $110^{\circ}$ - $180^{\circ}$	802	4.7	0.1

TABLE VI.1.2

Detectability statistics for the reconfigured NORSAR array for 15 geographic regions (see Berteussen et al, 1976). Within each region the table gives the number of reported events, the estimated cumulative 90% detection threshold in terms of NORSAR m<sub>b</sub> (from recurrence analysis) and the corresponding degradation relative to the 1972/75 performance (as estimated by Berteussen et al, 1976).

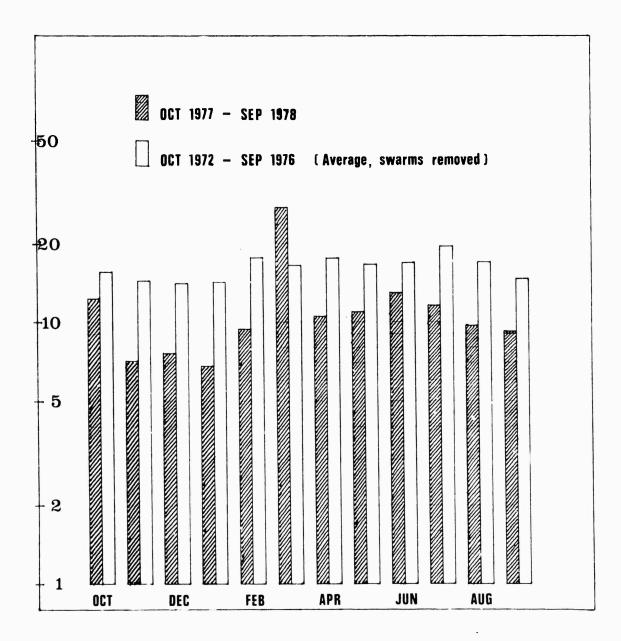


Fig. V.1.1 Monthly averages of reported events at NORSAR, corresponding to the last column of Table VI.1.1. Note that in computing the averages for the period Oct 1972 - Sep 1976 all months with significant earthquake swarms have been ignored. According to the same criterion, the month of March 1978 should be ignored when comparing the two periods.

### CUMULATIVE



# INCREMENTAL

A=  $4.72 \pm 0.14$ B=  $-0.56 \pm 0.03$  50% = 3.8390% = 4.12

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REGION 14

Fig. VI.1.2 Frequency-magnitude distribution for events reported in the NORSAR seismic bulletin for the one-year period October 1977-September 1978. The figure covers events in the distance range 30°-90° from NORSAR. Estimated cumulative and incremental detection thresholds are indicated.

# VI.2. Evaluation of the Current NORSAR Location Capabilities

The effect of the reduction in the array size has been investigated by comparing the NORSAR locations with the PDE (Preliminary Determination of Epicenters) from the USGS (United States Geological Survey). Because of the delay in the PDE bulletin, only three months of data have been available for comparison from October 1977 to December 1977.

Fig. VI.2.1 shows the NORSAR/USGS location difference for 286 events commonly reported within teleseismic distance from NORSAR (30°-90°). The median location difference is 230 km, which should be compared to the 130 km reported for the same region using the old and larger array (Berteussen et al, 1976). Results on a regionalized basis are shown in Table VI.2.1, where we can see that the increase in the median location difference ranges between 20% and 100% for the different regions. Estimates were not obtained for 6 of the regions because of the limited amounts of data available. We see from the table that while Japan-Kamchatka previously was the best region, it is now Central Asia, with a medium location difference of 170 km.

In considering these results it is important to notice that at the same time as the array was reduced in size (about 50% in diameter), the amount of processing for each event was also reduced, essentially by removing the previous epicentral refinement procedure and using only the beam location from the on-line detection. The effect of the array size reduction itself is therefore smaller than what is reflected in the numbers given in Table VI.2.1.

#### H. Bungum

#### Reference

Berteussen, K.-A., H. Bungum and F. Ringdal (1976): Re-evaluati of NORSAR detection and location capabilities, NORSAR Scientified Report 3-75/76.

Regions	Area of Coverage	Jan 73 - Events	- Mar 75 Median	Oct 77 - Events	- Dec 77 Median	Increase (%)
1	Aleutians-Alaska	195	110	38	180	79
2	Western North America	129	130	9	ŧ	ı
3	Central America	146	200	9	i	ı
<b>\</b> †	Mid-Atlantic Ridge	143	150	1	i	ı
5	Mediterranean-Middle East	389	300	29	520	73
9	Iran-Western Russia	182	170	11	ŧ	ı
7	Central Asia	349	120	32	170	42
80	Southern-Eastern Asia	205	150	21	180	20
6	Ryukuo-Philippines	454	200	39	250	25
10	Japan-Kamchatka	1062	100	82	200	100
11	New Guinea-Hebrides	263	210	10	ī	1
12	Fiji-Kermadec	208	230	74	400	74
13	South America	112	210	6	ı	I
14	Distance Range 30 <sup>0</sup> -90 <sup>0</sup>	3775	130	286	230	1,
15	Distance Range 110 <sup>o</sup> -180 <sup>o</sup>	1195	220	100	380	73

TABLE VI.2.1

Median location difference in km between USGS and MORSAR for the time period Jan 73-Mar 75 (Berteussen et al, 1976) and for Oct 77-Dec 77 (present study), and the increase in percentage.

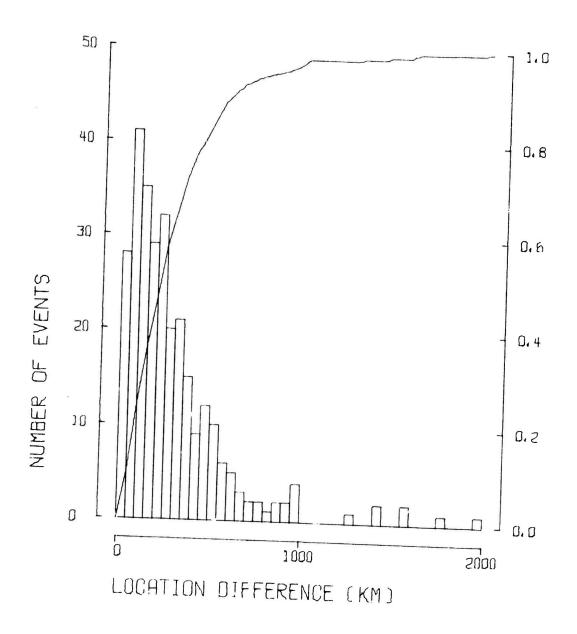


Fig. VI.2.1 Cumulative and incremental distribution of epicenter location differences between USGS and NORSAR for Region 14 (see Table VI.2.1) for the time period Oct 77 - Dec 77.

# VI.3 Magnitude Studies

Seismic event magnitude represents one of the most important parameters in the context of seismic discrimination due to the versatility of the m<sub>b</sub>:M<sub>c</sub> discriminant. A novel approach to the estimation of magnitude was introduced by Ringdal (1976), who pointed out the advantages of using truncated distribution theory in estimating network magnitudes of small events. This topic has been further elaborated by Christoffersson (1973), who developed a unified model for estimating magnitudes and detection thresholds. This approach has now been extended to estimate simultaneously  $\mathrm{M_e^{-m}_h}$  relation of earthquakes, the scattering in these observations together with detection thresholds for the arrays and individual seismograph stations used to form the data base. In the present study, we have adapted the maximum likelihood technique to assess the linearity or lack of such of the mh:Hg relationship - a problem critical for seismic source identification. Only preliminary results based on rather limited observational data have been obtained so far, and examples of the observed (m, M<sub>c</sub>) relations are shown in Figs. VI.3.1 and VI.3.2. These results are based on  $M_{\rm S}$ -values as reported by Uppsala, although we have also experimented with corresponding NOAA and NORSAR observations. In the latter cases, the results are similar to those displayed in Fig. VI.3.1 and VI.3.2. It should be noted here that Uppsala appears to be the only seismological station which consistently reports the  $\mathrm{M}_{\mathrm{c}}$ -parameter and also has done so over a very extensive period of time. Of course, other seismological agencies like ISC (International Seismological Centre), NOAA (National Oceanic and Atmospheric Administration, USA), Yoscow World Data Center and also the Berkeley (BKS) seismographic station often report  $M_S$ -magnitudes, but their observations constitute the average for a set of stations, while for BKS the reported  $M_{_{\mathbf{S}}}$  is the average of the truly observed Mg and the linearly transformed mg-to-1g values.

An illustration of Christoffersson's method applied to  $\mathbf{m}_{b}$  data from two stations is shown in Fig. VI.3.3, and it is seen that the apparent deviation from the expected slope of 1.00 can be satisfactorily explained by detectability considerations.

Our studies so far have verified the commonly observed appearance of  $M_s:m_b$  scatter plots: at high magnitudes, the  $M_s:m_b$  slope is significantly greater than 1.00 (typically around 2), while at lower magnitudes (below  $m_b \sim 6.0$ ) there is apparently a distinct curvature in the relationship between  $M_s$  and  $m_b$ . However, our results show that this behavior may be explained as a result of bias effects in the plots at low magnitude caused by detectability problems. Thus the hypothesis of an intrinsically linear  $m_b - M_s$  relationship with a slope greater than 1.00 even at low magnitudes cannot be rejected on the basis of these and other similar observations. The work reported above will be continued, and future plans include greatly extending the data base so as to allow more specific conclusions about the slope of the  $M_s - m_b$  relationship.

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- F. Ringdal
- C. Bjørck, Dept. of Statistics, Uppsala
- E.S. Husebye

### References

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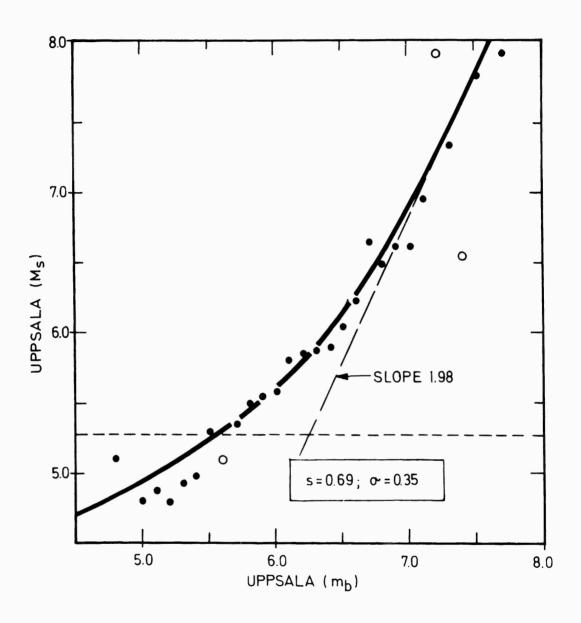


Fig. VI.3.1 Average M for given m values (both M and m reported by Uppsala for a reference data set of 412 earthquakes. The solid, curved line is a model fit based upon an assumed linear M m telation (slope as indicated) modified by an estimated M detectability curve. The parameters (µ,σ) of the detectability curve (Ringdal et al, 1977) are indicated (µ is shown by a stippled horizontal line). The standard deviation S of the inherent M m scatter is also estimated. The open circles are data points outride two standard deviations from the model.

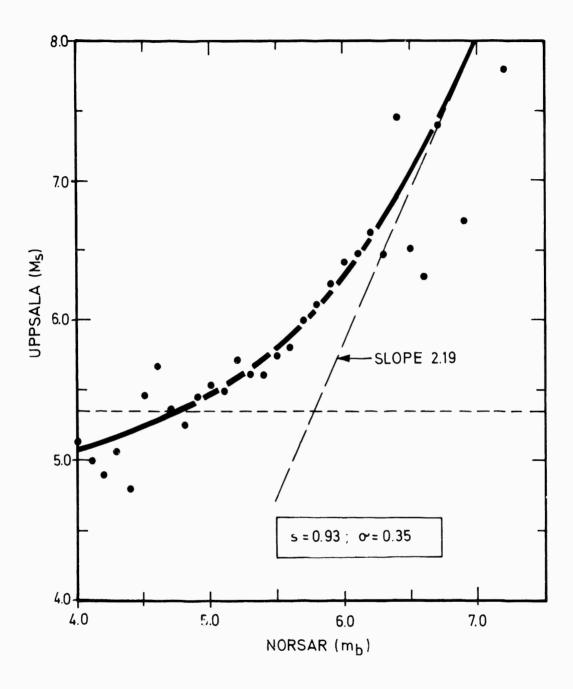


Fig. V1.3.2 Same as Fig. V1.3.1, but with m as reported by NORSAR, M by Uppsala and the data set containing 403 earthquakes.

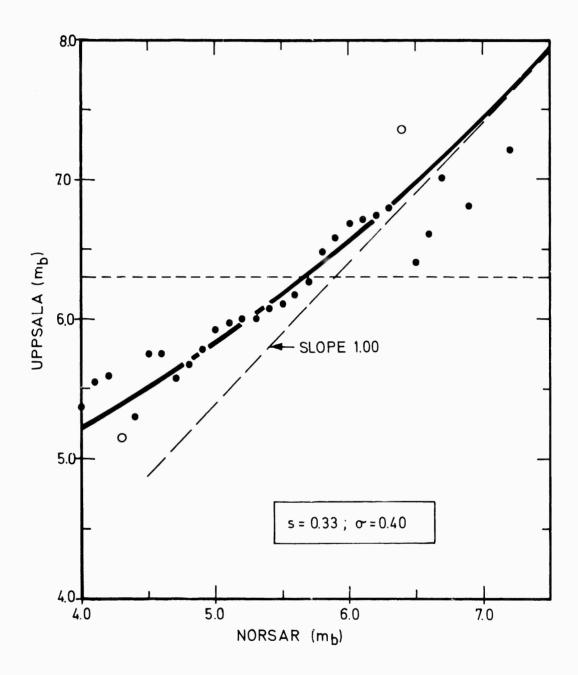


Fig. VI.3.3 Illustration of the model in Fig. VI.3.1 and VI.3.2 applied to estimate the relationship between m<sub>e</sub> (NORSAR) and m<sub>e</sub> (UPPSALA) (data base 386 earthquakes). The expected slope of 1.00 does not appear to fit the observed data. However, if one takes the effect of detection thresholds into account, while fixing the slope at 1.00, one arrives at the solid, curved line which is a much better fit.

# VI.4 Seismic Event Discrimination Based on Near-Field Observations

Up to now the NORSAR event discrimination efforts have to a large extent been based on information extended from teleseismic recordings corresponding to epicentral distances mostly in excess of 30 deg. Lately we have taken up research aimed at several aspects related to seismic event discrimination based on near-field recordings. Although only preliminary results are available at present, we consider it worthwhile to present in some detail the rationale behind these efforts and also the work done so far.

# Seismograph Station Detectability Studies in Near-Field Distance Cauges

Most event detectability studies previously undertaken have been based on the reporting performance of individual stations of P-wave recordings in the epicentral distance interval of 30-90 deg. The actual observational data necessary for such studies may be extracted from easily available files, and in this respect the ISC (International Seismological Centre, U.K.) bulletin files have been widely used, e.g., see Ringdal et al (1977). Now, in this approach a prerequisite is that a station reporting performance is somewhat lower than that of the reference station or reference reporting agency like ISC, a condition which with few exceptions is fulfilled for teleseismic distance ranges. For example, out of the 482 stations analyzed using the ISC focal parameters as a reference, only 4 stations had a P-wave reporting performance better than or equivalent to that of ISC, so that their detectability performance could not be estimated reliably. Problems of the latter type may become much more severe for near-field distance ranges, as in this case we have to actually consult individual station bulletins to check on possible mismatches in event detectability with that of ISC. As NORSAR's event detectability performance is superior to that of most other stations and networks for events occurring in large parts of Eurasia, we are considering options to replace ISC by NORSAR as a reference station in the detectability analysis. We are also investigating the possibility of obtaining local station bulletins so as to enable us to undertake specialized analysis of stations of particular interest in this respect.

Work accomplished so far here is as follows:

- All available ISC-bulletin files since 1971 have been transformed to a special compact format suitable for our particular kind of analysis.
- The corresponding data analysis routines have been adapted for detectability estimates for all stations consistently reporting to the ISC and for near-field distance intervals of 5, 10 and 15 deg to ensure that at least in some intervals sufficient data will be available.

We are also considering an extended detectability study based on the fact that stations in coastal and some other areas have an event detectability which is likely to be subjected to seasonal weather conditions. In other words, we are considering estimating station detectability as a function of time of year in intervals of 3 months. The outcome of this experiment will be used for checking whether there is a significant difference in event detectability of certain areas at high latitudes by using an appropriate network of high-quality seismograph stations.

Present status on the various detectability experiments is that the necessary software developments have been completed, and that the first analysis of real data has commenced.

J. Fyen

E.S. Husebye

F. Ringdal

#### Reference

Ringdal, F., E.S. Husebye and J. Fyen (1977): Earthquake detectability estimates for 478 globally distributed seismograph stations, Phys. Earth Planet. Inter., 15, P24-P32.

# VI.5 Near-Field Wave Propagation Problems

The event detectability study outlined in Section VI.4 is essentially an analysis of P-wave amplitude distribution as a function of epicentral distance, although the results also depend on local station factors like the geological setting at the site, noise conditions, operational quality and so on. From an event discrimination point of view, we are also interested in strong phases in the near-field recordings, not only the excitation level of the traditionally studied P- and Rayleigh-waves. For example, an unresolved question is whether or not so-called Lg-waves propagate unhampered across prominent tectonic features like the Urals and Himalayas, and also what the relative significance of this phase is in the near-field range. In order to answer these and related questions, we have started a relativ... j comprehensive analysis of near-field earthquake and explosion recordings from seismographic stations in Eurasia in addition to our own NORSAR recordings. In the following we will present the observational data presently under consideration, the method of analysis and finally some comments on preliminary results.

# WWSSN-station data base

Notwithstanding the many advantages of seismic tape recordings, on one account the analog WWSSN-records are superior, as they in a visual and compact form convey the essence of seismic recordings, namely, the relative energy distributions and the associated group velocities (or apparent group velocities). In this context we have collected, so far, about 320 ppies of Eurasian WWSSN recordings at the U.K. seismological data library in Edinburgh. It is already now clear that this data base has to be greatly extended in order to obtain reliable multidimensional earthquake/explosion discriminants as we have to know in detail the group velocity interval for the energetic parts of seismic recordings for the largest possible ensemble of source-station combinations (for further details, see Section VI.6).

The first step in analysis of these records is to measure arrival times and maximum amplitudes of all prominent phases - wave trains for epicentral distances less than exprox. 25°. In the second step of analysis we measure the group velocity interval associated with the most energetic wavetrains

in the records, and also check the mode of propagation, that is, fundamental and higher modes of Love and Rayleigh waves. For this particular task the WWSSN-analysis will be complemented by very detailed analysis or digital NORSAR and SRO-records.

# Preliminary results

Manual analysis of analog records is a rather time-consuming venture, so only preliminar results are available at present and are as follows:

- P-waves ( $P_g$ ,  $P_b$  or  $P_n$ ) are generally among the very strongest in the SP records. The most energetic phases here have in general velocities roughly linear with distance out to  $20^\circ$  in the bracket 7.0-7.5 km s<sup>-1</sup> generally associated with  $P_b$ .
- S-waves and/or higher mode Love-Rayleigh waves have velocities around 4.5 km s<sup>-1</sup> (Sn-waves), around 3.80 km s<sup>-1</sup> (Li-waves) and 3.35-3.54 km s<sup>-1</sup> (Lgl and Lg2 waves). Sn, Li, Lgl and Lg2 are now generally interpreted in terms of higher mode Love and Rayleigh waves, or as alternatives to the conventional Sn, Sb and Sg notations, for which the mode of propagation involves the uppermost part of the mantle. Specifically no low velocity layer below Moho is required for their explanation (e.g., see Knopoff et al, 1974; Mantovani et al, 1977; Panza and Calcagnile, 1975). We note in passing that initially Lg and Li waves conceptionally were associates with low velocity layers in the crust Lg for shear waves in the granite (g) layer and Li in the basalt or intermediate (i) layer (Bāth, 1962).
- Irrespective of source type, Sn (approx. 4.5 km s<sup>-1</sup>) and fundamental mode Rayleigh waves besides occasional P-phases dominate the LP-records. Li is seldom seen, and Lg almost never.
- In Sp records Lg-waves are prominent together with various P phases.

  In case of explosive sources Lg-waves sometimes completely dominate the records with amplitudes slightly larger on the horizontal components. The most efficient transmission paths for Central / sian events appear to be westward towards Fennoscandia, whereas propagation is less efficient towards India, Pakistan and Iran. For earthquake records usually exhibiting somewhat lower signal frequencies as compared to explosion sources Lg-waves from our preliminary observations appear to be less prominent. Beyond 12-15° the Lg-waves decrease rapidly.

For very short distances ( $\Delta$ <5°) the SP-records are relatively 'messy'. demonstrating the importance of scattering and mode conversion effects associated with crustal heterogeneities.

It is somewhat premature to speculate on the potential event discrimination power of the Lg-phase in a near-field context. What we know so far is that the observed Lg-excitations vary considerably with source type (preference to high-frequency radiation), and source-receiver paths. The relative attenuation efficiencies of tectonic barriers like the Urals and Himalayas as well as thick sedimentary basins are difficult to assess with our present data base. However, this can and will be done, given a sufficiently large data base. How to handle this kind of information in a discrimination context is the topic of the next section.

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## VI.6 General Purpose Program for Seismic Discrimination

One of the current projects at NORSAR is to investigate the discrimination power of the Lg-phase for near-field observations. In this connection a general program has been completed which is based upon a feature-extraction procedure combined with classification statistics. The idea is to extract as few parameters as possible from the records and still preserve the main information (information pertaining to the second order statistics of the time series like the autocorrelation function or equivalently the power spectrum. The program is general in the sense that all the available information about an event is loaded into the computer and then the information carrying parameters is extracted and subsequently used for classification (for references, see Sandvin and Tjøstheim, 1978).

The input data to the program should be single parameters like the  $\rm m_b^-$  parameter for body waves,  $\rm M_s^-$ parameter for surface waves and/or some  $\rm L_g^-$ parameters considered to have a substantial discrimination potential. Combined with these parameters different time windows with amplitude registration from the seismogram starting from the onset of the chosen phase, may be incorporated in the input data file.

With registrations of the actual phases/wave trains for the same event but from different stations, the various time series may be combined into one multidimensional time series. In case of different frequency response of the seismometers, the traces should be filtered to remove the instrument response from the registrations. The feature extraction procedure consists of two steps. The first step is accomplished by fitting a multivariate autoregressive model of order P to the combined phase registrations described above.

$$\underline{x}(t) - \Lambda_{l}\underline{x}(t-1) \dots - \Lambda_{p}\underline{x}(t-p) = \underline{z}(t)$$

Here  $\underline{x}(t)$  denotes the vector registration at time t,  $\Lambda_i$ , i=1,2,...p are n x n matrices where n is the number of individual records, and  $\underline{z}(t)$  an n-dimensional white noise vector.

The appropriate order p of the model should be determined from a criterion given by Akaike (1971). The fit of the model to the observations is evaluated by checking the whiteness of the residual process  $\underline{z}(t)$ . If the model is found appropriate, the second order statistics of the multidimensional time series as given by the individual spectra and the cospectra are completely specified from the matrices  $A_i$  and the variance matrix V of the white noise process.

The second step of the feature extraction procedure is to combine the parameters contained in the matrices  $A_i$  with the single input parameters  $(m_b, M_s,$  etc.) from the different stations into a vector  $\underline{Y}$  and then apply a principal component analysis to this vector. The idea in the principal component analysis of  $\underline{Y}$  is to pick vectors  $\underline{h}_i$ ,  $i=1,2,\ldots,m$  in such a way that the main part of the information as expressed by the variation of  $\underline{Y}$  is decomposed along a few of the vectors  $\underline{h}_i$ . The basis vector  $\underline{h}_i$  are given by the eigenvalue problem:

$$R h_i = \gamma_i h_i$$

where R is the covariance matrix  $E\{\underline{Y}\ \underline{Y}^T\}$ . The estimation of the covariance matrix for earthquakes,  $R_{EQ}$  and for explosions  $R_{EX}$  requires a data base of presumed earthquakes and of presumed explosions.

Now for each event the estimated vector  $\underline{Y}$  is decomposed along  $\underline{M}_{EQ}$  principal vectors  $\underline{h}_{EX}$  for the earthquake data base and along  $\underline{M}_{EX}$  principal vectors  $\underline{h}_{EX}$  for the explosion data base with components

$$Z_{EQ}(i) = \underline{Y}^{T} \cdot \underline{h}_{i,EQ}, \qquad i = 1,2,..., M_{EQ}$$

and

$$Z_{EX}(j) = \underline{Y}^{T} \cdot \underline{h}_{i,EX}, \qquad j = 1,2,..., M_{EX}$$

respectively.

The components are finally combined into one vector

$$\underline{z}^{T} = [z_{EQ}(1), ..., z_{EQ}(M_{EQ}), z_{EX}(1), ..., z_{EX}(M_{EQ})]$$

The vector  $\underline{Z}$  is then regarded as a stochastic variable with distribution function  $F_{EQ}$  or  $F_{EX}$  depending on whether the event is an earthquake or an explosion. It is assumed that  $F_{EQ}$  and  $F_{EX}$  are the multivariate Gaussian distribution with the mean value and covariance matrix for each population determined from the earthquake data base and the explosion data base respectively. The discrimination is then accomplished by a classification procedure where the event is assigned to the population having the highest probability of  $\underline{Z}$  occurring. As pointed out (Azen et al, 1975), the classification procedure is relatively robust to deviations from normality. Added flexibility to this particular discrimination is needed in order to handle missing observations and also changes in the number of reporting stations with changing source regions.

At present an extensive analysis of the Lg-phase from near-field (within 20°) observations obtained from WWSSN is going on. It is concluded that the Lg-phase observations as well as those of Sn are evident from these readings and the amplitudes should be included in the input data to the program described above.

Finally, if the Lg-phase may turn out to be a potential discrimination parameter, SRO-recordings should be provided in order to apply the feature extraction procedure to that section of the records where the Lg-phase is found. 'Identical Classification Procedures' should be applied to different two-dimensional discriminants, like the well-established m<sub>b</sub>:M<sub>s</sub> criterion, to have the opportunity to give a precise comparison of the different discriminants.

### O.A. Sandvin

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# VI.7 <u>Do Geological Surface Features Have a Counterpart in the Deeper Part</u> of the Lithosphere

Questions of the kind indicated in the heading of this section are frequently raised when efficiency of Lg-propagation along certain source-receiver paths are discussed. For example, mountain ranges like the Urals and Himalayas are intuitively associated with roots in the deeper lithosphere and thus act as barriers to Lg-propagation across such features. As regards the Himalayas there are, as also reported in a previous section, considerable observational evidence in support of the above hypothesis

In this context it may be appropriate to ask whether manifestations of less prominent tectonic activities like taphrogenesis have a counterpart in the lower lithosphere and thus affect the efficiency of Lg-propagation across such structures. As part of such an experiment we have tried to find out whether the Oslo graben has a seismic counterpart in the crust and upper mantle below the NORSAR array the problem was mainly unsettled in the Aki, Christoffersson and Husebye (ACH) (1977) travel time inversion experiment and this also applies to the amplitude modelling experiments by Haddon and Husebye (1978).

In order to answer the above question, we have modified the ACH-inversion techniques by imposing two types of restrictions on the parameter vector m:

- Some of the elements or blocks are zero, e.g., all m<sub>1</sub>'s within a particular layer are zero, which physically means that the layer represents homogeneous structures.
- Some of the elements are equalized, e.g., the blocks encompassed by the surface contours of the Oslo graben are made equal and thus constitute a large structural unit.

The above modified version of the ACH-inversion technique has been tested in the Oslo graben using the Haddon and Husebye time residual data base (more than 4000 observations). Relevant results are shown in Figs. VI.7.1 and VI.7.2 for which we have concluded that the surface graben contours have a seismic counterpart in the crust but probably not below Moho or at best it is very weakly represented here. This result is in good agreement with those derived from corresponding gravity observations as demonstrated by Husebye et al (1978).

The next step is, as mentioned above, to analyze proper observational data in order to check the propagation efficiency of high-frequency waves across a minor crustal tectonic feature.

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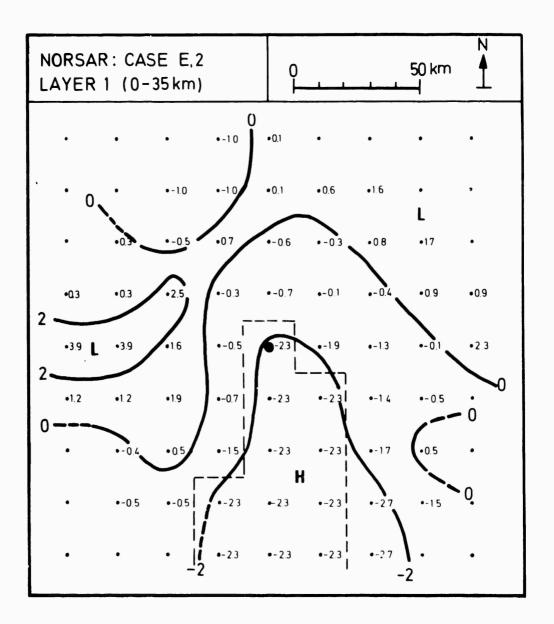


Fig. VI.7.1

a) Estimated seismic velocity anomalies in the crust of the NORSAR siting area. The Oslo graben contours are outlined and considered as one structural unit in the inversion experiment. A 3-layered standard earth model was used with Layer 2 removed or 'declared' homogeneous. Average layer velocity was 6.9 km s 1, block size or horizontal extent of blocks is 20 x 20 km 2, and non-hit blocks are marked by a dot. High and low velocity areas are marked by the letter H and L respectively. For computational details, we refer to Aki et al (1977) and Christoffersson and Husebye (in preparation). We take these results to indicate that the surface contours of the Oslo graben have a seismic counterpart in the crust.

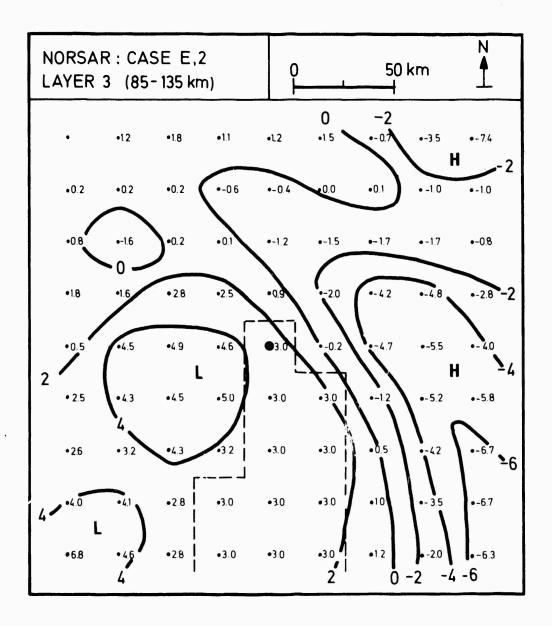


Fig. VI.7.2 These results plus estimated model fit parameters are taken to indicate that the graben imprints on the lower lithosphere are at best modest. Otherwise, caption as for Fig. VI.7.1.

# VI.8 <u>Description of and Preliminary Results from a Seismic Network for</u> Microearthquake Studies in Tanzania

The African continent has until now been fairly poorly covered with seismic stations, and this applies in particular to Eastern Africa. In our capacity as seismological consultants in connection with the planning of a 1200 MW hydroelectric power plant in the Rufiji Basin in Tanzania, NTNF/NORSAR has recently completed the installation of a modern network of 6 short period seismometers in the area. The installation, which is called the Stiegler's Gorge Seismic Network (SGSN), has an aperture of about 50 km (see Fig. VI.8.1), and is located around 8°S, 38°E (see Table VI.8.1). This is about 1000 km from any previously known seismic station. The individual stations of the SGSN are powered by solar panels and the data are transmitted by radio telemetry to a Central Recording Station near the future dam site (see Fig. VI.8.1), where the analog data are passing a voting detector (presently 2 out of 3) and subsequently recorded on digital magnetic tapes whenever the detection threshold is exceeded. A memory buffer (sampling in retrospect) provides a few seconds of noise data preceding the event on each channel.

The SGSN has been installed for two main purposes: (1) to provide local seismicity data for phase II of the seismic risk analysis for the planned dam (Phase I has been completed), and (2) to provide seismological background data for the possibility of induced seismicity (in accordance with, e.g., recommendations from the UNESCO International Committee on Large Dams). It is obvious, moreover, that the network will be of considerable interest also from a more general seismological point of view, in particular for the study of the East African Rift System, including the Gregory Rift in whose extension the network is located. The planned operational period for the array is 2 years, although one hopes for an extension.

Preceding the installation of the network (completed in September 1978), a portable analog seismograph was operated sporadically for a few months, mainly for site surveys. Two things became obvious quite soon: (1) the local seismicity level is quite high, and some of this activity is very close to the dam site, (2) the ambient noise level is very low (seasonal variations are possible). As to the latter point, the portable seismograph

could easily be operated at a magnification of 90 000 at 1 Hz, and an analog output from the permanent stations is usually kept at a magnification of about 120 000 at 1 Hz. This means that the RMS noise level at 1 Hz is not much above 1 nm.

Following the installation of the array, two timed explosions were fired near the dam site in order to provide an initial velocity model for the area. The results indicate significant velocity variations over the array (2 of the stations are on Basement, 4 are on Karroo), while on the average the data were best satisfied by a model with a P velocity of 5.0 km/s down to 4 km, followed by a layer with 5.9 km/s. So far only a few well-recorded events have been received and analyzed, one of which is presented in Fig. VI.8.2. The earthquake is located at a depth of 10 km about 4 km

NE of the dam site (see Fig. VI.8.1). A point from which other events have been recorded as well. Most of the quakes seem to occur at depths between 10 and 20 km, and connected to a fault system (Tagalala) which runs from NW to SE very close to seismometer 2 and 3 (Fig. VI.8.1). Finally, an example of a teleseismically recorded earthquake is given in Fig. VI.8.3, showing an analog recording at one of the network stations of the disastrous Iran earthquake on 16 September 1978; M<sub>e</sub>=7.3 measured at NOTSAR.

The installation, operation and data analysis of the Stiegler's Gorge Seismic Network are funded by the Norwegian Agency for International Development.

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Station	Lat.	Long.	Elev (m)
1	7 <sup>o</sup> 56.431'S	37 <sup>°</sup> 50.445'E	204
2	7 <sup>0</sup> 44.718 <b>'</b> S	37 <sup>°</sup> 53.189'E	334
3	7 <sup>o</sup> 51.213'S	38° 2.819'E	162
4	8° 7.226'S	37°50.791'E	268
5	7 <sup>o</sup> 55.327'S	37 <sup>°</sup> 35.577'E	283
6	7 <sup>0</sup> 45.009'S	37°39,516'E	275

TABLE VI.8.1

Coordinates for the Stiegler's Gorge Seismic Network (SGSN)

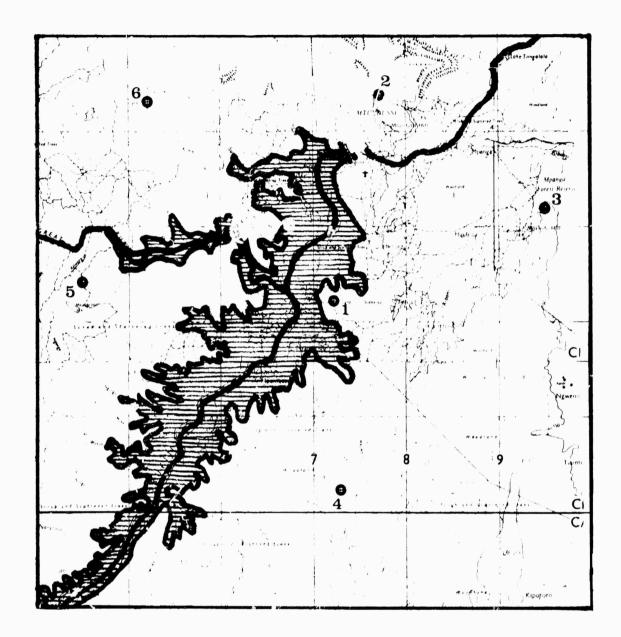
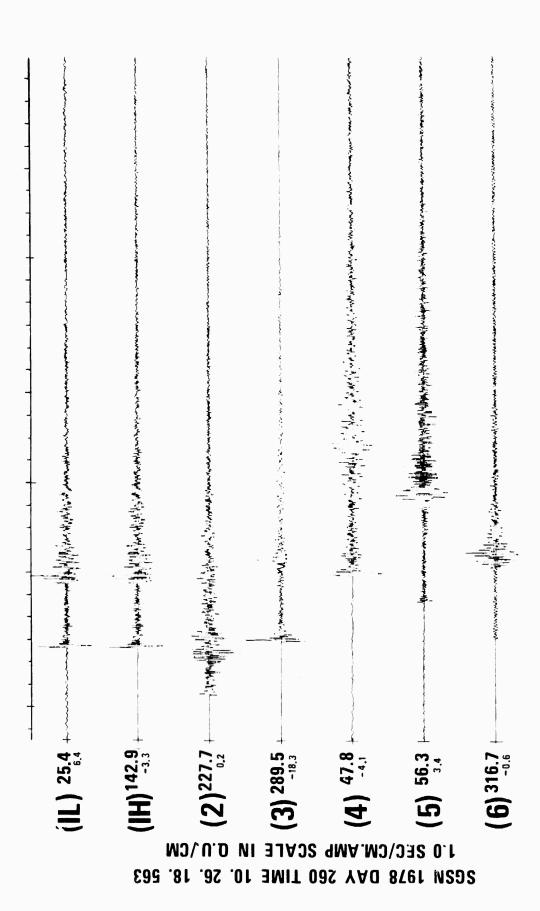
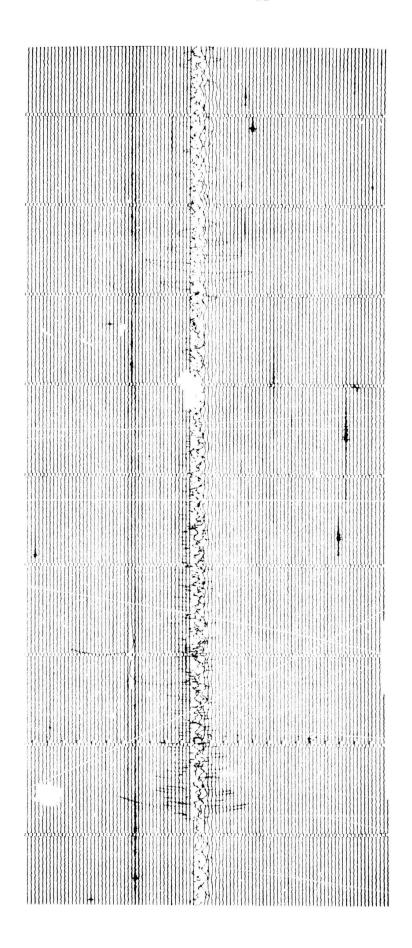


Fig. VI.8.1 Locations of the 6 stations of the Stiegler's Gorge Seismic Network (SGSN) in Tanzania. The Central Recording Station is situated just north of the dam site, close to seismometer No. 2. The water flows from SW to NE, and the hatched area is the future reservoir at a level of 150 m, which is 20-30 m below the planned maximum. The map covers 60 x 60 km.



The time marks on the top are one second apart, and the first trace is a low-gain version of Station 1. The earthquake is located near Station 2 at  $7^{0}46.36^{\circ}S$ ,  $37^{0}52.34^{\circ}E$  and at a depth of 10.1 km. Note especially the clear onsets and the changes in polarity. A local earthquake on 17 September 1978 recorded by the Stiegler's Gorge Saismic Network.

Fig. VI.8.2



disastrous Iran earthquake on 16 Sep 1978. Note especially the strong surface waves. Analog recording from Station 2 of the Stiegler's Gorge Seismic Network showing the Fig. VI.8.3

## VI.9 The Seismicity of Svalbard

Our investigation of the seismicity of Svalbard has continued as outlined in the previous Semiannual Technical Summary (STS). Extensive and detailed analysis of the recorded data have been completed up to the end of May 1978, giving us almost 6 months altogether. The stations BBG, LYR and PRD have been in operation most of the time, the WWSSN station KBS has been in operation (and available to us through the University of Bergen) all of the time, and a new station SWE became operational in the beginning of May 1978.

Some figures on the detectability of the stations are given in Table VI.9.1, where it is seen that 1258 events have been reported altogether. 878 or about 70% of these are local events from Svalbard and vicinity (including the mid-Atlantic Ridge west of Svalbard), while the rest are teleseismic. With the exception of SWE (where there have been severe problems with local noise from the mines) we see from Table VI.9.1 that the teleseismic detectability is around 2 events/day for all stations, while it is LYP which is the best station so far as local events are concerned. A breakdown of the detectability statistics on a weekly basis is shown in Fig. VI.9.1, where we see that about 30-60 events have been detected every week. The figure also shows the division between teleseismic and local events, and the number in the latter group which have been located. Altogether 566 of the earthquakes have been located, which amounts to 64% of the local ones and 45% of the total number.

Since our report in the previous STS, two important improvements have been added to our location procedure. First, a crustal model for the area has been developed using our recordings of the signals from a profiling survey performed last summer by the University of Bergen (Prof. Sellevoll, personal communication) in cooperation with the University of Hamburg and the Polish Academy of Sciences. Our preliminary model derived from these data consists of layers with P velocities of 5.7, 6.7 and 8.2 km/s, starting at depths of 0, 16 and 32 km, respectively. The P to S velocity ratio is 1.80. The second improvement in our locations is connected to the location method itself. We have tried to use location programs such as HYPO71 (from USGS) but cannot obtain satisfactory convergence because of the poor station

configuration with respect to the epicenter locations. A new method has therefore been developed which is based on the modified S-P method described in the previous STS, using the S-P location as a starting point and then refining the estimate using the absolute P times for the station for which reliable time corrections are available. Although the method in principle is similar to the one in HYPO71 (using the same intormation), it differs in the iterative procedure; our method has much stronger convergence properties for poor station configurations. The method is still being developed, and will be properly documented at a later stage.

An epicenter map from Svalbard using the new location method is given in Fig. VI.9.2, where only our most precise locations are plotted. 133 of our 566 located events are shown on this figure, and the number of phases used in each location ranges between 4 and 8 (S-P with no time correction counts as one phase). Reliable time corrections are available for at least one station for all of the events. The prominent feature is the now well-known Heerland earthquake zone, where any possible lineation still cannot be discerned, while it cannot be excluded either because of the precision of our epicenters. In comparing with the epicenter map in the previous STS, we see that the cluster of events now has been moved a little to the north-west, this is because of the new and imporved crustal model.

Besides the seismic activity along the mid-oceanic (Knipovich) ridge, there is one other feature in Fig. VI.9.2 which deserves attention, namely, the activity along the coast southwards from the WWSSN station KBS. Seven earthquakes are plotted in Fig. VI.9.2 along this (Forlandsundet) seismicity zone, which is reported here for the first time.

Some data on the measured local magnitudes are given in Fig. VI.9.3, which shows the incremental and cumulative distributions of magnitudes as measured from amplitudes. It is seen that the slope follows a value of b=1 down to a value of M=2.0, while the deviation below this value may be due to a combination of decreasing detectability, imprecise attenuation parameters in the magnitude formula, imprecise locations and the mixing of several

epicentral areas in one distribution. It is important to notice, moreover, that our Svalbard magnitude scale has so far not been calibrated with respect to absolute level, this may cause a later shift of all the magnitudes by a constant value.

The Svalbard microearthquake network is operated by the Norwegian Polar Research Institute in cooperation with NTNF/NORSAR and the Russian mining trust Arktikugol.

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- B. Kr. Hokland

	BBG	PRD	LYR	SWE	KBS	Total
Teles.	291	291	229	18	310	380
-Daily	1.8	2.1	1.9	1.2	1.7	
Local	640	511	653	58	363	873
-Daily	3.8	3.8	5.4	4.0	2.1	
Sum	931	802	882	76	673	1258
-Daily	5.6	5.9	7.3	5.2	3.8	

# TABLE VI.9.1

Detectability figures for the Svalbard microearthquake stations during the time period 8 Dec 1977 -31 May 1978. The daily averages are computed on the basis of the actual uptime for each station, which has been 95% for BBG, 77% for PRD, 67% for LYR, 8% for SWE (only part of May 1978) and 100% for the WWSSN station KBS. The locations of the stations are given in Fig. VI.9.2.

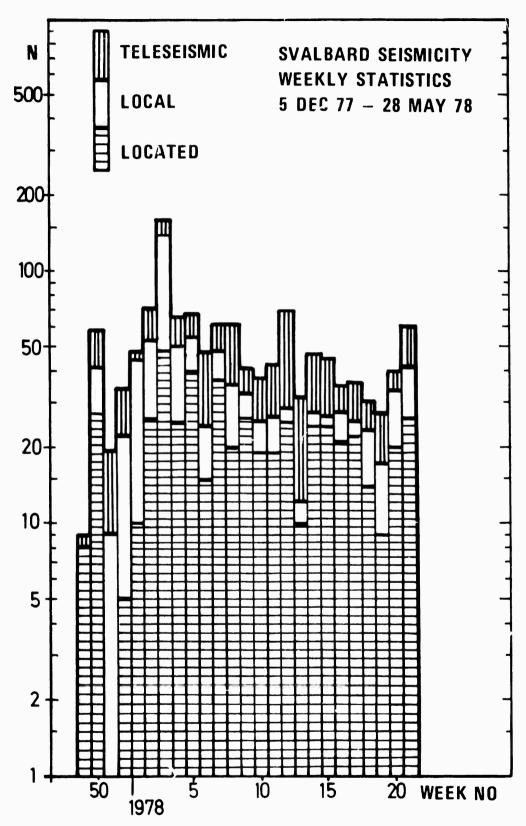
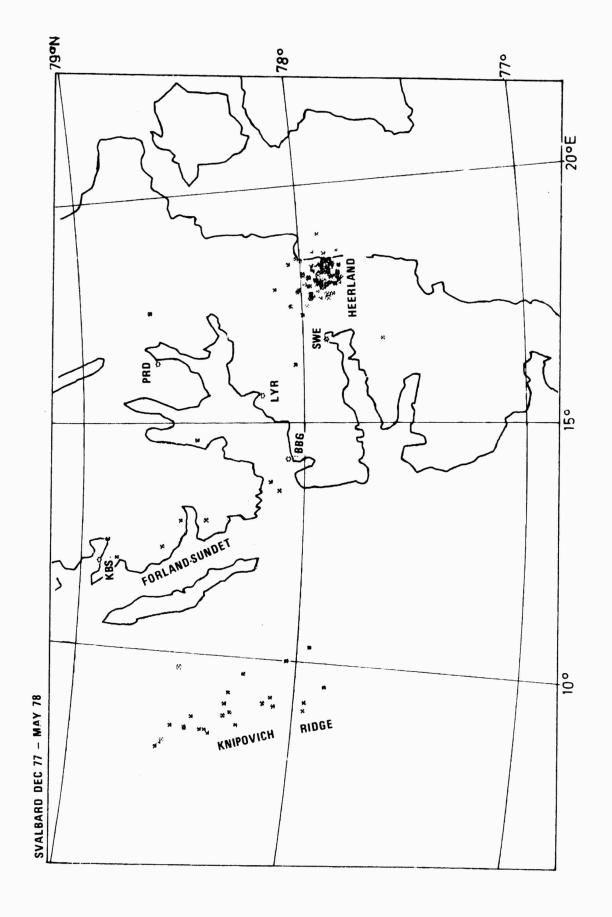


Fig. VI.9.1 Weekly breakdown of the number of earthquakes reported by the Svalbard microearthquake network in the 6 months between December 1977 and May 1978. The number of teleseismic, local and located (all local) events are shown separately. The peaks are due to swarms from the Knipovich Ridge and/or the Heerland earthquake zone (see Fig. VI.9.2).



Epicentral locations for the 133 most precisely located earthquakes from the Svalbard microearthquake network from December 1977 to May 1978. Fig. VI.9.2

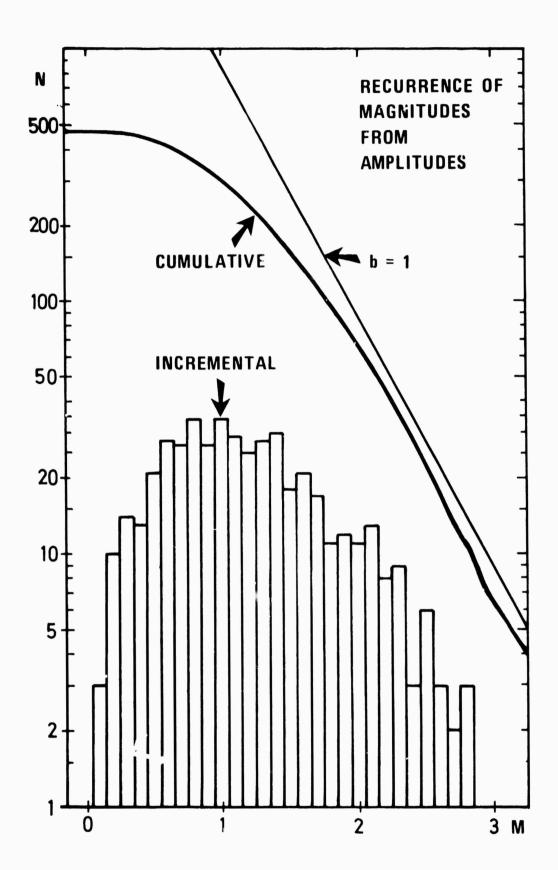


Fig. VI.9.3 Frequency-magnitude distribution for the located earthquakes in and around Svalbard, Dec 77 - May 78. The formula used is  $M=\log(\Lambda) = \alpha_1 + \alpha_2 \log \Delta + C$  where  $\Lambda$  is maximum amplitude,  $\Lambda$  is epicentral distance.

where A is maximum amplitude, A is epicentral distance,  $\alpha_1{}^{and}$   $\alpha_2$  are parameters, and C is a constant.