



Low Altitude Measurements of Precipitating Protons, Alpha Particles, and Heavy Ions During the Geomagnetic Storm of 26-27 March 1976

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Space and Missile Systems Organization 12. REPORT DATE 15 August 1978 Air Force Systems Command 3. NUMBER OF PAGES Los Angeles, Calif. 90009 32 14. MONITORING AGENCY NAME & ADDRESS(II different from Controlling Office) 15. SECURITY CLASS. (of this report) Unclassified 15a. DECLASSIFICATION/DOWNGRADING 16. DISTRIBUTION STATEMENT (of thie Report) Approved for public release; distribution unlimited. 17. DISTRIBUTION STATEMENT (of the abstract entered in Block 20, if different from Report) 18. SUPPLEMENTARY NOTES 19. KEY WORDS (Continue on reverse side if necessary and identify by block number) Heavy Ions Precipitating Ions -approx Magnetic Storm ABSTRACT (Continue on reverse side if necessary and identify by block number) During the geomagnetic storm of 26 March 1976, observations were made of precipitating protons, alpha particles, and heavy ions by two counter telescopes aboard the low-altitude S3-2 satellite. These observations are presented in this paper and their significance discussed. Briefly, a two zone precipitation region is observed on the night side, at $L \approx 2.7$ and $L \approx 4.0$ with an \Re/p ratio $\Re \times 10^{-3}$ in the low latitude region and $\Re/9 \times 10^{-4}$ at higher latitude at 0.4 MeV/nucleon. The heavier ions, CNO nuclei, are seen in the low latitude zone with a CNO/α ratio of 5.6 × 10-2 at 0.25 MeV/nucleon. DD FORM 1473 * .05 UNCLASSIFIED Cha SECURITY CLASSIFICATION OF THIS PAGE (When Date Entered) of about .008

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PREFACE

The design, construction, and integration of the instruments was the work of several people. The authors would like to acknowledge the substantial efforts of Erick Kunneth and Otto Vollmer of the Max-Planck-Institut, Norm Katz and Sam Imamoto of The Aerospace Corporation, and Ed Tums and John Cain of the University of Maryland.

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INTRODUCTION

Detailed low altitude measurements of precipitating protons at energies above 100 keV, and at lower, ring current energies have been reported by a number of authors (e.g., Amundsen et al., 1972; Mizera, 1974; Søraas and Berg, 1974; Hauge and Søraas, 1975; Søraas et al., 1977). From these measurements the following picture has emerged: During geomagnetically quiet conditions there exists at high latitudes in the evening-midnight sector a region where the protons have a roughly isotropic pitch angle distribution. Equatorward of the isotropic region there is a latitude interval where the pitch angle distribution is anisotropic (with no particles in the loss cone). At still lower latitudes there appear again protons in the loss cone, indicating weak to moderate pitch angle diffusion. During geomagnetically disturbed times the equatorward boundary of the isotropic region moves to lower latitudes and has been observed as low as L=3 (Lindalen and Egeland, 1972).

Low altitude measurements reveal only a small part of the equatorial pitch angle distribution. However, near equatorial measurements of energetic protons on the Explorer 45 (S^3) satellite by Williams and Lyons (1974a, 1974b) confirm the picture described above. During the geomagnetic storm recovery phase they observed isotropic pitch angle distributions with a nearly empty loss cone above the plasmapause region. In the region of the plasmapause they observed a transition from flat to rounded pitch angle distributions, indicating weak pitch angle scattering. Apogee of Explorer 45 was not high enough to observe the outer region of strong turbulence except during the geomagnetic storm main phase when this region moves to lower altitudes. During one geomagnetic storm Williams and Lyons (1974b) reported observations of flat pitch angle distributions with a full loss cone in the storm main phase at L ≈ 4.2 .

In this report we present observational data on precipitating protons which confirm these earlier results and, for the first time, data on precipitating alpha particles and heavy ions with Z > 4. Composition measurements of energetic particles in the magnetosphere are still rare (for recent reviews see, e.g., Fennell and Blake, 1976; Fritz, 1976), and new data on this subject should provide additional insight into the origin and dynamics of magnetospheric particles.

SATELLITE AND INSTRUMENTATION

The data presented in this paper were obtained from a proton and a heavy ion telescope on board the polar orbiting S3-2 satellite (inclination 96.3° , perigee ~ 230 km, apogee ~ 1500 km). The satellite is spin-stabilized with the spin axis maintained normal to the orbital plane; the spin period is 18.8 seconds. The two instruments are mounted with their apertures orthogonal to the spin axis.

Because of the large differences of proton, alpha particle and heavy ion fluxes in the radiation belt, the dynamical range required to measure all ion species simultaneously with reasonable statistical significance could not be achieved with a single detector system. Therefore two different instruments were employed.

The proton spectrometer uses a single silicon semiconductor detector of 200 μ m thickness for energy analysis, followed by a solid state anticoincidence detector. The particle acceptance angle is limited to 18° full angle, resulting in a geometrical factor of 0.04 cm² ster. The 5 integral proton channels p1 to p5 and the one integral alpha channel α 0 are listed in Table 1.

The alpha particle and heavy ion telescope is similar to the ULET instrument of the University of Maryland/Max-Planck-Institute experiment on IMP 7 amd IMP 8 (Hovestadt and Vollmer, 1971). The telescope is a three detector device. The first detector is a thin proportional counter filled with 125 μ g/cm² of isobutane. In this detector the specific energy loss Δ E of the penetrating particle is measured. The residual energy E is deposited in the second detector, a silicon semiconductor of 800 μ m thickness. Particles penetrating the second detector, are rejected by an anticoincidence

detector. The field-of-view is 49° full angle; the geometric factor is 0.52 cm^2 ster. This telescope provides a unique identification of protons, alpha particles and heavy ions (Z>4 and Z>16). The effective threshold for both C and O in the Z>4 channel is $\approx 0.25 \text{ MeV/nucleon}$. The channels are listed in Table 1. All channels are read out once per sec, except for the proton channels pi1 and pi2, the alpha particle channel $\alpha 6$ and the single rates in the three detectors. The latter channels are multiplexed in an 8-channel multiplexer with a 1 sec simultaneous accumulation time. A randomly selected sample of the heavy ion events is pulse height (2D) analyzed with up to 3 analyses per sec.

Because of the large geometrical factor of the ion telescope, the proton channels pil and pil saturate when passing through the outer radiation belt. Therefore, these channels are used only for a consistency check with the proton telescope measurements in regions of space where suitable fluxes were encountered. These checks reveal an excellent agreement of the proton fluxes measured by the two different systems. The saturation effects mentioned above occur in the proton logic and not in the detectors, amplifiers or discriminators; the saturation effects do not impair the heavy ion channels.

The alpha particle channels $\alpha 1$ to $\alpha 6$ consist of two groups with different thresholds in the ΔE signal. The low energy channels $\alpha 1$ to $\alpha 3$ have a ΔE -threshold 5.6 times higher than the ΔE -threshold utilized in the high energy channels $\alpha 4$ to $\alpha 6$. Where the proton single rates are very high in the radiation belts, some chance coincidence and pile-up effects contaminate the alpha particle channel $\alpha 4$. Only $\alpha 4$ is affected.

The excellent background rejection capability of the ion telescope is demonstrated by the following: In the inner radiation belt the $\alpha 1$ rate is < 1 count/sec in spite of the high singles rates in the ΔE and E channels of $\sim 5 \times 10^4/\text{sec}$ and $\sim 10^5/\text{sec}$, respectively.

In this paper we use only the rates with high ΔE thresholds. The Z>4 channel has a ΔE threshold of 4.5 times the $\alpha 1$ to $\alpha 3$ thresholds and is therefore even less sensitive to pile-up and chance coincidence background than $\alpha 1$ to $\alpha 3$.

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OBSERVATIONS

The data presented in this paper cover the period March 24-28, 1976. A large geomagnetic storm ($D_{st} = -230\gamma$) and an energetic solar particle event occurred within this time period. The top panel of Figure 1 shows the D_{st} value; the bars in the top panel indicate the time intervals of data retrieval.

The large geomagnetic storm occurred after the SSC on March 25, 2339 UT. The $D_{\rm st}$ index remained quiet until ~ 0200 UT on March 26, then started to decrease suddenly and reached -230 γ at 0800 UT. Thereafter the $D_{\rm st}$ index slowly recovered to reach -80 γ to -90 γ on March 28, 1976. In the following, we will investigate the time development of fluxes and pitch angle distributions of protons, alpha particles, and heavy ions measured at low altitudes during the formation and decay of the equatorial ring current. Examples of the particle measurements during various satellite passes are shown in Figures 2 to 4.

Figure 2a is an example of the fluxes measured during a pre-storm crossing of the outer radiation belt. The pitch angle distributions are strongly peaked at 90° to the geomagnetic field. The alpha particle and proton channels of the two instruments do not cover the identical energy per nucleon range. In order to determine the α/p ratios at the same energy per nucleon we have adopted the following procedure: The measured fluxes in adjacent channels are connected by power laws. The spectral exponents and the absolute intensities are determined by the fact that the integrals over the power laws between the channel boundaries must be equal to the measured fluxes. From the power laws we then derive α/p ratios at different energy per nucleon values. During the prestorm quiet time we find at L = 3.87, B = 0.307 G that the α/p ratios (perpendicular to the geomagnetic field) are $\sim 1.2 \times 10^{-3}$, 6.1×10^{-4} and 3.3×10^{-4} at 0.4, 0.5, and 0.6 meV/nucleon, respectively.

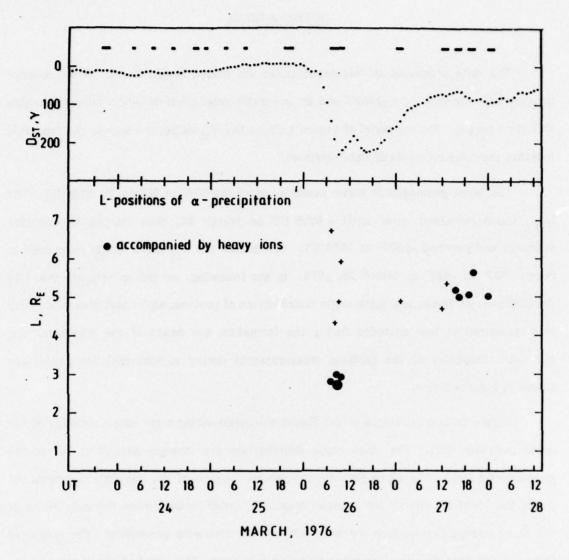


Fig. 1. Precipitation Zones of Alpha Particles and Heavy Ions as Function of UT. The value of D_{st} as a function of UT is also shown.

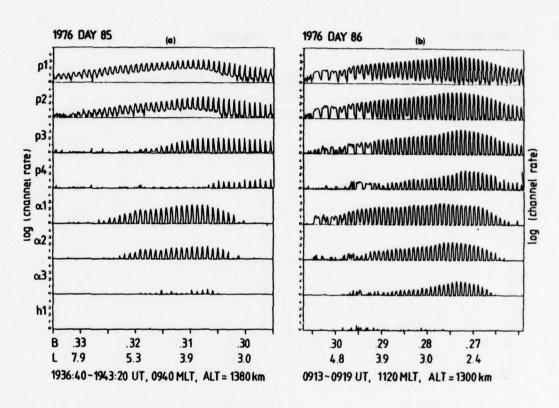


Fig. 2. Proton, Alpha Particle, and Heavy Ion Countrates for (a) Pre-Storm Pass, and (b) Dayside Pass at End of Storm Main Phase

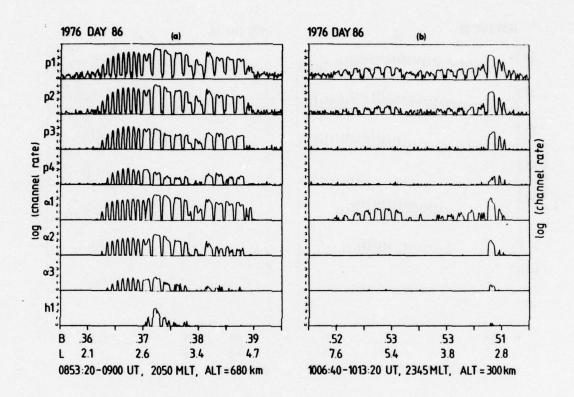


Fig. 3. Proton, Alpha Particle, and Heavy Ion Countrates for Two Nightside Passes at the End of the Storm Main Phase

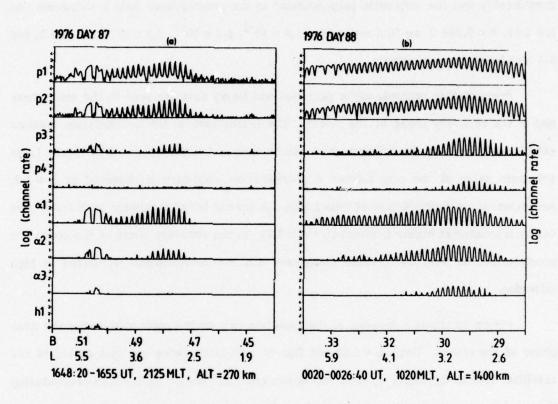


Fig. 4. Proton, Alpha Particle, and Heavy Ion Countrates for (a) Nightside Pass, and (b) Dayside Pass in Storm Recovery Phase

With the formation of the ring current, the L-location of the flux maximum of the trapped alpha particles moves to slightly lower L-values, the absolute fluxes increase considerably and the α/p ratio perpendicular to the geomagnetic field is enhanced. At L = 2.48, B = 0.368 G we find ratios of $\sim 1.8 \times 10^{-2}$, 5.9 x 10^{-3} , 2.3 x 10^{-3} at 0.4, 0.5, and 0.6 MeV/nucleon, respectively.

Precipitating protons, alpha particles and heavy ions are seen in the main phase and in the recovery phase of the storm. The L-locations of the precipitation maxima exhibit a strong time dependence which can be seen in the bottom panel of Figure 1. In the main phase of the ring current a precipitation maximum is observed at $L\approx 2.8$, accompanied by precipitation of heavy ions. A second broad maximum with lower flux values is located at higher L-values (L = 4.5-7.0). In the recovery phase of the storm the precipitation at the low L-value disappears and the precipitation is limited to high latitudes.

Figure 3a shows a crossing of the radiation belt on the night side during the main phase of the storm. Until L = 2.53 the flux is modulated twice per spin period of the satellite, indicating empty upward and downward loss cones. At L = 2.59 precipitating particles can be seen in the upper loss cone; from there poleward the upper loss cone is filled up, suggesting that the particle distribution is in strong pitch angle scattering. The α/p ratio in the maximum of the precipitation (L = 2.71) is $\sim 8 \times 10^{-3}$, 2.1 $\times 10^{-3}$, and 5.8 $\times 10^{-4}$ at 0.4, 0.5, and 0.6 MeV/nucleon, respectively. Figure 5 shows detailed pitch angle distributions of protons and alpha particles at the equatorial edge of the precipitation region. The transition from rounded (peaked at 90° to the geomagnetic field) pitch angle distributions to flat (filled upper loss cone) distributions occurs at L = 2.59. The higher energy protons (p4) and alpha particles (α 3) show filled loss cones,

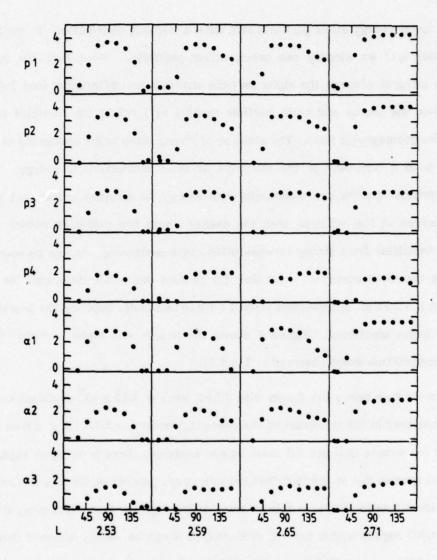


Fig. 5. Pitch Angle Distributions of Protons and Alpha Particles on 26 March 1976. Each individual plot shows log10 integral countrate (for protons) or log10 differential countrate (for alpha particles) versus local pitch angle for a specific energy and L value.

whereas the lower energy alpha particles still show a trapping distribution. In the lowest proton channel (p1) we already see precipitating particles. Note that the protons channels are integral whereas the alpha particle channels are differential (see Table 1). Figure 6 shows the proton and alpha particle spectra at L = 2.59 for particles moving parallel to the geomagnetic field. The absence of strong pitch angle scattering at lower energies leads to a turn-over of the spectrum at some characteristic energy. If the equatorial particle spectra are monotonic with energy to energies below that of the turnover observed at low altitude, then the energy where the turn-over occurs should indicate the transition from strong to weak pitch angle scattering. As can be seen from Figure 6, the turn-over occurs at ~ 2.0 MeV for protons and ~ 0.5 MeV/nucl. for alpha particles. At L = 2.71 all distributions exhibit a filled loss cone, implying the presence of strong pitch angle scattering. Figure 7 shows the proton and alpha particle spectra during the precipitation event observed at L = 2.71.

Extremely high heavy ion fluxes with J (>3 MeV) = 5.23×10^2 particles/cm² sec ster are encountered in the maximum of the precipitation at L = 2.71. In a Δ E vs E plot of the heavy ion events that are 2D pulse height analysed, there is no track separation into individual species due to the fact that the events are just above the sensor threshold. Thus an ion identification is impossible in spite of the high ion rate. However, the ions fall into the CNO regime within the Δ E vs E area, so it can be safely assumed that they belong to the CNO group. This heavy ion channel responds to the integral flux of CNO nuclei above 250 keV/nucleon. Integration of the alpha particle spectrum (Figure 7) above 250 keV/nucleon gives a CNO/ α ratio of $\sim 6.5 \times 10^{-2}$.

In the high latitude precipitation regions (L \approx 3.8) during the main phase, we derive α/p of \sim 9.1 x 10⁻⁴ and 2.1 x 10⁻⁴ at 0.4 and 0.5 MeV/nucleon, respectively.

TABLE 1

| | | SE | SENSOR PARAMETERS | S | |
|-------------|--|-------------|-------------------|--|----------------------|
| Proton tele | Proton telescope (g = 0.04 cm ² ster) | | Ion teles | Ion telescope ($g = 0.52 \text{ cm}^2 \text{ ster}$) | r) |
| Channel | Energy (MeV) | Channel | Particles | Energy (MeV) | Energy (MeV/nucleon) |
| p1 | 0.40 | α1 | alphas | 0.95 - 1.42 | 0.237 - 0.355 |
| p2 | 0.55 | 0.2 | • | 1.42 - 2.03 | 0.355 - 0.508 |
| p3 | 1.00 | G 33 | • | 2.03 - 2.89 | 0.508 - 0.722 |
| P4 | 2.00 | 94 | • | 2.89 - 4.75 | 0.722 - 1.19 |
| p5 | 5.11 | Q.5 | • | 4.75 - 7.33 | 1.19 - 1.83 |
| 020 | 99.6 | 98 | | 7.33 - 24 | 1.83 - 6.0 |
| | | hI | 2 > 4 | | >.250 |
| | 12 | h2 | Z > 16 | >12.0 | 716 |
| | | pi1 | protons | 0.48 - 0.76 | |
| | | pi2 | | 0.76 - 1.43 | |

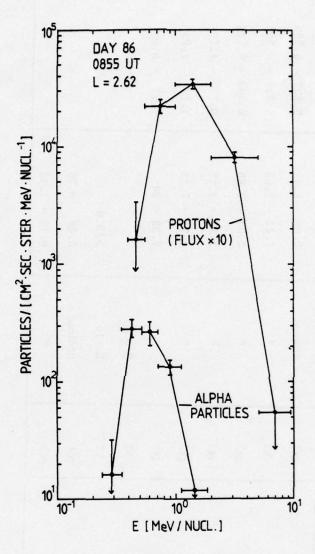


Fig. 6. Proton and Alpha Particle
Differential Spectra Parallel
to the Magnetic Field
(upward) at L = 2.59 on
26 March 1976 at 0855 UT

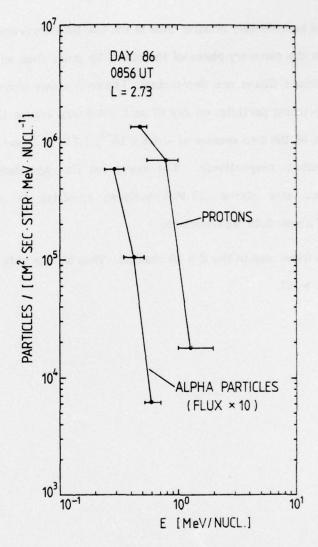


Fig. 7. Proton and Alpha Particle
Differential Spectra in the
Precipitation Event at
L = 2.7 on 26 March 1976
at 0856 UT

Thus, the ratio is considerably smaller than in the low latitude precipitation region. This ratio increases in the recovery phase of the storm by more than an order of magnitude, although the absolute fluxes are decreasing. Figure 8 shows proton and alpha particle spectra of precipitating particles on day 87 at L = 4.8 (see Figure 4). From Figure 7 we derive flux ratios of the two species of $\sim 4.6 \times 10^{-2}$, 1.7×10^{-2} and 6.7×10^{-3} at 0.4, 0.5 and 0.6 MeV/nucleon, respectively. The heavy ion flux has decreased to 2.3 ± 0.94 particles/cm² sec ster above .25 MeV/nucleon, resulting in a CNO/ α ratio of $(1.1 \pm 0.5) \times 10^{-2}$ above 0.25 MeV/nucleon.

No events were seen in the Z > 16 channel. Thus we can only set an upper limit of 0.5% for CNO/(Z > 16).

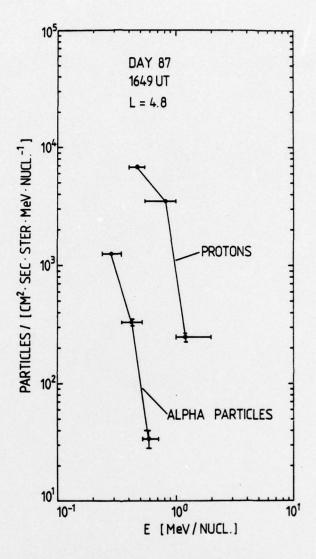


Fig. 8. Proton and Alpha Particle
Differential Spectra in the
High Latitude Precipitation
Zone at L = 4.8 on 26 March
1976 at 1649 UT

SUMMARY OF OBSERVATIONS

The low-altitude data presented in this report can be summarized as follows:

- 1. Precipitating protons, alpha particles, and heavy ions have been observed during the storm main phase down to L values of ≈ 2.7 at the night side. The pitch angle distributions are iostropic except for the downward loss cone, implying the existence of strong wave turbulence down to very low L values.
- 2. The transition from the isotropic to the rounded pitch angle distributions occurs at an L value that decreases with increasing proton and alpha particle energy. At the time the precipitation is seen at the lowest latitudes, at L = 2.6, ~ 0.6 MeV/nucleon particles are in strong pitch angle diffusion while 0.3 MeV/nucleon particles are only weakly scattered. The transition from strong to weak scattering occurs at about the same total energy for protons and alpha particles.
- 3. A two zone night side precipitation pattern exists during the storm main phase, with a maximum at L \approx 2.7 and a second broad maximum at higher latitudes (L \approx 4). The α/p ratio is \sim 8 x 10⁻³ in the equatorward region and \sim 9 x 10⁻⁴ in the high latitude precipitation zone at 0.4 MeV/nucleon. Precipitating heavy ions of energies > .25 MeV/nucleon, presumably CNO nuclei, are found in the lower latitude precipitation zone. A CNO/ α ratio of \sim 6.5 x 10⁻² above \approx 0.25 MeV/nucleon is derived.

- 4. The low latitude precipitation zone disappears in the storm recovery phase. The α/p ratio in the high latitude zone considerably increases during the recovery phase with values up to $\sim 5 \times 10^{-2}$ at 0.4 MeV/nucleon, and a CNO/ α ratio of $\sim 1.1 \times 10^{-2}$ above 0.25 MeV/nucleon has been found.
- 5. No counts were seen in the Z > 16 channel, giving an upper limit of 0.5% for CNO/(Z > 16) actually (Z > 4)/(Z > 16) at ≈ 0.25 MeV/nucleon.

DISCUSSION

The energy dependence of the low altitude equatorial boundary of the precipitation region has been previously described by Mizera (1974) and Søraas and Berg (1974) for protons. The relation of the equatorward extension of the strong precipitation to D_{st} has been found by Mizera (1974) and by Hauge and Søraas (1975). Williams and Lyons (1974b) report equatorial pitch angle measurements of protons from which they also conclude that the outer region of strong turbulence penetrates closer to the earth during the storm main phase than during the recovery phase. Søraas and Berg (1975) propose that the proton precipitation is caused by self-excited ion cyclotron waves. In this case the condition for resonance with a growing wave is

$$E_p > E_M A^{-2} (1 + A)^{-1}$$
 (1)

where E_p is the energy of the resonant protons, $E_M = B^2/8\pi N$ is the magnetic energy per particle, B is the magnetic field strength, N is the total plasma density and A is a measure of the proton pitch angle anisotropy. Since E_M decreases sharply with decreasing L at the plasmapause, proton precipitation by self-excited ion cyclotron waves would explain the energy dependence of the equatorward precipitation boundary. The equatorward precipitation zone should not be confused with the high latitude strong precipitation zone. The latter exists also during magnetic quiet times, and only a weak energy dependence of the low latitude boundary between one and several hundred keV (Søraas et al., 1977) has been observed there.

Let us now assume that the strong pitch angle scattering is due to a resonant interaction with electromagnetic waves. Whether these waves are self-generated by the protons or the interaction is parasitic, the resonance condition (see, e.g., Kennel and Petschek, 1966) reads

$$v_{||} = v_{ph} \left[1 - (Q^{\dagger}/\omega) \right]$$
 (2)

where $v_{||}$ is the component of the particle velocity parallel to the magnetic field, v_{ph} is the phase velocity of the waves causing pitch angle scattering, \mathcal{Q}^+ is the particle gyrofrequency and $\omega/2\pi$ is the wave frequency. In the low frequency approximation the phase velocity is the Alfven velocity v_{A} and the resonance frequency is

$$f = \omega/2 = \Omega^{+} v_{A}/2\pi v_{H}$$
 (3)

Since the gyrofrequency of protons and alpha particles differs by a factor 2, the resonance frequency for particles of the same total energy is equal. In the case of Alfven waves propagating parallel to the magnetic field the pitch angle diffusion coefficient is given (e.g., Hasselmann and Wibberenz, 1968) by

$$D_{\mu\mu} = \frac{1}{8} \left(\frac{Q^{+}}{B}\right)^{2} \frac{v_{A}}{v_{tt}} \left(\frac{1-\mu^{2}}{\mu}\right) P(f)$$
 (4)

where μ is the cosine of the pitch angle and P(f) is the power spectrum of the waves at the resonant frequency. This implies that the diffusion coefficient of alpha particles is half the diffusion coefficient of protons at the same energy. The minimal condition for strong scattering is

$$D_{\mu\mu} \text{ (strong)} \ge 2 \alpha_L^2 / \tau_B$$
 (5)

where α_L is the equatorial loss cone and τ_B is the bounce period. Since the bounce period is inversely proportional to velocity we have finally the result that if protons of a certain energy are in strong pitch angle diffusion, then the same is true for alpha particles of the same total energy and vice versa. This discussion is valid for the case of stripped helium ions, i.e., He^{++} . The radiation belt helium ions are expected to be both singly ionized and doubly ionized (Cornwall, 1972; Spjeldvik and Fritz, 1978a). What has been said above of H^+ and He^{++} is equally true of He^+ at 1/4 of the energy. Heavier ions, CNO ions in particular, will not be totally stripped in the energy range of ≈ 0.25 MeV/nucleon (Spjeldvik and Fritz, 1978b). Thus the value of e/m for the CNO ions will be < 1/2 and they will be in strong pitch angle diffusion at energies below that of He^{++} and H^+ . As can be seen from Figure 7, the turn-over in the spectra, i.e., the lowest energy where particles are still in strong diffusion, is at about the same energy for protons and alpha particles but the change in slope of the alpha particle spectrum is less than that of the proton spectrum. This result would be expected if the helium ions are a mixture of He^+ and He^{++} and there are significant abundances of both ionization states.

We have interpreted the isotropic distributions as being due to strong pitch angle diffusion. If the pitch angle diffusion process occurs in the near equatorial region, the ratios of different species should be characteristic of equatorial ratios. The α/p ratios of equatorially mirroring particles have been reported by Blake et al. (1973) and Fennell

and Blake (1976) for L \leq 2.5 and by Fritz and Williams (1973) for larger L values. At L = 3 Fritz and Williams (1973) observed a ratio of \sim 8 x 10^{-3} , which compares favorably well with our ratio measured at L = 2.7 at the end of the main phase. The ratio of $\sim 10^{-3}$ at L = 4 during the storm main phase is low compared with a ratio of $\sim 10^{-2}$ given by Fritz and Williams; however, during the storm recovery phase our ratio increases up to \sim 5 x 10^{-2} . Large increases of the α /p ratio at low altitudes coincident with geomagnetic disturbances have been observed (Van Allen and Randall, 1971; Krimigis and Verzariu, 1973; Randall, 1973). If our interpretation of the precipitation events as being due to near-equatorial strong scattering is correct, the α /p ratio near the equator also can show large variations. Large variations at low altitudes thus are not necessarily the result of relative changes of the pitch angle scattering strength for protons and alpha particles.

The first measurements of magnetospheric ions heavier than alpha particles were made by Krimigis et al. (1970) and Van Allen et al. (1970). From these low altitude measurements they derived (in the range 3.0 \leq L \leq 3.5, 0.15 \leq B \leq 0.2) a ratio of CNO/ α of \sim 3 x 10⁻³ above \approx 0.3 MeV/nucleon. Our measurements at the end of the storm main phase at L = 2.7 result in a CNO/ α ratio above \approx 0.25 MeV/nucleon which is about one order of magnitude higher. Interpreting the ratio in the precipitation event as characteristic of the near equatorially mirroring paricles, we would expect a higher CNO/ α ratio at large equatorial pitch angles than at small equatorial pitch angles during quiet conditions. Fritz and Wilken (1976) observed at the geostationary orbit an injection of heavy ions during an isolated substorm with a (Z > 2)/ α ratio of \sim 1.2 x 10⁻¹ at 0.15-0.23 MeV/nucleon; thus the equatorial CNO/ α ratio may well be time dependent. A definite conclusion about the origin of the heavy particles is not possible since the ratio of number densities of CNO to He is similar in the solar wind (Bame et al., 1975) and in the ionosphere (Taylor, 1973), as is also the H to He ratio. A conclusive test would be the measurement of the C to O ratio (Blake, 1973).

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