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THE INFLUENCE OF VARIATION IN GRAIN SIZE AND MINIMAL VARIATION IN ROCK TYPE ON THE QUALITY OF ROCK PROPERTY CORRELATIONS FOR INTACT IGNEOUS ROLKS

ARMY ENGINEER WATERWAYS EXPERIMENT STATION, Vicksburg, Mississippi

FEBRUARY 1971

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THE INFLUENCE OF VARIATION IN GRAIN SIZE AND MINIMAL VARIATION IN ROCK TYPE ON THE QUALITY OF ROCK PROPERTY CORRELATIONS FOR INTACT IGNEOUS ROCKS

R. W. Crisp

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February 1971

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FOREWORD

This report is based on a chesis prepared by Mr. Hebert W. Crisp of the Concrete Division of the U.S. Army Engineer Waterways Experiment Station (WES) in partial fulfillment of the requirements for the degree of Master of Spience in the Department of Civil Engineering, Mississippi State University.

The investigation was conducted at the Concrete Division, WHS, from June 1970 to September 1970. Mr. C. R. Mallford performed the petrographic analysis.

Directors of the WES during the conduct of the investigation and the proparation and publication of this report were COL Levi A. Brown, CE, and COL Elmost D. Peixette, CE. Technical Director was Kr. F. R. Brown.

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CHAPTER 1

INTRODUCTION

Background

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1

Efforts exerted ever the past several years to develop techniques and standards to allow for the compotent design and comstruction of engineering structures in rock have led to the establishment of several particular mechanical rock properties in positions of prominence. According to Obert (9)*, these "meat important physical properties for design purposes are density, Young's modulus, compressive strength, and flexural strength."

The determination of such physical properties, accompliabed through laboratory toring of samples of rock core and field testing of pertiens of the in-situ rock mass, is, as a rule, quite time comsuming and expensive. It is for this reason that correlations of physical properties of rock, and predictions of one property from an already determined value of another property, should be transmodens assets, provided the quality of the correlations is such that the element of doubt regarding test results and predicted properties is not of a magnitude necessitating an increase in the factor of safety.

One particular situation in which such correlations would prove of value might be the site evaluation and selection program.

*Numbers in parenthesis refer to references in the Bibliography.

Assuming several pites were being considered for a particular structural endeavor, preliminary elimination might well be fa-ilitated if one or two of the less time consuming and less expensive physical tests would yield data from which one could reasonably predist other, more difficult to determine, physical properties to be used in the evaluation and elimination process. It is also possible that correlations such as these, if of sufficient quality, would allow for reduction in the variety of tests required to determine the physical properties now deemed necessary for competent design and construction in rock media.

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Another significant application of physical property correlations might be the elimination of some destructive physical testing, to be replaced by nondestructive testing. As indicated by Obert (8), nondestructive tests can be repeated a number of times on the same specimens, facilitating determination of and compendation for procedural and instrument errors. This would allow one to separate the variations due to instrument error from the variations due to actual differences in mechanical properties of the specimens tested.

Previous Studies

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In the past, laboratory investigations and correlations of physical properties of rock have generally been limited to intact specimens, i.e., rock cores which are macroscopically homogeneous and free of discontinuities such as seams, joints, fractures, and inclusions. However, even when these investigations have been restricted to intact rock, thus eliminating the highly variable fracture parameter, data plots have frequently been highly scattered in nature resulting in correlations of questionable value.

Usually, investigations of this nature have encompassed wany rock types (3) (4) (6) (7), the determined physical properties for all rock types being lumped together and analyzed in mass. This precedure is eriented toward determining general relationships characteristic of the entire group of specimens examined. But the implication here, namely that large variations in mineral composition, geologic history, and grain size (which usually enter into the classification of rock materials) have little or no effect or the relationships between physical properties, is dubious (10).

Thus, a definite need exists to investigate the relationships between various rock properties, focusing attention on ...dividual rock types in an attempt to reduce the number of variables and eliminate some of the scatter typical of previous investigations. Hopefully, the resulting correlations will be of a quality which will allow for the elimination of repetitive testing and data reduction, facilitating more economic design of engineering structures.

Objectives of This Investigation

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This study will be primarily directed toward determining the influence of variation in rock type on the quality of correlations obtained through linear correlation analysis of various physical rock properties. An attempt will also be made to determine the effect of variation in grain size within a particular rock type on the nature of the correlations between physical properties determined for specimens of this rock type. Specimens of tonalite and granite will be propared and tested, the following physical properties being determined:

- (a) density
- (b) compressional pulse velocity
- (c) shear pulse velocity
- (d) Young's modulus of elasticity (static)
- (e) ultimate uniaxial compressive strength.

Pulse velocities will be determined according to the ASTM proposed Standard Method of Test for Laboratory Determination of Ultrasonic Pulse Velocities and Elastic Constants of Rock. Ultrasonic elastic constants will be computed from properties 1, 2, and 3. Static Young's modulus will be computed from stress-strain curves deter--mined during uniaxial compressive tests. Representative specimens will be subjected to petrographic examination (X-ray diffraction analysis, modal analysis, etc.). In us far as is possible, the specimens of tonalize and granite will represent samples of various grain sizes, and mineral composition will be varied within the limits of the classification system (11). All specimens will be intact, i.e., free of macroscopic discontinuities such as fractures, joints, seams, and vesicles.

The data accumulated will be grouped and analyzed according to rock type, and then, for comparative purposes, analyzed in mass. Correlations will be made between various pairs of the physical properties determined. In particular, ultimate uniaxial compressive strength and static Young's modulus will each be correlated with ultrasonic pulse velocities and the various uluraposic elastic constants. Comparisons will be made of the quality of correlations obtained from the data grauped according to rack type and from the data treated in their suffrety, and an effort made to determine the influence of data apalysis by rock type on quality of the correlations obtained.

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Correlations will be made using physical properties determined for three groups of specimens within one particular rock type (tonalite). These three groups will be essentially of the same mineral composition and geologic history, the variable being grain size. The intention will be to evaluate the contribution of variation in grain size to the nature of the physical property correletions obtained within the particular rock type.

CHAPTER 2

EXPERIMENTAL TECHNIQUE

General

Seventy-nine samples of rock core representing two rock types (granite and tonslite) were prepared and tested in the course of this investigation. These specimens were removed from 10 drill sites in six geographic localities. Generally, eight specimens of one particular rock type, either granite or tonslite, were selected from the core from each of the drill sites. All specimens tested were intact (contained no macroscopic joints or fractures) and essentially homogeneous NX-size (nominal 2-1/8-inch diameter) cylindrical cores.

Tests were conducted during this investigation to determine the following:

- (a) rock type and mineral composition
- (b) bulk density
- (c) ultrasonic pulse velocities (compressional and shear)
- (d) ultimate uniaxial compressive strength

(e) static stress-straip relations.

Ultrasonic elastic constants were computed from measured ultrasonic pulse velocities and specific gravities. Static Young's moduli were determined from the axial stress-strain relations observed and recorded during the uniaxial compressive tests.

Petrographic Examination

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One representative specimen from each of the 10 groups (10 drill sites) was selected for limited petrographic examination. These specimens were saved axially, one saved surface of each specimen being polished and photographed at normal size.

Composite samples were taken from the remaining pertions of the selected specimens, and ground into a fine pewder so as to pass a No. 325 sieve (44µ). X-ray diffraction patterns were made of each sample. These patterns were then examined to make mineralogical identifications and comparisons. Small portions of each of these powdered samples were tosted in dilute hydrochloric (HCl) acid and with a magnetized needle to detect the presence of carbonate minerals and magnetic minerals, respectively. All X-ray patterns were made with an XRD-5 diffractometer using mickel-filtered cooper radiation.

Thin sections were prepared from each specimen and examined with a Spencer pelarizing microscope. A point-count modal analysis was made on each thin section to determine the mineral composition by percent and grain size (13) of each of the rocks represented. The number of counts per section was held constant (500), but spacing of the counter was varied with grain size in an attempt to obtain a representative statistical average (2).

A summary of the results of the petrographic examination of representative samples from each of the 10 groups of core is given below. Detailed results are given in Appendix 1.

(a) <u>Tonalite</u> (Vermilion granite formation, Minnesota). Brownishgray, medium- to coarse-grained. Soctions were massive and unvesthered. Bistite was broken and altered to chlorite. Microcline was unaltered and unbroken. Very few microfractures were detected.

(b) <u>Granite</u> (Lucerne Pluton, Maine). Black and white. Coarsegrained perphyritic texture. Biotite was unaltered. Plagieclase wat slightly altered to sericite. Specimens were unweathered and contained very few microfractures.

(c) <u>Granite</u> (Granite Mountains Uplift, Wyoming). Unveathered, brownish-gray, coarse-grained. Microcline was unaltered. Plagioclase was altered to sericite. Biotite was slightly altered to chlorite. Microfractures were somewhat common.

(d) <u>Tomalite</u> (Sierra Nevada Batholith, California). Fine-grained, dark colored rock. Sections were fresh and contained no macrofractures. Contains principally plagioclase feldspar and biotite mica with smaller amounts of quartz and hormblends.

(e) <u>Tonalite</u> (Sierra Nevada Batholith, California). Medium- to coarsu-grained igneous rock. Sections were fresh and intact. Similar in composition to fine-grained rock discussed above.

(f) <u>Tenalite</u> (Sierra Nevada Bathelith, Califernia). Mediumgrained igneous rock; much finer grained than medium- to coarsegrained tenalite (e). Similur in mineral composition to the two tenalites discussed immediately before (d and e) except slightly more bietite and slightly less hornblende. Also contains very small amounts of magnetite.

(g) <u>Gramite</u> (Northwest of Lone Greve Fluton and Enchanted Nock Bathelith, Texas). kedium-grained, red gramite. Sections were intect and unveathered. Slight alteration of microcline and plagioclase. More muscovite mics present than biotite mics.

(b) <u>Gramite</u> (Shorman Gramite Facies of Southern Laramie Range, Wyeming). Coarse-grained, light-gray gramite. No preexisting fracture surfaces could be detected. Largely composed of quarts, petassium feldspar, plagioclase feldspar, and bietite, with lessor amounts of hormblende.

(i) <u>Gramite</u> (Laramie Range, Wyoming). Medium- to coarse-grained, pink gramite. Perphyritic texture. Sections were macroscopically free of fractures and were unvesthered. Predominately composed of quarts, plagioclass foldspar, petassium feldspar with lessor amounts of hermblende bistite and chlorite.

(j) <u>Tenalite</u> (Codar City Tenalite, Utah). Medium-grained, gray tenalite. Consisted primarily of plagioclase feldspar, quartz, and herablende with lesser amounts of petassium feldspar, bietite, and magnetite. Bietite was slightly altered to chlorite.

Specimen Preparation

Test specimens were prepared as suggested in the ASTM proposed "Standard Method of Test for Uncenfined Compressive Strength of Bock Core Specimens" and Corps of Engineers Standard Method of Test for Triaxial Strength of Undrained Rock Core Specimens (12), CRD-C 147. When propared according to the above specifications, specimen telerances were well within the limits required by the ASTM proposed "Standard Method of Test for Laboratory Deterministion of Ultrason's Palse Velecities and Klastic Constants of Rock. All samples were cut to lengths of approximately 4.32 inches with a Covington (Figure 1) slab saw (16-inch diameter diamond blade). This specimen length was selected in order to meet the specified length to diameter ratio requirements ($2.0 \le L/D \le 2.5$). Since specimen diameter ranged from 2.06 to 2.16 inches, probably due to variation in rock type, bit wear, and drilling technique, the actual length to diameter ratios also varied slightly, but were in all cases greater than 2.0 and less than 2.5.

During the cutting precess, specimens were secured in a vise (Figure 2) which aided in alignment and provided for cutting surfaces nearly perpendicular to the axis of the core. A solution of water and soluble oil was used as blade lubricant and coolant. All specimens were thorcughly washed immediately subsequent to cutting to remove any solution which might adhere to the specimen surface. Feed rate was adjusted such that one cut across a diameter required approximately 15 minutes.

After cutting, the ends of all specimens were ground smooth. parallel to each other, and perpendicular to the axis of the core with a Norton hydraulic surface grinder (Figure 3). This surface finishing was pursued in such a manner that specimen ends were flat to within 0 001 inches and did not depart from perpendicularity to the axis of the core by more than the allowable 0.01 inch in 2 inches (0.25 degrees). Subsequent to grinding, the specimens were again thoroughly washed to remove any of the oil-water grinding wheel coolant solution from the core surfaces.



Figure 1. Dismond Blade Slab Saw.





Figure 3. Hydraulic Surface Grinder.

Bulk Density

Bulk densities were determined according to U.S. Army Engineer Waterways Experiment Station, Concrete Division "T-2 Method of Determining Bulk Pensity of Bock Cores." The test procedure comsisted of:

(a) wash the core to remove dust and other coatings from the specimen

(b) air dry the specimen to constant weight, and weigh air-dried specimen to nearest 0.1 gram

(c) determine volume of specimen by liquid displacement in a pycnometer chamber (Figure 4) containing distilled water

(d) calculate the density of the core in the air-dried condition from the following formula:

$$G_{o} = \frac{W_{o}}{V_{o}}$$

where

G_a = density of the air-dried core

 W_0 = weight of the air-dried core in grams

V_a = volume of the core in cubic centimeters.

Temperature of the distilled water in the pycnometer chamber was taken into account when the volumes of the specimens were determined.

In this investigation, densities were computed from air-dried specimens rather than oven-dried speciment to avoid possible changes in physical properties due to oven-drying as have been observed in several previous studies. Obert (9), noted that oven-drying often produced prenounced and semetimes drastic changes in elastic constants, and that these changes were frequently permanent.



Figure 4. Pycaemeter Chamber.

<u>Ultrasonic Files</u> Velocities

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Ultramonic pulse velocities were determined according to the ASTM proposed "Standard Method of Test for Laboratory Determination of Ultrasonic Pulse Velocities and Elastic Constants of Rock." This method is valid for determination of compressional and shear wave velocities in both isotropic and anisotropic media.

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Barium titanate crystals and P2T-5A high capacitance leadzirconate-titanate crystals were used to produce compressional and shear pulses, respectively. These pulses were produced by applying short duration, high voltage pulses to the appropriate crystals. resulting in compressional or shear pulses, whichever the case may be, being generated in the specimen. The high voltage pulses, when applied to the X-cut barium titanate crystal, caused the crystal to expand and contract yielding compressional stress pulses, which were transmitted to one and of the specimen. when applied to the Y-cut lead-zirconate-titanate crystal, the pulses caused the shear crystal to vibrate in a direction perpendicular to the axis of the core creating shear pulses which were transmitted to one end of the specimen. The arrival of the pulses at the other end of the specimen were noted by a companion crystal affixed to that end, which acted as a mechanical-electrical transducer and generated equivalent electrical pulses. Pureline white petroleum jelly and pheryl salicylate were used respectively, between the compressional transducers and the rock specimens, and between the shear transducers and the rock specimens.

These electrical signals were recorded on an escillescope (Figure 5) as stationary wave forms, which allowed rather accurate



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measurement of the time of travel of the pulses through the individual specimens. Stationary wave forms, as displayed on a Hewlett-Packard model 1780-A oscilloscape, sere photographed (Figure 6) and pulse travel times read directly from the time serked photographs. These times were corrected to eliminate error due to pulse travel time through the transducer leads, transducers, and transducer-specimen connection materials (petroleum jelly or phenyl salicylate), thus, yielding pulse travel times through the rock core specimens above. Compressional and shear velocities were then determined from

$$V_p = \frac{L}{t_p}$$
 and $V_a = \frac{L}{t_a}$

where

V_p = compressional pulse velocity

V_a - shear pulse velocity

tp = travel time of the compressional pulse through the specimen alone

 t_{g} = travel time of the shear pulse through the specimen alone L = length of the specimen.

All compressional and shear pulse velocities determined in this investigation were measured with zero load on the specimens.

Static Axial Stress-Strain Measurements

To determine axial static stress-strain relations, Baldwin-Lima-Hamilton SR-4, Type A3-S-6, electrical resistance strain gages were affixed vertically to opposite sides of each specimen. The gages were located in a manner such that the midpoint of the



Shear Pulse



Compressional Pulse

Figure 6. Typical Photographs of Ultrasonic Wave-Ferms As Displayed on Hewlett-Packard Oscilloscope.

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resistance segment was at midheight of each specimen. The gage length was 13/16 inches, such that no portion would be effected by the nonuniform stress distributions noted by Fairhurst (5) to exist over the upper and lower 1/12 length of each specimen, i.e., the uppermost and lowermost 1/3 to 1/2 inches for all specimens used in this study.

All gages were bended directly to the rock specimens by using SR-4 cement, a fast drying mitro-collulose cement manufactured by Baldwin-Lima-Hamilton for the express purpose of application of SP-4 bonded strain gages. Prior to application of the cement, specimen surfaces were cleaned to remove substances such rs oil or dust which might impede development of a secure bond between the specimen and gage. A thin cost of cement was then applied both to the specimen and to the gage, after which the gage was mounted under moderate pressure and allowed to dry for 24 hours.

To determine stress-strain relations, the tr- gages were wired in series resulting in an output of the average strain registered by the two gages. Stress and strain were continuously plotted during the uniaxial compressive test by using a Meseley Autograf x-y recorder. A photograph of the recording equipment is given in Figure 7.

Uniazial Compressive Test

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All specimens were subjected to static compressive leading to determine axial stress-strair ~elations, as previously mentioned, and to determine ultimate uniaxial compressive strengths.


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Figure 7. Moseley Autograf, x-y Recorder and Other Stress-Strain Recording Equipment

A 440,000-pound universal, Baldwin hydraulic testing machine was used for leading the specimens. A photograph of this machine is given as Figure 8.

This test was conducted according to the ASTM proposed Standard Method of Test for Uncenfined Compressive Strength of Rock Core Specimens. All specimens were carefully aligned so that the axis of each core tested was coincident with the center of the ust of the spherically seated bearing block. An initial seating lock of approximately 100-200 pounds was applied very slewly while the spherical seated bearing block was adjusted. All tests were conducted st a leading rate of 35 ± 15 pai per second (constant for a particular specimen) so that catestrophic failure occurred within 5 to 15 minutes of commencement of leading. As noted by the propose: standard, such a rate of lead should provide values of ultimese uniaxial compressive strength which are relatively free from the effects of rapid leading.

Ultimate uniaxial compressive strengths were calculated by dividing the maximum lead carried by the specimen during the test by the average initial cross-sectional area of the specimen $C \in tor$ wined as suggested in the ASTM proposed standard. All strengths were expressed to the nearest 10 psi.



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Figure 8. Baldwin 440,000-pound Universal Testing Machine

CHAPTER 3

PRESENTATION AND DISCUSSION OF RESULTS

General

للقالة فالمحمد فلاقتماد ستعمينه مستعملا فاستقاقتهم وفالته فالمراح والمركز فتتأون والوغي وليريك ومرعد مسمد والممتحد منته مقدسة بالمسترفين فالمتعمل والمراحية وال

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Physical properties of the various granite and tonalite specimens tested were determined according to the procedures discussed in Chapter 2 and are presented in tabular form in Tables 1 and 2.

The static values of Young's moduli are tangent moduli of elasticity and were computed at 50 percent of ultimate uniaxial compressive strength. Static Young's moduli were determined for a minimum of six specimens from each drill site represented (ten sites), thus exceeding the optimum number of four specimens and minimum number of three recommended by the Bureau of Mines (1) for adequate evaluation of this particular property within a representative group of specimens.

Ultramonic elastic constants sere computed from individual values of density, ultramonic compressional pulse velocity, and ultramonic shear pulse velocity which wers determined for each specimen by procedures as discussed in Chapter 2. The equations used in the computation of ultramonic elastic constants are as follows:

(1)
$$E_{dyn} = \frac{2V_0^2(3V_p^2 - 4V_a^2)}{V_p^2 - V_a^2}$$

TABLE 1

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Physical Property Test Results

1 5		Petrograp! Deerrintit	hic	Ul timate Vniazial Co.upressive Strangth	Static Young'a Modulua	Ultramonic Pulme ' foa	Veloci ties
α- ! ?	No.	*()		pei	pei x 10-6	Compressional	Sheer
		Tonalite (29.200	10.0	16,160	9,190
		Tonalite		24,400	***	15,240	8,830
		Tonalite		25, 200	**1	16,200	8,66 0
		Tonalite		23,000	4.G	16,240	6, 8 30
		Tonalite		27,400	10.0	16,160	9,120
		Tonalite		24,600	0°3	14,180	8,550
		Tonalite		29, 800	10.0	15,980	8,950
	8	Tonalite		29,600	10.0	16,090	9,080
	•	Tonalite	(•)	22.500	7.4	15,650	8,570
		Tonalite		21.700	7.8	15,400	8,050
		Tonalite	•	17.600	7.6	15,180	8,570
		Tonalite		20,400	6.9	12,900	7,670
		Tonalite		14.200	******	14,790	8,030
		Tunalite		20, 500	6.9	12,680	7,670
		Tonalite		15.400	6.1	14,830	8,090
~ *		Tonalite		17,500	6.9	17,860	9,350
õ	•	Tanalita	()	22,000	6.3	15,900	8,690
-		Tanalita (24.000	6 .0	14,590	8,590
	16	Tonalite		24.200	8.1	13, 540	8,260
	46	Tonalite	(J	24,300	7.8	13,430	8,070

*Corresponds to petrographic descriptions given in Chapter 2, pages 7-9. **Static Young's moduli not determined for these specimens.

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\$peciar i No.	Petrographic Description ()*	Ultimate Uniazial Compressive Strength Dei	Static Young's Modulus Bei x 10-6	Ultresonic Pulse Compressional	Velocities Shear
8f	[onelite (f)	03 (10)			
2 C			0 -3 0 e		010 0 0
				10, 24 0	6, VUU
		24,000	2.2	14,710	8,450
10	(1) 0111000j	24,000		13,510	7,760
1 d	Toualite (d)	36,400	11.4	80,140	10.670
24	Tonalite (d)	47,200	11.6	80,180	10.590
3d	fonalite (d)	80,800	9.4	19,050	0.000
4d	[onelite (d)	43, 200	11.6	80,790	10.700
3 d	Tonalite (d)	45,500	11.6	21,160	10,970
0 d	Tonalite (d)	43,900	11.0	20,510	10.680
7d	Tonalite (d)	46,700	11.0	20,480	10,650
11	Tonalite (1)	0,800	4	12,590	6.560
18	Tonelite (1)	18,900	6. 3	12,340	6,790
31	Tonalite (1)	15,100	4.7	12,230	6,650
41	Tonalite (i)	17,800	6.4	18,750	6,730
51	Touslite (1)	12,500	4.0	12,610	7,050
61	Tonalite (1)	19,000	5.6	13,840	1,080
74	Tonalite (i)	12,300	4.5	12,660	6,950
72	Tonalite (i)	12,900	4.4	13,080	6,890

*Corresponds to petregraphic descriptions given in Chapter 2, pages 7-9. **Static Young's moduli not determined for these specimens.

TABLE 1 (Cont'd)

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	Petrographic Decentation	Ultimate Uniazial Gempressive Stranth	Static Young's Modulus	Ultrasonic Pulse fps	Velocities
No.	*()	pel	pei z 10-6	Compressionel	Shear
21	Granita (c)	34.000	4.0	17,430	0, 200
	Grant te (c)	31,100	9 .3	16,360	8,750
	Grantte (c)	34,600	0.9	16,830	9,340
	Grantte (c)	32,400	4.0	16,130	9,560
	Granita (c)	33,900	0.0	19,340	9,480
) y y	Granite (c)	32,300	9. 8	18,100	9,530
20	Granite	33,900	0.0	17,430	9,200
Hc Hc	Granite (c)	85,600	10.0	18,120	0,400
1	Granite (g)	23.600	9.6	11,900	9,530
. 0		21.100	0.0	18,520	9,210
10 t		17,100	*	14,040	0 6 0 0
10 T		25. × 00	10.2	17,680	9,450
10 1 7 12		23,000	10.0	18,350	9,630
		19,000	10.3	17,800	9,720
		26.500	9.0	18, 270	9,580
- 6 0	Granite (g)	20,800	‡ 1	18,510	9,820
1	Granita (h)	18.400	#	17,870	9,220
40	the strength	21.900	9.3	17,790	9,110
1	Grantte (b)	23.100	4.0	19,070	9,640
4 7	Oranite (h)	22, 200	8.3	18,370	9,540

*Corresponds to petrographic descriptions given in Chapter 2, pages 7-9. **Static Young's moduli not determined for these specimens.

TABLA 1 (Cont'd)

1

Velocities Sheer	9,630	8,970 9,180	8,930	1,010	7,000	7, 290	7,316	7,490	6,970	е, 500	1,310	9,200	6 .660	8,500	8,670	9,340	8,570	9,070	8,670	
Ul trasonic Pulas fpe Compressional	18,530	17,370	16,380	10,950	11,300	11,700	10,940	10,480	9,870	10,070	11,100	17,160	16,780	16,880	17,090	16,480	17,180	18,130	17,010	
Static Young'a Medulus Dai z 10-6	Ŧ	6. G	4.0	5.0	6.9	6.0	с. .	6°0	5.4	5.7	5.4	7.6	8.3	7.4	7.6	6.3	7.8	7.6	7.4	4
Ul timate Uniazial Compressive Strength pei	88,000	22,600 24,300	23,400	15,300	16,800	14,700	16,900	14,700	15,000	17,600	14,900	13,100	15,000	13,000	13,500	13,400	13,100	11,600	12,400	
Petrographic Description	Granite (h)	Granite (b) Granite (b)	Granite (b)	Granite (b)	Greatte (b)	Grenite (h)	Granite (b)	Grarite (b)	Greatte (b)	Granite (b)	Granite (b)	Oranite (j)	Granite (j)	Granite (j)	Grenite (j)	Granite (j)	Granite (j)	Grenite (j)	Gramite (j)	
gpectaen No.	Sh	40	ч8	91	gb B	3b	4 Þ	3b	0 b	7b	A	1.1	2	3.5	4	5	6.1		6.1	

*Corresponds to petrographic descriptions given in Chapter 2, pages 7-9. **Static Young's moduli not determined for these specimens.

TADLE 2

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Computed Ultrasonic Properties

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	[]] traeon(c Vodu)š, net s	, 10 ⁻⁶	Ul trasonic Poisson's
No.	Young' a	Bulk	Shear	Ratio
	1.6	5.3	3.0	0.26
9	0.9	4.6	8.8	0.25
		5.6	8.8	0.29
	1.0	5.7	2.8	0.29
	1.5	5.4	3.0	0.27
	6.3	3.7	2.6	0.21
	1.3	5.3	8.9	0.27
	7.4	5.3	2.9	0.27
	7.1	5.5	9 . 8	0.29
đ	4 • 9	5.6	2.4	0.31
•	7.0	5.0	8.8	0.87
	4.0	3.3	8.8	0.23
	0.0	4.9	8.4	0.29
		3.2	a.a	0.88
	6.9	4.8	8.4	0.20
• 20	6.9	7.6	3.3	0.31
16	7.1	5.6	8.8	0.89
5	9.9	4.2	2.7	0.23
er er	0.0	3.4	8.5	0.80
	3.0	3.4	8.4	0.88
51	0.1	4.8	8.8	0.86
81	7.6	5.8	3.0	0.26
11	8°8	4.4	8.6	0.85
8	5.5	3.7	8.8	0.25

			•	Ul trasonic
Specimen No.	VL traen Young's	ic Modula, pei Bulk	x 10-0	Poisson's Ratio
14	11.2	0.0	4.3	0.31
	11.2	Z		0.31
34	0.0	0.0	9. E	0.32
44	11.7	10.8	4.5	0.32
P S	12.1	11.0	4.6	0.38
64	11.4	11.4	4.4	0.32
7d	11.6	10.4	4.4	0.31
11	4.0	3.6	1.5	0.31
ī	9	3.2	1.6	0.28
31	•		1.6	0.29
	4	3.6	1.6	0.31
51	4	3.3	1.6	0.27
61	4.6	3.9	1.8	0.30
11	4.4	3.6	1.7	0.20
81	4.4	3.8	1.7	0.31
1c	8.0	6.7	3.1	0.30
90	7.1	5.9	8.1	0.30
36	9.5	و. ع م	3.4	0.31
40	5.5 2	7.3	3.8	0.31
56	0	b.7	3.5	0.32
6c	н. Ф	1.3	0. D	0.31
1c	7.8	6 . 8	3.0	0.31
La	Cl . 33	7.4	3.1	0.38
1 e	8.3	7.0	3.6	0.30
	7.8	8.8	3.0	0.34
	٤.ð	7.0	3.3	0.30
4	۶.8 ۲	7.0	3.1	0.31

TABLE 2 (Cont'd)

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			·	Ul trast
Jpecimen Vo	Ul treson	ic Moduli, pei ; Buit	2 10 ⁻⁰ Sheer	Poissou' Retio
	TINOT	Tha		
<u>i</u>	8 .3	7.5	8°9	0.31
	9.8	8.7	3.3	0.29
	- 4	7.4	3.8	0.31
- M	8.8	1.5	3.4	0.30
-	0	7,6	3.1	0.32
	1 c		3.0	0.32
1 4	0,0		4.0	0.33
1 4	2	0.1	3.3	0.32
42	 	6.0	3.4	0.32
64	7.7	7.1	8.9	0.32
4	8.1	7.8	3.1	0.38
αh	7.8	8. 4	8.9	0.34
-		0	7.7	0.14
	- 0 F 1		a	0. 0
			1.9	0.16
				0.10
			0.9	0.05
		a -	1.7	0.01
42		1.6	1.5	0.14
9 20	9.4	1.6	1.9	0.11
-	5	6 . 6	8.1	0.30
	7.1	6.6	0	0.38
		39.60	9	0.33
5-4	1.1	0.0	0.1	0.33
	- - 	6 . 9	3.8	0.33
	7.1	1.8	8.7	0.33
7,5	6.0	٤.0	3.0	0.33
EJ	1.9	1.0	1	0.33

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(3)
$$\mathbb{K}_{dyn} = \frac{o(3V_p^2 - 4V_a^2)}{3}$$

(4) $u_{dyn} = \frac{V_p^2 - gV_a^2}{2(V_p^2 - V_a^2)}$

where

Edyn = ultrasonic Young's modulus of elasticity in psi
Gdyn = ultrasonic shear modulus or modulus of rigidity in psi
Xdyn = ultrasonic bulk modulus in psi
^vdyn = ultrasonic Poissop's ratio
^o = density in pound-second² per inch⁴
^vp = ultrasonic compressional pulse velocity in inches per second

 V_6 = ultrasonic shear pulse velocity in inches per second. The number of specimens for which these constants were determined in all cases exceeded both the minimum and optimum number of tests per group (3 and 6, respectively) recommended by the Bureau of Mines (1).

Typical modes of failure exhibited by the seconty-nine rock core specimens tested in uniaxial compression are illustrated in the phytographs in Figure 9. Explosive type failures yielding the fragments illustrated in Figure 9(a) were typical of failures in the stronger tenalites, i.e., those yielding ultimate uniaxial compressive strengths greater than 35,000 psi. The conical type failure surfaces illustrated in Figure 9(b) were typical of the nature of failure exhibited by the remainder of the specimens tested.

In order to evaluate the linear degree of association between the various pairs of physical properties of interest in this investigation, correlation coefficients were computed for each of the



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a. Typical Explasive Type Failure Surfaces



b. Typical Conical Type Failure Surfaces

Figure 9. Typical Failures of Bock Cores in Uniaxial Compression

groups of data. A least-squares line was also fitted to each data group, the intention being to give a visual representation of the degree of linear correlation between the particular physical properties of interest.

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The correlation coefficient (:) is a measure of the degree of interdependence between the particular variables under study, with a coefficient of 1.0 or -1.0 indicating a perfect association, i.e., a perfect straight line relation, and a coefficient of 0.0 indicating no relation whatseever, a completely random association. The algebraic sign of the correlation coefficient reflects only the slope of the relationship, i.e., whether there is a trend toward increase or decrease in magnitude of one variable accompanying an increase ir the magnitude of other variables.

Once the correlation coefficients for the various groups of data were determined, it remained to evaluate the significance of these correlations, i.e., determine the probability of getting a correlation coefficient as large as the one actually obtained from the statistical sample if, in actuality, no correlation of this nature existed in the universe from which the statistical sample was taken. In this study, the probability level used to determine the minimum magnitude of a significant correlation coefficient for a given statistical sample size was 0.995. Thus, there would be only a 0.5 percent probability of obtaining a correlation coefficient of a magnitude greater than or equal to the predetermined value from a statistical sample taken from a universe in which ne cerrelation of this nature actually existed. A table of critical values for correlation coefficients corresponding to this chosen probability level and the various statistical sample sizes used in this study is presented as Appendix II.

The actual mathematical determination of the individual correlation coefficients and equations for least-squares lines was, for each data group, performed on a General Electric 400 computer.

Discussion of Correlations

لمكون والمشادلة المكركة بالمتاركة المقارلين فالممد فلأمتلأ أنكشف مقاله ولأراك مستعمل الملام للملاية

The remaining portion of this chapter will consist of the presentation of scatter diagrams (data plots) representing the various physical property correlations attempted. For each pair of physical properties examined, four data groups were arranged and correlated. These groups were comprised as follows:

Group 1; All tonalite specimens tested from the Sierra Nevada Batholith, California.

Group 2: All tonalite specimens tested.

Group 3: All granite specimens tested.

Group 4: All specimens tested.

Thus, for each pair of physical properties discussed, there will be four scatter diagrams, one for each group.

The correlations and scatter diagrams resulting from analysis of the data yielded by specimens comprising Group 1 (Sierra Navada Batholith tonalites) were compared to those resulting from analysis of the data yielded by the specimens comprising Group 2 (all tonalites). The objective of this comparison was to evaluate the degree of scatter which might be attributed to variation in grain size alone within the particular rock type. Variation in percentage mineral composition and geologic history were kept to a practical minisum for the Group 1 data (all specimens come from three holes within very close proximity of each other) and, for the Group 2 data, were allowed to vary within the confines imposed by restriction to the one particular rock type (tonalite). Therefore, close similarity between the correlation coefficients and scatter diagrams yielded for the two groups of data would suggest, for the particular pair of variables under consideration, that variation in grain size was of primary importance in the determination of degree of acatter typical of the association between those two physical properties for that rock type. On the other hand, significant dissimilarity in nature of the physical property correlations and correlation coefficients (i.e., a larger amount of scatter and lesser degree of linear association for the Group 2 data than for the Group 1 data) would indicate, for the particular pair of physical properties being examined, that variation in grain size, as opposed to variation in percentage mineral composition and geologic history as confined to the limits imposed by the particular rock type, was of lesser significance in the determination of the degree of scatter characteristic of the particular rock property correlation in question.

The correlations and scatter diagrams resulting from analysis of the data yielded by specimens comprising Groups 2 and 3 were compared to those resulting from analysis of the data yielded by specimens comprising Group 4 (all specimens in Groups 2 and 3 combined), the objective being to evaluate the effects of minimal variation in rock type on the nature of physical property correlations

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obtained for statistical data samples in olving more than one igneous rock type. As data Groups 2 and 3 represented individual rock types (tonalite and greatte, respectively) while data Group 4 represented both rock types, significant differences in the quality and orientation of the correlations determined for the three data groups should indicate the effects of slight variation in reck type on the nature of the resulting physical property correlations. Moreover, since the granites and tonalites involved in this investigation were relatively homogeneous and isotropic intrusive igneous rocks, and since the percentage mineral compositions of the two rock types are not extremely removed from one another (11), the rock property correlations determined for the data comprising Group 4 should indicate a reasonable maximum degree of linear association and minimum degree of scatter to be expected for the particular pairs of physical properties examined and for a statistical data sample which includes specimens representing more than one rock type.

 C_0 Veraus Ultrasonic Fulse Velocities (V_p and V_s). Scatter diagrams developed to illustrate the relationships between uniaxial compressive strength and iltrasonic compressional pulse velocity are presented in Figures 10 through 13. Figures 14 through 17 graphically illustrate the relationships found to exist between ultimate uniaxial compressive strength and ultrasonic shear pulse velocity as determined in this investigation.

The degrees of correlation characteristic of these two pairs of physical properties (C_0 versus V_p and C_0 versus V_B) were noticeably higher for the data yielded by the tonalite specimens alone then for the data yielded by the granite specimens alone. In particular, the



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Figure 11









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correlation coefficients (Table 3) (r) determined for the Group 2 tonalites were considerably larger than the corresponding critical values (given in Appendix II), while the correlation coefficients determined for the granites alone were only 0.44 and 0.49 for C_{\odot} versus $V_{\rm p}$ and C_{\odot} versus $V_{\rm g}$, respectively, low enough to be of questionable significance.

When the data yielded by the tonalites and granites were combined (for the particular pairs of physical properties) and examined as a single data sample, the result, in both instances, was a rather noticeable increase in scatter over and beyond the scatter typical of either the granite or tonalite data alone. Also, while the resulting correlation coefficients (for the Group 4 data) were slightly higher than those yielded for the Granite data alone, they were substantially lower than the values determined for the tonalite data alone. Thus the net effect, in these instances, appears to have been a considerable sacrifice of quality of the better scatter plots and physical property correlations (Group 2 tonalites) resulting from the introduction of data exhibiting trends of a somewhat different character and lesser quality (Group 3 granites).

A comparison of the correlation coefficients (Table 3) determined for the Groups 1 and 2 tonalites and examination of the scatter plots produced from data Fielded by these two groups of specimens revealed no significant differences in magnitude of the correlation coefficients or quality of the scatter diagrams for either C_0 versus V_p or C_0 versus V_8 . Thus, for these particular correlations, variation in grain size, rather than substantial variation in percentage

TABLE 3

Correlation Coefficients Obtained For Various Pairs of Physical Properties

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	Corre	lation Coeffic:	ients Obtained	Feri
Physical Properties Correlated	Sierra Nevada Batholith Tenalites (Group 1)	All Tonalites Tested (Group 2)	All Granites Tested (Group 3)	All Granites and Tomalites Tested (Group 4)
C va V	0.81	0.83	0.44	0.58
C va V	0.86	0.87	0.49	0.69
C vs E o dyn	0.86	0.67	0.49	0.71
C vs G dyn	0.87	0.85	0.50	0.73
C vs K o vs K	0.81	0.89	0.39	0.58
C ve dyn	0.49	0.20	0.33	0.24
E vs V	0.88	0.86	0.89	0.86
E vs V	0.92	0.92	0.91	0.92
E vs E dyn	0.92	0.90	0.90	0.90
E va G dyn	0.93	0.91	0.90	0.91
E vs K dyn	0.88	0.79	0.85	0.80
E vs udyn	0.58	0.11	0.80	0.45

mineral composition and geologic history as allowed within the confines of the specific rock type, appeared to be the primary factor contributing to the degree of scatter in the tonalite data plots.

 C_0 Versus Ultrasonic Moduli (E_{dyn} , G_{dyn} , and K_{dyn}). Scatter diagrams illustrating the general relationships existing between ultimate uniaxial compressive strength (C_0) and ultrasonic Young's modulus (E_{dyn}), ultrasonic shear modulus (G_{dyn}), and ultrasonic bulk modulus (K_{dyn}) are given in Figures 10 through 29. Correlations and data plots for comparable groups for each of the above three pairs of variables were generally similar, probably a reflection upon the similar origin of the various values of the three different ultrasonic moduli (all were computed from values of ultrasonic shear and compressional pulse velocities and specific gravities determined for the individual specimens).

Of particular interest was the fact that no appreciable changes in magnitudes of coefficients of linear correlation (as opposed to the magnitudes of those values determined for C_0 versus V_p and C_0 versus V_a) were effected by the use of ultrasonic elastic moduli in correlations with ultimate strength instead of the easier to determine pulse velocities (previously discussed). While correlations of ultimate uniaxial compressive strength (C_0) with ultrasonic shear modulus ($G_{\rm dyn}$) yielded the highest correlation coefficients of any cf the correlations in which ultimate strength was involved as a variable (0.87, 0 85, 0.50, and 0.73 for Groups 1, 2, 3, and 4, respectively), these coefficients were not sufficiently greater than those yielded by correlations of ultimate strength with ultrasonic



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Figure 19



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shear pulse velocity (0.86, 0.87, 0.49, and 0.89 for Groups 1, 2, 3, and 4, respectively) to alone justify the additional testing and computation necessary to the determination of ultrasonic shear moduli. This is not to say, however, that determination of ultrasonic elastic moduli would be undesirable.

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The same general trends prevalent in the correlations discussed previously (C_0 versus V_p , C_0 versus V_0) were also common to correlations of ultimate strength versus the various ultrasonic moduli. In particular, both groups of tonalite data yielded correlation coefficients of similar magnitudes (Table 3) and scatter diagrams of a similar nature. This apparently reinforces the previous indication that variation in grain size, rather than variation in mineral composition and geologic history as allowed within the confines of a specific rock type, appears to be of primary importance in the determination of the amount of scatter typical of correlations between various properties of this particular rock type.

Furthermore, examination of the correlation coefficients and scatter diagrams determined for the number 2, 3, and 4 data groups indicated, for all three physical property comparisons involving ultimate atrength and ultrasonic elastic moduli, that the correlation of data for both rock types as a single statistical sample was undesirable from the point of view that such a procedure resulted in an unnecessary sacrifice in quality of the scatter diagram and degree of correlation typical of the tonalite data alone without yielding an accompanying acceptable increase in quality of the scatter diagram and nature of correlation above those exhibited by the granite data alone. C_0 Versus v_{dyn} . Scatter diagrams illustrating the relationships observed to exist between ultimate uniaxial compressive strength and ultrasonic Poisson's ratio for the four data groups used in this investigation are given in Figures 30 through 33.

Correlation coefficients determined for these data groups (Table 3), were in general, very low, and in all cases but one, were less than the critical values (given in Appendix II), indicating no significant degree of linear association between the two variables. The coefficients were 0.49, 0.20, 0.33, and 0.24 for data groups 1, 2, 3, and 4, respectively.

Figures 30 through 33 further indicate the lack of linear association between the two variables under examination, and, in addition, reveal that the higher of the four correlation coefficients owe their larger magnitudes solely to the presence of a few points (eight) of questionable validity (unusually low values of ultrasonic Poisson's ratio) without which these coefficients would be even less significant.

B_{tan} Versus Ultrasonic Pulse Velocities (V_p and V_s). Figures 34 through 37 physically illustrate the relationships existing between values of tangent Young's modulus of elasticity (static) and values of ultrasonic compressional pulse velocity as determined in this study. Figures 35 through 41 depict the relationships existing between values of tangent Young's modulus and values of ultrasonic shear pulse velocity.

Interestingly, the eight correlations (four data groups) determined for these two particular pairs of variables were all quite







Figure 37

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Figure 13



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Figure 35

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Figure 19





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Figure 41

significant (correlation coefficients were such larger than the corresponding critical values as given in Appendix II) and very similar in nature, the quality and orientation of the relationships determined for the granite data alone being very nearly the same as those determined for the tonalite data alone, for comparable pairs of properties. This trend of similarity, which incidentally will be observed to exist generally throughout the correlations of tangent Young's modulus with the other various ultrasonic properties, was somewhat of a reversal of the earlier trend noted for correlations of ultimate uniaxial compressive strength with the various ultrasonic properties wherein the correlations for the granite relationships were frequently insignificant and noticeably inferior to those of the tonalite data. This "change" would appear to be a significant indication that the factors which contribute to ultimate uniaxial compressive strength characteristics and to the nature of the values of tangent Young's modulus of elasticity typical of a particular rock type do not necessarily contribute to the same properties of a similar rock type in the same manner.

A comparison of the scatter diagrams and correlation coefficients (Table 3) determined for the four groups of data representing each pair of physical properties resulted in two additional general observations: (1) correlations determined for the Group 1 and Group 2 tonalite data were similar both in orientation and degree of scatter, reinforcing the previous indication that variation in grain size within the particular rock type, rather than variation in mineral composition and geologic history as allowed within the confines of the specific rock type, was a primary influence upon the degree of scatter typical of the data plots for this rock type, and (2) amalgamation of the data for the tonalites and granites into a single group (Group 4) and correlation of this data as such yielded a scatter diagram and correlation coefficient which were of a nature and quality very similar to those determined for the parent data groups (Groups 2 and 3). This would appear to indicate, for the particular variables being examined, that such an axalgamation of data for such geologically similar rock types would not necessarily result in correlations of a quality substantially lower than the correlations determined for data grouped and correlated by individual rock type.

 E_{tan} Versus Ultrasonic Moduli (E_{dyn} , G_{dyn} , K_{dyn}). Scatter diagrams illustrating the general relationships found to exist between tangent Young's modulus of elasticity (E_{tan}) and ultrasonic values of Young's modulus of elasticity (E_{dyn}), shear modulus (G_{dyn}), and bulk modulus (K_{dyn}) are given in Figures 42 through 45. 46 through 49. and 50 through 53, respectively.

Nature and degree of linear correlations obtained for corresponding data groups yielded for all three pairs of variables were quite similar, probably due, as was the case with ultimate uniaxial compressive strength, to similarity in origin of the various values of the three ultrasonic elastic moduli ($E_{\rm dyn}$, $G_{\rm dyn}$, and $K_{\rm dyn}$). All were computed from equations involving values of ultrasonic shear and compressional pulse velocities and specific gravities of the individual specimens.

The physical property correlations determined for those pairs of variables involving ultrasonic moduli were generally quite good.



Static Young's Modulus ($E_{tan'}$, pai x 10⁻⁹

Figure 41



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Figure 44

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Correlation coefficients determined for two of these pairs (E_{tan} versus E_{dyn} and E_{tan} versus G_{dyn}) were among the highest yielded by correlation of any of the pairs of variables examined during the course of this investigation (Table 3). These two groups of correlations were not sufficiently superior to these involving the easier to determine values of shear and compressional pulse velocity, however, to alone justify the additional time and money which is most to say, however, that computation of ultraseries moduli. This is not to say, however, that computation of ultraseries moduli might not be desirable for purposes other than physical property correlations.

The general trends characteristic of the physical property correlations discussed in the section immediately prior to this one $(\mathbf{E}_{\text{tan}}$ versus V_p and V_g) were also noted to exist for these correlations. Specifically, degrees of correlation detormined for the Group 1 tonalite data groups were only slightly larger in megnitude than those yielded by the Group 2 tonalite date groups for corresponding pairs of physical properties, substantiating previous indications that variation in grain size, as opposed to significant variation in momental composition and geologic history as alleved within the confines of a specific rock type, appears to be of primary significance in the determination of the degree of scatter typical of various physical property correlations within the particular rock type.

As was mentioned in the previous section, combination of the data for both rock types, granite and tonalite, into a single statistical sample and correlation of this single mass of data did not, for either of the three poirs of physical properties considered in this section, result in an appreciable sacrifice in quality of the correlations obtained for the granite data alone or the tonalite data alone. Thus, for these pairs of physical properties and these rock types, variation in rock type appeared to have no significant effect on the quality of the correlations obtained.

 E_{tar} Versus v_{dyn} . Figures 54 through 57 graphically illustrate the general lack of linear association observed v_i exist between tangent Young's modulus of elasticity and ultrasonic Poisson's ratio.

Correlation coefficients determined for this pair of variables (four data groups) were, in most cases, appreciably larger than the unes obtained for the plots of ultimate uniaxial compressive strength versus PriceDu's ratio. Examination of the scatter plots (Figures 54 through 57), however, revealed no significant trends to exist between the two variables. The higher correlation coefficient (yielded by the Group 3 data) had resulted solely from the location of eight data peints, in this instance, points of questionable validity (values of Poisson's ratio were folt to be unrepresentative) in a manner so as to give definite orientation to the locat-squares line without increasing the quality of the plot to one of actual significance (See Figure 56).

Thus, as was the case with the correlation of ultimate uniaxial compressive strength with ultrasonic Pc'sson's ratio, the correlations determined for tangent Young's modulus of elasticity versus ultrasonic Poisson's ratio were all of little or no practical value from a property prediction point of view.

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CHAPTER 4

CONCLUSIONS AND RECOMMENDATIONS

Conclusions

Based upon results of the physical property tests conducted during this investigation, and on linear correlations of the various physical properties determined for these granites and tonalites, the following conclusions appear justified.

(a) Variation in grain size rather than substantial variation in percentage mineral composition and geolegic history as allowed within the confines of the specific rock type (tenalite) appears to be the primary factor responsible for the degree of scatter typical of data plots for the various pairs of physical rock properties examined for a specific intact rock type. It must be kept in mind, hewever, that the two rock types used in this investigation were both intrusive igneous rocks and were essentially of a homogeneous and isotropic nature. Thus, the above conclusion should be restricted basically to igneous rocks. It is felt that variation in geologic history would have substantial influence on the degree of scatter typical of rock property correlations for metamorphic and sedimentary rocks, particularly where there are vide variations in angles of inclination of planes of shistesity and sedimentation, and in degrees of cementation and recrystalization.

(b) Variation in rock type appears to have a generally undesirable influence on the quality of the correlation obtained for various pairs of rock properties. There correlations between two particular physical properties determined for various individual rock types are quite different in degree of linear association, analgamation of the smaller individual correlations into a single larger relationship inevitably results in a sacrifice of quality and usefullness of the good relationships due to introduction of data exhibiting lesser degrees of correlation. This type of situation was found to exist when the corresponding Group 2 and Group 3 correlations which involved ultimate uniaxial compressive strength as one variable were combined and evaluated as single groups (Group 4). A similar situation would result if the various physical property correlations for the individual rock types were of a differing nature (inclination or orientation) such that analgamation of the individual correlations for a single pair of properties into one large group would produce a large degree of scatter and thus a lesser degree of linear association. This would probably be more likely to occur when greatly different rock types were involved.

(c) A comparison of linear correlations involving ultrasonic pulse velocities with those involving ultrasonic elastic moduli indicated that linear relationships in which either ultrasonic shear moduli or ultrasonic Young's moduli were employed as one variable were slightly superior in quality to similar relationships in which one of the ultrasonic pulse velocities was involved. The degree of superiority did not appear great enough, however, to alone

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warrant the additional time and effort necessary to the determination of the ultrasonic elastic constants.

(d) Correlations which employ ultrasonic shear pulse velocity as one variable appear to be superior in quality to similar correlations involving ultrasonic compressional pulse velocity as one variable. Shear pulse velocity appears to be the ultrasonic physical property of those examined in chis study, which offers the best possibility for linear correlation with and proliminary prediction of ultimate unimial compressive strength and tangent (static) Young's modulus of elasticity.

(e) It does not appear likely that ultrasonic values of Peisson's ratio have any appreciable value in the area rock property correlation and prediction.

Recommendations

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As indicated previously, this investigation was confined to two types of igneous rocks. It is suggested that several varieties of metamorphic and sedimentary rocks be studied in a similar manner to determine whether or not the trends observed in this investigation are typical of all rocks or are merely typical of igneous rock types. In particular, the effects of variation in grain size on the degree of scatter typical of physical property correlations for metamorphic and sedimentary rocks should be investigated, as it is felt that variation geologic history may here be of primery importance rather than variation in grain size.

An investigation should be made to determine the extent to which physical property relations such as those investigated here might be better represented by curvilinear correlations rather than linear correlations. Several of the scatter diagrams in this study appeared to be of a nature that might be better represented by a second degree curve. In addition, a more extensive statistical analysis (determine confidence limits, etc.) might be performed to determine the practical value of such physical property correlations for property prediction purposes.

Ultrasonic shear pulse velocity and ultrasonic shear modulus should be more thoroughly examined for use in such physical property correlations. It would appear that these two ultrasonic properties offer the best persibilities of any of the ultrasonic properties used in this investigation.

ARSTRACT

Robert Wayne Crisp, Master of Science, 1971 Major: Civil Engineering

Title of Thesis: The Influence of Variation in Grain Size and Minimal Variation in Rock Type on the Quality of Rock Property Correlations for Intact Igneous Rocks

Directed by: Dr. Robert M. Scholtes, Head, Department of Civil Engineering

Pages in Ibesis: 116. Words in Abstract: 274

Correlations of physical properties of rock and predictions of one property from a previously determined value of another property should be of tremendous value in the field of civil engineering. In particular, such correlations and rock property predictions would expedite multiple site evaluation and selection programs, and possibly allow for reduction in the number of various tests required to determine the physical properties now deemed necessary for competent design and construction in rock modia.

Provious rock property correlations have generally encompassed many rock types, the objective being to determine general relationships typ'cal of all rock types. The data, however, have frequently exhibited such a great degree of scatter that subsequent correlations were of questionable value. Thus, in an effort to eliminate some of the scatter typical of many previous rock property correlations, this investigation was conducted to determine the inflaence of variation in grain size and minimal variation in rock type on the quality of rock property correlations for intact igneous rock types.

Physical property tests were conducted on 79 cylinderical specimens of granite and tonalite representing 10 drill sites. Values of ultimate uniaxial compressive strength and static Young's modulus of elasticity (tangent) were correlated with values of ultrasenic compressional pulse velocity, ultrasonic shear pulse velocity, ultrasonic Young's modulus, ultrasonic shear modulus, ultrasonic bulk modulus, and ultrasonic Poisson's ratio for each of the four fellowing groups of specimens:

Group 1: All tonalite specimens tested from the Sierra Nevada Satholith, California.

Group 2: All tenalite specimens tested.

Group 3: All granite specimens tested.

Group 4: All specimens tested.

Comparison of the nature and quality of these linear correlations revealed that variation in grain size, as opposed to variation in mineral composition and geologic history as allowed within the confines of a particular rock type, appears to have primary influence on the degree of scatter typical of rock property correlations for a particular intact igneous rock type. Moreover, rock property correlations involving only one igneous rock type are generally superior to those involving several igneous rock types, the

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latter case frequently suffering from an everell lack of quality due to a larger degree of scatter brought about by the analgumatier of data relationships exhibiting different trends and different degrees of these association.

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(a) Tenalite (Vermilien granite formation, Minnesota).
Brownish-gray medium- to coarse-gray tonalite. Biotite was broken and altered to chlorite. Mi.rocline was unaltered and unbroken.
Composed of 29% quarts, 48% plagioclase feldspar, 18% petassium feldspar (microcline), and 2% biotite with traces of magnitite, apatite, sphene, mircen, and calcite. Very few microfractures were detected.



(b) Granite (Lucerne Fluter, Maine). Black and white, cearsegrained granite. Porphyritic texture. Specimens were unweathered and contained very few microfractures. Contained 25% quartz, 30% plagieclase feldspar, 30% potassium feldspar (microcline), and 11% biotite with traces of magnitude, apatite, chlorite, epidete, and hematite. Plagieclase was slightly altered to sericite.



(c) Gramite (f amite Nountains Uplift, Fyoming). Unweathered brownish-gray, coarse-grained gramite. Microfractures were somewhat common. Contained 30% quarts, 30% plagioclase foldspar, 33% petassium foldspar (microcline), 5% biotite, 1% chlorite, and 1% magnetite and traces of epidete, apatite, and mircon. Amerihite content of the plagioclase was 15%. Plagioclase was slightly altered te mericite. Biotite was maintained to chlorite.

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(d) Tonalite (Sierra Nevada Batholith, California). Finegrained, dark colored rock. Sections were fresh and contained ne macrofractures. Contained 18% quartz, 42% plagioclase feldspar, 19% hornblende, 13% biotite, 4% chlorite and traces of microcline and other accessary minerals. The biotite was slightly altered to chlorite.



(e) Tenalite (Sierra Nevada Batholith, California). Nedium- te cearse-grained, black and white tenalit. Sections were fresh and intact. Percentage mineral compositions were 21% quartz, 45% plagieclase feldspar, 13% hernblende, 20% bietite, and 1% chlorite. Traces of microcline were also detected. The bietite was slightly altered to chlorite.

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(f) Tenalite (Sierra Nevada Bathelith, California). Mediumgrained, black and white tenalite; much finer grained than the medium- to coarse-grained tenalite (e). Sections were unweathered. Percentage mineral compositions were 19% quarts, 46% plagieclase feldspar, 2% microclime, 12% hernblende, and 21% bietite. Traces of chlorite and magnetite were also detected. The bietite had been slightly altered to chlorite. No macrofractures were detected.

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(g) Granite (Northwest of Lone Grove Pluton and Enchanted Nock Bathelith, Toxas). Medium-grained, red granite. Sections were intact and unweathered. More muscowite mica present than biotite mica. Percentage mineral composition is 30% quarts, 26% plagioclase feldspar, 53% petassium feldspar (microcline), and 9% biotite with traces of hornblende and chlorite.



(h) Granite (Sherman Granite Facies of Southern Laramie Range). Light-gray, coarse-grained granite. Composed of 24% quarts, 30% plagioclase foldspar, 32% petassium foldspar (microcline), 10% biotite, 4% hornblende, and a trace of chlorite. No proexisting fracture surfaces could be detected.



(i) Granite (Southern Laramie Range, Wyoming). Medium- to coarse-grained, pink granite, perphyritic texture. Percentage mineral composition is 20% quarts, 30% plagieclase feldspar, 34% petassium feldspar (microcline), 5% bietite, 1% hernblende, and 1% chlorite. Sections were unweathered and macroscopically free of fractures.



(j) Tenalite (Cedar City Tenalite, Utah). Light-gray, mediumgreined tenalite. Mineral composition is 20% quartz, 44% plagieclase feldapar, 3% potassium feldspar (micrecline), 21% hernblende, 5% bietite, 6% magnetite with traces of chlorite and other accessary minerals. Bietite was slightly altered to chlorite. Specimens were unweathered and macroscopically free of fractures.



(j) Tenalite (Cedar City Tenalite, Utah). Light-gray, mediumgrained tenalite. Mineral composition is 20% quartz, 44% plagieclase feldspar, 3% petassium feldspar (microcline), 21% hernblende, 5% bietite, 6% magnetite with traces of chlorite and other acceswary minerals. Biotite was slightly altered to chlorite. Specimens were unweathered and macroscopically free of fractures.

APPENDIX II

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CRITICAL VALUES

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CORRELATION COEFFICIENTS

Critical Values (r) of Correlation Coefficients

Statistical Sample Sige (n)	C. itical Value (rcr)
21	0.55
35	0.43
36	0.39
39	0.40
40	0.40
71	0.30
79	0.29

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IT SUPPLEMENTARY WOVES Report was also sub-	12. Proutonine		PITY
aitted to Mississippi State University,			
of Master of Science in Civil Engineering	-		
Previous rock property correlation	në hav- gene	rally enco	spassed many rock
Lypes, the objective being to determine gen- the data, however, have fremently avhibits:	eral relatio i much s are	nships tyr: af decree	ical of all rock type:
ment correlations were of mestionable value	ie. Thus, i	r al effor	t to eliminate some of
the scatter typical of many previous rock p	roperty corr	-lations,	this investigation was
in rock type on the mulity of rock property	lation in gr / correlatio	ain 312e a ne for ints	nd minimal Variation act innerns rock *vors
Physical property tests were conducted on 7	eylintrica	1 aveciment	s of granite and tona-
lite representing 10 drill sites. Values of static Voung's modulus of almaticity (tangen	Cultimate u atl ware nor	nindal co Martia di	mressive strength and
compressional pulse velocity, ultrasonic she	ar) were cor tor valze ve	locity. ali	trasonie Young's modu-
lus, ultrasonic shear modulus, ultrasonic bu	ilk zoi lus,	wij ultra	sonic Poisson's ratio
for each of the four following groups of spi from the Sierra Sevada Bathelith. Calif. A	'elmetr: 'al 1 all tonat	ul tonal: He comein	ite specimens tested -
granite specizens testei, and (4) all specis	wens trated.	Convaria	on of the nature and
mailty of these linear correlations reveale	el that vuri	willow in g	rain size, as opposed
se variation in mineral composition and good of a carticular rock type, appears to have a	ierir dition ministr infl	y se slijov uerce on ti	ea within the confined be Jegree of scatter
typical of rock property correlations for a	ver"i.uiar	intart ign	eous rock type. More
over, rock proverty correlations involving a	anly one light	ous reck	type are generally su-
ing from an overall lack of mulity due to	aa geer, 2 Marger ieg	an Lotter (ree of scal	usse irequently suffer ther brought about by
the meal remation of data relationships exhibit	biting diffe	rer* tren1	c and MIS" Frent de-
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Security Classification

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