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SIMULTANEOUS MEASUREMENT OF LASER EXTINCTION IN WARM FOG AT WAVELENGTHS OF 0.6328, 1.15 AND 10.6 MICRONS

Richard H. Munis, et al

Cold Regions Research and Engineering Laboratory

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Richard H. Munis and Allan J. Delaney

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PREFACE

This report was prepared by Dr. Richard H. Munis, Research Physicist, and Allan Delaney, Physical Science Technician, Physical Sciences Branch, Research Division, U.S. Army Cold Regions Research and Engineering Laboratory. The work was performed as part of the Arctic Surface Effect Vehicle program of the Advanced Research Projects Agency under ARPA Order 1615.

Technical review of the manuscript was performed by Dr. Yin-Chao Yen, Chief, Physical Sciences Branch.

ii

CONTENTS

	Page
Abstract	i
Preface	iii
Introduction	1
Experimental procedure	1
Experimental and theoretical equations	1
Results and discussion	3
Conclusions	6
Literature cited	7

ILLUSTRATIONS

TABLES

Table

1.	Experimental extinction coefficients of warm fog at -4°C	3
H.	Transmission values and corresponding optical depths at 0.6328 and 10.6 μ	6

SIMULTANEOUS MEASUREMENT OF LASER EXTINCTION IN WARM FOG AT WAVELENGTHS OF 0.6328, 1.15 and 10.6 MICRONS

by

R.H. Munis and A. Delaney

Introduction

A previous paper³ reported on the theoretical and experimental extinction coefficients in ice fog at wavelengths of 0.6328, 1.15 and 3.39 μ . These data were required by the Advanced Essearch Projects Agency (ARPA) for the design of an obstacle avoidance system for a surface effect vehicle (SEV). Also part of their requirements was the need to determine the extinction coefficients through warm fog at 0.6328, 1.15 and 10.6 μ . Since a direct comparison of the extinction coefficient at these wavelengths was desired, it was decided to set up an experiment in the fog chamber whereby all three lasers could be operated simultaneously. This report presents theoretical and experimental data on simultaneous laser extinction measurements through warm fog at 0.6328, 1.15 and 10.6 μ .

Experimental procedure

The simultaneous extinction measurements at 0.6328, 1.15 and 10.6 μ were made in a 4-m³ chamber whose temperature was maintained at -4°C (a detailed Jescription of this chamber is given in ref. 3). The three lasers were located so that, allowing for divergence of each beam at the detector, the distance between the 0.6328- and 1.15- μ beams was approximately one inch while the distance between these two beams and the 10.6- μ beam was approximately two inches. The three beams had to be near each other so that they would propagate through a volume of hydrometeors that could be sampled conveniently with the hand-operated impactor described in ref. 3. Thus the samples obtained with this instrument approximately described the hydrometeor size distributions which all three laser beams encountered simultaneously in the propagation path. These size distributions were then constructed using the technique discussed in ref. 3 to yield the so-called theoretical extinction coefficient. It should be pointed out that this extinction coefficient is not a true theoretically derived coefficient since the hydrometeor size distribution which is used in computing it is measured with the hand-operated impactor.

The procedure for measuring the transmission of the dissipating fog was identical with the technique discussed in ref. 3, except that in this experiment three laser beams were propagating simultaneously through the fog. This enabled three simultaneous equations to be written describing the nature of the theoretical extinction coefficients and their dependency on the hydrometeor size distribution.

Experimental and theoretical equations

Following the discussion in ref. 3, we write the general equation to calculate the experimental extinction coefficient:

MEASUREMENT OF LASER EXTINCTION IN WARM FOG

$$T[N(r), \lambda] = e^{-\alpha_{\lambda} |N(r), \lambda| \ell}$$

where $T[N(r), \lambda]$ = measured transmission, %

 α_{λ} = experimental extinction coefficient, m⁻¹

l = propagation path length, m

and N(r) describes the hydrometeor size distribution which exists in the fog chamber at the instant of time in which the transmission measurements are taken.

(1)

Three simultaneous equations may be written for the theoretical extinction coefficient:

$$\alpha_{0.6328} = \prod \sum_{r_{min}}^{r_{max}} N(r) r^2 \Delta r Q_{0.6238}(m, X)$$
⁽²⁾

$$\alpha_{1.15} = 11 \sum_{r_{\min}}^{r_{\max}} N(r) r^2 \Delta r Q_{1.15} (m, X)$$
(3)

$$\alpha_{10.6} = \prod \sum_{r_{min}}^{r_{max}} N(r) r^2 \Delta r Q_{10.6} (m, X)$$
(4)

where, as before,

 α_{λ} = theoretical extinction coefficient, m⁻¹ Q_{ext} = van de Hulst's efficiency factor for total extinction m = complex index of refraction X = particle size parameter, $2\pi r/\lambda$ and N(r) is as discussed above.

It should be noted from eq 2-4 that while the efficiency factor Q_{ext} is different at each value of λ the hydrometeor size distribution N(r) remains essentially identical because of the nature of the experiment. Figure 1 shows the behavior of Q_{ext} with particle radius for wavelengths of 0.6328, 1.15 and 10.6 μ . Note that the extinction efficiency of the complete particle spectrum is approximately



Figure 1. Q_{ext} vs particle radius at three wavelengths.

the same for 0.6328 and 1.15 μ , damping out eventually around a value of 2. However, the situation at 10.6 μ is quite different; the extinction efficiency is very low for small particles but grows considerably towards the larger end of the particle spectrum. From these graphs we could conclude that if the hydrometeor size spectra of a given experiment contained most of the particles in the range $0 < r < 10 \mu$ then the extinction coefficient at 10.6 μ would be expected to be considerably lower than at 0.6328 or 1.15 μ . The results of our experiment indicate that the above proposition is true for 0.6328 μ but not for 1.15 μ . More attention will be given to this matter in the following section.

Results and discussion

Table 1 compares the transmission and extinction coefficients at 0.6328, 1.15 and 10.6 μ for different hydrometeor concentrations. Figure 2 shows the hydrometeor spectra which resulted in the transmission values and extinction coefficients of Table 1.

Particle concentration (N/cm ³)	Experimental extinction coefficient (m ⁻¹)	Transmission (%)	Attenuation (db/m)
	a. 0.63	328 μ	
	0.1450	56	0.0629
1 30	0.0729	75	0.0312
125	0.0377	56	0.0164
68	0.0131	93	0.0079
41	0.0527	81	0.0229
35	0.0181	93	0.0079
227	0.5101	- 13	0.2214
116	0.0527	81	0.0229
258 .	0.6648	7	0.2885
184	C 2354	39	0.1022
	b. 1.	15 μ	
	0.749	5	0.325
1 30	0.802	9	0.261
125	0.337	26	0.146
68	0.168	51	0.073
41	0.492	14	0.213
35	0.104	66	0.045
227	0.978	2	0.425
116	0.379	22	0.164
258	0.877	3	0.381
184	0.805	4	0.349
	c. 10	.6 µ	
	0.1 362	58	0.0591
130	0.0753	74	0.0327
125	0.0291	89	0.0216
68	0.0051	98	0.0022
41	0.0496	82	0.0215
35	0.0181	93	0.0079
227	0.1941	46	0.0343
116	0.0291	89	0.0216
258	0.1783	49	0.0774
185	0.1733	50	0.0752

Table I. Experimental extinction coefficients of warm fog at $-4^{\circ}C$.

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Figure 2. Furticle concentration vs diameter at $-4^{\circ}C$.



Figure 3. Theoretical extinction coefficient at 10.6 µ versus theoretical coefficient at 0.6328 µ.

It is interesting to note that at 1.15 μ the values of transmission are considerably lower (consequently with significantly higher extinction coefficients) than at 0.6328 and 10.6 μ . As mentioned before, a comparison of the theoretical values of the extinction efficiency (Fig. 1) would seem to indicate that throughout the computed particle spectra the extinction coefficients at 0.6328 and 1.15 μ should have similar values. However, calculation of the theoretical extinction coefficient does not take into account the additional loss through water vapor which is present in the chamber during formation and dissipation of a fog. It is recognized that the wavelength of 1.15 μ appears in a wing of the 1.125- μ atmospheric H₂O vapor band. In a review of the optical properties of ice and water, Irvine and Pollock² indicate that the complex part of the index of refraction at 1.15 μ exceeds that at 0.6328 μ by approximately three orders of magnitude and that the absorption coefficient of water at 1.15 μ is about 400 times greater than that at 0.6328 μ . These values could therefore explain why the experimental extinction coefficients are so much higher at 1.15 μ than at 0.6328 μ . On the other hand experimental data obtained by Arnulf et al.¹ do not seem to corroborate the data obtained in this experiment.

Since the radiation at 1.15μ is adversely affected by atmospheric water vapor, it is of interest to examine the relationship between the extinction coefficients at 0.6328μ and 10.6μ . Figure 3 shows a plot of the theoretical extinction coefficients at 0.6328μ vs the coefficients at 10.6μ . These coefficients were calculated using eq 2 and 4 and the measured particle size distributions. The slope of the curve indicates that the extinction coefficient at 10.6μ should be somewhat smaller than that at 0.6328μ or, conversely, that transmission through the fog should be somewhat higher at 10.6μ . However, Figure 4 shows that the experimental results indicate that the extinction (or transmission) at 0.6328μ relative to that at 10.6μ becomes independent of particle concentration at approximately 200 cm⁻³. The slope of the linear portion of this curve indicates that the extinction coefficient at 0.6328μ is approximately equal to that at 10.6μ . Since the optical depth τ is quite large at both wavelengths with the particle concentration in the neighborhood of 200 cm⁻³ it is quite probable that the full effects of multiple scattering are dominating the scattering process. (Table II gives transmission values and optical depths τ at 0.6328 and 10.6μ .) Thus if only the linear portion of Figure 4 is taken into consideration the agreement with theory is not too bad.

MEASUREMENT OF LASER EXTINCTION IN WARM FOG

2 Amount

38 7 F.M.

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Figure 4. Experimental extinction coefficient at 10.6 µ versus experimental coefficient at 0.6328 µ.

Transmission (%)	Optical depth T	Transmission (%)	Optical depth t
58	0.544	S6	0.580
74	0.301	75	0.284
89	0.116	86	0.148
98	0.020	93	0.072
82	0.196	81	0.208
93	0.072	93	0.072
46	0.776	13	0.204
89	0.116	71	0.340
49	0.712	35	1.048
50	0.692	39	0.940

Table 11. Transmission values and corresponding optical depths at 0.6328 and 10.6 μ .

Conclusions

Experimental and theoretical data have been obtained on the simultaneous measurement of laser propagation through warm fog at 0.6328, 1.15 and 10.6 μ . It is theorized that due to high H₂O vapor concentrations in the chamber, propagation at 1.15 μ was severely reduced. This can be somewhat confirmed by the data in Figure 1, which indicate that for equal particle concentration the extinction coefficient should be approximately equal whether the particle spectrum peak is found at 7 or 12 μ . The 7- μ peak was measured for three concentrations during this experiment (warm fog) while the 12- μ peak was measured for approximately the same three concentrations (ice fog) and reported in ref. 3. The significant difference during these two experiments was that the ice fog propagation measurements were conducted at -43° C, which would tend to freeze out most of the water vapor, while the warm fog measurements were made at -4° C and hence a larger amount of H₂O vapor would be present in the fog chamber. This is illustrated dramatically when a comparison is made of the transmission values in Table 111 in ref. 3 and Table Ia in this report.

Theoretical calculations seem to indicate that the extinction coefficient at 10.6μ is somewhat smaller than that at 0.6328μ and thus should favor this wavelength if propagation through warm fog is the major concern of a design engineer. However, experimental data seem to contradict the theoretical calculations in that these data show virtually no difference between the extinction

7

coefficients at these two wavelengths for moderate fog concentrations, while for extremely large concentrations $\alpha_{10.6}$ assumes a constant value of approximately 0.2 while $\alpha_{0.6328}$ increases indefinitely.

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