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Atmospheric Attenuation of Laser Radiation From 0.76 to 31.25 μm

ROBERT A. McCLATCHEY JOHN E.A. SELBY

3 January 1974

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OPTICAL PHYSICS LABORATORY PROJECT 7670, 8603

L. G. HANSCOM FIELD, BEDFORD, MASSACHUSETTS 01730

AIR FORCE SYSTEMS COMMAND, USAF



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^{13.} ABSTRACT High resolution atmospheric transmittance curves are presented for the spectral region 320 to 13,200 cm ⁻¹ (0.7576 to 31.25 μ m). These spectra are useful as a guide for selecting laser wavelengths for atmospheric propagation studies in this spectral region. In addition, this report provides attenuation coefficients for those lines of the CO, HF, DF, and CO ₂ laser systems which suffer the least atmospheric attenuation. A new aerosol model is introduced here, taking into account recent measurements of the complex index of refrac- tion of aerosol particles.							
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Preface

This work was undertaken to provide extremely high resolution spectra as an aid to systems planning requiring a knowledge of laser propagation through the atmosphere. We have specifically addressed the problem of CO, HF, DF, and CO_2 laser systems and we have incorporated a new aerosol model.

We wish to acknowledge the Mie calculations performed by Dr. E. Shettle and the consultation with Dr. F. Volz and Dr. R. Fenn in the definition of the aerosol models described in this report. In addition, we acknowledge the efforts of Mr. J. Chetwynd in running computer programs and otherwise organizing the synthetic spectral plots.

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Atmospheric Attenuation of Laser Radiation from 0.76 to 31.25 μm

1. INTRODUCTION

Theoretical investigations of the attenuation of laser emission through the atmosphere require a knowledge of the molecular absorption of the atmosphere at very high spectral resolution. Absorption line widths of atmospheric molecules are typically of the order of 0.1 cm⁻¹ at one atmosphere pressure and decrease with decreasing pressure. Thus, considerations of laser propagation in the atmosphere require a knowledge of atmospheric transmittance with a spectral resolution of better than 0.1 cm⁻¹. In previous reports, calculations of synthetic atmospheric spectra were made for a spectral resolution of 0.01 cm⁻¹. The resulting spectra can thus be considered as representing an infinite resolution spectrum, limited only by the real width of the atmospheric absorption lines. One of the previous reports¹ provided spectra covering the region of CO emission - 1400 to 2120⁻¹. A second report² provided spectra covering the region of HF and DF emission from 2120 to 3740 cm⁻¹, and a third report³ provided spectra covering the region of CO₂ emission and beyond from 320 to 1400 cm⁻¹.

(Received for publication 3 January 1974)

- 1. McClatchey, R.A. (1970) <u>Atmospheric Attenuation of CO Laser Radiation</u>, AFCRL-71-0370, ERP 359.
- McClatchey, R.A. and Selby, J.E.A. (1972a) <u>Atmospheric Attenuation of HF</u> and DF Laser Radiation, AFCRL-72-0312, ERP 400.
- McClatchey, R.A. and Selby, J.E.A. (1972b) <u>Atmospheric Transmittance</u>, <u>7-30 μm: Attenuation of CO₂ Laser Radiation</u>, AFCRL-72-0611, ERP 419.

In addition to the "infinite" resolution spectra provided in these reports, specific laser attenuation charts have been provided for a great number of laser wavelengths in the CO, HF, DF and CO₂ systems. Although it is useful to have these laser attenuation coefficients immediately available, we have found the "infinite" resolution spectra of great value for a large number of purposes. For example, these spectra can be used directly as a guide to selecting other lasers which have lines that lie in the spectral interval in question.

Because of the growing interest in finding relatively transparent atmospheric windows for propagating new laser emission lines through the atmosphere, it was decided to extend the calculations reported earlier to shorter wavelengths and to provide in one report synthetic spectra for the entire spectral region from 320 to 13,200 cm⁻¹ (0.7576 to 31.25 μ m). The generation of accurate synthetic spectra requires a detailed knowledge of the spectroscopic parameters for each of the many thousands of molecular absorption features appearing in the infrared atmospheric spectrum. We are now in a position to perform these calculations due to the development of the AFCRL Compilation of Atmospheric Absorption Line Parameters described by McClatchey, et al.⁴

In addition to the absorption lines associated with water vapor, carbon dioxide, ozone, nitrous oxide, methane, carbon monoxide and oxygen, at low levels in the atmosphere there is the important water vapor continuum of particular importance in the 9- to 13- μ m region and between 16 μ m and 30 μ m.^{5,6} The pressure induced band at nitrogen in the region near 4.3 μ m has also been included.^{7,8} Absorption by each of the molecules mentioned here has been included in the calculation of synthetic spectra provided below.

For consistency with earlier reports on the problem of laser propagation in the atmosphere, synthetic spectra based only on molecular absorption have in all cases been provided for two different atmospheric paths: (1) A 10-km horizontal path at sea level, and (2) a 10-km horizontal path at an altitude of 12 km.

McClatchey, R.A., Benedict, W.S., Clough, S.A., Burch, D.E., Calfee, R.F., Fox, K., Rothman, L.S., and Garing, J.A. (1973) <u>AFCRL</u> <u>Atmospheric Absorption Line Parameters Compilation</u>, AFCRL-TR-73-0096.

^{5.} Burch, D. E. (1970) Semiannual Technical Report, <u>Investigation of the Absorp-</u> tion of Infrared Radiation by Atmospheric Gases U-4784, Jan. 1970.

^{6.} Bignell, K.J. (1970) Q.J.R.M.S., 96:409.

Burch, D. E., Gryvnak, D. A., and Pembrook, J. D. (1971) Philco-Ford Corporation, Aeronutronic Division, Contract No. F19628-69-C-0263, U-4897, ASTIA AD882876.

^{8.} Shapiro, M. M. and Gush, H. P. (1966) Canad, J. Phys. 44:949.

In addition to molecular absorption, three other sources of attenuation should be considered:⁹ molecular (or Rayleigh) scattering, aerosol scattering, and aerosol absorption. Quantitative data are also provided below on the basis of which aerosol attenuation can be estimated. It should be noted that aerosol attenuation and molecular scattering are very slowly varying functions of frequency and, therefore, provide a quasi-continuum attenuation over the whole spectral range of interest, whereas the molecular absorption is highly frequency-dependent. Thus, molecular absorption dominates in the determination of the relative "windows" where the transmittance of a laser beam is greatest.

2. ATMOSPHERIC MODELS

The atmospheric models used in the laser computations have been fully described,⁹ and so only a brief sketch will be provided here. Three model atmospheres for pressure, temperature, H_2O , and O_3 distributions have been used here and are referred to as Tropical, Midlatitude Winter, and Subarctic Winter. They refer to models of the same names defined in the Handbook of Geophysics and Space Environment.¹⁰ The major effect which these three different models have on the computations in this report is due to the differences in water vapor distribution. Table 1 provides the water vapor amounts in a 10-km sea level path, a 10km horizontal path at 12-km altitude, and in a vertical path through the entire atmosphere for the three models. The water vapor distribution in all models is identical above 11-km altitude. The ozone abundances have been included in Table 1 as ozone is the only other molecular species which is not assumed to be uniformly mixed in the atmosphere. All other absorbing gases were assumed uniformly mixed according to the mixing ratios indicated in Table 2.

In addition to the three models described above, computations were made for two aerosol models described as a "clear" and "hazy" atmosphere corresponding to a ground level visibility of 23 km and 5 km, respectively. The aerosol size distribution function for both models is the same at all altitudes and similar to one suggested by Deirmendjian¹¹ for continental haze. It differs from Deirmendjian's model "C" (and also from the model used by McClatchey et al⁷ in that the

- McClatchey, R.A., Fenn, R.W., Selby, J.E.A., Volz, F.E., and Garing, J.W. (1972) Optical Properties of the Atmosphere (Third Edition), AFCRL-72-0497, August 1972.
- 10. Valley, S.L., Ed., (1965) <u>Handbook of Geophysics and Space Environments</u>, AFCRL.
- 11. Deirmendjian, D. (1963) <u>Scattering and Polarization Properties of Polydis-persed Suspensions with Partial Absorption</u>, Proc. of the Interdisciplinary Conf. on Electromagnetic Scattering, Potsdam, NY, Milton Kerker, Ed., Pergamon Press.

Type at Path		Tropical	Midlatitude Winter	Subarctic Winter
10-km horizontal path at sea level	Н2О О3	$\begin{array}{r} 6.36 \times 10^{23} \\ 6.70 \times 10^{17} \end{array}$	$\begin{array}{c} 1.17 \times 10^{23} \\ 6.7 \times 10^{17} \end{array}$	$\begin{array}{c} 4.01 \times 10^{22} \\ 6.7 \times 10^{17} \end{array}$
10-km horizontal path at 12-km altitude	н ₂ О О ₃	2.00×10^{20} 5.40 x 10 ¹⁷	$\begin{array}{c} 2.00 \times 10^{20} \\ 3.23 \times 10^{18} \end{array}$	$\begin{array}{c} 2.00 \times 10^{20} \\ 5.4 \times 10^{18} \end{array}$
vertical path from sea level to space	н ₂ О О ₃	$\begin{array}{c} 1.38 \times 10^{23} \\ 6.62 \times 10^{18} \end{array}$	2.85 x 10^{22} 1.07 x 10^{19}	1.40 x 1022 1.29 x 10 ¹⁹

Table 1. Amount of Water Vapor and Ozone (molecules per square centimeter) in the Three Model Atmospheres for which Calculations Have Been Made

Table 2. Concentrations of Uniformly Mixed Gases

		Molecules/cm ²			
		Mi	dlatitude Winter M	odel	
Constituent	ppm by Volume	10 - km Sea Level	10-km Path at 12-km Altitude	Vertical Path From Sea Level	
N ₂	7.808×10^{5}	2.10 x 10^{25}	4.87×10^{24}	1.69×10^{25}	
02	2.095×10^{5}	5.63×10^{25}	$1.31 \times 10^{24^{-1}}$	4.52×10^{24}	
CO2	330	8.87×10^{21}	2.05×10^{21}	7.12 x 10^{21}	
CO	0.075	2.03×10^{18}	4.67 x 10^{17}	1.62×10^{18}	
N ₂ O	0.28	7.28×10^{18}	1.68×10^{18}	6.04×10^{18}	
СН ₄	1.6	4.30×10^{19}	9.92 x 10 ¹⁸	3.45×10^{19}	

large particle cut-off has been extended from 5 μ m to 100 μ m as indicated in Figure 1. The refractive index for the aerosol particles (Table 3) is based on experimental data published by Volz.¹² The attenuation coefficients were then determined as composites of 70 percent water soluble aerosol material and 30 percent dust-like substances which can be assumed representative of continental aerosol. The total numbers of aerosol particles per unit volume (Table 4) for the "clear" atmosphere have been adjusted to give an extinction coefficient at $\lambda =$ 0.55 μ m identical to the attenuation model of Elterman^{13, 14} at each altitude. The 12. Volz, F. E. (1972) <u>Appl. Opt. 11</u>:755.

- Elterman, L. (1968) <u>UV Visible, and IR Attenuation for Altitudes up to 50 km</u>, 1968, AFCRL, Environmental Res. Paper No. 285, AFCRL-68-0153.
- 14. Elterman, L. (1970) <u>Vertical-Attenuation Model with Eight Surface Meteoro-</u> logical Ranges 2 to 13 Kilometers, 1970, AFCRL, Environmental Research Paper No. 310, AFCRL-70-0200.

"clear" and "hazy" models are identical above 5 km. Below 5-km altitude, the number of aerosol particles in the "hazy" model increases exponentially to a value corresponding to a ground visibility of 5 km.



Figure 1. Aerosol Size Distribution Used in Computing Attenuation Coefficients

	Water Soluble	Dust - Liko
Wavelength	Refractive Index	Befractive Index
waverengtn	Terractive index	iten active index
.20000	1.530070*I	1.530070*I
.25000	1.530030*I	1.530030*I
.30000	1.530008*I	1.530008*I
.33710	1.530005*I	1.530008*I
.48800	1.530005*I	1.530008*I
.51450	1.530005*I	1.530008*I
.63280	1.530006*I	1.530008*I
.69430	1.530007*I	1.530008*I
.86000	1.520012*I	1.520008*I
1.06000	1.520017*I	1.520008*I
1.53600	1.510023*I	1.400008*I
2.00000	1.420 008*1	1.260008*I
2.50000	1.420012*I	1.180009*I
2.70000	1.400055*I	1.180013*I
3.00000	1.420022*I	1.160012*I
3.20000	1.430008*I	1.220010*I
3.39230	<u>1,430 - 007 *I</u>	<u>1.260 - 013*I</u>
3.50000	<u>1,450 - 005*I</u>	1.280011*I
3.75000	1.452 - 004*I	1,270011*I
4.00000	1.455005*I	1.260012*I
4.50000	1.460013*I	1.260014*I
5.50000	1.440018*I	1.220021*1
6.00000	1.410023*I	1,150037*I
6.50000	1.460033*I	1.130042*I
7.20000	1.400070*I	1.400055*I
7.90000	1.200065*I	1.150040*I
8.20000	1.010100*I	1.130074*I
8.50000	1.300215*I	1.300090*I
8.70000	2.400290*I	1.400100*I
9.00000	2.560370*I	1.700140*I
9.20000	2.200420*I	1.720150*I
9.50000	1.950160*I	1.730162*I
10.00000	1.820030*I	1.750 162*I
10.59100	1.760070*I	1.620120*I
11.00000	1.720050*I	1.620105*I
13.00000	1.620055*I	1.470100*I
14.80000	1.400100*I	1.570100*I
15.00000	1.420200*I	1.570100*I
17.20000	2.080240*I	1.630 0.100*I
18.50000	1.850170*I	1.648120*I
20.00000	2.120220*I	1.680220*I
25.00000		1.970248*I
27.90000	1.840290*1	1.890320*I
30.00000	1.820300*I	1.800420*I
35.00000	1.920400*I	1.900500*I
40.00000	1.860500*I	2.100600*I

Table 3. Aerosol Complex Index of Refraction (n-n'i): n = real (Scattering Part and n' = imaginary (absorption) Part

	PARTICLE DENSITY N (PARTICLES PER cm ³)			
Altitude	23-km Visibility	5-km Visibility		
(km)	Clear	Hazy		
$\begin{array}{c} 0\\ 1\\ 2\\ 3\\ 4\\ 5\\ 6\\ 7\\ 8\\ 9\\ 10\\ 11\\ 12\\ 13\\ 14\\ 15\\ 16\\ 17\\ 18\\ 19\\ 20\\ 21\\ \end{array}$	$\begin{array}{c} 2.828\pm +03\\ 1.244\pm +03\\ 5.371\pm +02\\ 2.256\pm +02\\ 1.192\pm +02\\ 8.987\pm +01\\ 6.337\pm +01\\ 5.890\pm +01\\ 6.069\pm +01\\ 5.818\pm +01\\ 5.675\pm +01\\ 5.317\pm +01\\ 5.585\pm +01\\ 5.156\pm +01\\ 4.744\pm +01\\ 4.511\pm +01\\ 4.511\pm +01\\ 4.513\pm +01\\ 3.634\pm +01\\ 2.667\pm +01\\ 1.933\pm +01\\ \end{array}$	1.378E+04 5.030E+03 1.844E+03 6.731E+02 2.453E+02 8.987E+01 6.337E+01 5.890E+01 5.818E+01 5.818E+01 5.585E+01 5.156E+01 5.048E+01 4.744E+01 4.511E+01 4.513E+01 3.634E+01 2.667E+01 1.933E+01		
22	1. $455 E+01$	1. $455 \pm +01$		
23	1. $113 E+01$	1. $113 \pm +01$		
24	8. $826 E+00$	8. $826 \pm +00$		
25	7. $429 E+00$	7. $429 \pm +00$		
30	2. $238 E+00$	2. $238 \pm +00$		
35	5. $890 E-01$	5. 890 ± -01		
40	1. $550 E-01$	1. 550 ± -01		
45	4. $082 E-02$	4. 082 ± -02		
50	1. $078 E-02$	1. 078 ± -02		
70	5. $550 E-05$	5. 550 ± -05		
100	1. $969 E-08$	1. 969 ± -08		

Table 4. Aerosol Models - Vertical Distributions for a "Clear" and "Hazy" Atmosphere

Through application of Mie scattering theory, attenuation coefficients were then extended to both longer and shorter wavelengths. The results of this extrapolation are contained in Figure 2 in which attenuation coefficients per kilometer are provided separately for absorption and total extinction (absorption plus scattering). The attenuation coefficients for molecular (Rayleigh) scattering is also given in Figure 2. The scale on the right hand side of Figure 2 is intended for use with some auxiliary curves provided in the report by McClatchey et al.⁹ The curve provided here is intended to be used as a replacement (containing more

recent information) for the curve contained in the earlier report. Using these attenuation coefficients, Figures 3a and 3b were constructed, providing the transmittance over a 10-km path at sea level and 12 km respectively, resulting from both the clear and hazy models. The transmittance due to Rayleigh scattering has also been included.



Figure 2. Attenuation Coefficients for Aerosol Transmittance (Absorption and Total Extinction)

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Figure 3a. Atmospheric Transmittance due to Aerosols and Rayleigh Scattering Through a 10-km Horizontal Path at Sea Level in a "Clear" and a "Hazy Atmosphere



Figure 3b. Atmospheric Transmittance due to Aerosols and Rayleigh Scattering Through a 10-km Horizontal Path at an Elevation of 12 km

3. COMPUTATIONAL TECHNIQUES FOR MOLECULAR ABSORPTION

A Lorentz line shape as given in Eq. (1) was assumed for each line.

$$k_{m} = \frac{S\alpha}{\pi [(\nu - \nu_{0})^{2} + \alpha^{2}]}$$
(1)

in which S is the line intensity, a is the half-width, ν_0 is the central line frequency, and ν is the laser frequency. For pressures less than 10 mb, a Voigt profile was used in the calculations.¹⁵ The laser frequency (ν) was assumed monochromatic for the purposes of this calculation. In general, a large number of absorption lines belonging to different molecules contribute to the attenuation at any specific laser frequency, so the total optical depth (O. D.) must be evaluated and is given by Eq. (2):

O.D. =
$$\sum_{j} \sum_{i} \frac{S_{ij} a_{ij} m_{j}}{\pi [(\nu - \nu_{ij})^{2} + a_{ij}^{2}]}$$
 (2)

where m_j represents the amount of the jth absorbing gas.

Pressure broadening enters through the a_{ij} values in Eq. (2). The Lorentz line width is given by $a = a_0 P/P_0 \sqrt{To/T}$. The monochromatic transmittance, τ_{ν} is thus given by

$$\tau_{\nu} = e^{-(O_{\bullet} D_{\bullet})}$$

The line intensity (S) is also temperature dependent through the population of the lower state of the transition and through the partition functions. These pressure and temperature effects have been included for all lines. The wings of all lines within \pm 20 cm⁻¹ of frequency, ν , were considered to contribute to the absorption coefficient at frequency ν .

In addition to this, absorption due to the water vapor continuum has been included based on the measurements of Burch et al⁷ and Bignell⁶ between 1250 and 320 cm⁻¹. Absorption due to the pressure-induced band of nitrogen was included in the 4- μ m region.^{7,8}

15. Young, C. (1965) J.Q.S.R.T. 5:549-552.

4. **RESULTS**

Figures 4 a through 4cl provide a high resolution (infinite resolution) transmittance spectrum for a 10-km horizontal path at sea level corresponding to the Midlatitude Winter model atmosphere. These curves cover the entire spectral region from 320 to 13,200 wavenumbers (0.76 to 31.25 μ m). Figures 5a through 5cl provide similar high resolution transmittance spectra for a 10-km horizontal path at a 12-km (approximately 40,000 ft) altitude for the same Midlatitude Winter model. The resulting curves in some portions of this spectral range were entirely opaque ($\tau_{\nu} \leq 10^{-3}$) and in portions were entirely transparent ($\tau_{\nu} \geq 0.999$). In these cases the spectra were omitted and are not included in Figures 4 or 5. However, the lettering sequence accounts for all plots whether or not they are included. This allows for an easy comparison between equivalent spectra at sea level (Figure 4) and at 12 km (Figure 5).

Table 5 indicates which curves have been omitted with the notation "opaque" or "transparent" as appropriate.

Figure No.	Spectral Range (cm ⁻¹)	,	Figure No.	Spectral Range (cm ⁻¹)	
4a	320-400	opaque	5ah	4220-4340	transparent
4b	400-560	opaque	5ai	4340-4460	transparent
4c	560-680	opaque	5aj	4460-4580	transparent
4j	✓1400 - 1520	opaque	5ak	4580-4700	transparent
4k	✓ 1520-1640	opaque	5au	5780-5900	transparent
41	✓ 1640-1760	opaque	5av	5900-6020	transparent
4m	✓ 1760-1880	opaque	5bn	8068-8180	transparent
4q	✓ 2240-2360	opaque	5bx	9260-9368	transparent
4ab	3560-3680	opaque	5by	9380-9500	transparent
4ac	3680-3740	opaque	5bz	9500-9620	transparent
4ad	3740-3860	opaque	5ca	9620-9740	transparent
4ap	5180-5300	opaque	5cb	9740-9860	transparent
4aq	5300-5420	opaque	5cc	9860-9980	transparent
4bf	7100-7220	opaque	5cd	9980-10040	transparent
4bg	7220-7340	opaque	5ci	11200-11500	transparent
			5cj	11500-11800	transparent
			5ck	11800-12100	transparent

Table 5.	Spectral	Plots Omitted	as	Being	Completely	Opaque	$(\tau_{\nu} \leq$	10	³) c	r
Franspare	ent $(\tau_{\nu} \geq$	0.999)		-			•			

In previous reports on laser propagation in the atmosphere, we have provided a large number of attenuation coefficient charts for specific laser lines of the CO, HF, DF, and $\rm CO_9$ systems. These charts provided attenuation coefficients as a function of altitude for several different atmospheric models. Our intent here is to provide the high spectral resolution curves described above and contained in Figures 4 and 5. However, during the last two or three years, some improvements in the molecular spectroscopic data have allowed us to make improved calculations for some of the laser wavelengths previously tabulated. In addition, interest has been indicated in the low vibration bands of CO and also in a number of additional DF lines. Consequently, we have compiled in Table 5 a large number of attenuation coefficients for laser emission lines belonging to these four molecular systems for which the attenuation coefficients per kilometer are the lowest. Although the laser frequencies are quoted to 0.001 cm^{-1} in Table 5, in most cases the probable accuracy is within +0.01 cm⁻¹ due to uncertainties in the molecular constants. Entries have been included in this table if the attenuation coefficients per kilometer for the Midlatitude Winter model are less than 0.25. In addition to these values, we have included attenuation coefficients per kilometer at sea level for the Tropical and the Subarctic Winter models and also for the Midlatitude Winter model at 12-km altitude. Table 6 contains attenuation coefficients for molecular absorption only. The effects of molecular (Rayleigh) scattering and of aerosol scattering and absorption would have to be added to these values if the total atmospheric attenuation is to be estimated. This can be accomplished by using Figure 2 as described above.

	CO LASER PARAMETERS			ATMOSPHERIC ABSORPTION COEFFICIENTS (km ⁻¹)					
				H = O km SEA LEVEL			H = 12 km		
	BAND	Rot. 1D	ν (cm ⁻¹)	k _{trop}	k mw	k sw	k mw		
*	1 - 0	P2 P14 P17 P21 P22 P25 P26 P27 P30	2135.549 2086.325 2073.267 2068.849 2055.402 2050.856 2037.027 2032.354 2027.651 2013.353	$\begin{array}{r}.661\\.409\\.608\\.268\\.141\\.152\\.441\\.178\\.757\\.548\end{array}$.249 .202 .159 .101 .0750 .0522 .0765 .0292 .137 .0784	.224 .176 .104 .0792 .0654 .0392 .0369 .0124 .0477 .0230	.266 .141 .0511 .0352 .0112 .00630 .00574 .000813 .000650 .000077		
*	2 - 1	P1 P2 P3 P4 P7 P8 P9 P11 P12 P15 P16 P17 P19 P21 P22 P25 P26 P27 P28	$\begin{array}{c} 2112.\ 977\\ 2109.\ 132\\ 2105.\ 256\\ 2101.\ 342\\ 2089.\ 393\\ 2085.\ 343\\ 2081.\ 258\\ 2072.\ 987\\ 2068.\ 802\\ 2056.\ 046\\ 2051.\ 729\\ 2047.\ 379\\ 2038.\ 582\\ 2029.\ 656\\ 2025.\ 145\\ 2011.\ 423\\ 2006.\ 786\\ 2002.\ 118\\ 1997.\ 419\\ \end{array}$.0935 .0525 .120 .122 1.52 .186 .151 .366 .240 .144 1.09 .350 .365 .213 .537 .407 .801 .320 .938	$\begin{array}{c} .0144\\ .0168\\ .0264\\ .0246\\ .191\\ .0346\\ .0276\\ .0733\\ .0761\\ .0218\\ .0846\\ .0718\\ .0542\\ .0314\\ .0746\\ .0577\\ .108\\ .0504\\ .157\end{array}$.00665 .0126 .0125 .0127 .0527 .0218 .0140 .0332 .0563 .0118 .0283 .0413 .0190 .00956 .0221 .0167 .0300 .0156 .0501	. 00035 . 00902 . 0038 . 0055 . 00671 . 00196 . 00196 . 00268 . 00427 . 000605 . 000769 . 00118 . 000178 . 000032 . 000079 . 000014 . 000020 . 000016 . 000045		
	3 - 2	P1 P2 P3 P5 P6 P7 P8 P10 P11 P12 P13	2086.594 2082.784 2078.940 2075.061 2071.148 2067.200 2063.218 2059.203 2051.071 2046.954 2042.804 2038.621	$\begin{array}{r} .479\\ .114\\ .630\\ .333\\ .123\\ .679\\ .801\\ .571\\ .414\\ .851\\ 1.49\\ .367\end{array}$	$\begin{array}{c} .0565\\ .0181\\ .171\\ .0558\\ .0235\\ .112\\ .130\\ .0937\\ .0581\\ .104\\ .225\\ .0525\\ \end{array}$.0263 .00920 .125 .0216 .0125 .0508 .0561 .0365 .0236 .0292 .0735 .0174	$\begin{array}{c} .00305\\ .00084\\ .045\\ .0064\\ .000861\\ .00181\\ .00152\\ .000655\\ .000598\\ .000119\\ .000429\\ .000122 \end{array}$		

Table 6. Attenuation Coefficients for Laser Frequencies

*Laser frequencies calculated using molecular constants of Young¹⁶.

16. Young, L.A. (1968) J. Quant, Spectrosc. Rad. Transfer 8:693.

	CO LASER PARAMETERS			ATMOSPHERIC ABSORPTION COEFFICIENTS (km ⁻¹)					
			<u> </u>	H - O	km SEA LE	EVEL	H = 12 km		
	BAND	Rot. ID	ν (cm ⁻¹)	^k trop	k _{mw}	k _{sw}	k mw		
꺄	3 - 2	P14	2034.405	. 882	.0896	.0217	.000239		
1	(Cont)	P15	2030.157	.317	.0406	.0116	.000073		
]		P16	2025.875	1.13	. 166	.0513	.000365		
		P17	2021.561	.734	.098	.0277	.000066		
		P19	2012.835	.739	. 102	.0290	.000077		
		P20	2008.424	1.68	. 231	.0654	.000044		
1		P21	2003.981	.299	.0410	.0127	.000117		
		P25	1985.891	1.06	.100	.0405	.000030		
		F20 1227	1901.290	.045	.0113	.0100	000011		
		P28	1971 995	607	.214	0290	000034		
		P30	1962.577	1.37	.216	.0660	.000058		
*	4 - 3	P2	2056.506	. 127	.0568	.0497	.00233		
		P3	2052.697	.0955	.0198	.0114	.000392		
		P4	2048.853	.283	.0616	.0406	.00151		
		P5	2044.975	.779	. 125	.0407	.000133		
			2037.110	. 308	.0802	.0305	.00110		
		· 120	2033,133	180	.0213	.00390			
		Г9 10	2025.121	503	0708	0214	000049		
		P11	2020.014	859	119	. 0338	000050		
		P13	2012.731	.581	.0816	.0234	.000022		
		P14	2008.550	1.43	.203	.0590	.000053		
		P15	2004.337	. 302	.0406	.0117	.000001		
		P17	1995.812	1.12	.170	.0513	.000039		
		P20	1982.783	.507	.0753	.0225	.000017		
		P21	1978.375	.281	.0446	.0141	.000048		
L		P22	1973.936	.386	.0607	.0187	.000016		
*	5 - 4	P2 De	2030.297	. 186	0.0236	.00682	.000011		
		F0 D7	2014.993	1.02	0.229	.0000	000120		
		P8	2011.002	1.02	0.225	0623	000219		
		P9	2003, 158	. 373	0.0502	.0144	.000018		
		P11	1995.100	1.61	.243	.0731	.000075		
		P14	1982.764	.496	.0730	.0217	.000017		
		P15	1978.586	.266	.0416	.0129	.000016		
		P16	1974.376	. 412	.0631	.0194	.000016		
		P21	1952.838	. 900	. 145	.0453	.000046		
		P25	1935.035	1.29	.205	.0681	.001500		
L		P26	1930.506	1.13	. 180	.0563	.000071		
幸	6 - 5	P2	2004.155	. 588	.0587	.0151	.000026		
		$\mathbf{P3}$	2000.415	.783	. 134	.0434	.000040		
		P4	1996.641	1.089	. 155	.0464	.00039		
			1985.115	. (38	. 108	.0319	.000024		
		гø	1901.203	1.00	• 118	.0207	.000013		

Table 6. A	ttenuation	Coefficients	for	Laser	Frequencies	(Cont)	l
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*Laser frequencies calculated using molecular constants of Young.¹⁶

	CO LAS	SER PARA	METERS	ATMOSPHERIC ABSORPTION COEFFICIENTS (km^{-1})				
				H · O k	m SEA LE	VEL	H 12 km	
	BAND	Rot. ID	ν (cm ⁻¹)	k _{trop}	k mw	k sw	k _{mw}	
*	6 - 5 (Cont)	P9 P10 P15 P19	1977.261 1973.284 1952.901 1936.007	. 437 . 432 . 917 1. 23	.0737 .0669 .147 .195	.0238 .0205 .0459 .0617	.000023 .000022 .000044 .000157	
a	7 - 6	P3 P4 P6 P7 P14	1974.409 1970.670 1963.089 1959.247 1931.380	.424 1.16 1.26 .969 1.36	.0641 .176 .195 .152 .212	.0196 .0529 .0594 .0469 .0653	.000016 .000042 .000052 .000048 .000106	

Table 6. Attenuation Coefficients for Laser Frequencies (Cont)

a Laser frequencies calculated using molecular constants of Mantz. ¹⁷
 * Laser frequencies calculated using molecular constants of Young. ¹⁶

17. Mantz, A.W., Nichols, E.R., Alpert, B.D. and Rao, K.N. (1970) J. Mol Spec. <u>35</u>:325.

	HF LAS	ER PARA	METER	ATMOSPHERIC ABSORPTION COEFFICIENTS (km^{-1})				
				H = O kr	n SEA LEV	EL	H = 12 km	
	BAND	Rot. ID	ν (cm ⁻¹)	k _{trop}	k mw	k sw	k mw	
b b	1 - 0	P11 P12	3436.12 3381.50	2.21 .496	.221 .0751	.0542 .0231	.0000287 .000022	
b	2 - 1	P8	3435.17	2.01	.209	.0512	.0000267	
b	3 - 2	P6	3373.46	.364	.0537	.0168	.000029	
с	4 - 3	P8 P9	3130.09 3083.83	.801 1.12	.148 .211	.0554 .0808	.000295 .000806	
с	5 - 4	P4	3150.67	.498	.126	.0736	.00229	
с	6 - 5	P6 P7 P8	2921.74 2880.70 2838.59	.586 .0430 .369	.0453 .00424 .0654	.0103 .00121 .0218	.000077 .000006 .000044	

Table 6.	Attenuation	Coefficients f	or Laser	Frequencies	(Cont)
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b Measured frequencies. 18
c Calculated frequencies. 19

18. Deutsch, T.F. (1968) Appl. Phys. Letters 10:234.

^{19.} Basov, N. G., Galochkin, V.T., Igoshin, V.I., Kulakov, L.V., Martin, E. P., Nikitin, A.I. and Oraevsky, A.N. (1971) Appl. Optics 10:1814.

	DF LAS	SER PARA	METER	ATMOSPHERIC ABSORPTION COEFFICIENTS					
				H = O k	m SEA LEV	ΈL.	H - 12 km		
	BAND	Rot. ID	ν (cm ⁻¹)	k _{trop}	k _{mw}	k _{sw}	^k mw		
d d	1 - 0	P1 P2 P3	2884.934 2862.652 2839.779	.414 .0540 0386	. 123 . 0115 . 007 25	.0772 .00485	.00316 .00316 .000038		
d d d		P4 P5 P6	2816.362 2792.437 2767.914	.0837 .0471 .0719	.0190 .0106 .0184	.0104 .00496 .00952	.00108 .000157 .000672		
d d d d d		P7 P8 P9 P10 P11 P12	2743.028 2717.536 2691.409 2665.20 2638.396 2611.125	.0352 .114 .0248 .0237 .337 .0133	.00801 .0204 .00485 .00752 .0664 .00394	.00352 .00718 .00252 .00489 .0247 .00302	.000043 .000034 .000053 .000307 .000187 .000090		
e c b b		P13 P14 P15 P16	2584.91 2557.09 2527.06 2498.02	.0145 .0176 .0145 .0261	.0102 .0180 .0155 .0282	.00981 .0185 .0161 .0295	.00390 .00335 .000565 .00103		
b b b b b b b b b b b b b b b b	2 - 1	P3 P4 P5 P6 P7 P8 P9 P10 P11 P12 P13 P16	2750.05 2727.38 2703.98 2680.28 2655.97 2631.09 2605.87 2580.16 2553.97 2527.47 2500.32 2417.27	.0401 .0378 .00528 .0600 .0535 .00950 .0311 .0282 .0144 .0140 .0240 .0811	$\begin{array}{c} . \ 00898 \\ . \ 00653 \\ . \ 00171 \\ . \ 0139 \\ . \ 0134 \\ . \ 00348 \\ . \ 00776 \\ . \ 0295 \\ . \ 0163 \\ . \ 0152 \\ . \ 0265 \\ . \ 0901 \end{array}$.00403 .00272 .00118 .00611 .00667 .00293 .00455 .0311 .0177 .0158 .0278 .0943	.000074 .000033 .0000307 .000069 .000733 .000761 .000110 .00180 .000883 .000554 .000972 .00330		
ხ ხ ხ ხ ხ ხ ხ ხ ხ ხ ხ ხ ხ ხ ხ ხ ხ ხ ხ	3 - 2	P3 P4 P5 P6 P7 P8 P9 P10 P11 P12 P13 P14	$\begin{array}{c} 2662. 17\\ 2640. 04\\ 2617. 41\\ 2594. 23\\ 2570. 51\\ 2546. 37\\ 2521. 81\\ 2496. 61\\ 2471. 34\\ 2445. 29\\ 2419. 02\\ 2392. 46\\ \end{array}$	$\begin{array}{c} .0354\\ .0437\\ .00490\\ .0118\\ .0507\\ .0322\\ .0150\\ .0319\\ .0509\\ .0659\\ .0797\\ .141 \end{array}$	$\begin{array}{c} . 007 \ 90 \\ . 009 \ 14 \\ . 00276 \\ . 00557 \\ . 0560 \\ . 0356 \\ . 0164 \\ . 0298 \\ . 0491 \\ . 07 \ 28 \\ . 0885 \\ . 119 \end{array}$	$\begin{array}{c} .00361 \\ .00424 \\ .00253 \\ .00480 \\ .0613 \\ .0379 \\ .0171 \\ .0307 \\ .0508 \\ .0756 \\ .0927 \\ .115 \end{array}$	$\begin{array}{c} .\ 000047 \\ .\ 000075 \\ .\ 000090 \\ .\ 000152 \\ .\ 00557 \\ .\ 00228 \\ .\ 00599 \\ .\ 00107 \\ .\ 00184 \\ .\ 00266 \\ .\ 00325 \\ .\ 00369 \end{array}$		
b b b b c c c	4 - 3	P5 P6 P7 P8 P9 P10	2532.502509.862486.832463.252439.292414.89	.0134 .0199 .0318 .0681 .0686 .0829	.0143 .0218 .0349 .0563 .0758 .0921	.0148 .0228 .0356 .0571 .0794 .0964	.000528 .000795 .00129 .00198 .00279 .00338		
c	5 - 4	P7	2404.63	.0878	.0965	.101	.00354		

Table 6. Attenuation Coefficients for Laser Frequencies (Cont)

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	DF LAS	SER PARA	METERS	ATMOSPHERIC ABSORPTION COEFFICIENTS (km ⁻¹)					
				H = O I	km SEA LE	EVEL	H = 12 km		
	BAND	Rot. ID	ν (cm ⁻¹)	^k trop	k mw	k sw	k _{mw}		
	7 - 6	P8 P10 P11 P12	2222.68 2177.99 2155.03 2131.68	.251 .123 .186 .272	.233 .0979 .0344 .187	.226 .0867 .0225 .195	.0102 .00297 .000846 .0311		
0 0 0 0 0 0	8 - 7	P7 P8 P9 P10 P12 P13	2165.93 2144.80 2123.24 2101.27 2056.14 2033.01	.0698 1.34 .187 .144 .114 .153	.0459 .129 .0296 .0322 .0222 .0198	.0466 .0349 .0169 .0180 .0131 .00580	.00258 .000357 .00410 .00599 .000494 .000100		
0 0 0 0 0 0	9 - 8	P6 P7 P8 P10 P11 P12	2108.48 2088.34 2067.76 2025.36 2003.56 1981.38	.0603 .444 .791 .646 .367 .476	.0172 .0567 .112 .0864 .0480 .0557	.0119 .0188 .0554 .0253 .0138 .0152	.00969 .00663 .00259 .000025 .000085 .000010		

Table 6. Attenuation Coefficients for Laser Frequencies (Cont)

b Measured, Deutsch. 18
c Calculated, Basov et al. 19
d Measured, Spanbauer et al. 20
e Calculated using Spanbauer et al. 20

20. Spanbauer, R. N., Rao, K. N. and Jones, L. H. (1965) J. Mol. Spec. 16:100.

CO ₂ LASER PARAMETERS	ATMOS	PHERIC AB	$SORPTION (km^{-1})$	COEFFICIENTS
	H + O k.	m SEA LEV	ЕΙ.	H 12 km
Rot. ID ν (cm ⁻¹)	k _{trop}	k _w	k _{sw}	k mw
1240 924.970	.514	0.0359	.0112	.000812
P38 927.004	.521	0.0423	.0154	.00164
P36 929.013	.744	0.0584	.0190	.00211
P34 930.997	.538	0.0536	.0227	.00311
P32 932.956	. 557	0.0650	.0302	.00520
P30 934.890	. 572	0.0737	.0360	.00677
P28 936.800	. 588	0.0852	.0440	.00887
P26 938.684	. 583	0.0853	.0447	.00955
P24 940.544	.603	0.0955	.0517	.0118
P22 942.380	.606	0.1021	.0569	.0136
P20 944.190	.609	0.0958	.0521	.0125
P18 945.976	.635	0.1223	.0717	.0186
P16 947.738	.572	0.0747	.0378	.00897
P14 949.476	.607	0.1101	.0642	.0173
P12 951.189	.591	0.1058	.0619	.0171
P10 952.877	.596	0.1008	.0580	.0161
P8 954.541	.553	0.0817	.0452	.0123
P6 956.181	. 513	0.0615	.0314	.00810
P4 957.797	. 484	0.0498	.0236	.00573
P2 959.388	.978	0.0753	.0282	.00609
R0 961.729	.456	0.0347	.0130	.00234
R2 963.260	.461	0.0401	.0170	.00367
R4 964.765	.478	0.0502	.0241	.00590
R6 966.247	.519	0.0614	.0308	.00783
R8 967.704	.505	0.0663	.0352	.00931
R10 969.136	.510	0.0714	.0389	.0104
R12 970.544	. 578	0.0788	.0418	.0109
R14 971.927	.556	0.0796	.0427	.0110
R16 973,285	. 554	0.0799	.0425	.0106
R18 974.618	. 522	0.0755	.0405	.0101
R2U. 975.927	. 194	0.2140	.0740	.0109
R22 977.210	. 6(4	0.08/1	.0398	.00803
R24 978.468	. 503	0.0641	.0318	.00099
	.484	0.0579	.0280	.00383
	• 4 (4	0.0529	.0240	.00471
	. 552	0.0587	.0240	.00378
152 983.248	.454	0.0430	.0183	.00324
	.400	0.0439	.0108	.00220
1,00 980.484 120 006.500	.4.30	0.0337	.0133	.001/0
1100 900,000 1240 007 £1£	,420	0.0020	0100	00133

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Table 6. Attenuation Coefficients for Laser Frequencies (Cont)
















































































Figure 4aa. Atmospheric Transmittance due to Molecular Absorption Through a 10-km Horizontal Path at Sea Level







































Figure 4an. Atmospheric Transmittance due to Molecular Absorption Through a 10-km Horizontal Path at Sea Level



















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Figure 4bn. Atmospheric Transmittance due to Molecular Absorption Through a 10-km Horizontal Path at Sea Level







































Figure 4by. Atmospheric Transmittance due to Molecular Absorption Through a 10-km Horizontal Path at Sea Level

































Figure 4cg. Atmospheric Transmittance due to Molecular Absorption Through a 10-km Horizontal Path at Sea Level



























Figure 5b. Atmospheric Transmittance due to Molecular Absorption Through a 10-km Horizontal Path at 12-km Altitude

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Figure 5f. Atmospheric Transmittance due to Molecular Absorption Through a 10-km Horizontal Path at 12-km Altitude



























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Figure 5n. Atmospheric Transmittance due to Molecular Absorption Through a 10-km Horizontal Path at 12-km Altitude

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Figure 5x. Atmospheric Transmittance due to Molecular Absorption Through a 10-km Horizontal Path at 12-km Altitude



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Figure 5ac. Atmospheric Transmittance due to Molecular Absorption Through a 10-km Horizontal Path at 12-km Altitude









































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Figure 5aw. Atmospheric Transmittance due to Molecular Absorption Through a 10-km Horizontal Path at 12-km Altitude





























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Figure 5br. Atmospheric Transmittance due to Molecular Absorption Through a 10-km Horizontal Path at 12-km Altitude

















Figure 5bw. Atmospheric Transmittance due to Molecular Absorption Through a 10-km Horizontal Path at 12-km Altitude























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