



20. (continued)

are also predicted by the model.

An additional feature is the optional inclusion in the vorticity equation of the terms which are usually considered negligible: (1) the advection of vorticity by the divergent wind, (2) the product of relative vorticity and divergence, (3) the vertical advection of vorticity, and (4) the twisting term.

The model can easily be adapted for different geographical areas with different grid lengths and also be integrated over different vertical layers. The lateral boundary values can be fixed or allowed to vary in time and thus it is possible to apply the so called nesting or telescoping technique. A special, relatively simple method to apply the nesting techniques is described.

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# A THREE-PARAMETER MODEL FOR LIMITED AREA FORECASTING

by

DR. L. BENGTSSON

MARCH 1974



ENVIRONMENTAL PREDICTION RESEARCH FACILITY  
NAVAL POSTGRADUATE SCHOOL  
MONTEREY, CALIFORNIA 93940

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## 1. INTRODUCTION

Numerical weather prediction with the aid of primitive equations has now been performed operationally for several years (Reiser, 1969; Shuman, 1968). Forecasts for one and two days by primitive equations indicate some improvement when compared with forecasts with the quasi-geostrophic and filtered equations. However, it has not been shown in a clear and convincing way whether this is solely due to the unfiltered part of the equations or to purely numerical improvements such as improved vertical resolution, an alternating grid, and the introduction of new physical effects such as sensible heat, latent heat, radiation, etc. When the forecasts with the primitive models are extended further in time, the improvements achieved with these models seem to be more obvious.

If a quasi-geostrophic model is compared with a primitive model with a resolution of three to four vertical layers or less, it will be found that the computational time (and cost) for the forecast with the primitive model is roughly about ten times as large as for the quasi-geostrophic model, if explicit time-integration schemes are used. The grid lengths now in use for operational weather prediction are still about 300 to 400 km.

Since significant weather disturbances have dimensions which are only three to five times that size, it is obvious that the truncation errors in the computation of the horizontal finite differences are much too large.

That grid distances of 300 km and more have been used for such a long time is naturally due to insufficient computational capacity, but may also partly be due to accustomed routine. However, a decrease in the horizontal grid length to half the size implies an increase in the number of grid points by a factor of four for the same computation area.

Due to the criterion of Courant, Friedrichs and Lewy, the time step must be decreased by a factor of two in order to maintain computational stability. If the computations can be organized in the same way as for the larger grid, a halving of the grid length, therefore, implies an eight-fold increase in the computation time. From the operational point of view, therefore, a primitive model should be compared with a quasi-geostrophic model where the horizontal grid length has been decreased to half the size.

When the forecasting areas are small and cover only one third or less of a hemisphere, the horizontal boundaries will fall in meteorologically active areas. This disadvantage does not create any large problem for the filtered models and a moderate horizontal smoothing in the neighborhood of the boundaries is sufficient. For the primitive models the

problem is much worse, since high-amplitude gravity waves are generated at the boundaries. These waves propagate into the area and greatly affect the meteorological information. Except for some successful experiments (Bushby, 1967, 1968; Gerrity, 1969), there has not been reported any adequate technique to avoid this in the general case. For this reason, forecasts for restricted areas with the complete equations in operational use imply considerable difficulties.

It may now be argued that it is of no use to apply a quasi-geostrophic model to a fine mesh, since that means the model will predict (or try to predict) scales of motion characterized by large Rossby numbers. For instance, when the Rossby number is on the order of one, all the terms in the vorticity equation are of equal magnitude. The quasi-geostrophic models will thus, according to this analysis, give rise to intolerably large errors for small and intense vortices, especially at low latitudes. However, it has been shown (Bengtsson and Moen, 1971) that substantial improvements in forecasts from quasi-geostrophic models are obtained if the grid size is reduced from 300 to 150 km.

The reason for this is that the higher order terms in the vorticity equation to a considerable degree cancel each other. It is only in the final stage of the cyclone development, when the flow becomes very deformed, that these terms become important.

It is also possible, as will be shown in this report, to include these terms in the integration and estimate them with the aid of quasi-geostrophic divergence.

It is the experience of the author that filtered models still are very useful for short-range predictions at medium and high latitudes. It is also very probable that unsuccessful predictions of especially rapid cyclogenesis are due to inaccuracies in the initial state and to unsatisfactory ways of including topographical effects, parameterization of dissipation, and the heating mechanisms. Recent comparisons performed in Sweden between filtered and primitive equation models support this view.

## 2. PROGNOSTIC EQUATIONS

The vorticity equation, thermodynamical equation and the continuity equation read:

$$\frac{\partial \zeta}{\partial t} = - \mathbf{V} \cdot \nabla \eta - f \quad D; \quad (2.1a)$$

$$\frac{\partial}{\partial t} \left( \frac{\partial \phi}{\partial p} \right) = - \mathbf{V} \cdot \nabla \left( \frac{\partial \phi}{\partial p} \right) - R(\Gamma_d - \Gamma) \frac{\omega}{p} - \frac{R}{c_p p} \left( \frac{\delta Q}{dt} \right); \quad (2.1b)$$

$$\frac{\partial \omega}{\partial p} = - D; \quad (2.1c)$$

where

$$\mathbf{V} = \mathbf{k} \times \nabla \phi,$$

$$D = \nabla \cdot \mathbf{V},$$

$$\Gamma_d = \frac{1}{\rho c_p}, \quad \text{and}$$

$$\Gamma = \frac{\partial T}{\partial p}.$$

We will now introduce some special model assumptions. We assume that the atmosphere is bounded by a pressure surface near the surface of the earth,  $p=p_0$ , and an upper pressure surface,  $p=p_1$ , near the tropopause. We will further assume a third interjacent pressure surface,  $p=p_m$ , which separates the atmosphere in two layers. We will now represent the wind field in the following way:

$$W = W_m - 2W_1 \frac{p-p_m}{p_0-p_m} ; \quad \text{layer 1}$$

$$W = W_m + 2W_2 \frac{p_m-p}{p_m-p_1} ; \quad \text{layer 2}$$

$$W = (W_m + 2W_2) \frac{p}{p_1} ; \quad \text{layer 3}$$

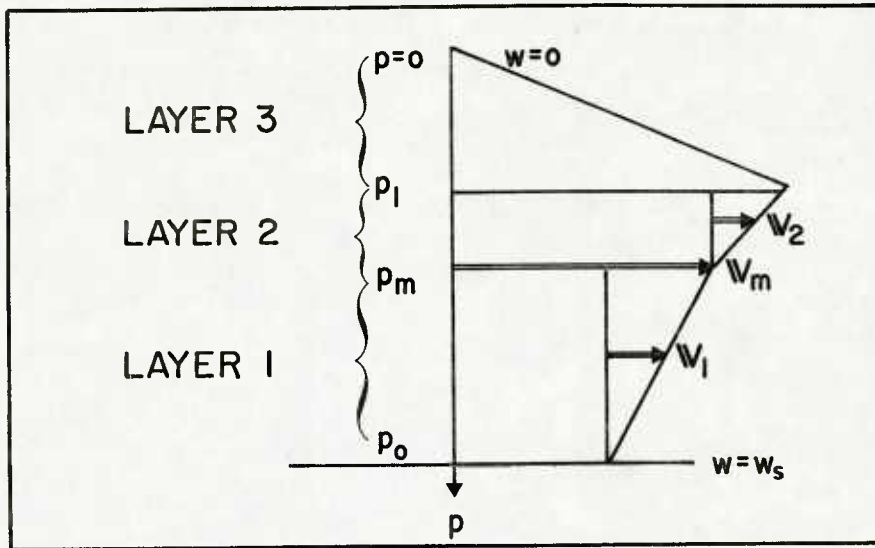


Figure 1. Vertical representation of the wind for a 3-parameter model.

According to the definition, equations for  $D$  and  $\zeta$  will have the same form as those for  $W$ . Equation (2.1a) can now be integrated between  $p=p_0$  and  $p=0$  with the boundary conditions  $\omega(p=0)=0$  and  $\omega(p_0)=\omega_s$ .

This gives a prognostic equation for the vertically integrated vorticity:

$$\begin{aligned} \frac{\partial}{\partial t} (\zeta_m + c_1 \zeta_2 - c_2 \zeta_1) = & -W_m \cdot \nabla (c_3 \eta_m - c_2 \zeta_1 + c_4 \zeta_2 + c_7 f) \\ & - W_1 \cdot \nabla (-c_2 \eta_m + c_5 \zeta_1) - W_2 \cdot \nabla (c_4 \eta_m + c_6 \zeta_2 + 2c_7 f) + c_8 f \omega_s \end{aligned} \quad (2.2)$$

where

$$\begin{aligned}
 c_1 &= \frac{2p_m}{2p_o - p_1} & c_4 &= \frac{6p_m - 2p_1}{6p_o - 3p_1} & c_7 &= \frac{p_1}{6p_o - 3p_1} \\
 c_2 &= \frac{2(p_o - p_m)}{2p_o - p_1} & c_5 &= \frac{8p_o - 8p_m}{6p_o - 3p_1} & c_8 &= \frac{2}{2p_o - p_1} \\
 c_3 &= \frac{6p_o - 4p_1}{6p_o - 3p_1} & c_6 &= \frac{8p_m}{6p_o - 3p_1} .
 \end{aligned}$$

The next two prognostic equations are computed from the difference between the vorticity equation at level  $p_m$  and  $p_o$  and the corresponding difference for  $p_1$  and  $p_m$ . We will now get two vorticity equations valid for layer 1 and layer 2:

$$\frac{\partial \zeta_1}{\partial t} + (W_m - W_1) \cdot \nabla \zeta_1 + W_1 \cdot \nabla (\eta_m - \zeta_1) + f D_1 = 0; \quad (2.3)$$

$$\frac{\partial \zeta_2}{\partial t} + (W_m + V_2) \cdot \nabla \zeta_2 + W_2 \cdot \nabla (\eta_m + \zeta_2) + f D_2 = 0. \quad (2.4)$$

From these two equations we will now eliminate the divergencies  $D_1$  and  $D_2$  with the aid of the continuity equation (2.1c) and the thermodynamical equation (2.1b). We will first integrate the continuity equation between the two levels  $p_a$  and  $p_b$ :

$$\omega(p_a) = \omega(p_b) + \int_{p_a}^{p_b} D \, dp. \quad (2.5)$$

We will now put  $p_a=p$ ,  $p_b=p_o$  and  $\omega(p_o)=\omega_s$  and will, thereby, get an expression for  $\omega(p)$  in layer 1:

$$\omega(p) = \omega_s + (p_o-p) D_m + \frac{(p-p_m)^2 - (p_o-p_m)^2}{p_o-p_m} D_1. \quad (2.6)$$

With  $p_a=p$ ,  $p_b=p_m$  and  $\omega(p_m)$  from (2.6) we get the following expression for  $\omega$  in layer 2:

$$\omega(p) = \omega_s + (p_o-p) D_m - (p_o-p_m) D_1 + \frac{(p_m-p)^2}{p_m-p_1} D_2. \quad (2.7)$$

With  $p_a=p$ ,  $p_b=0$  and  $\omega(p=0)=0$  we will have for layer 3:

$$\omega(p) = -\frac{p^2}{2p_1} (D_m + 2D_2). \quad (2.8)$$

Finally we get a relation between  $D_m$ ,  $D_1$ ,  $D_2$  and  $\omega_s$  with the aid of an integration of the continuity equation from the top to the bottom of the atmosphere:

$$D_m = c_2 D_1 - c_1 D_2 - c_8 \omega_s. \quad (2.9)$$

We now introduce the stream function into the thermodynamical equation and integrate through layers 1 and 2.

$(\Gamma_d - \Gamma)$  is assumed to be constant in every layer and  $\left(\frac{\delta Q}{dt}\right)_{pm} =$

$$\left(\frac{\delta Q}{dt}\right)_{p_1} = 0.$$



$$2f \frac{\partial \psi_1}{\partial t} = -2fW_m \cdot \nabla \psi_1 + R(\Gamma_d - \Gamma)_1 \int_{p_m}^{p_o} \frac{\omega}{p} dp + \frac{R}{2c_p} \ln\left(\frac{p_o}{p_m}\right) \left(\frac{\delta Q}{dt}\right)_{p_o};$$

$$2f \frac{\partial \psi_2}{\partial t} = -2fW_m \cdot \nabla \psi_2 + R(\Gamma_d - \Gamma)_2 \int_{p_1}^{p_m} \frac{\omega}{p} dp. \quad (2.10)$$

The integrals  $\int \frac{\omega}{p} dp$  can be computed with the aid of equations (2.6), 2.7) and (2.9), and the system (2.10) can be written

$$m_1 D_1 + m_2 D_2 = H_1;$$

$$n_1 D_1 + n_2 D_2 = H_2; \quad (2.11)$$

where

$$m_1 = \frac{R}{2} (\Gamma_d - \Gamma)_1 \left[ \frac{p_1 (p_o - p_m)^2 + p_m^2 (2p_o - p_1)}{(p_o - p_m) (2p_o - p_1)} \cdot \ln\left(\frac{p_o}{p_m}\right) - \frac{2(p_o - p_m)^2}{2p_o - p_1} + \frac{1}{2} (p_o + p_m) - 2p_m \right];$$

$$m_2 = \frac{R}{2} (\Gamma_d - \Gamma)_1 \left[ -\frac{2p_o p_m}{2p_o - p_1} \cdot \ln\left(\frac{p_o}{p_m}\right) + \frac{2p_m (p_o - p_m)}{2p_o - p_1} \right];$$

$$n_1 = \frac{R}{2} (\Gamma_d - \Gamma)_2 \left[ \frac{p_1 (p_o - p_m)}{2p_o - p_1} \ln\left(\frac{p_m}{p_1}\right) - \frac{2(p_o - p_m) (p_m - p_1)}{2p_o - p_1} \right];$$

$$n_2 = \frac{R}{2} (\Gamma_d - \Gamma)_2 \left[ \frac{p_1 p_m (2p_o - p_m)}{(2p_o - p_1)(p_m - p_1)} \ln \left( \frac{p_m}{p_1} \right) + \frac{2p_m (p_m - p_1)}{2p_o - p_1} + \frac{1}{2} (p_m + p_1) - 2p_m \right];$$

$$H_1 = f \frac{\partial \psi_1}{\partial t} + f \mathbf{V}_m \cdot \nabla \psi_1 - \frac{R}{4c_p} \ln \left( \frac{p_o}{p_m} \right) \cdot \left( \frac{\delta Q}{dt} \right)_{p_o} - \frac{R}{2} (\Gamma_d - \Gamma)_1 \left[ - \frac{p_1}{2p_o - p_1} \ln \left( \frac{p_o}{p_m} \right) + \frac{2(p_o - p_m)}{2p_o - p_1} \right] \omega_s ;$$

$$H_2 = f \frac{\partial \psi_2}{\partial t} + f \mathbf{V}_m \cdot \nabla \psi_2 - \frac{R}{2} (\Gamma_d - \Gamma)_2 \left[ - \frac{p_1}{2p_o - p_1} \ln \left( \frac{p_m}{p_1} \right) + \frac{2(p_m - p_1)}{2p_o - p_1} \right] \omega_s .$$

From the system (2.11) we can easily express the divergencies in an explicit way:

$$D_1 = -a_1 H_1 + a_2 H_2; \tag{2.12}$$

$$D_2 = b_1 H_1 - b_2 H_2;$$

where

$$a_1 = \frac{-n_2}{m_1 n_2 - n_1 m_2} ; \quad a_2 = \frac{-m_2}{m_1 n_2 - n_1 m_2} ;$$

$$b_1 = \frac{-n_1}{m_1 n_2 - n_1 m_2} ; \quad b_2 = \frac{-m_1}{m_1 n_2 - n_1 m_2} .$$

Introducing these expressions for  $D_1$  and  $D_2$  into the prognostic equations (2.3) and (2.4) and expressing the wind and vorticity in terms of the stream function gives:

$$\nabla^2 \frac{\partial \psi_1}{\partial t} - a_1 f^2 \frac{\partial \psi_1}{\partial t} + a_2 f^2 \frac{\partial \psi_2}{\partial t} = - J(\psi_m; \zeta_1) - J(\psi_1; n_m - 2\zeta_1) \\ + a_1 f^2 J(\psi_m; \psi_1) - a_2 f^2 J(\psi_m; \psi_2) - a_3 f \omega_S - a_1 f H \quad (2.13)$$

$$\nabla^2 \frac{\partial \psi_2}{\partial t} - b_2 f^2 \frac{\partial \psi_2}{\partial t} + b_1 f^2 \frac{\partial \psi_1}{\partial t} = - J(\psi_m; \zeta_2) - J(\psi_2; n_m + 2\zeta_2) \\ - b_1 f^2 J(\psi_m; \psi_1) + b_2 f^2 J(\psi_m; \psi_2) + b_3 f \omega_S + b_1 f H \quad (2.14)$$

Here we have

$$a_3 = a_1 s_1 - a_2 s_2;$$

$$b_3 = b_1 s_1 - b_2 s_2;$$

where

$$s_1 = \frac{R}{2} (\Gamma_d - \Gamma)_1 \left[ - \frac{p_1}{2p_o - p_1} \ln \left( \frac{p_o}{p_m} \right) + \frac{2(p_o - p_m)}{2p_o - p_1} \right]; \\ s_2 = \frac{R}{2} (\Gamma_d - \Gamma)_2 \left[ - \frac{p_1}{2p_o - p_1} \ln \left( \frac{p_m}{p_1} \right) + \frac{2(p_m - p_1)}{2p_o - p_1} \right]; \\ H = \frac{R}{4c_p} \ln \left( \frac{p_o}{p_m} \right) \left( \frac{\delta Q}{\delta t} \right)_{p_o} = h_1 \left( \frac{\delta Q}{\delta t} \right)_{p_o} \text{ with } h_1 = \frac{R}{4c_p} \ln \left( \frac{p_o}{p_m} \right).$$

We will also introduce the stream functions into equation (2.2) and define  $\psi_M$  by  $\psi_M = \psi_m - c_2 \psi_1 + c_1 \psi_2$ .

This results in the equation

$$\nabla^2 \frac{\partial \psi_M}{\partial t} = - J(\psi_m; c_3 n_m - c_2 \zeta_1 + c_4 \zeta_2 + c_7 f) - J(\psi_1; -c_2 n_m + c_5 \zeta_1) \\ - J(\psi_2; c_4 n_m + c_6 \zeta_2 + 2c_7 f) + c_8 f \omega_s. \quad (2.15)$$

The equations (2.13) through (2.15) now constitute our system of prognostic equations. The only thing which we now have to do is to find an expression for the non-adiabatic heat  $H$ , and the lower boundary condition,  $\omega_s$ .

### 3. BOUNDARY CONDITIONS

#### 3.1 LOWER BOUNDARY CONDITIONS

We now assume that  $\omega_s$  can be separated into two parts; one part which depends upon the dissipation in the boundary layer  $\omega'_s$  and one part which depends upon topography  $\omega''_s$ . We thereby assume:

$$\omega_s = \omega'_s + \omega''_s . \quad (3.1.1)$$

From the Ekman theory about the variation of the wind in the friction layer, we could easily derive the following expression:

$$\omega'_s = - g\rho_0 \sqrt{\frac{K}{2f}} \cdot F \cdot \zeta_0 \quad \text{where } F = 1 + c \cdot \sin\theta - c \cdot \cos\theta,$$

the wind in the surface layer has a magnitude  $c|V_0|$ , the angle between the geostrophic wind and the surface wind is given by  $\theta$ ,  $K$  is the turbulent coefficient of the viscosity, and  $\rho_0$  is the density of the air at the surface. Inserting the following numerical values:

$$K = 10 \text{ m/s}; \quad g = 9.81 \text{ m/s}^2; \quad f = 10^{-4} \text{ s}^{-1}; \quad T_0 = 280^\circ\text{K};$$

$$R = 287; \quad p_0 = 100 \text{ cb}; \quad \text{gives}$$

$$\omega'_s = - 2.729597 F \cdot \zeta_0 = -2.729597 \cdot F(\zeta_m - 2\zeta_1). \quad (3.1.2)$$

$F(c, \theta)$  can be given a constant value in the model or we can also assume different values over land and sea.

The following values will be used:

Over land  $\theta = 10^\circ$  and  $c = 0.78$  gives  $F = 0.36635$ ,

that is  $\omega'_s = -1. \zeta_0$ .

Over sea  $\theta = 5^\circ$ ;  $c = 0.85$  gives  $F = 0.22734$ ,

that is,  $\omega'_s = 0.62055 \zeta_0$ .

If  $k_2$  is assumed to be that part of the air which is forced over the mountains, we get

$$\omega''_s = k_2 \nabla_0 \cdot \nabla p_s = k_2 (\nabla_m - 2\nabla_1) \cdot \nabla p_s \quad (3.1.3)$$

$k_2 = 1$  will be used in the model.

The equations (3.1.1 - 3.1.3) now give the lower boundary condition for  $\omega_s$ . For further information see Bengtsson (1969). This way of treating the topographical effect as given by equation 3.1.3 is quite unrealistic and seems to underestimate the effect of the topography. (This will be especially true for steep and/or small scale mountains.) A new way to treat mountains as impenetrable vertical barriers has recently been published (Egger, 1972). A similar way to include mountains in vertically integrated balanced models will be described by the author in a coming investigation.

### 3.2 LATERAL BOUNDARY CONDITIONS

The model can use two different kinds of lateral boundary conditions namely:

a. Constant Inflow

$$\frac{\partial \psi_1}{\partial t} = \frac{\partial \psi_2}{\partial t} = \frac{\partial \psi_m}{\partial t} = 0$$

$$\zeta_1(t) = \zeta_1(t=0) \quad (= \text{constant in time})$$

$$\zeta_2(t) = \zeta_2(t=0) \quad (= \text{constant in time}) \quad (3.2.1)$$

$$\zeta_m(t) = \zeta_m(t=0) \quad (= \text{constant in time})$$

b. Variable Inflow

The values for  $\frac{\partial \psi_1}{\partial t}$ ,  $\frac{\partial \psi_2}{\partial t}$ ,  $\frac{\partial \psi_m}{\partial t}$ ,  $\zeta_1$ ,  $\zeta_2$ , and  $\zeta_m$  along the boundary are generated through an integration for a larger area which includes the actual area. Interpolation in time and space is necessary to synchronize the boundary values. A technical description of this is found in section 10.

#### 4. PARAMETERIZATION OF PHYSICAL PROCESSES

##### 4.1 SENSIBLE HEAT

In the computation of equation(2.10) we assume that  $\frac{\delta Q}{dt}$  decreased linearly to 0 from  $p_o$  to  $p_m$ . The sensible heat which is transported to the atmosphere from the underlying surface, is introduced in the following way:

$$\left(\frac{\delta Q}{dt}\right)_{p_o} = (A_1 |V_o| + A_2) (T_s - T_o) \quad \text{over sea, if } T_s > T_o; \quad (4.1.1)$$

$$\left(\frac{\delta Q}{dt}\right)_{p_o} = 0 \quad \text{over land and over ocean areas if } T_s \leq T_o;$$

where  $A_1$  and  $A_2$  are empirical constants  $A_2 = 10A_1$   
 $= 5 \cdot 10^{-3} \text{ m}(\text{sec})^{-2}(\text{deg})^{-1}$ .  $T_s$  is the sea surface temperature and  $T_o$  is the air temperature near the sea surface.

An approximative temperature at the level  $p_x = \frac{1}{2} (p_o + p_m)$  is obtained through an integration of the hydrostatic equation:

$$T_x = \frac{2f_o}{R \ln \left(\frac{p_o}{p_m}\right)} \psi_1. \quad (4.1.2)$$

We now assume that  $\Gamma$  is constant in the layer and is 75% of  $\Gamma_d$ . We therefore get

$$T_o = T_x + 0.75 \Gamma_d \frac{1}{2} (p_o - p_m) = \frac{2f_o}{R \ln \left(\frac{p_o}{p_m}\right)} \left[ 1 + \frac{0.75 R (p_o - p_m)}{c_p (p_o + p_m)} \right] \psi_1. \quad (4.1.3)$$



In equations (2.13) and (2.14) the sensible heating (H) was defined to be:

$$H = \frac{R}{4C_p} \ln \left( \frac{p_o}{p_m} \right) \left( \frac{\delta Q}{dt} \right)_{p_o} .$$

Substitution of  $\left( \frac{\delta Q}{dt} \right)_{p_o}$  from equation (4.1.1) and using the expression for  $T_o$  from equation (4.1.3) gives:

$$H = h_1 \cdot 10^{-2} (0.5 \cdot 10^{-1} |v_o| - 0.5) (T_s - h_2 \cdot \psi_1) \quad (4.1.4)$$

where

$$h_1 = \frac{R}{4c_p} \ln \left( \frac{p_o}{p_m} \right) ;$$

$$h_2 = \frac{2f_o}{R \ln \left( \frac{p_o}{p_m} \right)} \left[ 1 + \frac{0.75 R(p_o - p_m)}{c_p (p_o + p_m)} \right] .$$

## 4.2 PROGNOSTIC EQUATION FOR HUMIDITY

The specific humidity can be predicted by the following equation:

$$\frac{\partial q}{\partial t} = - \nabla \cdot \nabla q - \omega \frac{\partial q}{\partial p} + \epsilon - r + A \nabla^2 q . \quad (4.2.1)$$

Here  $\epsilon$  denotes evaporation,  $r$  condensation, and  $A$  is the coefficient of dissipation. We will disregard  $\epsilon$  and  $r$  will be introduced in a different way. Using the continuity equation we get the following prognostic equation:

$$\frac{\partial q}{\partial t} = - \nabla \cdot (q \nabla) - \frac{\partial}{\partial p} (q \omega) + A \nabla^2 q . \quad (4.2.2)$$

Since our model has a very low vertical resolution we have to parameterize the vertical distribution of humidity. We will, therefore, use precipitable water as the prognostic variable. The precipitable water is defined as

$$p_w = \int_0^{\infty} \rho_w dz = \frac{1}{g} \int_{p_T}^{p_0} q dp \quad (4.2.3)$$

where

$\rho_w$  = the density of water vapor and  $p_T$  is the pressure at the level over which we can disregard the humidity. (Here we have  $p_T = 30$  cb and  $p_0 = 100$  cb.)

A new quantity

$$w = \frac{1}{p_0 - p_T} p_w \quad (4.2.4)$$

which may be called normalized precipitable water is introduced. We assume that  $q(x,y,p,t) = gE(p)w(x,y,t)$  where  $E(p)$  describes the vertical variation of  $q$  computed from the standard atmosphere under the assumption that the relative humidity is 50%. The expression (4.2.4) is now introduced into (4.2.2) and we then integrate with respect to  $p$  from  $p_0$  to  $p_1$  to give

$$\frac{\partial w}{\partial t} = - \nabla \cdot (\tilde{V}w) - wd'\omega_s + A\nabla^2 w \quad (4.2.5)$$

where

$$\tilde{V} = \frac{1}{p_0 - p_T} \int_{p_T}^{p_0} V(p) E(p) dp \quad \text{and} \quad d' = \frac{E(p_0)}{p_0 - p_T} .$$

With  $\nabla = kx\nabla\psi + \nabla\chi$  equation (4.2.5) can be written

$$\frac{\partial w}{\partial t} = -J(\tilde{\psi}, w) - \nabla\tilde{\chi} \cdot \nabla w - w\nabla^2\tilde{\chi} - wd'\omega_s + A\nabla^2 w \quad (4.2.6)$$

where

$$\tilde{\psi} = e_m\psi_m + e_1\psi_1 + e_2\psi_2 \quad \text{and} \quad \nabla^2\tilde{\chi} = \tilde{D} = e_m D_m + e_1 D_1 + e_2 D_2.$$

#### 4.3 LATENT HEAT

There are two conditions for condensation:

- (a)  $\bar{\omega}_1 < \omega_{tol}$  (where  $\omega_{tol}$  is a given tolerance)
- (b) the relative humidity should exceed and be equal to 80%.

If these two conditions are valid, the latent heat is computed by the aid of expression (4.3.1)

$$H_{lat} = \frac{R}{4c_p} \ln \left( \frac{p_o}{p_m} \right) \left( \frac{\delta Q}{dt} \right)_{lat} = h_1 \left( \frac{\delta Q}{dt} \right)_{lat}$$

where

(4.3.1)

$$\left( \frac{\delta Q}{dt} \right)_{lat} = -L \cdot \bar{\omega}_1 F \quad \text{and} \quad F = \frac{1}{1 + \frac{L}{c_p} \left( \frac{\partial q^x}{\partial t} \right)_p} \left[ \left( \frac{\partial q^x}{\partial p} \right)_T + \frac{R}{c_p} \frac{T}{p} \left( \frac{\partial q^x}{\partial T} \right)_p \right]$$

$q^x$  is the maximum specific humidity. If the conditions (a) and (b) are not valid, we put  $H_{lat} = 0$ . Also see paragraph 8.4.

## 5. COMPUTATION OF THE VERTICAL MOTION

The integrated vertical motions in the two layers 1 and 2 are computed in the model. With the aid of the equation (2.6) and (2.9) we obtain:

$$\begin{aligned} \bar{\omega}_1 = \frac{1}{p_o - p_m} \int_{p_m}^{p_o} \omega dp = \frac{p_o + p_m - p_1}{2p_o - p_1} \omega_s + \frac{\frac{1}{3}(2p_1 - 3p_m - p_o)(p_o - p_m)}{2p_o - p_1} D_1 \\ - \frac{p_m(p_o - p_m)}{2p_o - p_1} D_2; \end{aligned} \quad (5.1)$$

$$\bar{\omega}_1 = t_1 \omega_s + t_2 D_1 - t_3 D_2.$$

In the same way we obtain from the equation (2.7) and (2.9):

$$\begin{aligned} \bar{\omega}_2 = \frac{1}{p_m - p_1} \int_{p_1}^{p_m} \omega dp = \frac{p_m}{2p_o - p_1} \omega_s - \frac{p_m(p_o - p_m)}{2p_o - p_1} D_1 + \\ + \frac{\frac{1}{3}(2p_o - p_1)(p_m - p_1) - p_m(2p_o - p_m - p_1)}{2p_o - p_1} D_2; \end{aligned} \quad (5.2)$$

$$\bar{\omega}_2 = t_4 \omega_s - t_3 D_1 + t_5 D_2 .$$

Here we have

$$t_1 = \frac{p_o + p_m - p_1}{2p_o - p_1}$$

$$t_2 = \frac{1}{3} \frac{(2p_1 - 3p_m - p_o)(p_o - p_m)}{2p_o - p_1}$$

$$t_3 = \frac{p_m(p_o - p_m)}{2p_o - p_1}$$

$$t_4 = \frac{p_m}{2p_o - p_1}$$

$$t_5 = \frac{\frac{1}{3}(2p_o - p_1)(p_m - p_1) - p_m(2p_o - p_m - p_1)}{(2p_o - p_1)}$$

The physical parameters are computed by the subroutine COEFF3P.

(See appendix B to this report)

## 6. NUMERICAL VALUES OF THE CONSTANTS

For the levels  $p_o = 100$  cb,  $p_m = 50$  cb,  $p_1 = 30$  cb and the stabilities  $(\Gamma_d - \Gamma)_1 = 0.422222$ ,  $(\Gamma_d - \Gamma)_2 = 0.511111$  we get the following numerical values for the constants:

$$\begin{array}{lll}
 c_1 = \frac{10}{17} = 0.588235 & m_1 = -826.330 & s_1 = 28.230856 \\
 c_2 = \frac{10}{17} = 0.588235 & m_2 = -688.40 & s_2 = 10.645832 \\
 c_3 = \frac{48}{51} = 0.941176 & n_1 = -532.92 & a_3 = 0.044374 \\
 c_4 = \frac{24}{51} = 0.470588 & n_2 = -10.58.36 & b_3 = 0.012258 \\
 c_5 = \frac{40}{51} = 0.784314 & a_1 = 2.0828 \cdot 10^{-3} & h_1 = 0.495351 \cdot 10^{-1} \\
 c_6 = \frac{40}{51} = 0.784314 & a_2 = 1.3546 \cdot 10^{-3} & h_2 = 0.110953 \cdot 10^{-5} \\
 & & h_3 = 1.03552 \cdot 10^{-6} \\
 & & h_4 = 0 \\
 & & h_5 = 0.492929 \cdot 10^{-1} \\
 & & h_6 = 1.03552 \cdot 10^{-6} \\
 c_7 = \frac{3}{51} = 0.588235 & b_1 = 1.0475 \cdot 10^{-3} & t_1 = 0.705882 \\
 c_8 = \frac{2}{170} = 0.0117647 & b_2 = 1.6261 \cdot 10^{-3} & t_2 = -18.627500 \\
 & & t_3 = 14.705900 \\
 & & t_4 = 0.294118 \\
 & & t_5 = -28.627500
 \end{array}$$

## 7. INTEGRATION OF THE COMPLETE VORTICITY EQUATION

### 7.1 GENERAL ASPECTS

Very little knowledge exists about the effect of the small order terms in the vorticity equation: the advection of vorticity by the divergent wind,  $\mathbb{W}\chi \cdot \nabla\zeta$ ; the product of relative vorticity and divergence,  $\zeta\nabla \cdot \mathbb{W}$ ; the vertical advection of vorticity,  $\omega \frac{\partial\zeta}{\partial p}$ ; and the twisting term  $\mathbb{k} \cdot \left(\frac{\partial\mathbb{W}}{\partial p} \times \nabla\omega\right)$ .

If these terms are used it is necessary, in order to conserve the total energy for an adiabatic model, to use the complete balance equation. If this is not the case, the model will not conserve total energy and after a certain time the development starts to deteriorate. This judgment has been mainly qualitative and we do not know the size of the error due to this inconsistency. It may be that this error is relatively small in comparison to other errors, as for instance, uncertainties of the initial state, and uncertainties in the description of the dissipation and the heating mechanisms. Therefore, it is necessary to perform a more detailed study of the problem and base our decision on a quantitative investigation.

It is by no means evident that we should use the same kind of assumptions in the formulation of models for short-range predictions (24 hours) as for models for medium- and long-range prediction. An example will illustrate this.

If one is interested in long-time integrations of the barotropic vorticity equation, it is necessary to use a finite difference expression which conserves kinetic energy, vorticity, and mean-squared vorticity. A finite difference expression which conserves these identities has been derived by Arakawa. However, if one is interested in short-range predictions, the so-called "Arakawa Jacobian" is not recommended since the phase-speed error is larger than the conventional finite difference analog to the Jacobian operator which only conserves vorticity. Experiments have shown that for forecasts up to four or five days it is not necessary to use an energy consistent Jacobian operator since the error in the kinetic energy is much smaller than errors due to other effects.

One of the problems which we have with the simplified vorticity equation is the over-prediction of anticyclogenesis. This seems due mainly to the lack of the term  $\zeta \nabla \cdot \mathbf{W}$  in the vorticity equation:

$$\frac{\partial \zeta}{\partial t} = \dots - (f + \zeta) \nabla \cdot \mathbf{W}. \quad (7.1.1)$$

In areas of convergence, relative vorticity is mostly positive, or will be after a short time. This means that the relative vorticity will increase faster in such areas if the complete expression (7.1.1) is included. On the other hand, in areas of divergence, the relative vorticity is mostly negative, or



will be after some hours. If the complete expression is used, the relative vorticity will decrease more slowly than if we use the simplified expression. It is easily seen that this term will create an asymmetry in the vorticity pattern which is also observed in reality.

Also the vertical advection of vorticity seems to play an important role, especially in cyclone development. During the development of the cyclone the activity is mainly concentrated in two different areas.

One area is in front of the warm front or, later in the development, the occluded part of the front. The other area is found below the upper-air low or trough, where a special center of activity is created in the later stages of the cyclone development. During the development of the cyclone, an area of sinking motion is concentrated under the upper-air low.

$$\frac{\partial \zeta}{\partial t} = -\omega \frac{\partial \zeta}{\partial p} \quad (7.1.2)$$

It is easily seen from the vorticity equation (7.1.2) that this effect will give an increase in the relative vorticity in areas of sinking motion and where the vorticity increases with height.

Upward motion over a surface low yields, in the same way, an increase in the vorticity for the levels above. Therefore the vertical advection of vorticity will increase the speed of occlusion. Figure 2 shows two different 12-hour

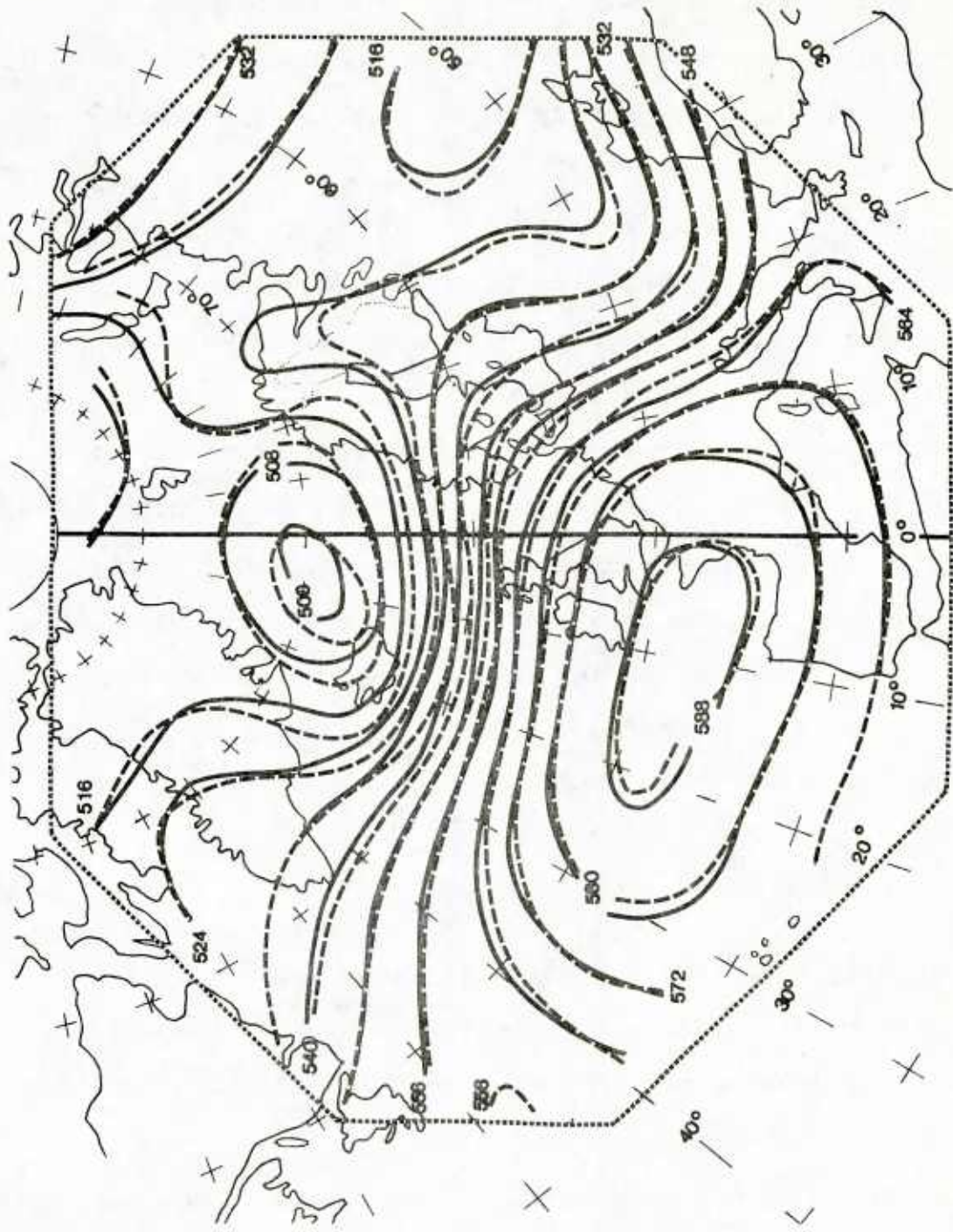


Figure 2. Two example 12-hour predictions with a 5-layer quasi-geostrophic model.

predictions with a 5-layer quasi-geostrophic model. Solid lines indicate a prediction performed with a quasi-geostrophic energy consistent model. Dashed lines indicate a prediction where the terms  $-(\zeta \nabla \cdot \mathbb{V} + \omega \frac{\partial \zeta}{\partial p})$  have been included.

## 7.2 DERIVATION OF THE FORECASTING EQUATIONS

The complete vorticity equation reads:

$$\frac{\partial \zeta}{\partial t} = - \mathbb{V}_\psi \cdot \nabla \eta - \underbrace{\mathbb{V}_\chi \cdot \nabla \eta}_1 - f \nabla \cdot \mathbb{V}_\chi - \underbrace{\zeta \nabla \cdot \mathbb{V}_\chi}_2 - \underbrace{\omega \frac{\partial \zeta}{\partial p}}_3 + \underbrace{|\mathbf{k} \cdot (\frac{\partial \mathbb{V}_\psi}{\partial p} \times \nabla \omega)}_4. \quad (7.2.1)$$

Here  $\mathbb{V}_\psi$  is the non-divergent wind and  $\mathbb{V}_\chi$  the divergent wind. The terms 1, 2, 3 and 4 are denoted non-geostrophic terms. Integrating through layer 1 by the representation of  $\mathbb{V}$ ,  $\zeta$  and  $D$  given in section 2 yields

$$\begin{aligned} (p_0 - p_m) \left[ \frac{\partial \zeta_m}{\partial t} - \frac{\partial \zeta_1}{\partial t} \right] = & (p_0 - p_m) \left[ - \mathbb{V}_m \cdot \nabla (\zeta_m - \zeta_1 + f) + \mathbb{V}_1 \cdot \nabla (\zeta_m + f) \right. \\ & \left. - \frac{4}{3} \mathbb{V}_1 \cdot \nabla \zeta_1 - f(D_m - D_1) - \zeta_m(D_m - D_1) + \zeta_1(D_m - \frac{4}{3}D_1) \right] \\ & + 2\zeta_1 \bar{\omega}_1 - 2|\mathbf{k} \cdot (\mathbb{V}_1 \times \nabla \bar{\omega}_1) \end{aligned} \quad (7.2.2)$$

Integration through layer 2 and layer 3 yields in a similar way

$$\begin{aligned} (p_m - p_1) \left[ \frac{\partial \zeta_m}{\partial t} + \frac{\partial \zeta_2}{\partial t} \right] = & (p_m - p_1) \left[ - \mathbb{V}_m \cdot \nabla (\zeta_m + \zeta_2 + f) - \mathbb{V}_2 \cdot \nabla (\zeta_m + f) \right. \\ & \left. - \frac{4}{3} \mathbb{V}_2 \cdot \nabla \zeta_2 - f(D_m + D_2) - \zeta_m(D_m + D_2) - \zeta_2(D_m + \frac{4}{3}D_2) \right] \\ & + 2\zeta_2 \bar{\omega}_2 - 2|\mathbf{k} \cdot (\mathbb{V}_2 \times \nabla \bar{\omega}_2) \end{aligned} \quad (7.2.3)$$

$$\begin{aligned}
\frac{1}{2}p_1 \left[ \frac{\partial \zeta_m}{\partial t} + 2 \frac{\partial \zeta_2}{\partial t} \right] = \frac{1}{2}p_1 \left[ - (V_m + 2V_2) \cdot \nabla f - \frac{2}{3} (V_m + 2V_2) \cdot \nabla (\zeta_m + 2\zeta_2) \right. \\
\left. - f(D_m + 2D_2) - \frac{2}{3} (\zeta_m + 2\zeta_2) (D_m + 2D_2) \right] - (\zeta_m + 2\zeta_2) \bar{\omega}_3 \\
+ |k \cdot \{ V_m + 2V_2 \} \times \nabla \bar{\omega}_3 \} \quad (7.2.4)
\end{aligned}$$

where  $\bar{\omega}_1$  and  $\bar{\omega}_2$  are given in (5.1) and (5.2) and  $\bar{\omega}_3$  is given by:

$$\bar{\omega}_3 = \frac{1}{p_1} \int_0^{p_1} \omega dp \quad (7.2.5)$$

If we now add (7.2.2), (7.2.3) and (7.2.4) and divide by  $\frac{2p_0 - p_1}{2}$ , we get a prognostic equation for the vertically integrated mean vorticity. From now on we will only keep the non-geostrophic terms on the right hand side of the equation.

$$\begin{aligned}
\frac{\partial}{\partial t} (\zeta_m + c_1 \zeta_2 - c_2 \zeta_1) = \{ - (V_\chi)_m \cdot \nabla (c_3 \eta_m - c_2 \zeta_1 + c_4 \zeta_2 + c_7 f) \\
- (V_\chi)_1 \cdot \nabla (-c_2 \zeta_m + c_5 \zeta_1) - (V_\chi)_2 \cdot \nabla (c_4 \eta_m + c_6 \zeta_2 + 2c_7 f) \}_1 \\
+ \{ c_8 \zeta_m \omega_s + c_2 \zeta_1 (D_m - \frac{4}{3} D_1) - \zeta_2 (c_4 D_m + c_6 D_2) + \zeta_m (c_7 D_m + 2c_7 D_2) \}_2 \\
+ c_8 \{ 2\zeta_1 \bar{\omega}_1 + 2\zeta_2 \bar{\omega}_2 - (\zeta_m + 2\zeta_2) \bar{\omega}_3 \}_3 + c_8 \{ |k \cdot [-2V_1 \times \nabla \bar{\omega}_1 - 2V_2 \times \nabla \bar{\omega}_2 \\
+ (V_m + 2V_2) \times \nabla \bar{\omega}_3] \}_4 \quad (7.2.6)
\end{aligned}$$

The two remaining prognostic equations are computed from the difference between the vorticity equation for  $p_m$  and  $p_0$  and the difference between the vorticity equation for  $p_1$  and  $p_m$ . We will then have two vorticity equations valid for layer 1 and layer 2 respectively. The geostrophic terms are omitted.

$$\begin{aligned}
\frac{\partial \zeta_1}{\partial t} = & - \{ ((\mathbb{V}_x)_m - (\mathbb{V}_x)_1) \cdot \nabla \zeta_1 + (\mathbb{V}_x)_1 \cdot \nabla (\eta_m - \zeta_1) \}_1 \\
& - \{ (\zeta_m - 2\zeta_1) D_1 + \zeta_1 D_m \}_2 + \left\{ \frac{\zeta_1}{p_0 - p_m} (\omega_m - \omega_s) \right\}_3 \\
& - \left\{ \frac{1}{(p_0 - p_m)} |\mathbf{k} \cdot [(\mathbb{V}_\psi)_1 \times \nabla (\omega_m - \omega_s)] \right\}_4
\end{aligned} \tag{7.2.7}$$

$$\begin{aligned}
\frac{\partial \zeta_2}{\partial t} = & - \{ ((\mathbb{V}_x)_m + (\mathbb{V}_x)_2) \cdot \nabla \zeta_2 + (\mathbb{V}_x)_2 \cdot \nabla (\eta_m + \zeta_2) \}_1 \\
& - \{ (\zeta_m + 2\zeta_2) D_2 + \zeta_2 D_m \}_2 + \left\{ \frac{1}{p_m - p_1} \zeta_2 (\omega_1 - \omega_m) \right\}_3 \\
& - \left\{ \frac{1}{p_m - p_1} |\mathbf{k} \cdot [(\mathbb{V}_\psi)_2 \times \nabla (\omega_1 - \omega_m)] \right\}_4
\end{aligned} \tag{7.2.8}$$

Subscript 1 indicates contribution from  $\mathbb{V}_x \cdot \nabla \eta$

Subscript 2 indicates contribution from  $\zeta \cdot \nabla \mathbb{V}$

Subscript 3 indicates contribution from  $\omega \frac{\partial \zeta}{\partial p}$

Subscript 4 indicates contribution from  $|\mathbf{k} \cdot (\frac{\partial \mathbb{V}}{\partial p} \times \nabla \omega)$

$\omega_m$  and  $\omega_1$  are the vertical motion at level  $p_m$  and  $p_1$  respectively.

The non-geostrophic terms will be approximated by the divergence computed from the geostrophic part of the equations for the preceding time step.

We can now express the forecasting equations in the following formal way (compare (2.2, 2.13 and 2.14).

$$\begin{aligned}
\nabla^2 \left\{ \frac{\partial}{\partial t} (\psi_m + c_1 \psi_2 - c_2 \psi_1) \right\} = & F_{mG} \\
+ F_{m1} + F_{m2} + F_{m3} + F_{m4}
\end{aligned} \tag{7.2.9}$$

$$\nabla^2 \frac{\partial \psi_1}{\partial t} - a_1 f^2 \frac{\partial \psi_1}{\partial t} + a_2 f^2 \frac{\partial \psi_2}{\partial t} = F_{1G} + F_{11} + F_{12} + F_{13} + F_{14} \quad (7.2.10)$$

$$\nabla^2 \frac{\partial \psi_2}{\partial t} - b_2 f^2 \frac{\partial \psi_2}{\partial t} + b_1 f^2 \frac{\partial \psi_1}{\partial t} = F_{2G} + F_{21} + F_{22} + F_{23} + F_{24} \quad (7.2.11)$$

$F_{mG}$ ,  $F_{1G}$  and  $F_{2G}$  are the geostrophic and non-adiabatic terms. They correspond to the right hand part of the equations 2.2, 2.13 and 2.14.

The non-geostrophic terms are the same as in the equations (7.2.6), (7.2.7) and (7.2.8).

## 8. NUMERICAL SOLUTION OF THE 3-PARAMETER MODEL

### 8.1 GENERAL ASPECTS

The equation of the model will be applied on a polar-stereographic projection. The polar-stereographic plane is assumed to cut the sphere at a latitude  $\phi_0$ . In the numerical computations  $\phi_0$  is put equal to  $60^\circ\text{N}$ , however, other values of  $\phi_0$  can easily be chosen. The grid-distance  $d$  can also be chosen arbitrarily. For further details see the program description in Appendix B.

The computational area can have any form, from an irregular octagon to a square. The only geometrical condition is that the inner angles of the area must be  $90^\circ$  or  $135^\circ$ .

The coordinate axis of the grid is positively oriented (see Figure 3). The computational area is specified by the coordinates of the corner points  $(x_1/y_1 \dots x_8/y_8)$  and by the coordinates of the north pole  $(x_{\text{pole}}/y_{\text{pole}})$ .

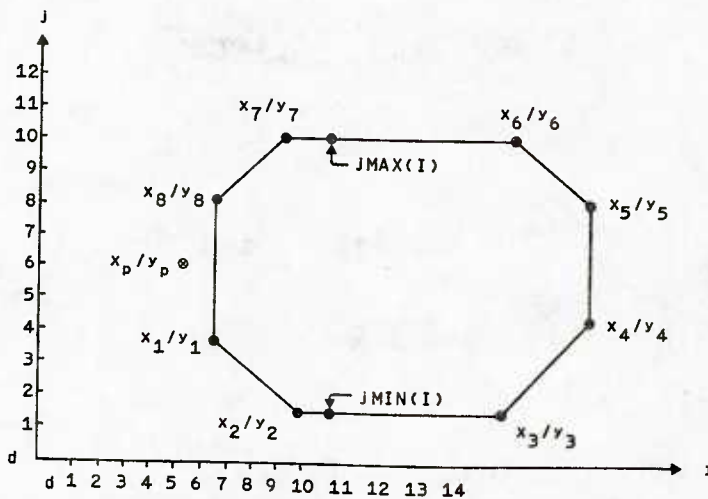


Figure 3. Example computational area.

If the area is reduced to a rectangle  $x_1 = x_2$ ,  $y_1 = y_2$ ,  
 $x_3 = x_4$ ,  $y_3 = y_4$ , etc.

## 8.2 FINITE-DIFFERENCES

In order to transform the differential equations to finite-difference equations we will introduce the following finite-difference notation and operators ( $\alpha$  and  $\beta$  are arbitrary quantities):

$$x \approx i\Delta s$$

$$y \approx j\Delta s$$

$$t \approx \tau\Delta t$$

$$f(x, y, t) \approx f_{i, j}^{\tau}$$

$$\nabla^2 \alpha \approx \frac{1}{d^2} \nabla^2 \alpha \quad (8.2.1a)$$

$$\mathbb{J}(\alpha, \beta) \approx \frac{1}{4d^2} \mathbb{J}(\alpha, \beta) \quad (8.2.1b)$$

$$\frac{\partial \alpha}{\partial t} \approx \frac{\Delta^{\tau} \alpha}{\epsilon \Delta t} \quad (8.2.1c)$$

$$(\nabla^2 \alpha)_{ij} = \alpha_{i+1, j} + \alpha_{i-1, j} + \alpha_{i, j+1} + \alpha_{i, j-1} - 4\alpha_{ij} \quad (8.2.2a)$$

$$\begin{aligned} \mathbb{J}(\alpha; \beta) = & (\alpha_{i+1} - \alpha_{i-1})_j (\beta_{j+1} - \beta_{j-1})_i \\ & - (\alpha_{j+1} - \alpha_{j-1})_i (\beta_{i+1} - \beta_{i-1})_j \end{aligned} \quad (8.2.2b)$$

$$\tau = 0 \text{ yields } \alpha^{\frac{1}{2}} - \alpha^0 \text{ and } \epsilon = \frac{1}{2}$$

$$\tau = \frac{1}{2} \text{ yields } \alpha^1 - \alpha^0 \text{ and } \epsilon = 1$$

$$\tau \geq 1 \text{ yields } \alpha^{t+1} - \alpha^{t-1} \text{ and } \epsilon = 2 \quad (8.2.2c)$$



Introducing the map-scale factor  $m = \frac{1+\sin\phi_0}{1+\sin\phi}$  and inserting the quantity  $\mu = \left(\frac{m}{d}\right)^2$  we get:

$$\nabla^2(\Delta\psi_1) - a_1 \frac{f^2}{\mu} (\Delta\psi_2) + a_2 \frac{f^2}{\mu} (\Delta\psi_2) = F_{1G}, \quad (8.2.3)$$

$$\nabla^2(\Delta\psi_2) - b_2 \frac{f^2}{\mu} (\Delta\psi_2) + b_1 \frac{f^2}{\mu} (\Delta\psi_1) = F_{2G}, \quad (8.2.4)$$

$$\nabla^2(\Delta\psi_M) - \frac{q}{\mu}(\Delta\psi_M) = F_{mG}. \quad (8.2.5)$$

$q$  is an empirical constant to adjust for the very long waves,  $q = 0.75 \cdot 10^{-12} m^{-2}$ .

Here we have (for simplicity we from now on put  $J = J$ ):

$$F_{1G} = F_1' + F_1'',$$

$$F_{2G} = F_2' + F_2'',$$

$$F_1' = - \varepsilon \Delta t \cdot \frac{f}{\mu} (a_3 \omega_s + a_1 H),$$

$$F_1'' = - \varepsilon \Delta t \frac{1}{4} [J_1 + J_2 + f^2 (a_2 J_4 - a_1 J_3)],$$

$$F_2' = - \varepsilon \Delta t \frac{f}{\mu} (b_3 \omega_s + b_1 H),$$

$$F_2'' = - \varepsilon \Delta t \frac{1}{4} [J_5 + J_6 + f^2 (b_1 J_3 - b_2 J_4)],$$

$$F_{mG} = - \varepsilon \Delta t \frac{1}{4} [J_7 + J_8 + J_9 - 4 \frac{f}{\mu} c_8 \omega_s].$$

We further have:

$$\beta_2 = \eta_m^{-2} \zeta_1,$$

$$\beta_6 = \eta_m^{+2} \zeta_2,$$

$$\beta_7 = c_3 \eta_m^{-c_2} \zeta_1 + c_4 \zeta_2 + c_7 f,$$

$$\beta_8 = -c_2 \eta_m^{+c_5} \zeta_1,$$

$$\beta_9 = c_4 \eta_m^{+c_6} \zeta_2 + 2c_7 f,$$

$$\begin{aligned}
J_1 &= J(\psi_m; \zeta_1), \\
J_2 &= J(\psi_1; \beta_2), \\
J_3 &= J(\psi_m; \psi_1), \\
J_4 &= J(\psi_m; \psi_2), \\
J_5 &= J(\psi_m; \zeta_2), \\
J_6 &= J(\psi_2; \beta_6), \\
J_7 &= J(\psi_m; \beta_7), \\
J_8 &= J(\psi_1; \beta_8), \\
J_9 &= J(\psi_2; \beta_9), \\
J_{10} &= J(\psi_m - 2\psi_1; p_s), \\
\omega_s &= \frac{H}{4} J_{10} - c_f (\zeta_m - 2\zeta_1).
\end{aligned}$$

$c_f$  is an empirical constant and a function of the exchange coefficient of momentum in the boundary layer.

### 8.3 COMPUTATION OF SENSIBLE HEAT

Equation (4.1.4) reads in finite-difference form:

$$H_{\text{SENS}} = 0.5 \cdot 10^{-2} h (0.1) \sqrt{\frac{H}{4} ((\Delta_x \psi_o)^2 + (\Delta_y \psi_o)^2) + 1} (T_s - h_2 \psi_1)$$

over ocean, if  $(T_s - h_2 \psi_1) > 0$ , (8.3.1)

$H_{\text{SENS}} = 0$  over land and if  $(T_s - h_2 \psi_1) < 0$ .

$$\Delta_x \psi_o = [\psi_m(i+1) - 2\psi_1(i+1)] - [\psi_m(i-1) - 2\psi_1(i-1)]$$

$$\Delta_y \psi_o = [\psi_m(j+1) - 2\psi_1(j+1)] - [\psi_m(j-1) - 2\psi_1(j-1)]$$

#### 8.4 COMPUTATION OF LATENT HEAT (See Gambo 1963)

The latent heat is introduced in the model by the expression:

$$\begin{aligned}
 H_{LAT} &= -hL\omega^*F^* && \text{if } \bar{\omega}_1 \leq -\delta_1 \\
 H_{LAT} &= 0 && \text{if } \bar{\omega}_1 > -\delta_1
 \end{aligned} \tag{8.4.1}$$

$$H_{LAT} = 0 \quad \text{if } \frac{r}{\epsilon \Delta t} < \text{tolerance}$$

$$\begin{aligned}
 \omega^* &= -(\delta_1 + \delta_2) \quad \text{if } \bar{\omega}_1 < -(\delta_1 + \delta_2) \\
 \omega^* &= -\left| \frac{\bar{\omega}_1^2}{\delta_1 + \delta_2} \right| \quad \text{if } -(\delta_1 + \delta_2) \leq \bar{\omega}_1 \leq -\delta_1
 \end{aligned} \tag{8.4.2}$$

$\delta_1$  and  $\delta_2$  are here two tolerances with the same dimension as  $\bar{\omega}_1$ ,  $r$  is the precipitation, and  $\delta_2$  is introduced for operational purposes.

$$F^* = F^*(p, T) = \frac{\epsilon \frac{T}{p} E(T) \left[ \frac{\epsilon L}{C_p} - T \right]}{pT^2 + \frac{\epsilon^2 L^2}{C_p R} E(T)} \tag{8.4.3}$$

$$E(T) = E_0 e^{\frac{\epsilon L}{R} \left( \frac{1}{T_0} - \frac{1}{T} \right)}$$

$$\epsilon = 0.622$$

$$L = 2.5 \cdot 10^6$$

$$P = P_{\text{mean}} = \frac{P_m + P_0}{2}$$

$$C_p = 1004$$

$$T_0 = 273$$

$$E_0 = 0.611$$

$$T = \frac{2f_0}{P} \psi_1 = h_6 \psi_1$$

$$R \ln \left( \frac{P_0}{P_m} \right)$$

#### 8.4.1 Computation of $\tilde{W}$ and $\tilde{D}$

We will compute the humidity in the layer

$P_0 = 100$  cb and  $P_T = 30$  cb; defining  $\tilde{W}$  as

$$\tilde{W} = \frac{1}{70} \int_{30}^{100} W(P) E(P) dp$$

and inserting the wind profile yields

$$\tilde{W} = e_m W_m + e_1 W_1 + e_2 W_2 \quad (8.4.4)$$

where

$$\begin{aligned} e_m &= \frac{1}{70} [15E(100)a_{100} + 25E(70)a_{70} + 20E(50)a_{50} + 10E(30)a_{30}], \\ e_1 &= \frac{1}{70} [15E(100)b_{100} + 25E(70)b_{70} + 20E(50)b_{50} + 10E(30)b_{30}], \\ e_2 &= \frac{1}{70} [15E(100)c_{100} + 25E(70)c_{70} + 20E(50)c_{50} + 10E(30)c_{30}]. \end{aligned} \quad (8.4.5)$$

The constants  $a_p$ ,  $b_p$  and  $c_p$ , where  $p$  assumes the values 100, 70, 50, 30 are computed for the three following alternatives:

I	II	III
$p_m \leq p \leq p_0$	$p_1 \leq p \leq p_m$	$0 \leq p \leq p_1$
$a_p = 1$	$a_p = 1$	$a_p = \frac{p}{p_1}$
$b_p = -\frac{2(p-p_m)}{p_0-p_m}$	$b_p = 0$	$b_p = 0$
$c_p = 0$	$c_p = \frac{2(p_m-p)}{p_m-p_1}$	$c_p = \frac{2p}{p_1}$

$$E(100)=2.5415, E(70)=0.9683, E(50)=0.3527, E(30)=0.0613$$

Correspondingly we get

$$\tilde{D} = e_m D_m + e_1 D_1 + e_2 D_2 = (e_1 + e_m c_2) D_1 + (e_2 - e_m c_1) D_2 - e_m c_8 w_s$$

#### 8.4.2 Computation of Precipitation

Equation (4.2.6) reads in finite-difference form (the term  $\nabla \tilde{\chi} \nabla w$  disregarded and  $w$  replaced with  $q$ ):

$$q^{\tau+1} = q^{\tau-1} + \varepsilon \Delta t H_q^\tau,$$

$$H_q^\tau = -\frac{\mu}{4} J(\tilde{\psi}^\tau; q^\tau) - q^\tau (\tilde{D}^\tau + dw_s^\tau) + A_{diff} \mu \nabla^2 q^\tau. \quad (8.4.6)$$

The precipitation is computed in the following way (precipitation is indicated by  $r$ ):

$$q_{SAT} = q(\bar{T}_\psi) = q(h_3 \psi_1 + h_4 \psi_2)$$

$$\text{If } (q^{\tau+1} - 0.8 q_{SAT}) > 0 \text{ then } q^{\tau+1} = 0.8 q_{SAT},$$

$$r = \Delta p (q^{\tau+1} - 0.8 q_{SAT}).$$

$$\text{If } (q^{\tau+1} - 0.8 q_{SAT}) < 0 \text{ then } r = 0 \text{ and if}$$

$$(q^{\tau+1} - 0.2 q_{SAT}) < 0 \text{ then } q_{mod}^{\tau+1} = 0.2 q_{SAT}. \quad (8.4.7)$$

$r$  is accumulated for every timestep and printed out at certain prescribed times. See Appendix B.

## 8.5 NUMERICAL SOLUTION OF THE NON-GEOSTROPHIC TERMS

The finite difference analogues for (7.2.9), (7.2.10) and (7.2.11) read:

$$\nabla^2 (\Delta\psi_M) - \frac{g}{\mu}(\Delta\psi_M) = F_{mG} + \frac{\varepsilon\Delta t}{\mu} \sum_{i=1}^4 F_{mi} \quad (8.5.1)$$

$$\nabla^2 (\Delta\psi_1) - a_1 \frac{f^2}{\mu} (\Delta\psi_1) + a_2 \frac{f^2}{\mu} (\Delta\psi_2) = F_{1G} + \frac{\varepsilon\Delta t}{\mu} \sum_{i=1}^4 F_{1i} \quad (8.5.2)$$

$$\nabla^2 (\Delta\psi_2) - b_2 \frac{f^2}{\mu} (\Delta\psi_2) + b_1 \frac{f^2}{\mu} (\Delta\psi_1) = F_{2G} + \frac{\varepsilon\Delta t}{\mu} \sum_{i=1}^4 F_{2i} \quad (8.5.3)$$

$$\begin{aligned} F_{m1} = & - \frac{\mu}{2} \{ \nabla\chi_m \cdot \nabla [ \underbrace{c_3(\zeta_m+f) - c_2\zeta_1 + c_4\zeta_2 + c_7f}_{\beta_7} ] \\ & + \nabla\chi_1 \cdot \nabla [ \underbrace{-c_2(\zeta_m+f) + c_5\zeta_1}_{\beta_8} ] \\ & + \nabla\chi_2 \cdot \nabla [ \underbrace{c_4(\zeta_m+f) + c_6\zeta_2 + 2c_7f}_{\beta_9} ] \} \end{aligned} \quad (8.5.4a)$$

$$\begin{aligned} F_{m2} = & [D_1(k_1\zeta_1 + k_2\zeta_2 + k_3\zeta_m) + D_2(k_4\zeta_1 + k_5\zeta_2 + k_6\zeta_m) \\ & + \omega_s(k_7\zeta_1 + k_8\zeta_2 + k_9\zeta_m)] \end{aligned} \quad (8.5.4b)$$

$$\begin{aligned} F_{m3} = & [D_1(k_{10}\zeta_1 + k_{11}\zeta_2 + k_{12}\zeta_m) + D_2(k_{13}\zeta_1 + k_{14}\zeta_2 + k_{15}\zeta_m) \\ & + \omega_s(k_{16}\zeta_1 + k_{17}\zeta_2 + k_{18}\zeta_m)] \end{aligned} \quad (8.5.4c)$$

$$F_{m4} = \frac{\mu}{2} \{ \nabla\psi_1 \cdot \nabla [ \underbrace{k_{19}D_1 + k_{20}D_2 + k_{21}\omega_s}_{\gamma_7} ] \}$$

$$\begin{aligned}
& + \nabla\psi_2 \cdot \nabla \underbrace{[k_{22}D_1 + k_{23}D_2 + k_{24}\omega_s]}_{\gamma_8} \\
& + \nabla\psi_m \cdot \nabla \underbrace{[k_{25}D_1 + k_{26}D_2 + k_{27}\omega_s]}_{\gamma_9}
\end{aligned} \tag{8.5.4d}$$

$$F_{11} = - \frac{\mu}{2} \{ \nabla\chi_m \cdot \nabla\zeta_1 + \nabla\chi_1 \cdot \nabla(\zeta_m + f - 2\zeta_1) \} \tag{8.5.5a}$$

$$F_{12} = [D_1(k_{28}\zeta_1 - \zeta_m) + D_2k_{29}\zeta_1 + \omega_s k_{30}\zeta_1] \tag{8.5.5b}$$

$$F_{13} = [\zeta_1(k_{31}D_1 + k_{32}D_2 + k_{33}\omega_s)] \tag{8.5.5c}$$

$$F_{14} = \frac{\mu}{2} \{ \nabla\psi_1 \cdot \nabla \underbrace{[k_{31}D_1 + k_{32}D_2 + k_{33}\omega_s]}_{\gamma_{10}} \} \tag{8.5.5d}$$

$$F_{21} = - \frac{\mu}{2} \{ \nabla\chi_m \cdot \nabla\zeta_2 + \nabla\chi_2 \cdot \nabla(\zeta_m + f + 2\zeta_2) \} \tag{8.5.6a}$$

$$F_{22} = \{ D_1k_{34}\zeta_2 + D_2[k_{35}\zeta_2 - \zeta_m] + \omega_s k_{30}\zeta_2 \} \tag{8.5.6b}$$

$$F_{23} = [\zeta_2(k_{36}D_1 + k_{37}D_2 + k_{38}\omega_s)] \tag{8.5.6c}$$

$$F_{24} = \frac{\mu}{2} \{ \nabla\psi_2 \cdot \nabla \underbrace{[k_{36}D_1 + k_{37}D_2 + k_{38}\omega_s]}_{\gamma_{11}} \} \tag{8.5.6d}$$

The computation of the non-geostrophic forcing function

$$\frac{\varepsilon\Delta t}{\mu} \sum_{i=1}^4 F_{mi}, \quad \frac{\varepsilon\Delta t}{\mu} \sum_{i=1}^4 F_{li} \quad \text{and} \quad \frac{\varepsilon\Delta t}{\mu} \sum_{i=1}^4 F_{2i}$$

are performed by a separate program. See Appendix B.

### 8.5.1 Numerical Coefficients for the Non-geostrophic Terms

Expressions for constants of the non-geostrophic terms  
read:

$$\bar{\omega}_3 = t_6 \omega_s + t_7 D_1 + t_8 D_2$$

$$\omega_m = t_9 \omega_s + t_{10} D_1 + t_{11} D_2$$

$$\omega_1 = t_{12} \omega_s + t_{13} D_1 + t_{14} D_2$$

$$t_6 = c_8 \frac{p_1}{6}$$

$$t_7 = -c_2 \frac{p_1}{6}$$

$$t_8 = (c_1 - 2) \frac{p_1}{6}$$

$$t_9 = 1 - c_8 (p_o - p_m)$$

$$t_{10} = (c_2 - 1) (p_o - p_m)$$

$$t_{11} = -c_1 (p_o - p_m)$$

$$t_{12} = c_8 \frac{p_1}{2}$$

$$t_{13} = -c_2 \frac{p_1}{2}$$

$$t_{14} = (c_1 - 2) \frac{p_1}{2}$$

$$k_1 = c_2 (c_2 - \frac{4}{3})$$

$$k_2 = -c_2 c_4$$

$$k_3 = c_2 c_7$$



$$k_4 = -c_1 c_2$$

$$k_5 = c_1 c_4 - c_6$$

$$k_6 = c_7 (2 - c_1)$$

$$k_7 = -c_2 c_8$$

$$k_8 = c_4 c_8$$

$$k_9 = c_8 (1 - c_7)$$

$$k_{10} = 2t_2 c_8$$

$$k_{11} = -2(t_3 + t_7) c_8$$

$$k_{12} = -t_7 c_8$$

$$k_{13} = -2t_3 c_8$$

$$k_{14} = 2(t_5 - t_8) c_8$$

$$k_{15} = -t_8 c_8$$

$$k_{16} = 2t_1 c_8$$

$$k_{17} = 2(t_4 - t_6) c_8$$

$$k_{18} = -t_6 c_8$$

$$k_{19} = 2t_2 c_8$$

$$k_{20} = -2t_3 c_8$$

$$k_{21} = 2t_1c_8$$

$$k_{22} = -2(t_3+t_7)c_8$$

$$k_{23} = 2(t_5-t_8)c_8$$

$$k_{24} = 2(t_4-t_6)c_8$$

$$k_{25} = -t_7c_8$$

$$k_{26} = -t_8c_8$$

$$k_{27} = -t_6c_8$$

$$k_{28} = 2-c_2$$

$$k_{29} = +c_1$$

$$k_{30} = +c_8$$

$$k_{31} = -\frac{t_{10}}{(p_o-p_m)}$$

$$k_{32} = -\frac{t_{11}}{(p_o-p_m)}$$

$$k_{33} = \frac{1-t_9}{(p_o-p_m)}$$

$$k_{34} = -c_2$$

$$k_{35} = c_1^{-2}$$

$$k_{36} = \frac{t_{10}-t_{13}}{(p_m-p_1)}$$

$$k_{37} = \frac{t_{11} - t_{14}}{(p_m - p_1)}$$

$$k_{38} = \frac{t_9 - t_{12}}{(p_m - p_1)}$$

For  $p_o=100$ ,  $p_m=50$  and  $p_1=30$  we have the following numerical values for the constants:

$$t_6 = 0.0588235$$

$$t_7 = -2.941175$$

$$t_8 = -7.058825$$

$$t_9 = 0.411765$$

$$t_{10} = -20.588250$$

$$t_{11} = -29.411750$$

$$t_{12} = 0.176471$$

$$t_{13} = -8.823525$$

$$t_{14} = -21.176475$$

$$k_1 = -0.438291$$

$$k_2 = -0.276816$$

$$k_3 = 0.034602$$

$$k_4 = -0.346020$$

$$k_5 = -0.507498$$

$$k_6 = 0.083045$$

$$k_7 = -0.006920$$

$$k_8 = 0.005536$$

$$k_9 = 0.011073$$

$$k_{10} = -0.438294$$

$$k_{11} = -0.276817$$

$k_{12} = 0.034602$   
 $k_{13} = -0.346021$   
 $k_{14} = -0.507498$   
 $k_{15} = 0.083045$   
 $k_{16} = 0.016609$   
 $k_{17} = 0.005536$   
 $k_{18} = -0.000692$   
 $k_{19} = -0.438293$   
 $k_{20} = -0.346021$   
 $k_{21} = 0.016609$   
 $k_{22} = -0.276817$   
 $k_{23} = -0.507497$   
 $k_{24} = 0.005536$   
 $k_{25} = 0.034602$   
 $k_{26} = 0.083045$   
 $k_{27} = -0.000692$   
 $k_{28} = 1.411765$   
 $k_{29} = +0.588235$   
 $k_{30} = +0.0117647$   
 $k_{31} = 0.411765$   
 $k_{32} = 0.588235$   
 $k_{33} = 0.0117647$   
 $k_{34} = -0.588235$   
 $k_{35} = -1.411765$

$$k_{36} = -0.588235$$

$$k_{37} = -0.411765$$

$$k_{38} = 0.0117647$$

## 9. INITIALIZATION

The stream functions  $\psi$  are computed from the geopotential  $Z$  by the relation

$$g\nabla^2 Z = \nabla \cdot (f\nabla\psi) \quad \text{or}$$

$$\nabla^2 \psi = \frac{g}{f}\nabla^2 Z - \frac{1}{f}\nabla f \cdot \nabla\psi; \quad (9.1)$$

inserting

$$\nabla\psi = \frac{g}{f}\nabla Z \quad \text{yields}$$

$$\nabla^2 \psi = \frac{g}{f}\nabla^2 Z - \frac{g}{f^2}\nabla f \cdot \nabla Z. \quad (9.2)$$

The boundary values of  $\psi$  are computed by the relation ( $\gamma$  is a constant)

$$\frac{\partial\psi}{\partial s} = \frac{g}{f} \frac{\partial Z}{\partial s} + \gamma. \quad (9.3)$$

If we assume that there is no net flow out of the area it is easily seen that  $\gamma = -\frac{1}{L} \oint \frac{g}{f} \frac{\partial Z}{\partial s} ds$ . The integration is performed along the boundary of the area in positive order. The integration starts in point  $x_1/y_1$  where we put

$$\psi(x_1; y_1) = \frac{g}{f(x_1; y_1)} Z(x_1; y_1). \quad (9.4)$$

$Z$  is computed from  $\psi$  in a similar way using the following equation:

$$\nabla^2 Z = \frac{f}{g}\nabla^2 \psi + \frac{1}{g}\nabla f \cdot \nabla\psi. \quad (9.5)$$

Boundary values for  $Z$  are computed initially and stored in a special string. See Appendix B.

9.1 NUMERICAL SOLUTION

Equation (9.2) reads in finite difference form:

$$\nabla^2 \psi = \frac{g}{f} \nabla^2 Z - \frac{g}{2f^2} \nabla f \cdot \nabla Z \tag{9.6}$$

where the standard five point formulas are used to compute  $\nabla^2 \psi$ ,  $\nabla^2 Z$  and  $\nabla f \cdot \nabla Z$  is defined as

$$\begin{aligned} \nabla f \cdot \nabla Z = & (f_{i+1} - f_i)_j (Z_{i+1} - Z_i)_j + (f_i - f_{i-1})_j (Z_i - Z_{i-1})_j \\ & + (f_{i+1} - f_j)_i (Z_{i+1} - Z_j)_i + (f_j - f_{j-1})_i (Z_j - Z_{j-1})_i. \end{aligned}$$

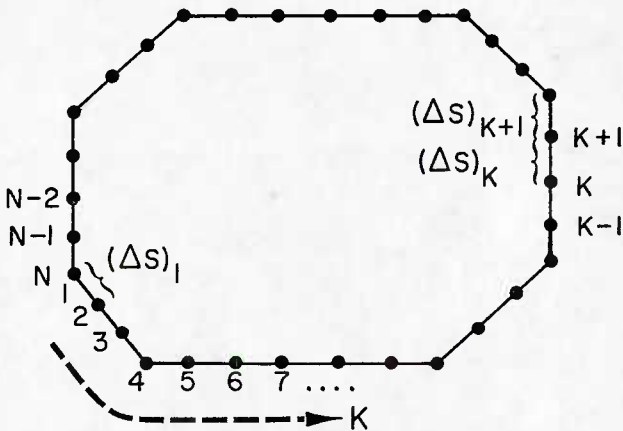
Equation (9.3) reads in finite difference form:

$$\psi_{k+1} = \psi_k + 2g \frac{Z_{k+1} - Z_k}{f_{k+1} + f_k} + \gamma (\Delta s)_k \tag{9.7}$$

where

$$\gamma = - \frac{1}{L} 2g \sum_{k=1}^{N-1} \frac{Z_{k+1} - Z_k}{f_{k+1} + f_k} \tag{9.8}$$

Figure 4 defines k and the order of integration.



Z to  $\psi$  and  $\psi$  to Z is computed by the program STREAMF (see Appendix B).

Figure 4. Definition of k and order of integration

## 9.2 INITIALIZATION OF THE SPECIFIC HUMIDITY

Since we are using the  $\psi$ -functions as the time-dependent variables, we have to modify the initial humidities in order to make them consistent with the time-integration.

Therefore, we put

$$q = q_{\text{analyzed}} \cdot \frac{q_{\text{sat}}(\bar{T}_{\psi})}{q_{\text{sat}}(\bar{T}_Z)} \quad (9.2.1)$$

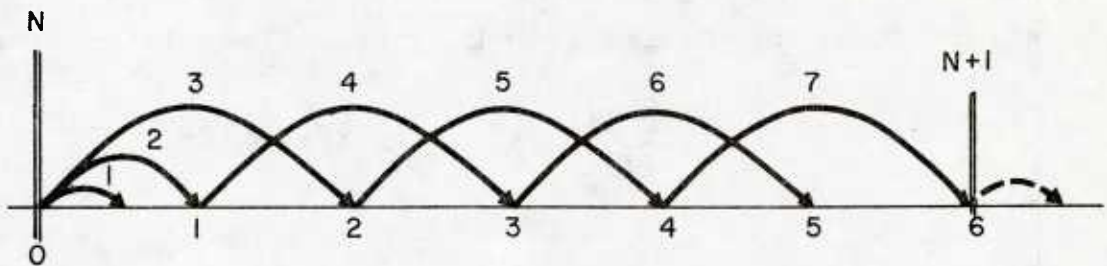
$$\bar{T}_Z = \frac{2g}{R \ln 2} \left( \frac{Z_{500} - Z_{1000}}{2} \right) = \frac{g}{f_0} (h_3 Z_1 + h_4 Z_2), \text{ and}$$

$$\bar{T}_{\psi} = h_3 \psi_1 + h_4 \psi_2.$$



## 10. TIME-INTEGRATION AND TREATMENT OF LATERAL BOUNDARY VALUES

The forecast is basically computed in 6-hour intervals but this interval can easily be changed. The first time step in each interval is non-centered. Smoothing, elliptization and printing of results (if so desired) are performed at the end of every interval.



$$\frac{\partial \psi}{\partial t} \sim \frac{\Delta \psi}{\epsilon \Delta t} \quad \Delta t = 1 \text{ hour in this example} \quad (10.1)$$

ND number of  $\Delta t$  for the interval

kT time step index;  $k = 1, 2, 3 \dots ND+1$

N index for every interval integration (e.g., 6-hour interval)

N	Forecast length
0	0
1	+6
2	+12
3	+18
⋮	⋮
⋮	⋮
⋮	⋮

Initial height fields  $Z^0$  are stored on secondary storage during the whole computation. In the case of variable boundary conditions,  $Z^0$  is followed by  $Z^N$  ( $N=1,2,3\dots$ ) (forecasts for each interval).

Assuming

$$\frac{\partial \psi}{\partial t} = \frac{g}{F} \frac{\partial Z}{\partial t} \quad \text{or} \quad \psi^N - \psi^0 = \frac{g}{F} (Z^N - Z^0) \quad (10.2)$$

the stream function at time step  $k$  can be interpolated from:

$$\psi^k = (\psi^0 - \frac{g}{F} Z^0) + \frac{g}{F} [(1-\alpha_k) Z^N + \alpha_k Z^{N+1}] \quad (10.3)$$

where

$$\alpha_k = \frac{kT-1}{ND}$$

At each time step this interpolated stream function is mixed with the forecasted stream function  $\psi_{\text{prog}}^k$  in the following way:

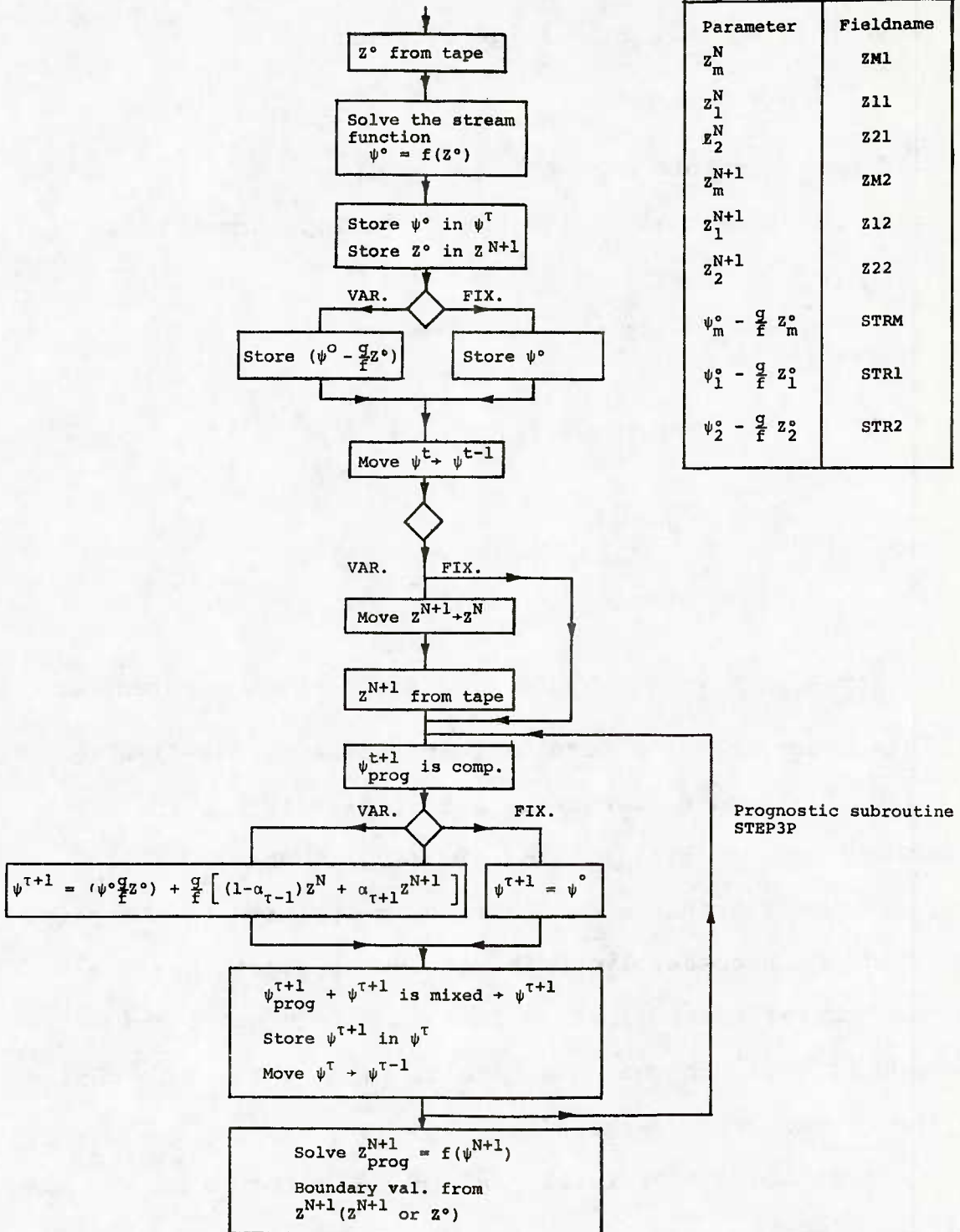
Boundary values	$\psi_{\text{mod}} = \psi_{\text{prog}}^k + w1(\psi^k - \psi_{\text{prog}}^k)$
1 grid point inside the boundary	$\psi_{\text{mod}} = \psi_{\text{prog}}^k + w2(\psi^k - \psi_{\text{prog}}^k)$
2 grid points inside the boundary	$\psi_{\text{mod}} = \psi_{\text{prog}}^k + w3(\psi^k - \psi_{\text{prog}}^k)$
3 grid points or more inside the boundary	$\psi_{\text{mod}} = \psi_{\text{prog}}^k$

(10.4)

$w1, w2$  and  $w3$  are 3 predetermined constants with  $0 \leq w1, w2, w3 \leq 1$

The following flow diagram illustrates the treatment of the boundary values:

Flow diagram: Treatment of boundary values.



Parameter	Fieldname
$Z_m^N$	ZM1
$Z_1^N$	Z11
$Z_2^N$	Z21
$Z_m^{N+1}$	ZM2
$Z_1^{N+1}$	Z12
$Z_2^{N+1}$	Z22
$\psi_m^\circ - \frac{q}{f} Z_m^\circ$	STRM
$\psi_1^\circ - \frac{q}{f} Z_1^\circ$	STR1
$\psi_2^\circ - \frac{q}{f} Z_2^\circ$	STR2

Prognostic subroutine  
STEP3P

## II. THE CRITERION OF ELLIPTICITY

An iterative procedure is used to modify a stream-function field so that the ellipticity criterion

$$\nabla^2 \psi + \frac{f}{2} > 0$$

is valid in all points of the field.

Each point is tested with the following formula:

$$\mu \nabla^2 \psi + \frac{f}{2} - \epsilon = \delta$$

where

$$\nabla^2 \psi = \psi_{i+1,j} + \psi_{i-1,j} + \psi_{i,j+1} + \psi_{i,j-1} - 4 \cdot \psi_{i,j}$$

and

$$\epsilon = 0.001 \cdot f.$$

If  $\delta < 0$ ,  $\psi_{i,j}$  is modified by

$$\psi_{i,j} = \psi_{i,j} - \frac{\epsilon - \delta}{2\mu} \cdot k \text{ where } k \text{ is a convergence parameter.}$$

This means that the vorticity increases by  $2(\epsilon - \delta) \cdot k$  in the point  $(i,j)$  and decreases by  $0.5(\epsilon - \delta) \cdot k$  in the four surrounding points  $(i+1,j)$ ,  $(i-1,j)$ ,  $(i,j+1)$  and  $(i,j-1)$ . This procedure guarantees that  $\delta$  becomes positive in the point  $(i,j)$ , but not necessarily in the surrounding points. The test must therefore be repeated for all points until the criterion is valid in the whole field. With a suitable choice of  $k$  the method is convergent.

$k$  can be found empirically and is estimated to be of the order  $k \approx 0.85$ .

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## APPENDIXES

APPENDIX A  
GENERAL PRINCIPLES  
IN THE  
PROGRAMMING OF THE THREE PARAMETER MODEL

The programming has been performed in a very general way. It is therefore possible:

- a. to generate initial fields and run the model in a channel with variable cyclic boundary conditions (this is the case for the actual program).
- b. to run the model with constant boundary values on a given area (see above).
- c. to run the model with variable boundary values, the boundary values taken from another model.

The model needs 13 fields in the core;

1 field for the coriolis parameter	F,
1 field for $\mu = \frac{m^2}{d^2}$	MY,
1 special field indicator	MARK,
10 fields for the model computations	F1-F10.

In addition to this, 40 fields (28 for the quasi-geostrophic part of the model) are specified on secondary memories (disks, drums or large extended core).

"Household" programs:

PUT1:        This program computes FPAR (see common,  
                 Subroutine JMIMA)

- MARKF: This program computes a special integer field MARK. The program MARK specifies status of the field. Points outside the area = 0, points inside the are <0 and points on the boundary >0. The corner points etc. are specified according to given examples.
- MYFF: This program computes  $f$  and  $\mu$  and puts the result in the fields F and MY.
- RANWT: Writes fields on secondary storage e.g., disk drums, or extended core storage.
- RANRD: Reads fields from secondary storage e.g., disk, drums, or extended core storage.
- MAP: Prints pattern on line printer; 0 (zero) points, resolution and map scale are specified.
- GENCH: Generates initial state for a channel flow.
- BMOVE: Administrates computation of cyclic boundary conditions for a channel flow.

Since the axis of the grid is defined differently from what is used in Fortran, the programming of finite difference operators should be performed in a special way. This is not necessary, but speeds up the computation considerably. As an example, subroutine JACOB reads:



.  
.  
.  
.  
.  
MI = M-1

Dφ 10 I = 2,MI

J1 = JMIN(I)+1

J2 = JMAX(I)-1

K = (J1-1)\*M+I

Dφ 10 J = J1,J2

C(K) = (A(K+1) - A(K-1))\*(B(K+M) - B(K-M)).....

10 K = K + M

.  
.  
.

APPENDIX B  
PROGRAM SPECIFICATIONS

The following programs are written for the 3-parameter model.

Level 1: Main program

Level 2: Subroutines called by level 1

Level 3: Subroutines called by level 2

Level 4: Subroutines called by level 3

<u>Level</u>	<u>Program</u>
1	PROG3P
2	PUT1
3	JMIMA
2	COEFF3P
2	MARKF
2	MYFF
2	STREAMF
3	BDRGDR
3	BDRVAL
2	SATUR
2	RANWT (RANRD)
2	ELLIPT
2	GENCH
2	ASMUT
2	MAP3P
3	MAP
2	STEP3P

<u>Level</u>	<u>Program</u>
3	JACOB
3	ABSVOR (RELVOR)
3	MIXF
3	HELM (POIS)
3	HELMSYS
3	BMOVE
3	STEPEXT
4	GRADPR
4	VELPOT

Program PROG3 (Main program for 3-parameter model, see flow diagram B-1 and description.)

Arrays

<u>Name</u>	<u>Dimensions</u>	<u>Description</u>
F1-F10	2	Working fields in core. See flow diagram B-1.
F,MY (real)	2	Fields for Coriolis parameter, $f$ , and $\mu = \left(\frac{m}{d}\right)^2$ . $m$ is the mapscale factor for a polar stereographic projection and $d$ is the grid length. F and MY are generated by the subroutine MYFF.
MARK	2	Marking of each grid point by a special integer (index). See description. MARK is generated by the subroutine MARKF.
UPS, UPM, UPL	1	Zonal wind profiles at the levels $p_s$ , $p_m$ and $p_l$ for the initial fields in the channel case. The values are introduced by DATA statements.
WX	1	Working array for generation of initial fields in the channel case. Dimension shall be at least equal to the number of grid-points across the channel.

<u>Name</u>	<u>Dimensions</u>	<u>Description</u>
NX,NY	1	Wave numbers for perturbations of the initial fields in the channel case in x and y directions. The values are introduced by DATA statements.
PSIC, PSIS, LAMC, LAMS	1	Amplitudes (PSI) and phase lags (LAM) perturbations of the initial field with cosine and sine profiles in the y direction respectively. For both profiles a sine wave is used in the x-direction. The values are introduced by DATA statements. See further description of GENCH.
ZB,PSIB,FB,SB	1	Boundary strings for storage of boundary values in counterclockwise order, starting with corner point 1 (see FPAR). Used by the subroutine STREAMF.
IT	2	Storage of the number of iterations for the solution of a): the system of two Helmholtz equations and b): the Helmholtz equation for the mean field. Values for an integration interval are stored in the array by the subroutine STEP3P.

<u>Name</u>	<u>Dimensions</u>	<u>Description</u>
MAPHOUR (1)	1	Integers giving the hours for mapping (2).
NSMUTT (1)	1	Integers giving the hours for smoothing (2).
MELLIPT (1)	1	Integers giving the hours for ellipticity (2).
LETACC (1)	1	Integers giving the hours at which the accumulation of precipitation shall be interrupted (2).

---

(1) The values shall be given in DATA statements and must be multiples of the integration interval (e.g., 6 hours). Unused positions are indicated by -1.

(2) This value can be increased if this is necessary.



Common areas

/FPAR/

<u>Name</u>	<u>Type</u>	<u>Description</u>
IC(8),JC(8)	integer	i- and j-coordinates for the corner points of the area.
XPOL,YPOL	real	i- and j-coordinates for the north pole. Specified relative to the origin of the grid.
R	real	radius of the earth.
RE	real	Distance from the pole to the equator on the map in grid-length units.
DS	real	grid-length (in meters).
JMIN(100), JMAX(100)	integer	minimum and maximum j-coordinate for each i-column in the field.
M,N	integer	Dimension of the field arrays.
KIND	integer	Channel indicator. Channel = 1, No channel = 0. Value shall be given by DATA statement.

---

/FORM1/

IC1(8),JC1(8)	integer	Same as IC(8),JC(8) for actual field. Values introduced by DATA statement.
---------------	---------	--



<u>Name</u>	<u>Type</u>	<u>Description</u>
JMIN1(100), JMAX1(100)	integer	Same as JMIN(100), JMAX(100)
XP,YP,D1	real	Same as XPOL, YPOL, DS. Values shall be introduced by DATA statements. D1 is given in km.
/KANAL/		
FIM	real	Used for channel computations. Latitude for the middle of the channel. Used for computation of B-plane. To be given by DATA statement.
/DRM/*		
MN	integer	=M*N. To be computed in main program
NDIM	integer	Core memory space available for field arrays. The arrays F1-F10, F,MY,MARK must be included in this area. Additional space is used for storage of fields in COMMON area//.
NFLD	integer	Number of field arrays in core
SL	real	Working area for core memory fields. The dimension must be at least SL(NDIM)

---

Arrays stored at extended core, disks or drums. 1 in the end  
of the name indicate timestep  $\tau$ , 2 in the end of the name  
timestep  $\tau-1$ .

---

\* Not necessary to specify if only extended core storage is used.

PSIM1,PSI11,PSI21 PSIM2,PSI12,PSI22	real	Storage arrays for stream functions. See flow diagram B-1.
HUM1,HUM2,DIV1 DIV2,WS,HEAT	real	Storage arrays for humidity, divergence, vertical velocity at the lower boundary, and heat.
J789,J12,J56,J3		Storage arrays for Jacobians.
PS,TS,PREC		Storage arrays for surface standard pressure, sea surface temperature, and accumulated precipitation.
STRM,STR1,STR2,ZM1 Z11,Z21,ZM2,Z12,Z22	real	Storage arrays. See Chapter 10 (lateral boundary mixing).
H11,H21,HM1		Fields for advection of vorticity by the divergent wind, thickness field (1,2) and mean field (M).
H12,H22,HM2		$\omega \frac{\partial \zeta}{\partial t} + \zeta \nabla \cdot \mathbf{W}$ for thickness fields (1,2) and mean field (M).
H13,H23,HM3		Twisting term for thickness fields (1,2) and for mean field, M.
V1,V2,VM		Velocity potential for thickness fields (1,2) and for mean field (M).
ID(200)*	integer	Catalog array for direct memory access.

---

\* Not necessary to specify if only extended core storage is used.

/COEFF/

A1,A2.....,T5      real      Constants used in the model.  
Computed by the subroutine  
COEFF3P from given values on  
stability and pressure levels.

/COEFF2/

T6,TF.....,T38      real      Constants used for the computation  
of the non-geostrophic terms.  
Computed by COEFF3P from given  
values on stability and pressure  
levels.

/RUNPAR/

DELTA                      real      Timestep in sec. computed in the  
main program.

NTSTEP                    integer      Number of timesteps for an inte-  
gration interval (6 hours).  
To be defined in a DATA statement.

ALFASYS, ALFAM            real      Overrelaxation coefficients (ALFA)  
ALFAZ, ALFAPSI,            and maximum residual in the solu-  
RESSYS, RESM,              tion for the system of the two  
RESZ, RESPSI                Helmholtz equations (SYS), Helmholtz  
equation for the mean field equa-  
tion (M), solution of Z from  $\psi$   
(Z) and solution of  $\psi$  from Z  
(PSI). To be defined by a DATA  
statement.

Q,FOCEAN,FCONT	real	The Helmholtz term for the mean field equation, friction coefficients over ocean and land. To be defined by a DATA statement.
WGT1,WGT2,WGT3	real	Weights for mixing near the boundary of boundary fields with forecasted fields. To be defined by a DATA statement.
ADIFF	real	Diffusion coefficient for humidity; it is defined by a DATA statement
Parameters only specified in Data statements.		
STAB1	real	Static stability of layer 1
STAB2	real	Static stability of layer 2
PNIVS*	real	Pressure level $P_0$
PNIVM	real	Pressure level $P_m$
PNIV1	real	Pressure level $P_1$
IVAR	integer	Indicates if variable lateral boundaries are used; IVAR=0, constant boundary values; IVAR=1, variable boundary values.
KIND	integer	Indicates if channel is used (cyclic boundary conditions); KIND=1, channel; KIND=0, no channel (polar stereographic projection).

Subroutine STEP3P (F1,F2,F3,F4,F5,F6,F7,F8,F9,F10,MY,F,MARK,  
IT,IVAR,M,N)

The subroutine performs a computation for a time interval (6 hours) by the 3P-model including humidity and precipitation prediction. (See Flow Diagram B-2.)

<u>Subroutine Parameters</u>	<u>Description</u>
F1-F10	Working fields.
MY	$\mu$ -parameter field.
F	Coriolis parameter field.
MARK	Field indicator field.
IT	Array to store number of iterations for every timestep.
IVAR	Indicator for constant or variable boundary conditions; IVAR = 0 constant boundary; IVAR = 1 variable boundary.
M,N	Field vector.

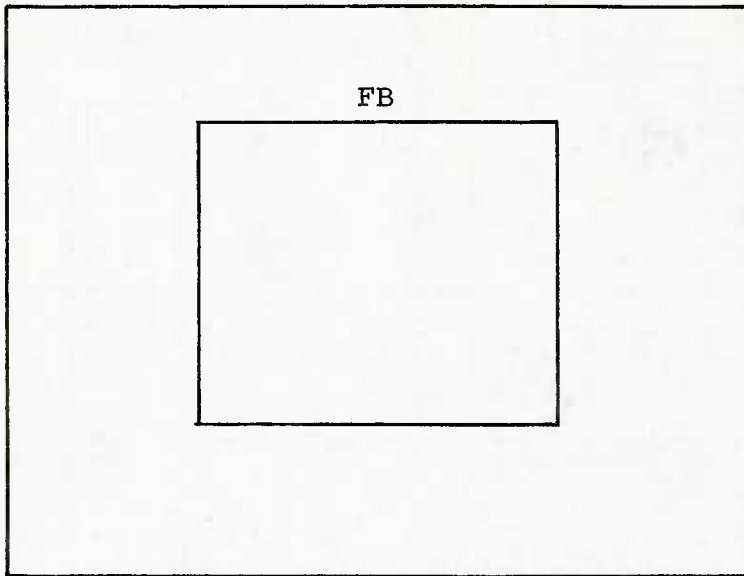
Parameters specified only in DATA statements

RGAS	R = 287
EE	E = 0.622
HL	L = $2.5 \cdot 10^6$
CP	Cp = 1004
TO	To = 273
EO	Eo = 0.611
DEL1	$\delta_1$
DEL2	$\delta_2$
TOL	Tolerance

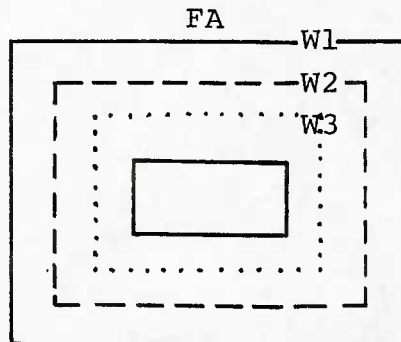
For explanation see chapter 8.4.

Subroutine MARK

The subroutine mixes field FA (corresponding to limited area) with FB (corresponding to a field taken from a large area).



large comp.  
area



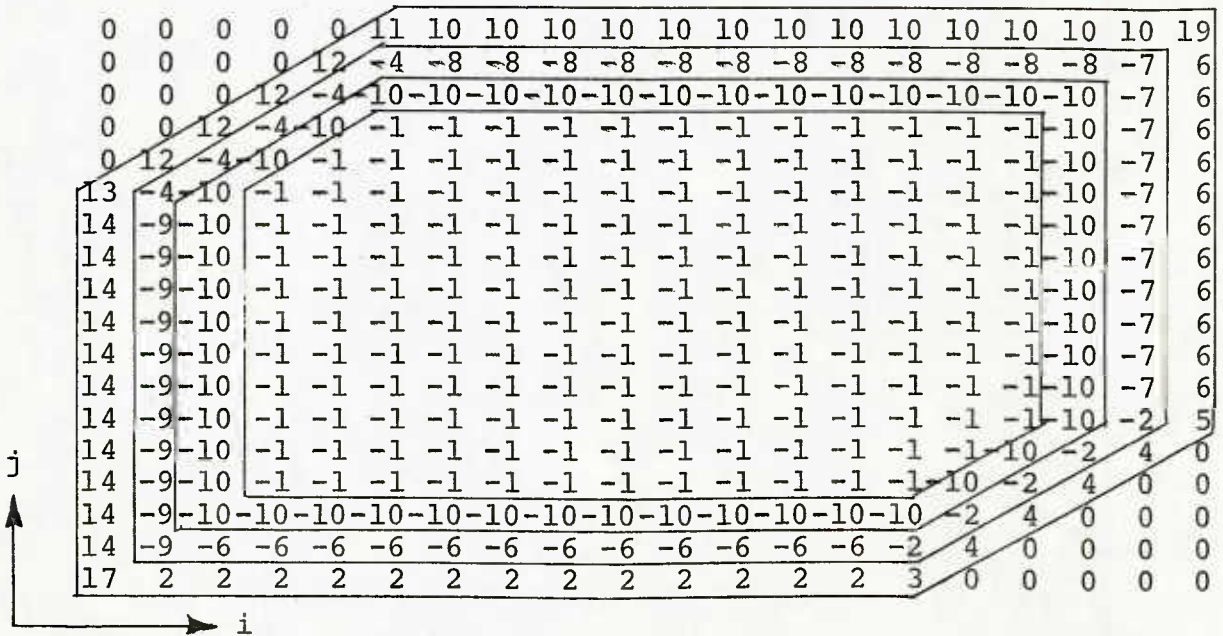
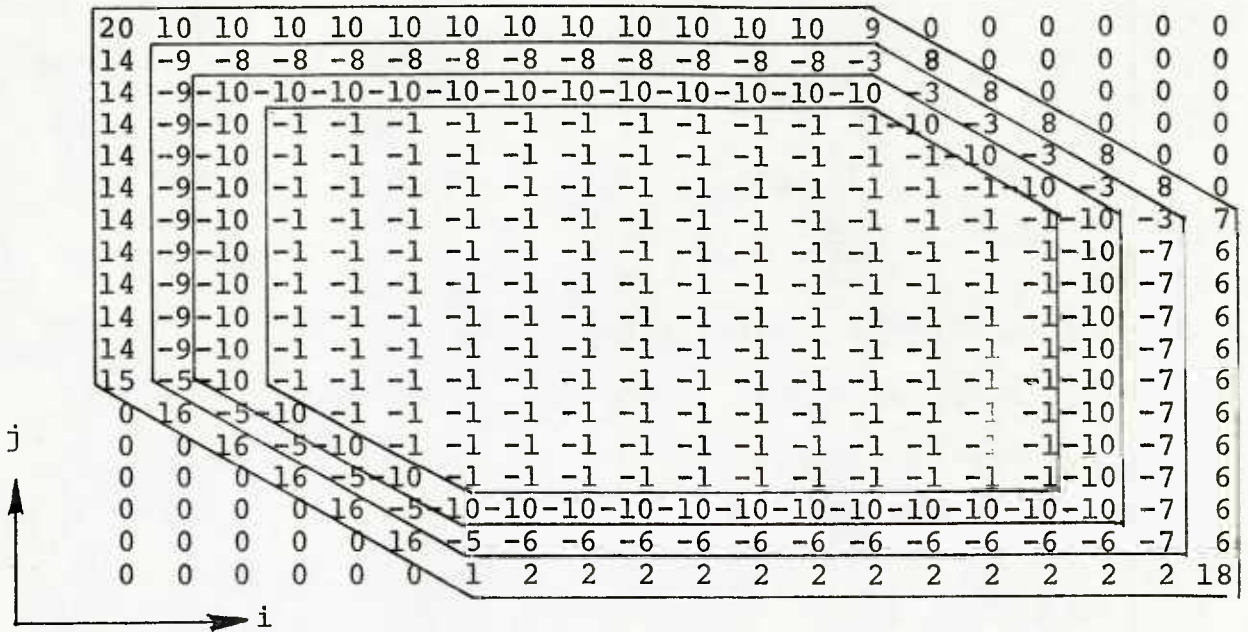
limited  
comp. area

$W_1, W_2, W_3$   
MN

Weight factors  
Field vectors



3. Fields with 2 rectangular points.





Subroutine RANWT (A,ALFA)

This subroutine writes field A (in core) to field ALFA (in secondary storage).

Subroutine RANRD (A,ALFA)

This subroutine reads field ALFA (in secondary storage) into field A (in core).

Subroutine MIXF(FA,FB,MARK,W1,W2,W3,M,N)

This subroutine mixes field FA (corresponding to limited area) with FB (corresponding to a field taken from a large area).

Subroutine HELMSYS (Z1,Z2,FORC1,FORC2,F,MY,FMY,A1,A2,B1,B2,  
ALFA,RESIDUE,IT,M,N)

This subroutine solves the following system of Helmholtz equations by relaxation:

$$\nabla^2(Z1) - A1 \frac{f^2}{\mu}(Z1) + A2 \frac{f^2}{\mu}(Z2) = \text{FORC1}$$

$$\nabla^2(Z2) - B2 \frac{f^2}{\mu}(Z2) + B1 \frac{f^2}{\mu}(Z1) = \text{FORC2}$$

Subroutine  
Parameter

Description

Z1,Z2

Fields to be solved by relaxation.  
First guesses in Z1,Z2 before  
using subroutine.

FORC1,FORC2

Forcing functions.

F

Coriolis parameter field.

MY

$\mu$ -parameter field.

FMY

Working field.

A1,A2,A3,A4

Physical parameters (constants).

ALPHA

$\alpha$ -overrelaxation coefficient.

RESIDUE

Residual, R, in the solution  
of the system.

IT

Number of iterations.

M,N

Field vectors.

Subroutine MAP3P(I1,I2,I3,I4,I5,I6,I7,I8,I9,I10,I11,F1,F2,  
 F3,F4,F5,F6,F7,F8,F9,F10,F,MARK,M,N,IDAY,ITIME,NTIME,PNIVS,  
 PNIVM,PNIV1,ZB1,ZB2,ZB3)

Printing on line-printer of forecast fields in zebra patterns. Heights are computed from stream functions.

<u>Subroutine Parameters</u>	<u>Description</u>
I1-I11	Indicators. I = 0: no printing I = 1: printing. The numbers refer to the following fields: I1: Surface pressure. I2: Height for level $p_m$ . I3: Height for level $p_1^m$ . I4: Thickness ( $p_m - p_s$ ). I5: Lower vertical velocity $\overline{\omega}_1$ . I6: Precipitable water. I7: Accumulated precipitation. I8: Relative humidity. I9: Stream function for level $p_s$ . I10: Stream function for level $p_m$ . I11: Stream function for level $p_1^m$ .
F1-F10	Working fields. (See flow diagram B-2)
F	Coriolis parameter field.
MARK	Indicator field (see MARKF).
M,N	Field dimension.
IDAY	Year, month and day as one integer.
ITIME,NTIME	Initial and forecast time.
PNIVS,PNIVM,PNIV1	Pressure levels in the model.
ZB1,ZB2,ZB3	Working arrays for boundary strings.

Flow Diagram B-2: MAP3P

Variables to and from secondary storage

{PSIM1,PSI11  
{PSI21

{WS,DIV1,DIV2  
{HUM1,PREC

{ZM2,Z12,Z22

F1	F2	F3	F4	F5	F6	F7	F8	F9	F10
$\psi(P_S)$ •	$\psi(P_M)$ •	$\psi(P_L)$ •	$\omega_S$ x $\frac{\omega_L}{\omega_L}$	$D_1$ x	$D_2$ x R.H.	q x •	$\psi_m$ x $\Sigma r$ x •	$\psi_1$ x $Z_1^T$	$\psi_2$ x $Z_2^T$
				$Z(P_S)$ x	$Z(P_M)$ x	$Z(P_L)$ •	$Z_m^T$ •		
		p •	$Z_L$ •						

MAP

MAP

STREAMF

MAP

Subroutine STREAMF (Z,PSI,R,F,MARK,ZB,PSIB,FB,SB,GAMMA,IND,  
IT,M,N)

The subroutine computes the linearized balance equation.

<u>Subroutine Parameter</u>	<u>Description</u>
Z	Z field.
PSI	PSI field.
R	Forcing function for the Poisson equation in the solution of $\nabla^2\psi = F(Z)$ or $\nabla^2Z = Z(\psi)$ .
F	Coriolis parameter.
MARK	MARK-field.
ZB	String of boundary values for $Z_k$ .
PSIB	String of boundary values for $\psi_k$ .
FB	String of boundary values for $f_k$ .
SB	String of boundary values for $(\Delta S)_k$ .
GAMMA	$\gamma$ .
IND	See subroutine comments.
IT	Number of iterations for solving the Poisson equation by relaxation.
M,N	Field vectors.

Subroutines called by STREAMF

BDRGRD(SB)	Computes SB.
BRDVAL(A,AB,M,IND)	Computes boundary values from field A and stores the boundary values in string AB or reverse: IND = 0, AB = A; IND = 1, A = AB.
M	Field parameter
IND	See subroutine comments

POIS is an entry point in HELM for the solution of a Helmholtz equation by Liebmann relaxation.

HELM(Z,FORC,Q,ALFA,RESIDUE,IT,M,N)

<u>Subroutine Parameter</u>	<u>Description</u>
Z	Z-field.
FORC	Forcing function.
Q	Helmholtz coefficient.
ALFA	Overrelaxation coefficient.
RESIDUE	Residual.
M,N	Field vectors.
IT	Counts the number of iterations required to obtain a solution.

Subroutine SATUR(TM, QSAT)

This subroutine computes QSAT, integrated mixing ratio at saturation as a function of the mean temperature between 500 and 1000 mb. TM values of QSAT are given for each whole degree between -50°C to +20°C in the subroutine. Linear interpolation between whole degrees is employed and the result QSAT is given in TON/m<sup>2</sup>/cb.

Subroutine MAP(Q,M,N,DS,QZ,QD,SCALE)

Prints a field in "zebra pattern" on line-printer.

Subroutine Parameters

Description

Q	Field to be printed.
M,N	Field vectors.
DS	Grid distance (meters)
QZ	Isoline corresponding to 000, the line towards 999 on printer (indicated by heavy line below).
QD	Resolution indicator (see below).

000000000 } QD  
000000000 } QD  
                  } QD

111111111  
111111111

222222222  
222222222

SCALE                   Map scale factor.  
                          (unit  $10^6$  m)



Subroutine JACOB(A,B,C,M,N)

This subroutine computes a Jacobian operator (of the form)

$$C = J(A,B) = (A_{i+1} - A_{i-1})_j (B_{j+1} - B_{j-1})_i \\ - (A_{j+1} - A_{j-1})_i (B_{i+1} - B_{i-1})_j .$$

<u>Subroutine parameters</u>	<u>Description</u>
A,B,C	Fields according to formula.
M,N	Field vectors .

Subroutine ABSVOR(PHI,VOR,F,MY,MARK,M,N)

This subroutine computes the relative or the absolute vorticity.

- a.  $\eta = \mu \nabla^2 \psi + f$                       ABSVOR  
b.  $\zeta = \mu \nabla^2 \psi$                               RELVOR (via entry point)

<u>Subroutine Parameters</u>	<u>Description</u>
PHI	$\psi$ field .
VOR	$\eta$ or $\zeta$ field .
F	Coriolis parameter field .
MY	$\mu$ parameter field .
MARK	Field indicator array.
M,N	Field vectors .

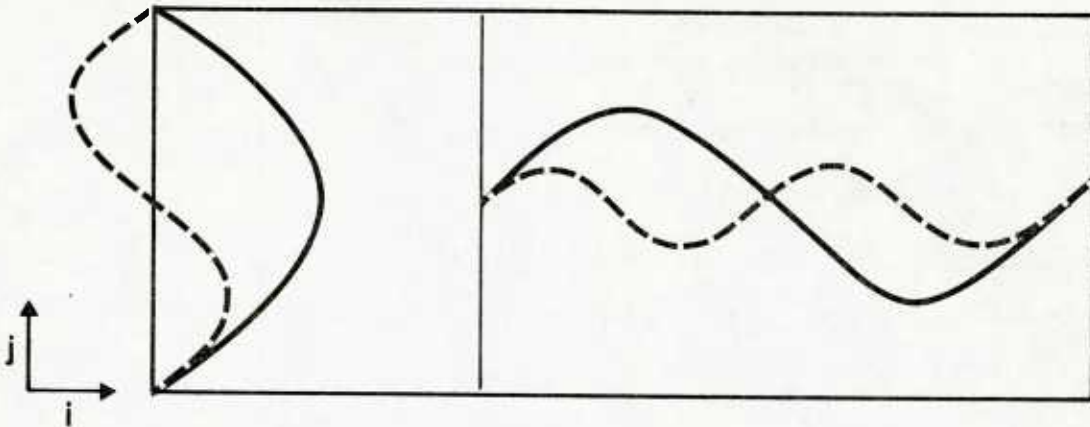
Subroutine GENCH(PSI,M,N,PSIO,U,NU,NWAVE,NX,NY,PSIC,LAMC,  
PSIS,LAMS,WX)

<u>Subroutine Parameters</u>	<u>Description</u>
PSI	Result field.
M,N	Field vectors.
PSIO	Constant value ( $\psi_0$ ).
U	Zonal wind speed.
NU	Resolution of zonal wind speed.
NWAVE	Number of waves.
NX	Wave numbers as a function of channel length in x-direction (nx).
NY	Wave numbers as a function of channel width in y-direction (ny).
PSIC	Amplitudes of the cosine function ( $\psi_c$ ).
LAMC	Phase differences for the cosine functions ( $\psi_c$ ) in whole degrees.
PSIS	Amplitudes of the sine functions ( $\psi_s$ ) in whole degrees.
LAMS	Phase differences for the sine functions ( $\lambda_s$ ) in whole degrees.
WX	Adjustment of the wave near the rigid boundaries: Boundary, WX = 0; +1 row from the boundary, WX = 0.33; +2 row from the boundary, WX = 0.64; +3 and more, WX = 1.

Fields are generated according to the formula:

$$\psi = \psi_0 - Uy + \sum_{v=1}^N [(\psi_c)_v \sin\{(nx)_v(x) - \frac{\pi}{180}(\lambda_c)_v\} \cos(ny)_v] \\ + (\psi_s)_v \sin\{(nx)_v(x) - \frac{\pi}{180}(\lambda_s)_v\} \sin(ny)_v]$$

#### Example Prediction Area



- Nx = 1      1 wave in x-direction over the area (solid line)
- Nx = 2      2 waves in x-direction over the area (dashed line)
- Nx = N      N waves in x-direction over the area
- Ny = 1      1 half wave in y-direction (solid line)
- Ny = 2      2 half waves in y-direction (dashed line)
- Ny = N      N half waves in y-direction

The zonal wind speed is specified with an arbitrary resolution across the channel. All parameters are given by a DATA statement.

Subroutine STEPEXT(F1,F2,F3,F4,F5,F6,F7,F8,F9,F10,MY,F,MARK, M,N,I1,I2,I3,I4)

The subroutine computes the forcing functions

$$\sum_{i=1}^4 F_{mi}, \quad \sum_{i=1}^4 F_{li}, \quad \text{and} \quad \sum_{i=1}^4 F_{2i}$$

and stores the result in the fields HM3, H13 and H23 respectively on secondary memories (see Flow Diagram B-3).

The program is a subroutine to STEP3P.

<u>Subroutine Parameters</u>	<u>Description</u>
F1-F10	Working field.
MY	$\mu$ -parameter field.
F	Coriolis parameter field.
MARK	Field indicator array.
I1,I2,I3,I4	Computational parameters.
	I1 = 0 $W_{\chi} \cdot \nabla \eta = 0$
	I1 = 1 $W_{\chi} \cdot \nabla \eta \neq 0.$
	I2 = 0 $\zeta \cdot \nabla W = 0.$
	I2 = 1 $\zeta \cdot \nabla W \neq 0.$
	I3 = 0 $\omega \frac{\partial \zeta}{\partial p} = 0.$
	I3 = 1 $\omega \frac{\partial \zeta}{\partial p} \neq 0.$
	I4 = 0 $ k \cdot (\frac{\partial V}{\partial p} \times \nabla \omega) = 0.$
	I4 = 1 $ k \cdot (\frac{\partial V}{\partial p} \times \nabla \omega) \neq 0.$
M,N	Field vectors.

Flow Diagram B-3: STEPEXT

F1	F2	F3	F4	F5	F6	F7	F8	F9	F10	To second storage	From second storage	Comments
$\psi_m^T$	$\psi_1^T$	$\psi_2^T$	$W_8$	Heat	-	-	-	$J_4$		J4=F8	F5=DIV1 F6=DIV2	
				$D_1^{-1}$	$D_2^{-1}$							
					$\gamma_{10}$	$\gamma_{11}$	$\gamma_{10}$	$\psi_1^T \gamma_{10}$	$\psi_2^T \gamma_{11}$	H13=F9 H23=F10		Twisting term for thickness fields to second storage
					$\gamma_7$	$\gamma_8$	$\psi_1^T \gamma_7$	$\psi_2^T \gamma_8$				
					$\gamma_9$	$\psi_m^T \gamma_9$				H13=F8		Twisting term for mean fields to second storage.
					$F_m^4$	$\zeta_1$	$\zeta_2$					
$F_{m2}$	$F_{12}$	$F_{22}$										
$F_{m2}^* F_{m3}$	$F_{12}^* F_{13}$	$F_{22}^* F_{23}$								H12=F1 H12=F2 H22=F3		$J_1^* \psi_1^T$ to secondary storage
$\psi_m^2$	$\psi^2 x_1$	$\psi^2 x_2$	$(x_m)^g$	$(x_1)^g$	$(x_2)^g$							Compute forcing function for velocity potential in order to get velocity
			$x_m$	$x_1$	$x_2$					V1=F4 V1=F5 V2=F6		
$\eta_m^2 c$	$\eta_m^2 c_2$											
		$\psi_m^T \psi_1^T$						$\psi_1^T \psi_{13}$				$\gamma_{13}^{m-n-2} \zeta_1$
		$\psi_m^T \psi_2^T$						$F_{11}$		H11=F10		$\psi_1^T \psi_n$ to secondary storage
$\beta_7$	$\beta_8$	$\beta_9$						$\psi_2^T \psi_{14}$				$\gamma_{14}^{m-n-2} \zeta_2$
								$F_{21}$		H21=F10		$\psi_2^T \psi_n$ to secondary storage
$F_{m1}$					$\psi_m^T \psi_7$	$\psi_1^T \psi_8$	$\psi_2^T \psi_9$					
$F_{m2}^* F_{m3}$	$F_{m4}$	$F_{11}$	$F_{12}^* F_{13}$	$F_{14}$	$F_{21}$	$F_{22}^* F_{23}$	$F_{24}$					
$F_{m1}$		$F_{11}$			$F_{21}$					H13=F1 H13=F4 H23=F7		Sum of higher order terms to H13, H14 and R23
$\psi_m^T$	$\psi_1^T$	$\psi_2^T$	$W_8$	Heat	0	0	$J_4$					Restore initial fields

Subroutine VELPOT(KSI, FORC, M, N, RESIDUE, ALFA)

This subroutine computes the velocity potential from a known divergence field

$$\nabla^2 \chi = D.$$

In finite-difference form

$$\nabla^2 \chi = \frac{D}{\mu}.$$

The solution is performed by Liebmann relaxation with an overrelaxation coefficient ALFA equal approximately to 1.4, but its size depends on the area and mesh width. The residual RESIDUE must also be given ( $0.5 \cdot 10^{+6}$  recommended value).

<u>Subroutine Parameters</u>	<u>Description</u>
KSI	2D array for the $\chi$ field.
FORC	2D array for the forcing function.
ALFA	Overrelaxation coefficient.
RESIDUE	Residual.
M, N	Field vectors.

Remark: A first guess must be put in  $\chi$  field before the execution.

Subroutine GRADPR(A,B,C,MARK,M,N)

This subroutine computes a finite difference operation of the form  $\nabla A \cdot \nabla B$ .

$$\begin{aligned} (\nabla A \nabla B)_{ij} = & (A_{i+1} - A_i) (B_{i+1} - B_i)_j + (A_i - A_{i-1}) (B_i - B_{i-1})_j \\ & + (A_{j+1} - A_j) (B_{j+1} - B_j)_i + (A_j - A_{j-1}) (B_j - B_{j-1})_i \end{aligned}$$

<u>Subroutine Parameters</u>	<u>Description</u>
A,B	Fields for operational vector.
C	Result field.
MARK	Field indicator array.
M,N	Field vectors.



APPENDIX C  
PROGRAM LISTINGS

Three programs for the three-parameter model in Fortran IV are presented: PROG3P, STEP3P, and STEPEXT. The other subroutines may be obtained from ENVPREDRSCHFAC by request.

PROGRAM PROG3P

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PROGRAM PROG3P(OUTPUT,TAPE6=OUTPUT,DATA,INPUT=DATA)
* BAROCLINIC BALANCED INTEGRATED 3-PARAMETER MODEL INCLUDING HUMIDITY
* AND PRECIPITATION. BOUNDARY VALUES CAN BE VARIABLE, CONSTANT OR OF
* CHANNEL TYPE.
  DIMENSION F1(57,57),F2(57,57),F3(57,57),F4(57,57),F5(57,57),
  1 F6(57,57),F7(57,57),F8(57,57),F9(57,57),F10(57,57),
  2 F(57,57),MY(57,57),MARK(57,57)
  DIMENSION UPS(10),UPM(10),UP1(10),WX(15),NX(20),NY(20),PSIC(20),
  X PSIS(20),LAMC(20),LAMS(20)
  DIMENSION ZB(250),PSIB(250),FB(250),SB(250),IT(2,25)
  DIMENSION MAPHOUR(10),LETACC(10),NSMUTT(10),NELLIPT(10)
  DIMENSION MSMUTT(10)
  DIMENSION MELLIPT(10)
COMMON/FPAR/IC(8),JC(8),XPOL,YPOL,R,RE,DS,JMIN(100),JMAX(100),
X M,N,KIND
COMMON/FORM1/IC1(8),JC1(8),JMIN1(100),JMAX1(100),XP1,YP1,D1
COMMON/DRM/MN,NDIM,NFLD
COMMON/KANAL/FIM
COMMON F
COMMON/ECS/ PSIM1,PSI11,PSI21,PSIM2,PSI12,PSI22,HUM1,HUM2,DIV1,
2 DIV2,WS,HEAT,J789,J12,J56,J3,PS,TS,PREC,STRM,STR1,STR2,ZM1,Z11,Z21
3,ZM2,Z12,Z22,H13,H23,HM3,HM2,H12,H22,H11,H21,J4,VM,V1,V2
COMMON/COEFF/A1,A2,A3,B1,B2,B3,C1,C2,C3,C4,C5,C6,C7,C8,D,DELP,EM,
X E1,E2,H1,H2,H3,H4,H5,H6,PMEAN,S1,S2,T1,T2,T3,T4,T5,
X P0,PM,P1
COMMON/COEFF2/T6,T7,T8,T9,T10,T11,T12,T13,T14,
X K1,K2,K3,K4,K5,K6,K7,K8,K9,K10,K11,K12,K13,K14,K15,
X K16,K17,K18,K19,K20,K21,K22,K23,K24,K25,K26,K27,K28,
X K29,K30,K31,K32,K33,K34,K35,K36,K37,K38
COMMON/RUNPAR/DELT,NTSTEP,ALFASYS,ALFAM,ALFAZ,ALFAPSI,RESSYS,RESM,
X RESZ,RESPSI,Q,FOCEAN,FCONT,WGT1,WGT2,WGT3,ADIFF
REAL MY
REAL K1,K2,K3,K4,K5,K6,K7,K8,K9,K10,K11,K12,K13,K14,K15,
X K16,K17,K18,K19,K20,K21,K22,K23,K24,K25,K26,K27,K28,
X K29,K30,K31,K32,K33,K34,K35,K36,K37,K38
DATA PSIM1,PSI11,PSI21,PSIM2,PSI12,PSI22,HUM1,HUM2,DIV1,
*DIV2,WS,HEAT,J789,J12,J56,J3,PS,TS,PREC,STRM,STR1,STR2,ZM1,Z11,Z21
*,ZM2,Z12,Z22,H13,H23,HM3,HM2,H12,H22,H11,H21,J4,VM,V1,V2
*/1,3250,6499,9748,12997,16246,19495,22744,25993,29242,32491,35740,
*38989,42238,45487,48736,51985,55234,58483,61732,64981,68230,
*71479,74728,77977,81226,84475,87724,90973,94222,97471,100720,
*103969,107218,110467,113716,116965,120214,123463,126712/
NAMELIST/FPARR/IC,JC,XPOL,YPOL,R,RE,DS,JMIN,JMAX,M,N,KIND
DATA(IC1(I),I=1,8)/1,1,57,57,57,57,1,1/
DATA(JC1(I),I=1,8)/1,1,1,1,57,57,57,57/
DATA XP1,YP1,D1 /-21,,29,,95.25/
DATA FIM/ 50. /
DATA KIND/0/
DATA ICASE,IDAY,ITIME/1,660922,00/
DATA ALFASYS,ALFAM,ALFAZ,ALFAPSI/,.85,1.45,1,4,1,8/
DATA RESSYS,RESM,RESZ,RESPSI/,.5E5,.5, .5E5 /
DATA Q,FOCEAN,FCONT,WGT1,WGT2,WGT3,ADIFF/ 0,0 ,.62,1.,1.,.5.,.2,
* 0,0/
DATA STAB1,STAB2,PNIVS,PNIVM,PNIV1/,42222,,51111,100,,50,,30./
C NTSTEP IS NUMBER OF TIME STEPS IN 3 HOURS
DATA NTSTEP/4/
DATA NEND/48/
DATA IVAR/ 0 /

```

PROG3P.2  
 PROG3P.3  
 PROG3P.4  
 PROG3P.5  
 PROG3P.6  
 PROG3P.7  
 PROG3P.8  
 PROG3P.9  
 PROG3P.10  
 PROG3P.11  
 PROG3P.12  
 JUN12,2  
 JUN12,3  
 PROG3P.13  
 PROG3P.14  
 PROG3P.15  
 PROG3P.16  
 PROG3P.17  
 PROG3P.18  
 PROG3P.19  
 APR14,2  
 PROG3P.20  
 PROG3P.30  
 PROG3P.31  
 PROG3P.32  
 PROG3P.33  
 PROG3P.34  
 PROG3P.35  
 PROG3P.36  
 PROG3P.37  
 PROG3P.38  
 PROG3P.39  
 PROG3P.40  
 PROG3P.41  
 PROG3P.42  
 PROG3P.43  
 PROG3P.44  
 APR14,3  
 APR14,4  
 APR14,5  
 APR14,6  
 APR14,7  
 APR14,8  
 APR14,9  
 PROG3P.45  
 PROG3P.46  
 PROG3P.47  
 PROG3P.48  
 PROG3P.49  
 PROG3P.50  
 PROG3P.51  
 JUN12,4  
 PROG3P.53  
 PROG3P.54  
 PROG3P.55  
 PROG3P.56  
 APR14,10  
 JUN12,5  
 APR14,12  
 PROG3P.59

PROG3P (Continued)

C Q AND ADIFF HAVE TO BE DEFINED ***	PROG3P.60
C MAPHOUR GIVES THE HOURS(MULTIPLE OF 6) FOR MAP=PRINTING.	PROG3P.61
C LETACC GIVES THE HOURS(MULTIPLE OF 6) AT WHICH THE ACCUMULATED PRECIPITATION IS PRINTED.	PROG3P.62
C SHALL BE PUT =0 AGAIN.	PROG3P.63
C NSMUTT GIVES THE HOURS(MULTIPLE OF 6) FOR SMOOTHING.	PROG3P.64
C NELLIPT GIVES THE HOURS (MULTIPLE OF 6) FOR ELLIPTICITY TEST.	PROG3P.65
DATA (MAPHOUR/0,12,24,36,48,-1,-1,-1,-1/	JUN12,6
DATA (LETACC(I),I=1,10)/0,12,24,36,48,-1,-1,-1,-1/	JUN12,7
DATA(NSMUTT(I),I=1,10)/0,3,6,9,12,15,18,21,24,-1/	PROG3P.6A
DATA (MSMUTT(I),I=1,10)/27,30,33,36,39,42,45,48,-1,-1/	JUN12,8
DATA(NELLIPT(I),I=1,10)/0,3,6,9,12,15,18,21,24,-1/	PROG3P.69
DATA (NELLIPT(I),I=1,10)/27,30,33,36,39,42,45,48,-1,-1/	JUN12,9
3=9,806	PROG3P.70
F0=1.03E-4	PROG3P.71
C TIME STEP IN SECONDS	APR14,14
DELT=1.*3.6E3/NTSTEP	JUN12,10
LABEL = 10H PUT1	PROG3P.73
BTIME = SECOND(FAKE)	PROG3P.74
WRITE(6,8000) BTIME	PROG3P.75
8000 FORMAT(1H0, *BTIME=*, F10,4)	PROG3P.76
CALL PUT1	PROG3P.77
DTIME = SECOND(FAKE) - BTIME	PROG3P.78
WRITE(6,8005) LABEL, DTIME	PROG3P.79
8005 FORMAT(1H0, *TIME TO EXECUTE*, A10, F10,4)	PROG3P.80
LABEL = 10H COEFF3P	PROG3P.81
BTIME = SECOND(FAKE)	PROG3P.82
WRITE(6,8000) BTIME	PROG3P.83
CALL COEFF3P(STAB1,STAR2,PNIVS,PNIVM,PNIV1)	PROG3P.84
DTIME = SECOND(FAKE) - BTIME	PROG3P.85
WRITE(6,8005) LABEL, DTIME	PROG3P.86
MN = M*N	PROG3P.87
NDIM = 20000	PROG3P.88
LABEL = 10H MARKF	PROG3P.89
BTIME = SECOND(FAKE)	PROG3P.90
WRITE(6,8000) BTIME	PROG3P.91
CALL MARKF(MARK,M,N)	PROG3P.92
DTIME = SECOND(FAKE) - BTIME	PROG3P.93
WRITE(6,8005) LABEL, DTIME	PROG3P.94
LABEL = 10H MYFF	PROG3P.95
BTIME = SECOND(FAKE)	PROG3P.96
WRITE(6,8000) BTIME	PROG3P.97
CALL MYFF(MY,M,N)	APR14,16
DTIME = SECOND(FAKE) - BTIME	PROG3P.98
WRITE(6,8005) LABEL, DTIME	PROG3P.100
DO 9 I=1,MN	PROG3P.101
9 F1(I)=0.0	PROG3P.102
LABEL = 10H RANWT	PROG3P.103
BTIME = SECOND(FAKE)	PROG3P.104
WRITE(6,8000) BTIME	PROG3P.105
CALL RANWT(VM,F1)	PROG3P.106
CALL RANWT(V1,F1)	PROG3P.107
CALL RANWT(V2,F1)	PROG3P.108
DTIME = SECOND(FAKE) - BTIME	PROG3P.109
WRITE(6,8005) LABEL, DTIME	PROG3P.110
VTIME = 0	PROG3P.111
IF(KIND,NE,0) GO TO 45	PROG3P.112
C INITIAL FIELDS, NO CHANNEL.	PROG3P.113
C THE SUBROUTINE INITF HAS TO BE WRITTEN ***	PROG3P.114
C INITIAL Z-FIELDS AT THE LEVELS PS,PM AND P1 TO F1,F2 AND F3, THE HUMIDITY F4, F5 AND F6, THE WIND SPEED F7, F8 AND F9, THE SURFACE PRESSURE F10, F11 AND F12, THE SEA LEVEL TEMPERATURE F13, F14 AND F15, THE SURFACE WIND DIRECTION F16, F17 AND F18, THE SURFACE WIND SPEED F19, F20 AND F21, THE SURFACE WIND DIRECTION F22, F23 AND F24, THE SURFACE WIND SPEED F25, F26 AND F27, THE SURFACE WIND DIRECTION F28, F29 AND F30, THE SURFACE WIND SPEED F31, F32 AND F33, THE SURFACE WIND DIRECTION F34, F35 AND F36, THE SURFACE WIND SPEED F37, F38 AND F39, THE SURFACE WIND DIRECTION F40, F41 AND F42, THE SURFACE WIND SPEED F43, F44 AND F45, THE SURFACE WIND DIRECTION F46, F47 AND F48, THE SURFACE WIND SPEED F49, F50 AND F51, THE SURFACE WIND DIRECTION F52, F53 AND F54, THE SURFACE WIND SPEED F55, F56 AND F57, THE SURFACE WIND DIRECTION F58, F59 AND F60, THE SURFACE WIND SPEED F61, F62 AND F63, THE SURFACE WIND DIRECTION F64, F65 AND F66, THE SURFACE WIND SPEED F67, F68 AND F69, THE SURFACE WIND DIRECTION F70, F71 AND F72, THE SURFACE WIND SPEED F73, F74 AND F75, THE SURFACE WIND DIRECTION F76, F77 AND F78, THE SURFACE WIND SPEED F79, F80 AND F81, THE SURFACE WIND DIRECTION F82, F83 AND F84, THE SURFACE WIND SPEED F85, F86 AND F87, THE SURFACE WIND DIRECTION F88, F89 AND F90, THE SURFACE WIND SPEED F91, F92 AND F93, THE SURFACE WIND DIRECTION F94, F95 AND F96, THE SURFACE WIND SPEED F97, F98 AND F99, THE SURFACE WIND DIRECTION F100, F101 AND F102, THE SURFACE WIND SPEED F103, F104 AND F105, THE SURFACE WIND DIRECTION F106, F107 AND F108, THE SURFACE WIND SPEED F109, F110 AND F111, THE SURFACE WIND DIRECTION F112, F113 AND F114, THE SURFACE WIND SPEED F115, F116 AND F117, THE SURFACE WIND DIRECTION F118, F119 AND F120, THE SURFACE WIND SPEED F121, F122 AND F123, THE SURFACE WIND DIRECTION F124, F125 AND F126, THE SURFACE WIND SPEED F127, F128 AND F129, THE SURFACE WIND DIRECTION F130, F131 AND F132, THE SURFACE WIND SPEED F133, F134 AND F135, THE SURFACE WIND DIRECTION F136, F137 AND F138, THE SURFACE WIND SPEED F139, F140 AND F141, THE SURFACE WIND DIRECTION F142, F143 AND F144, THE SURFACE WIND SPEED F145, F146 AND F147, THE SURFACE WIND DIRECTION F148, F149 AND F150, THE SURFACE WIND SPEED F151, F152 AND F153, THE SURFACE WIND DIRECTION F154, F155 AND F156, THE SURFACE WIND SPEED F157, F158 AND F159, THE SURFACE WIND DIRECTION F160, F161 AND F162, THE SURFACE WIND SPEED F163, F164 AND F165, THE SURFACE WIND DIRECTION F166, F167 AND F168, THE SURFACE WIND SPEED F169, F170 AND F171, THE SURFACE WIND DIRECTION F172, F173 AND F174, THE SURFACE WIND SPEED F175, F176 AND F177, THE SURFACE WIND DIRECTION F178, F179 AND F180, THE SURFACE WIND SPEED F181, F182 AND F183, THE SURFACE WIND DIRECTION F184, F185 AND F186, THE SURFACE WIND SPEED F187, F188 AND F189, THE SURFACE WIND DIRECTION F190, F191 AND F192, THE SURFACE WIND SPEED F193, F194 AND F195, THE SURFACE WIND DIRECTION F196, F197 AND F198, THE SURFACE WIND SPEED F199, F200 AND F201, THE SURFACE WIND DIRECTION F202, F203 AND F204, THE SURFACE WIND SPEED F205, F206 AND F207, THE SURFACE WIND DIRECTION F208, F209 AND F210, THE SURFACE WIND SPEED F211, F212 AND F213, THE SURFACE WIND DIRECTION F214, F215 AND F216, THE SURFACE WIND SPEED F217, F218 AND F219, THE SURFACE WIND DIRECTION F220, F221 AND F222, THE SURFACE WIND SPEED F223, F224 AND F225, THE SURFACE WIND DIRECTION F226, F227 AND F228, THE SURFACE WIND SPEED F229, F230 AND F231, THE SURFACE WIND DIRECTION F232, F233 AND F234, THE SURFACE WIND SPEED F235, F236 AND F237, THE SURFACE WIND DIRECTION F238, F239 AND F240, THE SURFACE WIND SPEED F241, F242 AND F243, THE SURFACE WIND DIRECTION F244, F245 AND F246, THE SURFACE WIND SPEED F247, F248 AND F249, THE SURFACE WIND DIRECTION F250, F251 AND F252, THE SURFACE WIND SPEED F253, F254 AND F255, THE SURFACE WIND DIRECTION F256, F257 AND F258, THE SURFACE WIND SPEED F259, F260 AND F261, THE SURFACE WIND DIRECTION F262, F263 AND F264, THE SURFACE WIND SPEED F265, F266 AND F267, THE SURFACE WIND DIRECTION F268, F269 AND F270, THE SURFACE WIND SPEED F271, F272 AND F273, THE SURFACE WIND DIRECTION F274, F275 AND F276, THE SURFACE WIND SPEED F277, F278 AND F279, THE SURFACE WIND DIRECTION F280, F281 AND F282, THE SURFACE WIND SPEED F283, F284 AND F285, THE SURFACE WIND DIRECTION F286, F287 AND F288, THE SURFACE WIND SPEED F289, F290 AND F291, THE SURFACE WIND DIRECTION F292, F293 AND F294, THE SURFACE WIND SPEED F295, F296 AND F297, THE SURFACE WIND DIRECTION F298, F299 AND F300, THE SURFACE WIND SPEED F301, F302 AND F303, THE SURFACE WIND DIRECTION F304, F305 AND F306, THE SURFACE WIND SPEED F307, F308 AND F309, THE SURFACE WIND DIRECTION F310, F311 AND F312, THE SURFACE WIND SPEED F313, F314 AND F315, THE SURFACE WIND DIRECTION F316, F317 AND F318, THE SURFACE WIND SPEED F319, F320 AND F321, THE SURFACE WIND DIRECTION F322, F323 AND F324, THE SURFACE WIND SPEED F325, F326 AND F327, THE SURFACE WIND DIRECTION F328, F329 AND F330, THE SURFACE WIND SPEED F331, F332 AND F333, THE SURFACE WIND DIRECTION F334, F335 AND F336, THE SURFACE WIND SPEED F337, F338 AND F339, THE SURFACE WIND DIRECTION F340, F341 AND F342, THE SURFACE WIND SPEED F343, F344 AND F345, THE SURFACE WIND DIRECTION F346, F347 AND F348, THE SURFACE WIND SPEED F349, F350 AND F351, THE SURFACE WIND DIRECTION F352, F353 AND F354, THE SURFACE WIND SPEED F355, F356 AND F357, THE SURFACE WIND DIRECTION F358, F359 AND F360, THE SURFACE WIND SPEED F361, F362 AND F363, THE SURFACE WIND DIRECTION F364, F365 AND F366, THE SURFACE WIND SPEED F367, F368 AND F369, THE SURFACE WIND DIRECTION F370, F371 AND F372, THE SURFACE WIND SPEED F373, F374 AND F375, THE SURFACE WIND DIRECTION F376, F377 AND F378, THE SURFACE WIND SPEED F379, F380 AND F381, THE SURFACE WIND DIRECTION F382, F383 AND F384, THE SURFACE WIND SPEED F385, F386 AND F387, THE SURFACE WIND DIRECTION F388, F389 AND F390, THE SURFACE WIND SPEED F391, F392 AND F393, THE SURFACE WIND DIRECTION F394, F395 AND F396, THE SURFACE WIND SPEED F397, F398 AND F399, THE SURFACE WIND DIRECTION F400, F401 AND F402, THE SURFACE WIND SPEED F403, F404 AND F405, THE SURFACE WIND DIRECTION F406, F407 AND F408, THE SURFACE WIND SPEED F409, F410 AND F411, THE SURFACE WIND DIRECTION F412, F413 AND F414, THE SURFACE WIND SPEED F415, F416 AND F417, THE SURFACE WIND DIRECTION F418, F419 AND F420, THE SURFACE WIND SPEED F421, F422 AND F423, THE SURFACE WIND DIRECTION F424, F425 AND F426, THE SURFACE WIND SPEED F427, F428 AND F429, THE SURFACE WIND DIRECTION F430, F431 AND F432, THE SURFACE WIND SPEED F433, F434 AND F435, THE SURFACE WIND DIRECTION F436, F437 AND F438, THE SURFACE WIND SPEED F439, F440 AND F441, THE SURFACE WIND DIRECTION F442, F443 AND F444, THE SURFACE WIND SPEED F445, F446 AND F447, THE SURFACE WIND DIRECTION F448, F449 AND F450, THE SURFACE WIND SPEED F451, F452 AND F453, THE SURFACE WIND DIRECTION F454, F455 AND F456, THE SURFACE WIND SPEED F457, F458 AND F459, THE SURFACE WIND DIRECTION F460, F461 AND F462, THE SURFACE WIND SPEED F463, F464 AND F465, THE SURFACE WIND DIRECTION F466, F467 AND F468, THE SURFACE WIND SPEED F469, F470 AND F471, THE SURFACE WIND DIRECTION F472, F473 AND F474, THE SURFACE WIND SPEED F475, F476 AND F477, THE SURFACE WIND DIRECTION F478, F479 AND F480, THE SURFACE WIND SPEED F481, F482 AND F483, THE SURFACE WIND DIRECTION F484, F485 AND F486, THE SURFACE WIND SPEED F487, F488 AND F489, THE SURFACE WIND DIRECTION F490, F491 AND F492, THE SURFACE WIND SPEED F493, F494 AND F495, THE SURFACE WIND DIRECTION F496, F497 AND F498, THE SURFACE WIND SPEED F499, F500 AND F501, THE SURFACE WIND DIRECTION F502, F503 AND F504, THE SURFACE WIND SPEED F505, F506 AND F507, THE SURFACE WIND DIRECTION F508, F509 AND F510, THE SURFACE WIND SPEED F511, F512 AND F513, THE SURFACE WIND DIRECTION F514, F515 AND F516, THE SURFACE WIND SPEED F517, F518 AND F519, THE SURFACE WIND DIRECTION F520, F521 AND F522, THE SURFACE WIND SPEED F523, F524 AND F525, THE SURFACE WIND DIRECTION F526, F527 AND F528, THE SURFACE WIND SPEED F529, F530 AND F531, THE SURFACE WIND DIRECTION F532, F533 AND F534, THE SURFACE WIND SPEED F535, F536 AND F537, THE SURFACE WIND DIRECTION F538, F539 AND F540, THE SURFACE WIND SPEED F541, F542 AND F543, THE SURFACE WIND DIRECTION F544, F545 AND F546, THE SURFACE WIND SPEED F547, F548 AND F549, THE SURFACE WIND DIRECTION F550, F551 AND F552, THE SURFACE WIND SPEED F553, F554 AND F555, THE SURFACE WIND DIRECTION F556, F557 AND F558, THE SURFACE WIND SPEED F559, F560 AND F561, THE SURFACE WIND DIRECTION F562, F563 AND F564, THE SURFACE WIND SPEED F565, F566 AND F567, THE SURFACE WIND DIRECTION F568, F569 AND F570, THE SURFACE WIND SPEED F571, F572 AND F573, THE SURFACE WIND DIRECTION F574, F575 AND F576, THE SURFACE WIND SPEED F577, F578 AND F579, THE SURFACE WIND DIRECTION F580, F581 AND F582, THE SURFACE WIND SPEED F583, F584 AND F585, THE SURFACE WIND DIRECTION F586, F587 AND F588, THE SURFACE WIND SPEED F589, F590 AND F591, THE SURFACE WIND DIRECTION F592, F593 AND F594, THE SURFACE WIND SPEED F595, F596 AND F597, THE SURFACE WIND DIRECTION F598, F599 AND F600, THE SURFACE WIND SPEED F601, F602 AND F603, THE SURFACE WIND DIRECTION F604, F605 AND F606, THE SURFACE WIND SPEED F607, F608 AND F609, THE SURFACE WIND DIRECTION F610, F611 AND F612, THE SURFACE WIND SPEED F613, F614 AND F615, THE SURFACE WIND DIRECTION F616, F617 AND F618, THE SURFACE WIND SPEED F619, F620 AND F621, THE SURFACE WIND DIRECTION F622, F623 AND F624, THE SURFACE WIND SPEED F625, F626 AND F627, THE SURFACE WIND DIRECTION F628, F629 AND F630, THE SURFACE WIND SPEED F631, F632 AND F633, THE SURFACE WIND DIRECTION F634, F635 AND F636, THE SURFACE WIND SPEED F637, F638 AND F639, THE SURFACE WIND DIRECTION F640, F641 AND F642, THE SURFACE WIND SPEED F643, F644 AND F645, THE SURFACE WIND DIRECTION F646, F647 AND F648, THE SURFACE WIND SPEED F649, F650 AND F651, THE SURFACE WIND DIRECTION F652, F653 AND F654, THE SURFACE WIND SPEED F655, F656 AND F657, THE SURFACE WIND DIRECTION F658, F659 AND F660, THE SURFACE WIND SPEED F661, F662 AND F663, THE SURFACE WIND DIRECTION F664, F665 AND F666, THE SURFACE WIND SPEED F667, F668 AND F669, THE SURFACE WIND DIRECTION F670, F671 AND F672, THE SURFACE WIND SPEED F673, F674 AND F675, THE SURFACE WIND DIRECTION F676, F677 AND F678, THE SURFACE WIND SPEED F679, F680 AND F681, THE SURFACE WIND DIRECTION F682, F683 AND F684, THE SURFACE WIND SPEED F685, F686 AND F687, THE SURFACE WIND DIRECTION F688, F689 AND F690, THE SURFACE WIND SPEED F691, F692 AND F693, THE SURFACE WIND DIRECTION F694, F695 AND F696, THE SURFACE WIND SPEED F697, F698 AND F699, THE SURFACE WIND DIRECTION F700, F701 AND F702, THE SURFACE WIND SPEED F703, F704 AND F705, THE SURFACE WIND DIRECTION F706, F707 AND F708, THE SURFACE WIND SPEED F709, F710 AND F711, THE SURFACE WIND DIRECTION F712, F713 AND F714, THE SURFACE WIND SPEED F715, F716 AND F717, THE SURFACE WIND DIRECTION F718, F719 AND F720, THE SURFACE WIND SPEED F721, F722 AND F723, THE SURFACE WIND DIRECTION F724, F725 AND F726, THE SURFACE WIND SPEED F727, F728 AND F729, THE SURFACE WIND DIRECTION F730, F731 AND F732, THE SURFACE WIND SPEED F733, F734 AND F735, THE SURFACE WIND DIRECTION F736, F737 AND F738, THE SURFACE WIND SPEED F739, F740 AND F741, THE SURFACE WIND DIRECTION F742, F743 AND F744, THE SURFACE WIND SPEED F745, F746 AND F747, THE SURFACE WIND DIRECTION F748, F749 AND F750, THE SURFACE WIND SPEED F751, F752 AND F753, THE SURFACE WIND DIRECTION F754, F755 AND F756, THE SURFACE WIND SPEED F757, F758 AND F759, THE SURFACE WIND DIRECTION F760, F761 AND F762, THE SURFACE WIND SPEED F763, F764 AND F765, THE SURFACE WIND DIRECTION F766, F767 AND F768, THE SURFACE WIND SPEED F769, F770 AND F771, THE SURFACE WIND DIRECTION F772, F773 AND F774, THE SURFACE WIND SPEED F775, F776 AND F777, THE SURFACE WIND DIRECTION F778, F779 AND F780, THE SURFACE WIND SPEED F781, F782 AND F783, THE SURFACE WIND DIRECTION F784, F785 AND F786, THE SURFACE WIND SPEED F787, F788 AND F789, THE SURFACE WIND DIRECTION F790, F791 AND F792, THE SURFACE WIND SPEED F793, F794 AND F795, THE SURFACE WIND DIRECTION F796, F797 AND F798, THE SURFACE WIND SPEED F799, F800 AND F801, THE SURFACE WIND DIRECTION F802, F803 AND F804, THE SURFACE WIND SPEED F805, F806 AND F807, THE SURFACE WIND DIRECTION F808, F809 AND F810, THE SURFACE WIND SPEED F811, F812 AND F813, THE SURFACE WIND DIRECTION F814, F815 AND F816, THE SURFACE WIND SPEED F817, F818 AND F819, THE SURFACE WIND DIRECTION F820, F821 AND F822, THE SURFACE WIND SPEED F823, F824 AND F825, THE SURFACE WIND DIRECTION F826, F827 AND F828, THE SURFACE WIND SPEED F829, F830 AND F831, THE SURFACE WIND DIRECTION F832, F833 AND F834, THE SURFACE WIND SPEED F835, F836 AND F837, THE SURFACE WIND DIRECTION F838, F839 AND F840, THE SURFACE WIND SPEED F841, F842 AND F843, THE SURFACE WIND DIRECTION F844, F845 AND F846, THE SURFACE WIND SPEED F847, F848 AND F849, THE SURFACE WIND DIRECTION F850, F851 AND F852, THE SURFACE WIND SPEED F853, F854 AND F855, THE SURFACE WIND DIRECTION F856, F857 AND F858, THE SURFACE WIND SPEED F859, F860 AND F861, THE SURFACE WIND DIRECTION F862, F863 AND F864, THE SURFACE WIND SPEED F865, F866 AND F867, THE SURFACE WIND DIRECTION F868, F869 AND F870, THE SURFACE WIND SPEED F871, F872 AND F873, THE SURFACE WIND DIRECTION F874, F875 AND F876, THE SURFACE WIND SPEED F877, F878 AND F879, THE SURFACE WIND DIRECTION F880, F881 AND F882, THE SURFACE WIND SPEED F883, F884 AND F885, THE SURFACE WIND DIRECTION F886, F887 AND F888, THE SURFACE WIND SPEED F889, F890 AND F891, THE SURFACE WIND DIRECTION F892, F893 AND F894, THE SURFACE WIND SPEED F895, F896 AND F897, THE SURFACE WIND DIRECTION F898, F899 AND F900, THE SURFACE WIND SPEED F901, F902 AND F903, THE SURFACE WIND DIRECTION F904, F905 AND F906, THE SURFACE WIND SPEED F907, F908 AND F909, THE SURFACE WIND DIRECTION F910, F911 AND F912, THE SURFACE WIND SPEED F913, F914 AND F915, THE SURFACE WIND DIRECTION F916, F917 AND F918, THE SURFACE WIND SPEED F919, F920 AND F921, THE SURFACE WIND DIRECTION F922, F923 AND F924, THE SURFACE WIND SPEED F925, F926 AND F927, THE SURFACE WIND DIRECTION F928, F929 AND F930, THE SURFACE WIND SPEED F931, F932 AND F933, THE SURFACE WIND DIRECTION F934, F935 AND F936, THE SURFACE WIND SPEED F937, F938 AND F939, THE SURFACE WIND DIRECTION F940, F941 AND F942, THE SURFACE WIND SPEED F943, F944 AND F945, THE SURFACE WIND DIRECTION F946, F947 AND F948, THE SURFACE WIND SPEED F949, F950 AND F951, THE SURFACE WIND DIRECTION F952, F953 AND F954, THE SURFACE WIND SPEED F955, F956 AND F957, THE SURFACE WIND DIRECTION F958, F959 AND F960, THE SURFACE WIND SPEED F961, F962 AND F963, THE SURFACE WIND DIRECTION F964, F965 AND F966, THE SURFACE WIND SPEED F967, F968 AND F969, THE SURFACE WIND DIRECTION F970, F971 AND F972, THE SURFACE WIND SPEED F973, F974 AND F975, THE SURFACE WIND DIRECTION F976, F977 AND F978, THE SURFACE WIND SPEED F979, F980 AND F981, THE SURFACE WIND DIRECTION F982, F983 AND F984, THE SURFACE WIND SPEED F985, F986 AND F987, THE SURFACE WIND DIRECTION F988, F989 AND F990, THE SURFACE WIND SPEED F991, F992 AND F993, THE SURFACE WIND DIRECTION F994, F995 AND F996, THE SURFACE WIND SPEED F997, F998 AND F999, THE SURFACE WIND DIRECTION F1000, F1001 AND F1002, THE SURFACE WIND SPEED F1003, F1004 AND F1005, THE SURFACE WIND DIRECTION F1006, F1007 AND F1008, THE SURFACE WIND SPEED F1009, F1010 AND F1011, THE SURFACE WIND DIRECTION F1012, F1013 AND F1014, THE SURFACE WIND SPEED F1015, F1016 AND F1017, THE SURFACE WIND DIRECTION F1018, F1019 AND F1020, THE SURFACE WIND SPEED F1021, F1022 AND F1023, THE SURFACE WIND DIRECTION F1024, F1025 AND F1026, THE SURFACE WIND SPEED F1027, F1028 AND F1029, THE SURFACE WIND DIRECTION F1030, F1031 AND F1032, THE SURFACE WIND SPEED F1033, F1034 AND F1035, THE SURFACE WIND DIRECTION F1036, F1037 AND F1038, THE SURFACE WIND SPEED F1039, F1040 AND F1041, THE SURFACE WIND DIRECTION F1042, F1043 AND F1044, THE SURFACE WIND SPEED F1045, F1046 AND F1047, THE SURFACE WIND DIRECTION F1048, F1049 AND F1050, THE SURFACE WIND SPEED F1051, F1052 AND F1053, THE SURFACE WIND DIRECTION F1054, F1055 AND F1056, THE SURFACE WIND SPEED F1057, F1058 AND F1059, THE SURFACE WIND DIRECTION F1060, F1061 AND F1062, THE SURFACE WIND SPEED F1063, F1064 AND F1065, THE SURFACE WIND DIRECTION F1066, F1067 AND F1068, THE SURFACE WIND SPEED F1069, F1070 AND F1071, THE SURFACE WIND DIRECTION F1072, F1073 AND F1074, THE SURFACE WIND SPEED F1075, F1076 AND F1077, THE SURFACE WIND DIRECTION F1078, F1079 AND F1080, THE SURFACE WIND SPEED F1081, F1082 AND F1083, THE SURFACE WIND DIRECTION F1084, F1085 AND F1086, THE SURFACE WIND SPEED F1087, F1088 AND F1089, THE SURFACE WIND DIRECTION F1090, F1091 AND F1092, THE SURFACE WIND SPEED F1093, F1094 AND F1095, THE SURFACE WIND DIRECTION F1096, F1097 AND F1098, THE SURFACE WIND SPEED F1099, F1100 AND F1101, THE SURFACE WIND DIRECTION F1102, F1103 AND F1104, THE SURFACE WIND SPEED F1105, F1106 AND F1107, THE SURFACE WIND DIRECTION F1108, F1109 AND F1110, THE SURFACE WIND SPEED F1111, F1112 AND F1113, THE SURFACE WIND DIRECTION F1114, F1115 AND F1116, THE SURFACE WIND SPEED F1117, F1118 AND F1119, THE SURFACE WIND DIRECTION F1120, F1121 AND F1122, THE SURFACE WIND SPEED F1123, F1124 AND F1125, THE SURFACE WIND DIRECTION F1126, F1127 AND F1128, THE SURFACE WIND SPEED F1129, F1130 AND F1131, THE SURFACE WIND DIRECTION F1132, F1133 AND F1134, THE SURFACE WIND SPEED F1135, F1136 AND F1137, THE SURFACE WIND DIRECTION F1138, F1139 AND F1140, THE SURFACE WIND SPEED F1141, F1142 AND F1143, THE SURFACE WIND DIRECTION F1144, F1145 AND F1146, THE SURFACE WIND SPEED F1147, F1148 AND F1149, THE SURFACE WIND DIRECTION F1150, F1151 AND F1152, THE SURFACE WIND SPEED F1153, F1154 AND F1155, THE SURFACE WIND DIRECTION F1156, F1157 AND F1158, THE SURFACE WIND SPEED F1159, F1160 AND F1161, THE SURFACE WIND DIRECTION F1162, F1163 AND F1164, THE SURFACE WIND SPEED F1165, F1166 AND F1167, THE SURFACE WIND DIRECTION F1168, F1169 AND F1170, THE SURFACE WIND SPEED F1171, F1172 AND F1173, THE SURFACE WIND DIRECTION F1174, F1175 AND F1176, THE SURFACE WIND SPEED F1177, F1178 AND F1179, THE SURFACE WIND DIRECTION F1180, F1181 AND F1182, THE SURFACE WIND SPEED F1183, F1184 AND F1185, THE SURFACE WIND DIRECTION F1186, F1187 AND F1188, THE SURFACE WIND SPEED F1189, F1190 AND F1191, THE SURFACE WIND DIRECTION F1192, F1193 AND F1194, THE SURFACE WIND SPEED F1195, F1196 AND F1197, THE SURFACE WIND DIRECTION F1198, F1199 AND F1200, THE SURFACE WIND SPEED F1201, F1202 AND F1203, THE SURFACE WIND DIRECTION F1204, F1205 AND F1206, THE SURFACE WIND SPEED F1207, F1208 AND F1209, THE SURFACE WIND DIRECTION F1210, F1211 AND F1212, THE SURFACE WIND SPEED F1213, F1214 AND F1215, THE SURFACE WIND DIRECTION F1216, F1217 AND F1218, THE SURFACE WIND SPEED F1219, F1220 AND F1221, THE SURFACE WIND DIRECTION F1222, F1223 AND F1224, THE SURFACE WIND SPEED F1225, F1226 AND F1227, THE SURFACE WIND DIRECTION F1228, F1229 AND F1230, THE SURFACE WIND SPEED F1231, F1232 AND F1233, THE SURFACE WIND DIRECTION F1234, F1235 AND F1236, THE SURFACE WIND SPEED F1237, F1238 AND F1239, THE SURFACE WIND DIRECTION F1240, F1241 AND F1242, THE SURFACE WIND SPEED F1243, F1244 AND F1245, THE SURFACE WIND DIRECTION F1246, F1247 AND F1248, THE SURFACE WIND SPEED F1249, F1250 AND F1251, THE SURFACE WIND DIRECTION F1252, F1253 AND F1254, THE SURFACE WIND SPEED F1255, F1256 AND F1257, THE SURFACE WIND DIRECTION F1258, F1259 AND F1260, THE SURFACE WIND SPEED F1261, F1262 AND F1263, THE SURFACE WIND DIRECTION F1264, F1265 AND F1266, THE SURFACE WIND SPEED F1267, F1268 AND F1269, THE SURFACE WIND DIRECTION F1270, F1271 AND F1272, THE SURFACE WIND SPEED F1273, F1274 AND F1275, THE SURFACE WIND DIRECTION F1276, F1277 AND F1278, THE SURFACE WIND SPEED F1279, F1280 AND F1281, THE SURFACE WIND DIRECTION F1282, F1283 AND F1284, THE SURFACE WIND SPEED F1285, F1286 AND F1287, THE SURFACE WIND DIRECTION F1288, F1289 AND F1290, THE SURFACE WIND SPEED F1291, F1292 AND F1293, THE SURFACE WIND DIRECTION F1294, F1295 AND F1296, THE SURFACE WIND SPEED F1297, F1298 AND F1299, THE SURFACE WIND DIRECTION F1300, F1301 AND F1302, THE SURFACE WIND SPEED F1303, F1304 AND F1305, THE SURFACE WIND DIRECTION F1306, F1307 AND F1308, THE SURFACE WIND SPEED F1309, F1310 AND F1311, THE SURFACE WIND DIRECTION F1312, F1313 AND F1314, THE SURFACE WIND SPEED F1315, F1316 AND F1317, THE SURFACE WIND DIRECTION F1318, F1319 AND F1320, THE SURFACE WIND SPEED F1321, F1322 AND F1323, THE SURFACE WIND DIRECTION F1324, F1325 AND F1326, THE SURFACE WIND SPEED F1327, F1328 AND F1329, THE SURFACE WIND DIRECTION F1330, F1331 AND F1332, THE SURFACE WIND SPEED F1333, F1334 AND F1335, THE SURFACE WIND DIRECTION F1336, F1337 AND F1338, THE SURFACE WIND SPEED F1339, F1340 AND F1341, THE SURFACE WIND DIRECTION F1342, F1343 AND F1344, THE SURFACE WIND SPEED F1345, F1346 AND F1347, THE SURFACE WIND DIRECTION F1348, F1349 AND F1350, THE SURFACE WIND SPEED F1351, F1352 AND F1353, THE SURFACE WIND DIRECTION F1354, F1355 AND F1356, THE SURFACE WIND SPEED F1357, F1358 AND F1359, THE SURFACE	

PROG3P (Continued)

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C TS TO F6, DEFINITION Q=0,0334892*M(850)+0,0363689*M(700)+0,0290951*M(
C 0,0145476*M(300) WHERE M(P) ARE MIXING RATIOS IN MTS=UNITS, PS=101,325
C AT SEA SURFACE,
  LABEL = 10H INITF
  BTIME = SECOND(FAKE)
  WRITE(6,8000) BTIME
  CALL INITF(F1,F2,F3,F4,F5,F6,F7(1,1),F7(1,5),F7(1,9),F7(1,13),
  *F7(1,17),F7(1,21),F7(1,25),F8,M,N)
  SCALE=20.
  DTIME = SECOND(FAKE) * BTIME
  WRITE(6,8005) LABEL, DTIME
  LABEL = 10H MAP 6X
  BTIME = SECOND(FAKE)
  WRITE(6,8000) BTIME
C SOLUTION OF STREAMFUNCTIONS
  LABEL = 10H STREAMF 1
  BTIME = SECOND(FAKE)
  WRITE(6,8000) BTIME
  LABEL = 10H STREAMF
  BTIME = SECOND(FAKE)
  WRITE(6,8000) BTIME
  15 CALL STREAMF(F1,F7,F10,MARK,ZB,PSIB,FB,SB,GAMMA,0,IT1,M,N)
  DTIME = SECOND(FAKE) * BTIME
  WRITE(6,8005) LABEL, DTIME
  LABEL = 10H STREAMF 2
  BTIME = SECOND(FAKE)
  WRITE(6,8000) BTIME
  CALL STREAMF(F2,F8,F10,MARK,ZB,PSIB,FB,SB,GAMMA,0,IT2,M,N)
  DTIME = SECOND(FAKE) * BTIME
  WRITE(6,8005) LABEL, DTIME
  LABEL = 10H STREAMF 3
  BTIME = SECOND(FAKE)
  WRITE(6,8000) BTIME
  CALL STREAMF(F3,F9,F10,MARK,ZB,PSIB,FB,SB,GAMMA,0,IT3,M,N)
  DTIME = SECOND(FAKE) * BTIME
  WRITE(6,8005) LABEL, DTIME
  DO 20 I=1,MN
  F5(I)=F5(I)*0,1
  Z=F2(I)
  F2(I)=,5*(Z+F1(I))
  F3(I)=,5*(F3(I)-Z)
  F1(I)=Z
  20 CONTINUE
C ADAPTION OF HUMIDITY FIELD
  DO 21 I=1,MN
  TZ = G*(H3*F2(I)+H4*F3(I))/F0
  TPSI = ,5*(H3*(F8(I)-F7(I))+H4*(F9(I)-F8(I)))
  CALL SATUR(TZ,QZ)
  CALL SATUR(TPSI,QPSI)
C CHANGE IN F4 SINCE F4 IS RELATIVE HUMIDITY
  F4(I)=F4(I)*QPSI
  F4(I)=F4(I)*QPSI/QZ
  Z = F4(I)-,8*QPSI
  IF(Z,GT,0,0) F4(I)=,8*QPSI
  21 CONTINUE
C STORE Z-FIELDS FOR BOUNDARY MIXING
  LABEL = 10H RANWT 3X
  BTIME = SECOND(FAKE)
  WRITE(6,8000) BTIME
  CALL RANWT(ZM2,F1)
  CALL RANWT(Z12,F2)

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PROG3P,117  
 PROG3P,118  
 PROG3P,119  
 PROG3P,120  
 PROG3P,121  
 PROG3P,122  
 PROG3P,123  
 APR14,17  
 PROG3P,125  
 PROG3P,126  
 PROG3P,127  
 PROG3P,128  
 PROG3P,129  
 PROG3P,130  
 PROG3P,139  
 PROG3P,140  
 PROG3P,141  
 PROG3P,142  
 PROG3P,143  
 PROG3P,144  
 PROG3P,145  
 APR14,18  
 PROG3P,147  
 PROG3P,148  
 PROG3P,149  
 PROG3P,150  
 PROG3P,151  
 APR14,19  
 PROG3P,153  
 PROG3P,154  
 PROG3P,155  
 PROG3P,156  
 PROG3P,157  
 APR14,20  
 PROG3P,159  
 PROG3P,160  
 PROG3P,161  
 PROG3P,162  
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 PROG3P,177  
 PROG3P,178  
 PROG3P,179  
 PROG3P,180  
 PROG3P,181  
 PROG3P,182  
 PROG3P,183  
 PROG3P,184  
 PROG3P,185

PROG3P (Continued)

CALL RANWT(Z22,F3)	PROG3P.186
DTIME = SECOND(FAKE) = BTIME	PROG3P.187
WRITE(6,8005) LABEL, DTIME	PROG3P.188
C TEST FOR ELLIPTICITY OF PSI-FIELDS	PROG3P.189
LABEL = 10H ELLIPT 3X	PROG3P.190
BTIME = SECOND(FAKE)	PROG3P.191
WRITE(6,8000) BTIME	PROG3P.192
CALL ELLIPT(F7,MY,M,N)	APR14,21
CALL ELLIPT(F8,MY,M,N)	APR14,22
CALL ELLIPT(F9,MY,M,N)	APR14,23
DTIME = SECOND(FAKE) = BTIME	PROG3P.196
WRITE(6,8005) LABEL, DTIME	PROG3P.197
C COMPUTATION OF PSIM,PSI1 AND PSI2	PROG3P.198
DO 25 I=1,MN	PROG3P.199
Z=F8(I)	PROG3P.200
F8(I)=.5*(Z-F7(I))	PROG3P.201
F9(I)=.5*(F9(I)-Z)	PROG3P.202
F7(I)=Z	PROG3P.203
25 CONTINUE	PROG3P.204
IF(IVAR.EQ.0) GO TO 31	PROG3P.205
C STORE FIELDS FOR BOUNDARY MIXING	PROG3P.206
DO 30 I=1,MN	PROG3P.207
Z=F1(I)	PROG3P.208
F1(I)=F7(I)	PROG3P.209
F7(I)=F7(I)-G*Z/F0	PROG3P.210
Z=F2(I)	PROG3P.211
F2(I)=F8(I)	PROG3P.212
F8(I)=F8(I)-G*Z/F0	PROG3P.213
Z=F3(I)	PROG3P.214
F3(I)=F9(I)	PROG3P.215
F9(I)=F9(I)-G*Z/F0	PROG3P.216
30 CONTINUE	PROG3P.217
GO TO 36	PROG3P.218
31 DO 35 I=1,MN	PROG3P.219
F1(I)=F7(I)	PROG3P.220
F2(I)=F8(I)	PROG3P.221
F3(I)=F9(I)	PROG3P.222
35 CONTINUE	PROG3P.223
LABEL = 10H RANWT 3X	PROG3P.224
BTIME = SECOND(FAKE)	PROG3P.225
WRITE(6,8000) BTIME	PROG3P.226
36 CALL RANWT(STRM,F7)	PROG3P.227
CALL RANWT(STR1,F8)	PROG3P.228
CALL RANWT(STR2,F9)	PROG3P.229
DTIME = SECOND(FAKE) = BTIME	PROG3P.230
WRITE(6,8005) LABEL, DTIME	PROG3P.231
C PRINT NUMBER OF ITERATIONS IN THE SOLUTION OF PSI	PROG3P.232
41 WRITE(6,202) IT1,IT2,IT3	PROG3P.233
202 FORMAT(///1X,33HNUMBER OF ITERATIONS IN STREAMF ,3(14))	PROG3P.234
GO TO 53	PROG3P.235
C GENERATION OF INITIAL FIELDS FOR THE CHANNEL VERSION, PSI-FIELDS ARE	PROG3P.236
C GENERATED IN THE SUBROUTINE GENCH FROM PARAMETERS GIVEN IN DATA=STATEM	PROG3P.237
C HUMIDITY-FIELD CORRESPONDS TO 50 PERCENT RELATIVE HUMIDITY, PS IS 100	PROG3P.238
C AND TS 273 DEGREES EVERYWHERE.	PROG3P.239
45 CALL GENCH(F7,M,N,PSIPS,UPS,NU,NWAVE,NX,NY,PSIC,LAMC,PSIS,LAMS,WX)	PROG3P.240
WRITE(6,FPARR)	PROG3P.241
CALL GENCH(F8,M,N,PSIPM,UPM,NU,NWAVE,NX,NY,PSIC,LAMC,PSIS,LAMS,WX)	PROG3P.242
CALL GENCH(F9,M,N,PSIP1,UP1,NU,NWAVE,NX,NY,PSIC,LAMC,PSIS,LAMS,WX)	PROG3P.243
C COMPUTATION OF PSIM,PSI1,PSI2,HUMIDITY,PS AND TS	PROG3P.244
46 DO 50 I=1,MN	PROG3P.245

PROG3P (Continued)

F1(I)=F8(I)	PROG3P,246
F2(I)=.5*(F8(I)+F7(I))	PROG3P,247
F3(I)=.5*(F9(I)+F8(I))	PROG3P,248
F5(I)= 100.	PROG3P,249
F6(I)= 273.	PROG3P,250
TPSI= H3*F2(I)+H4*F3(I)	PROG3P,251
CALL SATUR(TPSI,QPSI)	PROG3P,252
F4(I)=.5+QPSI	PROG3P,253
50 CONTINUE	PROG3P,254
C STORE PSI-FIELDS IN STRM,STR1 AND STR2	PROG3P,255
CALL RANWT(STRM,F1)	PROG3P,256
CALL RANWT(STR1,F2)	PROG3P,257
CALL RANWT(STR2,F3)	PROG3P,258
C SMOOTHING OF INITIAL FIELDS	PROG3P,259
53 IF(NSMUTT(1).NE.0) GO TO 55	PROG3P,260
LABEL = 10H ASMUT 6X	PROG3P,261
BTIME = SECOND(FAKE)	PROG3P,262
WRITE(6,8000) BTIME	PROG3P,263
CALL ASMUT(F1,F10,M,N,.5)	PROG3P,264
CALL ASMUT(F1,F10,M,N,-.5)	PROG3P,265
CALL ASMUT(F2,F10,M,N,.5)	PROG3P,266
CALL ASMUT(F2,F10,M,N,-.5)	PROG3P,267
CALL ASMUT(F3,F10,M,N,.5)	PROG3P,268
CALL ASMUT(F3,F10,M,N,-.5)	PROG3P,269
DTIME = SECOND(FAKE) - BTIME	PROG3P,270
WRITE(6,8005) LABEL, DTIME	PROG3P,271
C LOADING OF INITIAL FIELDS, ZERO TO ACCUMULATED PRECIPITATION	PROG3P,272
55 DO 56 I=1,MN	PROG3P,273
F10(I)=0.0	PROG3P,274
56 CONTINUE	PROG3P,275
LABEL = 10H RANWT 10X	PROG3P,276
BTIME = SECOND(FAKE)	PROG3P,277
WRITE(6,8000) BTIME	PROG3P,278
CALL RANWT(PSIM1,F1)	PROG3P,279
CALL RANWT(PSI11,F2)	PROG3P,280
CALL RANWT(PSI21,F3)	PROG3P,281
CALL RANWT(HUM1, F4)	PROG3P,282
CALL RANWT(PS, F5)	PROG3P,283
CALL RANWT(TS, F6)	PROG3P,284
CALL RANWT(PREC, F10)	PROG3P,285
CALL RANWT(WS, F10)	PROG3P,286
CALL RANWT(DIV1, F10)	PROG3P,287
CALL RANWT(DIV2, F10)	PROG3P,288
DTIME = SECOND(FAKE) - BTIME	PROG3P,289
WRITE(6,8005) LABEL, DTIME	PROG3P,290
IF(MAPHOUR(1).NE.0) GO TO 60	PROG3P,291
C MAPPRINTING OF FIRST TIMESTEP	PROG3P,292
IF(KIND.NE.0) GO TO 58	PROG3P,293
LABEL = 10H MAP3P	PROG3P,294
BTIME = SECOND(FAKE)	PROG3P,295
WRITE(6,8000) BTIME	PROG3P,296
CALL MAP3P(1,1,1,1,0,2,0,2,0,0,0,F1,F2,F3,F4,F5,F6,F7,F8,F9,F10,	APR14,24
X MARK,M,N,IDAY,ITIME,NTIME,PNIVS,PNIVM,PNIV1,ZB,PSIR,FB)	PROG3P,298
DTIME = SECOND(FAKE) - BTIME	PROG3P,299
WRITE(6,8005) LABEL, DTIME	PROG3P,300
GO TO 60	PROG3P,301
58 CALL MAP3P(0,0,0,0,0,0,0,0,1,1,1,F1,F2,F3,F4,F5,F6,F7,F8,F9,F10,	APR14,25
X MARK,M,N,IDAY,ITIME,NTIME,PNIVS,PNIVM,PNIV1,ZB,ZB,ZB)	PROG3P,303
LABEL = 10H FORECASTS	PROG3P,304
BTIME = SECOND(FAKE)	PROG3P,305

PROG3P (Continued)

WRITE(6,8000) BTIME	PROG3P,306
C THREE HOURS FORECAST STARTS HERE	PROG3P,307
60 DO 61 I=1,MN	PROG3P,308
F1(I)= 0.0	PROG3P,309
61 CONTINUE	PROG3P,310
CALL RANWT(DIV1,F1)	PROG3P,311
CALL RANWT(DIV2,F1)	PROG3P,312
CALL RANRD(HUM1,F2)	PROG3P,313
CALL RANWT(HUM2,F2)	PROG3P,314
CALL RANRD(PSIM1,F2)	PROG3P,315
CALL RANWT(PSIM2,F2)	PROG3P,316
CALL RANRD(PSI11,F2)	PROG3P,317
CALL RANWT(PSI12,F2)	PROG3P,318
CALL RANRD(PSI21,F2)	PROG3P,319
CALL RANWT(PSI22,F2)	PROG3P,320
IF(IVAR.EQ.0) GO TO 70	PROG3P,321
IF(KIND.NE.0) GO TO 70	PROG3P,322
C INPUT OF BOUNDARY FIELD EACH SIX HOUR	PROG3P,323
C THE SUBROUTINE ZINPUT HAS TO BE WRITTEN ***	PROG3P,324
CALL RANRD(ZM2,F2)	PROG3P,325
CALL RANWT(ZM1,F2)	PROG3P,326
CALL RANRD(Z12,F2)	PROG3P,327
CALL RANWT(Z11,F2)	PROG3P,328
CALL RANRD(Z22,F2)	PROG3P,329
CALL RANWT(Z21,F2)	PROG3P,330
C CALL ZINPUT( ) ZPS,ZPM,ZP1 TO FIELDS F1,F2 AND F3,	PROG3P,331
DO 65 I=1,MN	PROG3P,332
Z=F2(I)	PROG3P,333
F2(I)=.5*(Z-F1(I))	PROG3P,334
F3(I)=.5*(F3(I)-Z)	PROG3P,335
F1(I)=Z	PROG3P,336
65 CONTINUE	PROG3P,337
CALL RANWT(ZM2,F1)	PROG3P,338
CALL RANWT(Z12,F2)	PROG3P,339
CALL RANWT(Z22,F3)	PROG3P,340
C GENERAL TIMESTEP THREE HOURS AHEAD	PROG3P,341
70 CALL STEP3P(F1,F2,F3,F4,F5,F6,F7,F8,F9,F10,MY,MARK,IT,IVAR,M,N)	APR14,26
NCOUNT =NTIME	PROG3P,343
NTIME=NTIME+1	JUN12,11
I1=NTSTEP+1	PROG3P,345
C PRINT NUMBER OF ITERATIONS	PROG3P,346
WRITE(6,200) NCOUNT,NTIME	PROG3P,347
DO 71 I=1,I1	PROG3P,348
WRITE(6,204) IT(1,I),IT(2,I)	PROG3P,349
71 CONTINUE	PROG3P,350
C SMOOTHING AND ELLIPTICITY TEST	PROG3P,351
I1=0	PROG3P,352
I2=0	PROG3P,353
DO 80 I=1,10	PROG3P,354
IF(ABS(FLOAT(NSMUTT(I))-NTIME).LT.0.1) I1=1	JUN12,12
IF(ABS(FLOAT(MSMUTT(I))-NTIME).LT.0.1) I1=1	JUN12,13
IF(ABS(FLOAT(MELLIPT(I))-NTIME).LT.0.1) I2=1	JUN12,14
IF(ABS(FLOAT(NELLIPT(I))-NTIME).LT.0.1) I2=1	JUN12,15
80 CONTINUE	PROG3P,357
IF(I1.EQ.0.AND.I2.EQ.0) GO TO 95	PROG3P,358
CALL RANPD(PSIM1,F1)	PROG3P,359
CALL RANRD(PSI11,F2)	PROG3P,360
CALL RANRD(PSI21,F3)	PROG3P,361
DO 85 I=1,MN	PROG3P,362
Z=F1(I)	PROG3P,363

PROG3P (Continued)

F1(I)=Z+2*F2(I)	PROG3P,364
F3(I) = Z+2.*F3(I)	PROG3P,365
F2(I)=Z	PROG3P,366
85 CONTINUE	PROG3P,367
IF(I1.EQ.0) GO TO 86	PROG3P,368
CALL ASMUT(F1,F4,M,N,.5)	PROG3P,369
CALL ASMUT(F1,F4,M,N,=.5)	PROG3P,370
CALL ASMUT(F2,F4,M,N,.5)	PROG3P,371
CALL ASMUT(F2,F4*M,N,=.5)	PROG3P,372
CALL ASMUT(F3,F4,M,N,.5)	PROG3P,373
CALL ASMUT(F3,F4,M,N,=.5)	PROG3P,374
86 IF(I2.EQ.0) GO TO 90	PROG3P,375
IF(KIND.NE.0) GO TO 90	PROG3P,376
CALL ELLIPT(F1,MY,M,N)	APR14,27
CALL ELLIPT(F2,MY,M,N)	APR14,28
CALL ELLIPT(F3,MY,M,N)	APR14,29
90 DO 93 I=1,MN	PROG3P,380
Z=F2(I)	PROG3P,381
F2(I)=.5*(Z-F1(I))	PROG3P,382
F3(I)=.5*(F3(I)-Z)	PROG3P,383
F1(I)=Z	PROG3P,384
93 CONTINUE	PROG3P,385
CALL RANWT(PSIM1,F1)	PROG3P,386
CALL RANWT(PSI11,F2)	PROG3P,387
CALL RANWT(PSI21,F3)	PROG3P,388
C MAPPRINTING	PROG3P,389
95 DO 96 I=1,10	PROG3P,390
IF(MAPHOUR(I).EQ.NTIME) GO TO 97	PROG3P,391
96 CONTINUE	PROG3P,392
GO TO 100	PROG3P,393
97 IF(KIND.NE.0) GO TO 98	PROG3P,394
CALL MAP3P(1,1,1,0,1,1,1,0,0,0,F1,F2,F3,F4,F5,F6,F7,F8,F9,F10,	APR14,30
X MARK,M,N,IDAY,ITIME,NTIME,PNIVS,PNIVM,PNIV1,ZB,FB,SB)	PROG3P,396
GO TO 100	PROG3P,397
98 CALL MAP3P(0,0,0,0,1,0,0,0,1,1,0,F1,F2,F3,F4,F5,F6,F7,F8,F9,F10,	APR14,31
X MARK,M,N,IDAY,ITIME,NTIME,PNIVS,PNIVM,PNIV1,ZB,ZB,ZB)	PROG3P,399
C ZERO TO ACCUMULATED PRECIPITATION	PROG3P,400
100 DO 105 I=1,10	PROG3P,401
IF(LETACC(I).EQ.NTIME) GO TO 106	PROG3P,402
105 CONTINUE	PROG3P,403
GO TO 115	PROG3P,404
106 DO 110 I=1,MN	PROG3P,405
F1(I)=0.0	PROG3P,406
110 CONTINUE	PROG3P,407
CALL RANWT(PREC,F1)	PROG3P,408
200 FORMAT(29H1NUMBER OF ITERATIONS BETWEEN,13,4H AND,13,6H HOURS)	PROG3P,409
201 FORMAT(4X,2(I4))	PROG3P,410
115 IF(NEND.LE.NTIME) STOP	PROG3P,411
GO TO 60	PROG3P,412
END	PROG3P,413



SUBROUTINE STEP3P

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SUBROUTINE STEP3P(F1,F2,F3,F4,F5,F6,F7,F8,F9,F10,MY,MARK,IT,
X          IVAR,M,N)
C SIX HOURS FORECAST BY 3P=MODEL
DIMENSION F1(1),F2(1),F3(1),F4(1),F5(1),F6(1),F7(1),F8(1),F9(1),
X          F10(1),MY(1),F(1),MARK(1),IT(2,1)
COMMON/FPAR/IC(8),JC(8),XPOL,YPOL,R,RE,DS,JMIN(100),JMAX(100),
X          MX,NX,KIND
COMMON/ECS/ PSIM1,PSI11,PSI21,PSIM2,PSI12,PSI22,HUM1,HUM2,DIV1,
2DIV2,WS,HEAT,J789,J12,J56,J3,PS,TS,PREC,STRM,STR1,STR2,ZM1,Z11,Z21
3,ZM2,Z12,Z22,H13,H23,HM3,HM2,H12,H22,H11,H21,J4,VM,V1,V2
COMMON/COEFF/A1,A2,A3,B1,B2,B3,C1,C2,C3,C4,C5,C6,C7,C8,D,DELP,EM,
X          E1,E2,H1,H2,H3,H4,H5,H6,PMEAN,S1,S2,T1,T2,T3,T4,T5,
X          P0,PM,P1
COMMON/COEFF2/T6,T7,T8,T9,T10,T11,T12,T13,T14,
X          K1,K2,K3,K4,K5,K6,K7,K8,K9,K10,K11,K12,K13,K14,K15,
X          K16,K17,K18,K19,K20,K21,K22,K23,K24,K25,K26,K27,K28,
X          K29,K30,K31,K32,K33,K34,K35,K36,K37,K38
COMMON/RUNPAR/DELT,NTSTEP,ALFASYS,ALFAM,ALFAZ,ALFAPSI,RESSYS,RESM,
X          RESZ,RESPS1,Q,FOCEAN,FCONT,WGT1,WGT2,WGT3,ADIFF
COMMON F
REAL MY,KEFF
REAL          K1,K2,K3,K4,K5,K6,K7,K8,K9,K10,K11,K12,K13,K14,K15,
X          K16,K17,K18,K19,K20,K21,K22,K23,K24,K25,K26,K27,K28,
X          K29,K30,K31,K32,K33,K34,K35,K36,K37,K38
DATA RGAS,EE,HL,CP,T0,E0,DEL1,DEL2,TOL/287,..622,2.5E6,1004,..
X          273,..611,0.0,0.0,0.0,0.0/
C DEL1,DEL2, AND TOL HAVE TO BE DEFINED *****
DEL1=0.0001
DEL2=0.001
TOL=0.0
C CONSTANTS FOR COMPUTATION OF LATENT HEAT
CC1 = 1./T0
CC2 = EE*HL/RGAS
CC3 = EE*HL/CP
CC4 = CC2*CC3
CC5 = DEL1+DEL2
C
IF(KIND.EQ.0) GO TO 5
WGT1=1.0
WGT2=.67
WGT3=.33
5 MN=4*N
ND=NTSTEP+1
M1=M-1
DO 170 KT=1,ND
EPS = 2.
IF(KT.LT.3) EPS = .5*KT
C
C*****JACOBIAN COMPUTATIONS*****
C ALL JACOBIANS ARE COMPUTED AND STORED
CALL RANRD(PSIM1,F1)
CALL RANRD(PSI11,F2)
CALL RANRD(PSI21,F3)
C
CALL ABSVOR(F1,F4,MY,MARK,M,N)
CALL RELVOR(F2,F5,MY,MARK,M,N)
CALL RELVOR(F3,F6,MY,MARK,M,N)
KEFF=1.0
CALL RANRD(PS,MY)
CALL DLOG(F5,MY,M,N,KEFF)
C

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APR14,46  
STEP3P,3  
STEP3P,4  
STEP3P,5  
STEP3P,6  
STEP3P,7  
STEP3P,8  
STEP3P,18  
STEP3P,16  
STEP3P,20  
STEP3P,21  
STEP3P,22  
STEP3P,23  
STEP3P,24  
STEP3P,25  
STEP3P,26  
STEP3P,27  
STEP3P,28  
STEP3P,29  
APR14,47  
STEP3P,30  
STEP3P,31  
STEP3P,32  
STEP3P,33  
STEP3P,34  
STEP3P,35  
STEP3P,36  
STEP3P,37  
STEP3P,38  
STEP3P,39  
STEP3P,40  
STEP3P,41  
STEP3P,42  
STEP3P,43  
STEP3P,44  
STEP3P,45  
STEP3P,46  
STEP3P,47  
STEP3P,48  
JUN12,16  
JUN12,17  
STEP3P,51  
STEP3P,52  
STEP3P,53  
STEP3P,54  
STEP3P,55  
STEP3P,56  
STEP3P,57  
STEP3P,58  
STEP3P,59  
STEP3P,60  
STEP3P,61  
STEP3P,62  
STEP3P,63  
APR14,48  
APR14,49  
APR14,50  
STEP3P,67  
STEP3P,68  
STEP3P,69  
STEP3P,70

STEP3P (Continued)

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DO 10 I=1,MN
F7(I) = C3*F4(I)-C2*F5(I)+C4*F6(I) + C7*F(I)
F8(I) = C2*F4(I)+C9*F5(I)
F9(I) = C4*F4(I) +C6*F6(I)+2*C7*F(I)
10 CONTINUE
C
CALL JACOB(F1,F7,F10,M,N)
CALL JACOB(F2,F8,F7,M,N)
CALL JACOB(F3,F9,F8,M,N)
CALL OROG(F7,MY,M,N,KEFF)
C
DO 20 I=1,MN
F10(I) = F10(I)+F8(I)+F7(I)
F9(I) = F4(I)-2*F5(I)
20 CONTINUE
C
CALL RANWT(J789,F10)
C
CALL JACOB(F2,F9,F7,N,N)
CALL OROG(F7,MY,M,N,KEFF)
CALL JACOB(F1,F5,F8,M,N)
C
DO 30 I=1,MN
F10(I) = F7(I)+F8(I)
F9(I) = F4(I)+2*F6(I)
30 CONTINUE
C
CALL RANWT(J12,F10)
C
CALL JACOB(F3,F9,F7,M,N)
CALL JACOB(F1,F6,F8,M,N)
CALL JACOB(F1,F3,F9,M,N)
CALL JACOB(F1,F2,F10,M,N)
CALL OROG(F10,MY,M,N,KEFF)
C
CALL RANWT(J3,F10)
C
DO 40 I=1,MN
40 F10(I)=F7(I)+F8(I)
CALL RANWT(J56,F10)
CALL RANRD(PS,F10)
PPM=2/(P0-PM)
DO 44 I=1,MN
F7(I)=F1(I)-F2(I)*PPM*(F10(I)-PM)
F8(I)=PPM*F5(I)*(F10(I)-PM)-F4(I)+F(I)
44 F4(I)=F10(I)
C
CALL JACOB(F7,F4,F5,M,N)
CALL MYFF(MY,M,N)
C
C*****LOWER BOUNDARY CONDITION*****STEP3P.120
C INFLUENCE FROM TOPOGRAPHY AND FRICTIO OVERR LAND OR OCEAN SURFACE STEP3P.121
C OCEAN SURFACE IS ASSUMED WHERE STANDARD PRESSURE PS IS 101.35 CB OR MOSTEP3P.123
DO 50 I=1,MN
IF(MARK(I)) 49,50,50
49 CF=FOCEAN
PP=F4(I)
IF(PP,LT,101.35) CF=FCONT
F4(I) = .25*MY(I)*F5(I)+CF*F8(I)
F10(I)=PP
50 CONTINUE
STEP3P.71
STEP3P.72
STEP3P.73
STEP3P.74
STEP3P.75
STEP3P.76
STEP3P.77
STEP3P.78
STEP3P.79
STEP3P.80
STEP3P.81
STEP3P.82
STEP3P.83
STEP3P.84
STEP3P.85
STEP3P.86
STEP3P.87
STEP3P.88
STEP3P.89
STEP3P.90
STEP3P.91
STEP3P.92
STEP3P.93
STEP3P.94
STEP3P.95
STEP3P.96
STEP3P.97
STEP3P.98
STEP3P.99
STEP3P.100
STEP3P.101
STEP3P.102
STEP3P.103
STEP3P.104
STEP3P.105
STEP3P.106
STEP3P.107
STEP3P.108
STEP3P.109
STEP3P.110
STEP3P.111
STEP3P.112
STEP3P.113
STEP3P.114
STEP3P.115
STEP3P.116
STEP3P.117
STEP3P.118
APR14,51
STEP3P.120
STEP3P.121
STEP3P.122
STEP3P.123
STEP3P.124
STEP3P.125
STEP3P.126
STEP3P.127
STEP3P.128
STEP3P.129
STEP3P.130
STEP3P.131

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STEP3P (Continued)

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C                                                    STEP3P,132
  CALL RANWT(WS,F4)                                STEP3P,133
  CALL RANRD(TS,F5)                                STEP3P,134
C*****SENSIBLE HEAT*****STEP3P,135
C HEATING FROM OCEAN SURFACE WHEN THE AIR IS COLDER, PQ IS THE TEMP DIFFSTEP3P,136
C OCEAN SURFACE IS ASSUMED WHERE STANDARD PRESSURE PS IS 101.35 CB OR MOSTSTEP3P,137
  DO 59 I=2,M1                                     STEP3P,138
    J1=JMIN(I)+1                                   STEP3P,139
    J2=JMAX(I)-1                                   STEP3P,140
    K=(J1-1)*M+1                                   STEP3P,141
    DO 59 J=J1,J2                                  STEP3P,142
      PP=F10(K)                                     STEP3P,143
      PQ=F5(K)-H2*F2(K)                             STEP3P,144
      IF(PP.LT.101.324) GO TO 57                    STEP3P,145
      IF(PQ.LE.0.0) GO TO 57                        STEP3P,146
      PR = SQRT(.25*MY(K)*((F7(K+1)-F7(K-1))*(F7(K+1)-F7(K-1))+
X      (F7(K+M)-F7(K-M))*(F7(K+M)-F7(K-M))))      STEP3P,148
      F5(K) = .5E-2*H1*(1+PR-1,)*PQ                STEP3P,149
      GO TO 58                                       STEP3P,150
  57 F5(K)=0.0                                       STEP3P,151
  58 K=K+M                                           STEP3P,152
  59 CONTINUE                                        STEP3P,153
C*****HUMIDITY FORECAST*****STEP3P,154
C EM,E1,E2 ARE COEFF FOR COMP OF MEAN STREAMFUNCTION STEP3P,155
C SS1,SS2,SS3 ARE COEFF FOR COMP OF MEAN DIVERGENCE, D IS ZERO OVER LANDSTEP3P,156
C THE HUMIDITY IS GIVEN IN TON PER SQUAREMETER AND CENTIBAR STEP3P,157
  CALL RANRD(DIV1,F6)                                STEP3P,158
  CALL RANRD(DIV2,F7)                                STEP3P,159
  CALL RANRD(HUM1,F8)                                STEP3P,160
C                                                    STEP3P,161
  SS1 = E1+EM*C2                                     STEP3P,162
  SS2 = E2-EM*C1                                     STEP3P,163
  SS3 = -EM*C8                                       STEP3P,164
  DO 61 I=1,MN                                       STEP3P,165
    PQ = D                                           STEP3P,166
    PP = F10(I)                                       STEP3P,167
    IF(PP.LT.101.35) P0=0.0                          STEP3P,168
    F7(I) = F8(I)*(SS1*F6(I)+SS2*F7(I)+F4(I)*(PQ+SS3)) STEP3P,169
    F6(I) = EM*F1(I)+E1*F2(I)+E2*F3(I)              STEP3P,170
  61 CONTINUE                                        STEP3P,171
  CALL JACOB(F6,F8,F10,M,N)                          STEP3P,172
  DO 62 I=2,M1                                       STEP3P,173
    J1=JMIN(I)+1                                   STEP3P,174
    J2=JMAX(I)-1                                   STEP3P,175
    K=(J1-1)*M+1                                   STEP3P,176
    DO 62 J=J1,J2                                  STEP3P,177
      F6(K) = F8(K+M)+F8(K-M)+F8(K+1)+F8(K-1)-4*F8(K) STEP3P,178
      K=K+M                                         STEP3P,179
  62 CONTINUE                                        STEP3P,180
C                                                    STEP3P,181
  DEPS = EPS*DELT                                    STEP3P,182
  DO 63 I=1,MN                                       STEP3P,183
    IF(MARK(I).GE.0) GO TO 63                       STEP3P,184
    F8(I)=-DEPS*(.25*MY(I)*F10(I)+F7(I)-ADIFF*MY(I)*F6(I)) STEP3P,185
  63 CONTINUE                                        STEP3P,186
C                                                    STEP3P,187
  CALL RANRD(HUM2,F6)                                STEP3P,188
  DO 65 I=1,MN                                       STEP3P,189
    IF(MARK(I)) 64,65,65                             STEP3P,190
  64 F6(I) = F6(I)+F8(I)                             STEP3P,191
  65 CONTINUE                                        STEP3P,192
C*****PRECIPITATION*****STEP3P,193

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STEP3P (Continued)

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C THE RAIN FOR ONE TIMESTEP IS ACCUMULATED. THE RAIN IS GIVEN IN MM OR KSTEP3P.104
C PER SQUAREMETER - STEP3P.105
DO 70 I=1,MN STEP3P.196
IF(MARK(I)) 66,70,70 STEP3P.197
66 TEMP = H3*F2(I)+H4*F3(I) STEP3P.198
CALL SATUR(TEMP,QSAT) STEP3P.199
PQ = F6(I)-.8*QSAT STEP3P.200
IF(PQ) 68,68,67 STEP3P.201
67 F6(I) = .8*QSAT STEP3P.202
F7(I) = .5E3*DELP*PQ STEP3P.203
IF(KT,EQ.1) F7(I)=0,0 STEP3P.204
IF(KT,EQ.2) F7(I)=2*F7(I) STEP3P.205
GO TO 70 STEP3P.206
68 F7(I) = 0 STEP3P.207
PQ = F6(I)-.2*QSAT STEP3P.208
IF(PQ) 69,70,70 STEP3P.209
69 F6(I) = .2*QSAT STEP3P.210
70 CONTINUE STEP3P.211
CALL RANRD(PREC,F8) STEP3P.212
DO 71 I=1,IN STEP3P.213
F8(I) = F8(I) + F7(I) STEP3P.214
71 CONTINUE STEP3P.215
CALL RANNT(PREC,F8) STEP3P.216
CALL RANRD(HUM1,F8) STEP3P.217
CALL RANNT(HUM1,F6) STEP3P.218
CALL RANNT(HUM2,F8) STEP3P.219
C*****LATENT HEAT*****STEP3P.220
CALL RANRD(DIV1,F6) STEP3P.221
CALL RANRD(DIV2,F8) STEP3P.222
DO 180 I=1,MN STEP3P.223
IF(MARK(I)) 179,180,180 STEP3P.224
179 F6(I) = T1*F4(I)+T2*F6(I)+T3*F8(I) STEP3P.225
180 CONTINUE STEP3P.226
C STEP3P.227
DO 190 I=1,MN STEP3P.228
IF(MARK(I)) 187,190,190 STEP3P.229
187 RAIN = F7(I)/EPS/DELT STEP3P.230
IF(RAIN,LT,TOL) GO TO 188 STEP3P.231
VERT = F6(I) STEP3P.232
IF(VERT,GT,-DEL1) GO TO 188 STEP3P.233
OSTAR = VERT STEP3P.234
IF(VERT,GE,-CC5,AND,VERT,LE,-DEL1) OSTAR = -ABS(VERT*VERT/CC5) STEP3P.235
TEMP = H6*F2(I) STEP3P.236
X = CC2*(CC1-1./TEMP) STEP3P.237
E = E0*EXP(X) STEP3P.238
FSTAR = EE*TEMP*E*(CC3-TEMP)/PMEAN/(PMEAN*TEMP*TEMP + CC4*E) STEP3P.239
HLAT = H1*HL*OSTAR*FSTAR STEP3P.240
GO TO 189 STEP3P.241
188 HLAT = 0,0 STEP3P.242
189 F5(I) = F5(I) + HLAT STEP3P.243
190 CONTINUE STEP3P.244
CALL RANNT(HEAT,F5) STEP3P.245
C STEP3P.246
I1=0 STEP3P.247
I2=0 STEP3P.248
I3=0 STEP3P.249
I4=0 STEP3P.250
CALL STEPEXT(F1,F2,F3,F4,F5,F6,F7,F8,F9,F10,MY,MARK,M,N, APR14.56
X I1,I2,I3,I4) STEP3P.253
C*****FORCING FUNCTIONS*****STEP3P.254

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STEP3P (Continued)

CALL RANRD(J789,F10)	STEP3P.255
DEPS = .25*DELT*EPS	STEP3P.256
DO 73 I=1,MN	STEP3P.257
IF(MARK(I)) 72,73,73	STEP3P.258
72 PQ = 4*F(I)/MY(I)	STEP3P.259
F8(I) =-DEPS*(F10(I) - CR*F4(I)+PQ)	STEP3P.260
F6(I) = PQ*(A3*F4(I)+A1*F5(I))	STEP3P.261
F7(I)=PQ*(B3*F4(I)+B1*F5(I))	STEP3P.262
73 CONTINUE	STEP3P.263
C	STEP3P.264
CALL RANRD(J3,F10)	STEP3P.265
CALL RANRD(J12,F4)	STEP3P.266
CALL RANRD(J56,F5)	STEP3P.267
C	STEP3P.268
DO 80 I=1,MN	STEP3P.269
IF(MARK(I)) 79,80,80	STEP3P.270
79 PQ = F(I)*F(I)	STEP3P.271
F6(I) =-DEPS*(F4(I)+PQ+(A2*F9(I)-A1*F10(I))+F6(I))	STEP3P.272
F7(I) =-DEPS*(F5(I)+PQ*(A1*F10(I)-B2*F9(I))+F7(I))	STEP3P.273
80 CONTINUE	STEP3P.274
C*****SOLUTION OF FORECAST F0,*****	STEP3P.275
CALL RANRD(HM3,F4)	STEP3P.276
CALL RANRD(H13,F5)	STEP3P.277
CALL RANRD(H23,F10)	STEP3P.278
DEPS1=DELT*EPS	STEP3P.279
DO 81 I=1,MN	STEP3P.280
F8(I)=F8(I)+DEPS1/MY(I)*F4(I)	STEP3P.281
F6(I)=F6(I)+DEPS1/MY(I)*F5(I)	STEP3P.282
81 F7(I)=F7(I)+DEPS1/MY(I)*F10(I)	STEP3P.283
CALL RANRD(PS112,F5)	STEP3P.284
CALL RANRD(PS122,F10)	STEP3P.285
C	STEP3P.286
DO 90 I=1,MN	STEP3P.287
89 F2(I)=2*(F2(I)-F5(I))	STEP3P.288
F3(I)=2*(F3(I)-F10(I))	STEP3P.289
90 CONTINUE	STEP3P.290
C	STEP3P.291
LABEL = 10H HELMSYS	APR14.57
BTIME = SECOND(DUMMY)	APR14.58
WRITE(6,8000) BTIME	APR14.59
8000 FORMAT(1H , *BTIME= *, F10.4)	APR14.60
CALL HEL ISYS(F2,F3,F6,F7,MY,F4,A1,A2,B1,B2,ALFASYS,RESSYS,	APR14.61
X ITSYS,M,N)	STEP3P.293
DTIME = BTIME - SECOND(DUMMY)	APR14.62
WRITE(6,8005) LABEL, DTIME	APR14.63
8005 FORMAT(1H , *TIME TO EXECUTE *, A10, F10.4)	APR14.64
IT(1,KT) = ITSYS	STEP3P.294
CALL ASMUT(F2,F4,M,N,.,5)	JUN12.18
CALL ASMUT(F2,F4,M,N,.,5)	JUN12.19
CALL ASMUT(F3,F4,M,N,.,5)	JUN12.20
CALL ASMUT(F3,F4,M,N,.,5)	JUN12.21
C	STEP3P.295
DO 100 I=1,MN	STEP3P.296
IF(MARK(I)) 99,100,100	STEP3P.297
99 TFILT=0.4	JUN12.22
IF(MARK(I).EQ.-10) TFILT=0.7	JUN12.23
IF(MARK(I).EQ.-1) TFILT=1.	JUN12.24
F5(I)=F5(I) + TFILT*F2(I)	JUN12.25
F10(I)=F10(I) + TFILT*F3(I)	JUN12.26
F4(I) = 0/MY(I)	STEP3P.300
100 CONTINUE	STEP3P.301
C	STEP3P.302

STEP3P (Continued)

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C      CALL RANRD(PSIM2,F6)
C      DO 110 I=1,MN
109  F1(I) = 2*(F1(I)-F6(I)) -C2*F2(I) +C1*F3(I)
110  CONTINUE
C      CALL HELM(F1,F8,F4,ALFAM,RESM,ITM,M,N)
      IT(2,KT) = ITM
C      CALL ASMUT(F1,F4,M,N,,5)
      CALL ASHUT(F1,F4*M,N,,-.5)
      DO 120 I=1,MN
119  IF(MARK(I))119,120,120
      TFILT=.4
      IF(MARK(I).EQ.-10) TFILT=0.7
      IF(MARK(I).EQ.-1) TFILT=1.
      F6(I)=F6(I) + TFILT*F1(I) + C2*F2(I)
      S=C1*F3(I)
120  CONTINUE
C*****MIXING WITH BOUNDARY FIELDS*****
      CALL RANRD(STRM,F4)
      IF(IVAR.EQ.0) GO TO 156
      IF(KIND.NE.0) GO TO 156
      CALL RANRD(ZM1,F7)
      CALL RANRD(ZM2,F8)
C      WF = (KT-1)/ND
      FW = 1.-WF
      G = 9.806
      DO 130 I=1,MN
129  IF(MARK(I)) 129,130,130
      F7(I) = G*(FW*F7(I)+WF*F8(I))/F(I)
130  CONTINUE
      CALL MIXF(F6,F7,MARK,WGT1,WGT2,WGT3,M,N)
      CALL RANRD(STR1,F4)
C      CALL RANRD(Z11,F7)
      CALL RANRD(Z12,F8)
      DO 140 I=1,MN
139  IF(MARK(I)) 139,140,140
      F7(I) = G*(FW*F7(I)+WF*F8(I))/F(I)
140  CONTINUE
      CALL MIXF(F5,F7,MARK,WGT1,WGT2,WGT3,M,N)
      CALL RANRD(STR2,F4)
C      CALL RANRD(Z21,F7)
      CALL RANRD(Z22,F8)
      DO 150 I=1,MN
149  IF(MARK(I)) 149,150,150
      F7(I) = G*(FW*F7(I)+WF*F8(I))/F(I)
150  CONTINUE
      CALL MIXF(F10,F7,MARK,WGT1,WGT2,WGT3,M,N)
      GO TO 156
151  CALL MIXF(F6,F4,MARK,WGT1,WGT2,WGT3,M,N)
      CALL RANRD(STRM,F4)
      CALL MIXF(F5,F4,MARK,WGT1,WGT2,WGT3,M,N)
      CALL RANRD(STR1,F4)
      CALL MIXF(F10,F4,MARK,WGT1,WGT2,WGT3,M,N)
      CALL RANRD(STR2,F4)
C*****STORE NEW TIMESTEP*****

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STEP3P.303  
STEP3P.304  
STEP3P.305  
STEP3P.306  
STEP3P.307  
STEP3P.308  
STEP3P.309  
STEP3P.310  
STEP3P.311  
JUN12,27  
JUN12,28  
STEP3P.312  
JUN12,29  
JUN12,30  
JUN12,31  
JUN12,32  
JUN12,33  
JUN12,34  
STEP3P.314  
STEP3P.315  
STEP3P.316  
STEP3P.317  
STEP3P.318  
STEP3P.319  
STEP3P.320  
STEP3P.321  
STEP3P.322  
STEP3P.323  
STEP3P.324  
STEP3P.325  
STEP3P.326  
JUN12,35  
STEP3P.328  
STEP3P.329  
STEP3P.330  
STEP3P.331  
STEP3P.332  
STEP3P.333  
STEP3P.334  
STEP3P.335  
JUN12,36  
STEP3P.337  
STEP3P.338  
STEP3P.339  
STEP3P.340  
STEP3P.341  
STEP3P.342  
STEP3P.343  
STEP3P.344  
JUN12,37  
STEP3P.346  
STEP3P.347  
STEP3P.348  
STEP3P.349  
STEP3P.350  
STEP3P.351  
STEP3P.352  
STEP3P.353  
STEP3P.354  
STEP3P.355

STEP3P (Continued)

156	IF(KT.LT.3) GO TO 157	STEP3P,356
	CALL RANRD(PSIM1,F1)	STEP3P,357
	CALL RANRD(PSI11,F4)	STEP3P,358
	CALL RANRD(PSI21,F7)	STEP3P,359
157	CALL RANWT(PSIM1,F6)	STEP3P,360
	CALL RANWT(PSI11,F5)	STEP3P,361
	CALL RANWT(PSI21,F10)	STEP3P,362
	IF(KT.LT.3) GO TO 158	STEP3P,363
	CALL RANWT(PSIM2,F1)	STEP3P,364
	CALL RANWT(PSI12,F4)	STEP3P,365
	CALL RANWT(PSI22,F7)	STEP3P,366
C	*****COMPUTATION OF DIVERGENCE*****	STEP3P,367
158	CALL RANRD(J3,F10)	STEP3P,368
	CALL RANRD(HEAT,F5)	STEP3P,369
	CALL RANRD(WS,F8)	STEP3P,370
C		STEP3P,371
	PQ = 1./EPS/DELT	STEP3P,372
	DO 160 I=1,MM	STEP3P,373
	IF(MARK(I)) 159,160,160	STEP3P,374
159	TERM1 = F(I)*(PQ *F2(I)+.25*MY(I)*F10(I))-F5(I)-S1 *F8(I)	STEP3P,375
	TERM2 = F(I)*(PQ *F3(I)+.25*MY(I)*F9 (I)) -S2 *F8(I)	STEP3P,376
	F2(I) =- A1*TERM1 + A2*TERM2	STEP3P,377
	F3(I) =- B1*TERM1 - B2*TERM2	STEP3P,378
160	CONTINUE	STEP3P,379
C		STEP3P,380
	CALL RANWT(DIV1,F2)	STEP3P,381
	CALL RANWT(DIV2,F3)	STEP3P,382
	CALL RANRD(PSIM1,F1)	STEP3P,383
	CALL RANRD(PSI11,F5)	STEP3P,384
	CALL RANRD(PSI21,F10)	STEP3P,385
C		STEP3P,386
	PRINT 7512,KT	STEP3P,387
7512	FORMAT(1X,3HKT=,I3)	STEP3P,388
170	CONTINUE	STEP3P,389
	RETURN	STEP3P,390
	END	STEP3P,391

# SUBROUTINE STEPEXT

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SUBROUTINE STEPEXT(F1,F2,F3,F4,F5,F6,F7,F8,F9,F10,MY,MARK,M,N, APR14,67
1 11,I2,I3,I4) STEPEXT,3
C THIS SUBROUTINE COMPUTES THE CONTRIBUTION FROM THE HIGHER ORDER TERMS STEPEXT,4
C IN THE VORTICITY EQUATION, STEPEXT,5
C I1 INDICATES THE VORTICITY ADVECTION BY THE DIVERGENT VIND STEPEXT,6
C I2 INDICATES THE RELATIVE VORTICITY* DIVERGENCE STEPEXT,7
C I3 INDICATES THE VERTICAL ADVECTION OF VORTICITY STEPEXT,8
C I4 INDICATES THE TWISTINGTERM STEPEXT,9
C I1=0 NO CONTRIBUTION I1 DIFFERENT FROM 0 CONTRIBUTION FROM I1 STEPEXT,10
C I2=0 NO CONTRIBUTION I2 DIFFERENT FROM 0 CONTRIBUTION FROM I2 STEPEXT,11
C I3=0 NO CONTRIBUTION I3 DIFFERENT FROM 0 CONTRIBUTION FROM I3 STEPEXT,12
C I4=0 NO CONTRIBUTION I4 DIFFERENT FROM 0 CONTRIBUTION FROM I4 STEPEXT,13
C PSIM IS IN F1,PSI1 IS IN F2,PSI2 IS IN F3,WS IS IN F4 STEPEXT,14
C STEPEXT NEEDS 10 FIELDS IN THE FAST CORE MEMORY F,MY AND MARK MUST STEPEXT,15
C ALSO BE IN FAST CORE MEMORY STEPEXT,16
DIMENSION F1(1),F2(1),F3(1),F4(1),F5(1),F6(1),F7(1),F8(1),F9(1), STEPEXT,17
1 F10(1),F(1),MARK(1),MY(1) STEPEXT,18
COMMON F APR14,68
COMMON/COEFF/A1,A2,A3,B1,B2,B3,C1,C2,C3,C4,C5,C6,C7,C8,D,DELP,EM, STEPEXT,19
1 E1,E2,H1,H2,H3,H4,H5,H6,PMEAN,S1,S2,T1,T2,T3,T4,T5, STEPEXT,20
2 P0,PM,P1 STEPEXT,21
COMMON/COEFF/T6,T7,T8,T9,T10,T11,T12,T13,T14, STEPEXT,22
X K1,K2,K3,K4,K5,K6,K7,K8,K9,K10,K11,K12,K13,K14,K15, STEPEXT,23
X K16,K17,K18,K19,K20,K21,K22,K23,K24,K25,K26,K27,K28, STEPEXT,24
X K29,K30,K31,K32,K33,K34,K35,K36,K37,K38 STEPEXT,25
COMMON/ECS/PSIM1,PSI11,PSI21,PSI12,PSI22,HUM1,HUM2,DIV1, APR14,69
2 DIV2,WS,HEAT,J789,J12,J56,J3,PS,TS,PREC,STRM,STR1,STR2,ZH1,Z11,Z21 APR14,70
3,ZM2,Z12,Z22,H13,H23,H43,H2,H12,H22,H11,H21,J4,VM,V1,V2 APR14,71
REAL MY STEPEXT,34
REAL K1,K2,K3,K4,K5,K6,K7,K8,K9,K10,K11,K12,K13,K14,K15, STEPEXT,35
X K16,K17,K18,K19,K20,K21,K22,K23,K24,K25,K26,K27,K28, STEPEXT,36
X K29,K30,K31,K32,K33,K34,K35,K36,K37,K38 STEPEXT,37
KIND=0 APR14,72
MN=1+N STEPEXT,38
RESIDUE =.5E4 STEPEXT,39
ALFA=1.4 STEPEXT,40
C JACOBIAN J4 TO SECONDARY STORAGE,DIVERGENCIES TO FAST MEMORY STEPEXT,41
CALL RANHT(J4,F9) STEPEXT,42
CALL RANRD(DIV1,F5) STEPEXT,43
CALL RANRD(DIV2,F6) STEPEXT,44
DO 9 I=1,MN STEPEXT,45
IF(MARK(I)) 9,7,7 STEPEXT,46
7 F4(I)=0.0 STEPEXT,47
F5(I)=0.0 STEPEXT,48
F6(I)=0.0 STEPEXT,49
9 CONTINUE STEPEXT,50
IF(KIND.EQ.0) GO TO 8 STEPEXT,51
CALL BMOVE(F4,M,N) STEPEXT,52
CALL BMOVE(F5,M,N) STEPEXT,53
CALL BMOVE(F6,M,N) STEPEXT,54
8 CONTINUE STEPEXT,55
IF(I4.EQ.0.AND.I3.EQ.0.AND.I2.EQ.0.AND.I1.EQ.0) GO TO 170 STEPEXT,56
IF(I4.EQ.0) GO TO 44 STEPEXT,57
C COMPUTE THE TWISTINGTERM,I4 STEPEXT,58
DO 10 I=1,MN STEPEXT,59
F7(I)=K31*F5(I)+K32*F6(I)+K33*F4(I) STEPEXT,60
10 F8(I)=K36*F5(I)+K37*F6(I)+K38*F4(I) STEPEXT,61
CALL GRADPR(F2,F7,F9,MARK,M,N) STEPEXT,62
CALL GRADPR(F3,F8,F10,MARK,M,N) STEPEXT,63
DO 11 I=1,MN STEPEXT,64
F9(I)=0.5*MY(I)*F9(I) STEPEXT,65

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STEPEXT (Continued)

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11 F10(I)=0.5*MY(I)*F10(I) STEPEXT.66
   CALL RANWT(H13,F9) STEPEXT.67
   CALL RANWT(H23,F10) STEPEXT.68
   DO 20 I=1,MN STEPEXT.69
20 F7(I)=K19*F5(I)+K20*F6(I)+K21*F4(I) STEPEXT.70
   F8(I)=K22*F5(I)+K23*F6(I)+K24*F4(I) STEPEXT.71
   CALL GRADPR(F2,F7,F9,MARK,M,N) STEPEXT.72
   CALL GRADPR(F3,F8,F10,MARK,M,N) STEPEXT.73
   DO 30 I=1,MN STEPEXT.74
30 F7(I)=K25*F5(I)+K26*F6(I)+K27*F4(I) STEPEXT.75
   CALL GRADPR(F1,F7,F8,MARK,M,N) STEPEXT.76
   DO 40 I=1,MN STEPEXT.77
40 F8(I)=(F8(I)+F9(I)+F10(I))*0.5*MY(I) STEPEXT.78
   CALL RANWT(HM3,F8) STEPEXT.79
   GO TO 45 STEPEXT.80
44 DO 46 I=1,MN STEPEXT.81
46 F8(I)=0.0 STEPEXT.82
   CALL RANWT(H13,F8) STEPEXT.83
   CALL RANWT(H23,F8) STEPEXT.84
   CALL RANWT(HM3,F8) STEPEXT.85
45 IF(I2.EQ.0.AND.I3.EQ.0.AND.I1.EQ.0) GO TO 163 STEPEXT.86
50 CALL RELVOR(F1,F7,MY,MARK,M,N) APR14.73
   CALL RELVOR(F2,F8,MY,MARK,M,N) APR14.74
   CALL RELVOR(F3,F9,MY,MARK,M,N) APR14.75
   IF(I2.EQ.0.AND.I3.EQ.0) GO TO 94 STEPEXT.90
   IF(I2.EQ.0) GO TO 65 STEPEXT.91
C COMPUTATION OF THE RELATIVE VORTICITY* DIVERGENCE, I2 STEPEXT.92
   DO 60 I=1,MN STEPEXT.93
60 F1(I)=F5(I)*(K1*F8(I)+K2*F9(I)+K3*F7(I))+F6(I)*(K4*F8(I)+K5*F9(I)+
   K6*F7(I))+F4(I)*(K7*F8(I)+K8*F9(I)+K9*F7(I)) STEPEXT.94
   F2(I)=F5(I)*(K28*F8(I)+F7(I))+F6(I)*K29*F8(I)+F4(I)*K30*F8(I) STEPEXT.95
60 F3(I)=F5(I)*K34*F9(I)+F6(I)*( K35*F9(I)+F7(I))+F4(I)*K30*F9(I) STEPEXT.96
C COMPUTATION OF THE VERTICAL ADVECTION OF VORTICITY, I3 STEPEXT.97
   IF(I3.EQ.0) GO TO 91 STEPEXT.98
   GO TO 80 STEPEXT.99
65 DO 70 I=1,MN STEPEXT.101
70 F1(I)=0.0 STEPEXT.102
   F2(I)=0.0 STEPEXT.103
70 F3(I)=0.0 STEPEXT.104
80 DO 90 I=1,MN STEPEXT.105
90 F1(I)=F1(I)+F5(I)*(K10*F8(I)+K11*F9(I)+K12*F7(I))+
   F6(I)*(K13*F8(I)+K14*F9(I)+K15*F7(I))+
   F4(I)*(K16*F8(I)+K17*F9(I)+K18*F7(I)) STEPEXT.106
90 F2(I)=F2(I)+F8(I)*(K31*F5(I)+K32*F6(I)+K33*F4(I)) STEPEXT.107
90 F3(I)=F3(I)+F9(I)*(K36*F5(I)+K37*F6(I)+K38*F4(I)) STEPEXT.108
   GO TO 91 STEPEXT.109
94 DO 92 I=1,MN STEPEXT.110
94 F1(I)=0.0 STEPEXT.111
   F2(I)=0.0 STEPEXT.112
94 F3(I)=0.0 STEPEXT.113
91 CALL RANWT(HM2,F1) STEPEXT.114
   CALL RANWT(H12,F2) STEPEXT.115
   CALL RANWT(H22,F3) STEPEXT.116
C COMPUTATION OF THE ADVECTION OF VORTICITY BY THE DIVERGENT WIND -I1 STEPEXT.117
C COMPUTE FORCINGFUNCTION FOR THE VELOCITYPOTENTIAL STEPEXT.118
   IF(I1.EQ.0) GO TO 166 STEPEXT.119
95 DO 100 I=1,MN STEPEXT.120
95 F1(I)=(C2*F5(I)-C1*F6(I)-C8*F4(I))/MY(I) STEPEXT.121
   F2(I) = F5(I)/MY(I) STEPEXT.122
100 F3(I) = F6(I)/MY(I) STEPEXT.123
C SOLVE THE POISSONEQUATION BY RELAXATION IN ORDER TO GET VELOCITYPOT. STEPEXT.124
   CALL RANRD(VN,F4) STEPEXT.125

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# STEPEXT (Continued)

	CALL RANRD(V1,F5)	STEPEXT,128
	CALL RANRD(V2,F6)	STEPEXT,129
	CALL VELPOT(F4,F1,M,N,RESIDUE,ALFA)	STEPEXT,130
	CALL VELPOT(F5,F2,M,N,RESIDUE,ALFA)	STEPEXT,131
	CALL VELPOT(F6,F3,M,N,RESIDUE,ALFA)	STEPEXT,132
	CALL RANWT(VH,F4)	STEPEXT,133
	CALL RANWT(V2,F6)	STEPEXT,134
	CALL RANWT(V1,F5)	STEPEXT,135
	DO 110 I=1,MN	STEPEXT,136
	F1(I)=F7(I)+F(I)-2.*F8(I)	STEPEXT,137
110	F2(I)=F7(I)+F(I)+2.*F9(I)	STEPEXT,138
	CALL GRADPR(F4,F8,F3,MARK,M,N)	STEPEXT,139
	CALL GRADPR(F5,F1,F10,MARK,M,N)	STEPEXT,140
	DO 120 I=1,MN	STEPEXT,141
120	F10(I)=-0.5*MY(I)*(F3(I)+F10(I))	STEPEXT,142
	CALL RANWT(H11,F10)	STEPEXT,143
	CALL GRADPR(F4,F9,F3,MARK,M,N)	STEPEXT,144
	CALL GRADPR(F6,F2,F10,MARK,M,N)	STEPEXT,145
	DO 130 I=1,MN	STEPEXT,146
130	F10(I)=-0.5*MY(I)*(F3(I)+F10(I))	STEPEXT,147
	CALL RANWT(H21,F10)	STEPEXT,148
	DO 140 I=1,MN	STEPEXT,149
	F1(I)=C3*(F7(I)+F(I))-C2*F8(I)+C4*F9(I)+C7*F(I)	STEPEXT,150
	F2(I)=-C2*(F7(I)+F(I))+C5*F8(I)	STEPEXT,151
140	F3(I)=C4*(F7(I)+F(I))+C6*F9(I)+2.*C7*F(I)	STEPEXT,152
	CALL GRADPR(F4,F1,F7,MARK,M,N)	STEPEXT,153
	CALL GRADPR(F5,F2,F8,MARK,M,N)	STEPEXT,154
	CALL GRADPR(F6,F3,F9,MARK,M,N)	STEPEXT,155
	DO 150 I=1,MN	STEPEXT,156
150	F1(I)=-0.5*MY(I)*(F7(I)+F8(I)+F9(I))	APR14,76
	CALL RANRD(HM2,F2)	STEPEXT,158
	CALL RANRD(HM3,F3)	STEPEXT,159
	CALL RANRD(H11,F4)	STEPEXT,160
	CALL RANRD(H12,F5)	STEPEXT,161
	CALL RANRD(H13,F6)	STEPEXT,162
	CALL RANRD(H21,F7)	STEPEXT,163
	CALL RANRD(H22,F8)	STEPEXT,164
	CALL RANRD(H23,F9)	STEPEXT,165
	DO 160 I=1,MN	STEPEXT,166
	F1(I)=F1(I)+F2(I)+F3(I)	STEPEXT,167
	F4(I)=F4(I)+F5(I)+F6(I)	STEPEXT,168
160	F7(I)=F7(I)+F8(I)+F9(I)	STEPEXT,169
	GO TO 190	STEPEXT,170
163	CALL RANRD(HM3,F1)	STEPEXT,171
	CALL RANRD(H13,F4)	STEPEXT,172
	CALL RANRD(H23,F7)	STEPEXT,173
	GO TO 190	STEPEXT,174
166	CALL RANRD(HM2,F2)	STEPEXT,175
	CALL RANRD(HM3,F3)	STEPEXT,176
	CALL RANRD(H12,F5)	STEPEXT,177
	CALL RANRD(H13,F6)	STEPEXT,178
	CALL RANRD(H22,F8)	STEPEXT,179
	CALL RANRD(H23,F9)	STEPEXT,180
	DO 167 I=1,MN	STEPEXT,181
	F1(I)=F2(I)+F3(I)	STEPEXT,182
	F4(I)=F5(I)+F6(I)	STEPEXT,183
167	F7(I)=F8(I)+F9(I)	STEPEXT,184
	GO TO 190	STEPEXT,185
170	DO 180 I=1,MN	STEPEXT,186
	F1(I)=0.0	STEPEXT,187
	F4(I)=0.0	STEPEXT,188
180	F7(I)=0.0	STEPEXT,189

STEPEXT (Continued)

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190 CALL RANWT(HM3,F1)
    CALL RANWT(H13,F4)
    CALL RANWT(H23,F7)
    CALL RANRD(PSIM1,F1)
    CALL RANRD(PSI11,F2)
    CALL RANRD(PSI21,F3)
    CALL RANRD(WS,F4)
    CALL RANRD(HEAT,F5)
200 CALL RANRD(J4,F9)
    RETURN
    END
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STEPEXT,190
STEPEXT,191
STEPEXT,192
STEPEXT,193
STEPEXT,194
STEPEXT,195
STEPEXT,196
STEPEXT,197
STEPEXT,198
STEPEXT,199
STEPEXT,200
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