



Х



Electron Density Profiles of Wavemotions in the Ionosphere Caused by Nuclear Detonations

GEORGE J. GASSMANN



IONOSPHERIC PHYSICS LABORATORY PROJECT 8653

AIR FOPCE CAMBRIDGE RESEARCH LABORATORIES, OFFICE OF AEROSPACE RESEARCH, UNITED STATES AIR FORCE, L.G. HANSCOM FIELD, MASS.

AFCRL--63-440 APRIL 1963

1-

Research Report

. . نتر

1444

- Right and a second design

- 7

1 25 3

Electron Density Profiles of Wavemotions in the lonosphere Caused by Nuclear Detonations

GEORGE J. GASSMANN

IONOSPHERIC PHYSICS LABORATORY PROJECT 8653

AIR FORCE CAMBRIDGE RESEARCH LABORATORIES, OFFICE OF AEROSPACE RESEARCH, UNITED STATES AIR FORCE, L.G. HANSCOM FIELD, MASS.

Abstract

Horizontally traveling waves in the ionosphere. occurring naturally and from nuclear detonations, cause nonvertical reflections and, therefore, abnormal ionospheric recordings at stations during overhead passage. It is shown how an electron density cross-section in the vertical plane through an ionospheric wave can be constructed from a single station's ionospheric recordings, provided those are taken at time intervals not exceeding 5 minutes and provided the general direction of travel is known. The described analysis yields also an approximate value for the velocity. Figures 6 and 7 show examples of profiles of ionospheric waves from two different nuclear detonations, observed at great distances. The ionospheric wave of 30 October 1961 from Novaya Zemlya appears to be caused by a gravity wave, as may be inferred from Table 3 which gives a review of world-wide observations.

÷

いたななく

A NUMBER

An an and an and a new or here a second the second

Acknowledgements

Š

This research has been sponsored in part by the Defense Atomic Support Agency, Washington, D. C. The airborne ionograms were produced under the supervision of R. Gowell of AFCRL, with Major R. Creager of the 3245th Operations Group, L. G. Hanscom Field, Bedford, Massachusetts as aircraft commander. Helpful discussions were held with T. Barrett of American Science and Engineering, Inc., Cambridge, Massachusetts.

Contents

`_-

		Page
	List of Illustrations	vi
	List of Tables	vi
1	INTRODUCTION	1
2	THE RAW DATA	2
3	THE ELECTRON DENSITY PROFILES	10
4	ESTIMATE OF VELOCITY	10
5	CONSTRUCTION OF VERTICAL PROFILES OF IONOSPHERIC	
	WAVES	13
6	DISCUSSION	15
	References	21

v

ŧ

Illustrations

4

Figure	I	Page
1	Outline Map – Novaya Zemlya Detonation (in both geographic and magnetic coordinates)	3
2a, b	Sequence of Ionograms Taken on Aircraft on 30 October 1961, 0850 to 0920 GMT	4,5
3	Ionograms of Figure 2 Retraced, and Extraordinary Components and Multiples Eliminated	6
4	Outline Map – Bikini Detonation (in both geographical and magnetic coordinates)	7
5a, b	Sequence of Ionograms taken at Kusaie Island after the Bikini Detonation	8,9
6	Cross-Section of Ionospheric Wave from Novaya Zemlya Passing over Aircraft. Dotted contour lines are not measured and represent suggested supplementation.	14
7	Cross-Section of Ionospheric Wave from Bikini Passing over Kusaie. Dotted contour lines are not measured and represent suggested supplementation.	15
8	Graphical Construction of Curved Ray Paths	16

Tables

Table		Page
1	The Ionograms of Figure 3 reduced to true 'height'	11
2	The Ionograms of Figure 5 reduced to true 'height'	12
3	Review of Observations on Ionospheric Wave of 30 October 1961	17

۰.

Electron Density Profiles of Wavemotions in the lonosphere Caused by Nuclear Detonations

1. INTRODUCTION

Jonospheric h'f recordings, especially those taken at arctic stations, frequently during one 24-hour period show echoes that cannot be explained by assuming horizontally stratified layers. It has been obvious that some of the echoes are not strictly vertical soundings but arise from large-scale irregularities near the zenith of the station. From comparing consecutive h'f recordings it has also been obvious that many of those natural irregularities travel with a speed of the order of sound waves. The determination of dimensions, speed, and direction is possible, in principle, if one utilizes several sounding stations whose spacings are large enough to obtain, for triangulation, measurable differences in echo-ranges at each station and whose spacings are small enough to obtain recordings that can be correlated.

Such stringent experimental requirements for investigating natural large-scale movements in the ionosphere, requiring three ionospheric stations at close distance (100 km), are considerably eased for the case of man-made disturbances. In the latter case the direction of movement at a particular station and the speed of travel may be deduced approximately from the geographical location and the time where and when the man-made disturbance was introduced. With this information available the ionograms from a single station are sufficient to obtain a vertical crosssection of the ionospheric layers along the direction of travel at a time when the disturbance passes through. How this can be done is outlined below and is demonstrated by two examples. One is the wave from the nuclear detonation at Novaya Zemlya on 30 October 1961, as observed near Norway; the other example utilizes

(Received for publication, 8 March 1963)

ionospheric data taken at Kusaie Island, South Pacific, after a detonation above Bikini Atoll.

2. THE RAW DATA

2

The data on 30 October 1961 were taken from a USAF KC-135 jet aircraft at 67°N-10°E while on a heading toward magnetic north (Figure 1). The distance to Novaya Zemlya was approximately 1600 km. The yield of the detonation was estimated as being between 35 and 90 MT; and the altitude of the burst was estimated as being below the tropopause.¹ Burst time was 0833 GMT. The equipment in the aircraft consisted of a modified Step Frequency Ionospheric Recorder* that sweeps from 1 to 25 Mc in 100 kc steps within a 4-minute time period. Nominal transmitter output was above 10 kw. The transmitting-receiving antenna was of the traveling-wave type; three wires approximately 23 m long extended from a common tiepoint on the aircraft's upper fuselage to three separate points on the vertical stabilizer, each wire terminated by a resistor. The antenna pattern favored upward radiation in a wide cone of about 90° for most frequencies. Anterna efficiency was below but approaching that of the conventional delta antenna of an ionospheric ground station

* most frequencies.

The sequence of ionograms on 35-mm film used in this analysis is reproduced in Figures 2a and 2b. The same ionograms retraced and supplemented, eliminating extraordinary component and multiples, appear in Figure 3.

Data from the Bikini detonation were taken at Kusaie Island 740 km south of Bikini (Figure 4).² The equipment was an Ionospheric Recorder Type C3, which swept the frequency range from 1 to 25 Mc within 30 sec, using a standard delta antenna. Detonation occurred approximately 1 hour before local sunrise. The detonation above Bikini Atoll was below the tropopause and equivalent to several megatons of TNT.

Figures 5a and 5c show the sequence of ionograms used in this analysis. This sequence covers the first disturbance arriving at Kusaie. (A second disturbance, not covered in this analysis, arrived overhead at Kusaie at H+56 minutes.)

* Manufactured by Philips Electronics Industries Ltd, Toronto, Canada



Figure 1. Outline Map - Novaya Zemlya Detonation (in both geographic and magnetic coordinates)

ŝ

-. `

-

.

tana tan

1

٠ •



11、白油四日 ---

Y

-



,





Figure 3. Ionograms of Figure 2 Retraced, and Extraordinary Components and Multiples Eliminated



Э



ŧ

Figure 4. Outline Map - Bikini Detonation (in both geographical and magnetic co-ordinates)

7

**



Figure 52. Sequence of Ionograms taken at Kusaie Island after the Bikini Detonation

8

a more dependent of the second s



:

Figure 5b. Sequence of Ionograms taken at Kusaie Island after the Bikini Deconation

3. THE ELECTRON DENSITY PROFILES

The reproduced sequences of ionograms show that the single echo traces near the start of the sequence develop into two or sometimes more branches (marked I, II, and III on Figure 3). This means that for certain frequencies reflections occur at more than one distance. This indicates clearly the existence of separate ray paths for each reflection at one particular frequency, since in one single ray path the closest reflection surface would obscure all others at a farther listance. (Nonobscuring partial reflections, which produce echoes at virtual ranges lacking the feature of critical frequency and associated range increase (cusp), are not seen on the ionograms and are not considered here.) Each branch in the ionograms is the result of a particular set of ray paths (from the station to the reflection level and back). Each set may consist of a bundle of single rays. For simplification, however, it has been assumed that each br nch is created in the same way as are vertical ionograms, i.e., the ray paths for all frequencies of a branch are identical except for different penetration heights into the layers. With this simplification, reducing the cone angle of the bundle of ray paths to zero, each branch of the ionogram traces can be reduced from virtual to true height. This procedure is standard for vertical ionograms and is done with the notion that the obtained true 'height' is actually the distance to the layer at an unknown elevation angle. For the reduction to true height the simple method of Kelso³ was used since it allows rapid hand calculations and yields sufficient accuracy compatible with that of the other steps in the analysis.

Tables 1 and 2 contain results of the true height reduction, referring to the ionograms of Figures 3 and 5, respectively.

4. ESTIMATE OF VELOCITY

The next step in this analysis is the postulation that the irregularity travels horizontally as a whole and with a particular velocity. The better this is fulfilled and the better the velocity is estimated, the faster the analysis described under Section 5 leads to results.

The effect of the Novaya Zemlya detonation on the ionospheric recordings at the aircraft shows up first at 0900 GMT This yields an average speed of 1600 km/27 min or 987 m/second. On the other hand, other widely-differing velocity values ranging from 330 to 910 m/sec from ground observation were measured (Table 3 page 17). From these data it appeared obvious that the ionospheric disturbance had started with higher velocity, slowing down at medium distances, and that the aircraft happened to be flying at a transition distance. The analysis, therefore, had been started with

			0920		0916		0912			8060	0000	0852			
	والمتعاد والمراجع والمحالي وا			Branch III	Branch II	Branch III	Branch II	Branch I	Branch				GMT	Plasma-Pi	over the same of t
Conve			270	255	255	230	230	220		210	210	210	^f N 2.0	equency.	TAB
rsion:			300	295	295	240	240	240		215	215	215	3.0		
$f_{\rm N}^2$ (Mc) = 1		tru	305	310	390	300	320	330		220	220	220	3. 5		lonograms o
7.98 × 10 ⁻⁵		e 'height' s			410			370					3. 7		of Figure 3
		02	310	320	490	320	370	8 1 3		1 1 1	230	230	40		reduced
ctrons m ³					8 5 5							280	<u>4</u> .5		to true he
			320	330		390					295	1	50		ight s
						1 1 1					330		5.5		
	km		360	400							8 1 1		5.7		
			385	1 t									5.8		

		46	38	36	34		32		30		28		26	24	22	12	After d	Minutes
						Branch III	Branch II	Branch III	Branch II	Branch II	Branch II	Bry h III	Branch II				etonation f	Plas
Conve		240	265	260	260		260		255		245	-	240	235	230	215	N 2.0	ma-Freq
rsion: f _r		250	275	270	270		270		265		260		250	245	240	230	.3.0	uency
2 [Mc] =	t	255	280	285	290		290	320	280	,	270		260	250	250	240	4.0	
- 7. 98 × 10	ue 'height'	260	295	300	310	320	t 1 1	340	1 L J		330		290	290	280	265	5.0	
Ne Ne	a2										8 8 8		1 1 1	320	290	1 1 1	5.1	
ectrons cm ³														8 1 5	315		5.3	
		280	305	310	320	340		350		360		360			1 2 1		6.0	
		5 1 8	315	2 2 1	330	345						1 L 1					6, 4	
								375		390							6.5	
	km		340		350	360		1 1 1		1 1 1			<u></u>				6.8	

-

، انتاج

ì

× 1 . .

TABLE 2. Ionograms of Figure 5 reduced to true height s

an estimated velocity of 675 m/sec which, however, yielded inconsistent results. Several lower velocities down to 420 m/sec were then assumed, and it was found that most consistent results were obtained with a velocity of 500 m/second.

The arrival at Kusaie of the ionospheric disturbance from the Bikini detonation yielded an average speed of 514 m/second. In the analysis, however, consistent results were obtained with velocities centering at 420 m/second.

5. CONSTRUCTION OF VERTICAL PROFILES OF IONOSPHERIC WAVES

Figure 6 presents a two-dimensional coordinate system showing vertical distances above ground and horizontal distances from an arbitrary zero-point. The abscissa also shows the time when ionospheric records were taken. The relative position of the time marks to the kilometer marks on the abscissa is arbitrary, however, the value $\partial(distance)/\partial(time)$ equals 500 m/sec representing the speed postulated above. The zero-km mark on the abscissa may be considered as being fixed with the horizontally-moving ionospheric disturbance; the ionospheric station consequently moves along the zero height mark from left to right. The true height values (s of Table 1) were now plotted as segments of circles, with that particular radius above that point on the ground which is marked with the particular time when the record was taken. After all values of Table 1 were plotted in this manner, a first generation of contour lines of the plasma frequency f_N was drawn by connecting all segments of circles of the same plasma frequency by smooth curves. Several inconsistencies remained, especially in the upper levels of f_N , because in this illustration each ray path from the ionospheric station to the reflection point is represented by a straight line, namely, the radius of the particular segment to the point where it touches the contour line.

A second generation of contour curves was then obtained by replacing those straight ray paths of obviously incorrect shape with curved ones. Samples of curved ray paths were obtained by a graphical construction according to Figure 8, which is based on Snell's Law and which approximates the refraction process by a step-wise procedure. The demand for sample ray paths can be kept to a minimum if one is experienced in selecting ray paths (choosing the right departure angle) whose reflection points are both consistent with others and which approximately fulfill the specular reflection requirement (180° return). The drawing of the contour lines is then a matter of interpolation.

A third generation of contour lines may now be obtained, commencing at the lowest levels of uncertainty and proceeding to higher levels, by calculating the virtual height $s' = \sum_{\mu}^{s} \frac{ds}{\mu}$ along each ray path. If this value differs by $\Delta s'$ from the experimental s' ('virtual height'), the entire procedure might be repeated with new values

- .

ŝ



・ ・ と * ・

••••

-



ę :

rear and the constant and any provide the constant of the constant of the

Į



Figure 6. Cross-Section of Ionospheric Wave from Novaya Zemlya Passing over Aircraft. Dotted contour lines are not measured and represent suggested supplementation.

ş

1.040.0

the assessments to another second and

ŧΊ

7

:



Figure 8. Graphical Construction of Curved Ray Paths

of s individually adjusted according to Δ s'. This refinement was used here on only a few points as a check.

The described construction assumed a velocity of 500 ± 80 m/second. In trials with velocities beyond these limits certain inconsistencies could not be removed.

The ionospheric wave over Kusaie was constructed in the same way. The tabulated values of Table 2 yield contour lines as seen in J igure 7, in which a velocity of 420 ± 80 m/sec is assumed. Assuming velocities b youd the indicated limits led to inconsistencies during the construction of the contour lines.

The above described construction of contour lines may be applied to any sequence of ionospheric recordings. As a general rule, however, it is felt the time interval between recordings should not exceed 5 minutes.

6. DISCUSSION

The traveling ionospheric disturbance from the 30 October 1961 detonation has been observed at several locations. Table 3 shows a review of the results with velocities deduced from the various data and assumptions. The values bear varying validity margins caused by the uncertainties of distance and travel time. The determination of the latter is affected by the period of data taking and by the uncertainty

Osaka, Sodankyla,	Gardiner ⁹ Reykjavik, Lulea, Inverness, Uppsala, <u>Slough,</u> Lindau, Roma	Dieminger Landau et al. 8	Stoffregen ⁷ Kiruna, Lyksele, <u>Uppsala</u> , Lindau, Dourbes, Athens	Hultqvist Kiruna et al 1	Lindquist ⁶ Kiruna	Rose, et al. ³ Sodankyla	Oksman ⁴ Sodankyla	Source Location
East Disturbances of F layer, period ≥ 10 min Excessive velocity near detonation, then nearly	Amplitudes Δ foF2 increasing with decreasing magnetic latitude. Microbarographic (round-the-world) signals appear to travel at velocities v _{west} = 296 m/sec and v _{enct} = 314 m/sec	Marked change of f _O F2 (1/2 hr records) consecutive oscillations of f _O F2, period approx. 1.5 hr Disturbance of F layer 31.5 hr after detonation (round-the-world, NE or SW uncertain).	Sudden change of f_0F^2 with consecutive oscillations, period/hr increasing with distance. Amplitudes Δf_0F^2 increasing with distance.	In addition to the above: at H+25 hr strong disturbance of F layer (attributed to N.Z detonation of 31 Oct 62, 0830 UT) If inter- preted as round-the-world.	Sudden change of $f_{\chi}F2$ at H+24 min	Disturbance of F layer 26.3 hr after detonation (round-the-world NE). Disturbance of F layer 27.2 hr after detonation (round-the-world SW) Microbarographic (round-the-world) signal travels at a speed of 311 m/sec On magnetometers no pronounced effect noticed during ionospheric disturbance.	Sudden change of f _o F2 at H+22 min	Results
		3,200 40,000-3,200 40,000+3,200		40,000-1,350	1,350	40,000-1,200 40,000+1,200	1,200	Approx. Distance to Novaya Zemlya, (km)
420	(Same as Dieminger)	460 to 610 330 380	630 onset (430) 1st max (260) 2nd max (170) 3rd max	430	938	413 418	910	Average Velocity of F Layer Disturbance (m/sec)

TABLE 3. Review of observations on ionospheric wave of 30 October 1961

21

•

ţ

TABLE 3. (Contd)

- 7

~~ <u>`</u>

N.

		1 Aug 58	-
	serva- F layer disturbance inferred from magnetic Pacific. observations. nation	Magn. Obst tories in Pa Teak Deton	Cummack et al. 13
1, 30 ton	Change of ioF2.	Maui, Orange Det 12 Aug 58	Matsushita ¹²
740 740	Sudden indication of F layer disturbance at H+24 min Profile analysis from sequence of ionospheric soundings.	Kusaie, Bikini Deto	Gassmann (Figures 5a and 7)
	OTHER RELATED OBSERVATIONS		•
1,600	min Profile analysis from sequence of ionospheric soundings.	Norway	(Figures 2a and 6)
1,600	r Sudden indication of F layer disturbance at H+27	KC135 near	Gassmann
	ons as Decrease of f ₀ F2 followed by oscillations, period: 1.5 to 2.5 hr. No effect of geomagnetic slus field on propagation of F layer disturbance undaring, noticeable. elen	Same statio analyzed by Obayashi pl Quetta, Mur and Kerguel	Lichtman et al. ¹¹
ns.	ichwarz, Wavelength in F layer approx. 200-300 km, kkanai, (Gravity Wave). nugawa No marked changes in F2 at North American Stat Microbarographic (round-the-world) signals travel at 310 m/sec	Dourbes, Sc Roma, Wakł Akita, Yama	
Appr Dista Nova Zeml	r Results	Location	Source

at which we wanted

. T. Tarker

Real Sector

81

and the second secon

of the time of origin. If, for example, the ionospheric wave was originated by the rising fireball, a delay after detonation of as much as 5 minutes should be assumed, raising some of the velocities listed to even higher values.

The F-layer disturbance appears to have started with a velocity exceeding 1000 m/sec, slowing down after a few thousand kilometers of travel to velocities between 330 and 430 m/second. Simultaneously, a microbarographic wave propagated to far distances at lower velocities near 300 m/sec as a separate phenomenon. The F-layer disturbance consisted of a sudden onset followed by consecutive oscillations that lasted for several hours. The wave appears to have excluded certain directions of propagation; this might explain the discrepancy in the round-the-world velocity values that are based on great circle paths. The observations of Stoffregen and of Gardiner, namely, the increase of the amplitude of the wave with distance and/or with decreasing magnetic latitude, appear significant.

The observed wavelength of >> 100 km indicates that compressional forces play a minor role in this wave since only acoustic waves of frequencies > 0.03 cps or of wavelength < 100 km can propagate.¹⁴ Obayashi suggests that the ionospheric disturbance was caused by a gravitational wave.¹⁰ Such a wave would have large amplitudes at great heights. However, at F-layer maximum and above, the coupling of motion of neutral and charged particles is appreciable and imposes an attenuation on the gravity wave. By this effect Obayashi explains why the largest amplitude of the wave is observed in the lower F layer.

It is obvious that, under the influence of either a gravity wave or a compressional wave, the motion of the F-layer plasma will not coincide with that of the neutral particles, the former creating dynamo currents and/or electric fields. The cross-sections shown in Figures 6 and 7 must therefore be regarded as mere examples of a variety of possible forms. It must be expected that the shapes of the ionization contour lines will vary greatly, depending on the orientation of the geomagnetic vector relative to the direction of propagation.

19

. . .

21

References

 Hultquist, B., Egeland, A., Gustafsson, G., Liszka, L., Ortner, J., Heikkila, K. E., Olsen, S., Makitalo, T., and Eriksson, B., "Report on Observations Made at Kiruna Geophysical Observatory During the Series of Nuclear Weapon Tests Carried Out at Novaja Zemlja Between 10 Sept and 4 Nov 1961," Kiruna Geophysical Observatory, Kiruna C, Sweden Scientific Report No. KGO 611, (11 Dec 1961).

The state of the second s

- 2. Data taken in 1956 during Operation Redwing by Harris, A.K., Gassmann, G.J., Daniels, F.B., Pfister, W., Caplan, P.J., and Winkelman, R.E.
- 3. Kelso, J., "A Procedure for Determination of the Vertical Distribution of the Electron Density in the Ionosphere," Journ. Geoph. Res. 57, 357 (1952).
- 4. Oksman, J., Geophysical Observatory, Sodankyla, Finland, private communication.
- 5. Rose, G., Oksman, J., and Kataja, E., "Round the World Sound Waves Produced by the Nuclear Explosion on 30 October 1961 and Their Effect on the Ionosphere at Sodankyla," <u>Nature</u>, 192, 1173 (1961).
- L'ndquist, R., Research Institute of National Defense, Stockholm 80, Sweden, private communication.
- Stoffregen, W., "Ionospheric Effects Observed in Connection with Nuclear Explosions at Nova Semlja on 23 and 30 October 1961," Res. Inst. of Nat'l Defense, Stockholm 8G, Sweden, Report FOA3/A517, (May 1962).
- 8. Dieminger, W. and Kohl, H., "Effects of Nuclear Explosions on the Ionosphere," Nature, 193 (No. 4819), 963 (Mar 1962).
- 9. Gardiner, G. W., "Effects of the Nuclear Explosion of 30 October 1961," Jour. Atm. Terr. Ph. 24, 990 (Nov 1962).
- 10. Obayashi, T., "Wide-Spread Ionospheric Disturbances Due to Nuclear Explosions During October 1961," <u>Nature</u>, <u>196</u>, 4849 (Oct 1962).
- 11. Lichtman, S. W. and Anderson, E. J., "Ionospheric Effects of Nuclear Detonations in the Atmosphere," Hughes Aircraft Company, Culver City, California, Report OP-36.
- 12. Matsushita, S., "On Artificial Geomagnetic and Ionospheric Storms Associated with High-Altitude Explosions," Jour. Geophy. Res. 64, 1149 (1959).
- 13. Cummack, C. H. and King, G. A. M., "Disturbances in the Ionospheric F-Regionfollowing the Johnston Island Nuclear Explosion," <u>Nature</u> 184, 33 (Sept 1959).
- 14. Eckart, C., <u>Hydrodynamics of Oceans and Atmospheres</u>, Pergamon Press, New York (1960).