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VARIATIONS IN ORBITAL ELEMENTS

James A. Ward, Jr., RCA International Service Corporation

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Instrumentation and Data Processing Division
Directorate of Range Operations
Air Force Eastern Test Range (AFSC)
Patrick Air Force Base, Florida
Due to perturbative forces acting on a near earth satellite, the associated classical Keplerian orbital elements vary with time. These variations are divided into secular, short period and long period. The mathematical equations expressing these variations are presented without derivation along with numerical examples. A discussion of the practical applications of these variations to trajectory generation and orbit determination is included.
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ABSTRACT

Due to perturbative forces acting on a near earth satellite, the associated classical Keplerian orbital elements vary with time. These variations are divided into secular, short period and long period. The mathematical equations expressing these variations are presented without derivation along with numerical examples. A discussion of the practical applications of these variations to trajectory generation and orbit determination is included.
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James A. Ward, Jr.
RCA Missile Test Project
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This technical report has been reviewed and approved.

Robert W. Schmeling
Lt. Colonel, USAF
Chief, Instrumentation and Data Processing Division
Directorate of Range Operations
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SECTION I
INTRODUCTION

If the gravity field of a homogeneous spherical earth were the only force acting on an object, the classical Keplerian orbital elements describing the size, shape and orientation of the orbit of the object would remain constant in time. However, the presence of various perturbative forces causes the classical (osculating) elements to vary with time. These variations can be divided into three categories: secular, short period and long period. The short period variations appear to be functions of the position of the object in its orbit. The long period variations appear to be functions of the position of perigee in space. Keplerian elements from which the short and long period variations have been removed are referred to herein as mean elements. The mean elements change monotonically with time and the rates of these changes are the secular variations.

Analysts making decisions concerning orbits of various objects must be aware of the fact that Keplerian elements vary with time. For example, if comparisons are desired among orbit determination solutions at various epochs, these comparisons must be made in mean elements. If classical elements are used, it may be erroneously inferred that a maneuver had occurred or that one or another solution was unreliable. Since mean elements are well behaved in time, mean elements at any epoch can be computed from those at any other epoch if the secular variations are known. However, it must be remembered that the mean elements do not represent the actual position of an object. For example, if observation station look angles are desired, they must be computed from osculating elements. Hence, it is necessary to be able to convert back and forth between mean and osculating elements.

In summary, the classical elements represent the actual position and velocity of the object but are poorly behaved in time. The mean elements do not represent the actual position and velocity but are well behaved in time. Since each of these characteristics is useful in orbit support, some means of transformation from one to the other is desired. The following sections give the equations for these transformations and show examples of their applications to both theoretical and real data.
The Keplerian elements used in this discussion are defined by:

- \( a \) = semi-major axis
- \( e \) = eccentricity
- \( i \) = inclination
- \( \omega \) = argument of perigee
- \( \Omega \) = right ascension of ascending node
- \( M \) = mean anomaly

The symbols for the classical elements are those used above, while the symbols for the equivalent mean elements are those above in conjunction with a bar superscript (e.g. \( \bar{a} \)).

It is not the purpose of this report to derive all the mathematical expressions used for the various computations, but rather to show what the computational procedures are, together with numerical results obtained. The references are recommended for derivations of the techniques. This report shows examples of the practical applications of these techniques.
SECTION II

SECULAR VARIATIONS

When using mean elements for orbit support it is necessary to have the secular variations or rates of change of these elements with time. Often when mean elements are obtained from an outside agency, the secular variations are included in the transmitted information. A typical example of this case is the Space Defense Center (SDC) 5-card bulletin, an example of which is shown in Appendix I. This bulletin has been recognized as a useful technique for transmitting orbital information and now enjoys wide-spread use over the Air Force Eastern Test Range (AFETR).

When secular rates are not furnished with the mean elements, or when the mean elements have been derived internally, it is necessary to be able to compute these rates as accurately as possible. This can be accomplished analytically if the equations are known or numerically if mean elements at several different epochs are available. The analytical technique will be discussed here, while the numerical is discussed in Section IV.3. The analytical secular rates are computed as functions of the earth’s second zonal gravity harmonic coefficient, $J_2$, and atmospheric drag.

The secular variation of the mean mean anomaly is the mean mean motion. This (perturbed by Kozai’s factor) is given by (Ref. 1):

$$ \frac{\mu e}{\bar{a}^3} \left[ 1 - \frac{3}{2} \frac{J_2 a^2 (1-e^2)^{3/2}}{\bar{p}} \left( 1 - \frac{3}{2} \sin^2 I \right) \right]^{1/2} $$

where $\bar{p}$ is the mean semi-latus rectum given by:

$$ \bar{p} = \bar{a} (1-e^2) $$

The rates of change of the mean argument of perigee and right ascension of ascending node are given by (Ref. 2 and 3):

$$ \dot{\omega} = - \frac{3\bar{a} J_2 a^2 (5 \cos^2 I - 1)}{4\bar{p}^2} $$

$$ \dot{\Omega} = - \frac{3\bar{a} J_2 a^2 \cos I}{2\bar{p}^2} $$

The rate of change of the mean inclination is assumed to be zero. All element rates are mean element rates.
\[
\frac{e}{2} = \left[ \frac{a}{2} \right]^{1+\frac{e}{a}}
\]

(II.12)

\[
\frac{n}{6} = -3n\left[ \frac{a}{2} \right] + 5e\left[ \frac{n}{2} \right]
\]

(II.13)

where

\[
c = \frac{e}{2} \frac{1}{n}
\]

\[
d = Ac\left[ 1 + \frac{n_0}{3(n_0 - n)} \right]
\]

where

\[
n_0 = 16,667 \text{ revs/day}
\]

\[
A = 0, \text{ if } e > 0.06
\]

\[
A = 4, \text{ if } e < 0.06 \text{ and } n < 16.204
\]

\[
A = 13, \text{ if } e < 0.06 \text{ and } n > 16.204
\]

Thus, the mean element secular variations can be obtained analytically for whatever use may be desired of them. The capabilities of this technique can be demonstrated by attempting to reproduce the element rates appearing on a SDC 5-card bulletin. The results of such a test are shown in Table I. The six mean elements \( s, d, a \) were input and the other rates were computed from them. Notice the difference between SDC and analytical values for \( \frac{d}{2} \) and \( \frac{s_0}{2} \). Drag effects should cause negative rather than positive values.

Analytical secular variations are used for trajectory generation in Section IV.2 where it is shown that better results are obtained when \( \frac{N}{2}, \frac{\theta}{2}, \frac{d}{2} \), and \( \frac{s}{6} \) are zero for that particular example. It could be that these terms are necessary for use with objects more significantly affected by drag. In any case, it is probably true that the empirical expression for \( \frac{N}{4} \) could be improved. Analytically and numerically computed secular variations are compared in Section IV.3. Appendix III shows the FORTRAN coding for computing all the secular rates described above. All cases studied have shown that best results are obtained with the above-mentioned acceleration terms set to zero.
### TABLE I

COMPARISONS OF SDC AND ANALYTICALLY COMPUTED SECULAR VARIATIONS

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<th>Parameter</th>
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<tr>
<td>( \ddot{a} ) (earth radii)</td>
<td>1.06351376</td>
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</tr>
<tr>
<td>( \dot{e} )</td>
<td>0.032704</td>
<td>Input</td>
</tr>
<tr>
<td>( \dot{\Omega} ) (deg)</td>
<td>48.3932</td>
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</tr>
<tr>
<td>( \dot{\epsilon} ) (deg)</td>
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</tr>
<tr>
<td>( \dot{\Omega} ) (deg)</td>
<td>247.0671</td>
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</tr>
<tr>
<td>( \dot{\Omega} ) (deg)</td>
<td>233.5949</td>
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<tr>
<td>( \ddot{a} ) (ER/day)</td>
<td>-1.157451xl0^-3</td>
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<tr>
<td>( \ddot{a} ) (ER/day^2)</td>
<td>1.574609xl0^-6</td>
<td>-3.198782xl0^-5</td>
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<tr>
<td>( \ddot{a} ) (ER/day^2)</td>
<td>-1.0527xl0^-3</td>
<td>-1.0527xl0^-3</td>
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<tr>
<td>( \ddot{a} ) (ER/day^2)</td>
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<tr>
<td>( \ddot{a} ) (deg/day^2)</td>
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<tr>
<td>( \ddot{a} ) (deg/day^2)</td>
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<td>4.84601</td>
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<tr>
<td>( \ddot{a} ) (deg/day^2)</td>
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<td>( \ddot{a} ) (deg/day^2)</td>
<td>-5.34388</td>
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<td>15.53805068</td>
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SECTION III
PERIODIC VARIATIONS

III.1 Introduction
The short and long period variations in the classical Keplerian elements can be approximated as functions of the earth's second and third zonal gravity harmonic coefficients. Two methods for computing the short period variations have been used at the AFETR Real Time Computer System (RTCS), one developed by Y. Kozai (Ref 5) and the other by J. B. Frazer (Ref 7). The Frazer method has also been used to compute long period variations. The two methods are presented below along with discussions of their applicability.

III.2 Kozai Periodic Variations
This method has been used at the RTCS to compute short period variations in the Keplerian elements as functions of the earth's second zonal gravity harmonic and is valid for elliptical orbits only. This method computes the short period variations from the mean elements.

First the eccentric anomaly is computed by solving Kepler's equation iteratively:

\[ E^1 = \tilde{E} \quad \text{on first iteration} \]
\[ E = E^1 - \frac{\tilde{E} \sin E^1 - \tilde{M}}{1 - \tilde{e} \cos E^1} \quad \text{on succeeding iterations} \]

The process is converged when two successive estimates are within some epsilon of each other. Then the true anomaly is given by

\[ \tilde{v} = 2 \tan^{-1} \left( \frac{1 + \tilde{e}}{1 - \tilde{e}} \sin \frac{E^1}{2} \right) \frac{\sin \frac{E^1}{2}}{\cos \frac{E^1}{2}} \]

The mean radius from the center of the earth to the object is given by:

\[ \tilde{r} = \tilde{a}(1 - \tilde{e} \cos \tilde{E}) \]
and the mean semi-latus rectum is given by:

\[ \tilde{p} = \tilde{a}(1 - \tilde{e}^2) \]
Kozai applies an extra perturbation to the mean semi-major axis in order to satisfy the following relationship:

$$\frac{n^2 a^3}{\mu_e} = 1 - \frac{3J_2 a^2}{2\mu_e} \left[ 1 - \frac{3}{2} \sin^2 I \right] \left[ 1 - e^2 \right]^2$$

where $a_e$ = earth's semi-major axis
$\mu_e$ = earth's gravity constant

and $\bar{n}$ is the mean mean motion. When computing osculating from mean elements, first the extra perturbation is removed then the short period variations are computed and added to the mean elements. The perturbed mean semi-major axis is given by:

$$\bar{a} = \frac{\bar{n}}{\mu_e} \left( 1 - \frac{3J_2 a^2}{2\mu_e} \left[ 1 - \frac{3}{2} \sin^2 I \right] \left[ 1 - e^2 \right]^2 \right)$$

Then the short period Keplerian element variations are given by (Ref 1):

$$\delta a_s = \frac{J e}{a} \left\{ \frac{2}{3} \left[ 1 - \frac{3}{2} \sin^2 I \right] \left[ \left( \frac{a}{a'} \right)^3 (1 - e^2) - \frac{3}{2} \right] \right\}$$

$$+ \left( \frac{a}{a'} \right)^3 \sin^2 I \cos 2 (\bar{v} + \bar{w})$$

$$\delta e_s = \frac{1 - e^2}{e} \left\{ \frac{J e}{a^2} \left\{ \frac{1}{3} \left[ 1 - \frac{3}{2} \sin^2 I \right] \left[ \left( \frac{a}{a'} \right)^3 (1 - e^2) - \frac{3}{2} \right] \right\} \right. + \frac{1}{2} \left( \frac{a}{a'} \right)^3 \sin^2 I \cos 2 (\bar{v} + \bar{w})$$

$$- \frac{\sin^2 I}{2 e} \left( \frac{J e}{a^2} \right) \cos 2 (\bar{v} + \bar{w}) + \bar{e} \cos (\bar{v} + 2\bar{w})$$

$$+ \frac{1}{3} \bar{e} \cos (3\bar{v} + 2\bar{w})$$

$$\delta i_s = \frac{J e}{b^2} \sin 2I \left[ \cos 2 (\bar{v} + \bar{w}) + \bar{e} \cos (\bar{v} + 2\bar{w}) \right] + \frac{\bar{e}}{3} \cos (3\bar{v} + 2\bar{w})$$

(III.1, III.2, III.3, III.4)
\[
\delta \omega = \frac{J_e}{p^2 \bar{e}} \left\{ \bar{e} \left( 2 - \frac{5}{2} \sin^2 \bar{I} \left( \bar{\nu} - \bar{M} + \bar{e} \sin \bar{\nu} \right) + (1 - \frac{3}{2} \sin^2 \bar{I}) \left( 1 - \frac{e^2}{4} \right) \sin \bar{\nu} + \frac{e^2}{2} \sin 2 \bar{\nu} \right) + \frac{e^2}{12} \sin 3 \bar{\nu} \right\} \\
- \sin (\bar{\nu} + 2 \bar{\omega}) \left[ \frac{1}{4} \sin^2 \bar{I} + \bar{e}^2 \left( \frac{1}{2} - \frac{15}{16} \sin^2 \bar{I} \right) \right] + \frac{\bar{e}^2}{16} \sin^2 \bar{I} \sin (\bar{\nu} - 2 \bar{\omega}) - \frac{e^2}{2} (1 - \frac{5}{2} \sin^2 \bar{I}) \sin 2 (\bar{\nu} + \bar{\omega}) + \sin (3 \bar{\nu} + 2 \bar{\omega}) \left[ \frac{7}{12} \sin^2 \bar{I} - \frac{\bar{e}^2}{6} (1 - \frac{19}{8} \sin^2 \bar{I}) \right] + \frac{3 \bar{e}^2}{8} \sin^2 \bar{I} \sin (4 \bar{\nu} + 2 \bar{\omega}) + \frac{\bar{e}^2}{16} \sin^2 \bar{I} \sin (5 \bar{\nu} + 2 \bar{\omega}) \right\} \tag{III.5}
\]

\[
\delta \Omega = - \frac{J_e}{p^2 \bar{e}} \cos \bar{I} \left\{ \left( \bar{\nu} - \bar{M} + \bar{e} \sin \bar{\nu} \right) - \frac{1}{2} \sin 2 (\bar{\nu} + \bar{\omega}) - \bar{e} \sin (\bar{\nu} + 2 \bar{\omega}) - \frac{e}{6} \sin (3 \bar{\nu} + 2 \bar{\omega}) \right\} \tag{III.6}
\]

\[
\delta \bar{M} = \frac{J_e}{p^2 \bar{e}} \left( 1 - \bar{e}^2 \right) \left\{ (1 - \frac{3}{2} \sin^2 \bar{I}) \left[ (1 - \frac{\bar{e}^2}{4}) \sin \bar{\nu} + \frac{\bar{e}^2}{2} \sin 2 \bar{\nu} + \frac{\bar{e}^2}{12} \sin 3 \bar{\nu} \right] + \frac{1}{4} (1 + \frac{5 \bar{e}^2}{4}) \sin (\bar{\nu} + 2 \bar{\omega}) - \frac{\bar{e}^2}{16} \sin (\bar{\nu} - 2 \bar{\omega}) \right. \\
- \frac{7}{12} (1 - \frac{\bar{e}^2}{28}) \sin (3 \bar{\nu} + 2 \bar{\omega}) - \frac{3 \bar{e}^2}{8} \sin (4 \bar{\nu} + 2 \bar{\omega}) - \frac{\bar{e}^2}{16} \sin (5 \bar{\nu} + 2 \bar{\omega}) \left. \right\} \tag{III.7}
\]

where \( J_e = \frac{3}{2} J \ \bar{a}^2 \).

Notice that in the above expressions for \( \delta \omega_s \) and \( \delta \bar{M}_s \) there are eccentricity divisors. The variations in these elements are undefined for circular orbits because the elements themselves are undefined.

These variations are added to the mean elements to yield quasi-osculating (short period variations only).
Appendix IV gives the FORTRAN coding for this algorithm and Figures 1 through 6 show theoretical results obtained by the method. The figures show the classical Keplerian elements and quasi mean elements (short period variations removed) for one orbital revolution of a satellite. An orbit was numerically generated with initial conditions in the classical elements given in Table II, with no drag and using the gravity model shown in Table III. The Kozai algorithm was applied to points taken from the theoretical trajectory at five minute intervals. In this case osculating elements were being transformed to quasi mean elements. Since the variation equations are functions of the mean elements, the solution was performed iteratively. Secular variations in the elements were removed for ease of plotting.

Due to the problems of this algorithm associated with near circular orbits, this method is not recommended for general orbit support.

The effect of Kozai's extra perturbation can be seen in Figure 1, where the quasi mean is displaced from the center of the plot of the osculating semi-major axis.
TABLE II
THEORETICAL TRAJECTORY INITIAL CONDITIONS

Classical Keplerian Elements

<table>
<thead>
<tr>
<th>Symbol</th>
<th>Value</th>
</tr>
</thead>
<tbody>
<tr>
<td>a</td>
<td>22981800. feet</td>
</tr>
<tr>
<td>e</td>
<td>.02</td>
</tr>
<tr>
<td>i</td>
<td>28. deg</td>
</tr>
<tr>
<td>ω</td>
<td>9. deg</td>
</tr>
<tr>
<td>Ω</td>
<td>115. deg</td>
</tr>
<tr>
<td>M</td>
<td>0. deg</td>
</tr>
</tbody>
</table>
TABLE III
THEORETICAL TRAJECTORY GRAVITY MODEL

\[ a_e = 6378.165 \text{ km} \]
\[ \mu_e = 398601.2 \text{ km}^3/\text{sec}^2 \]

<table>
<thead>
<tr>
<th>( n )</th>
<th>( m )</th>
<th>( C_{\text{nm}} )</th>
<th>( S_{\text{nm}} )</th>
</tr>
</thead>
<tbody>
<tr>
<td>2</td>
<td>0</td>
<td>-1082.3</td>
<td>0</td>
</tr>
<tr>
<td>3</td>
<td>0</td>
<td>2.3</td>
<td>0</td>
</tr>
<tr>
<td>4</td>
<td>0</td>
<td>1.8</td>
<td>0</td>
</tr>
<tr>
<td>2</td>
<td>1</td>
<td>0</td>
<td>0</td>
</tr>
<tr>
<td>2</td>
<td>2</td>
<td>1.68</td>
<td>-0.64</td>
</tr>
<tr>
<td>3</td>
<td>1</td>
<td>1.77</td>
<td>0.19</td>
</tr>
<tr>
<td>3</td>
<td>2</td>
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<tr>
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<td>3</td>
<td>0.15</td>
<td>0.14</td>
</tr>
<tr>
<td>4</td>
<td>1</td>
<td>-0.57</td>
<td>-0.46</td>
</tr>
<tr>
<td>4</td>
<td>2</td>
<td>0.06</td>
<td>0.26</td>
</tr>
<tr>
<td>4</td>
<td>3</td>
<td>0.08</td>
<td>-0.003</td>
</tr>
<tr>
<td>4</td>
<td>4</td>
<td>-0.008</td>
<td>0.006</td>
</tr>
</tbody>
</table>

*Multiply all values shown by \( 10^{-6} \)
Figure 1. Kozai Semi-Major Axis Short Period Variations

Figure 2. Kozai Eccentricity Short Period Variations
Figure 3. Kozai Inclination Short Period Variations

Figure 4. Kozai Argument of Perigee Short Period Variations
Figure 5. Kozai Right Ascension of Ascending Node Short Period Variations

Figure 6. Kozai Mean Anomaly Short Period Variations
III.3 Frazer Periodic Variations

III.3.1 Introduction

The method developed by J. B. Frazer (Ref. 7) does not have a problem with circular orbits as does the Kozai because the variations are computed in Cartesian elements. This method is valid for circular and elliptical orbits only and as described herein is invalid for equatorial orbits due to sine inclination divisors.

First, the mean Keplerian elements are transformed to mean Cartesian elements as follows:

\[
\mathbf{r} = \mathbf{r} + \mathbf{r} \dot{V} \mathbf{V}
\]

where

\[
\mathbf{r} = \begin{bmatrix} x \\ y \\ z \end{bmatrix}, \quad \mathbf{r} = \begin{bmatrix} x' \\ y' \\ z' \end{bmatrix}
\]

The osculating Cartesian system is defined as follows: x, y, z are position components in an inertial geocentric righthanded system with x and y in the plane of the mean equator of date, x toward the mean Vernal equinox of date, and z toward the north pole ("mean" here refers to the absence of nutations).

\[
\mathbf{\dot{U}} = \begin{bmatrix} \dot{x} \\ \dot{y} \\ \dot{z} \end{bmatrix} = \begin{bmatrix} \cos \tilde{u} \cos \tilde{\Omega} - \sin \tilde{u} \sin \tilde{\Omega} \cos \tilde{I} \\ \cos \tilde{u} \sin \tilde{\Omega} + \sin \tilde{u} \cos \tilde{\Omega} \cos \tilde{I} \\ \sin \tilde{u} \sin \tilde{I} \end{bmatrix}
\]

\[
\mathbf{\dot{V}} = \begin{bmatrix} \dot{V} \\ \dot{V} \\ \dot{W} \end{bmatrix} = \begin{bmatrix} -\sin \tilde{u} \cos \tilde{\Omega} - \cos \tilde{u} \sin \tilde{\Omega} \cos \tilde{I} \\ -\sin \tilde{u} \sin \tilde{\Omega} + \cos \tilde{u} \cos \tilde{\Omega} \cos \tilde{I} \\ \cos \tilde{u} \sin \tilde{I} \end{bmatrix}
\]

where \( \tilde{u} = \) mean argument of latitude = \( \tilde{\omega} + \tilde{\nu} \) and where \( \mathbf{r} \) is the radius from the earth's center to the object given by

\[
\mathbf{r} = \frac{p}{1+e \cos \nu}
\]

\( \mathbf{\dot{r}} \) is the rate of change of \( \mathbf{r} \) given by

\[
\mathbf{\dot{r}} = \frac{\dot{u} - \frac{h}{p}}{\mathbf{e} \sin \nu} \mathbf{r}
\]

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and \( rv \) is the product of the radius and the true anomaly rate of change given by

\[
rv = \frac{\dot{r}}{p} (1 + e \cos \nu)
\]

where \( p \) = mean semi-latus rectum = \( a(1-e^2) \)

The perturbed or osculating Cartesian vector due to short and long period variations is given by

\[
\dot{r} = (\dot{r} + \delta r)(\ddot{u} + \delta \dot{u})
\]

\[
\dot{r} = (\dot{u} + \delta u)(\ddot{u} + \delta \dot{u}) + (\ddot{r} + \delta r)(\dot{u} + \delta \dot{u})
\]

or the variations alone are given by

\[
\delta \dot{r} = \delta \dot{u}^2 + \delta \dot{u} \delta r^2
\]

\[
\delta \dot{x} = \delta \dot{x} u + \delta \dot{u} \delta r u^2 + \delta r u \dot{u}^2
\]

Using

\[
\begin{bmatrix}
\dot{x} \\
\dot{y} \\
\dot{z}
\end{bmatrix} = \begin{bmatrix}
\sin u \sin \iota \\
-\cos u \sin \iota \\
\cos \iota
\end{bmatrix}
\]

then

\[
\delta \dot{u} = (\delta u + \delta u \cos \iota) \dot{u} + (\delta \dot{u} - \delta \dot{u} \cos \iota) \dot{u}
\]

\[
\delta \dot{u} = - (\delta u + \delta u \cos \iota) \dot{u} + (\delta \dot{u} + \delta \dot{u} \cos \iota) \dot{u}
\]

and hence

\[
\delta \dot{r} = \delta \dot{r} \dot{u} + \dot{r}(\delta u + \delta u \cos \iota) \dot{u}
\]

\[
+ \dot{r}(\delta \dot{u} - \delta \dot{u} \cos \iota) \dot{u}
\]

\[
\delta \dot{r} = \left[ \delta \dot{r} - \dot{r}(\delta u + \delta u \cos \iota) \dot{u} \right.
\]

\[
+ \dot{r}(\delta \dot{u} - \delta \dot{u} \cos \iota) \dot{u}
\]

+ \left. \dot{r}(\delta \dot{u} \cos \iota + \delta \dot{u} \sin \iota) \right] \dot{u}
\]

These variations are composed of short and long period components as

\[
\delta \dot{r} = \delta \dot{r}_s + \delta \dot{r}_L
\]

\[
\dot{r} = \dot{r}_s + \dot{r}_L
\]

which are computed separately as shown in the following sections. Once the short and long period variations have been computed, osculating Cartesian elements can be obtained. The osculating Cartesian elements can then be transformed into osculating Keplerian elements by standard techniques.
III.3.2 Frazer Long Period Variations

Although Frazer develops the variations in a Cartesian coordinate system, it is interesting to note the forms of the equations for the long period variations in the Keplerian elements (Ref. 7). For all zonal harmonics, $J_n$, $n>2$:

\[
\begin{align*}
\delta a_L &= 0 \\
\delta e_L &= S_n(1-e^2)\sin i \cos \xi \\
\delta i_L &= -S_n\bar{e} \cos i \cos \xi \\
\delta \omega_L &= -S_n\bar{e} \sin i \left[ 5 - 2n - \frac{1 - e^2}{e^2} - \frac{10\cos^2 I}{1 - 5\cos^2 I} \right] \\
&\quad + \frac{\nu \cos^2 I}{\sin^2 I} \sin \lambda \xi \\
\delta \Omega_L &= -S_n\bar{e} \sin i \cos i \left[ -\frac{10}{1 - 5\cos^2 I} - \frac{\nu}{\sin^2 I} \right] \sin \lambda \xi \\
\delta M_L &= -S_n \frac{(1-e^2)^2}{e} \sin i \mu \sin \xi \\
\end{align*}
\]

where

\[
S_n = -\frac{8(n-1)!}{3\cdot 2} \frac{J_n a^{n-2}}{J_0 b^{n-2}(1-5\cos^2 I)} \sum_{\lambda, \mu, \nu} Q
\]

\[
Q = \frac{\lambda - \nu}{2} \frac{(n-\nu)!}{(n+\nu)!} \frac{\mu-1}{\sin \frac{I}{2}} \frac{\nu-1}{(n-\mu-1)!} \frac{(n-\mu)!}{(n+\nu)!} \frac{(\lambda-\nu)!}{(\lambda+\nu)!} \frac{(\lambda+\nu)!}{(\lambda-\nu)!}
\]

\[
\sin \lambda \xi = (-1)^{\frac{\nu}{2}} \sin \lambda \bar{w}, \lambda \text{ even}
\]

\[
\sin \lambda \xi = -(-1)^{\frac{\mu-1}{2}} \cos \lambda \bar{w}, \lambda \text{ odd}
\]

\[
\cos \lambda \xi = (-1)^{\frac{\nu}{2}} \cos \lambda \bar{w}, \lambda \text{ even}
\]

\[
\cos \lambda \xi = (-1)^{\frac{\mu-1}{2}} \sin \lambda \bar{w}, \lambda \text{ odd}
\]

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As can be seen in the above equations, there is no first order long period variation in the semi-major axis. Since there is an eccentricity multiplier in the expressions for the long period variations in inclination and right ascension of the ascending node, these variations are insignificant for near-circular orbits. There is also an eccentricity multiplier in \( S_n \) for all \( n \) except \( n=3(J_3) \). For near-circular orbits, then, \( J_3 \) is the only zonal gravity harmonic coefficient which contributes significantly to long period variations, and then only in eccentricity, argument of perigee and mean anomaly. But argument of perigee, mean anomaly and their variations (note eccentricity divisors) are undefined for circular orbits.

In the above equation for \( S_n \) there is a divisor given as

\[ 1-5\cos^2I \]

This expression becomes zero when inclination is approximately 63.5 or 116.5 degrees. This is known as the critical inclination. At the critical inclination the long period variations in the orbital elements are undefined.

After considerable manipulation, Frazer reduces the long period variations as functions of the zonal harmonic coefficient \( J_3 \) to the following in the inertial geocentric Cartesian elements (Ref. 7):

\[
\delta r_L = \bar{r} a_2 \left[ \sin I (1+\epsilon \cos v) \sin \bar{u} 
+ \sin I (2+\epsilon \cos v) \cos \bar{v} + \cos I \bar{e} \cos v \bar{w} \right] \quad (III.10)
\]

\[
\delta \bar{r} = -\left(\frac{\mu \epsilon}{\bar{p}}\right) \frac{1}{a_2} \left[ \sin I (1+\epsilon \cos v) \cos \bar{u} 
+ \sin I (\sin \bar{u} + \bar{e} \sin \bar{w}) \bar{v} + \cos I \bar{e} \sin \bar{w} \right] \quad (III.11)
\]

where

\[
a_2 = \frac{J_3 \epsilon}{2J_2 \bar{p}}
\]

\[
\lambda = 2,4,\ldots,(n-2) \quad n \text{ even}
\]

\[
= 1,3,\ldots,(n-2) \quad n \text{ odd}
\]

\[
\nu = \lambda, \lambda+2,\ldots, (n-2)
\]

\[
\nu = \lambda, \lambda+2,\ldots, n
\]

\[ a_e = \text{earth's semi-major axis} \]
Notice that in the expressions for the Cartesian variations due to \( J_3 \), the effect of the critical inclination does not appear.

Appendix V gives the FORTRAN coding for the Frazer algorithm (includes short period variations) and Figures 7 through 13 show theoretical results obtained in the Keplerian elements by applying the long period variations as described above. The example used is the same as that previously given for the Kozai method in Tables II and III. The mean longitude variations, \( \delta L_L \), given in Figure 13 are defined by:

\[
\delta L_L = \delta \omega_L + \delta M_L
\]

The figures show mean elements and quasi mean elements (short period variations removed) for twenty days of satellite motion. Secular variations were removed for ease of plotting.

Originally, long period variations due to \( J_2 \) and \( J_4 \) were included in the computations, but numerical results obtained when comparing trajectory generations via mean elements versus numerical integration (Section IV.2) were much poorer than those obtained with \( J_2 \) and \( J_4 \) effects eliminated.

If the computations of these variations are restricted to near-circular orbits, the above equations (with \( \tilde{e}=0 \)) reduce to:

\[
\begin{align*}
\delta \varpi_L &= \alpha_2 \, \tilde{r} \sin \tilde{I} \, (\sin \tilde{U} + 2 \cos \tilde{V}) \\
\delta T_L &= -\alpha_2 \left( \frac{\tilde{u}}{\tilde{p}} \right)^{1/2} \sin \tilde{I} \, (\cos \tilde{U} + \sin \tilde{V})
\end{align*}
\]

These are the resulting approximate equations even when \( J_2 \) and \( J_4 \) are included in the full equations.

These approximate equations yield long period variations with errors varying as a function of the eccentricity. Table IV shows some typical errors, encountered at various eccentricities, caused by using the approximate rather than the full equations. Since these are just isolated examples, they must not by any means be considered as maximum errors which could be obtained. The errors given are the differences between elements obtained after applying the equations III.10-11 and the approximate equations above to an input vector.
Figure 7. Frazer Semi-Major Axis Long Period Variations

Figure 8. Frazer Eccentricity Long Period Variations
Figure 9. Frazer Inclination Long Period Variations

Figure 10. Frazer Argument of Perigee Long Period Variations
Figure 11. Frazer Right Ascension of Ascending Node Long Period Variations

Figure 12. Frazer Mean Anomaly Long Period Variations
Figure 13. Frazer Mean Longitude Long Period Variations
TABLE IV

Examples of Errors Caused by Using Approximate Long Period Variations

<table>
<thead>
<tr>
<th></th>
<th>Eccentricity</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td>0.02</td>
</tr>
<tr>
<td>$\delta x$ (ft.)</td>
<td>0.</td>
</tr>
<tr>
<td>$\delta y$ (ft.)</td>
<td>0.</td>
</tr>
<tr>
<td>$\delta z$ (ft.)</td>
<td>207.</td>
</tr>
<tr>
<td>$\delta x$ (ft./sec.)</td>
<td>0.18</td>
</tr>
<tr>
<td>$\delta y$ (ft./sec.)</td>
<td>0.04</td>
</tr>
<tr>
<td>$\delta z$ (ft./sec.)</td>
<td>0.32</td>
</tr>
<tr>
<td>$\delta a$ (ft.)</td>
<td>249.</td>
</tr>
<tr>
<td>$\delta e$</td>
<td>0.000004</td>
</tr>
<tr>
<td>$\delta i$ (deg.)</td>
<td>0.002</td>
</tr>
<tr>
<td>$\delta w$ (deg.)</td>
<td>0.0233</td>
</tr>
<tr>
<td>$\delta \Omega$ (deg.)</td>
<td>0.0019</td>
</tr>
<tr>
<td>$\delta M$ (deg.)</td>
<td>0.0215</td>
</tr>
</tbody>
</table>
III.3.3 Frazer Short Period Variations

The short period variations are computed as functions of the earth's second gravity zonal harmonic coefficient in the elements \( r, \bar{u}, \bar{N}, I, \bar{e}, \bar{\nu}, \) and then these variations are transformed to the inertial Cartesian system as shown in Section III.3.1. That is (Ref. 7):

\[
\delta r_s = a_0 \left\{ \sin^2 \bar{I} \cos 2\bar{u} + (1 - 3 \cos^2 \bar{I}) \left[ 1 - \frac{(1 - \bar{e}^2)^{1/2}}{1 + \bar{e} \cos \bar{\nu}} \right] \right. \\
+ \left. \frac{\bar{e} \cos \bar{\nu}}{1 + (1 - \bar{e}^2)^{1/2}} \right\} 
\]

\[
(III.12)
\]

\[
\delta \bar{u}_s = -a_1 \left\{ 2 \sin^2 \bar{I} (1 + \bar{e} \cos \bar{\nu})^2 \sin 2\bar{u} \\
+ (1 - 3 \cos^2 \bar{I}) \bar{e} \sin \bar{\nu} \left[ \frac{(1 - \bar{e}^2)^{1/2}}{2} + \frac{(1 + \bar{e} \cos \bar{\nu})^2}{1 + (1 - \bar{e}^2)^{1/2}} \right] \right\} 
\]

\[
(III.13)
\]

\[
\delta \bar{v}_s = a_1 \left\{ \sin^2 \bar{I} \left[ 2 \cos 2\bar{u} + 2 \bar{e} \cos (2\bar{u} - \bar{\nu}) \right] \\
+ \bar{e} \cos \bar{\nu} \cos 2\bar{u} \left( 1 + \bar{e} \cos \bar{\nu} \right) \\
- (1 + \bar{e} \cos \bar{\nu}) (1 - 3 \cos^2 \bar{I}) \left[ \frac{3}{2} + \bar{e} \cos \bar{\nu} \right] \frac{2 + (1 - \bar{e}^2)^{1/2}}{1 + (1 - \bar{e}^2)^{1/2}} \\
+ \frac{\bar{e}^2 - 2 (\bar{e} \sin \bar{\nu})^2}{2 [1 + (1 - \bar{e}^2)^{1/2}]} \right\} 
\]

\[
(III.14)
\]

\[
\delta \bar{I}_s = a \sin \bar{I} \cos \bar{I} \left\{ 3 \cos 2\bar{u} + \bar{e} \cos (2\bar{u} - \bar{\nu}) \right\} 
\]

\[
(III.15)
\]

\[
\delta \bar{\nu}_s = -a \cos \bar{I} \left\{ 6 (\bar{\nu} - \bar{\nu} + \bar{e} \sin \bar{\nu}) \\
- 3 \left[ \sin 2\bar{u} + \bar{e} \sin (2\bar{u} - \bar{\nu}) \right] - \bar{e} \sin (2\bar{u} + \bar{\nu}) \right\} 
\]

\[
(III.16)
\]
\[ \delta u_s = -\frac{a}{2} \left( 6(1-5\cos^2 I)(\tilde{v}-\tilde{N}) + 4 \left[ 1-6\cos^2 I + \frac{1-3\cos^2 I}{1+(1-e^2)^{\frac{1}{2}}} \right] \bar{e}\sin v \right. \\
+ \left. (1-3\cos^2 I) \frac{2(\bar{e}\sin v)(\bar{e}\cos v)}{1+(1-e^2)^{\frac{1}{2}}} \right) + 2(5\cos^2 I-2) \bar{e}\sin (2\bar{v}-v) \\
+ (7\cos^2 I-1)\sin 2\bar{u} + 2\cos^2 I \bar{e}\sin(2\bar{u}-\bar{v}) \right) \tag{III.17} \]

where

\[ a = \frac{J_2 a^2}{4r^2} \]

\[ a_1 = a \left[ \frac{\mu_e}{P} \right] \]

The difference between the true and mean anomalies needed to compute \( \delta N_s \) and \( \delta u_s \) is given by:

\[ (\bar{v}-\tilde{N}) = (\bar{v}-\bar{E}) + (\bar{E}-\tilde{N}) \]

where \( \bar{E} \) is the eccentric anomaly and (Ref. 8):

\[ (\bar{v}-\bar{E}) = \sin^{-1} \left[ \left( \frac{\bar{e}\sin v}{1+\bar{e}\cos v} \right) \left( \frac{0-e^2}{1+\bar{e}\cos v} \right) \right] \]

and

\[ (\bar{E}-\tilde{N}) = \frac{(1-e^2)^{\frac{1}{2}} \bar{e}\sin v}{1+\bar{e}\cos v} \]

These short period variations (specifically \( \delta r_s \), \( \delta I_s \) and \( \delta v_s \)) include the effect of Kozai's perturbed semi-major axis (Ref. 1) as given previously in Section III.2. This additional variation does not have to be accounted for separately.
Appendix V gives the FORTRAN coding for the Frazer algorithm (includes long period variations) and Figures 14 through 20 show theoretical results obtained in the Keplerian elements by applying the short period portion as described above. The example used is the same as that previously given for the Kozai method in Tables II and III. The mean longitude variations, $\delta L_s$, given in Figure 20 are defined by:

$$\delta L_s = \delta \omega_s + \delta M_s$$

The figures show quasi-mean (short period variations removed) and classical Keplerian elements for one orbital revolution of a satellite. Secular variations were removed for ease of plotting.

If the computations of these variations are restricted to near-circular orbits, the above equations (with $\epsilon=0$) reduce to:

$$\delta r_s = \alpha r \sin^2 i \cos 2u$$
$$\delta t_s = -2 \alpha i \sin^2 i \sin 2u$$
$$\delta \nu_s = \alpha i \left[2 \sin^2 i \cos 2u - \frac{3}{2} \left(1 - 3 \cos^2 i\right)\right]$$
$$\delta i_s = 3 \alpha \sin i \cos i \cos 2u$$
$$\delta \Omega_s = 3 \alpha \cos i \sin 2u$$
$$\delta u_s = -\frac{\alpha}{2} \left(7 \cos^2 i - 1\right) \sin 2u$$

These approximate equations yield short period variations with errors varying as a function of the eccentricity. Table V shows some typical errors, encountered at various eccentricities, caused by using the approximate rather than the full equations. Since these are just isolated examples, they must not by any means be considered as maximum errors which could be obtained. The errors given are the differences between elements obtained after applying the equations III.12-17 and the approximate equations above to an input vector.
Figure 14. Frazer Semi-Major Axis Short Period Variations

Figure 15. Frazer Eccentricity Short Period Variations
Figure 16. Frazer Inclination Short Period Variations

Figure 17. Frazer Argument of Perigee Short Period Variations
Figure 18. Frazer Right Ascension of Ascending Node Short Period Variations

Figure 19. Frazer Mean Anomaly Short Period Variations
Figure 20. Frazer Mean Longitude Short Period Variations
TABLE V
Examples of Errors Caused by Using Approximate Short Period Variations

<table>
<thead>
<tr>
<th></th>
<th>Eccentricity</th>
</tr>
</thead>
<tbody>
<tr>
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<td>0.02</td>
</tr>
<tr>
<td>δr (ft.)</td>
<td>98.</td>
</tr>
<tr>
<td>δf (ft./sec.)</td>
<td>0.05</td>
</tr>
<tr>
<td>δr' (ft./sec.)</td>
<td>0.26</td>
</tr>
<tr>
<td>δι (deg.)</td>
<td>0.0002</td>
</tr>
<tr>
<td>δΩ (deg.)</td>
<td>0.0030</td>
</tr>
<tr>
<td>δω (deg.)</td>
<td>0.0053</td>
</tr>
<tr>
<td>δx (ft.)</td>
<td>628.</td>
</tr>
<tr>
<td>δy (ft.)</td>
<td>633.</td>
</tr>
<tr>
<td>δz (ft.)</td>
<td>725.</td>
</tr>
<tr>
<td>δx (ft./sec.)</td>
<td>0.04</td>
</tr>
<tr>
<td>δy (ft./sec.)</td>
<td>1.12</td>
</tr>
<tr>
<td>δz (ft./sec.)</td>
<td>0.62</td>
</tr>
<tr>
<td>δa (ft.)</td>
<td>257.</td>
</tr>
<tr>
<td>δe</td>
<td>0.000006</td>
</tr>
<tr>
<td>δω (deg.)</td>
<td>0.0482</td>
</tr>
<tr>
<td>δM (deg.)</td>
<td>0.0426</td>
</tr>
</tbody>
</table>
III.4 Inverse Computations

In both the Kozai and Frazer forms of the variation equations, the variations are computed as functions of the mean elements. If it is desired to transform osculating to mean elements the equations cannot be used directly. Instead, they are used in an iterative sense as described below.

First an approximate set of mean elements is obtained (usually the osculating elements). These are used to compute variations and approximate osculating elements are obtained and compared to the known osculating elements. The differences between the two sets are the errors in the computed osculating elements caused by errors in the approximate mean elements. These mean elements are then corrected by the amounts of the differences concluding the first iteration. This process is repeated until some convergence criteria are satisfied.

When using the Frazer method and convergence criteria of 1.0 feet and 0.001 ft/sec, convergence is almost always achieved in three iterations. When using the Kozai method, convergence is sometimes difficult to achieve for near-circular orbits. After forty iterations, significant errors may still be present in argument of perigee and mean anomaly. However, a method involving matrix inversion (Ref. 9) can be used to yield quick solutions for the Kozai method.

Appendix VI shows the FORTRAN coding which performs the iteration control when going from osculating to mean elements with the Frazer method. Table VI shows the results of using the iterative technique to compute mean from osculating elements using the Frazer method. Although in the example shown, the long period variations are larger than the short period in the Cartesian elements, the reverse is true in the Keplerian elements. Also, the long period variations in the Cartesian elements make no significant contribution toward variation in the semi-major axis.
### Table VI

**Example of Frazer Inverse Computations**

<table>
<thead>
<tr>
<th>INITIAL OSCULATING VECTORS</th>
<th>(a)</th>
<th>(e)</th>
<th>(i)</th>
<th>(\omega)</th>
<th>(\Omega)</th>
<th>(M)</th>
</tr>
</thead>
<tbody>
<tr>
<td>22974394</td>
<td>.019253</td>
<td>27.984616</td>
<td>16.819571</td>
<td>108.579818</td>
<td>301.112438</td>
<td></td>
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</tbody>
</table>

<table>
<thead>
<tr>
<th></th>
<th>(x)</th>
<th>(y)</th>
<th>(z)</th>
<th>(\dot{x})</th>
<th>(\dot{y})</th>
<th>(\dot{z})</th>
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</thead>
<tbody>
<tr>
<td>(x)</td>
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<td>19964142.</td>
<td>-7413602.</td>
<td>-20727.6632</td>
<td>11024.1029</td>
<td>8573.4654</td>
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</tbody>
</table>

#### FIRST ITERATION

**VAARIATIONS**

<table>
<thead>
<tr>
<th></th>
<th>SHORT PERIOD</th>
<th>LONG PERIOD</th>
<th>TOTAL</th>
</tr>
</thead>
<tbody>
<tr>
<td>SHORT PERIOD</td>
<td>3068.</td>
<td>-194.</td>
<td>2968.</td>
</tr>
<tr>
<td>LONG PERIOD</td>
<td>14777.</td>
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<td>-7568.</td>
</tr>
<tr>
<td>TOTAL</td>
<td>17846.</td>
<td>-676.</td>
<td>-4600.</td>
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</tbody>
</table>

**NEW ESTIMATED MEAN VECTOR**

|                        | 7990451. | 19964818. | -7409002. | -20730.7911 | 11010.8306 | 8572.5043 |

#### SECOND ITERATION

**VAARIATIONS**

<table>
<thead>
<tr>
<th></th>
<th>SHORT PERIOD</th>
<th>LONG PERIOD</th>
<th>TOTAL</th>
</tr>
</thead>
<tbody>
<tr>
<td>SHORT PERIOD</td>
<td>3064.</td>
<td>-179.</td>
<td>2978.</td>
</tr>
<tr>
<td>LONG PERIOD</td>
<td>14789.</td>
<td>-473.</td>
<td>-7561.</td>
</tr>
<tr>
<td>TOTAL</td>
<td>17852.</td>
<td>-652.</td>
<td>-4583.</td>
</tr>
</tbody>
</table>

**OSCULATING ERRORS**

|                        | -6.         | -23.        | -17.   | 0.0099  | -0.0052 | -0.0200 |

**NEW ESTIMATED MEAN VECTOR**

|                        | 7990445. | 19964794. | -7409019. | -20730.7812 | 11010.8255 | 8572.4844 |

#### THIRD ITERATION

**VAARIATIONS**

<table>
<thead>
<tr>
<th></th>
<th>SHORT PERIOD</th>
<th>LONG PERIOD</th>
<th>TOTAL</th>
</tr>
</thead>
<tbody>
<tr>
<td>SHORT PERIOD</td>
<td>3063.</td>
<td>-179.</td>
<td>2978.</td>
</tr>
<tr>
<td>LONG PERIOD</td>
<td>14789.</td>
<td>-473.</td>
<td>-7561.</td>
</tr>
<tr>
<td>TOTAL</td>
<td>17852.</td>
<td>-652.</td>
<td>-4583.</td>
</tr>
</tbody>
</table>

**OSCULATING ERRORS**

|                        | -0.04       | -0.04       | -0.04  | -0.0002 | -0.0002 | 0.0005  |

<table>
<thead>
<tr>
<th></th>
<th>(\bar{a})</th>
<th>(\bar{e})</th>
<th>(\bar{i})</th>
<th>(\bar{\omega})</th>
<th>(\bar{\Omega})</th>
<th>(\bar{M})</th>
</tr>
</thead>
<tbody>
<tr>
<td>(\bar{a})</td>
<td>22952977.</td>
<td>0.018754</td>
<td>27.984143</td>
<td>18.085366</td>
<td>108.609094</td>
<td>299.836032</td>
</tr>
</tbody>
</table>
SECTION IV
APPLICATIONS

IV.1 Introduction

The preceding sections have presented the mathematical expressions describing variations in the orbital elements and have shown numerical examples obtained from the applications of those expressions. The purpose of this section is to show how knowledge of these variations can be used to advantage in orbital support. At the ATR RTCS, two major uses of these variations have been exploited. These two uses are described below.

IV.2 Trajectory Generation

The major use to which element variations have been applied at the RTCS is that pertaining to trajectory generation. This is based upon the premise that, given elements at epoch, elements at any time can be obtained if the element variations can be computed. This can be easily accomplished using the mathematical expressions given in the preceding sections.

For example, suppose that the position and velocity of an object is given in the osculating geocentric inertial Cartesian system at epoch and the position and velocity is desired in the same system at some other time. One way to accomplish this trajectory generation would be by numerical integration of the equations of motion where accelerations in the Cartesian system are computed based on models of the earth's gravity and atmosphere. But such a trajectory generation can be accomplished through element variations, also. Such a generation using the methods presented in the preceding sections would require a four-step process. First, mean elements would be obtained from the osculating (periodic variations removed) using inverse computations (Section III.4). Second, element rates (secular variations) would be computed (Section II). Third, these rates would be applied to the mean elements at epoch to obtain the mean elements at the time of interest. Fourth, the mean elements at the time of interest are converted to osculating (periodic variations added) thus completing the trajectory generation. This procedure is shown pictorially by:
Mathematical expressions for accomplishing the first, second and fourth parts above have been presented. The third part, comprising the update of the mean elements from epoch, \( t_0 \), to time of interest, \( t \), is described below. Often, the mean elements and their rates at epoch are given in the form of an Air Force Space Defense Center 5-card element set (SDC bulletin). Information contained on this bulletin includes the following mean elements and their time derivatives at epoch (see Appendix I):

\[
\begin{align*}
\bar{a} &= \text{mean semi-major axis} \\
\bar{\dot{a}} &= \text{derivative of } \bar{a} \\
\bar{\frac{\Delta}{2}} &= \text{half of derivative of } \bar{\dot{a}} \\
\bar{e} &= \text{mean eccentricity} \\
\bar{\dot{e}} &= \text{derivative of } \bar{e} \\
\bar{\frac{\Delta}{2}} &= \text{half of derivative of } \bar{\dot{e}} \\
\bar{\Omega} &= \text{mean inclination} \\
\bar{\dot{\Omega}} &= \text{derivative of } \bar{\Omega} \\
\bar{\omega} &= \text{mean argument of perigee} \\
\bar{\dot{\omega}} &= \text{derivative of } \bar{\omega} \\
\bar{\frac{\Delta}{2}} &= \text{half of derivative of } \bar{\dot{\omega}} \\
\bar{M} &= \text{mean right ascension of ascending node} \\
\bar{\dot{M}} &= \text{derivative of } \bar{\dot{M}} \\
\bar{\frac{\Delta}{2}} &= \text{half of derivative of } \bar{\dot{M}} \\
\bar{\dot{\gamma}} &= \text{mean mean anomaly} \\
\bar{\dot{\gamma}} &= \text{mean mean motion (derivative of } \bar{\dot{\gamma}}) \\
\bar{\Delta} &= \text{half of derivative of } \bar{\dot{\gamma}}
\end{align*}
\]
As can be seen, these bulletin parameters are nothing more than the coefficients of a Maclaurin's series expansion about epoch. The mean element update can then be accomplished by:

\[ \Delta t = t - t_0 \]
\[ \bar{a} = \bar{a}_0 + \dot{a} \Delta t + \frac{a_2}{12} \Delta t^2 \]  
\[ \bar{e} = \bar{e}_0 + \dot{e} \Delta t + \frac{e_2}{12} \Delta t^2 \]  
\[ \bar{I} = \bar{I}_0 + \dot{I} \Delta t \]  
\[ \bar{\omega} = \bar{\omega}_0 + \dot{\omega} \Delta t + \frac{\omega_2}{12} \Delta t^2 \]  
\[ \bar{n} = \bar{n}_0 + \dot{n} \Delta t + \frac{n_2}{24} \Delta t^2 \]  
\[ \bar{\mu} = \bar{\mu}_0 + \dot{\mu} \Delta t + \frac{\mu_2}{60} \Delta t^3 + \frac{\mu_4}{24} \Delta t^4 \]  

In practice \( \bar{\omega} \) and \( \bar{n} \) are usually zero.

The above method for updating mean eccentricity can sometimes cause problems. The secular variation in eccentricity for most objects is caused principally by atmospheric drag which has the effect of circularizing the orbits. In other words \( \dot{\bar{e}} \) and \( \ddot{\bar{e}} \) are usually both negative. When \( \bar{e}_0 \) is small, \( \bar{e} \) can be negative for some values of \( \Delta t \). This creates problems when trying to transform these mean elements into osculating Cartesian elements. This problem can be circumvented if it is assumed that the eccentricity decay is due to a decrease in apogee height and that perigee height remains constant. This is usually a valid assumption, especially if the element rates were computed based on equations derived from that assumption (Section II). Then:

\[ \bar{e} = 1 - \bar{a}_0 \left( 1 - \bar{e}_0 \right) \]  

This method of trajectory generation has a major advantage over numerical integration. Since it is not a step-by-step procedure, the vector at time of interest can be immediately computed with no regard for vectors at intermediate times. When \( \Delta t \) is large, this advantage results in a considerable saving in computation time. Being a step-by-step procedure, numerical integration is subject to round-off and truncation errors which can be fatal when \( \Delta t \) is large.
The mean element update method of trajectory generation has proved to be sufficiently accurate for satellite acquisition at AFETR Mipir radar sites. The FORTRAN coding for this procedure is given in Appendix VII. This technique was compared against numerical integration for two cases: with and without drag. Table VII shows the initial conditions used for both methods. The resulting orbital characteristics are also shown as well as the ballistic coefficient used with the U. S. Standard Atmosphere 1962 model for drag computation. The gravity model used with the numerical integration is given in Table III. The central gravity term, \( \mu \), and the zonal coefficients, \( J_2 \) and \( J_3 \) were also used for the mean element computations.

The mean elements and their rates at epoch, obtained from the initial conditions as shown in Section II, are given in Table VIII. The differences between the trajectories are presented in intrack, crosstrack and radial components in Tables IX and X. Differences are shown at ten minute intervals for the first revolution, the last revolution of the first day and the last revolution of the week. The differences obtained in these examples are due to the use of tesseral harmonics in the integrated trajectory and to inadequacies inherent in the mean elements trajectory generation scheme. The differences are obtained by subtracting the mean element updated trajectory points from the numerically integrated trajectory. Note that in both the vacuum and drag cases, the radial and crosstrack errors oscillate about zero with ever increasing amplitude and very little secular variation, while the intrack oscillates about an ever-growing mean value.

Drag trajectory generations were also performed with zero values for \( \delta, \phi, \omega, \) and \( \Omega \). The resulting trajectory errors in radial and intrack components were considerably smaller than those shown in Table X. For example, mean differences at the end of one week were:

\[
\begin{align*}
\text{radial} &= -2226 \, \text{ft} \\
\text{crosstrack} &= 2179 \, \text{ft} \\
\text{intrack} &= -131289 \, \text{ft}
\end{align*}
\]

For some reason, these terms adversely affect the accuracy of this drag trajectory generation. It could be that they are useful for objects with high ballistic coefficients or in extremely low orbits. However, all cases studied have shown that optimum results are obtained with the above mentioned parameters set to zero.
**TABLE VII**

**THEORETICAL TRAJECTORY INITIAL CONDITIONS**

\[
\begin{align*}
x &= 17837622. \text{ ft.} \\
y &= 11170525. \text{ ft.} \\
z &= 4559553. \text{ ft.} \\
x' &= -14197. \text{ ft./sec.} \\
y' &= 17945. \text{ ft./sec.} \\
z' &= 11635. \text{ ft./sec.}
\end{align*}
\]

Trajectory Characteristics:

\[
\begin{align*}
h_p &= 100 \text{ nm} \\
h_a &= 150 \text{ nm} \\
P &= 89 \text{ min.}
\end{align*}
\]

Ballistic Coefficient:

\[
B = \frac{C_p A}{W} = 0.01 \text{ ft}^2/\text{lb}
\]
TABLE VIII

THEORETICAL MEAN ELEMENTS AND RATES

\[ \begin{align*}
\bar{a} & = 1.03625778 \text{ earth radii (ER)} \\
\bar{e} & = .006798 \\
\bar{I} & = 29.999538 \text{ deg} \\
\bar{\omega} & = 16.219243 \text{ deg} \\
\bar{\Omega} & = 9.998874 \text{ deg} \\
\bar{\Omega} & = 8.781831 \\
\dot{\bar{a}} & = -2.843243 \times 10^{-4} \text{ ER/day} \\
\dot{\bar{e}} & = -4.020530 \times 10^{-6} \\
\dot{\bar{I}} & = -2.725106 \times 10^{-4} \\
\dot{\bar{\omega}} & = -3.853477 \times 10^{-6} \\
\dot{\bar{\Omega}} & = 12.086671 \text{ deg/day} \\
\dot{\bar{\Omega}} & = 5.758724 \times 10^{-3} \\
\dot{\bar{\Omega}} & = -7.612567 \text{ deg/day} \\
\ddot{\bar{\Omega}} & = -0.003627 \text{ deg/day}^2 \\
\dot{\bar{\Omega}} & = 16.24925696 \text{ revs/day} \\
\ddot{\bar{\Omega}} & = 0.00332323 \text{ revs/day}^2 \\
\dddot{\bar{\Omega}} & = 3.208824 \times 10^{-5}
\end{align*} \]

*These rates set to zero for vacuum trajectory
<table>
<thead>
<tr>
<th>Prediction Time (min)</th>
<th>Prediction Errors (ft)</th>
<th></th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td>Radial</td>
<td>Crosstrack</td>
</tr>
<tr>
<td>0</td>
<td>0.0</td>
<td>0.0</td>
</tr>
<tr>
<td>10</td>
<td>16.0</td>
<td>16.0</td>
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<td>67.0</td>
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<td>78.0</td>
<td>97.0</td>
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<td>50</td>
<td>108.0</td>
<td>22.0</td>
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<td>60</td>
<td>99.0</td>
<td>-26.0</td>
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<td>70</td>
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<td>80</td>
<td>-63.0</td>
<td>201.0</td>
</tr>
<tr>
<td>(1 rev)</td>
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<td>(1 day) 1440</td>
<td>-32.0</td>
<td>108.0</td>
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</tr>
<tr>
<td>(1 week) 10080</td>
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</tr>
<tr>
<td>Last rev mean</td>
<td>-93.0</td>
<td>1318.0</td>
</tr>
</tbody>
</table>
## TABLE X

**DRAG TRAJECTORY DIFFERENCES**

*(POLYNOMIAL UPDATE)*

<table>
<thead>
<tr>
<th>Prediction Time (min)</th>
<th>Prediction Errors (ft)</th>
<th></th>
<th></th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td>Radial</td>
<td>Crosstrack</td>
<td>Intrack</td>
</tr>
<tr>
<td>0</td>
<td>0.</td>
<td>0.</td>
<td>0.</td>
</tr>
<tr>
<td>10</td>
<td>24.</td>
<td>16.</td>
<td>3.</td>
</tr>
<tr>
<td>20</td>
<td>34.</td>
<td>32.</td>
<td>4.</td>
</tr>
<tr>
<td>30</td>
<td>112.</td>
<td>67.</td>
<td>172.</td>
</tr>
<tr>
<td>40</td>
<td>208.</td>
<td>97.</td>
<td>367.</td>
</tr>
<tr>
<td>50</td>
<td>238.</td>
<td>22.</td>
<td>664.</td>
</tr>
<tr>
<td>60</td>
<td>162.</td>
<td>27.</td>
<td>951.</td>
</tr>
<tr>
<td>70</td>
<td>-38.</td>
<td>91.</td>
<td>1111.</td>
</tr>
<tr>
<td>80</td>
<td>-208.</td>
<td>201.</td>
<td>927.</td>
</tr>
<tr>
<td>(1 rev)</td>
<td>90</td>
<td>391.</td>
<td>194.</td>
</tr>
<tr>
<td>1350</td>
<td>941.</td>
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<td>29273.</td>
</tr>
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<td>1126.</td>
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</tr>
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</tr>
<tr>
<td>1390</td>
<td>2904.</td>
<td>817.</td>
<td>41387.</td>
</tr>
<tr>
<td>1400</td>
<td>968.</td>
<td>-967.</td>
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</tr>
<tr>
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<tr>
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<td>-2612.</td>
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<td>36347.</td>
</tr>
<tr>
<td>(1 day)</td>
<td>1440</td>
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<td>209.</td>
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<td>-1654791.</td>
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</tr>
<tr>
<td>(1 week)</td>
<td>10080</td>
<td>-72103.</td>
<td>19268.</td>
</tr>
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</table>

**Last rev mean**

|  | -66228. | 2241. | -1710972. |

46
The differences shown above resulting from the use of analytical secular variations are considerably larger than those obtained from numerical computations using real data (Section IV.3, Table XIV).

There is good agreement between numerical and mean element trajectories for the vacuum case (Table IX). Therefore the large intrack differences obtained for the drag case must be due to either drag-affected secular rates or their application to the trajectory generation problem. If a logarithmic rather than polynomial update scheme is used the intrack differences are considerably decreased. The logarithmic scheme is based upon the fact that the mean mean motion is a logarithmic function of time:

\[
\bar{n} = \bar{n}_0 + \log_e \left(1 + 2\frac{\delta}{2} \Delta t\right)
\]

and the mean semi-major axis is given by:

\[
\bar{a} = \bar{a}_0 \left[\frac{\bar{n}}{\bar{n}_0} + \frac{\Delta n}{\bar{n}_0}\right]^{\frac{5}{3}}
\]

The mean mean anomaly can then be updated by:

\[
\bar{M} = \bar{M}_0 + \bar{n}_a \Delta t
\]

where

\[
\bar{n}_a = \bar{n}_0 + \Delta n
\]

and

\[
\Delta n = \int_0^{\Delta t} \log_e \left(1 + \frac{\bar{n}}{2} t\right) dt
\]

which, after changing limits and applying a four-point Gauss' method of approximate quadratures, becomes (Ref. 10):

\[
\Delta n = \sum_{i=1}^{4} A_i \log_e \left[1 + \frac{\bar{n}_a}{2} \Delta t X_i\right]
\]

where \(A_i\) and \(X_i\) are quadrature values:

<table>
<thead>
<tr>
<th></th>
<th>(A_i)</th>
<th>(X_i)</th>
</tr>
</thead>
<tbody>
<tr>
<td>1</td>
<td>0.34785485</td>
<td>0.93056815</td>
</tr>
<tr>
<td>2</td>
<td>0.65214515</td>
<td>0.66999052</td>
</tr>
<tr>
<td>3</td>
<td>0.65214515</td>
<td>0.33000947</td>
</tr>
<tr>
<td>4</td>
<td>0.34785485</td>
<td>0.06943184</td>
</tr>
</tbody>
</table>

Using this logarithmic scheme of mean mean anomaly update, the differences shown in Table XI were obtained. As can be seen, the intrack differences are much improved over those given above which suggests that this method should be extensively studied both experimentally and analytically. However, preliminary results indicate that the improvement shown is typical of what can be achieved.
### TABLE XI

**DRAG TRAJECTORY DIFFERENCES**

*(LOGARITHMIC UPDATE)*

<table>
<thead>
<tr>
<th>Prediction Time (min)</th>
<th>Prediction Errors (ft)</th>
<th>Radial</th>
<th>Crosstrack</th>
<th>Intrack</th>
</tr>
</thead>
<tbody>
<tr>
<td>0</td>
<td>0</td>
<td>0.</td>
<td>0.</td>
<td>0.</td>
</tr>
<tr>
<td>1.0</td>
<td>24.</td>
<td>16.</td>
<td>-3.</td>
<td>41.</td>
</tr>
<tr>
<td>20</td>
<td>34.</td>
<td>32.</td>
<td>-172.</td>
<td>367.</td>
</tr>
<tr>
<td>30</td>
<td>112.</td>
<td>67.</td>
<td>-664.</td>
<td>951.</td>
</tr>
<tr>
<td>40</td>
<td>208.</td>
<td>97.</td>
<td>-926.</td>
<td>1111.</td>
</tr>
<tr>
<td>50</td>
<td>238.</td>
<td>22.</td>
<td>-208.</td>
<td>91.</td>
</tr>
<tr>
<td>60</td>
<td>162.</td>
<td>-27.</td>
<td>-391.</td>
<td>194.</td>
</tr>
<tr>
<td>70</td>
<td>-38.</td>
<td>91.</td>
<td>-1088.</td>
<td>-373.</td>
</tr>
<tr>
<td>80</td>
<td>-208.</td>
<td>201.</td>
<td>-2638.</td>
<td>-926.</td>
</tr>
<tr>
<td>(1 rev) 90</td>
<td></td>
<td>-391.</td>
<td>-26677.</td>
<td>373.</td>
</tr>
<tr>
<td>1350</td>
<td>-959.</td>
<td>22.</td>
<td>-23497.</td>
<td></td>
</tr>
<tr>
<td>1360</td>
<td>1045.</td>
<td>1885.</td>
<td>-23532.</td>
<td></td>
</tr>
<tr>
<td>1370</td>
<td>2824.</td>
<td>2852.</td>
<td>-26214.</td>
<td></td>
</tr>
<tr>
<td>1380</td>
<td>3479.</td>
<td>2384.</td>
<td>-30588.</td>
<td></td>
</tr>
<tr>
<td>1390</td>
<td>2722.</td>
<td>816.</td>
<td>-34860.</td>
<td></td>
</tr>
<tr>
<td>1400</td>
<td>838.</td>
<td>-966.</td>
<td>-37246.</td>
<td></td>
</tr>
<tr>
<td>1410</td>
<td>-1250.</td>
<td>-2265.</td>
<td>-36704.</td>
<td></td>
</tr>
<tr>
<td>1420</td>
<td>-2638.</td>
<td>-2554.</td>
<td>-33329.</td>
<td></td>
</tr>
<tr>
<td>1430</td>
<td>-2596.</td>
<td>-1572.</td>
<td>-29494.</td>
<td></td>
</tr>
<tr>
<td>(1 day) 1440</td>
<td>-1088.</td>
<td>209.</td>
<td>-26677.</td>
<td></td>
</tr>
<tr>
<td>9990</td>
<td>-6191.</td>
<td>18850.</td>
<td>61152.</td>
<td></td>
</tr>
<tr>
<td>10000</td>
<td>7573.</td>
<td>14260.</td>
<td>58903.</td>
<td></td>
</tr>
<tr>
<td>10020</td>
<td>16693.</td>
<td>-9791.</td>
<td>14265.</td>
<td></td>
</tr>
<tr>
<td>10030</td>
<td>7703.</td>
<td>-17655.</td>
<td>-5041.</td>
<td></td>
</tr>
<tr>
<td>10040</td>
<td>-6151.</td>
<td>-17014.</td>
<td>-7541.</td>
<td></td>
</tr>
<tr>
<td>10050</td>
<td>-18414.</td>
<td>-8062.</td>
<td>9340.</td>
<td></td>
</tr>
<tr>
<td>10060</td>
<td>-22944.</td>
<td>4920.</td>
<td>38732.</td>
<td></td>
</tr>
<tr>
<td>10070</td>
<td>-17505.</td>
<td>15544.</td>
<td>67476.</td>
<td></td>
</tr>
<tr>
<td>(1 week) 10080</td>
<td>-4795.</td>
<td>18673.</td>
<td>82652.</td>
<td></td>
</tr>
<tr>
<td>Last rev mean</td>
<td>-2788.</td>
<td>2172.</td>
<td>35977.</td>
<td></td>
</tr>
</tbody>
</table>
Another use to which element variations have been applied at the RTCS is that pertaining to orbit determination. This technique has been put to full use in the CDC 3100 computer program, CASS (Ref 10).

Orbit determination using mean elements is accomplished as follows. First, each pass of tracking data is reduced to a vector (position and velocity) using some method of orbit determination. Second, all such vectors are transformed to mean Keplerian elements using the inverse computations described in Section III.4. Third, polynomials in time are passed in a least squares sense through like elements, one from each pass. The polynomial coefficients thus determined are identical in definition to the parameters of an SDC bulletin (Appendix I). These coefficients then represent the multi-pass solution and can be used to predict future positions of the object as shown in Section IV.2.

In CASS, however, the polynomial coefficients are not the ones appearing on the final transmitted bulletin. It is necessary to analytically recompute some of the parameters in order to preserve known relationships among them (e.g. $a$ and $e$). Some of these relationships are given in Section II. Examples of actual single-pass mean elements used and the polynomial and analytical bulletins thus obtained are given in Tables XII, XIII and XIV. Prediction errors for the analytical bulletin are given in Table XV in radial, cross-track and intrack components. The solution and prediction errors were obtained using actual tracking data from AFETR Mipir radars. These errors are caused by several factors. First, the mean elements and their rates at epoch contain errors due to their computations being based on imperfect tracking data. Second, the vectors against which the predictions are compared to produce residuals are based on imperfect tracking data and contain errors. Third, the actual forces influencing the object are not completely modeled in the mean elements updating scheme. Examples of some not included in the computations are higher order gravity terms and atmospheric density variations. Approximately 20.7 days after epoch (17 August 1970) a large geomagnetic disturbance occurred which significantly effected the prediction errors (Table XV).

Notice in Table XIII that the single-pass mean element residuals from the polynomial solution show periodic errors in mean longitude despite the modeling of periodic variations in CASS. The cause of the residual long period variation is not known. It may be that they are due to unmodeled gravity or drag forces.
TABLE XII

CASS SINGLE-PASS MEAN ELEMENTS FOR ORBIT DETERMINATION

<table>
<thead>
<tr>
<th>DAY</th>
<th>T + F</th>
<th>MA</th>
<th>HE</th>
<th>N1</th>
<th>MAP</th>
<th>MRAN</th>
<th>MPA</th>
<th>STATION</th>
</tr>
</thead>
<tbody>
<tr>
<td>100</td>
<td>1.04652</td>
<td>1.09305797</td>
<td>31.7610568</td>
<td>346.935052</td>
<td>322.73834</td>
<td>129.23311</td>
<td>91A</td>
<td></td>
</tr>
<tr>
<td>101</td>
<td>1.0475227</td>
<td>1.09263407</td>
<td>31.7610627</td>
<td>61.294520</td>
<td>273.10596</td>
<td>-75.20777</td>
<td>00A</td>
<td></td>
</tr>
<tr>
<td>102</td>
<td>1.0475327</td>
<td>1.09257924</td>
<td>31.7741414</td>
<td>137.103627</td>
<td>224.87523</td>
<td>-94.33455</td>
<td>12A</td>
<td></td>
</tr>
<tr>
<td>103</td>
<td>1.0485182</td>
<td>1.09361356</td>
<td>31.7741509</td>
<td>195.729214</td>
<td>286.66719</td>
<td>-125.44498</td>
<td>12A</td>
<td></td>
</tr>
</tbody>
</table>
### TABLE XIII

**CASS POLYNOMIAL MEAN ELEMENT RESIDUALS**

<table>
<thead>
<tr>
<th>xA</th>
<th>xE</th>
<th>+E</th>
<th>+AP</th>
<th>+RAP</th>
<th>A = +AP + +A</th>
</tr>
</thead>
<tbody>
<tr>
<td>0.37903</td>
<td>0.00001</td>
<td>0.01198</td>
<td>0.0589090</td>
<td>0.0000843</td>
<td>0.0237936</td>
</tr>
<tr>
<td>0.084914+</td>
<td>0.0002124+</td>
<td>0.0630399</td>
<td>-1.6744449+</td>
<td>0.0178202+</td>
<td>-0.0655779</td>
</tr>
<tr>
<td>0.000326</td>
<td>0.0000009</td>
<td>0.036542</td>
<td>-0.0781195</td>
<td>0.0006999</td>
<td>0.0259819</td>
</tr>
<tr>
<td>0.0000008</td>
<td>0.0000012</td>
<td>0.014967</td>
<td>0.0177944</td>
<td>0.0019278</td>
<td>0.032518</td>
</tr>
<tr>
<td>0.000015</td>
<td>0.003756</td>
<td>0.010576</td>
<td>0.1369923</td>
<td>0.004199</td>
<td>0.032412</td>
</tr>
</tbody>
</table>

**Point edited from solution**
### TABLE XIV

CASS POLYNOMIAL AND ANALYTICAL ORBIT SOLUTIONS

<table>
<thead>
<tr>
<th>Element</th>
<th>Value</th>
<th>Derivatives</th>
<th>First</th>
<th>Second</th>
</tr>
</thead>
<tbody>
<tr>
<td>Semi-Major Axis</td>
<td>1.52998677E+01</td>
<td>-6.44654047E-06</td>
<td>-2.06021578E-07</td>
<td></td>
</tr>
<tr>
<td>Eccentricity</td>
<td>1.46559925E-05</td>
<td>-1.97589699E-05</td>
<td>-4.5230954E-07</td>
<td></td>
</tr>
<tr>
<td>Inclination</td>
<td>3.1763579E-01</td>
<td>0</td>
<td></td>
<td></td>
</tr>
<tr>
<td>Argument of Perigee</td>
<td>2.53127577E+01</td>
<td>9.54468554E-08</td>
<td></td>
<td>0</td>
</tr>
<tr>
<td>Right Ascension of Ascending Node</td>
<td>1.49353359E-02</td>
<td>-6.21794185E00</td>
<td>-8.41913575E-05</td>
<td></td>
</tr>
<tr>
<td>True Anomaly</td>
<td>4.42061337E-01</td>
<td>-4.49092993E-07</td>
<td>-1.6379541E-05</td>
<td></td>
</tr>
</tbody>
</table>

---

<p>| Semi-Major Axis                 | 1.09299865E+00      | -1.60194893E-06        | -3.24272797E-11       |                         |
| Eccentricity                     | 1.46959952E-02      | -1.44328835E-06        | -2.9232171E-11        |                         |
| Inclination                      | 3.1783579E-01       | 0                       |                        |                         |
| Argument of Perigee              | 2.53127577E+00      | 9.54004701E 00         |                        | 0                       |
| Right Ascension of Ascending Node| 1.49353359E-02      | -6.21794185E00         | -8.41913575E-05       |                         |
| True Anomaly                     | 8.42081332E-01      | 1.49093121E 01         | 1.6379541E-05         |                         |</p>
<table>
<thead>
<tr>
<th>Time From Epoch (days)</th>
<th>Prediction Errors (ft)</th>
<th></th>
<th></th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td>Radial</td>
<td>Crosstrack</td>
<td></td>
</tr>
<tr>
<td>4.2</td>
<td>466.</td>
<td>-2778.</td>
<td>5892.</td>
</tr>
<tr>
<td>8.1</td>
<td>-3206.</td>
<td>3673.</td>
<td>8994.</td>
</tr>
<tr>
<td>9.0</td>
<td>-4083.</td>
<td>3768.</td>
<td>9482.</td>
</tr>
<tr>
<td>15.1</td>
<td>-4344.</td>
<td>2393.</td>
<td>8799.</td>
</tr>
<tr>
<td>17.1</td>
<td>-3757.</td>
<td>775.</td>
<td>9588.</td>
</tr>
<tr>
<td>21.2</td>
<td>-1490.</td>
<td>-13790.</td>
<td>25658.</td>
</tr>
<tr>
<td>22.1</td>
<td>-4807.</td>
<td>-8062.</td>
<td>24480.</td>
</tr>
<tr>
<td>22.2</td>
<td>-4287.</td>
<td>-14080.</td>
<td>31676.</td>
</tr>
<tr>
<td>24.1</td>
<td>-3066.</td>
<td>-2328.</td>
<td>51068.</td>
</tr>
<tr>
<td>25.0</td>
<td>-1906.</td>
<td>2045.</td>
<td>67755.</td>
</tr>
<tr>
<td>28.0</td>
<td>-2452.</td>
<td>-9273.</td>
<td>144416.</td>
</tr>
</tbody>
</table>
Comparisons between SDC bulletin parameters and computed secular variations were presented in Section II, Table I. Here note the differences between the polynomial coefficients and the analytical parameters in Table XIV. Those parameters analytically recomputed were: \( \hat{a} \), \( \hat{b} \), \( \hat{n} \), \( \hat{e} \), \( \hat{\epsilon} \), \( \hat{\omega} \), and \( \hat{n} \) using methods similar to those given in Section II (Ref 11).
APPENDIX I

SPACE DEFENSE CENTER 5-CARD BULLETIN

<p>| Key on following page |</p>
<table>
<thead>
<tr>
<th>Card</th>
<th>Column</th>
<th>Description</th>
<th>Value</th>
</tr>
</thead>
<tbody>
<tr>
<td>1</td>
<td>1</td>
<td>Object Number</td>
<td>80148</td>
</tr>
<tr>
<td>2</td>
<td>2</td>
<td>Class Number</td>
<td>100147</td>
</tr>
<tr>
<td>3</td>
<td>3</td>
<td>Bulletin Number</td>
<td>610</td>
</tr>
<tr>
<td>4</td>
<td>4</td>
<td>Tr</td>
<td>00148</td>
</tr>
<tr>
<td>5</td>
<td>5</td>
<td>Name</td>
<td>a</td>
</tr>
<tr>
<td>6</td>
<td>6</td>
<td>M</td>
<td>00148</td>
</tr>
<tr>
<td>7</td>
<td>7</td>
<td>P</td>
<td>00148</td>
</tr>
<tr>
<td>8</td>
<td>8</td>
<td>Element</td>
<td>00148</td>
</tr>
<tr>
<td>9</td>
<td>9</td>
<td>Life From Epoch</td>
<td>00148</td>
</tr>
<tr>
<td>10</td>
<td>10</td>
<td>Revs. From Launch At Epoch</td>
<td>00148</td>
</tr>
<tr>
<td>11</td>
<td>11</td>
<td>TDD</td>
<td>00148</td>
</tr>
<tr>
<td>12</td>
<td>12</td>
<td>ID</td>
<td>00148</td>
</tr>
<tr>
<td>13</td>
<td>13</td>
<td>Reflectivity Factor</td>
<td>00148</td>
</tr>
</tbody>
</table>

**Card 2: Modified Julian Date of Epoch**
- **Object Number**: 80148
- **Modified Julian Date of Epoch**: 40281.78295506
- **Mean Anomaly (Deg)**: 045.0356
- **Right Ascension of Asc. Node (Deg)**: 348.1962
- **Argument of Perigee (Deg)**: 313.3246
- **Eccentricity (no units)**: 0.0210250
- **Inclination (Deg)**: 065.7968

**Card 3: Mean Motion (Revs/MSD)**
- **Object Number**: 80148
- **Mean Motion (Revs/MSD)**: 13.77087298
- **Mean Motion (Revs/MSD^2)**: b.0000006037
- **Mean Motion (Deg/MSD)**: -2.48478
- **Mean Motion (Deg/MSD^2)**: -0.48369
- **Mean Motion (Deg/MSD^3)**: -57230.6

**Card 4: Semi Major Axis (Earth Radii)**
- **Object Number**: 80148
- **Semi Major Axis (Earth Radii)**: b00000000000
- **Semi Major Axis (ER/MSD)**: -24822.5
- **Semi Major Axis (ER/MSD^2)**: -48319.6

**Card 5: Note**
- **Object Number**: 80148
- **Semi Major Axis (Earth Radii)**: 01.15286023
- **Semi Major Axis (ER/MSD)**: -6739576.6
- **Semi Major Axis (ER/MSD^2)**: 492491.12

**Notes**
- **MSD**: Mean Solar Day
- **b**: blank column
- **s**: sign of value/exponent
APPENDIX II

SUBROUTINE DIPDA

(Semi-Major Axis Decay Rate Computation - See SECTION II)
SUBROUTINE DRPNT(FI,AF,E2,A,E,F0,RHOEFF,ADDY)

THE PURPOSE OF THIS ROUTINE IS TO COMPUTE DRAG DECAY RATES OF APOGEE, PERIGEE AND SEMI-MAJOR AXIS BY INTEGRATION USING SIMPSONS RULE.

INPUTS ARE:
- ROOM GRAVITY CONSTANT (Ft^3/Sec^2)
- EARTH SEMI-MAJOR AXIS (FT)
- EARTH ECCENTRICITY SQUARED
- ORBIT SEMI-MAJOR AXIS (FT)
- ORBIT ECCENTRICITY
- ORBIT INCLINATION (RAD)
- ORBIT ARGUMENT OF PERIGEE (RAD)
- B0=OBJECT BALLISTIC COEFFICIENT (CD*A7W) (FT^2/LB)

WHERE:
- CD=COEFFICIENT OF DRAG
- W=OBJECT WEIGHT

OUTPUT IS:
- ADDY=DECAY RATE OF ORBIT SEMI-MAJOR AXIS (FT/DAY)

ATMOSPHERE MODEL SUBROUTINE IS ATMOSP

DATA(PT=3,1415926536)

TPI = 3. * PI
SUNPE=96400. * SQRF(A*E2)/TPI
F0=E2/(1.+E2)
D=4.*(1.-E^2)
SO=SINF(F0)
CO=COSF(F0)
SI=SINF(F1)

COMPUTE TRUE ANOMALY STEP SIZE

NV=PI/36.

V=0.
Y=0.
YP=0.
1 SV=SINF(V)
CV=COSF(V)

COMPUTE SATELLITE HEIGHT

SH=SI*(SO*CV+CO*SV)
RE=AF/SQRTF(1.+E*E+SP)
REP/1.*E*CV)

58
H = H - HE
CALL ATMOS(H,RH)
F=SQRT(F(1.+2.*E+*E)*RHO)
F=2.*F(1.+CV)
F=2.*F(1.-CV)
IF(1.E0.(1/2)*2) GO TO 3
2 FA=2.*FA
FP=2.*FP
3 YA=YA+FA
YP=YP+FP
IF(1.E0.35) GO TO 4
I=I+1
V=V+DV
GO TO 1
4 YA=YA+DV*SQRT((1.+E)/(1.-E)*3)/3.
YP=YP+DV*SQRT((1.-E)/(1.+E)*3)/3.
C
C COMPUTE APDOGEF DECAY RATE
C
ADHDUT=SOTP*YP*COEF.
C
COMPUTE PERIGEE DECAY RATE
C
PHDOUT=SOTP*YP*COEF.
C
COMPUTE SEMI-MAJOR AXIS DECAY RATE
C
ADOTE,4*(ADHDUT-PHDOUT)
END

NOT REPRODUCIBLE
APPENDIX III

SUBROUTINE ELRAT

(Mean Elements Secular Rates Computation - See SECTION II)
THE PURPOSE OF THIS ROUTINE IS TO COMPUTE ELEMENT RATES

INPUTS ARE

AF= EARTH SEMI-MAJOR AXIS (FT)
F? = EARTH ECCENTRICITY SQUARED
F? = EARTH GRAVITY CONSTANT (FT**3/SEC**2)
FJ2=EARTH GRAVITY SECOND ZONAL HARMONIC COEFFICIENT
= OBJECT BALLISTIC COEFFICIENT (FT**2/LP)
FM(1)*(6)=MEAN ELEMENTS AT EPOCH
FM(1)=MEAN SEMI-MAJOR AXIS (FT)
FM(2)=MEAN ECCENTRICITY
FM(3)=MEAN INCLINATION (DEG)
FM(4)=MEAN ARGUMENT OF PERIGEE (DEG)
FM(5)=MEAN RIGHT ASCENSION OF ASCENDING NODE (DEG)
FM(6)=MEAN MEAN ANOMALY (DEG)

OUTPUTS ARE

FM(1)*(12)=MEAN ELEMENT RATES AT EPOCH
FM(1)=MEAN SEMI-MAJOR AXIS RATE (ER= EARTH RADII)
FM(2)=AJOIT=SEMl-MAJOR AXIS RATE (ER/DAY)
FM(3)=AJ02 = (FT/DAY**2)
FM(4)=EDOT= ECCENTRICITY RATE (/DAY)
FM(5)=JU02 = (DEG/DAY**2)
FM(6)=JOIT=ARGUMENT OF PERIGEE RATE (DEG/DAY)
FM(7)=JU02 = (DEG/DAY**2)
FM(8)=JOIT=RIGHT ASCENSION OF ASCENDING NODE RATE (DEG/DAY)
FM(9)=JU02 = (DEG/DAY**2)
FM(10)=VN=MEAN MOTION (REV/DAY)
FM(11)=V02 = (REV/DAY**2)
FM(12)=V03 = (REV/DAY**3)

DIMENSION FM(6),FM(12)
TYPE REAL NV,IR,08,JOIT,V001,NN06,N002,N002
DATA(N)GRA=57.7957775131,(FK3=360.)
COMPUTE SEMI-MAJOR AXIS, ECCENTRICITY AND MEAN MOTION

xy=1.5*J2*SORTF(1.0-8.0)(1.0-1.5*S1*S1)/p2
2=M/SORTF(FM)*F<K1*(1.0)*M(AE=AE+AH))

COMPUTE SECULAR RATES OF ARG. OF PERIGEE AND R.A. ASCEN.

60 OUT=75*V+N*J2*F*K3*(5.*C1+C1-1.0)/p2
70 OUT=1.0*V+N*J2*F*K3*(5.0)/p2

COMPUTE RATES OF SEMI-MAJOR AXIS, ECCENTRICITY AND MEAN MOTION

CALL HYPHA(FMU,AE=2.0,EB,18,10,3,AE)
ADD=OCT/AC

10 ADDT=ADDT+(1.0-AT)/AV

DEQ=2.0+AV/(4.*AV)

COMPUTE ARGUMENT OF PERIGEE AND RIGHT ASCENSION OF ASCENDING NODE

ACCELERATION TERMS

TEMP=ADD2*(1.0-AT)/(3.*(1.0-AT))/AV

4002=ADD*TEMP

41002=ADD*TEMP

IF(420.0>0.0) 50 TO 41

COMPUTE SEMI-MAJOR AXIS, ECCENTRICITY AND MEAN MOTION ACCELERATION TERMS

C=AV02/AV

AC04=1.0

IF(VV/VV,16,2)4) AC0V=1.0.

C=AV04+C*1.0-16.667/(3.*(16.667-N1))

11 ADD=ADD-(-2.*1+20.*C=7.9)

12 ADD=ADD-(3.0*ADD+5.*ADD+VU2)/(6.*AV)

ADD2=(1.-AT)*ADD/AN

60 100

41 ADDJ2=0.

61 ADDJ2=0.

ENDJ2=0.

50 CONTINUE

SET UP OUTPUT ARRAY

FN(1)=AV

FN(2)=ADD

FN(3)=ADD2

FN(4)=ADD2
APPENDIX IV
SUBROUTINE DELTA

(Aozai Short Period Element Variations - See SECTION III.2)
SUBROUTINE DELTA (AE, J2, M, I)

THE PURPOSE OF THIS SUBROUTINE IS TO COMPUTE SHORT PERIOD ELEMENT EVARIATIONS USING THE KOZAI METHOD.

INPUTS ARE

AE = EARTH SEMI-MAJOR AXIS (FT)
J2 = EARTH GRAVITY SECOND ZONAL HARMONIC COEFFICIENT
M(1) = MEAN SEMI-MAJOR AXIS (FT)
M(2) = MEAN ECCENTRICITY
M(3) = MEAN INCLINATION (DEG)
M(4) = MEAN ARGUMENT OF PERIGEE (DEG)
M(5) = MEAN RIGHT ASCENSION OF ASCENDING NODE (DEG)
M(6) = MEAN ANOMALY (DEG)

OUTPUTS ARE DIFFERENCES (OSCULATING - MEAN) OF ABOVE ELEMENTS IN THE SAME UNITS STORED IN D(1) THRU D(6)

DIMENSION M(6), D(6)

TYPE REAL M, J2, J, E

DATA (PI=3.1415926536), (F=5.295795111), (D=.33333333333)
DATA (T=.66868666667)

C

IF I = 2, L PI
JF = 1.5 * M * AE * AF
E = M(2)
IF (E<1,2,2
1 E = 0
M(2) = E
GO TO 4
2 IF (E<1,4,3,3
3 E = 9999999999999
M(2) = E
4 E = (F-F)* E
OME2 = 1. - E
SOME = 1. / (OME2)**1.5
FM = M(6) / FK
9 IF (AHSF(FM)-PI) 10,10,11
C REDUCE MEAN ANOMALY TO RANGE OF -PI TO +PI
10 FM = FM - PI * AHSF(FM) / FM
GO TO 9
C COMPUTE ECCENTRIC ANOMALY
10 EP = FM
DO 14 J = 1, 10
IF (AHSF(EA-EP), LT, 1, E = 10) GO TO 20
14 EP = EA

C COMPUTE TRUE ANOMALY
20 V = 2. * ATAN2(SINTF(1.,E)/(1.-E)) * SINF(AE*5), COSF(AE*5))
IF (FM) 30 31, 31
30 FM=FM+TP1
31 IF (V) 32, 33, 33
32 V=V+TP1
33 IF (V+TP1) 34, 35, 35
35 V=V+TP1
34 AB=M(1)
   FI=M(3)/FK
   FO=M(4)/FK
   FN=M(5)/FK

COMPUTE RADIO AND SFIN-LATUS RECTUM
R=AB*(1.-H(2)*COSF(EA))
P=AB*OME2
P2=P*P
AOR3=(AB/R)**3

COMPUTE VARIOUS FUNCTIONS OF SIN- AND COSING OF C INCLINATION, TRUE ANOMALY, AND ARGUMENT OF PERIGEE
S12=SINF(FI)**2
SFI=1.-1.5*S12
SV=SINF(V)
CV=COSF(V)
SO=SINF(F0)
CO=COSF(F0)
SV0=SV*CO+CV*SO
CV0=CV*CO-SV*SO
S2V0.*,SV*CV
C2V0=.*,CV*CV-1.
S3V0=SV*(3.-4.*SV*SV)
C3V0=CV*(4.*CV*CV-3.)
S4V0=SV*C3V-CV*S3V
C4V0=CV*C3V-SV*S3V
S5V0=S2V0*C3V+C2V0*S3V
C5V0=C2V0*C3V-S2V0*S3V
S2V0=.*,SO*CO
C2V0=.*,CO*CO-1.
S2V0=2.*SV0*CVO
C2V0=2.*CV0*CV0-1.
S2V0=SV0*C20+CV0*S20
C2V0=CV0*C20-SV0*S20
S3V0=S3V0*C20+C3V0*S20
C3V0=C3V0*C20-S3V0*S20
S4V0=S4V0*C20+C4V0*S20
S5V0=S5V0*C20+C5V0*S20

COMPUTE K07A1 SHORT PERIOD VARIATIONS
0(1)=JE*(1.)*SFI+TT*(2*OM3-SOM)*AOR3*S12*C2V0
DE1=JE*OM2/(1-AB*AB+E)*(1.*SFI+(AOR3-SOM)+.5*AOR3*S12*C2V0)
DE2=JE*S12/(2.-AB*AB+E)*C2V0+C20+CT-E*C3V0
0(2)=DE1-DE2
D(3)*JE*SF/(2.*FI)/(4.*P2)*(C2*V0+E*V20+O*T=F*C3*V20)*FK
D01=(2.*-2.*5.*S12)*(V=F+*E*SV)+SF+((1.*-25*ESQ)*SV/F+
1.*S2*V+*C.S5*V/12.)
D02=+SV20/(4.*S12+ESQ*(5.*-9375*S12))+.0625*E*SI2*SV*V20
D03=.5*(1.-2.*5.*S12)*S2*V+SV*V20/E*(7.*S12/12.-ESQ)
1 (1.*-19.*S12/8.)/6.1
D04=.375*SI2*SV*V20+.0625*E*SI2*SV*V20-.375*SI2*SV20
D(4)=JE/P2*(D01+D02+D03+D04)*FK
D(5)=JE*COSF(FI)/P2*(V=F*E*SV-.5*S2*V0-.5*E*SV*V20-
1 E*SV*V20/6.)*FK
D(6)=SF+I*((1.*-25*ESQ)*SV+.5*E*SV+ESQ*S3*V/12.)
D(6)=SI2+*(25*SV*V20+1.*+1.*ESQ)-.0625*ESQ+S3*V0-
1 7.*S3*V0/12.+
2 (1.*-ESQ/28.)-.375*E*SV*V20+.0625*ESQ*S5*V20+.375*E*SV20
D(6)=JE/(E*P2)*SORTF(0*ME2)*(SM1+SM2)*FK
END

NOT REPRODUCIBLE
APPENDIX V

SUBROUTINE DELXYZ

(Frazer Long and Short Period Element Variations - See SECTION III.3)
SUBROUTINE DFLXYZ(A, B, C, X, Y, Z)
C THIS SUBROUTINE ACCEPTS MEAN CARTESIAN ELEMENTS AND
C COMPUTES THE DIFFERENCES BETWEEN MEAN AND OSCULATING
C CARTESIAN ELEMENTS. REFERENCE IS FRASER, SEPT. 1966.
C
C INPUTS ARE
C A = EARTH SEMI-MAJOR AXIS (FT)
C F4 = EARTH GRAVITY CONSTANT (FT**3/SEC**2)
C F2 = EARTH GRAVITY SECOND ZONAL HARMONIC COEFFICIENT
C F3 = EARTH GRAVITY THIRD ZONAL HARMONIC COEFFICIENT
C XZ1, YZ1, XZ2, YZ2 = MEAN POSITION COMPONENTS IN AN
C EARTH-CENTRED, EQUATORIAL, INERTIAL, RIGHT-HANDED,
C CARTESIAN SYSTEM, WITH (1) TOWARD THE VERTICAL
C EQUINOX AND (5) THRU THE NORTH POLE (FT)
C FT = (4, 5, 6) - VELOCITY COMPONENTS IN
C ABOVE SYSTEM (FT/SEC)
C
C OUTPUT ARE
C (OSCI-MAN) VARIATIONS IN CARTESIAN ELEMENTS
C (OSCU-MAN) (FT) AND (FT/SEC)
C
C DIMENSION XZ(6), YZ(6)
C IF (J, J = 1, 10) 10
C 10 X = XZ(I)
C Y = XZ(J)
C Z = XZ(K)
C 10
C X0 = XZ(4)
C Y0 = XZ(5)
C Z0 = XZ(6)
C R0 = SQRT(F)
C
C COMPUTE RADIUS
C R0 = SQRT((X+Y+Z)*2)
C
C COMPUTE INVERSE OF SEMI-MAJOR AXIS
C X2 = X*Y*Z/(Y+Z)*X
C AINV = 0.2*F4*RA*RA*(X2)/F3*RA)
C IF (AINV .LE. 1, .GT. 1) GO TO 10
C
C NOT REPRODUCIBLE
C
C H = SQRT(F)
C
C COMPUTE INCLINATION
C SINT = SQRT(F)*Y
C IF (SINT .LT. 0.017) GO TO 10
C
C NOT REPRODUCIBLE
C
C R = Y/SINT
C SINT = Y/H
C SINT = -SINT
C
C NOT REPRODUCIBLE
C
C END
COMPUTE ARG. OF LATITUDE
\[ \sin u = \frac{X + \eta Y + \chi}{(\alpha + \psi \sin u)} \]
\[ \cos u = \frac{X + \eta Y + \chi}{(\alpha + \psi \sin u)} \]
\[ \sin^2 u = \frac{\sin u}{\cos u} \]
\[ \cos^2 u = 1 - 2 \sin^2 u \]
\[ \sqrt{v} = \sqrt{X + Y} + \frac{\pi}{2} \]

COMPUTE RADIUS RATE
\[ \theta = \theta \sin u / \varphi \]
\[ \varphi = \theta \cos u / \psi \]

COMPUTE SEMI-LATUS RECTUM
\[ \vartheta = \varphi / \varphi \]
\[ \sqrt{p} = \sqrt{p} \sin u / \cos u \]
\[ \cos \psi = \cos (\psi \sin u / \cos u) \]

COMPUTE ECCENTRICITY SQUARED
\[ e^2 = \sin^2 u + \cos^2 u \]
\[ \text{IF}(2 \cdot \eta \cdot \psi, \text{GO TO 10}) \]
\[ \eta = \frac{\psi \cos u}{(1 + \eta)} \]
\[ \text{IF}(\alpha \neq \psi, \text{GO TO 10}) \]
\[ \text{NE}=\text{ALPHA} \cdot (\psi + \eta) \]
\[ \text{NE} = \text{ALPHA} \cdot (\psi + \eta) \]

COMPUTE SHORT PERIOD PERBUBATIONS IN SPECIAL ELEMENTS
\[ S_{12} = \sin^2 u \]
\[ S_{12} = \cos^2 u \]
\[ S_{12} = \sin^2 u \]
\[ S_{12} = \cos^2 u \]
\[ S_{12} = \sin^2 u \]
\[ S_{12} = \cos^2 u \]
\[ S_{12} = \sin^2 u \]
\[ S_{12} = \cos^2 u \]

NOT REPRODUCIBLE
COMPUTE SHORT PERIOD PERTURBATIONS IN CARTESIAN ELEMENTS

\[ X = \cos \iota \cdot \cos \omega - \sin \iota \cdot \sin \omega \cdot \cos \delta \]
\[ Y = \cos \iota \cdot \sin \omega + \sin \iota \cdot \cos \omega \cdot \cos \delta \]
\[ Z = \sin \iota \cdot \sin \delta \]
\[ X' = \cos \iota \cdot \cos \omega - \sin \iota \cdot \sin \omega \cdot \sin \delta \]
\[ Y' = \cos \iota \cdot \sin \omega + \sin \iota \cdot \cos \omega \cdot \sin \delta \]
\[ Z' = \cos \omega \cdot \cos \delta \]

\[ \text{TEMP} = \text{NSU} \cdot \cos \iota \cdot \text{DES} \]
\[ \text{TM} = \text{SU} \cdot \text{DES} \cdot \text{CSU} \cdot \sin \iota \cdot \text{DES} \]
\[ \text{NSX} = \text{NSR} \cdot \text{YU} \cdot (\text{TM} \cdot \text{XX} \cdot \text{TM} \cdot \text{XX}) \]
\[ \text{NSY} = \text{NSR} \cdot \text{YU} \cdot \text{RAI} \cdot (\text{TM} \cdot \text{YY} \cdot \text{TM} \cdot \text{YY}) \]
\[ \text{NSZ} = \text{NSR} \cdot \text{JU} \cdot \text{RAI} \cdot (\text{TM} \cdot \text{ZZ} \cdot \text{TM} \cdot \text{ZZ}) \]
\[ \text{TC} = \text{NSR} \cdot \text{YD} \cdot \text{TM} \cdot \text{TP} \]

COMPUTE LONG PERIOD PERTURBATIONS IN CARTESIAN ELEMENTS

\[ \text{ALFA} = \frac{\text{JSJ}}{\text{FJ} \cdot \text{JF} + \text{FJ} \cdot \text{FJ}} \]
\[ \text{ES0} = \text{SU} \cdot \text{CSU} \cdot \cos \iota \cdot \text{ESW} \]
\[ \text{TA} = \text{SNT} \cdot \text{OPECV} \cdot (\text{ALFA} \cdot \text{SNU}) \]
\[ \text{TC} = \text{COST} \cdot (\text{ALFA} \cdot \cos \iota) \]
\[ \text{TD} = \text{SNI} \cdot \text{OPECV} \cdot (\text{ALFA} \cdot \text{CSU}) \]
\[ \text{TE} = \text{SNI} \cdot (\text{ALFA} \cdot (\text{SNU} \cdot \text{ESW})) \]
\[ \text{Tf} = -\text{COST} \cdot (\text{ALFA} \cdot \text{ESW}) \]
\[ \text{DELX} = \text{RAD} \cdot (\text{TA} \cdot \text{XX} \cdot \text{TH} \cdot \text{XX} \cdot \text{TC} \cdot \text{XX}) \]
\[ \text{DELY} = \text{RAD} \cdot (\text{TA} \cdot \text{YY} \cdot \text{TH} \cdot \text{YY} \cdot \text{TC} \cdot \text{YY}) \]
\[ \text{DELZ} = \text{RAD} \cdot (\text{TA} \cdot \text{ZZ} \cdot \text{TH} \cdot \text{ZZ} \cdot \text{TC} \cdot \text{ZZ}) \]
\[ \text{TEM} = \text{SMI} \cdot \text{SPT} \]
\[ \text{DELX} = \text{TEM} \cdot (\text{TM} \cdot \text{XX} \cdot \text{TH} \cdot \text{XX} \cdot \text{TC} \cdot \text{XX}) \]
\[ \text{DELY} = \text{TEM} \cdot (\text{TM} \cdot \text{YY} \cdot \text{TH} \cdot \text{YY} \cdot \text{TC} \cdot \text{YY}) \]
\[ \text{DELZ} = \text{TEM} \cdot (\text{TM} \cdot \text{ZZ} \cdot \text{TH} \cdot \text{ZZ} \cdot \text{TC} \cdot \text{ZZ}) \]

10 CONTINUE

20 CONTINUE

END

NOT REPRODUCIBLE
APPENDIX VI

SUBROUTINE OSMCE

(Iteration Control for Computing Mean From Osculating Elements - See SECTION III.4)
SUBROUTINE DECE(AE,FM1,FMJ,FJ3,XY7,XYZ)

THE PURPOSE OF THIS SUBROUTINE IS TO COMPUTE MEAN FROM
OSCATULATING CARTESIAN ELEMENTS ITERATIVELY USING DELXYZ.

INPUTS ARE

AE = EARTH SEMI-MAJOR AXIS (FT)
FM1 = EARTH GRAVITY CONSTANT (FT/SEC**2)
FJ3 = EARTH GRAVITY SECOND ZONAL HARMONIC COEFFICIENT
FJ1 = EARTH GRAVITY THIRD ZONAL HARMONIC COEFFICIENT

XYZ(1), XYZ(2), XYZ(3) = OSCULATING POSITION COMPONENTS
IN AN EARTH-CENTERED, EQUATORIAL, IDEALIZED,
RIGHT-HANDED CARTESIAN SYSTEM, WITH (1) TANGENT
THE VELOCITY AT INERTIAL AND (3) TO THE NORTH POLE (FT)

XYZ(4), XYZ(5), XYZ(6) = VELOCITY COMPONENTS IN
ABOVE SYSTEM (FT/SEC).

OUTPUTS ARE

XYZ(1) - XYZ(6) = MEAN POSITION AND VELOCITY COMPONENTS IN
SYSTEM SIMILAR TO ABOVE (FT) AND (FT/SEC).

ESTIMATING INITIAL ESTIMATE OF MEAN ELEMENTS

DIMENSION XYZ(6), XYZ(6), XYZ(6), XYZ(6)

1 CONTINUE

XYZ(J)=XYZ(J)

2 CONTINUE

1=1+1

3 CONTINUE

CALL DELXYZ(AE,FM1,FJ3,XYZ,DES)

4 CONTINUE

IMPROVE ESTIMATE OF MEAN ELEMENTS

J=1+1

5 CONTINUE

IMPROVE ESTIMATE OF MEAN ELEMENTS

J=1+1

6 CONTINUE

TEST FOR CONVERGENCE

J=1+1

7 CONTINUE

IF(AESF(3)(XY7(J)),GT.1.) GO TO 2

8 CONTINUE

IF(AESF(3)(XY7(J)),GT.0.01) GO TO 2

9 CONTINUE

10 CONTINUE

11 CONTINUE

12 CONTINUE

13 CONTINUE

14 CONTINUE

15 CONTINUE

16 CONTINUE

17 CONTINUE

18 CONTINUE

19 CONTINUE

20 CONTINUE

21 CONTINUE

22 CONTINUE

23 CONTINUE

24 CONTINUE

25 CONTINUE

26 CONTINUE

27 CONTINUE

28 CONTINUE

29 CONTINUE

30 CONTINUE

31 CONTINUE

32 CONTINUE

33 CONTINUE

34 CONTINUE

35 CONTINUE

36 CONTINUE

37 CONTINUE

38 CONTINUE

39 CONTINUE

40 CONTINUE

41 CONTINUE

42 CONTINUE

43 CONTINUE

44 CONTINUE

45 CONTINUE

46 CONTINUE

47 CONTINUE

48 CONTINUE

49 CONTINUE

50 CONTINUE

51 CONTINUE

52 CONTINUE

53 CONTINUE

54 CONTINUE

55 CONTINUE

56 CONTINUE

57 CONTINUE

58 CONTINUE

59 CONTINUE

60 CONTINUE

61 CONTINUE

62 CONTINUE

63 CONTINUE

64 CONTINUE

65 CONTINUE

66 CONTINUE

67 CONTINUE

68 CONTINUE

69 CONTINUE

70 CONTINUE

71 CONTINUE

72 CONTINUE

73 CONTINUE

74
CONTINUE

END
APPENDIX VII

SUBROUTINE UPDAT

(Analytical Trajectory Generator - See SECTION IV.2)
SUBROUTINE UPDATE(TH, XM, XN, DAY, FRAY, A, E, FI, FO, FU, FN)

THE PURPOSE OF THIS SUBROUTINE IS TO PERFORM TRAJECTORY
GENERATION USING MEAN ELEMENT SECULAR VARIATIONS AND A
MACLAURIN'S SERIES EXPANSION.

INPUTS ARE

TH = FRACTIONAL PORTION OF EPOCH (DAYS)

XM(1)- (12) ARE THE MEAN ELEMENTS AND
THEIR RATES AT EPOCH.

XM(1) = MEAN SEMI-MAJOR AXIS (FT)
XM(2) = MEAN ECCENTRICITY
XM(3) = MEAN INCLINATION (DEG)
XM(4) = MEAN ARGUMENT OF PERIGEE (DEG)
XM(5) = MEAN RIGHT ASCENSION OF ASCENDING NODE (DEG)
XM(6) = MEAN PARANALY (DEG)
XM(7) = MEAN SEMI-MAJOR AXIS (EARTH RADIUS)
XM(8) = DOT=SEMIAJOR AXIS RATE (FT/DAY)
XM(9) = DOT=ECCENTRICITY RATE (/DAY)
XM(10) = DOT=INCLINATION RATE (DEG/DAY)
XM(11) = DOT=ARGUMENT OF PERIGEE RATE (DEG/DAY)
XM(12) = DOT=RIGHT ASCENSION OF ASCENDING NODE RATE (DEG/DAY)

FM = FRAC. PORTION OF TIME OF UPDATED ELEMENTS (DAYS)

OUTPUTS ARE

NM = UPDATED MEAN SEMI-MAJOR AXIS (FT)
N E = UPDATED MEAN ECCENTRICITY
NF = UPDATED MEAN INCLINATION (DEG)
NE = UPDATED MEAN ARGUMENT OF PERIGEE (DEG)
NF = UPDATED MEAN RIGHT ASCENSION OF ASCENDING NODE (DEG)

DIMENSION IN(2), XM(1), XM(12)

DEL = "DAY-T(1)]+(FMAY-T(2])
DEL = DEL*ELT
A=XM(1)+DEL+XM(2)+ELT2*XM(3)
F=XM(1)*XM(2)+XM(3)/A
F = XM(1)

NOT REPRODUCIBLE 78
FD = XM(4) + XM(6) + DELT * XM(7) * DEFT

FN = XM(9) + XM(2) * DELT * XM(9) * DEFT

FM = XM(8) + 360 * (XM(1) + DELT * XM(11) - DELT + XM(12) + DELT * DEFT)

C REDUCE MEAN ANOMALY TO RANGE OF 0 TO 360

FM = MOD(FM, 360)

[FM, 1.0]. FM = FM * 360

END

NOT REPRODUCIBLE
REFERENCES


